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**The influence of rock heterogeneity on the scaling properties of simulated
and natural stylolites**

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Abstract

Stylolites are among the most prominent deformation patterns in sedimentary rocks that document localized pressure solution. Recent studies revealed that stylolite roughness is characterized by two distinct scaling regimes. The main goal of the present study is to decipher whether this complex scaling behavior of stylolites is caused by the composition of the host rock, *i.e.* heterogeneities in the material, or is governed by inherent processes on respective scales, namely the transition from a surface energy to an elastic energy dominated regime, as theoretically predicted. For this purpose we have developed a discrete numerical technique, based on a lattice spring model, to simulate the competition between stress, strain, and dissolution during stylolite roughening. We varied systematically the quenched noise, initially present in the material, which controls the roughening. We also changed the size, amount, and dissolution rate of the heterogeneities introduced in our model and evaluated the influence on the scaling exponents. Our findings demonstrate that the roughness and growth exponents are independent of the exact nature of the heterogeneities. We discovered two coinciding crossover phenomena in space and time that separate length and timescales for which the roughening process is either balanced by surface and elastic energies. Our observations are coherent with analytical predictions and with investigations quantifying precisely the scaling laws in the morphology of natural stylolites. The findings presented here can further be used to refine volume loss estimates from the finite strain pattern of stylolites.

1. Introduction

Pressure solution in sedimentary rocks results in either intergranular or localized dissolution of material (e.g. Tada & Siever, 1989). The latter is responsible for the formation of stylolites, a frequent deformation pattern in sedimentary rocks (e. g. Stockdale, 1922; Dunnigton 1954, Heald, 1955, Park & Schot, 1968, Buxton & Sibley 1981, Rutter, 1983; Railsback, 1993). Stylolites are rough interfaces that frequently contain insoluble material (Fig. 1), which is

49 considered to be the residuum of the dissolved rock (Railsback, 1993; and references cited
50 therein). Stylolite initiation is still highly debated (e.g. Tada & Siever, 1989) but several
51 mechanisms have been proposed that are in agreement with field observations: Formation (I)
52 along preexisting anisotropies (Bathurst, 1987) (II) as anticracks (Fletcher and Pollard, 1981)
53 that propagate due to stress concentrations at anticrack tips (even though this idea was
54 challenged recently by Katsman et al. 2006) and (III) by stress induced self-organization
55 (Merino, 1992, Railsback, 1998; Merino et al., 2006).

56 In the present study we focus on a quantitative description and characterization of the
57 roughness of simulated stylolites and study their dynamic development independent of the
58 process leading to the initial development of the localization of dissolution along a plane.
59 Based on recent quantitative methods of stylolite roughness characterization (Renard et al.,
60 2004; Schmittbuhl et al., 2004, Koehn et al., 2007, Ebner et al., submitted) we will use
61 statistical tools to compare simulated and natural stylolites. In particular we study the
62 influence of initial heterogeneity concentration in the host-rock on a) stylolite roughness, b)
63 dynamic roughness growth and c) the correlation of crossover phenomena in space and time.
64 To integrate the results of our study in the context of quantitative characterization we will first
65 review some of the major findings and basic principles, used in recent studies, necessary for
66 the understanding of our approach.

67 The exact classification of stylolites in the field is a difficult task because there is a wide
68 range of geometries (e.g. Park and Schott, 1968) that are often transitional even within a
69 single outcrop. Many previous studies (Park & Schot, 1968; Buxton & Sibley, 1981,
70 Guzzetta, 1984; Tada & Siver, 1989; Railsback, 1993) used classification schemes that were
71 based on visual descriptions of macroscopic features of stylolites. These classification
72 schemes are however not quantitative and hard to compare since these studies focused on a
73 variety of different aspects of stylolite formation. Recent studies, however (Drummond &
74 Sexton, 1998, Karcz & Scholz, 2003) took a more quantitative approach using fractal

75 concepts to describe the stylolite roughness in a statistical sense. They could describe stylolite
76 roughness with a fractal scaling over several orders of magnitude, which means that their
77 roughness is not dominated by a certain wavelength.

78 Renard et al. (2004) and Schmittbuhl et al. (2004) went one step further and revealed that
79 bedding parallel stylolite surfaces show a self-affine scaling invariance with characteristic
80 Hurst exponents (also called roughness exponents). A self-affine rough surface is
81 characterized statistically by the fact that points along the surface separated by a distance Δx
82 from each other are typically distant in the direction transverse to the surface by $\Delta h = \Delta x^\alpha$,
83 where α is the roughness exponent. It was further noticed that two distinct scaling regimes
84 exist that were characterized by two different Hurst or roughness exponents separated by a
85 crossover-length (L), around the millimeter scale for the analyzed natural stylolites. Above
86 this crossover, all investigated stylolites exhibit a Hurst exponent of about 0.5 meaning that
87 they change relatively fast from being flat features on larger scale to being rough features on
88 the smaller scale. Below the crossover-length the Hurst exponent is about 1.0, which means
89 that the slopes, or aspect ratio $\Delta z / \Delta x$, stays more or less constant. Schmittbuhl et al. (2004)
90 and Renard et al. (2004) established from first principles of mechanics and chemistry a model
91 for stylolite growth under the form of a stochastic partial differential equation (called in this
92 case a generalized Langevin equation). This equation simulates the roughening of a stylolite
93 surface as a competition between stabilizing forces (that keep the surface flat), which are
94 controlled by long range elastic and local surface tension effects, and destabilizing forces (that
95 roughen the interface) that are induced by pinning effects of material heterogeneities. The
96 analytical solution of Schmittbuhl et al. (2004) reproduced the observed scaling behavior of
97 natural stylolites and demonstrated that the two scaling regimes (characterized by the two
98 different Hurst exponents) correspond to two thermodynamic regimes that are dominated by
99 either surface or elastic energies on small and large scales, respectively (Renard et al., 2004;
100 Schmittbuhl et al., 2004, Gratier et al., 2005). Based on the work of Schmittbuhl et al., (2004)

it was demonstrated for the first time by Ebner et al., (submitted) that the crossover-length of natural stylolites, which should be a function of the stress during stylolite growth, can be used to determine stress magnitudes and burial depth in sedimentary basins. The discrete numerical simulation technique of Koehn et al. (2007) enabled to study the dynamics of the roughening process through time revealing that the stylolite interface width w (defined in detail below) grows as a power law with time ($w \sim t^\beta$) with a growth exponent β of 0.5 in the surface energy dominated regime and a growth exponent of 0.8 in the elastic energy dominated regime. In addition the roughness growth may saturate so that the stylolites lose their memory for compaction or finite strain. It is important to notice that the roughness of simulated stylolites in this contribution is produced by heterogeneities in the material that pin the stylolitic interface due to slower dissolution rate constants, which are in competition with the surface and elastic energies which tend to flatten the surface (Koehn et al., 2007). Therefore the obvious question to ask is whether a variation of the quenched noise changes the scaling properties of the stylolitic interface?

Thus, in the present contribution we investigate the influence of different heterogeneities (namely the percentage of pinning particles, their pinning factor (defined below), and their size) on the scaling behavior, dynamic growth, and determined crossover length of simulated stylolites.

2. Setup numerical model

The numerical technique that we use to simulate stylolite roughening is based on a lattice-spring model coupled with a dissolution routine (Koehn et al., 2004, 2006, 2007). The model itself is embedded as a module in the “*Elle*” modeling-platform (Bons et al., 2008).

For computational reasons, to access large systems and analyze scaling laws over a large ratio between the system size and the resolution, we will consider situations spatially invariant along one of the directions tangential to the stylolite – and effectively treat systems with two

spatial dimensions. The initial configuration of this 2D model, as shown in Fig. 2a, contains a predefined flat interface that is considered to be filled with a confined fluid. Two blocks of particles are separated by a fluid pocket. Such an approximate configuration would be expected for example, for a fluid pocket embedded between two lowly permeable sedimentary layers. This model system represents two solids or rocks that are pressed together by inward moving top and bottom boundaries, whereas the side boundaries remain fixed (uniaxial strain). A quenched noise (denoted by darker particles in Fig 2a & b) is introduced by a change of the dissolution rate constant of a certain fraction of the particles (= pinning particles) and represents material heterogeneities initially present in the host rock of natural stylolites.

2.1. Theory

This section provides only a cursory review on the governing equations of the dissolution process used in the model, for a detailed description and implementation the reader is referred to Koehn et al., (2007) and Bons et al., (2008).

The pressure solution process is discretized in steps of dissolution of entire particles, following a linear rate law (Koehn et al., 2007 and references cited therein) according to

$$D = kV \left[1 - \exp \left(- \frac{\{\Delta\psi + \Delta\sigma_n\}V}{RT} \right) \right], \quad (1)$$

where D is the dissolution velocity of the interface (m s^{-1}), k a dissolution kinetics rate constant ($\text{mol m}^{-2} \text{s}^{-1}$), V the molecular volume of the solid ($\text{m}^3 \text{mol}^{-1}$), R the universal gas constant ($8.314 \text{ J mol}^{-1} \text{ }^\circ\text{K}^{-1}$), T the temperature ($^\circ\text{K}$), $\Delta\psi$ (Pa) the changes in Helmholtz free energy density (which accounts for the variations in elastic and surface energies) of the solid during dissolution of a solid element, and $\Delta\sigma_n$ (Pa) the differences between the average normal stress along the interface and the local normal stress at a specific location which is due to the repulsion of the solids (Koehn et al., 2007).

Surface energies (E^s) of particles are calculated from the local curvature of the interface around each particle, which can be expressed as

$$E^s = \frac{\gamma}{\rho}, \quad (2)$$

where γ is the surface free energy and ρ is the local radius of curvature of the interface. We consider a plane strain situation, i.e. an invariance along the third spatial dimension, so one radius along the 2D plane investigated entirely characterizes the curvature of the interface – the radius of curvature along the direction of invariance is infinite, and no surface energy is associated to this direction. The surface energies of individual particles are averaged over their neighbors to avoid artifacts from the discreteness of the model (for details see Koehn et al., 2007).

In the lattice spring model every particle (i) is connected to its neighbors (j) via a triangular linear elastic spring network. The elastic energy (E^{el}) of a single element is given by

$$E^{el} = \frac{1}{4} \sum_{(j)} \kappa (|x_i - x_j| - l)^2, \quad (3)$$

where the sum is over all neighbors (j), κ is a spring constant, l is the equilibrium distance between elements i and j .

2.2. Basic numerical step

The constitutive equations stated above are implemented as follows:

- Top and bottom walls are moved inwards simultaneously at a given time/deformation step.
- For every deformation step the rate law (Eq. 1) is used to calculate if individual particles at the interface can dissolve in the given time as soon as the two solids meet.
- When elements dissolve they are removed completely and the system can relax. Relaxation is accomplished by an over-relaxation algorithm that finds the new

equilibrium configuration for the lattice. Dissolution of particles can take place as long as the given time for the individual deformation step is not consumed.

- If the deformation time is used up or no particles can dissolve within the given time the system is stressed again by a deformation step.

2.3. Parameters, boundary conditions, limitations

The material parameters we use resemble those of a limestone, these values are analog to those found in reference works (e.g. Clark, 1966) and in line with values used for the analytical solutions of Renard et al. (2004) and Schmittbuhl et al. (2004): a molar volume of $0.00004 \text{ m}^3/\text{mol}$, a Young's Modulus of 80 GPa, a Poisson's ratio of 0.33 (this number is given by the triangular lattice configuration), a surface free energy of 0.27 J/m^2 , a temperature of 300 K and a dissolution rate constant of $0.0001 \text{ mol}/(\text{m}^2\text{-s})$. In addition, the displacement rate of the upper and lower boundaries are fixed at a constant value corresponding to velocities between 10^{-10} and $10^{-12} \text{ m s}^{-1}$.

The boundary condition can be seen as equivalent to a constant load boundary condition since the dissolution process is fast enough to relax the stresses that build up during a single deformation step. The sidewalls remain fixed during the model runs and there is no wrapping of particles in the x-direction of the model.

We use three basic model setups for which we systematically vary the heterogeneities in the structure. All boxes used have the same number of particles (400 particles) in the x-direction but three different particles-sizes were used 0.01 mm, 0.1 mm and 1mm, which corresponds to absolute box-sizes of 4, 40 and 400 mm.

To introduce the quenched noise in the simulations a pseudorandom algorithm is used to create a spatial Gaussian distribution of particles that dissolve slower (pinning particles). We varied three parameters of the quenched noise in this study: (i) number of pinning particles in a range from 1-20%, (ii) dissolution rate constant of pinning particles (from 0.1-0.99

normalized to the dissolution rate constant of the matrix, which is 1), which determines the pinning factor and (iii) the absolute size of the heterogeneities which varies with the particle size in the range of 0.01-1 mm.

3. Data analysis & results

The individual model runs are grouped with respect to the particle-size of the model and are termed *surface*, *intermediate* and *elastic class* according to the dominance of the energy regime during the roughening process. The *surface class* has a particle size of 0.01 mm (box size of 4 mm), the *intermediate class* a particle size of 0.1 mm (box size of 4 cm) and the *elastic class* a particle size of 1mm (box size of 40 cm). In nature the particles may resemble actual grains so that the grain size of the rock varies between the different classes.

Unfortunately we are restricted by the resolution of the numerical model to 400 particles in the x-direction, with this resolution a single simulation may run from 10 to 15 days on 4 cores of a recent workstation. Figure 3 shows the roughening of stylolites of these three classes, each with identical quenched noise (5 % pinning particles with half the dissolution rate of the matrix) to demonstrate the influence of the absolute box/particle size. The differences in the roughness and the roughness growth can easily be seen when individual steps of different classes are compared but also by following the growth of individual stylolite peaks along the time axis. In the *elastic class* individual peaks are growing very persistently whereas in the *surface class* the growth is often disrupted, due to dissolution of pinning particles as a result of high surface energies along pronounced peaks.

In the following sections we concentrate on the influence of the noise (amount and pinning factor of the noise particles) on 1) the roughness exponents, 2) the growth exponents and 3) the crossover length.

3.1. Roughness exponents

To quantitatively characterize the roughness of an individual 1D profile of a stylolite we used concepts from statistical physics (Barabasi & Stanley, 1995), which are briefly introduced in the first part of this section. The methods used here are the same as those of previous studies of natural stylolites (Renard et al., 2004; Schmittbuhl et al., 2004; Gratier et al., 2005; Ebner et al., submitted), which facilitates comparison.

The prerequisite for the application of these scaling methods is that the 1D signal of the numerical stylolite obeys a self-affine scaling invariance, which is given by (e.g. Barabasi & Stanley, 1995)

$$f(bx) \sim b^\alpha f(x), \quad (4)$$

where $f(x)$ is a single valued function and the power-law exponent α is called *roughness* or *Hurst* exponent and provides a quantitative measurement of the roughness of the signal. A self-affine function must be rescaled differently in x and y directions to obtain a scaling invariance i.e. horizontal rescaling of the form $x \rightarrow bx$, b being a dilation factor, has to be rescaled in the vertical direction by $y \rightarrow b^{-\alpha}y$ to obtain a scaling invariance. Different statistical methods can be used to evaluate the self-affine character of a signal and to determine the associated roughness exponent. We apply two independent methods in this contribution, the *Fourier method* and the *Average Wavelet Coefficient method*.

The *Fourier method* (e.g. Barabasi & Stanley, 1995; Schmittbuhl et al., 1995) is based on a Fourier transform of the original 1D signal (Fig. 4a). For every 1D signal (every deformation step) the Fourier power spectrum $P(k)$ i.e., the square of the modulus of the Fourier transform, was calculated as a function of the wave-number k . Plotting $P(k)$ as a function of k in log-log space reveals a linear trend for a self-affine function (Fig. 4b), and the slope is a function of the Hurst exponent through (Renard et al., 2004; Schmittbuhl et al., 2004):

$$P(k) \sim k^{-1-2\alpha} . \quad (5)$$

The *Average wavelet coefficient method* (*AWC*) was used as a second independent method to confirm the scaling results (Simonsen et al., 1998; Hansen et al., 2000). This method is again based on a decomposition of the 1D signal into wavelets, whose amplitude depends on scale and the position. The wavelet transform is defined after Simonsen et al. (1998) by

$$W_{a,b} = \frac{1}{\sqrt{a}} \int_{-\infty}^{\infty} \psi\left(\frac{x-b}{a}\right) f(x) dx, \quad (6)$$

where ψ is the wavelet basis (Daubechies wavelet of order 12) which is parameterized by a scale parameter a and a translation parameter b , and f is the single-valued original function. Finally the wavelet coefficients are averaged over the translation parameter b for every a to obtain the average wavelet coefficient $W(a)$. If the input signal is self-affine, the wavelet transform verifies that the average wavelet coefficient $W(a)$ scales as (Simonsen et al., 1998)

$$W(a) \sim a^{\alpha+1/2}. \quad (7)$$

Plotting the average wavelet coefficients as a function of the scale parameter a in log-log space (Fig. 4c), the slope of the linear regression through the data is again a function of the Hurst exponent.

Using these two statistical methods, we first study the dynamics of the roughness exponents through time during stylolite growth, and then concentrate on their stability with respect to variations of the noise. The roughness exponents increase relatively quickly in the course of a simulation run (Fig. 4d) and become stable after model step 3000 with only minor fluctuations. The *Fourier*- as well as the *AWC-method* show consistent evolutions and similar values of the roughness exponents. Averages of the plateau values reached (after step 3000) for individual model runs are used as a characteristic value for the roughness exponent for a specific setup. Error bars underline the standard deviation around this average (Fig. 5).

The *surface class* is characterized by consistently high values for the Hurst exponent i.e. $\alpha \sim 1$ -1.1, independent of the pinning factor (i.e. the dissolution rate constant k in Eq 1; cp Fig. 5a) or the amount of pinning particles (Fig. 5b).

The exponents only decrease when the pinning factor of particles is very low (dissolution rate constant > 0.9) and the rock becomes very homogeneous. The most stable roughness exponents for the surface class are reached in the range 0.1 to 0.8 for the pinning factor and 1 to 20 % of pinning particles.

The *elastic class* reveals lower Hurst exponents ($\alpha \sim 0.6$ to 0.9) than the *surface class*. If the pinning factor of particles is very strong (below the value 0.4 for the relative dissolution rate constant in Fig. 5a) stress concentrations are locally too high once two pinning particles meet and artifacts develop (usually anticracks that grow laterally emerge from these concentrations, hence modifying the surface topography) in the numerical model within the *elastic class*. Therefore we did not include values below 0.4 from elastic simulations in Fig. 5a. Generally the roughness exponents in the elastic class show stronger fluctuations than those of the surface class. They are relatively stable within a pinning factor range of 0.5 to 0.8 (Fig. 5a) and 1 to 20 % of pinning particles.

The *surface* and *elastic classes* correspond well to the two scaling regimes found in natural stylolites (Renard et al., 2004; Schmittbuhl et al., 2004) that are separated by a crossover-length at the millimeter scale. The roughness exponents of the surface class ($\alpha \sim 1.1$) are in very good agreement with analytical predictions and experimental observations (e.g. Gratier et al., 2005). The *elastic class* displays values for the Hurst exponent ($\alpha \sim 0.6$ - 0.9) that are higher than exponents from natural examples (Renard et al., 2004; Schmittbuhl et al., 2004; Ebner et al., submitted) or analytical predictions, which are generally around 0.5. These analytical predictions are usually based on linear approximations, which are strictly speaking valid as long as the surface morphology is not too developed. The present model does not present any such limitations, and the fully developed situation can thus present a different Hurst exponent from the initial one. The discrepancy with natural data may arise from the large particle (or grain) size that we use in the setup for the elastic class. In nature the grain size is much smaller and corresponds to the values that we use in the surface class.

3.2. Interface growth

In addition to the dynamic development of the roughness exponents the simulations allow us to study how fast the amplitude of the stylolite roughness grows through (model-) time or as a function of the finite strain. First we concentrate on different growth regimes of stylolites, the associated growth exponents and prefactors of scaling functions and then study the variation of these factors as a function of host-rock heterogeneities. In order to quantify the amplitude of the roughness we use the *interface width* (w) that is defined as the root-mean-square fluctuation of the height of the interface for a given time-step (Barabasi & Stanley, 1995)

$$w(L,t) \equiv \sqrt{\frac{1}{L} \sum_{i=1}^L [h(i,t) - \bar{h}(t)]^2}, \quad (8)$$

where w is the interface width as a function of system size L and time t , h is the height of point i on the interface at time t and \bar{h} the average height of the interface at time t given by

$$\bar{h}(t) \equiv \frac{1}{L} \sum_{i=1}^L h(i,t). \quad (9)$$

In our simulations, the system size L is defined as the number of elements in the x-direction, which is constant for all simulations, i. e. 400 particles. Roughening processes of interfaces in a wide range of fields have been demonstrated to follow a power law in time (e.g. Barabasi & Stanley, 1995) defined by a growth-exponent β (given by, $w(L,t) \sim t^\beta$). This initial phase of interface growth is usually followed by a second regime during which the interface width reaches a saturation value, w_{sat} , which is directly related to the system-size.

Both, the growth and saturation regimes can be seen in Fig. 6a for an experiment of the surface class with a characteristic growth exponent of $\beta \sim 0.5$. The arrow in Figure 6a marks the transition from the power law growth regime to the regime where the interface width saturates and stays constant. The *intermediate class* simulations show a similar growth exponent (around 0.5) but do not saturate in the given deformation time. The *elastic class*

(Fig. 6b) shows two successive growth regimes, the first being defined by a growth exponent of $\beta \sim 0.5$ up to a crossover interface width followed by a second regime with $\beta \sim 0.8$ without reaching the saturation regime. We suggest the following schematic growth regimes (Fig 6c) for stylolites: (i) growth in the surface energy dominated regime with an exponential growth defined by $\beta \sim 0.5$ followed by (ii) growth in the elastic energy dominated regime with $\beta \sim 0.8$ and finally reaching (iii) a saturation regime where the interface width stays constant.

This strict non-linearity of the interface growth in our simulations suggests that estimated amounts of compaction (here used as synonymous with volume/area loss due to pressure solution) from stylolite amplitude heights (e.g. Tada & Siever, 1989 and references cited therein) only capture a small part of the actual compaction. To cope with this problem, Koehn et al. (2007) demonstrated that the actual displacement can be expressed for the elastic or surface energy dominated growth regimes as a function of the interface width and the growth exponent (as long as the critical saturation time is not reached), given by

$$A \sim (w/l)^{1/\beta} l \quad (10)$$

where, A is the compaction displacement, w the interface width, β the growth exponent (for a certain class) and l the particle size. The slope of this function gives a prefactor for the scaling relation which should remain constant until the saturation time is reached. We call the slope of this relation here and in the subsequent sections *compaction prefactor* because it relates the interface width to the total compaction (Fig. 7). As soon as the saturation time is reached the relation does not hold any more and the function deviates from the linear trend (Fig. 7a). This effect can be observed in Figure 7a where the arrow marks the onset of the saturation of the interface width, compaction goes on but the interface width (x-axis) remains constant. However, if the saturation is not attained (Fig. 7b) the actual compaction can be calculated accurately from the interface width and the growth exponent using Eq. 9.

Looking at the growth exponent as a function of the quenched noise it can be clearly demonstrated that neither a variation in the pinning factor of particles (Fig. 8a) nor in the

amount of pinning particles (Fig. 8b) has a significant influence. The values for the growth exponent cluster around $\beta \sim 0.5$ for the *surface* and *intermediate classes* and around $\beta \sim 0.8$ for the *elastic class*. The compaction prefactors display a higher variability than the growth exponents but no systematic trend can be seen that relates this variation to the pinning factor or amount of pinning particles in the host-rock. Figures 8c & 8d show that there is no significant difference between the three classes of particle sizes used with values for the compaction prefactor in a range between ~ 12 and ~ 25 . As a third quantity we compare the maximum interface width normalized by the particle size (w_{max}) that develops during simulations with different heterogeneities (Fig. 8e & f). The largest interface widths are achieved in the *elastic class* with $w_{max} \sim 20$ in contrast to $w_{max} \sim 10$ reached in the *intermediate* and *surface class*. Hence the interface growth displays twice the displacement in the *elastic class* in the given simulation time due to the larger growth exponents than those of the *intermediate* or *surface class*, respectively. For the *surface*, *intermediate* and *elastic classes* the variation of the pinning factor of particles (Fig. 8e) have no considerable influence on the maximum interface width. However the amount of pinning particles has a significant influence on the *surface* and *intermediate class* (Fig. 8f). Both classes show an evident decrease in the maximum interface width with increasing amount of pinning particles. This trend cannot be observed in the *elastic class* (Fig. 8f).

We also tested the influence of the initial shape of the predefined interface separating the two blocks that are pressed together during an experimental run (compare Fig. 2a), which is flat in all the simulation data shown in the preceding sections. To investigate the dynamic roughness evolution of an already rough interface we arbitrarily choose a time/deformation step (t_n) of a simulation run (Fig 9a) and subtracted the topography (h) of this step from the subsequent time steps similar to $h(t) = h(t_n + m \cdot \Delta t) - h(t_n)$. This procedure allows to investigate the dynamic evolution of a rough interface but statistically evaluating the difference of the evolving roughness from time t_n onwards (Fig 9b). Departing from an already rough interface does not

change the scaling parameters (growth and roughness exponent) as depicted in Fig 9. Hence the model setup we choose in this work (i.e. a flat initial interface) can also account for complex initial topographies.

3.3. Crossover length scales

The crossover-length of the stylolite roughness that separates the surface energy regime from the elastic energy dominated regime is a function of the stress during stylolite growth and can be used as paleo-stress gauge. It is of fundamental importance to know if this crossover is constant when the heterogeneities in the host-rock vary. In the following section we want to explore how sensitive the crossover-length scale is on variations in the amount and strength of pinning particles. The *surface* and *elastic classes* of our simulations reveal roughness exponents that are characteristic for the two regimes where surface or elastic energy are dominant. The crossover-length between the two regimes can be found in the *intermediate class* of our simulations that reveals two different roughness exponents and hence the transition between the two scaling regimes (Fig. 10). The Fourier power spectrum of the 1D signal of a stylolite in the intermediate class (Fig. 10b) shows a change from a shallow to a steep slope, i.e. small and large roughness exponents on large and small scales, respectively. To avoid bias due to improper fitting of the crossover-length we used a nonlinear least square curve fitting algorithm in logarithmic space to model our scaling function (Ebner et al., submitted)

$$f(x) = (a_L x + m_L)(1 - w(x)) + (a_S x + m_S)w(x) \quad (11a)$$

and

$$w(x) = \frac{(\tanh(x + L) + 1)}{2}, \quad (11b)$$

where $a_{L,S}$ are the exponents of the scaling function for large and small scales, $m_{L,S}$ the corresponding intercepts with the ordinate and $w(x)$ the weighting function. During this

procedure the roughness exponents of our nonlinear model function were fixed according to the roughness exponents ($\alpha_S \sim 1.1$; $\alpha_L \sim 0.5$) reported from natural stylolites (Renard et al., 2004; Schmittbuhl et al., 2004, Ebner et al., submitted).

The crossover-lengths obtained for all experiments of the intermediate class are in a range of $L \sim 1.33 \pm 0.09$ mm, the crossover length usually develops simultaneously with the achievement of the plateau values (compare Fig. 4d). Neither of the quenched noise parameters varied in the experiments influences the crossover-length (Fig. 11).

The time evolution of the roughness presented in the previous section showed that the surface energy dominated regime is characterized by a growth exponent of 0.5 whereas the elastic energy dominated regime is characterized by a growth exponent of 0.8. Looking at the roughness growth in the elastic class one could still see the slow surface energy dominated growth in the beginning of the roughness evolution (Fig. 6). Therefore the growth exponents also show a transition (that we term the crossover interface width) between growth in the surface energy dominated regime and growth in the elastic energy dominated regime, similar to the two roughness exponents that are characteristic for these two regimes. The crossover interface width is very consistent for all experiments with $w \sim 1.23 \pm 0.04$ mm, independent of the quenched noise introduced in the system. We did not observe a crossover in the interface growth of the *intermediate* and *surface class* because the interface width in these classes is simply not large enough to reach the elastic growth regime. Due to the very good correlation between the magnitudes of the crossover-length and the crossover interface width (Fig. 11) we argue that both crossovers arise from the same process, namely the transition from a surface energy to an elastic energy dominated regime.

4. Discussion

In the following section we first discuss the influence of the quenched noise on the scaling parameters in our numerical simulations, deal with the relevance of the noise and compare the

results to natural stylolites. Secondly, we will focus on the crossover phenomena and their significance for the estimate of volume loss along a finite natural stylolite.

The influence of the exact nature of the heterogeneities (i.e. pinning factor, amount and size of pinning particles) on the scaling exponents can be directly investigated by a close examination of Figures 5, 8 and 11. Doing so, it is immediately evident that the pinning factor has the least influence on the roughness and growth exponents as well as on the crossover scaling (L and w_{cross}), with resulting values that are very consistent over wide ranges of the parameter space. The amount of pinning particles only shows an influence on the maximum interface width w_{max} , which decreases with increasing amount of heterogeneities (Fig. 8f). This fact is in good agreement with observations made on natural stylolites that stylolite amplitudes decrease with the amount of heterogeneities i.e. clay particles (e.g. Tada & Siever, 1989, and references cited therein). The scaling exponents themselves are independent of the amount of pinning particles. The biggest influence exists between the different experiment classes i.e. particle sizes, which reveal different roughness and growth exponents that are on the other hand very consistent within an individual class. But it has been demonstrated (Renard et al., 2004; Schmittbuhl et al., 2004, Koehn et al., 2007), that these differences arise from a transition from a surface energy dominated regime (where surface energies smoothen the interface) to an elastic energy dominated regime (where elastic energies smoothen the interface). The differences in the scaling exponents between the experimental classes are therefore not caused by the heterogeneities themselves but by the processes that govern the roughening on respective scales. We are thus convinced that the influence of the exact nature of the heterogeneities plays a minor role and that roughening is dominated by an inherent process that depends on the length-scale.

The quenched noise we introduced in the different simulations, i.e. changes in the dissolution rate constant that influence the dissolution velocity of a particle (cp. Eq. 1), is a simple chemical noise as pointed out by Koehn et al. (2007). We are aware that along natural

stylolitic interfaces the elastic parameters, surface energies, crystallographic orientations are very likely to change, in addition to chemical variations. However the effects of a change of these other parameters in our model also result in a change of the dissolution velocity. We therefore argue that for the developing structure it should make no difference what the exact nature of the noise is, since any particle with a slower dissolution velocity will pin the surface and therefore cause a roughening of the interface.

It was demonstrated that individual natural stylolites from different outcrop localities and lithologies, i.e. different host-rock compositions, reveal the same scaling behavior (e.g. Renard et al., 2004; Schmittbuhl et al., 2004, Ebner et al., submitted). Consequently the investigation of natural stylolites corroborates the fact that a common underlying mechanism for stylolite roughening can be assumed rather than a roughening that is dominated by the composition of the host rock.

Nevertheless we do not claim that knowledge of the exact nature or distribution of the material heterogeneities is unimportant. Brouste et al., (2007) have shown that a changing amount of heterogeneities might cause a stylolite to become a non-stationary signal with alternating wavy and flat portions along the interface. We have not investigated the effects of irregularly distributed noise since the heterogeneities are distributed equally in our model setup.

The roughness data of simulated stylolites presented in this study reveal two self-affine scaling regimes that are separated by a distinct crossover-length of $L \sim 1.3$ mm, which is well in line with investigations of natural stylolites (Renard et al., 2004; Schmittbuhl et al., 2004, Ebner et al., submitted). Additionally we have detected a crossover in the growth of the interface during which the initial growth exponent of $\beta \sim 0.5$ up to a crossover interface width of $w \sim 1.23$ mm is replaced by a growth regime with an exponent of $\beta \sim 0.8$. Due to the very good correlation between the magnitudes of the crossover-length and the crossover interface width we argue that both crossovers arise from the same process, namely the transitions from

a surface energy to an elastic energy dominated regime. The knowledge of the crossover length L which can be derived from the finite pattern of a natural stylolite with the above methods is thus equivalent to knowledge of the crossover interface width. This fact has important consequences regarding the assessment of the amount of total compaction of individual stylolites. Substituting the growth exponents and the compaction prefactors found for the two growth regimes in combination with the crossover-length, which separates the two growth regimes (cp. Fig. 10c), into Eq. 10 should allow an exact reconstruction of the amount of total compaction from finite pattern of a natural stylolites.

5. Conclusions

In the course of this contribution we evaluated the scaling properties of simulated stylolites, which facilitate a quantitative comparison with natural examples, exactly reproducing their scaling. We observed only minor correlation between the exact nature of the noise introduced in the model or the topography of the predefined interface and the scaling parameters investigated, concluding that inherent processes i.e. the transition from a surface to an elastic energy dominated regime control the roughening process.

Nevertheless the amount of heterogeneities has a negative effect on the maximum interface width (w_{max}) achieved during deformation revealing increasing interface width with decreasing amount of quenched noise. The absolute particle/noise size influences the roughness and growth exponents, which in turn is caused by the transition from a surface to an elastic energy dominated regime. Therefore it is important to know how large the noise or pinning particles are in natural systems. The transition from surface energy as the dominant stabilizing force of the interface to the dominance of elastic energies is the causation for the most significant scaling transitions: (i) the roughness is characterized by two distinct spatial scaling regimes on small and large length scales, respectively; (ii) the interface growth reveals two growth regimes with a growth exponent of $\beta \sim 0.5$ up to a crossover interface width that

coincides with the crossover-length L followed by a growth regime with an exponent of $\beta \sim 0.8$ that eventually saturates due to finite size effects; (iii) the crossover interface width w coincides with the crossover-length L and thus allows to accurately reconstruct the compaction history of finite stylolite patterns.

Our study corroborates the evidence that the simple mechanisms summarized above support analytical predictions and natural observations given in previous studies, and are a convincing causation for the formation of stylolite roughness. However a detailed study on the exact nature and distribution of quenched noise in the host rocks of natural stylolites would shed light on the origin and initiation of these complex structures.

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References

- Barabasi, A. L. & Stanley, H. E. 1995. *Fractal concepts in surface growth*. Cambridge University Press.
- Bathurst, R. G. C. 1987. Diagenetically Enhanced Bedding in Argillaceous Platform Limestones - Stratified Cementation and Selective Compaction. *Sedimentology* **34**(5), 749-778.
- Bons, P. D. D., Koehn, D. & Jessell, M. W. 2008. Microdynamics Simulation. In: *Lecture Notes in Earth Sciences* (edited by Friedmann, G. M. & Seilacher, A.) **106**. Springer, Berlin, 406.
- Brouste, A., Renard, F., Gratier, J. P. & Schmittbuhl, J. 2007. Variety of stylolites' morphologies and statistical characterization of the amount of heterogeneities in the rock. *Journal of Structural Geology* **29**(3), 422-434.
- Buxton, T. M. & Sibley, D. F. 1981. Pressure Solution Features in a Shallow Buried Limestone. *Journal of Sedimentary Petrology* **51**(1), 19-26.
- Clark, S. P. J. 1966. *Handbook of Physical Constants*. Geological Society of America, New York.
- Dunnington, H. V. 1954. Stylolite development post-dates rock induration. *Journal of sedimentary Petrology* **24**(1), 27-49.
- Ebner, M., Koehn, D., Toussaint, R., Renard, F. & Schmittbuhl, J. submitted. Scaling of natural stylolites and their use as stress-depth gauges. *Geophysical Research Letters*.

544 Fletcher, R. C. & Pollard, D. D. 1981. Anti-Crack Model for Pressure Solution Surfaces.
545 *Geology* **9**(9), 419-424.

546 Gratier, J. P., Muquet, L., Hassani, R. & Renard, F. 2005. Experimental microstylolites in
547 quartz and modeled application to natural stylolitic structures. *Journal of Structural*
548 *Geology* **27**(1), 89-100.

549 Guzzetta, G. 1984. Kinematics of Stylolite Formation and Physics of the Pressure-Solution
550 Process. *Tectonophysics* **101**(3-4), 383-394.

551 Hansen, A., Schmittbuhl, J., Batrouni, G. G. & de Oliveira, F. A. 2000. Normal stress
552 distribution of rough surfaces in contact. *Geophysical Research Letters* **27**(22), 3639-
553 3642.

554 Heald, M. T. 1955. Stylolites in Sandstones. *Journal of Geology* **63**(2), 101-114.

555 Katsman, R., Aharonov, E. & Scher, H. 2006. A numerical study on localized volume
556 reduction in elastic media: Some insights on the mechanics of anticracks. *Journal of*
557 *Geophysical Research-Solid Earth* **111**(B3), B03204.

558 Karcz, Z. & Scholz, C. H. 2003. The fractal geometry of some stylolites from the Calcare
559 Massiccio Formation, Italy. *Journal of Structural Geology* **25**(8), 1301-1316.

560 Koehn, D., Dysthe, D. K. & Jamtveit, B. 2004. Transient dissolution patterns on stressed
561 crystal surfaces. *Geochimica et Cosmochimica Acta* **68**(16), 3317-3325.

562 Koehn, D., Malthé-Sorensen, A. & Passchier, C. W. 2006. The structure of reactive grain-
563 boundaries under stress containing confined fluids. *Chemical Geology* **230**(3-4), 207-
564 219.

565 Koehn, D., Renard, F., Toussaint, R. & Passchier, C. W. 2007. Growth of stylolite teeth
566 patterns depending on normal stress and finite compaction. *Earth and Planetary*
567 *Science Letters* **257**(3-4), 582-595.

568 Koehn, D., Ebner, M., Renard, F., Toussaint, R. & Passchier C. W. in preparation. The use of
569 stylolites as paleo-stress-gauges.

570 Merino, E. 1992. Self-organization in stylolites. *American Scientist* **80**(5), 466.

571 Merino, E., Calas, A. & Fletcher, R. C. 2006. Genesis of self-organized zebra textures in
572 burial dolomites: Displacive veins, induced stress, and dolomitization. *Geologica Acta*
573 **4**(3), 383-393.

574 Park, W. C. & Schot, E. H. 1968. Stylolites: their nature and origin. *Journal of sedimentary*
575 *Petrology* **38**(1), 175-191.

576 Railsback, L. B. 1993. Lithologic Controls on Morphology of Pressure-Dissolution Surfaces
577 (Stylolites and Dissolution Seams) in Paleozoic Carbonate Rocks from the Mideastern
578 United-States. *Journal of Sedimentary Petrology* **63**(3), 513-522.

579 Railsback, L. B. 1998. Evaluation of spacing of stylolites and its implications for self-
580 organizations of pressure dissolution. *Journal of Sedimentary Research* **68**(1), 2-7.

581 Renard, F., Schmittbuhl, J., Gratier, J. P., Meakin, P. & Merino, E. 2004. Three-dimensional
582 roughness of stylolites in limestones. *Journal of Geophysical Research-Solid Earth*
583 **109**(B3), -.

584 Rutter, E. H. 1983. Pressure solution in nature, theory and experiment. *Journal of the*
585 *Geological Society of London* **140**(5), 725-740.

586 Schmittbuhl, J., Vilotte, J. P. & Roux, S. 1995. Reliability of Self-Affine Measurements.
587 *Physical Review E* **51**(1), 131-147.

588 Schmittbuhl, J., Renard, F., Gratier, J. P. & Toussaint, R. 2004. Roughness of stylolites:
589 Implications of 3D high resolution topography measurements. *Physical Review Letters*
590 **93**(23), -.

591 Simonsen, I., Hansen, A. & Nes, O. M. 1998. Determination of the Hurst exponent by use of
592 wavelet transforms. *Physical Review E* **58**(3), 2779-2787.

593 Stockdale, P. B. 1922. Stylolites: their nature and origin. *Indiana University Studies* **9**, 1-97.

Tada, R. & Siever, R. 1989. Pressure Solution during Diagenesis. *Annual Review of Earth and Planetary Sciences* **17**, 89-118.

Figure captions

Figure 1: Plane section of a bedding parallel stylolite in a jurassic limestone from Cirque de Navacelles (southern France). The rough interface is accentuated by a thin clay layer that is considered to be the residuum of the dissolved rock mass.

Figure 2: Simplified sketch of the setup of the numerical model (modified after Koehn et al., 2007). The top and bottom walls of the box are moved inwards simultaneously to stress the system and initiate dissolution along the interface. **a)** Initial configuration of the setup showing a flat interface (dashed line). **b)** Configuration after a certain amount of compaction. The interface (dashed line) has developed a distinct roughness, note that the heterogeneities (darker spheres) accumulate along the interface.

Figure 3: Growth of three stylolites with similar heterogeneities but different lengths and discretization resolution. Left panels show 3D plots of the stylolite growth from a flat interface until the end of the experiment. Right panels show 3 individual deformation steps (step 400, 2000 & 4000) corresponding to the solid lines in the 3D plot **a)** surface class (box size=0.4 cm) **b)** intermediate class (box size =4 cm) **c)** elastic class (box size =40 cm); Notice the disrupted growth of the surface class whereas the elastic class exhibits very continuous growth and pronounced peaks and teeth.

Figure 4: Roughness characterization of single deformation steps by means of the Hurst exponent. **a)** 1D profile (deformation step 3000) of the surface class with 5% pinning particles and a dissolution rate constant of $k=0.4$; **b)** Fourier power spectrum $P(k)$ of the signal from

Fig. 4a plotted as a function of the wavenumber k . Linear regression (solid line) is calculated from the binned spectra (crosses), the slope of the regression is a function of the Hurst exponent, see section 3.1.1. for detailed description. **c)** Wavelet spectra $W(a)$ plotted (crosses) as a function of the scaling parameter a (see average wavelet coefficient-method in section 3.1.1.). The slope of the linear regression (solid line) is again a function of the Hurst exponent. **d)** Evolution of the Hurst exponent in the course of an entire experiment for the Fourier and AWC methods. Stars indicate the Hurst exponent of the 1D signal shown in Fig. 4a for the two methods used. Notice the leveling off at a specific Hurst exponent (plateau value) which is characteristic for all simulations.

Figure 5: Composite plots showing the roughness data (Hurst exponents) for the surface and elastic classes. **a)** Hurst exponent plotted versus dissolution rate constants k with a fixed amount of pinning particles of 5% (where $k=1$ is the dissolution rate of the matrix). **b)** Hurst exponent plotted versus amount of pinning particles with a dissolution rate constant of $k=0.5$ for all experiments. The error bars correspond to the standard deviation given by the fluctuation around the plateau values, compare Fig 4d.

Figure 6: Interface growth depicted by means of the interface width (Eq. 8). **a)** Log-log plot of the interface width as a function of time in the surface class (5% pinning particles & dissolution rate constant of 0.4). Growth exponent $\beta=0.5$ is given by the slope of the linear regression (solid line). Notice the saturation point (indicated by arrow) i.e. interface width remains constant during ongoing deformation. **b)** Log-log plot of the interface width as a function of time in the elastic class (5% pinning particles & dissolution rate constant of 0.4). Notice the two successive growth regimes characterized by an initial growth exponent of $\beta=0.5$ up to a crossover width (indicated by arrow) followed by an exponent of $\beta=0.8$; no

saturation can be observed within the given simulation time. **c)** Proposed composite graph of the interface growth of simulated stylolites. Two successive growth regimes separated by a narrow crossover width that are dominated either by surface or elastic energies are followed by a saturation of the interface growth due correlation introduced by finite size effects.

Figure 7: Total compaction displacement expressed as a linear function (solid line) of the interface width and the compaction prefactor (compare Eq. 10). **a)** Surface class experiment (5% pinning particles & dissolution rate constant of $k=0.4$) demonstrates that the relationship is only valid as long as the interface is not saturated i.e. strong deviation from linear trend (compare Fig 6a). **b)** Elastic class experiment (5% pinning particles & dissolution rate constant of 0.4) reveals an accurate reproduction of the linear relationship (solid line) stated in Eq. 10. No saturation was observed in the given simulation time.

Figure 8: Cumulative data for the surface, intermediate and elastic classes for: **a)** Growth exponent versus dissolution rate constant (pinning particles fixed at 5%); **b)** Growth exponent versus amount of pinning particles (dissolution rate constant fixed at $k=0.5$); **c)** Compaction prefactor versus dissolution rate constant (pinning particles fixed at 5%); **d)** Compaction prefactor versus amount of pinning particles (dissolution rate constant fixed at $k=0.5$). Maximum interface width normalized by the particle size attained during experimental runs for the *surface*, *intermediate* and *elastic* class **e)** with changing dissolution rate constant (with 5% pinning particles); **f)** with changing amount of pinning particles (dissolution rate constant is fixed at 0.5).

Figure 9: The influence of the topography of the initial predefined interface on the dynamic roughness evolution. **a)** Rough initial interface used as starting point to evaluate the interface width evolution on top of this roughness for a *surface class* simulation with 5% pinning

particles and a pinning factor of 0.5 (for details see text). **b)** Comparison of the interface width evolution of an initially flat and rough (see Fig 9a) interface. Both interfaces show a similar evolution with growth exponents of 0.54 and 0.51 for the flat and rough initial interface respectively. Notice that both growth and the roughness exponents (not shown) are independent of the initial topography of the predefined interface.

Figure 10: Crossover phenomena in the roughness scaling and interface growth. **a)** 1D signal of an experiment from the intermediate class (5% pinning particles; dissolution rate constant $k=0.4$); **b)** Fourier power spectrum (inset) of the signal and binned spectra (crosses). A nonlinear model function (for explanation see text) used to minimize the original data (solid line) is used to accurately locate the position of the crossover length (triangle) $L=1.27$ mm; **c)** Log-log plot of the interface width versus time of the elastic class (5% pinning particles & dissolution rate constant $k=0.4$). Notice the two successive growth regimes characterized by an initial growth exponent of $\beta=0.5$ up to a crossover width $w=1.24$ mm followed by an exponent of $\beta=0.8$. Notice that both crossover scales correspond to the transition from a surface to an elastic energy dominated regime.

Figure 11: Crossover length L and crossover interface width w plotted as a function of the quenched noise. **a)** Crossover length (circles) calculated from the Fourier power spectrum (cp. Fig. 10b) and crossover interface width (diamonds) for simulations with different dissolution rate constants. **b)** Crossover length (circles) and crossover interface width (diamonds) for simulations with different amounts of pinning particles.

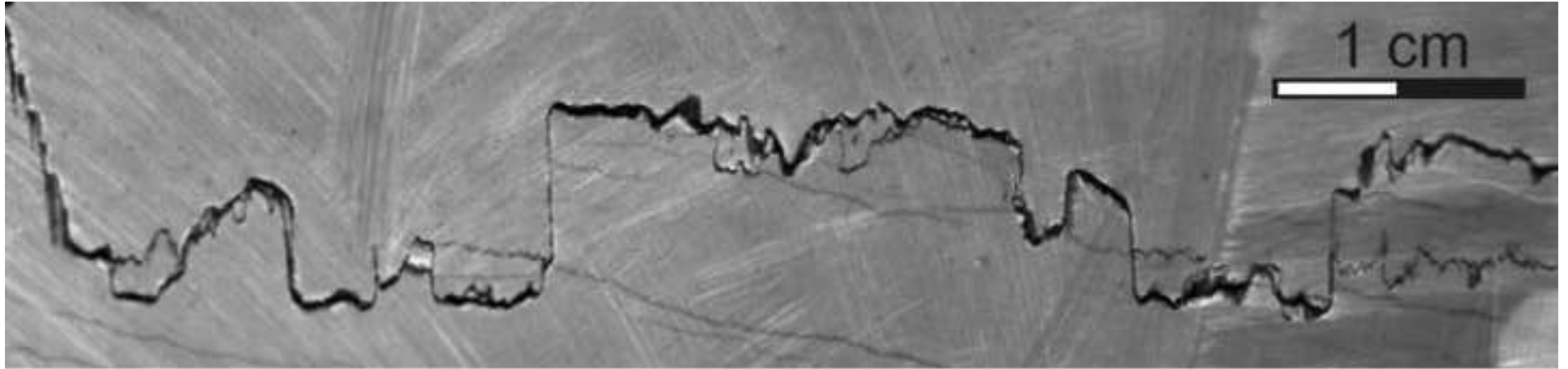


Figure 1

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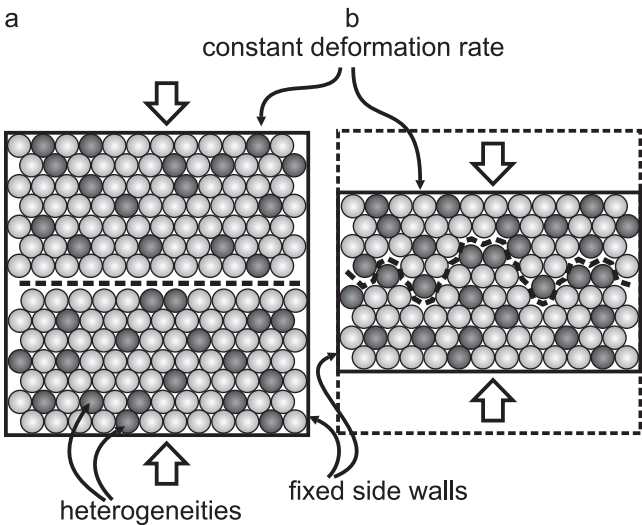


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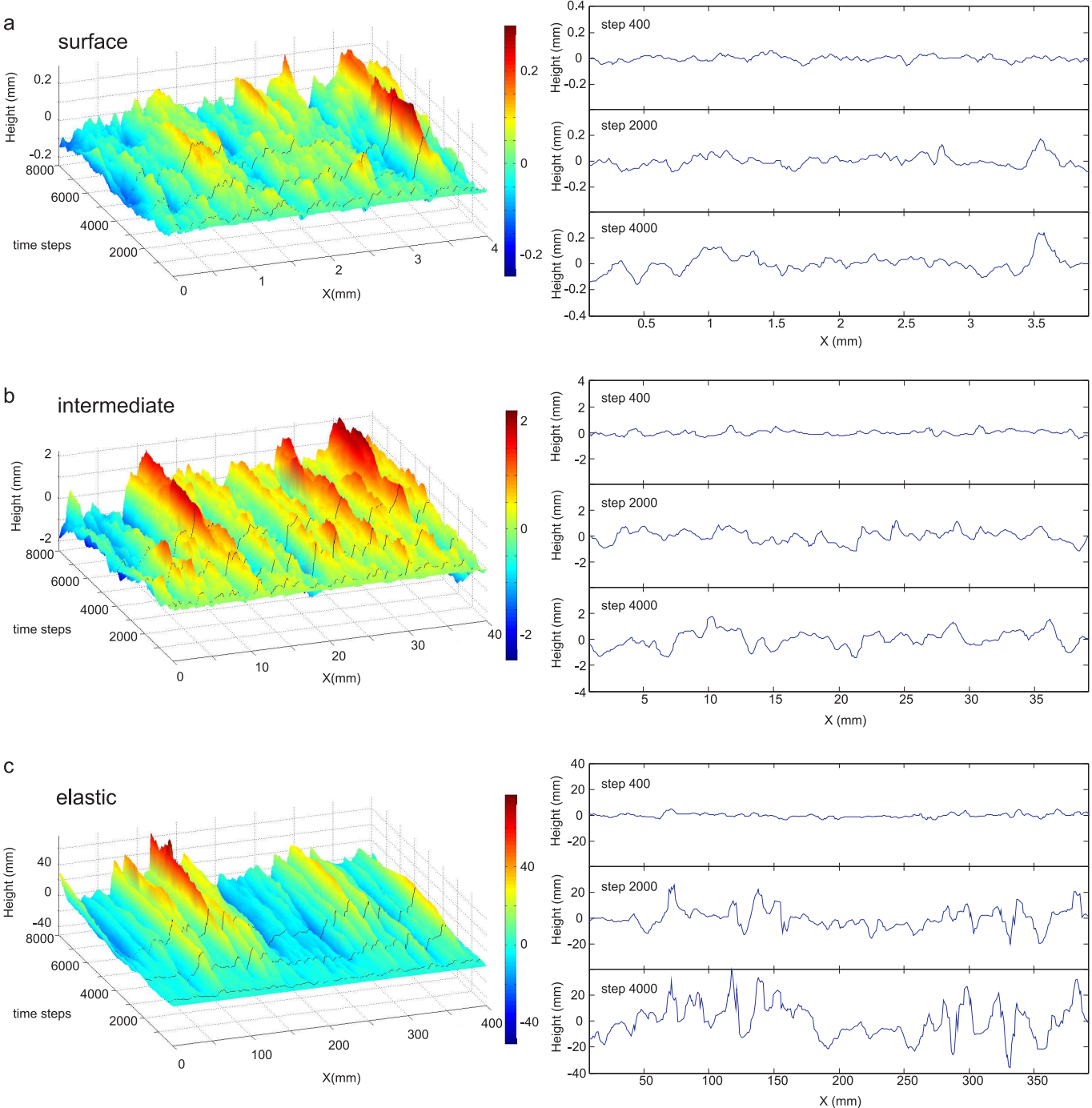


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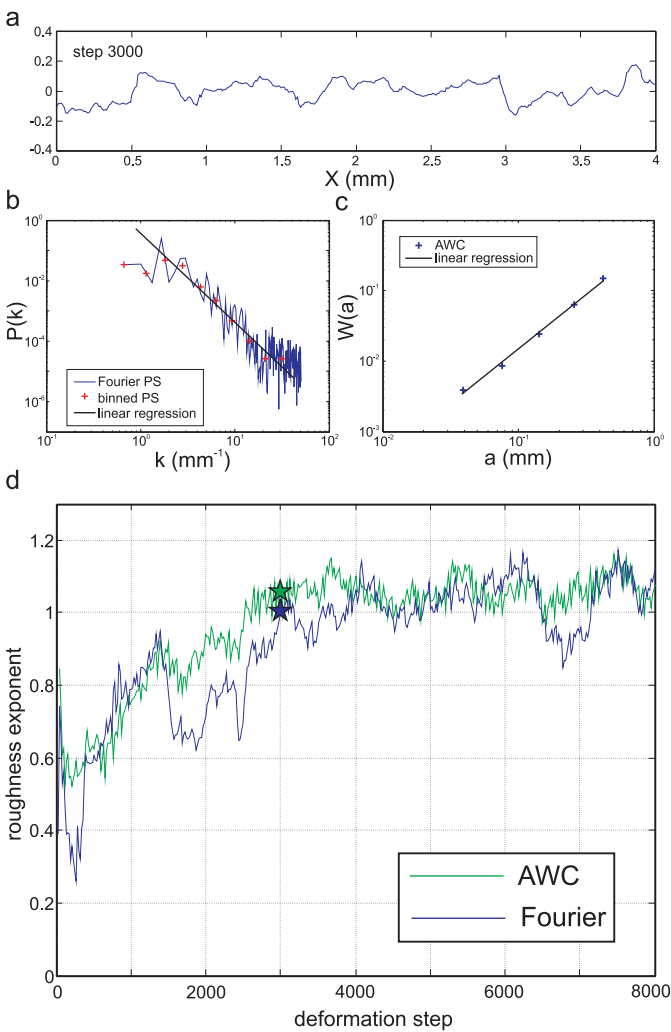


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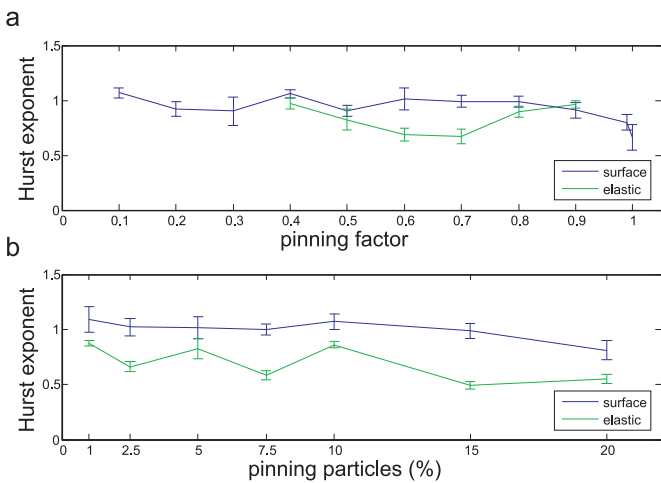


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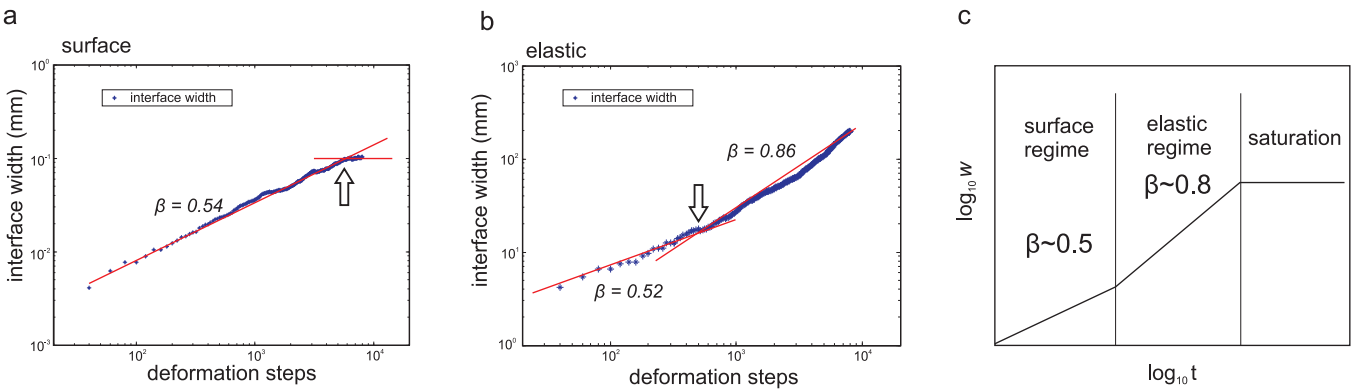


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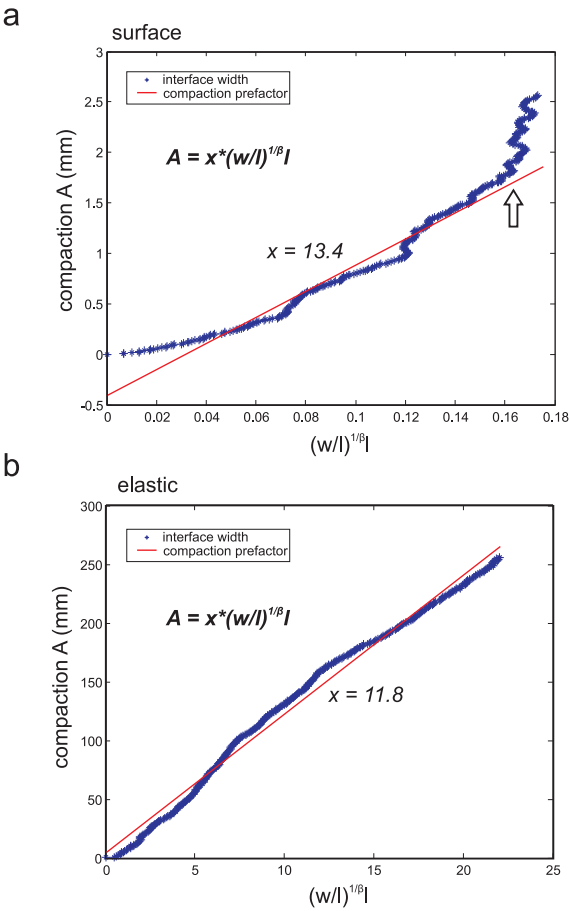


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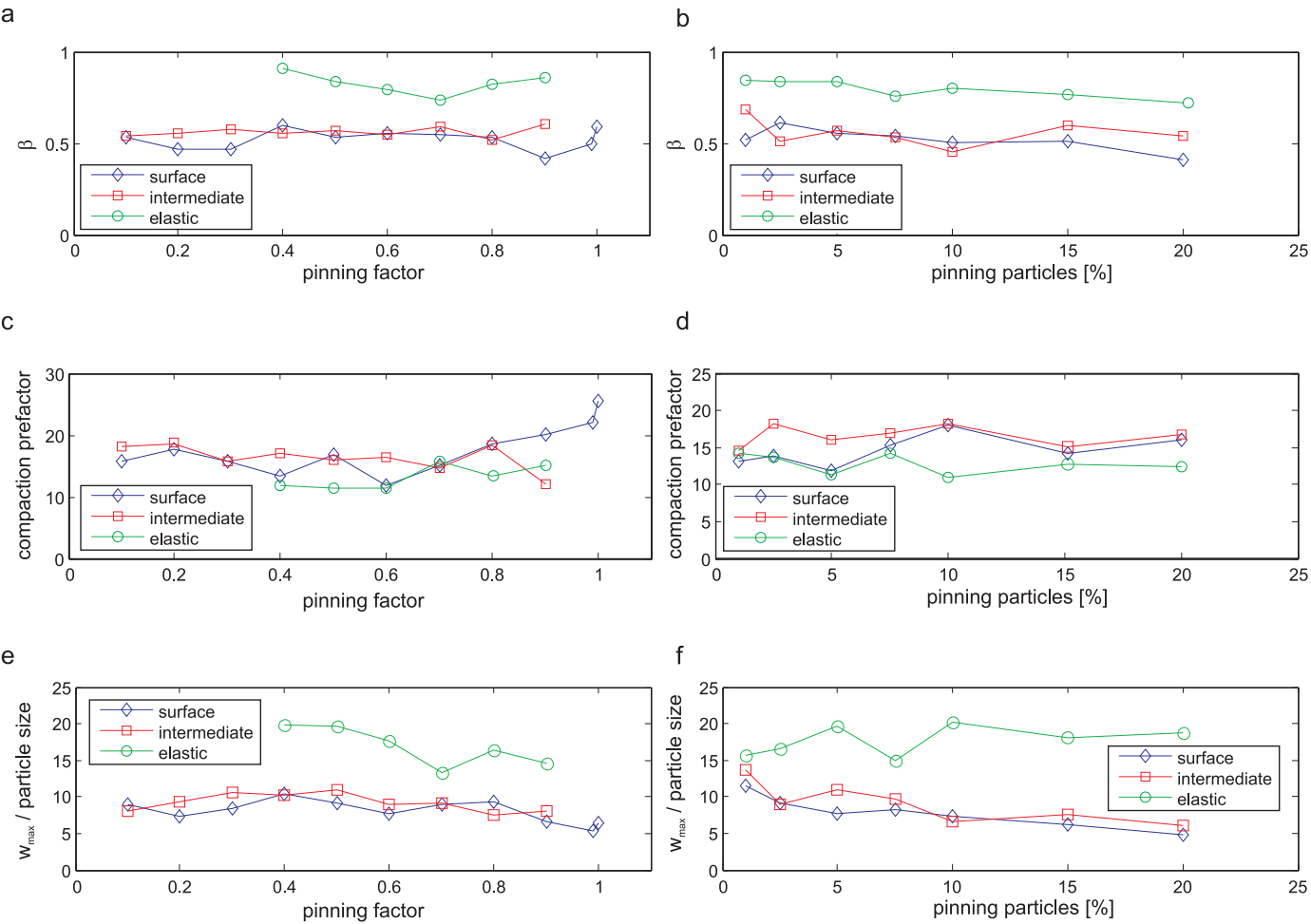


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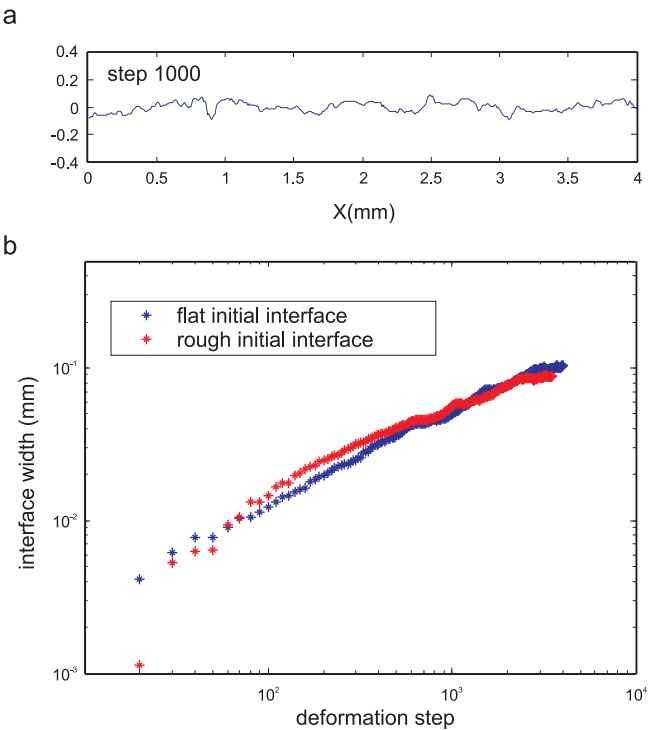


Figure 9

Figure 10

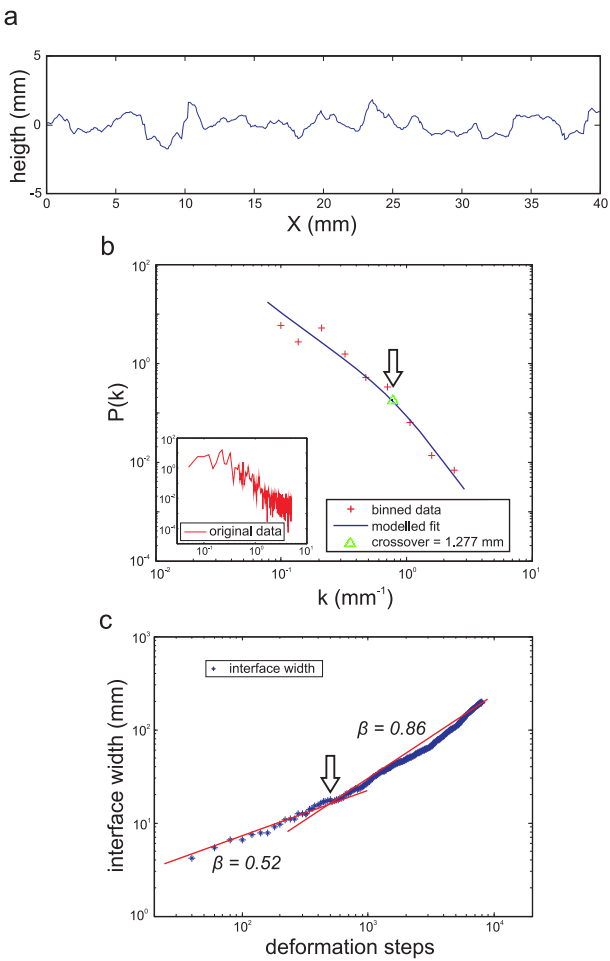


Figure 10

Figure 11

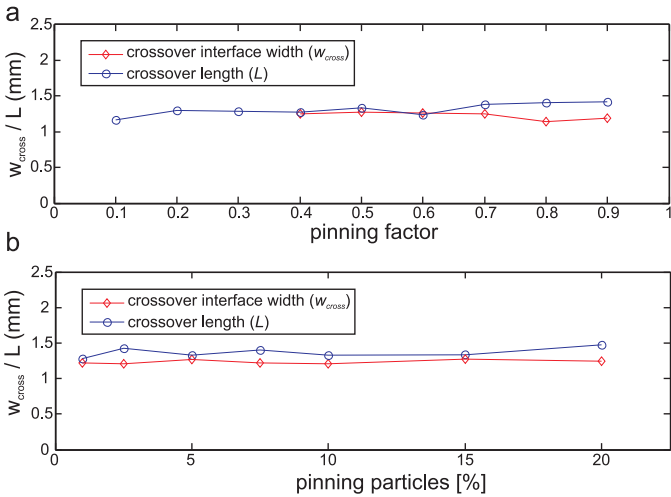


Figure 11