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► **To cite this version:**

Sébastien Gogo, Patrick Albéric, Fatima Laggoun-Défarge, Stéphane Binet, Axel Aurouet. Spatial and temporal variations of dissolved organic carbon and inorganic carbon concentrations and  $\delta^{13}\text{C}$  in a peatland-stream continuum: implications of peatland invasion by vascular plants. *Biogeosciences Discussions*, 2012, 9, pp.3515-3544. 10.5194/bgd-9-3515-2012 . insu-00681971

**HAL Id: insu-00681971**

**<https://insu.hal.science/insu-00681971v1>**

Submitted on 23 Mar 2012

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This discussion paper is/has been under review for the journal Biogeosciences (BG).  
Please refer to the corresponding final paper in BG if available.

# Spatial and temporal variations of dissolved organic carbon and inorganic carbon concentrations and $\delta^{13}\text{C}$ in a peatland-stream continuum: implications of peatland invasion by vascular plants

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Received: 5 March 2012 – Accepted: 6 March 2012 – Published: 20 March 2012

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Published by Copernicus Publications on behalf of the European Geosciences Union.

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## Abstract

Dissolved fluxes of C which are an important component of the global C budget of peatlands may be affected by global change in different ways. The evolution of peatland vegetation is an issue caused by the indirect effect of global change that still needs to be addressed. The aim of this study was to study the temporal and spatial variations in the quantity and the characteristics of dissolved C within in a peatland stream continuum, the peatland being invaded by *Molinia caerulea* and *Betula* spp.

The study highlighted the following key points:

- (i) Vegetation change tends to increase the DOC concentration of peat surface water.
- (ii) DOC exports were in the range of those found in the literature, and it was clearly established that peatland DOC can reach second order streams.
- (iii) Peat water was supersaturated in CO<sub>2</sub> especially in summer during the peak of microbial activity.
- (iv) the CO<sub>2</sub> critical layer at the surface of the soil may promote CO<sub>2</sub> excess build-up in the peat water. This critical layer may promote CO<sub>2</sub> export out of the peatland.

The effect of vegetation change on the DOC content of peat water deserves further attention as it may have a deep impact on downstream river water quality. In particular, it is necessary to clarify whether the concentration of C dynamics is affected by the invading species. More attention should be paid to hydrological processes (e.g. pumping capacity of the different plants) and the microbial activity in the rhizosphere of the invading plants. Also, the relation between the CO<sub>2</sub> critical zone at the surface of the soil and the CO<sub>2</sub> concentration in water should be studied in detail.

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## 1 Introduction

Peatland ecosystems are the major natural continental carbon (C) store as well as an important CH<sub>4</sub> source (Gorham, 1991). Their response to global change is uncertain, as positive or negative feedback can be triggered (Lashof et al., 1997). There is a need to better assess and predict their C source or sink functioning to be able ultimately to take the peatland contribution into account in the global climate model (Limpens et al., 2008). To achieve this goal, C budgets at the ecosystem level under different climatic conditions have to be monitored. Practically, each C input and output has to be quantified. Greenhouse gas (CO<sub>2</sub> and CH<sub>4</sub>) fluxes between peatland and atmosphere are quantitatively the most important fluxes and, thus, the most frequently measured. This is particularly true for CO<sub>2</sub> which can be monitored at high frequencies by eddy covariance techniques (Hendricks et al., 2010). These studies rarely measure fluxes of C through dissolved forms. However, dissolved organic and inorganic C (DOC and DIC respectively) fluxes from the peatland to downstream rivers can also represent an important part of the ecosystem C budget (Worrall et al., 2003; Dinsmore et al., 2011). It is only recently that investigation into DIC and DOC export has received more attention (Billett and Moore, 2008; Dinsmore et al., 2009; Dinsmore et al., 2011). Many studies on stream water quality acknowledged that peatlands were present in the watershed and had an impact on the rivers' dissolved C content. However, such studies did not always specifically investigate the characteristics of the water within the peatlands themselves (Telmer and Veizer, 1999; Dawson et al., 2004; Doctor et al., 2008). Thus, it seems relevant to study dissolved C export in a peatland-stream continuum, rather than in streams alone, to better understand both the peatland C budget and the river water quality.

Global changes, in addition to their direct effects, also modify peatland vegetation towards an increase in vascular plant contribution (Bubier et al., 2007; Limpens et al., 2008). The vascular plants affect soil organic matter dynamics (Mitchell et al., 2007; Gogo et al., 2011) and DOC production (Freeman et al., 2004; Gogo et al., 2012).

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## 2.2 Sample collection and conservation, climatic conditions

Water was sampled at 16 different sites: 4 within the peatland (1, 2, 3 and 4 in Fig. 1), 5 on the fringe of the peatland corresponding to water drained out of the peatland (6, 7, 8, 9 and 10 in Fig. 1), 4 in first order streams relative to the peatland (5, 11, 12 and 13 in Fig. 1) and 3 in second order streams relative to the peatland (14, 15 and 16 in Fig. 1). At each sampling site and at each date, pH, temperature and conductivity of the free water were measured with a multi-analyzer WTW 350i.

Within the peatland two sites were located in the western part: site 1 is dominated by open (without trees) vegetation and site 2 is dominated by *Molinia caerulea* and especially *Betula spp.* vegetation, which tend to “close” the system. Two sites were located in the eastern part of the peatland: site 3 has open vegetation and site 4 has a closed vegetation cover (by *Molinia caerulea* and *Betula spp.*). Site 5 corresponds to a drain going into the peatland. Water in this drain was only seen during the flooding event in May 2008. Site 6 collects water drained out of the peatland and corresponds to the limit between the peatland and the riparian zone surrounding the peatland. An intense sampling effort was made in the most intensive discharge area of the peatland: sites 7, 8, 9 and 10. Sites 11 and 12 correspond to the first order streams relative to the peatland and more precisely to the river La Guette. Site 13 can be assimilated to the outlet of the peatland as this stream receives the water of an artificial drain running parallel to the road D926 which collects water coming out of the sites 7, 8, 9 and 10. In a flooding event, such as during the first sampling date (May 2008), La Guette river and the drain along the road D926 can be connected. La Guette River and the peatland outlet flow into the second order stream, relative to the peatland, called the Barangeon. Site 14 corresponds to the upstream, site 15 is located after the confluence with La Guette river and site 16 is located after the confluence with the peatland outlet.

Triplicate samples at each sampling location were collected at four different dates: 25 May 2008, 1 September 2008, 2 December 2008 and 9 March 2009. Note that the first collection was made during a flooding event. For DIC measurements, water was

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collected with a syringe mounted with a filter (pore size: 0.45  $\mu\text{m}$ ), filtered and stored at field in a 7 ml headspace-free vial (Labco Exetainer). The tube was completely filled to obtain a convex water lens just above the edge of tube and closed with a screw cap mounted with a rubber septum. The flexibility of the rubber septum allowed water from the lens to be kept within the tube. This prevented any air bubble from being trapped in the tube, which avoided air  $\text{CO}_2$  equilibration with the liquid phase. For DOC measurements, water was collected with a syringe mounted with a filter (pore size: 0.45  $\mu\text{m}$ ), filtered and stored in situ in a 12 ml container (Falcon). Two drops of concentrated phosphoric acid were added to DOC samples within 24 h after collection. All the samples were brought back to the laboratory and kept in the cold room (4  $^\circ\text{C}$ ) until measurement.

### 2.3 Dissolved Inorganic and Organic Carbon concentration measurement

Dissolved Inorganic Carbon (DIC) concentrations were evaluated during the same run as DIC  $\delta^{13}\text{C}$  analysis (see below) using the Thermo LC-Isolink interface coupled to a Delta V isotopic ratio mass spectrometer (IRMS). Peak intensity was quantified versus  $\text{NaHCO}_3$  calibrating solutions.

Dissolved Organic Carbon (DOC) concentrations were determined using a Shimadzu<sup>®</sup> TOC 5000A total organic carbon analyzer, which is a high-temperature catalytic oxidation type of device with infrared detection, equipped with an ASI 5000A auto-sampler. DOC concentrations were measured in triplicate with a coefficient of variations better than 3 %.

### 2.4 $\delta^{13}\text{C}$ of Dissolved Inorganic and Organic Carbon analyses

Both DIC and DOC isotopic ratios were measured following the same procedure as that described by Albéric (2011) using the Thermo LC-Isolink IRMS interface. Each run started by three monitor gas pulses ( $\text{CO}_2$ ) for  $\delta^{13}\text{C}$  calculation. For DIC  $\delta^{13}\text{C}$  calibration, two standard solutions were prepared just before measurement by dissolving

carbonate salts with degassed MilliQ water: LSVEC (Li carbonate) international standard with a  $\delta^{13}\text{C}$  of  $-46.6\text{‰}$  and a working (Na bicarbonate) standard with a  $\delta^{13}\text{C}$  of  $-4.3\text{‰}$ . A calibration curve was calculated with the theoretical and measured values and the data were corrected using the equation obtained by linear regression. All the samples were injected manually using the 6 way valve of the LC-Isolink interface fitted with a 100  $\mu\text{l}$  loop.

For DOC  $\delta^{13}\text{C}$  calibration was corrected against simple standard compounds: IAEA-CH6 (sucrose), IAEA-C6 (sucrose), IAEA-C8 (oxalic acid), and USGS- 40 (glutamic acid) between  $-26.6$  and  $-10.45\text{‰}$ . DOC  $\delta^{13}\text{C}$  of macromolecular natural humic substances (Suwannee River fulvic acid IHSS-1S101F;  $-27.6\text{‰}$ ) was found to fit the calibration curves within  $0.3\text{‰}$  (Albéric, 2011). All the samples were acidified and He purged before autosampler injection and analysis (Albéric, 2011).

## 2.5 Calculation of excess dissolved $\text{CO}_2$ : $ep\text{CO}_2$

Excess partial pressure of  $\text{CO}_2$  ( $ep\text{CO}_2$ ) is the ratio of the partial pressure of  $\text{CO}_2$  ( $p\text{CO}_2$ ) in water to the atmospheric  $\text{CO}_2$  partial pressure (Neal, 1988; Doctor et al., 2008). Measurement of DIC, in situ pH and in situ temperature enable the calculation of  $p\text{CO}_2$  (Doctor et al., 2008; Steinmann et al., 2008):

$$p\text{CO}_2 = \frac{C_t}{\left(1 + \frac{K_1}{10^{-\text{pH}}} + \frac{K_1 K_2}{(10^{-\text{pH}})^2}\right) K_H} \quad (1)$$

Where  $C_t$  is the concentration of DIC,  $K_1$  and  $K_2$  is the first and second dissociation constants of  $\text{H}_2\text{CO}_3$  respectively, calculated from algebraic expressions (Telmer and Veizer, 1999), and  $K_H$  is the Henry constant for  $\text{CO}_2$ , corrected for temperature using the Van't Hoff equation. The atmospheric value of  $\text{CO}_2$  concentration used for the calculation was 378 ppm (Doctor et al., 2008).

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## 2.6 Atmospheric CO<sub>2</sub> concentration

In May 2011, atmospheric CO<sub>2</sub> concentration at different heights above three plots dominated by *Molinia caerulea* and Ericaceous shrubs was undertaken with a Vaisala Carbocap probe GMP 343.

## 2.7 Statistics

For each sampling plot and for each variable, the average of the three replicates collected was calculated. Then, averages were calculated and used to undertake statistical analyses. Paired *t*-tests were performed to assess differences between times within the peatland and the rivers. To test differences between vegetation types (open: plots 1 and 3; closed: plots 2 and 4), the multivariate Wilk test was performed. The significance level was set at 0.05 and *P* values between 0.05 and 0.1 were considered to indicate a trend. When the averages of the variables were compared between dates, different letters indicate significant differences (*P* < 0.05). These tests were performed with Statistica (2008). Model II regression (Legendre and Legendre, 1998) was undertaken with an Excel macro.

## 3 Results

### 3.1 Global trend of DIC and DOC concentrations and $\delta^{13}\text{C}$

In peatland water, both DIC and DOC concentrations were higher during summer than during the other seasons (Fig. 2a). Rivers tended to be richer in DIC and DOC during the flood event in May 2008 (Fig. 2b) than during the other seasons. DOC concentrations in peatland were always higher than in rivers (Fig. 2).

Within the peatland, waters collected in closed plots (invaded by *Molinia caerulea* and *Betula* spp.) tended to contain more DOC than in open plots (Table 2; Fig. 3a). The isotopic analysis showed that the DOC of the water collected in open plots tended

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to be enriched in  $^{13}\text{C}$  compared to the water collected in the closed plots (Table 1; Fig. 3b). For both DIC concentration and isotopic signature, no trends could be detected (Table 1).

### 3.2 Seasonal patterns of DOC $\delta^{13}\text{C}$ and concentrations

In spring (20 May 2008) and especially in summer (3 September 2008), the river DOC  $\delta^{13}\text{C}$  decreased from upstream to downstream (Fig. 4a and b). At these dates, but especially in September, the DOC in the Barangeon river (downstream) was nearly as depleted in  $^{13}\text{C}$  as the DOC in peatland (respectively  $-27.5/-28.5\text{‰}$  in spring, and  $-27.0/-28.0\text{‰}$  in summer), whereas upstream, la Guette river contained DOC relatively enriched in  $^{13}\text{C}$  compared to all the other sampling locations at the same date ( $-27\text{‰}$  in spring and  $-24.5\text{‰}$  in summer, Fig. 4a and b). In autumn (11 December 2008) and winter (3 March), the  $\delta^{13}\text{C}$  signatures of the DOC in rivers and peatland water were in the same range (Fig. 4c and d).

### 3.3 Patterns of $\delta^{13}\text{C}$ -DIC, excess $\rho\text{CO}_2$ , and $\text{CO}_2$ atmospheric concentration

During the flood event in May 2008, waters from all streams and from all the peatland sites had relatively similar characteristics in terms of  $\text{CO}_2$  saturation and isotopic signature (Fig. 5a). Other than during the flood event,  $\delta^{13}\text{C}$  of the DIC of the la Guette and Barangeon rivers were distinctly higher than those of the peatland (Fig. 5b, c, d). At these dates, the  $\delta^{13}\text{C}$  signatures of the DIC of the outlet water were always closer to those of the peatland than those of the rivers (Fig. 5b, c, d). Pooling the data of all dates revealed a negative correlation between the isotopic signature of the DIC and the excess of  $\text{CO}_2$  in both rivers and peatland waters (Fig. 6). The outlet data were in an intermediate position between rivers and peatland data. The slope in the peatland was steeper than in the river.  $\text{CO}_2$  acidifies the water and as it leaves the water, and the pH increases (Fig. 7). Air  $\text{CO}_2$  concentration was much higher just above the soil than in the air above (Fig. 8).

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### 3.4 Exported dissolved C fluxes

Exported dissolved C fluxes were calculated using the concentrations in the water in the periphery of the peatland ( $n = 4$ ) and in the outlet water ( $n = 1$ ). The flux values used in the calculations were those of Binet et al. (2012). DOC export ranged from 12.6 to 25.0  $\text{mg h}^{-1} \text{m}^{-2}$  and DIC export ranged from 6.03 to 14.9  $\text{mg h}^{-1} \text{m}^{-2}$  (Table 3). The only noteworthy differences between the two calculations were observed in September fluxes, with higher fluxes in the peatland data than the outlet data (Table 3).

## 4 Discussion

### 4.1 Peatland DOC export

La Guette peatland waters contained more DOC than the regional rivers. Thus, as already observed in other studies (e.g. Hemond, 1990; Kortelainen, 1999; Sasche et al., 2005), this peatland can be a potential source of DOC to downstream river systems. In this study, such an export was evidenced through the contrasted  $\delta^{13}\text{C}$  isotopic signature of the DOC in the La Guette river (upstream, location 11, red circle in Fig. 4) and the Barangeon river (downstream, location 16, blue circle in Fig. 4). In May 2008 and September 2008, La Guette upstream river contained DOC enriched in  $^{13}\text{C}$  compared to the DOC from the peatland sampling sites. Moreover, in the Barangeon downstream water, DOC was strongly depleted in  $^{13}\text{C}$  and the isotopic signature was similar to that assessed for the peatland DOC. Elder et al. (2000) found that peatland vegetation was much more depleted in  $^{13}\text{C}$  compared to aquatic plants (waterweed, flatstem pondweed) and algae. The DOC isotopic signature reported by Elder et al. (2000) actually reflected these differences in the sources, with a  $\text{DOC-}\delta^{13}\text{C}$  of  $-28.2$  and  $-25.0\%$  in the streams influenced by the peatland and the valley ecosystems composed of submerged and aquatic plants, respectively. In upstream La Guette river, the  $\delta^{13}\text{C}$  depleted DOC observed during spring and especially summer could reflect the influence

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of aquatic vegetation growing in the river at these times. Such an influence was not observed in the Barangeon, although aquatic plants were also present. Thus, it can be inferred that, at least in May and September 2008, part of the DOC of the Barangeon downstream river originated from the La Guette peatland. Furthermore, the fluxes of DOC estimated using the outlet data were the highest in December 2008 and March 2009. This suggests that the La Guette peatland is a substantial source of DOC to the downstream river system.

Peat solids do not exert a strong sorptive retention for DOC and, as water passes mainly through the very first centimetres of the peat profile, newly formed DOC through decomposition in this layer can be efficiently transferred to the streams (Tipping et al., 2010). In May 2008, because of heavy rainfall, all the sampling sites were flooded. DOC concentration in the downstream Barangeon in May was the highest recorded in the river system during all time course of this survey ( $14 \text{ mg l}^{-1}$ ). This suggests that large amounts of DOC were rinsed out of the peatland to the downstream river system.

The DOC fluxes calculated with the peatland periphery data were dramatically higher than the fluxes calculated with the outlet data in September only (Table 3). A drain collects the water going out of the peatland and the outlet sampling point was located after this drain. The results suggest that a significant amount of DOC was either mineralised or precipitated through a flocculation reaction in the drain.

Thus, when we consider the DOC isotopic signature and concentration results together, the peatland could represent a major source of DOC not only to the first order stream, but also to the larger second order stream. At the moment, DOC exports out of the peatland are in the medium range of values found in the literature, such as those from Dyson et al. (2011) in Finnish peatland (between  $21.7$  and  $42.4 \text{ mg d}^{-1} \text{ h}^{-1}$ ). Elder et al. (2000) estimated lower DOC fluxes (between  $3.2$  and  $15.1 \text{ mg d}^{-1} \text{ h}^{-1}$ ), whereas Clark et al. (2007) found much higher export fluxes (between  $46.6$  and  $74.0 \text{ mg d}^{-1} \text{ h}^{-1}$ ). Any increase in DOC export from the peatland to the downstream river system can have important implications (i) for the C budget estimation as punctual events can remove substantial amounts of C from the peatland, (ii) in terms of water quality as any

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perturbation that promotes an increased concentration of DOC in the peatland would thus increase the DOC content in the downstream river system.

## 4.2 Within peatland spatial variation of $\delta^{13}\text{C}$ isotopic signature and concentration of DOC

5 Water within the closed vegetation plots tended to have more DOC than waters from the open vegetation plots. Such a tendency could result from (i) a concentration through water uptake by invading plants, (ii) carbon dynamics promoting DOC accumulation, (iii) a combination of both processes. Untangling this issue is beyond the scope of this work. However it is worthy of interest as, in any case, a higher DOC concentration could  
10 originate from a modification in the functioning of the system caused by the invading species: perturbation of the water cycle and/or modification in the soil C dynamic.

Furthermore, DOC in open vegetation plots tended to be enriched in  $^{13}\text{C}$  (compared to the closed vegetation plots). In La Gulette peatland, the difference in vegetation is characterised by a switch in dominance from *Sphagnum* species in open vegetation  
15 plots to vascular plants (*Molinia caerulea* and *Betula* spp.) in the closed plots (Gogo et al., 2011). Compared to *Sphagnum* species, vascular vegetation litter contains less carbohydrates (Jia et al., 2008), a biochemical family known to have a higher  $\delta^{13}\text{C}$ . In contrast to *Sphagnum* species, the vascular plant litter is composed of true lignin, which tends to be depleted in  $^{13}\text{C}$  (Gleixner et al., 1993; Kracht and Gleixner, 2000;  
20 van Dongen et al., 2002). Thus, the slight difference in  $\delta^{13}\text{C}$  may reflect the influence of the vegetation change on the DOC composition: the decomposition of the lignin rich vascular litter could lead to  $^{13}\text{C}$  depleted DOC.

Although the composition and amount of DOC between open and closed vegetations tended to be different ( $P < 0.10$ ), these results suggest that further effort should be  
25 devoted to confirm or infirm these tendencies. This is crucial as, if closed vegetation does indeed increase DOC concentrations, the vegetation change would ultimately increase the DOC export to downstream rivers.

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### 4.3 Excess of CO<sub>2</sub> in peatland waters

Throughout the year, and especially in summer, water samples from the peatland (all plots taken together) were more supersaturated in CO<sub>2</sub> than those from the rivers (Fig. 5). The higher CO<sub>2</sub> supersaturation of peatland water during the summer certainly reflects higher microbial respiration activity during this season (Billett and Moore, 2008). Although in the same order of magnitude, the values of *ep*CO<sub>2</sub> for peatland waters reported in this work were higher than those recorded by Billett and Moore (2008) in a North-American peatland. However, our values fall within the range of the Ottawa river basin soils *p*CO<sub>2</sub> calculated by Telmer and Viezer (1999). They unfortunately did not detail the characteristics of the soils. As La Guelle peatland experiences warmer conditions than Canadian peatlands, this higher temperature in the French site would promote microbial activity, which may explain the high *ep*CO<sub>2</sub> measured in this study.

However, in addition to high microbial activity, another mechanism could contribute to generating such high excesses of dissolved CO<sub>2</sub>. We hypothesize that an air layer is maintained isolated from the atmospheric air by the vegetation and CO<sub>2</sub> produced by the microbial biomass can accumulate in this layer. Following Henry's law, as the CO<sub>2</sub> accumulates in the air just above the soil surface, CO<sub>2</sub> would also accumulate in the water. This hypothesis was supported by measurements of CO<sub>2</sub> concentrations in air above vegetation plots (dominated by *Molinia caerulea* and Ericaceous shrubs) and just above the soil (Fig. 8). These measurements, although done only once, clearly showed that such a CO<sub>2</sub> rich air layer just above the soil does exist. Further measurement has to be done (i) to highlight the seasonal pattern of this layer, (ii) to assess whether such a layer exists in *Sphagnum* carpets, and (iii) to assess whether this critical zone is specific to histosols.

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#### 4.4 Evidence of peatland-CO<sub>2</sub> export in downstream river system

Outgassing of excess CO<sub>2</sub> in water could induce an isotopic fractionation, associated with the equilibrium between CO<sub>2</sub> and HCO<sub>3</sub><sup>-</sup> (HCO<sub>3</sub><sup>-</sup> being much more enriched in <sup>13</sup>C than CO<sub>2</sub>). This could result in a linear relationship between the natural Log of *ep*CO<sub>2</sub> and δ<sup>13</sup>C-DIC in stream waters (Doctor et al., 2008). Such a fractionation is assumed to occur when the *p*CO<sub>2</sub> of the solution is greater than twice that of the atmosphere (Hendy, 1971). This situation actually occurs at all times in the streams and peatland waters, with some exceptions in December 2009. A relationship between *ep*CO<sub>2</sub> and δ<sup>13</sup>C-DIC was found in both peatland and river water (Fig. 6). In streams, for 1 natural Log unit decrease in *ep*CO<sub>2</sub>, δ<sup>13</sup>C-DIC increased to 2.32‰. This is in the range observed by Doctor et al. (2008). Outgassing of this nature is supported by the significant relationship between *ep*CO<sub>2</sub> and the pH. This relationship showed that in peatland water and direct tributaries, dissolved CO<sub>2</sub> is a major control of water pH. An increased flux of dissolved CO<sub>2</sub> from the peatland to the proximal downstream system would thus significantly affect the pH.

In the peatland water, the relationship between *ep*CO<sub>2</sub> and δ<sup>13</sup>C-DIC was also significant, but the slope was much steeper than in streams and the regression explained less variability than in the case of stream waters. The same could be said for the relationship between *ep*CO<sub>2</sub> and pH in river and peatland waters (Fig. 7). This probably reflects a significant production and accumulation of <sup>13</sup>C depleted CO<sub>2</sub> from microbial activity, with much less outgassing than in streams. This occurs especially during summer, when δ<sup>13</sup>C-DIC was the closest to the corresponding δ<sup>13</sup>C-DOC from which it proceeds through microbial assimilation and degradation, without much isotopic fractionation (Park and Epstein, 1961; Lin and Ehleringer, 1997). These results suggest that the CO<sub>2</sub> produced it is not readily outgassed, probably because of the concentration of CO<sub>2</sub> at the soil surface as described above, and accumulates in water.

Other than during the flood event, the DIC isotopic signature and the *ep*CO<sub>2</sub> in the outlet water (location 13) were in the same range as those found in the peatland waters.

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5 A few meters downstream, in the Barangeon river (location 16), the DIC isotopic signature decreased to values equivalent to those of the la Guette upstream river. The  $epCO_2$  also decreased markedly. DOC results showed that in May and September, significant amounts of water from the peatland flood into the Barangeon river. Thus, in  
10 terms of DIC, peatland water did not influence the Barangeon river. This suggests that a substantial amount of  $CO_2$  produced within the peatland is degassed just outside the peatland in the proximal downstream river. Such a process was particularly obvious in summer (Table 3), when  $CO_2$  produced by microbial activity accumulates in water more than during any other season. The outgassed flux through this pathway was estimated to be about  $10 \text{ mg d}^{-1} \text{ m}^{-2}$ .

## 5 Conclusions

15 The results of this study suggest that both DIC and DOC produced within the peatland are exported to the downstream hydrological system. This means that, for the La Guette peatland, a C budget based only on C exchange between the soil and the atmosphere would miss imported C fluxes. Dissolved exports were particularly high in summer for DIC and during the extreme flood event for DOC. Furthermore, the scope of the peatland influence on the downstream river system is greater in terms of DOC than DIC, because the latter is affected by the outgassing of  $CO_2$ . DIC export may be promoted by the occurrence of a  $CO_2$  critical zone just at the surface of the soil that  
20 prevents the  $CO_2$  from being degassed within the peatland. The results highlighted that further work should focus on (i) the effect of the peatland vegetation change on both the amount and characteristics of DOC, and (ii) the effect of different vegetation on the gas concentration gradient in air above the soil, as this gradient may be crucial in determining the amount of DIC exported out of the system.

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*Acknowledgements.* The authors would like to thank Emilie Rivoire and Nicolas Parou for their assistance during fieldwork. S. Gogo's salary was funded by the CG 45, the Spatioflux CPER research program (Resonat – FEDER) and the Labex VOLTAIRE.



The publication of this article is financed by CNRS-INSU.

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**Table 1.** Annual average ( $n = 4$ ) of pH and conductivity ( $\mu\text{S cm}^{-1}$ ) of the peatland and river waters.

Number Fig. 1	Sampling station	pH	Conductivity
1	West Open	$3.98 \pm 0.21$	$49.0 \pm 6.47$
2	West Close	$3.73 \pm 0.22$	$65.0 \pm 9.36$
3	East Open	$4.03 \pm 0.34$	$51.9 \pm 8.25$
4	East Close	$3.93 \pm 0.23$	$63.9 \pm 12.6$
5	Drain in 1	4.63	23.0
6	Drain out 1	$4.62 \pm 0.22$	$42.8 \pm 5.33$
7	Drain out 2	$4.64 \pm 0.34$	$40.2 \pm 6.25$
8	Drain out 3	$4.69 \pm 0.45$	$53.6 \pm 5.93$
9	Drain out 4	$4.16 \pm 0.35$	$54.3 \pm 10.4$
10	Drain out 5	$5.56 \pm 0.43$	$49.6 \pm 5.82$
11	La Guette upstream	$6.35 \pm 0.38$	$147 \pm 22.7$
12	La Guette downstream	$6.41 \pm 0.33$	$136 \pm 25.6$
13	Peatland outlet	$6.15 \pm 0.11$	$76.6 \pm 19.2$
14	Barangeon upstream	$7.24 \pm 0.65$	$199 \pm 28.0$
15	Barangeon downstream 1	$7.04 \pm 0.42$	$200 \pm 28.5$
16	Barangeon downstream 2	$6.91 \pm 0.24$	$186 \pm 35.5$

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**Table 2.** Results of the Wilk test on the vegetation effect on [DOC],  $\delta^{13}\text{C}$  DOC, [DIC] and  $\delta^{13}\text{C}$  DIC.

	Wilk test value	F	df	<i>P</i>
[DOC]	0.005	105	1	0.07
$\delta^{13}\text{C}$ DOC	0.003	147	1	0.06
[DIC]	0.47	0.57	1	0.68
$\delta^{13}\text{C}$ DIC	0.52	0.47	1	0.72

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**Table 3.** DOC and DIC export out of the peatland fluxes ( $\text{mg d}^{-1} \text{m}^{-2}$ ) estimated using the mean DOC concentrations of the water at the periphery of the peatland ( $n = 4$ ) and the DOC concentrations of the outlet water ( $n = 1$ ). The water flow values used in the calculations were those of Binet et al. (2012).  $\Delta$  corresponds to the difference between fluxes estimated with peatland periphery data and fluxes estimated with outlet data.

		Peatland periphery	Outlet	$\Delta$
DOC	May 2008	$12.6 \pm 0.68$	11.8	0.83
	September 2008	$25.0 \pm 4.67$	7.61	17.4
	December 2008	$19.8 \pm 6.03$	14.5	5.28
	March 2009	$20.8 \pm 5.00$	21.6	-0.76
DIC	May 2008	$6.03 \pm 0.91$	8.01	-1.98
	September 2008	$14.9 \pm 2.61$	4.4	10.5
	December 2008	$5.86 \pm 1.60$	5.98	-0.11
	March 2009	$6.86 \pm 0.83$	8.94	-2.09

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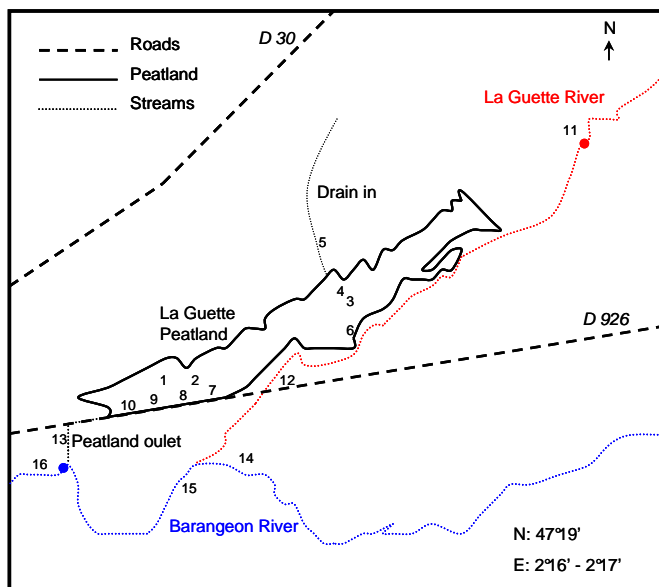
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**Fig. 1.** Map of the La Guette peatland and sampling points. The red dot corresponds to the upstream La Guette river (sample point taken as a reference for stream water characteristics) and the blue dot corresponds to the downstream Barangeon river (the river sampling point that may be influenced by peatland waters). The colour code is the same as Fig. 4.

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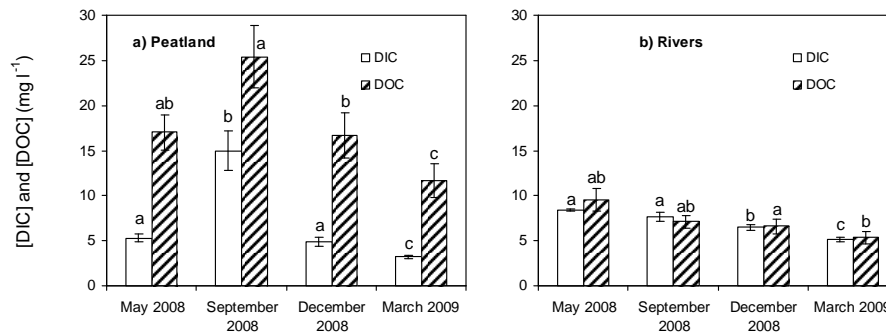
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**Fig. 2.** Average DIC and DOC concentrations at each sampling time in **(a)** the peatland water ( $n = 9$ ) and in **(b)** the streams ( $n = 4$  in May and September,  $n = 6$  in December and March). Different letters indicate significant differences ( $P > 0.05$ ).

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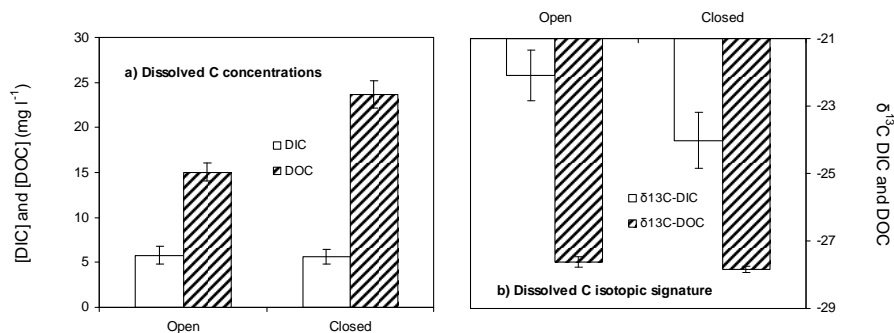
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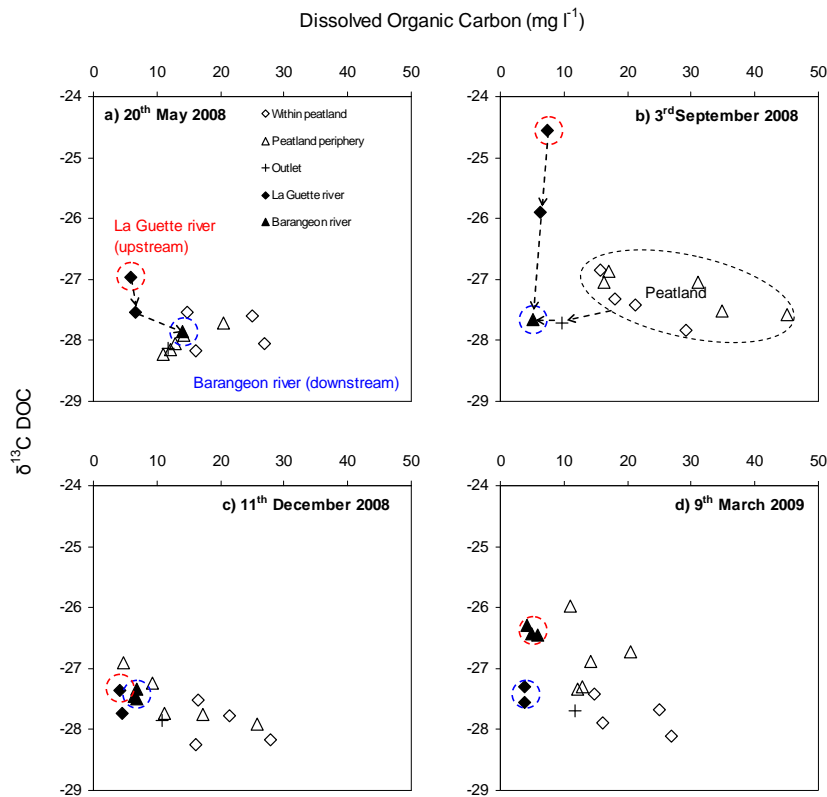
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**Fig. 3.** Average DIC and DOC concentrations **(a)** and  $\delta^{13}\text{C}$  **(b)** in open and closed vegetation ( $n = 8$ ).

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**Fig. 4.** Seasonal and spatial variation of the relationship of the DOC isotopic signature and concentration. The red circles correspond to the upstream La Guette river sample and the blue circles correspond to the downstream Barangeon river sample.

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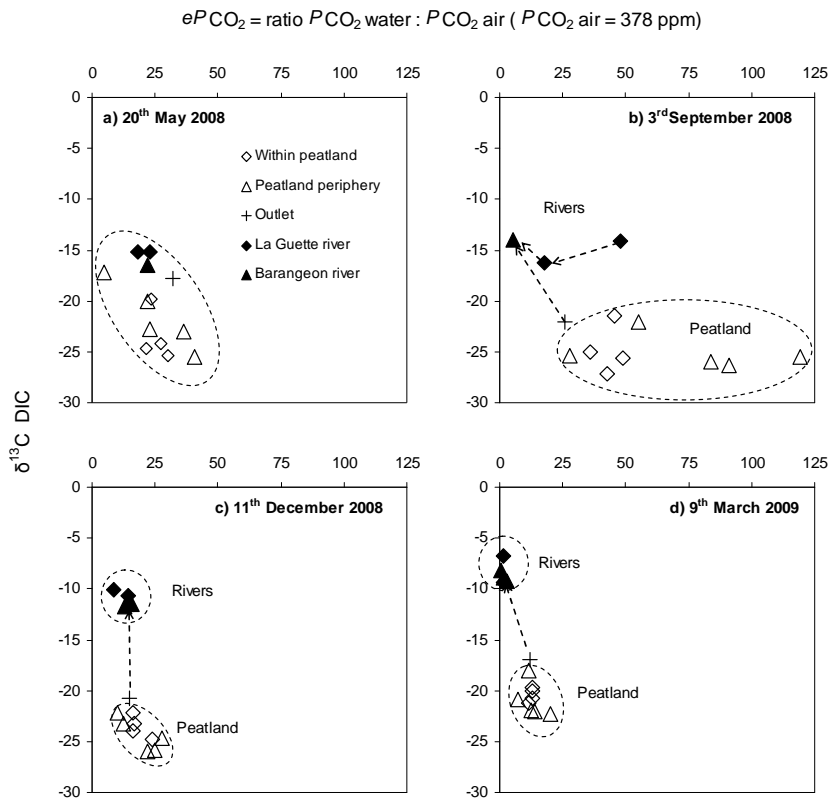


Fig. 5. Seasonal and spatial variation in the relationship of the DIC isotopic signature and  $ePCO_2$ .

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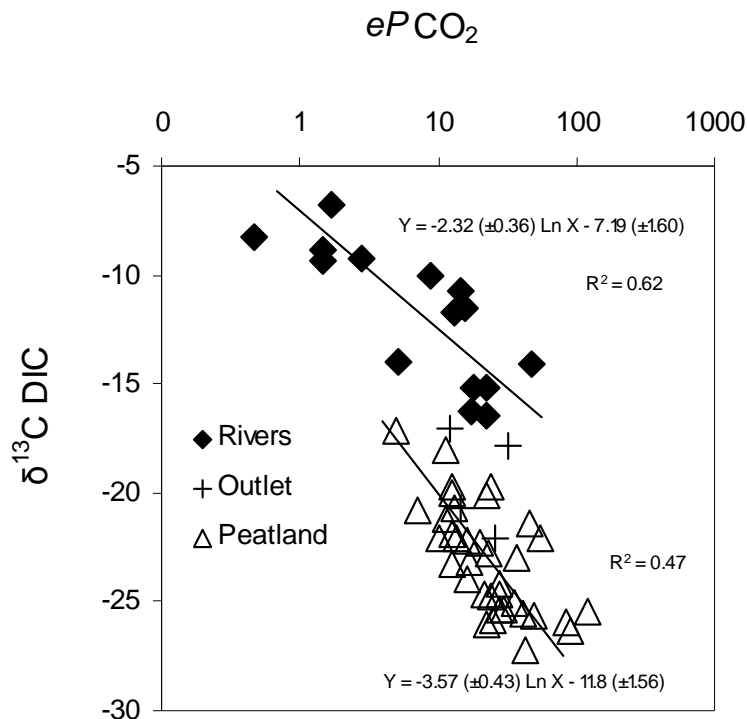
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**Fig. 6.** Relationship between the natural Log of the excess of dissolved CO<sub>2</sub> and the DIC isotopic signature (closed diamonds = river samples, cross = outlet, open triangle = peatland samples). Regression lines were calculated with the model II regression.

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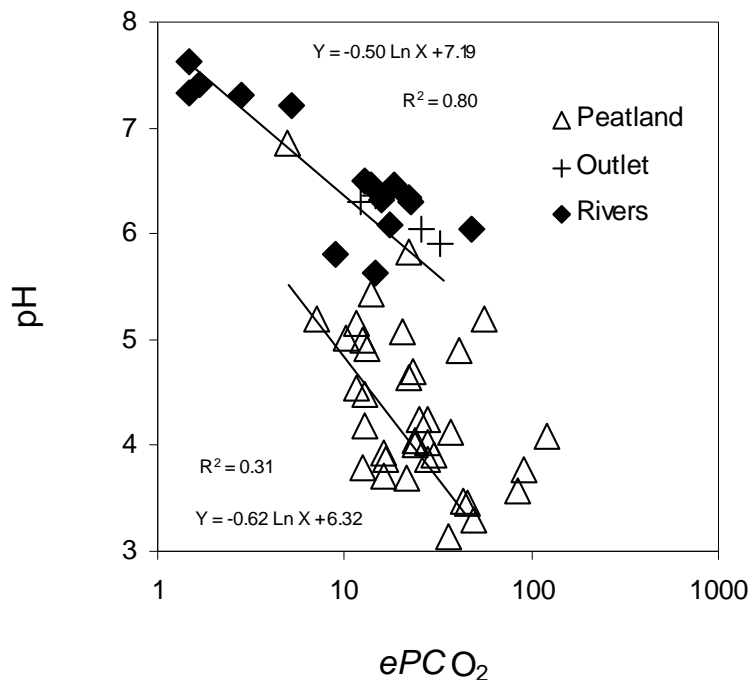
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**Fig. 7.** Relationship between the natural Log of the excess of dissolved CO<sub>2</sub> and the pH (closed diamonds = river samples, cross = outlet, open triangle = peatland samples). Regression lines were calculated with the model II regression.

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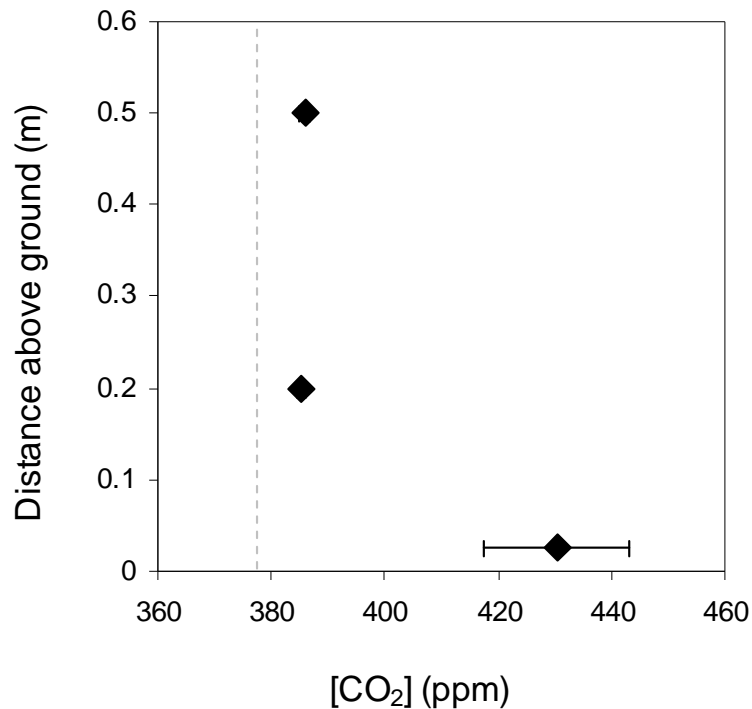
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**Fig. 8.** CO<sub>2</sub> partial pressure (in ppm) in air above the peatland ground in a *Molina caerulea*-Ericaceous shrub vegetation ( $n = 3$ ). The dotted line represent a mean atmospheric CO<sub>2</sub> concentration (378 ppm).

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