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Evidence for intense REE scavenging at cold seeps from the Niger Delta Margin

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Abstract :

For many trace elements, continental margins are the location of intense exchange processes between sediment and seawater, which control their distribution in the water column, but have yet to be fully understood. In this study, we have investigated the impact of fluid seepage at cold seeps on the marine cycle of neodymium. We determined dissolved and total dissolvable (TD) concentrations for REE and well-established tracers of fluid seepage (CH₄, TDFe, TDMn), and Nd isotopic compositions in seawater samples collected above cold seeps and a reference site (i.e. away from any fluid venting area) from the Niger Delta margin. We also analyzed cold seep authigenic phases and various core-top sediment fractions (pore water, detrital component, easily leachable phases, unclean d foraminifera) recovered near the hydrocast stations.

Methane, TDFe and TDMn concentrations clearly indicate active fluid venting at the studied seeps, with plumes rising up to about 100 m above the seafloor. Depth profiles show pronounced REE enrichments in the non-filtered samples (TD concentrations) within plumes, whereas filtered samples (dissolved concentrations) exhibit slight REE depletion in plumes relative to the overlying water column and display typical seawater REE patterns. These results suggest that the net flux of REE emitted into seawater at cold seeps is controlled by the presence of particulate phases, most probably Fe–Mn oxyhydroxides associated to resuspended sediments. At the reference site, however, our data reveal significant enrichment for dissolved REE in bottom waters, that clearly relates to diffusive benthic fluxes from surface sediments.

Neodymium isotopic ratios measured in the water column range from ϵ_{Nd} –15.7 to –10.4. Evidence that the ϵ_{Nd} values for Antarctic Intermediate waters (AAIW) differed from those reported for the same water mass at open ocean settings shows that sediment/water interactions take place in the Gulf of Guinea. At

each site, however, the bottom water ϵ_{Nd} signature generally differs from that for cold seep minerals, easily leachable sediment phases, and detrital fractions from local sediments, ruling out the possibility that seepage of methane-rich fluids and sediment dissolution act as a substantial source of dissolved Nd to seawater in the Gulf of Guinea. Taken together, our data hence suggest that coprecipitation of Fe–Mn oxyhydroxide phases in sub-surface sediments leads to quantitative scavenging of dissolved REE at cold seeps, preventing their emission into bottom waters. Most probably, it is likely that diffusion from suboxic surface sediments dominates the exchange processes affecting the marine Nd cycle at the Niger Delta margin.

Keywords : rare earth elements; neodymium isotopes; seawater; cold seeps; Fe–Mn oxyhydroxides; benthic fluxes

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38 **1 – Introduction**

39 **1.1. The sources of dissolved neodymium to the ocean**

40 The distribution of neodymium isotope ratios in seawater matches remarkably well global
41 ocean circulation patterns (see Frank, 2002; Goldstein and Hemming, 2003 for summaries).
42 On this basis, neodymium isotopes have been increasingly used as water-mass tracers in
43 marine authigenic precipitates and biogenic sediments to improve understanding of past ocean
44 circulation (e.g., Rutberg et al., 2000; Piotrowski et al., 2005, 2009; Scher and Martin, 2004;
45 Pucéat et al., 2005; van de Flierdt et al., 2006; Haley et al., 2008; Gutjahr et al., 2008;
46 Robinson and van de Flierdt, 2009). Despite significant interest in using Nd isotopes for
47 paleoceanographic studies, the way water masses acquire their Nd isotopic composition is not
48 fully understood yet. In fact, the sources of dissolved Nd and other rare earth elements (REE)
49 to the ocean are still being debated. Hydrothermal systems probably do not contribute much
50 to the dissolved Nd oceanic budget, because Nd and other rare earth elements emitted at vent
51 sites are efficiently scavenged by iron-rich plumes (e.g., Michard et al., 1983; German et al.,
52 1990; Halliday et al., 1992; Sherrel et al., 1999). To a first approximation, therefore,
53 dissolved neodymium in seawater is derived from continental inputs, with possible
54 contributions from rivers (e.g., Goldstein and Jacobsen, 1987; Elderfield et al., 1990;
55 Sholkovitz, 1995; Sholkovitz et al., 1999; Sholkovitz and Szymczak, 2000), dissolution of
56 settling particles (e.g., German and Elderfield, 1990; Greaves et al., 1994; Tachikawa et al.,
57 1999; Nozaki and Alibo, 2002; Bayon et al., 2004; Jacobson and Holmden, 2006), submarine
58 groundwater discharge (Johannesson and Burdige, 2007), and benthic fluxes (e.g., Elderfield
59 and Sholkovitz, 1987; Sholkovitz et al., 1992; Amakawa et al., 2000; Lacan and Jeandel,
60 2005; Arsouze et al., 2007, 2009). The Nd isotopic composition in ocean basins hence
61 globally reflects the age of surrounding terranes.

62

63 A major advance in the understanding of the marine Nd cycle has been the recognition over
64 recent years that the Nd isotopic signature of water masses could be modified along
65 continental and island margins, without any significant additional input of dissolved Nd
66 (Jeandel et al., 1998; Tachikawa et al., 1999; Lacan and Jeandel, 2001; Lacan and Jeandel,
67 2004a; Lacan and Jeandel, 2004b; Lacan and Jeandel, 2005; Andersson et al., 2008;
68 Amakawa et al., 2009). Lacan and Jeandel (2005) referred to this process as ‘boundary
69 exchange’, suggesting that ocean margins were an important component of the oceanic Nd
70 cycle. Recent modeling studies even proposed that exchange processes at margins could
71 represent the dominant source of dissolved Nd to the ocean (up to ~90%), far more important
72 than inputs from rivers and aeolian particles taken together (Arzouse et al., 2007; Arzouse et
73 al., 2009). However, despite the evidence that sediment/water interactions at margins play a
74 key role in the marine Nd geochemistry, the mechanisms of this exchange are not well
75 understood. Dedicated studies are now needed to better constrain the processes behind
76 boundary exchange. There are few sources of dissolved REE at margins that could possibly
77 impact the oceanic Nd budget at a global scale, which include dissolution of lithogenic
78 sediments, benthic fluxes from sub-surface sediments, and venting of methane-rich fluids
79 from reducing sediments. The goal of the present work is to assess, for the first time, the
80 potential importance of this latter source (i.e. fluid seepage) in the marine Nd cycle.

81

82 **1.2. Cold seeps and emission of methane-rich fluids on margins**

83 Venting of methane-rich fluids is a widespread phenomenon at ocean margins. Although
84 there are large uncertainties in estimating the mass of methane stored in marine sediment
85 (Judd et al., 2002), it is likely that methanogenesis occurs over at least 30% of the world’s
86 continental margins (Hovland and Judd, 1992). Seafloor expressions of focused fluid venting
87 are commonly referred to as cold seeps, which include a large range of geological structures

88 such as pockmarks, mud volcanoes, gas chimneys, and brine pools. In marine sediment,
89 methane is typically produced through microbial degradation of organic matter under anoxic
90 conditions, after a specific sequence of reactions, which greatly affect pore water chemistry
91 (e.g. Froelich et al., 1979; Thomson et al., 1993). In particular, organic matter degradation in
92 reducing sediments can lead to significant enrichments (from 10 to 1,000 times) in the rare
93 earth element contents of pore waters relative to seawater (Elderfield and Sholkovitz, 1987;
94 Haley et al., 2004).

95
96 Because methane, as a greenhouse gas, plays a key role in the Earth's climate, there have been
97 significant efforts to quantify methane fluxes at continental margins, and assess their
98 relevance to the global carbon budget (e.g., Judd et al., 2002; Milkov et al., 2003; Kopf, 2003;
99 Wallmann et al., 2006). In marked contrast, however, very little is known about trace element
100 biogeochemistry at cold seeps, and the impact of fluid seepage on ocean chemistry. A few
101 dedicated studies have focused on the geochemical cycling of barium at cold seeps from the
102 Peru and California margins (Torres et al., 1996; 2002; Castellini et al., 2006; McQuay et al.,
103 2008). These studies showed that emission of dissolved Ba at vent sites had significant local
104 impact on the marine Ba budget. Similarly, fluid seepage on continental margins could also
105 represent a potential source of dissolved Nd to the ocean, but to the best of our knowledge,
106 there has been no comparable work for the rare earth elements.

107
108 Here, we report dissolved and total dissolvable (TD) REE concentrations, Nd isotopic
109 compositions, and data for well-established tracers of fluid seepage (CH₄, TDFe, TDMn) for
110 seawater samples collected in the water column above deep-sea fluid-escape structures from
111 the Niger Delta (Gulf of Guinea, West African margin). In addition, we also present data for
112 a series of pore water samples, sub-surface sediments and associated authigenic precipitates

113 from the same area. Our data demonstrate that fluid seepage at cold seeps is not accompanied
114 by emission of dissolved REE into bottom waters, because Fe-oxyhydroxide co-precipitation
115 leads to quantitative REE scavenging at vent sites.

116

117 **2 – Regional setting**

118 **2.1. Studied sites**

119 The area investigated in this study is located on the Niger Delta, between 500 m and 1800 m
120 water depth (Fig. 1). A large number of seafloor structures related to fluid venting (i.e., mud
121 volcanoes, diapirs, pockmarks) were reported previously on the Niger Delta deep province
122 (Masclé et al., 1973; Brooks et al., 1994; Cohen and McClay, 1996; Bayon et al., 2007; Sultan
123 et al., 2010). In this study, all water and sediment samples were collected from three distinct
124 areas (Fig. 1). 1) A pockmark-rich area (water depth: ~ 550m; hereafter referred to as
125 Pockmark Field), characterized by the presence of large seafloor depressions with irregular
126 shapes (Fig. 2a). 2) A mud volcano (~ 680 m water depth; about 1km wide) situated on the
127 north flank of a dome, composed of two distinct volcanic cones with a mean elevation of
128 about 40m (Fig. 2b). The dome also exhibits a wide range of fluid venting structures related
129 to the presence of faults and/or gas hydrate reservoirs. 3) An area located at ~ 1780 m water
130 depth (Reference Site), where several submarine slope failures were reported previously
131 (Sultan et al., 2007), but which is not characterized by any active fluid seepage. In addition, a
132 few pore water samples were collected from sub-surface sediments recovered from other
133 active pockmarks of this Niger Delta area (see Bayon et al., 2007).

134

135 **2.2. Hydrography of the Gulf of Guinea**

136 The surface layer of the eastern tropical Atlantic is composed of warm and poorly salted
137 Tropical Surface Water (TSW; Fig. 3). The low salinity of TSW is largely attributable to

138 intense river runoff and rainfall in the Gulf of Guinea (Fig. 3A). At about 70 m depth, the
139 base of TSW is marked by a broad salinity maximum in the temperature range 17- 22 °C (Fig.
140 3B), which corresponds to Subtropical Underwater (STUW). Below STUW, the South
141 Atlantic Central Water masses (SACW) extend up to ~500 m depth, characterized by a nearly
142 linear temperature - salinity relationships (Fig. 3B). The water mass below SACW
143 corresponds to colder (~5°C) and fresher (salinity ~34.5) Antarctic Intermediate Water
144 (AAIW), centered at about 800 m depth. Finally, the deeper water masses in the study area
145 are dominated by southward-flowing North Atlantic Deep Water (NADW). Circulation
146 patterns of the upper water masses are quite complex in the Niger Delta area (Fig. 1). Surface
147 waters are transported eastward by the Guinea Current (GC), while circulation of water
148 masses below 100 m is dominated by the westward-flowing Northern South Equatorial
149 Current (nSEC).

150

151 **3 – Sampling and methods**

152 Samples were collected during previous expeditions to the Niger Delta aboard N/O *Atalante*
153 (NERIS project, 2004) and N/O *Pourquoi Pas?* (ERIG-3D project, 2008). All seawater
154 samples were collected during the ERIG-3D cruise using 8 l PVC-bottles mounted on a CTD-
155 rosette assembly. For determination of methane concentrations, aliquots of 125 ml were
156 collected in glass bulbs on board, and stored in a cold room to await transportation to the
157 laboratory in Brest. Then, methane was analysed using a chromatographic purge/trap
158 technique (Charlou and Donval, 1993; Charlou et al., 1998). For total dissolvable trace
159 element analyses (TDFe, TDMn, TDREE), a 60 ml aliquot of non filtered seawater was
160 transferred into acid-cleaned polyethylene bottles, and acidified to ~ pH 2 with ultra-pure
161 twice sub-boiled HNO₃. For dissolved REE studies, 250 ml seawater samples were filtered
162 through 0.45µm cellulose filters. After filtration, seawater samples were acidified to ~ pH 2

163 with ultra-pure twice sub-boiled HNO₃, prior to addition of Tm spike. The REE were then
164 extracted from the filtered samples by ferric-hydroxide co-precipitation, after addition of NH₄
165 (Bayon et al., 2011). For Nd isotope measurements, between ~ 5 and 20 l of seawater were
166 filtered and acidified to ~ pH 2. At Brest, Nd and other REE were then pre-concentrated by
167 ferric-hydroxide co-precipitation, followed by purification using cation exchange (AG 50W-
168 X8) and Ln-resin columns.

169

170 A series of sub-surface sediment samples recovered by either piston or gravity coring near the
171 hydrocast stations were also analysed in this study (see core location in Fig. 2). Pore waters
172 were extracted from bulk sediments on board by centrifugation and filtered (0.45 µm)
173 immediately. Upon availability, ~ 3 to 40 ml aliquots of pore waters were processed for
174 determination of REE concentrations, following the procedure described above (Bayon et al.,
175 2011). Uncleaned foraminifera fractions (mainly *Globigerinoides ruber*) were analysed to
176 gain additional information on the ε_{Nd} signature of bottom waters, as demonstrated recently
177 by Roberts et al. (2010). Foraminifera fractions were cleaned in ultrasonic bath with ultra
178 pure water, prior to dissolution using dilute HNO₃ acid. The terrigenous fraction of every
179 studied sediment sample was also analysed after removal of carbonate and Fe-oxyhydroxide
180 phases from the bulk sediment (Bayon et al., 2002). In addition, the fine-grained (< 45 µm)
181 fraction of each core-top sediment sample was leached (room T°C, 24 h) using ultra-pure
182 dilute (0.05% v/v) nitric solution (i.e. easily leachable fraction), in order to assess the
183 potential contribution of sediment dissolution to the non-filtered seawater samples. The acid
184 strenght of this dilute nitric solution exactly matches that of the solution (pH ~ 2) in which
185 non-filtered seawater samples were stored prior to analysis. Then, dilute HNO₃ leachates
186 were filtered (0.45 µm) before processing for REE and Nd isotope measurements. Finally,
187 two methane-derived carbonate concretions and authigenic gypsum were hand-picked from

188 the Pockmark Field and Mud Volcano sediments, cleaned using ultra pure water, and
189 analysed to provide direct information on the pore water ϵ_{Nd} signature at the studied cold seep
190 sites.

191

192 All measurements were made at the Pôle Spectrométrie Océan (PSO), Brest. Rare earth
193 element, Fe and Mn concentrations were measured with an ELEMENT 2 ICP-SFMS. The
194 REE were analysed with the low resolution mode to enhance sensitivity, but were corrected
195 for interferences following the procedure of Bayon et al. (2009). Rare earth element
196 concentrations were calculated using the Tm addition method (Barrat et al., 1996; Bayon et
197 al., 2009). Details on the applicability of this method for determining REE abundances in
198 seawater are given elsewhere (Bayon et al., 2011; Freslon et al., 2011). For Fe and Mn, the
199 ELEMENT2 was operated in medium resolution mode. Procedural blanks for Fe and Mn
200 corresponded to ~ 1.5 nM and ~ 0.5 nM, respectively. Neodymium isotopic ratios were
201 determined by Neptune MC-ICP-MS. Analysis of the JNdi-1 standard during the analytical
202 session gave $^{143}Nd/^{144}Nd$ of 0.512115 ± 0.000011 (2 s.d., n=12), which corresponds in epsilon
203 notation (DePaolo and Wasserburg, 1976) to an ϵ_{Nd} value of $= -10.16 \pm 0.21$. Total
204 procedural blanks were less than 1 ng for Nd, which represented less than 6% of the mass of
205 Nd in the measured fraction of seawater samples.

206

207 **4 – Results and Discussion**

208

209 **4.1. Depth profiles at the active venting sites: Pockmark Field and Mud Volcano sites**

210 The bottom water samples at the Pockmark Field (CTD-08) and Mud Volcano (CTD-06)
211 stations exhibit CH_4 values with concentrations up to ~ 2000 nl/l and ~ 330 nl/l respectively,
212 much higher than background seawater values (in the range ~ 15 and 40 nl/l), which clearly

213 indicate active fluid venting (Table 1). At these two sites, methane plumes rise up to about
214 100 m above the seafloor (Fig. 4). Iron and manganese oxyhydroxide precipitation typically
215 occurs above methane seeps at submarine hydrothermal systems (e.g. German et al., 1990),
216 but also on continental margins (Charlou et al., 2004), when Fe-rich vent fluids mixed with
217 high pH (pH ~ 8) and oxygen-rich bottom waters. Similarly, here, the plumes at Pockmark
218 Field and Mud Volcano also exhibit distinctive anomalies for both TDMn (up to 8 nmol/L;
219 Table 1) and TDFe (up to ~ 50 nmol/L), which could hence reflect the presence of Fe-Mn
220 oxyhydroxide particulates. Alternatively, the occurrence of Fe and Mn anomalies in non-
221 filtered seawater samples could also indicate partial dissolution of suspended particles
222 entrained within the plumes. High levels of TDREE concentrations were also determined in
223 the methane plumes at both sites (Table 1), with depth profiles for TDNd closely resembling
224 those for TDFe (Fig. 4). Interestingly, while TDNd concentrations are significantly enriched
225 in the methane plumes (up to 62 pmol/kg), the dissolved Nd contents for the same samples are
226 much lower (around 22 pmol/kg; Table 2, Fig. 5), and do not exhibit any significant
227 enrichment relative to the overlying water column (Fig. 5). Overall, these results suggest that
228 venting of methane-rich fluids at cold seeps does not lead to significant emission of dissolved
229 REE into the water column. Our data show however that fluid venting is accompanied by a
230 flux of REE associated with iron-rich particulate phases, which could indicate either co-
231 precipitation of Fe-Mn oxyhydroxides in bottom waters or re-suspension of local sediments.

232

233 **4.2. Evidence for benthic fluxes at the Reference Site**

234 At the Reference Site, the bottommost water sample (CTD03-B1) exhibits higher TDFe and
235 TDMn concentrations (i.e. the highest TDMn value measured during the course of this study;
236 10.3 nmol/kg, Table 1) than the overlying water column (Fig. 4). In contrast with the
237 Pockmark Field and Mud Volcano sites, however, these anomalies are most probably due to

238 diffusion from surface sediments at this location, rather than to active fluid venting.
239 Similarly, the same bottommost sample also displays the highest dissolved REE
240 concentrations determined in this study (e.g. [Nd] ~ 27.9 pmol/kg, Table 1). Taken together,
241 these results could suggest that benthic fluxes at the Reference Site (i.e. away from any active
242 fluid venting area) lead to diffusive emission of REE into bottom waters. Evidence that both
243 TDNd and dissolved Nd exhibit similar concentrations at this site, as shown in Fig. 5,
244 indicates however the absence of any significant Fe-oxyhydroxide co-precipitation or
245 sediment resuspension at this station.

246

247 **4.3. Deciphering REE provenance in the filtered and non-filtered seawater samples**

248 To gain further constraints on the origin of REE sources in the methane plumes, we
249 considered shale-normalised REE patterns for both non-filtered (TD data) and filtered
250 (dissolved concentrations) samples (Fig. 6), and compared them to data for pore waters (Table
251 3) and easily leachable sediment fractions (Table 4). For clarity, only REE patterns for
252 selected seawater samples from the Pockmark Field and the Reference Site are shown in Fig.
253 6, but note that similar conclusions could be also drawn using samples from the Mud
254 Volcano.

255

256 At the Pockmark Field, filtered samples collected from within the plume (sample CTD08-B1
257 to –B8; Table 2) all display very similar seawater-like REE patterns (Fig. 6A), characterized
258 by a pronounced negative Ce-anomaly and progressively increasing shale-normalized values
259 from the light- (LREE) to the heavy-REE (HREE). These patterns are very similar to those
260 determined for the seawater samples at the Reference hydrocast station (Fig. 6B). In marked
261 contrast, non-filtered samples collected at the same water depths at the Pockmark Field show

262 a larger range of REE patterns, with variable Ce-anomalies and various mid-REE (MREE)
263 over LREE enrichments (Fig. 6A). In comparison, pore waters from sub-surface sediments at
264 the Mud Volcano and other active venting sites of the Niger Delta area exhibit REE
265 concentrations about one order of magnitude higher than those for seawater samples (Table
266 3). These pore water samples display shale-normalized patterns characterized by a positive
267 Ce-anomaly and a MREE enrichment relative to LREE and HREE (Fig. 6A). This MREE-
268 bulge type pattern is a typical feature of anoxic pore waters in marine sediments, interpreted
269 as the consequence of the reduction of sedimentary Fe-oxyhydroxide phases during early
270 diagenesis (Haley et al., 2004). Although we did not analyse any pore water sample from the
271 Pockmark Field area, the carbonate concretion collected from core ER-CS-38 also displays a
272 similar REE pattern (Table 4, pattern not shown here), which suggests that it was formed
273 from fluids having similar REE signature (Rongemaille et al., 2011). Here, the evidence that
274 filtered samples collected from within the methane plume exhibit seawater-like REE patterns
275 that are well distinct from those for local sub-surface pore waters provides strong support that
276 active venting at these seeps does not represent any substantial source of dissolved REE to
277 bottom waters.

278

279 As discussed earlier, one explanation accounting for the TDFe, TDMn and TDREE anomalies
280 at both the Pockmark Field and Mud Volcano was that they were due to co-precipitation of
281 Fe-Mn oxyhydroxide phases in bottom waters above venting sites. If this was the case,
282 however, one would expect the filtered samples collected from the methane plumes to have
283 inherited, at least partly, the distinctive REE signature of pore waters. Instead, it is more
284 likely that these anomalies indicate partial dissolution of resuspended particles entrained
285 within the methane plumes. This hypothesis can be demonstrated using simple mass balance
286 calculations with REE concentrations for typical bottom water (e.g., filtered sample CTD08-

287 B3) and the easily leachable fractions of core-top sediments (Table 4). Comparatively, the
288 REE concentrations determined in the dilute nitric leachates are much higher (i.e., about a
289 factor 10^8) than seawater values. The leaching experiments with dilute HNO_3 led to the
290 extraction of about 20 wt% of the initial mass of sediment. This implies that the presence of
291 even a very small amount of suspended particles in any of our non-filtered seawater samples
292 could have a significant impact on its REE composition. In Fig. 6C, we show that the REE
293 patterns for non-filtered samples from within the plume at the Pockmark Field can be
294 generated by partial dissolution of sediments in seawater samples having total suspended
295 matter loadings (TSM) of about 0.1 to 0.2 mg/l. For comparison, this range of values is
296 similar to the maxima TSM concentrations measured in hydrothermal plumes (i.e., up to 90
297 $\mu\text{g/l}$; Trocine and Trefry, 1988; Feely et al., 1994).

298

299 Interestingly, the shale-normalized REE patterns of these easily leachable sediment fractions
300 are also characterized by a strong positive Ce-anomaly and a marked MREE enrichment (see
301 the theoretical pattern for a non-filtered seawater sample with TSM of 1 mg/l, Fig. 6C). As
302 mentioned earlier, this pattern is typical of sedimentary Fe-oxyhydroxide phases (e.g.,
303 Bayon et al., 2004). This suggests that a significant fraction of the REE extracted from our
304 core-top sediments during our leaching experiments is derived from the dissolution of REE-
305 rich Fe-oxyhydroxide phases. By analogy, it is very likely that the measured TDFe, TDMn
306 and TDREE anomalies determined in the non-filtered seawater samples above venting sites
307 were due to dissolution of Fe-Mn oxyhydroxide phases associated to suspended particles
308 within the plumes. Taking a further logical step, we propose that Fe-oxyhydroxide co-
309 precipitation in the near surface environment is responsible for the net removal of pore water
310 REE in sub-surface sediments at active vent sites, thereby leading to the absence of
311 significant emission of dissolved REE into bottom waters.

312

313 Interestingly, careful examination of the vertical profiles at both Pockmark Field and Mud
314 Volcano sites shows that dissolved Nd concentrations are actually slightly depleted in
315 methane plumes relative to the overlying water column (Fig. 5). By analogy with what was
316 shown at hydrothermal systems (e.g. Michard et al., 1983; German et al., 1990; Sherrell et al.,
317 1999; Edmonds and German, 2004), this could suggest that additional scavenging of seawater
318 REE take place within the plume, perhaps through continuous adsorption onto Fe-Mn
319 oxyhydroxide phases or any other suspended particulates. Considering the Nd concentrations
320 measured at these two sites (Table 2), one can calculate that Fe-rich particles within the
321 plumes can incorporate up to ~ 7% of the dissolved REE content of ambient seawater.
322 Importantly, this also suggests that fluid seepage at cold seeps could act as a net sink in the
323 global ocean budget of the REE.

324

325 **4.4. Nd isotope constraints on processes controlling dissolved REE profiles in the Gulf of** 326 **Guinea**

327 Neodymium isotopic measurements provide further constraints on the processes controlling
328 the distribution of dissolved REE at the studied CTD hydrocast stations. The Nd isotope
329 ratios measured in this study encompass a large range of ϵ_{Nd} values from about -10.7 to -15.7
330 (Table 2). Surface waters (TSW) exhibit ϵ_{Nd} values of ~ -12.5 (CTD08-B13/14, 57m depth),
331 while the underlying subtropical underwater waters (STUW) are characterized by
332 unradiogenic values (~ -15.7; CTD3-B13/14, 60-180m depth). At the transition between
333 South Atlantic central waters and Antarctic Intermediate water, values are centered around ~ -
334 12.5, with the exception of one sample (~ -10.7; CTD06-B10/12, 460-500m depth,). The core
335 of AAIW displays lower ϵ_{Nd} values (~ -13.3; CTD03-B9/10, 990-1190m depth), while

336 NADW at the Reference Site is characterized by ϵ_{Nd} of ~ -12.5 . Note that the ϵ_{Nd} values for
337 the uncleaned foraminifera separates from core-top sediments at the Pockmark Field and
338 Reference sites (Table 5; taken as an indirect measurement of the Nd isotope composition of
339 bottom waters; Roberts and al., 2010) are also in very good agreement with the ϵ_{Nd} signature
340 determined for deep waters at these sites. Clearly, the large ϵ_{Nd} variability in the Niger Delta
341 water column indicates various sources of dissolved Nd. Below, we investigate several
342 possible mechanisms (i.e. isotopic exchange at cold seeps, sediment dissolution, lateral
343 advection), which could account for the vertical distribution of Nd isotopes at the three CTD
344 hydrocast stations.

345

346 First, although there is clear evidence for a net removal of REE at cold seeps (see previous
347 section), isotopic exchange processes between methane-rich fluids and/or associated particles
348 and seawater could possibly affect the Nd isotopic composition of the Gulf of Guinea bottom
349 waters. To test this hypothesis, we measured the Nd isotopic composition of cold seep
350 carbonate concretions and/or authigenic gypsum from sediments at the Mud Volcano and
351 Pockmark Field, to estimate the ϵ_{Nd} signature of fluids expelled at these sites. Authigenic
352 gypsum typically forms in reduced sediments after opening of the core sections, as a result of
353 the oxidation of sulfides to sulfate. During precipitation, it probably incorporates a number of
354 dissolved trace elements (including REE) from pore waters, and can hence be used to infer the
355 Nd isotopic composition of surrounding pore waters. At the Pockmark Field, the authigenic
356 carbonate concretion exhibits a ϵ_{Nd} value (-12.0 ± 0.3) similar to the measured bottom water
357 signature (-12.1 ± 0.6), but slightly lower than that for uncleaned foraminifera (-12.5 ± 0.1).
358 At the Mud Volcano, however, the obtained ϵ_{Nd} values for authigenic minerals (ϵ_{Nd} from $\sim -$
359 11.5 to -11.3 ± 0.2 ; Table 3) differ significantly from that of local bottom waters (-12.3 ± 0.4).

360 In agreement with our REE data, this suggests that fluid seepage at cold seeps do not modify
361 significantly the Nd isotopic composition of bottom water masses at ocean margins.

362

363 Second, as suggested previously for other areas of high sedimentary inputs (Nozaki and
364 Alibo, 2002; Tachikawa et al., 1999), partial dissolution of detrital particles settling through
365 the water column could play a significant role in controlling the vertical distribution of
366 dissolved REE. In the study area, however, both detrital sediments (average $\epsilon_{Nd} \sim -11.6 \pm$
367 0.3) and easily leachable fractions (i.e., dilute HNO_3 leachates; average -11.3 ± 0.3) are
368 characterized by a Nd isotopic signature significantly different from the seawater ϵ_{Nd} values
369 throughout the water column (Fig. 4). Clearly, this shows that interaction between seawater
370 and settling particles in this part of the Gulf of Guinea is unlikely to play any significant role
371 in the REE oceanic cycling.

372

373 Finally, based on these results, our preferred explanation is that lateral advection (i.e. ocean
374 circulation patterns) controls the observed vertical distribution of Nd isotope ratios at our
375 CTD hydrocast stations. This hypothesis is supported by evidence that 1) each water mass is
376 characterized by a well-distinct ϵ_{Nd} signature (Fig. 4), and 2) that the composite vertical
377 profile for ϵ_{Nd} closely resembles those for dissolved Nd concentrations (Fig. 4). In addition,
378 lateral advection would explain well why the Nd isotopic composition for STUW is very
379 unradiogenic ($\epsilon_{Nd} \sim -15.7 \pm 0.5$). In the Gulf of Guinea, the STUW is transported by the
380 northern Equatorial current (Fig. 1), which mainly receives its water from the northward
381 flowing Equatorial undercurrent (EUC). The main rivers draining western equatorial Africa
382 are delivering to the Atlantic Ocean suspended and/or dissolved loads characterized by very
383 low ϵ_{Nd} values (Congo ~ -16 ; Allègre et al ., 1996; Bayon et al., 2009; Ogooué ~ -24 ; G.

384 Bayon, unpubl. data; Ntem \sim -28; Weldeab et al., 2011). Therefore, if any significant
385 sediment/seawater interaction takes place at the western African ocean margin, one would
386 accordingly expect the water masses transported by the EUC to acquire a unradiogenic ϵ_{Nd}
387 signature, thereby explaining the low value measured in this study for STUW.

388

389 **4.5. Implications for the marine Nd cycle at continental margins**

390 As already mentioned in the Introduction, there are numerous evidence that the Nd isotopic
391 signature of water masses can be modified on ocean margins (e.g. Lacan and Jeandel, 2005;
392 Andersson et al., 2008; Amakawa et al., 2009), which suggest that sediment-seawater
393 interaction at margins could represent a major component of the oceanic Nd cycle (e.g.
394 Arsouze et al., 2009). Similarly, our Nd isotope seawater data for the Niger Delta margin also
395 provide another evidence for 'boundary exchange' (see discussion above for STUW). In
396 addition, the ϵ_{Nd} value determined for AAIW in our study area (between -13.3 ± 0.3 and -12.4
397 ± 0.4 ; Table 2) markedly differs from that reported for the same water mass at a nearby
398 station, in the western part of the Gulf of Guinea (-11.5 ± 0.3 ; Rickli et al., 2010). Similarly
399 to what was proposed above for explaining the unradiogenic signature of STUW, the lower
400 ϵ_{Nd} value measured here for AAIW most probably indicate sediment/seawater interactions at
401 the western equatorial African margin (Fig. 1). As discussed above, venting of reduced fluids
402 at cold seeps and dissolution of settling lithogenic particles both are unlikely to account for
403 the observed differences. Alternatively, one possible explanation accounting for the shift of
404 AAIW towards unradiogenic ϵ_{Nd} signature during its northward flow trajectory in this part of
405 the Gulf of Guinea would be that it was modified by diffusive benthic fluxes from organic-
406 rich sediments.

407

408 Of course, we cannot rule out the possibility that in certain parts of the ocean, dissolution of
409 settling particles, for example, represents the dominant input of dissolved REE to the ocean.
410 Additional case studies would also be clearly needed to confirm the results presented here.
411 However, our data suggest that diffusive benthic fluxes from suboxic settings could represent
412 a substantial source of dissolved REE in the Gulf of Guinea. Earlier works already suggested
413 that diffusion from marine sediments was likely to play a significant role in the marine REE
414 cycle (e.g. Elderfield and Greaves, 1982). Although only few studies have examined the
415 distribution of REE in interstitial waters of marine sediments (Elderfield and Sholkovitz,
416 1987; Sholkovitz et al., 1989; Haley et al., 2004), these works clearly showed that REE were
417 significantly enriched in pore waters relative to bottom waters, in agreement with the data
418 presented here, establishing strong chemical gradients in the near surface environment.
419 Certainly, the relatively high REE contents in pore waters are derived from the degradation of
420 potentially REE-rich phases (e.g., organic material, Fe-Mn oxyhydroxides) during early
421 diagenetic processes (e.g. Haley et al., 2004), which, in turn, is closely related to the amount
422 of organic compounds accumulated in subsurface sediments. As a first approximation,
423 therefore, one could suggest that the benthic fluxes of dissolved REE from marine sediments
424 are positively correlated with organic material contents. Because accumulation rates of
425 organic material in marine sediments are typically much higher on continental margins than in
426 open ocean settings, this would be entirely consistent with the proposed hypothesis that
427 sediment-seawater interactions at margins (in this case, benthic fluxes from suboxic
428 sediments) may represent an important component of the marine Nd cycle.

429

430 **4 – Conclusion**

431 The data presented here indicate that seepage of methane-rich fluids on continental margins
432 do not represent a source of dissolved Nd to the ocean. Similarly to what was previously

433 reported at submarine hydrothermal systems, it is very likely that Fe-Mn oxyhydroxide
434 precipitation in sub-surface sediments leads to quantitative removal of dissolved REE
435 whenever reduced (anoxic) fluids are emitted at cold seeps, acting possibly as a net sink for
436 REE in the ocean. In contrast, we suggest that diffusive benthic fluxes from suboxic surface
437 sediments could play a significant role in the marine Nd cycle, at least at the Niger Delta
438 margin.

439

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446

447

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660

661

662 **Figure captions**

663

664 **Figure 1. Location of the three studied areas on the Niger Delta margin.**

665 Location map showing the location of the study areas in the Gulf of Guinea. GC (Guinea
666 Current) transports Tropical Surface Water (TSW) eastward in the Gulf of Guinea, whereas
667 STUW (Subtropical Underwater) is advected by EUC (Equatorial Undercurrent) and NSEC
668 (Northern Equatorial Current), respectively. NICC (Northern Intermediate Countercurrent)
669 transports AAIW (Antarctic Intermediate Water) in the study area.

670

671 **Figure 2. Shaded bathymetric map for the Pockmark Field and the Mud Volcano areas.**

672 The location of the hydrocast stations and studied sediment cores is represented with large
673 white circles and small red circles, respectively. A) The pockmark field area is characterized
674 by the presence of large seafloor depressions with irregular shapes. B) The studied mud
675 volcano (about 1km wide) is composed of two distinct volcanic cones. Note the presence of a
676 well-characterized depression at the periphery of the mud volcano.

677

678 **Figure 3. Hydrography at the Niger Delta margin.**

679 A) Salinity versus depth profiles and B) Temperature-Salinity diagram for the three studied
680 hydrocast stations. The positions corresponding to the seawater samples analysed for Nd
681 isotopes are shown in the Temperature-Salinity diagram. TSW: Tropical Surface Water;
682 STUW: Subtropical Underwater; SACW: South Atlantic Central Water; AAIW: Antarctic
683 Intermediate Water; NADW: North Atlantic Deep Water (NADW).

684

685 **Figure 4. Depth profiles for methane, total dissolvable (TD) concentrations for Mn, Fe**
686 **and Nd, dissolved Nd and ϵ_{Nd} .**

687 The methane, TDFe and TDMn concentrations show that active fluid venting occurs at the
688 Pockmark Field and Mud volcano sites, with methane plumes rising up to about 100 m above
689 the seafloor. Note that ϵ_{Nd} represents the relative deviation of the $^{143}Nd/^{144}Nd$ ratios of a
690 sample, in parts per 10^4 , from that of the CHUR reference (CHondritic Uniform Reservoir):
691 $[(^{143}Nd/^{144}Nd)_{sample} / (^{143}Nd/^{144}Nd)_{CHUR} - 1] \times 10^4$. The average Nd isotopic composition of
692 leachable and detrital sediment fractions from Niger Delta core-top sediments is shown for
693 comparison.

694

695 **Figure 5. Relationships between total dissolvable (TD) and dissolved Nd concentrations**
696 **in the bottom part of the water column.**

697 The TDNd concentrations at the Pockmark Field and Mud volcano stations are significantly
698 enriched in the methane plumes. In contrast, dissolved Nd contents for the same samples are
699 much lower, and do not exhibit any significant enrichment relative to the overlying water
700 column. In the upper panel (Pockmark Field), note the small depletion in dissolved Nd at the
701 bottom part of the plume relative to the upper part.

702

703 **Figure 6. Shale-normalised REE patterns for seawater (both non-filtered and filtered)**
704 **and pore water samples.**

705 A) At the Pockmark Field, filtered samples collected from within the plume exhibit similar
706 shale-normalized patterns, while non-filtered samples collected at the same water depths show
707 a large range of REE patterns. B) At the Reference site, in contrast, both filtered and non-
708 filtered samples display similar REE patterns. C) Theoretical REE patterns generated by
709 partial dissolution of sediments in seawater samples having total suspended matter loadings
710 (TSM) of about 0.1, 0.2 and 1 mg/l. The theoretical REE concentra are generated using
711 simple mass balance calculations with REE concentrations for typical bottom water (filtered
712 sample CTD08-B3) and the easily leachable fraction of core ER-CS-38.

713

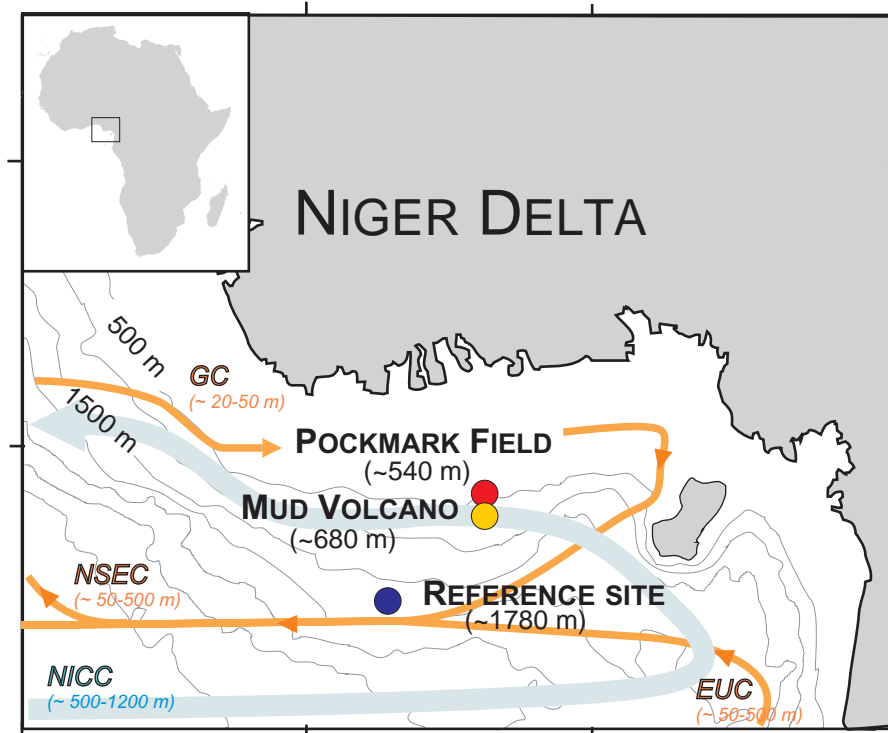


Fig. 1

A) POCKMARK FIELD

B) MUD VOLCANO

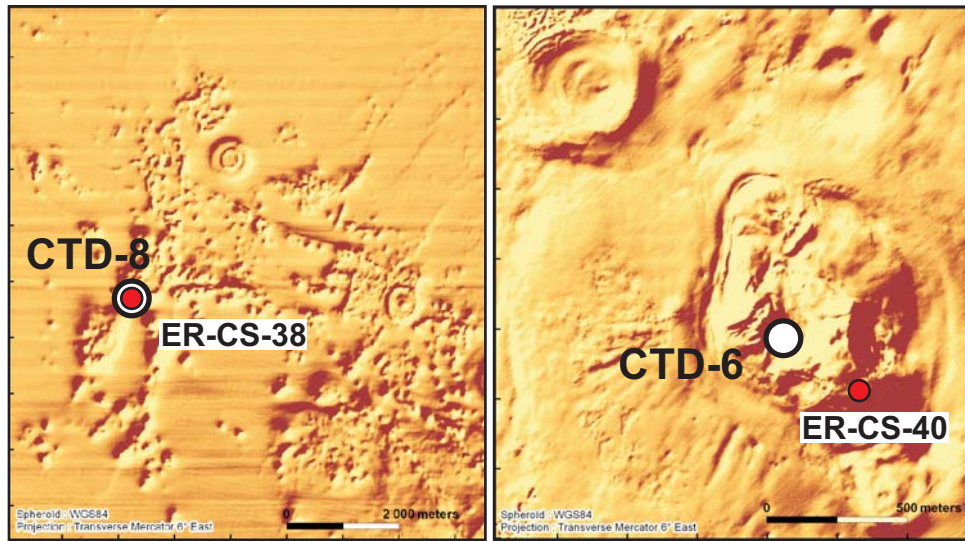
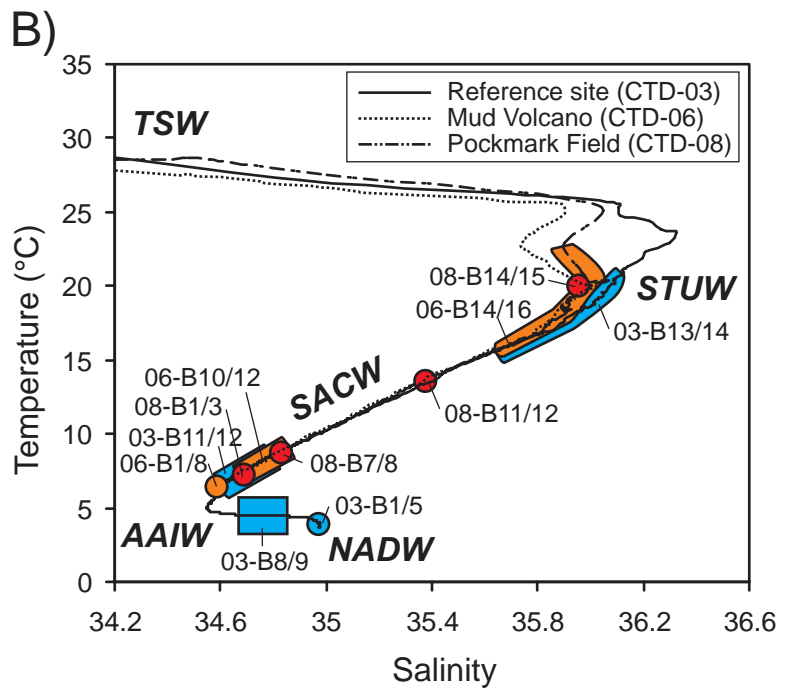
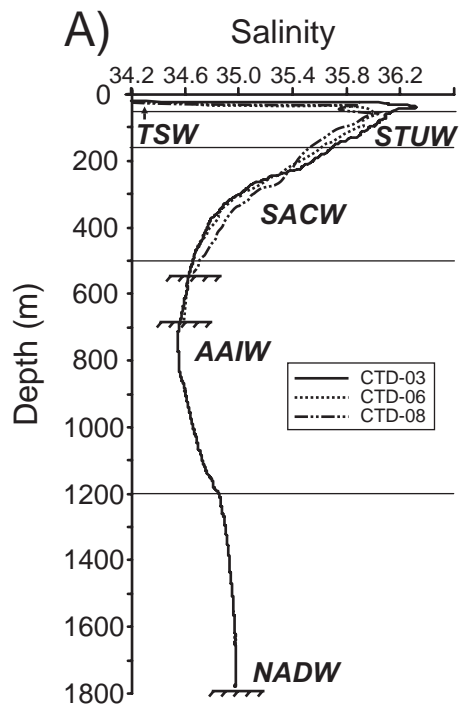


Fig. 2



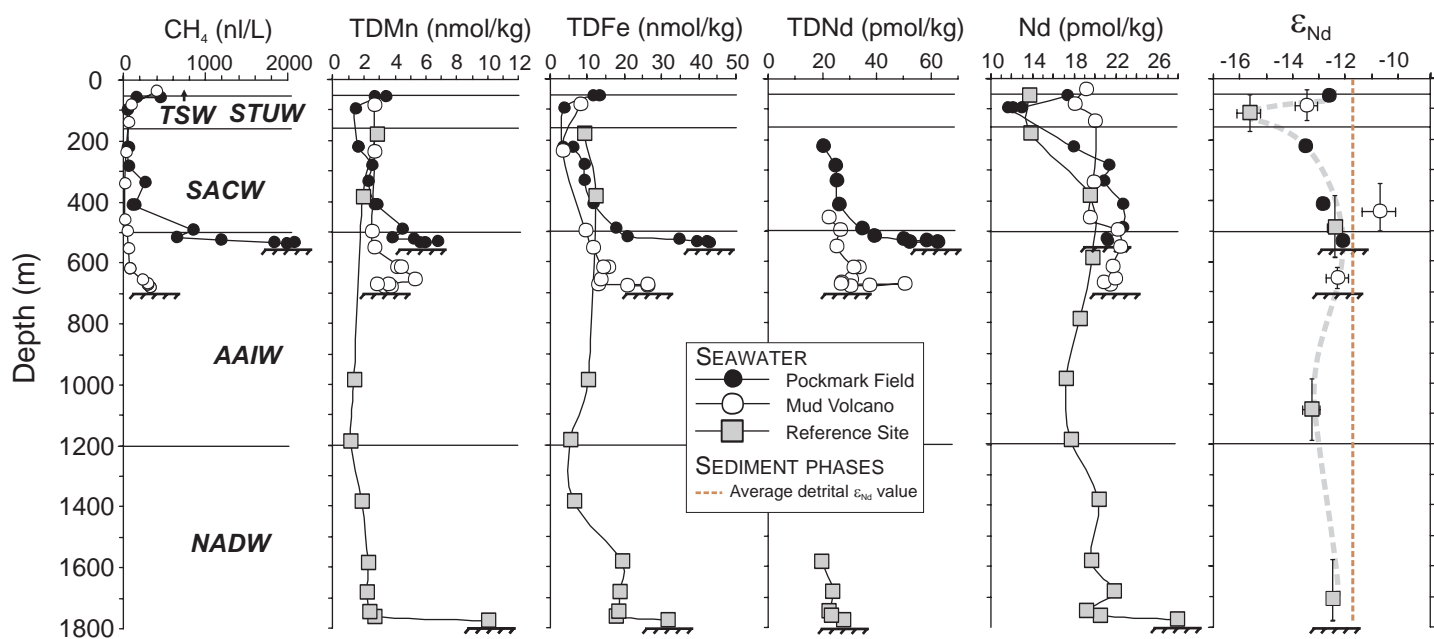


Fig. 4

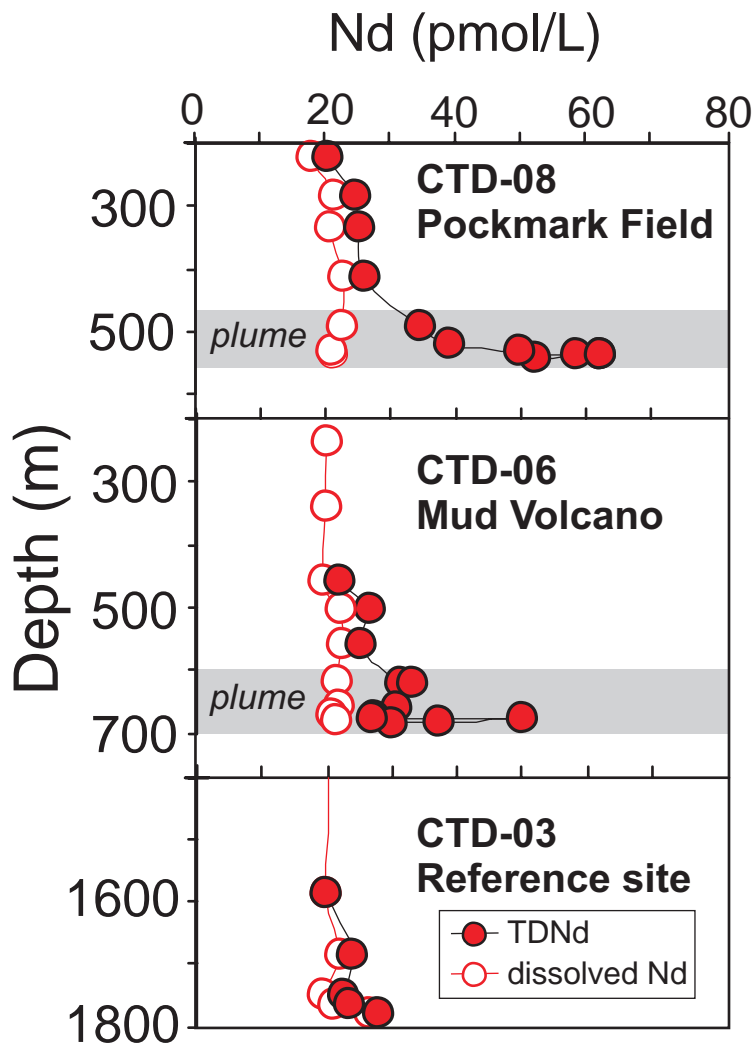


Fig. 5

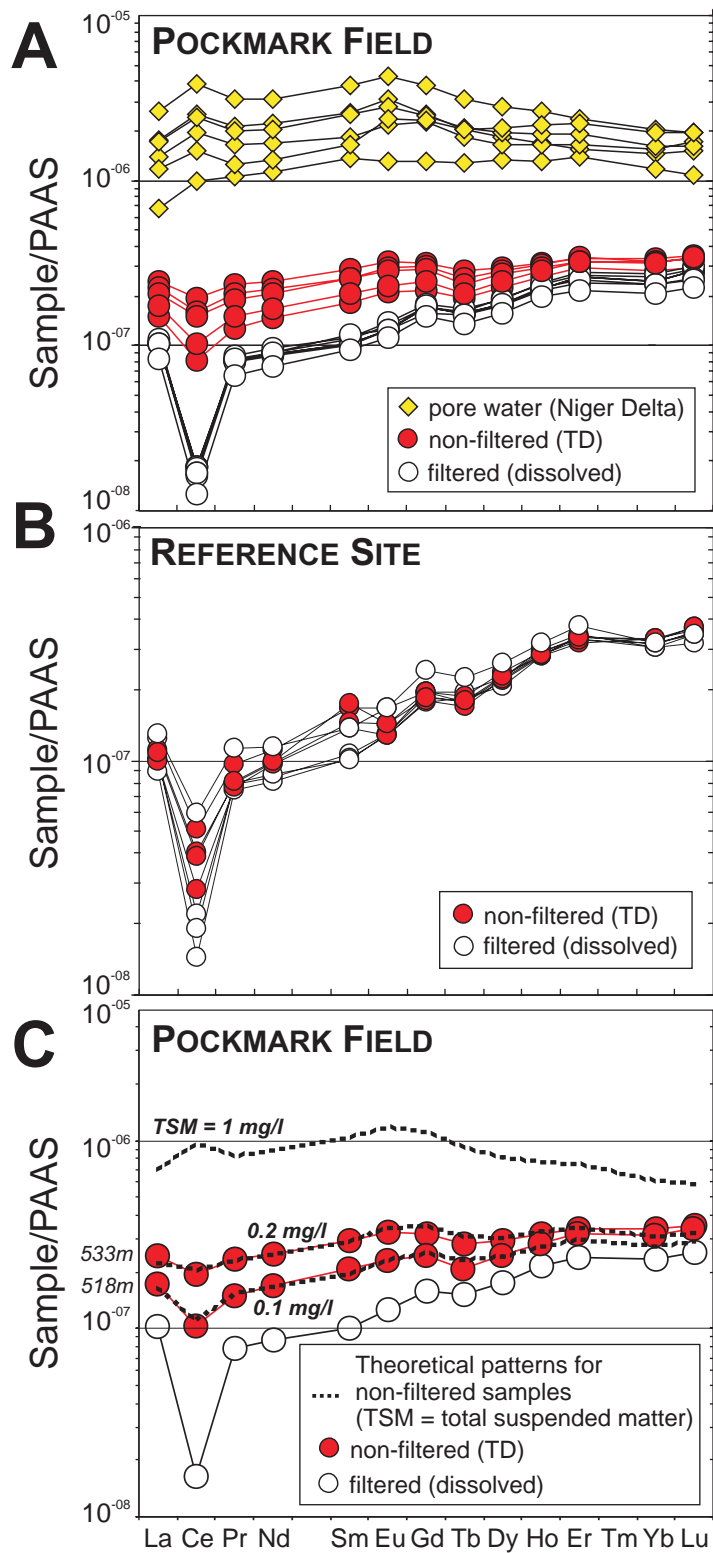


Fig. 6

Table 2
Dissolved REE concentrations and Nd isotope data for Niger Delta seawater samples.

Sample	Depth (m)	La (pmol/kg)	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu	Depth average	Nd	ϵ_{Nd}	2 s
<i>ER-CTDR-03 – Reference Area</i>																		
CTD3-B1	1776	36.03	34.09	7.12	27.85	5.36	1.21	7.30	1.10	7.58	1.92	6.44	5.22	0.87				
CTD3-B2	1762	27.93	12.57	5.02	20.42	4.06	0.93	5.63	0.88	6.57	1.78	5.69	4.99	0.79				
CTD3-B3	1747	24.85	8.25	4.72	19.16	3.91	0.92	5.31	0.90	6.07	1.70	5.83	5.16	0.89	1710	21.8	-12.5	0.2
CTD3-B4	1682	28.34	10.91	5.04	21.83	3.79	0.93	5.84	0.96	6.59	1.81	5.97	4.98	0.87				
CTD3-B5	1584	27.56	8.79	4.97	19.61	4.09	1.14	5.08	0.88	5.91	1.72	5.57	4.94	0.81				
CTD3-B7	1385	28.72	9.84	5.07	20.35	3.82	1.35	5.84	0.84	6.30	1.70	5.64	5.08	0.82				
CTD3-B8	1187	23.53	5.55	4.40	17.63	3.33	1.03	4.46	0.73	5.18	1.53	4.79	4.90	0.83	1087	17.4	-13.3	0.3
CTD3-B9	987	27.75	4.59	4.31	17.20	3.26	0.99	4.06	0.74	5.34	1.56	5.44	4.97	0.79				
CTD3-B10	788	21.92	4.54	4.48	18.55	3.76	0.91	4.37	0.69	5.04	1.52	5.11	4.56	0.77				
CTD3-B11	588	25.30	5.76	4.78	19.73	3.74	1.12	4.43	0.71	4.94	1.34	4.44	3.73	0.65	487	19.6	-12.4	0.3
CTD3-B12	385	25.11	8.39	4.59	19.51	3.74	1.12	4.56	0.70	4.71	1.24	4.02	2.97	0.51				
CTD3-B13	181	15.84	13.34	3.20	13.77	2.85	0.69	3.51	0.55	3.94	1.09	3.31	2.62	0.38	113	13.7	-15.7	0.5
CTD3-B14	58	17.92	16.65	3.29	13.65	2.65	0.80	3.71	0.54	3.95	1.08	3.25	2.49	0.39				
<i>ER-CTDR-06 – Mud Volcano</i>																		
CTD6-B1	679	28.64	17.38	5.31	21.38	4.04	0.70	5.29	0.78	5.33	1.38	4.68	4.37	0.72				
CTD6-B2	679																	
CTD6-B3	673																	
CTD6-B4	673																	
CTD6-B5	668	29.53	10.91	5.13	20.81	3.91	1.13	5.20	0.80	5.40	1.40	4.70	4.36	0.69	656	21.4	-12.3	0.4
CTD6-B6	658	30.98	11.12	5.37	21.92	3.96	1.24	5.42	0.80	5.60	1.43	4.64	4.04	0.68				
CTD6-B7	618	31.88	11.14	5.33	21.67	3.90	1.17	5.34	0.80	5.40	1.51	4.66	4.15	0.69				
CTD6-B8	618																	
CTD6-B9	553	34.76	16.10	5.68	22.38	3.66	0.40	6.07	0.88	5.78	1.49	4.76	3.91	0.70				
CTD6-B10	498	31.17	11.85	5.60	22.16	4.49	0.79	5.51	0.80	5.45	1.42	4.56	4.11	0.70				
CTD6-B11	458	26.56	8.68	4.68	19.45	3.79	0.72	4.75	0.72	4.89	1.31	4.42	4.02	0.68	435	21.5	-10.7	0.6
CTD6-B12	340	28.49	9.72	4.89	19.90	3.64	1.01	4.99	0.75	5.01	1.33	4.22	3.74	0.60				
CTD6-B14	141	24.82	11.01	4.77	19.96	3.57	1.10	4.81	0.71	4.87	1.25	3.92	3.03	0.49				
CTD6-B15	84	22.73	13.36	4.31	18.04	3.72	1.12	4.70	0.71	5.03	1.25	3.98	3.05	0.48	90	19.6	-13.5	0.4
CTD6-B16	38	23.72	18.79	4.56	19.08	3.82	1.18	5.44	0.81	5.69	1.37	4.31	3.37	0.52				
<i>ER-CTDR-08 – Pockmark Field</i>																		
CTD8-B1	535														534	21.2	-12.1	0.6
CTD8-B3	533	28.13	10.55	5.06	21.16	3.81	0.88	5.09	0.77	5.18	1.33	4.43	3.97	0.70				
CTD8-B4	525	28.88	10.37	5.11	21.02	3.98	0.94	5.12	0.78	5.54	1.42	4.54	4.26	0.73				
CTD8-B6	490	30.22	9.60	5.46	22.68	4.20	0.99	5.29	0.81	5.52	1.43	4.74	4.42	0.74				
CTD8-B7	411														411	22.7	-12.9	0.2
CTD8-B8	411	30.86	9.81	5.49	22.70	4.09	1.00	5.11	0.79	5.38	1.39	4.41	4.04	0.66				
CTD8-B9	335	28.14	9.23	4.96	20.79	3.63	0.89	4.70	0.75	5.05	1.31	4.11	3.81	0.63				
CTD8-B10	283	28.48	9.56	5.20	21.33	4.25	0.89	5.19	0.75	5.19	1.34	4.24	3.85	0.61				
CTD8-B11	223	23.03	7.12	4.12	17.91	3.50	0.80	4.45	0.65	4.53	1.19	3.72	3.36	0.56	223	17.9	-13.5	0.2
CTD8-B12	223																	
CTD8-B13-1	97	14.83	7.02	2.75	11.64	2.43	0.52	3.28	0.52	3.78	1.04	3.44	3.02	0.48				
CTD8-B13-2	97	17.13	7.38	2.84	12.12	2.55	0.56	3.13	0.53	3.87	1.09	3.62	3.09	0.49				
CTD8-B13-3	97	16.24	7.77	3.05	12.92	2.61	0.61	3.52	0.55	4.10	1.11	3.61	3.17	0.49				
CTD8-B14	57	19.97	12.58	4.02	17.28	3.42	0.81	4.48	0.71	4.96	1.25	3.85	3.32	0.51				
CTD8-B15	57														57	17.3	-12.6	0.2

The errors reported here correspond to the measurement errors (note that the external reproducibility is 0.2 ϵ units).

Table 3
Dissolved REE concentrations for pore water samples at active Niger Delta seeps.

Sample	Core depth (cm)	La (pmol/kg)	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu
<i>Mud Volcano</i>														
ER CS 40	0-2	185	563	65.9	266	49.9	9.3	39.1	6.2	38.5	7.9	23.8	19.0	2.6
<i>Pockmarks</i>														
N2-KI-41	0-2.5	485	1439	133	518	95.1	22.0	74.5	10.0	52.6	10.1	26.7	23.9	3.7
N2-KI-41	15-20	382	1117	103	388	68.5	15.5	67.6	9.0	47.3	9.9	28.2	25.4	4.3
N2-KI-20	0-2.5	723	2177	193	746	139	30.7	111.6	15.3	81.5	15.8	40.8	33.3	4.8
N2-KI-20	5-10	472	1386	124	486	90.7	19.9	73.9	10.2	56.7	11.6	32.6	26.5	4.0
N2-KI-20	75-80	326	870	79.2	319	62.5	16.8	68.5	9.9	59.6	13.2	38.0	31.6	4.9

Table 4

REE concentrations of easily leachable sediment fractions and cold seep carbonates.

Sample	Core depth (cm)	La (ppm)	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu
<i>Dilute HNO₃ sediment (<45 μm fraction) leachates</i>														
N1-KSF-01	0-2	71	238	19.4	74	144	3.0	12.8	1.9	9.8	1.8	5.1	3.9	0.52
N1-KSF-42	0-2	79	261	22.6	88	173	4.0	15.6	2.1	11.2	2.1	5.8	4.5	0.63
ER-CS-38	0-2	113	371	32.8	132	259	5.7	21.6	3.0	14.7	2.6	7.1	5.2	0.70
ER-CS-40	0-2	22	80	9.1	41	103	2.4	11.0	1.7	9.1	1.8	4.6	3.6	0.49
<i>Authigenic carbonates</i>														
ER-CS-38	500	8.2	21.0	2.17	8.5	1.59	0.35	1.35	0.21	1.08	0.21	0.56	0.45	0.06
ER-CS-40	5	3.3	7.7	0.85	4.1	0.94	0.21	0.72	0.09	0.45	0.08	0.21	0.15	0.02

Table 5

Nd isotope data for core-top sediment fractions and cold seep carbonate samples.

Sample	Core depth (cm)	ϵ_{Nd}	2 s
Reference Area			
N1-KSF-01			
Uncleaned forams	0-2	-12.54	0.08
Dilute HNO ₃ leachate	0-2	-11.9	0.2
Detrital sediment	0-2	-11.79	0.08
N1-KSF-42			
Uncleaned forams	0-2	-12.69	0.08
Dilute HNO ₃ leachate	0-2	-11.2	0.5
Detrital sediment	0-2	-11.70	0.08
Mud Volcano			
ER-CS-40			
Dilute HNO ₃ leachate	0-2	-11.2	0.2
Detrital sediment	0-2	-11.23	0.10
Authigenic gypsum	0-2	-11.3	0.2
Authigenic carbonate	5	-11.5	0.2
Pockmark Field			
ER-CS-38			
Uncleaned forams	0-2	-12.50	0.12
Dilute HNO ₃ leachate	0-2	-11.1	0.2
Detrital sediment	0-2	-11.79	0.07
Authigenic carbonate	500	-12.0	0.3