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On the gas-ice depth difference (Δdepth) along the EPICA Dome C ice core

F. Parrenin^{1,2}, S. Barker³, T. Blunier⁴, J. Chappellaz², J. Jouzel⁵, A. Landais⁵, V. Masson-Delmotte⁵, J. Schwander⁶, and D. Veres^{2,7}

¹Laboratoire Chrono-Environnement, Besançon, France

²UJF Grenoble 1, UMR5183, CNRS – Laboratoire de Glaciologie et Géophysique de l'Environnement (LGGE), Grenoble 38041, France

³School of Earth and Ocean Sciences, Cardiff University, UK

⁴Centre for Ice and Climate, Niels Bohr Institute, University of Copenhagen, Copenhagen, Denmark

⁵Laboratoire des Sciences du Climat et de l'Environnement, Gif-sur-Yvette, France

⁶Physics Institute, University of Bern, Bern, Switzerland

⁷Institute of Speleology, Romanian Academy, Cluj-Napoca, Romania

Correspondence to: F. Parrenin (parrenin@ujf-grenoble.fr)

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Abstract. We compare a variety of methods for estimating the gas/ice depth offset (Δdepth) at EPICA Dome C (EDC, East Antarctica). (1) Purely based on modelling efforts, Δdepth can be estimated combining a firn densification with an ice flow model. (2) The diffusive column height can be estimated from $\delta^{15}\text{N}$ and converted to Δdepth using an ice flow model and assumptions about past average firn density and thickness of the convective zone. (3) Ice and gas synchronisation of the EDC ice core to the GRIP, EDML and TALDICE ice cores shifts the ice/gas offset problem into higher accumulation ice cores where it can be more accurately evaluated. (4) Finally, the bipolar seesaw hypothesis allows us to synchronise the ice isotopic record with the gas CH_4 record, the later being taken as a proxy of Greenland temperature. The general agreement of method 4 with methods 2 and 3 confirms that the bipolar seesaw antiphase happened during the last 140 kyr. Applying method 4 to the deeper section of the EDC core confirms that the ice flow is complex and can help to improve our reconstruction of the thinning function and thus, of the EDC age scale. We confirm that method 1 overestimates the glacial Δdepth at EDC and we suggest that it is due to an overestimation of the glacial lock-in depth (LID) by the firn densification model. In contrast, we find that method 1 very likely underestimates Δdepth during Termination II, due either to

an underestimated thinning function or to an underestimated LID. Finally, method 2 gives estimates within a few metres of methods 3 and 4 during the last deglacial warming, suggesting that the convective zone at Dome C cannot have been very large at this time, if it existed at all.

1 Introduction

Ice cores provide a wealth of information on past climatic variations (Jouzel et al., 2007; Pol et al., 2011) and on past greenhouse gases concentrations (Lüthi et al., 2008; Loulergue et al., 2008) at time scales ranging from centennial to orbital (Earth orbit variations, 10^4 – 10^5 yr).

To interpret the records from ice cores, it is essential to derive accurate chronologies (e.g., Parrenin et al., 2007b). One of the peculiarities of ice core dating is that two age scales need to be derived: one for the ice matrix and one for the gas phase. Gas bubbles are always younger than the surrounding ice because they close off and trap the air at 50–120 m (depending on site conditions) below the surface, after the snow has densified into ice (Schwander and Stauffer, 1984). What is important for paleoclimatic studies is the lock-in depth (LID) where gas diffusion becomes negligible, and which is slightly smaller than the close-off depth (COD) where it is

not possible to pump air (Witrant et al., 2011). The determination of this ice/gas offset is essential, however, to derive the phase relationship between proxies recorded in the ice phase and in the gas bubbles. As an example, CO_2 was estimated to lag Antarctic temperature changes by 800 ± 600 yr during the last deglaciation (Monnin et al., 2001) by 800 ± 200 yr during Termination III (Caillon et al., 2003), and on average by 600 ± 400 yr during the last three deglaciations (Fischer et al., 1999). This finding suggests that CO_2 was an amplifier rather than the initial trigger of glacial terminations, although this view has been challenged by more recent studies (Loulergue et al., 2007; Pedro et al., 2012).

The gas/ice offset can be characterized in two different ways. Δage measures the difference in age between the ice and gas phases at any given depth. Δdepth , on the other hand, represents the depth difference between gas and ice of the same age. Each parameter has advantages and drawbacks. Δage is fixed when the gas is locked in and does not evolve with time because there is no relative movement of the gas bubbles or hydrates with respect to the surrounding ice. Δage is, however, strongly dependent on the rate of surface snow accumulation at the site, which is poorly constrained for the past; for a given LID and a given density profile, Δage is inversely proportional to the accumulation rate. By contrast, Δdepth is independent of the reference age scale used. It is only weakly dependent on accumulation rate since both the LID and the thinning function are only weakly affected by changes in accumulation rate. However, it continually evolves as the ice thins, which complicates its evaluation. Inversely, having observations of Δdepth from ice and gas proxies can provide useful information on the past flow of ice. Moreover, an error on Δdepth (expressed in metres) should be converted into years (using a prior chronology) to give an uncertainty on the ice–gas phasing (e.g. $\text{CO}_2/\text{temperature}$).

In this paper we focus on the evaluation of the Δdepth along the EDC (EPICA Dome C) ice core using different approaches. They fall into two broad categories: (1) estimation of the initial LID of gas bubbles and estimation of the thinning of snow/ice layers; and (2) determination of synchronous events in gas and ice proxy records. The results will be inter-compared and discussed.

Note that in the following, we have to deal with datasets on both the EDC96 and EDC99 ice cores. We systematically transfer all EDC96 datasets to EDC99 depths using a linear interpolation of the volcanic tie points between both cores (Parrenin et al., 2012). We use the same depth–depth relationship for both gas and ice datasets, i.e. we assume that Δage as a function of age is the same for both cores.

2 Methods

2.1 Δdepth from ice flow and densification models

From a mechanical point of view, Δdepth is given by:

$$\int_0^h \frac{D(z', t)}{\tau(z', t)} dz' = \int_{z-\Delta\text{depth}(z)}^z \frac{D(z', t=0)}{\tau(z', t=0)} dz', \quad (1)$$

where $D(z')$ and $\tau(z')$ are respectively the density of the material relative to pure ice and the thinning as a function (the ratio of a layer thickness to its initial thickness) of the depth z' and h is the lock-in depth (LID) at the time t when the initial snow layer, which is now ice at depth $z - \Delta\text{depth}(z)$, was at surface. We further define h_{ie} the lock-in depth in ice equivalent (LIDIE):

$$h_{\text{ie}} = \int D(y, t) dy, \quad (2)$$

Using this formalism, Δdepth is a function of the gas depth. We implicitly assumed that at a given depth z , the gas age is uniquely defined. In reality, because of gas diffusion in the firn and because of the gradual bubble close off process, any depth contains a distribution of age which can be accurately approximated by a log-normal function (Köhler, 2010). We implicitly use here the median of this log-normal distribution as the so-called gas age for any given depth.

$D(z', t=0)$ can be estimated by measuring the weight and the volume of the drilled cores. However, no reliable quantitative proxy has been proposed for $\tau(z', t=0)$, $\tau(z', t)$ and $D(z', t)$ and their evaluation usually relies on ice flow (e.g., Reeh, 1989; Parrenin et al., 2007a; Salamatin et al., 2009b) and firn densification modelling (e.g., Herron and Langway, 1980; Arnaud et al., 2000; Salamatin et al., 2009a). We will detail in the following the ice flow model and firn densification model used in this study.

2.1.1 Ice flow model

A one-dimensional (1-D) ice flow model has been used to construct the modelled age scale χ_m at the EDC drilling site and to derive a modelled thinning function τ_m (Parrenin et al., 2007a). In this model, the vertical velocity $u_{\bar{z}}$ of the ice relative to the bedrock is expressed as:

$$u_{\bar{z}}(\bar{z}) = - \left[m + \left(a - \frac{\partial H}{\partial t} - m \right) \omega(\zeta) \right] \quad (3)$$

where z is the vertical coordinate of the ice particle (oriented toward the top), $\bar{z} = z - B$ is the distance to the bedrock (B is the bedrock elevation), $\zeta = \bar{z}/H$ is the non-dimensional vertical coordinate, m is the melting rate at the base of the ice sheet, a is the surface accumulation rate, H is the ice thickness and $\frac{\partial H}{\partial t}$ is its temporal variation. $\omega(\zeta)$, called the flux shape function (Parrenin and Hindmarsh, 2007), depends on

the non-dimensional vertical coordinate and is the contribution of one sliding term and one deformation term:

$$\omega(\zeta) = s\zeta + (1-s)\omega_D(\zeta) \quad (4)$$

where s is the sliding ratio (ratio of the basal horizontal velocity to the vertically averaged horizontal velocity; it is 0 for no sliding and 1 for full sliding) and $\omega_D(\zeta)$ can be approximated by (Lliboutry, 1979):

$$\omega_D(\zeta) = 1 - \frac{p+2}{p+1}(1-\zeta) + \frac{1}{p+1}(1-\zeta)^{p+2} \quad (5)$$

where p is a parameter for the vertical profile of deformation $\omega_D(\zeta)$. The values of p , m and s are assumed constant through time.

The past variations of ice thickness $H(t)$ are obtained from a 1-D model (Parrenin et al., 2007a) fitted onto the results of a 3-D model of the Antarctic ice sheet (Ritz et al., 2001). The main process is the reduced accumulation rate during glacial times, which induces a lower elevation (and reduced ice thickness) at Dome C during glacial periods. However, preliminary results with an improved 3-D model with increased spatial resolution suggests that the presence of an ice sheet in the Ross Embayment might limit the impact of a reduced accumulation on the elevation at the EDC site, at least during the last glacial maximum (C. Ritz, personal communication, 2012). This is why, in the following, we will also test the hypothesis of zero ice thickness variations at the EDC site. This appears as an extreme case, given that some geomorphological data in the Transantarctic Mountains show little elevation change of the Antarctic plateau for the last glacial maximum (Denton et al., 1989), despite the presence of the ice sheet in the Ross Embayment.

Modelled accumulation a_m and temperature T are deduced from the deuterium content of the ice extracted from the drill core, through the following relationships:

$$a_m = a^0 \exp(\beta \Delta\delta D_{\text{smo}}) \quad (6)$$

$$T = T^0 + \alpha \Delta\delta D_{\text{cor}} \quad (7)$$

where a^0 and T^0 are surface accumulation and temperature for a reference deuterium content of -396.5‰ (roughly corresponding to the present-day value). $\Delta\delta D_{\text{smo}}$ is a 50-yr smoothed version of $\Delta\delta D_{\text{cor}}$ because the accumulation rate a_m is supposed to be related to the isotope content of the deposited snow only over a certain time interval (high frequency variations of deuterium may be affected by post depositional processes such as wind scouring). The poorly constrained glaciological parameters $p = 2.30$, $m = 0.066\text{ cm-of-ice yr}^{-1}$, $s = 2.23\%$, $a^0 = 2.841\text{ cm-of-ice yr}^{-1}$ and $\beta = 0.0157$ were obtained by fitting independent age markers identified within the core (Parrenin et al., 2007a). The inferred value for β appears consistent with modern spatial gradients in central East Antarctica (Masson-Delmotte et al.,

2008). Given the inability of the model to fit some age markers (Dreyfus et al., 2007; Parrenin et al., 2007b), the thinning function τ and surface accumulation rate a were tuned a posteriori using a spline method so that the tuned age scale χ fit these age markers (see Appendices of Parrenin et al., 2007b; Dreyfus et al., 2007).

It is difficult to quantify the uncertainty on the modelled thinning function $\tau_m(z)$ because we do not know which processes are not accounted for in the ice flow models. Here we consider only the non-laminar ice flow effects and assume that the error they induce on $\ln(\tau_m)$ is:

$$\sigma_{\ln(\tau)}(z) = \frac{k}{H} \int \frac{D(z')}{\tau_m(z')} dz', \quad (8)$$

where $D(z)$ is the density of the material relative to pure ice and k is a proportionality coefficient. We infer the value of k with a residual approach using the multiplicative correction for the thinning function $C(z)$ which has been inferred from the orbital tuning of $\delta^{18}\text{O}_{\text{atm}}$ in the 2700–3200 m interval of the EDC ice core (Dreyfus et al., 2007). k is simply given by the standard deviation of the following function $f(z)$:

$$f(z) = \ln(C(z)) \left(\frac{1}{H} \int \frac{D(z')}{\tau_m(z')} dz' \right)^{-1}, \quad (9)$$

which is represented in Fig. 1. This gives $k = 0.0974$ and the resulting $\sigma_{\ln(\tau)}$ function is plotted in Fig. 2.

2.1.2 Firn densification model

For the firn densification modelling exercise and the determination of $D(z)$, we used the Arnaud/Goujon model (Goujon et al., 2003). Arnaud et al. (2000) developed an advanced densification model which considers two densification stages: pure sliding of snow grains for density lower than $\sim 0.55\text{ g cm}^{-3}$, and pure deformation of grains for density higher than $\sim 0.55\text{ g cm}^{-3}$. Goujon et al. (2003) then incorporated heat transfer into this model. In the applications below we used a surface density of 0.35 g cm^{-3} .

The lock-in density $D(z=h)$ is determined from the total air content (TAC) of the ice (Martinerie et al., 1992, 1994; Raynaud et al., 2007) corrected for local atmospheric pressure changes (due in particular to elevation changes) using the perfect gas law. In the applications below, for simplification and based on the approach by Martinerie et al. (1992, 1994), we use a conventional linear empirical relationship between the volume of pores at close-off (V_c , $\text{cm}^3\text{ g}^{-1}$) and surface temperature T_S (K) as:

$$V_c = 6.95 \times 10^{-4} T_S - 0.043 \quad (10)$$

The evolution of closed porosity in the firn, P_{closed} , is deduced from the following relationship:

$$P_{\text{closed}} = 0.37 P_{\text{total}} \left(\frac{P_{\text{total}}}{P_{\text{close-off}}} \right)^{-7.6}, \quad (11)$$

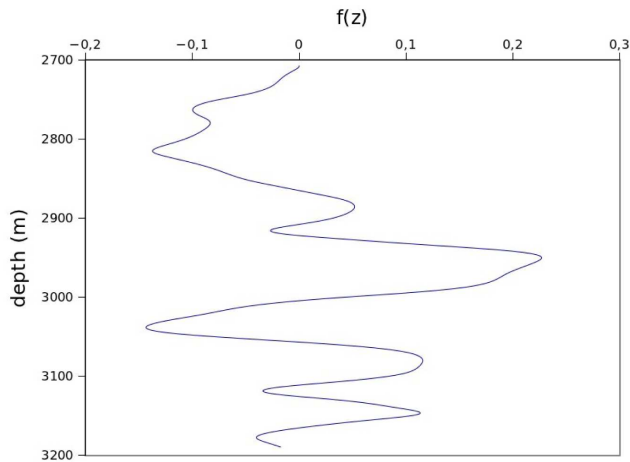


Fig. 1. Function $f(z)$ in EDC as described in Eq. (9).

which has been calibrated with P_{closed} and P_{total} measurements on several ice cores from Greenland and Antarctica (J.-M. Barnola, personal communication, 2009). This relationship means that at close-off, 37 % of the pores are closed. The LID is further defined, at EDC, when 20 % of the pores are closed.

The steady LID, Δage and LIDIE/LID ratio simulated by the Goujon/Arnaud model are represented in Fig. 3 in a surface temperature–surface accumulation diagram for a surface density of 0.35 g cm^{-3} . Qualitatively speaking, LID is greater for a greater accumulation or for a lower temperature, whereas Δage is greater for a lower accumulation or for a lower temperature. The LIDIE/LID ratio is practically constant: it only varies between 0.682 and 0.702 for the accumulation and temperature ranges considered in Fig. 3.

2.2 Δdepth from ice flow modelling and $\delta^{15}\text{N}$ -based estimates of firn thickness

The ice flow modelling part of this estimate has been described above. We use here the variations of ice thickness at EDC as derived for the EDC3 age scale (Parrenin et al., 2007a, b).

Now h is estimated using the fact that, below a convective zone of height h_{conv} where the air is mixed (Colbeck, 1989, 1997; Sowers et al., 1992; Bender et al., 1994; Kawamura et al., 2006), gravitational settling enriches heavy isotopes of inert gases (such as $\delta^{15}\text{N}$ of N_2 and $\delta^{40}\text{Ar}$) proportionally to the diffusive column height h_{diff} (Craig et al., 1988; Sowers et al., 1989; Dreyfus et al., 2010) until gases no longer diffuse in the open pores. We implicitly assume here that all gases stop diffusing at the same depth. Note that a recent study suggested that some trace gases continue to diffuse below the LID defined by the start of the $\delta^{15}\text{N}$ plateau (Buizert et al., 2012). In delta notation, this gravitational fractionation is expressed as:

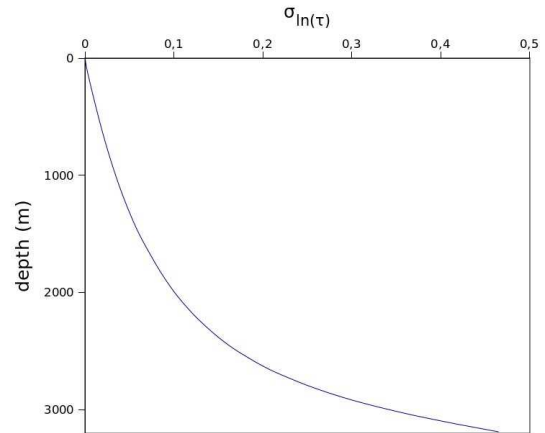


Fig. 2. Evolution of the error in the logarithm of thinning function as a function of the depth in the EDC ice core.

$$\delta_{\text{grav}} = \left[\exp\left(\frac{\Delta m g h_{\text{diff}}}{RT}\right) - 1 \right] \times 1000, \quad (12)$$

where Δm is the mass difference between species (kg mol^{-1}), g is the gravitational acceleration (9.825 m s^{-2} for Antarctica), R is the universal gas constant ($8.314 \text{ J mol}^{-1} \text{ K}^{-1}$) and T is the firn temperature (K). Equation (12) can be approximated within 0.02 % with:

$$\delta_{\text{grav}} \simeq \left(\frac{\Delta m g h_{\text{diff}}}{RT} \right) \times 1000. \quad (13)$$

Thermal fractionation of $\delta^{15}\text{N}$ occurs because of the temperature difference ΔT between the surface and the LID:

$$\delta_{\text{th}} = \Omega(T) \Delta T_{\text{diff}}, \quad (14)$$

where ΔT_{diff} is the temperature difference between the top and the bottom of the diffusive zone. $\Omega(T)$ has been estimated from laboratory measurements (Grachev and Severinghaus, 2003).

Conversely, one can deduce h from the $\delta^{15}\text{N}$ data at EDC, measured on the last three glacial terminations and five glacial-interglacial cycles between 300 and 800 kyr BP (Dreyfus et al., 2010):

$$h = h_{\text{conv}} + \left(\delta^{15}\text{N} - \Omega(T) \Delta T_{\text{diff}} \right) \left(\frac{\Delta m g \times 1000}{RT} \right)^{-1}. \quad (15)$$

One of the reasons for variations in the convective height is the change of wind stress. GCM experiments for the LGM show little variations in wind on the East Antarctic plateau (Krinner et al., 2000). Note that we have evidence for a large convective zone at present at some sites (Bender et al., 2006; Severinghaus et al., 2010). In the applications below, we will assume that there was no convective zone at EDC during the last 800 kyr, in agreement with current observations (Landais

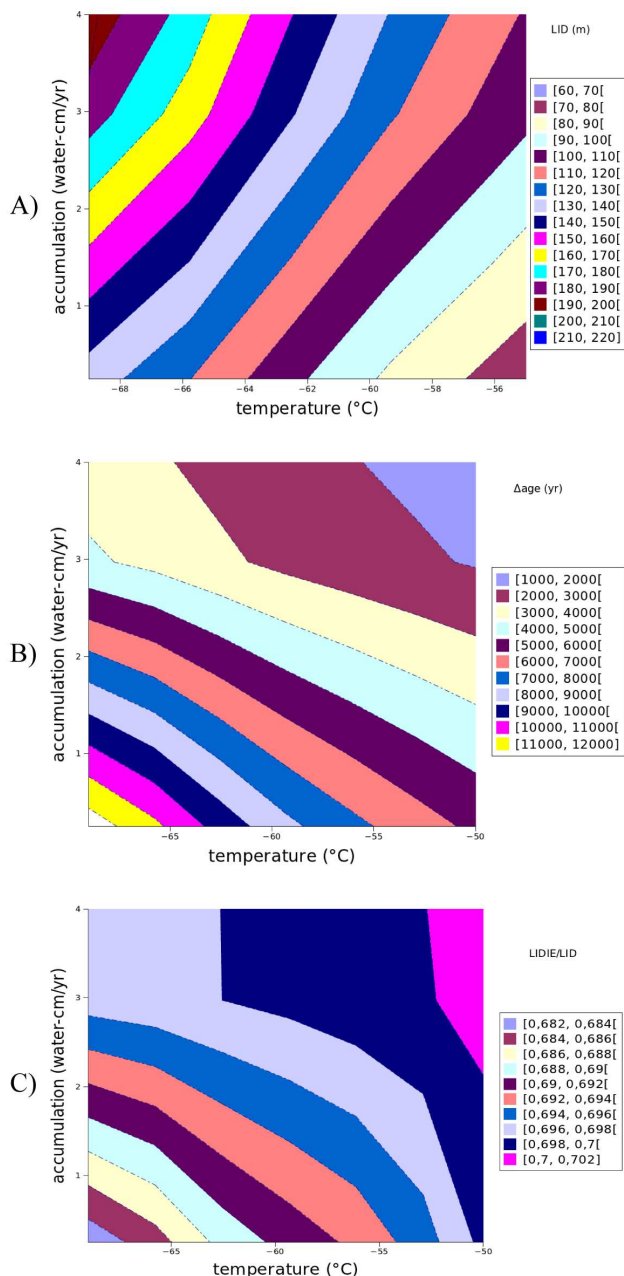


Fig. 3. Simulations of the Goujon/Arnaud firn densification model (Goujon et al., 2003), in a (surface temperature–surface accumulation) diagram. Surface density is 0.35 g cm^{-3} . LID is taken at 20 % of closed porosity. **(A)** Simulated LID; **(B)** Simulated Δage ; **(C)** Simulated LIDIE/LID ratio.

et al., 2006). We will be able to assess the limits of our assumption about the convective zone when we compare the results from the different methods. We further deduce the temperature difference in the firn ΔT_{diff} from simulations of the temperature field by the Goujon/Arnaud model (Loulergue et al., 2007, scenario 1). We take the surface temperature as computed for the EDC3 age scale (Parrenin et al., 2007a),

as a function of the ice depth. We also need a prior Δdepth estimate to convert the ice depths to gas depths and we use the EDC3 scenario 1 estimate (Loulergue et al., 2007). We estimate the uncertainty on the temperature estimate to be $< 4 \text{ K}$ which translates into a $< 2 \%$ 2σ error on h . We also estimate the uncertainty on the temperature gradient to be $< 0.003 \text{ K m}^{-1}$ which translates into a $< 1 \%$ 2σ error on h .

Following the Goujon/Arnaud model simulations (see Fig. 3c), Eq. (2) is simplified into:

$$h_{\text{ie}} \simeq h \times 0.698. \quad (16)$$

Bender et al. (2006) also used a similar approximation. It corresponds to assuming that the average density of the firn is correctly predicted by the Goujon/Arnaud model. If this model would not predict the right average densification velocity but would predict instead the right densification profile shape, this approximation would still be valid. So this leaves us with mainly two reasons why this approximation would not be valid: (1) a variable surface density and/or (2) a variable lock-in density. Option (1) cannot be ruled out since, depending on the characteristics of the surface (glazed surface, megadunes, etc.), surface densities $> 0.4 \text{ g cm}^{-3}$ are observed on the East Antarctic plateau (Courville et al., 2007). Note that because the densification velocity is greater at surface than in depth, an error of x on the surface density has a relatively low impact of $\sim x/3$ on the average density. Because the density at the COD does not change very much with time (it is well constrained by the measured total air content of the ice), option (2) would imply a varying difference between LID and COD. In total, we estimate the 2σ error of Eq. (16) to be 5 %. Equation (1) is solved assuming the thinning function expressed as a function of z_{ie} , the ice equivalent depth, was the same at the time of deposition as for the present. The error of this approximation is due to the varying basal melting/accumulation ratio and to the varying ice thickness (Parrenin et al., 2007a), but we evaluate it to be $< 0.1 \%$ on Δdepth .

2.3 Δdepth from ice and gas synchronisation to GRIP

Δdepth at the depth of the ^{10}Be peak (Raisbeck et al., 2007) in the EDC ice core can be estimated by linking the ice and gas signals to GRIP (Loulergue et al., 2007). The ice link is obtained by ^{10}Be synchronisation of EDC and GRIP for two ^{10}Be sub-peaks corresponding to the Laschamp geomagnetic event (Raisbeck et al., 2007). The gas link is obtained by matching the EDC CH_4 record to the GRIP ice isotopic record (Fig. 4), assuming that these two records are synchronous during the rapid Dansgaard–Oeschger (DO) transitions (Flückiger et al., 2004; Huber et al., 2006). In Fig. 4, we align the onsets of DO9 and DO11 and remark that the onset of DO10 is also aligned. Linear interpolations allow obtaining the corresponding gas depths in EDC99. Our horizontal scale is expressed in GRIP ages, since there are significant variations of accumulation which would make the

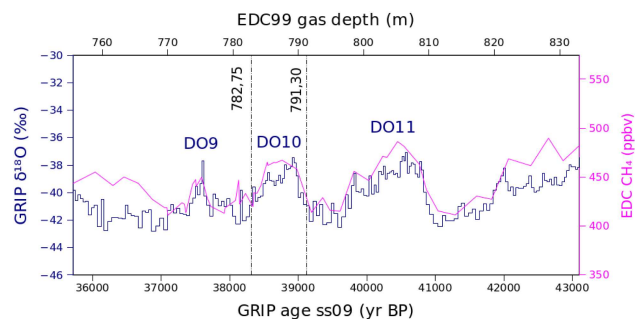


Fig. 4. Synchronisation of GRIP $\delta^{18}\text{O}$ and EDC CH_4 during DO9–11. The two vertical dashed lines mark the position of the ^{10}Be sub-peaks in the GRIP core and give the corresponding gas depths in the EDC99 ice core.

interpolation inaccurate in depths. It is not necessary to use an age scale for EDC since the variations of accumulation are small and the variations of thinning are $< 2\%$ over the considered interval.

This finally gives these two estimates of the Δ depth at EDC99: Δ depth at 782.9 m = 48.9 ± 2 m (2σ); and Δ depth at 791.5 m = 48.2 ± 2 m (2σ). The uncertainty accounts for the uncertainty in the ^{10}Be sub-peaks positions in EDC (1.1 m) and in GRIP (1.1 m EDC equivalent depth) and for the GRIP–EDC synchronisation (1.3 m).

2.4 Δ depth from ice and gas synchronisation to EDML

Another approach to deduce EDC Δ depth is to synchronise the ice core records, both in the ice and gas phases, to a higher accumulation Antarctic ice core, such as EDML (see Fig. 5) which has a better constrained ice/gas offset (Louergue et al., 2007). Such an approach has already been applied to constrain the Vostok gas/ice offset using data from the Byrd ice core record (Blunier et al., 2004; Bender et al., 2006) and the EDC gas/ice offset using EDML (Louergue et al., 2007). The EDC and EDML ice cores have been synchronised (Severi et al., 2007; Ruth et al., 2007) using volcanic stratigraphic markers recorded in the ice phase. Here we also derive 20 new CH_4 tie points (see Table 1 and Fig. 6) (Louergue et al., 2007, 2008; Schilt et al., 2010), mainly at the onsets of Greenland Interstadials (GI) over the period 0–140 kyr BP. We did not use systematically the tie points of Louergue et al. (2007) or of Schilt et al. (2010) since (1) we do not use ends of GIs because they are less well marked than the onsets and therefore bring little information with respect to the neighbouring onsets, (2) we do not use GI2 and GI9 because we reckon their identification is too ambiguous, and (3) we choose the tie points exactly at the mid-transitions. The evaluation of Δ depth at EDC now relies on its evaluation at EDML. We derived the latter from Eq. (1). For the thinning, we did not use estimates based on an ice flow model (Huybrechts et al., 2007) because for the same age, the depth is larger at EDML than at EDC and the evaluation of the

Table 1. CH_4 synchronisation tie points between EDC and EDML and corresponding Δ depth estimates at EDC.

Description	EDC99 gas depth (m)	EDML gas depth (m)	Δ depth (m)	1σ uncertainty (m)
YD-Hol	418.5	718.0	58.4	3.84
BA-YD	443.5	766.4	56.2	3.79
onset BA	476.1	830.2	55.4	3.92
onset DO3	639.1	1155.1	49.7	3.32
onset DO4	651.9	1174.2	50.2	3.37
onset DO5	688.1	1233.7	49.3	3.26
onset DO6	702.1	1260.8	46.3	3.05
onset DO7	719.7	1286.4	48.3	3.15
onset DO8	751.3	1338.0	47.9	3.23
onset DO10	791.0	1404.6	48.0	3.10
onset DO11	809.2	1439.0	42.0	3.32
onset DO12	848.5	1490.0	44.0	3.09
onset DO14	925.0	1601.8	43.6	2.99
onset DO17	986.5	1688.2	43.0	2.92
onset DO18	1038.5	1760.5	43.1	2.82
onset DO19	1105.0	1861.3	41.2	3.05
onset DO20	1142.0	1915.0	39.4	2.73
onset DO21	1239.0	2023.2	34.2	3.01
onset DO24	1473.0	2230.3	28.3	2.32
onset MIS5.5	1722.0	2369.0	38.2	3.18

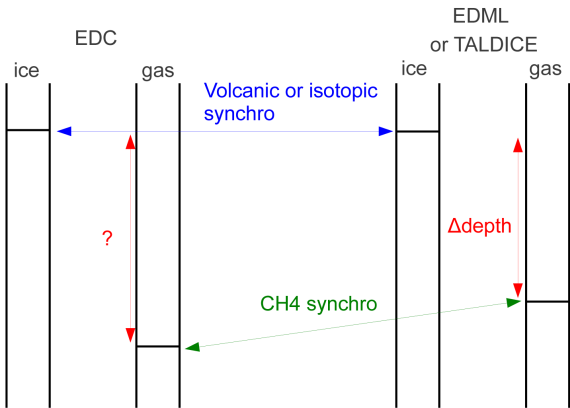


Fig. 5. Scheme illustrating the deduction of the Δ depth at EDC from ice (volcanic) and gas (CH_4) synchronisation to the EDML or TALDICE ice cores and evaluation of Δ depth at EDML or TALDICE.

thinning function from ice flow modelling thus becomes inaccurate. Instead, we fixed the EDML1 age scale (synchronised to EDC3, Ruth et al., 2007), used EDML accumulation rates from Louergue et al. (2007, scenario 1) and deduced an EDML thinning function. For the LID, we used the Goujon et al. (2003) densification model forced with the temperature and accumulation estimates as derived from Louergue et al. (2007, scenario 1). The LID is taken at 5 % of closed porosity.

This method is more precise than a direct evaluation of the Δ depth at EDC from modelling. Indeed, the accumulation rate is 3 times higher at EDML than at EDC. An error in EDML LID thus has a 3 times lower impact than at EDC in terms of ages.

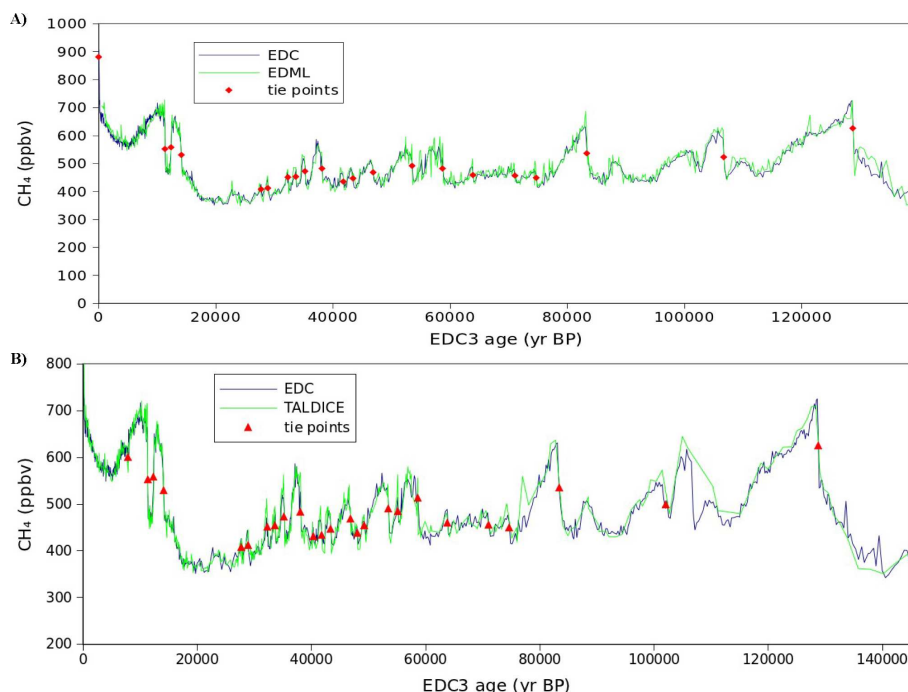


Fig. 6. Gas synchronisation between EDC and EDML (**A**) and between EDC and TALDICE (**B**) using the CH_4 records (Loulergue et al., 2007, 2008; Schilt et al., 2010; Buiron et al., 2011).

There are 4 sources of uncertainty in this EDML-synchro based approach: (1) the uncertainty in the gas (CH_4) synchronisation, (2) the uncertainty in the LID estimate at EDML, (3) the uncertainty in the thinning function at EDML, and (4) the uncertainty in the ice (volcanic) synchronisation (including the interpolation between two neighbouring volcanic tie points). We estimate uncertainty (1) (2σ) as half the duration of the CH_4 transition. Uncertainty (2) (2σ) is thought to be $< 20\%$ at EDML (Landais et al., 2006), i.e. $< 7\%$ at EDC. Based on the relative duration of events in different glaciological time scales (Parrenin et al., 2007b), we deduce that uncertainty (3) (2σ) is $< 10\%$ (Parrenin et al., 2007b). With the same argument, uncertainty (4) (2σ) is estimated to be $< 10\%$ of the distance to the nearest tie point, i.e. we neglected the uncertainty in the tie points. In order to compute the total uncertainty, we assume uncertainties (1), (2), (3) and (4) to be independent.

Note that a recent study (Köhler, 2010) suggested that aligning the mid-transitions of CH_4 in different ice cores induces an error because of the different diffusion times of the gas signals. We consider that this conclusion only applies if one defines the gas age as the minimum gas age of the distribution. We defined here the gas age as the median of the distribution and we are therefore free from such an error.

2.5 Δ depth from ice and gas synchronisation to TALDICE

The method is the same as in the previous sub-section (see Fig. 5). Its advantage is also based on the fact that TALDICE accumulation rate is 3 times higher than EDC accumulation rate. The EDC and TALDICE ice cores can be synchronised in the ice phase using volcanic markers for the last 42 kyr (Severi et al., 2012) and isotopic records for older time periods (Jouzel et al., 2007; Stenni et al., 2011, see Table 2) as well as in the gas phase using the CH_4 records (Loulergue et al., 2008; Buiron et al., 2011, see Table 3 and Fig. 6). We do not use the tie points of Buiron et al. (2011) since they are not always placed exactly at mid-transitions. We also restricted the tie point selection to the part of CH_4 records bearing the less disputable common structure. We use Δ age at TALDICE as computed by Buiron et al. (2011).

The uncertainty is calculated in the very same way as for the synchronisation to EDML.

2.6 Δ depth from the thermal bipolar seesaw hypothesis

Following the so-called thermal bipolar seesaw hypothesis (Stocker and Johnsen, 2003), Greenland temperature is related to the derivative of the Antarctic temperature derived from EDC isotopic record (Barker et al., 2011). The most viable mechanism for abrupt climate changes in the North Atlantic region involves reorganizations of the ocean circulation (Stommel, 1961; Ruddiman and McIntyre, 1981) but atmospheric mechanisms may also be at play in the antiphase

Table 2. Isotopic synchronisation tie points between the TALDICE and EDC99 ice cores.

Depth-TALDICE	Depth-EDC99
1160	800
1220	880
1260	950
1285	1002
1303	1059
1312	1100
1331	1200
1353	1320
1375	1449
1411	1700
1440	1855
1446	1889
1463	1906
1471	1935
1477.5	1952.5
1485	1981
1493.7	2027
1497	2048
1508	2093
1522.5	2170
1528	2222
1534	2235
1545	2295
1582	2500

relationship proven for the last glacial period (Blunier et al., 1998; Blunier and Brook, 2001; EPICA community members, 2006; Capron et al., 2010). Using the seesaw hypothesis, we can consequently synchronise the deuterium content of the EDC ice (a proxy for Antarctic temperature) with the CH_4 (Louergue et al., 2008) content of the EDC gas bubbles (a proxy of Greenland temperature) and produce Δ depth estimates during periods of fast CH_4 variations corresponding to the various maxima or minima in the deuterium record.

To localize the multiple maxima or minima in the deuterium record, we used the synthetic Greenland temperature curve GL_{T_syn} from Barker et al. (2011) calculated from the deuterium record of ice at EDC. This curve was constructed by a mathematical process which involves separation of the high and low frequency components of the Antarctic temperature record and differentiation of the high frequency component before it is recombined with the low frequency component. Identifying an extremum in the EDC deuterium record thus corresponds to identifying a fast transition in the GL_{T_syn} curve. The fact that Barker et al. (2011) were able to reconstruct a curve from an Antarctic ice isotope record which resembles the Greenland ice isotope record and in particular exhibits similar fast transitions is another congruent proof for the validity of the seesaw mechanism.

In Fig. 7, we compare the GL_{T_syn} curve of Barker et al. (2011) and the EDC deuterium record of Jouzel et al. (2007) with the CH_4 record from EDC (Louergue et al.,

Table 3. CH_4 synchronisation tie points between EDC and TALDICE and corresponding Δ depth estimates at EDC.

Description	EDC gas depth (m)	Talos gas depth (m)	Δ depth (m)	1σ uncertainty (m)
8.2 kyr event	323	553.0	64.5	4.19
YD-Holo	418.5	695.4	64.0	4.29
BO/YD	443.5	736.8	61.1	4.07
onset BA	476.1	785.5	57.4	4.06
onset DO3	639.1	941.0	55.9	3.72
onset DO4	651.9	953.0	58.3	3.81
onset DO5	688.1	990.0	52.9	3.55
onset DO6	702.1	1004.0	53.6	3.53
onset DO7	719.7	1025.0	50.0	3.29
onset DO8	751.3	1059.2	50.0	3.43
onset DO9	776	1085.0	55.4	3.57
onset DO10	791.0	1098.6	56.7	3.63
onset DO11	809.2	1125.0	53.2	3.56
onset DO12	848.5	1170.7	48.4	3.44
onset DO13a	862.0	1183.0	46.4	3.64
onset DO13b	876.0	1192.0	48.7	4.37
onset DO14	925.0	1229.0	44.9	3.45
onset DO15	945.0	1239.4	48.2	3.58
onset DO17	986.5	1263.3	42.0	3.05
onset DO18	1038.5	1286.7	42.4	3.20
onset DO19	1105.0	1307.2	38.3	3.52
onset DO20	1142.0	1315.1	37.3	3.67
onset DO21	1239.0	1333.5	35.9	4.89
onset DO23	1431.5	1369.0	34.4	6.63
onset MIS5.5	1722.0	1411.2	28.8	3.42

2008) on a depth scale. Using these constraints, 82 tie points (see Table 4) are derived between the two records, mainly at times of fast variations in Greenland temperature. These tie points correspond to maxima and minima in the EDC deuterium record (see Fig. 7). Δ depth estimates are simply computed as the depth of the transition in the methane record minus the depth of the transition in the GL_{T_syn} curve (or equivalently to the depth of the maxima or minima in the deuterium record).

There are two sources of error in this procedure. First, the identified transitions in GL_{T_syn} and CH_4 may not correspond to the same event. We therefore tagged the pairs of tie points as “virtually certain” or “tentative”. Second, even if the transitions in GL_{T_syn} and CH_4 correspond to the same event, there is an error linked to the determination of the depth of the transitions in both curves. To evaluate this (2σ) error of these Δ depth estimates, we added the error of the depth estimates of the transition in the methane and GL_{T_syn} curves, respectively. These errors are evaluated as half of the duration of the transition.

The reasons why we used GL_{T_syn} and not the raw deuterium record are: (1) it is easier and more accurate to select a mid-transition than an extremum, and (2) it is also easier to estimate the error in the tie point picking as half of the duration of the transition.

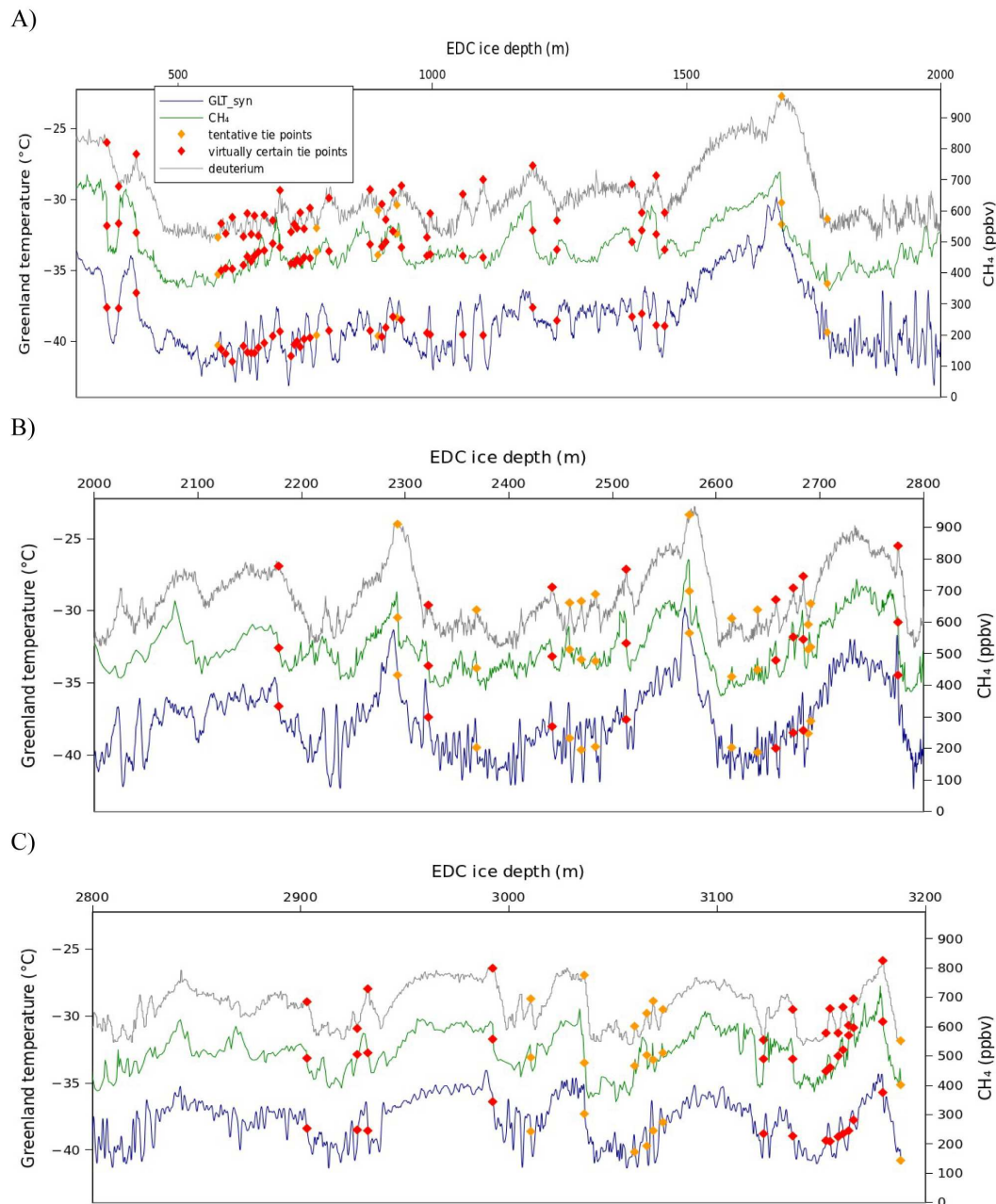


Fig. 7. Comparison between the GLT_{syn} curve (Barker et al., 2011), the EDC CH_4 record (Louergue et al., 2008) and the EDC δD record (Jouzel et al., 2007): **(A)** in the depth interval 0–2000 m, **(B)** in the depth interval 2000–2800 m, **(C)** in the depth interval 2800–3200 m. The deuterium record has been resampled on 100-yr intervals.

3 Discussion

In Fig. 8, we compare the various Δdepth estimates derived previously. Several conclusions can be outlined.

3.1 Confirmation of the bipolar seesaw antiphase

The GRIP-synchro based estimate of Δdepth during the Laschamp geomagnetic event is the most accurate (± 2 m)

and robust (it does not rely on controversial hypotheses) estimate available among the EDC ice core data. It is important to note that this estimate is fully compatible with the bipolar seesaw-based estimate. In other words, the seesaw phasing between EDC and GRIP is observed during the Laschamp geomagnetic event, as was already concluded by Raisbeck et al. (2007).

A second noteworthy remark is that the EDML ice core record mainly confirms the seesaw hypothesis, as was also

Table 4. Ice-depth/gas-depth synchronisation based on the seesaw method and corresponding Δ depth along the EDC ice core.

Description	Ice depth (m)	Gas depth (m)	Δ depth (m)	1σ (m)	Tentative or not
YD-Holo	359.82	418.50	58.68	5.12	
BA-YD	383.36	443.50	60.14	4.22	
onset BA	418.83	476.10	57.27	4.86	
end DO3	578.22	627.00	48.78	2.66	tentative
onset DO3	585.44	639.10	53.66	2.34	
end DO4	593.79	647.00	53.21	2.50	
onset DO4	607.26	651.90	44.64	2.10	tentative
end DO5	629.07	681.00	51.93	2.36	
onset DO5	636.90	688.10	51.20	1.80	
end DO6	644.07	696.00	51.93	1.95	
onset DO6	651.06	702.10	51.04	1.72	
end DO7	658.55	712.00	53.45	2.12	
onset DO7	669.51	719.70	50.19	2.59	
end DO8	686.46	732.50	46.04	3.05	
onset DO8	700.51	751.30	50.79	1.90	
end DO9	722.40	773.30	50.90	2.24	
onset DO9	729.37	776.00	46.63	1.46	
end DO10	733.82	784.20	50.38	1.68	
onset DO10	740.83	791.00	50.17	1.42	
end DO11	748.32	800.00	51.68	2.50	
onset DO11	759.81	809.20	49.39	1.98	
end DO12	772.34	823.00	50.66	2.83	tentative
onset DO12	797.00	848.50	51.50	2.76	
onset DO14	878.00	925.00	47.00	1.90	
end DO15	894.00	935.00	41.00	1.80	tentative
onset DO15	901.00	945.00	44.00	1.25	
end DO16	909.00	953.00	44.00	2.12	
onset DO16	923.00	969.00	46.00	2.12	
end DO17	931.00	974.00	43.00	2.12	tentative
onset DO17	940.00	986.50	46.50	1.90	
end DO18	990.50	1030.00	39.50	2.30	
onset DO18	997.00	1038.50	41.50	1.72	
onset DO19	1061.00	1105.00	44.00	1.85	
onset DO20	1101.00	1142.00	41.00	2.27	
onset DO21	1197.50	1239.00	41.50	2.50	
end DO22	1245.50	1281.00	35.50	1.95	
onset DO23	1393.00	1431.50	38.50	1.80	
end DO24	1413.00	1445.00	32.00	2.50	
onset DO24	1441.00	1473.00	32.00	1.95	
end DO25	1458.00	1485.00	27.00	2.12	
onset MIS5.5	1687.00	1722.00	35.00	3.64	tentative
	1777.50	1812.00	34.50	2.02	tentative
	2178.00	2194.00	16.00	2.12	
	2293.00	2307.00	14.00	1.58	tentative
	2322.50	2337.00	14.50	1.60	
	2369.00	2380.00	11.00	1.25	tentative
	2442.00	2454.00	12.00	0.90	
	2459.00	2470.00	11.00	1.12	tentative
	2470.20	2480.50	10.30	0.71	tentative
	2484.00	2493.50	9.50	0.71	tentative
	2513.30	2523.30	10.00	0.71	
	2574.50	2584.70	10.20	1.30	tentative
	2615.70	2623.30	7.60	0.38	tentative
	2640.30	2647.70	7.40	0.25	tentative
	2658.10	2665.50	7.40	0.39	

Table 4. Continued.

Description	Ice depth (m)	Gas depth (m)	Δ depth (m)	1σ (m)	Tentative or not
	2674.80	2682.70	7.90	0.49	
	2684.20	2691.00	6.80	0.43	
	2688.80	2694.90	6.10	0.57	tentative
	2691.70	2697.00	5.30	0.61	tentative
	2775.50	2780.70	5.20	0.71	
	2903.00	2907.00	4.00	0.71	
	2927.50	2935.00	7.50	1.12	
	2932.30	2941.00	8.70	0.64	
	2992.50	2997.80	5.30	0.64	
	3010.50	3013.20	2.70	0.35	tentative
	3036.50	3038.50	2.00	0.71	tentative
	3060.70	3064.90	4.20	0.29	tentative
	3066.30	3070.30	4.00	0.46	tentative
	3069.30	3074.30	5.00	0.43	tentative
	3074.00	3078.70	4.70	0.40	tentative
	3122.40	3124.00	1.60	0.35	
	3136.50	3139.50	3.00	0.56	
	3152.50	3157.50	5.00	0.35	
	3154.50	3159.10	4.60	0.35	
	3158.20	3162.20	4.00	0.57	
	3160.70	3164.40	3.70	0.43	
	3163.30	3167.00	3.70	0.61	
	3165.70	3169.00	3.30	0.35	
	3179.70	3182.60	2.90	0.39	

stated recently (EPICA community members, 2006; Capron et al., 2010). One can, however, remark that the EDML synchro estimates tend to underestimate Δ depth during the last glacial period with respect to the seesaw-based estimates by ~ 2 – 3 m in average, probably resulting from an underestimation of Δ depth at EDML, as systematic offsets in both the CH_4 and volcanic synchronisations are unlikely. An underestimated LID by the densification model is also unlikely because $\delta^{15}\text{N}$ data shows the contrary (Landais et al., 2006). It thus leaves us only with an underestimation of EDML thinning, which may be due to an overestimation of EDML accumulation rates during the glacial. We indeed did not take into account the fact that accumulation rates are lower upstream of the EDML site from where the ice supposedly originates (Huybrechts et al., 2007).

A third remark, arising from the present study, is that the TALDICE ice core records confirm the seesaw hypothesis. Contrary to the EDML ice core, we did not observe for the TALDICE ice core a systematic offset with respect to the seesaw-based estimates. As a consequence, the TALDICE-synchro method overestimates Δ depth with respect to the EDML-synchro method.

In the detail, there are EDML-based or TALDICE-based estimates which significantly deviate from the seesaw-based estimates. Note that we used exactly the same depths for the CH_4 transitions in all three methods. At ~ 651.90 m (onset of DO4), the seesaw-based estimate is very small

(44.64 m) compared to the EDML-based and TALDICE-based estimates. The corresponding maxima in the deuterium curve is ambiguous and it is why this tie point has been tagged as “tentative”. Another possible explanation is that EDC3 underestimates the duration of events in this interval, leading to overestimated thinning function at EDML and TALDICE. At 809.2 m (onset DO11) and 848.5 m (onset DO12), the TALDICE-based estimates are very small (53.23 m and 48.42 m) compared to the EDML-based and seesaw-based estimates. We note however that we are here beyond the EDC–TALDICE volcanic synchronisation, so this discrepancy can come from a poor EDC–TALDICE ice synchronisation. At 1105 m, 1142 m, 1239 m, 1431.5 m and 1473 m (onsets DO19, 20, 21, 23 and 24), the seesaw-based estimates are systematically higher than the EDML-based and TALDICE-based estimates. One possible explanation is that EDML and TALDICE Δ depths are underestimated due to an overestimation of durations in EDC3.

3.2 The “glacial Δ depth paradox at EDC”

Focusing now on the modelling estimates of Δ depth during the last glacial period (the last deglaciation and the last glacial period), they are on average $\sim 15\%$ larger than the seesaw-based estimates or the EDML-synchro and TALDICE-synchro based estimates. It is very likely that the modelled Δ depth is inaccurate during this time period. We

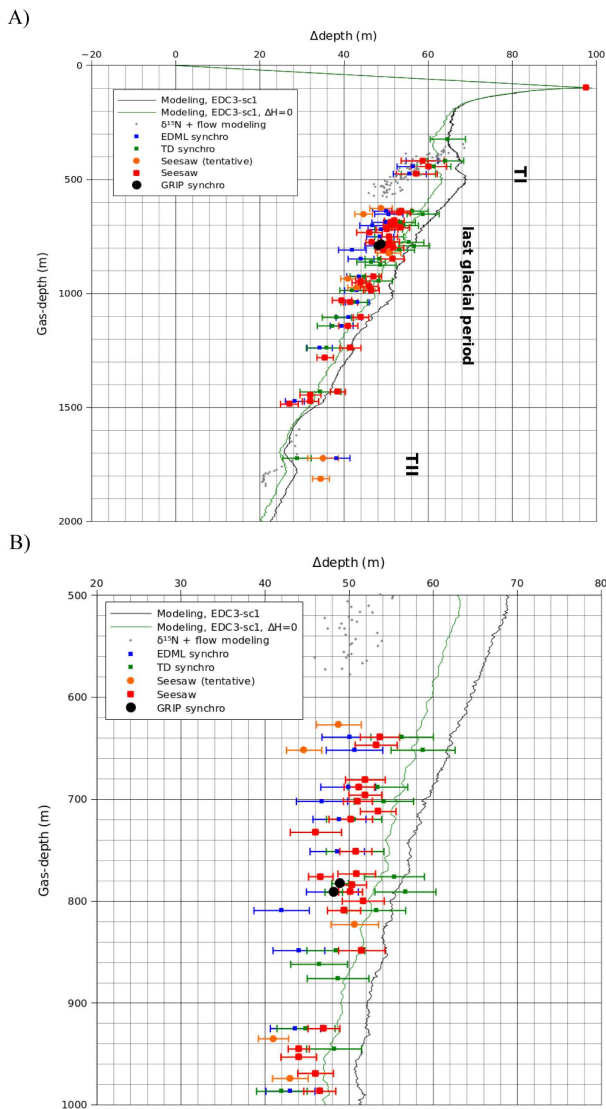


Fig. 8. Δdepth estimates along the EDC ice core. **(A)** Zoom on the 0–2000 m depth interval, which shows the glacial and termination II Δdepth paradoxes. **(B)** Zoom on the 500–1000 m depth interval encompassing the Laschamp event. The error bars have to be taken as 1σ uncertainty.

call this model-data discrepancy the “glacial Δdepth paradox at EDC”. This may be due either to an overestimation of the thinning function or to an overestimation of the LID.

One factor influencing the thinning function is the existence of non-laminar flow effects. However, the amplitude of the Δdepth paradox is greater than the uncertainty on the thinning function due to non-laminar flow effects (Fig. 2).

Another factor which has a significant impact on the thinning function in the upper part of the ice sheet is the past variation of ice thickness. As previously explained, apart from the EDC3 scenario of past ice thickness variations, we tested a scenario without ice thickness variations as an extreme case

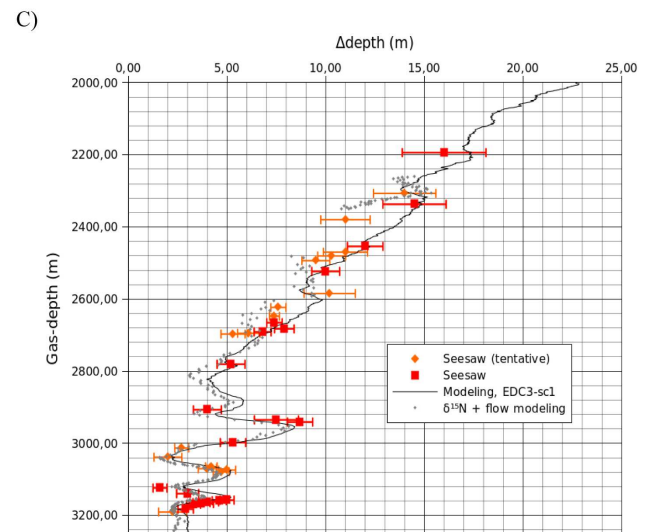


Fig. 8. (C) Zoom on the 2000–3200 m depth interval.

(see Fig. 8). This scenario only partially solves the glacial Δdepth paradox at EDC.

The last factor that has a significant impact on the thinning function in the shallow part is the ice thickness at the site of snow deposition. The glacial Δdepth paradox at EDC could be solved if one assumes that the ice flow was not vertical in the past and that the ice originates from a site with greater ice thickness. This hypothesis is difficult to test from a modelling point of view because there are many unknown parameters in 3-D ice flow models of the Antarctic ice sheet influencing the position of the ridges and of the domes. We however remark here that the glacial Δdepth paradox at EDC only concerns the glacial part and is not present (and is even inverted) for the Eemian ice. This paradox therefore seems to have a climatic origin. Our conclusion is thus that, for reasons which are beyond the scope of the present manuscript, the firn densification model overestimates the glacial LID at EDC. A possible explanation is the effect of impurities on the densification process (Hörhold et al., 2012).

This remark is in contradiction with the conclusions of Caillon et al. (2003), who stated that the firn densification model correctly estimated Δage at Vostok during Termination III. However, Caillon et al. (2003) based their conclusions on the assumption that, at Vostok, $\delta^{40}\text{Ar}$ varies in phase with temperature, which has never been demonstrated. If, as proposed by Hörhold et al. (2012), the densification velocity (and thus the gravitational fractionation) is influenced by impurities, $\delta^{40}\text{Ar}$ should be better correlated with the impurity record than with the ice isotopic record when both are not in phase (Röthlisberger et al., 2008). Moreover, their conclusion suffers from the poorly known thinning function at Vostok at the depth corresponding to Termination III.

3.3 The “Termination II Δdepth paradox at EDC”

During Termination II, we have two different estimates of the Δdepth (see Fig. 8a, depth interval 1700–1800 m). On one hand, the model-based estimate suggests a relatively low Δdepth . On the other hand, the seesaw method roughly agrees with the EDML-synchro and TALDICE-synchro methods and suggests a relatively high Δdepth (we should note however that the two seesaw points are only tentative at this stage). We call this discrepancy the “Termination II Δdepth paradox at EDC”.

There are two possible explanations for this discrepancy. Either the seesaw phenomenon is not at work during the penultimate deglaciation, unlike the last glacial period, and the EDML-synchro and TALDICE-synchro methods are not precise during this time period; or the modelled estimates are too low.

One possibility to reconcile the model with the seesaw-based estimates would be to increase the thinning function, which would have the side effect to decrease the duration of the penultimate interglacial in EDC and to give a better agreement with the duration of this stage in the Dome Fuji ice core (Parrenin et al., 2007b). Another possibility would be to increase the LIDIE, for example by assuming that the surface temperature has been underestimated during this time period. It is indeed not clear that the deuterium-temperature spatial relationships used remain valid for climates warmer than the present (Sime et al., 2009).

3.4 Using Δdepth estimates in the deepest part to improve the EDC age scale

For the depth interval 2000–2800 m, Δdepth estimates regularly decrease from ~ 16 m to ~ 5 m. In this depth interval, the agreement between the seesaw-based and model-based estimates is surprisingly good. The fact that the model does not systematically overestimate the Δdepth in this depth interval, contrary to the last glacial period, may be just a coincidence: an overestimated LIDIE may be exactly compensated by an underestimated thinning function.

For the depth interval 2800–3200 m, the agreement is also good. In particular, the seesaw-based Δdepth estimates reproduce well the bumps in the thinning function, which have been suggested based on the comparison between $^{18}\text{O}_{\text{atm}}$ measurements and insolation variations and based on the phasing between δD and CO_2 (Dreyfus et al., 2007). We thus independently confirm the hypothesis that the flow in the bottom part of the EDC ice core is complex.

There are however several seesaw-based estimates which deviate significantly from the model-based estimates at ~ 2360 m (tentative tie point, ~ 260 kyr), ~ 2900 m (~ 527 kyr), ~ 2930 m (554 kyr), ~ 3000 m (~ 585 kyr) and ~ 3120 m (~ 719 kyr). This is somewhat in agreement with results by Landais et al. (2012) based on the tuning of the O_2/N_2 record on local insolation variations. They indeed

suggested corrections of up to ~ 5 kyr for the time period 390–460 kyr, 550 kyr, 650 kyr and 750 kyr BP.

In conclusion, we thus suggest that using these seesaw-based estimates associated with new O_2/N_2 measurements could improve the reconstruction of the thinning function and thus the evaluation of the EDC chronology in the deep part.

3.5 Validity of the $\delta^{15}\text{N}$ firn thickness estimate for the last deglaciation

The $\delta^{15}\text{N}$ record in association with the thinning model gives an evaluation of the Δdepth decreasing from 67 m to 45 m during the course of the last deglaciation. In this upper part of the EDC ice core, the uncertainty in the thinning function is thought to be small (Fig. 2). These Δdepth estimates are in good agreement with the estimates based on the synchronisation to EDML and TALDICE or based on the seesaw hypothesis (Fig. 8). If there are no other compensating effects, the convective zone cannot be more than a few metres at EDC during the last deglaciation. Consequently, we conclude that the model- $\delta^{15}\text{N}$ data mismatch observed at EDC during the last deglaciation (Dreyfus et al., 2010) probably results from an incorrect representation of the densification process in firn models (Hörhold et al., 2012), and not to a varying convective height or to poorly known $\delta^{15}\text{N}$ fractionation processes (Dreyfus et al., 2010).

This conclusion seems in contradiction with a study on the Vostok ice core using ice and gas synchronisation to Byrd (Bender et al., 2006), which concluded that $\delta^{15}\text{N}$ underestimates the LID during the last glacial period. However, concerning this study, we note that: (1) there is no estimate of uncertainty in the Byrd-synchro method; (2) the thinning function reconstruction is a lot more uncertain than in our case because the coring point at Vostok is not situated on a dome (Parrenin et al., 2004); (3) the ice synchronisation is less precise than in our study because it is based on the ice isotope data; and (4) the gas synchronisation is also less precise because it is based on a smaller number of CH_4 measurements.

Consequently, the $\delta^{15}\text{N}$ data seems to be a more appropriate tool as compared to the current densification models used to estimate the LID during the last deglaciation at EDC. Is this conclusion valid for other time periods and for other Antarctic sites where a model- $\delta^{15}\text{N}$ data mismatch has been observed, such as Vostok (Sowers et al., 1992; Caillon et al., 2003), EDML (Landais et al., 2006), Law Dome (Landais et al., 2006) and Dome Fuji (Severinghaus et al., 2010, based on the data by Kawamura, 2000)? Further studies following the approach presented here are needed to answer this question.

4 Conclusions

We have shown that the bipolar seesaw antiphase relationship is generally supported by the ice–gas cross synchronisation of EDC to the GRIP, EDML and TALDICE ice cores. The

glaciological model overestimates the glacial Δ depth at EDC (we called this the “glacial Δ depth paradox at EDC”) and this is probably due to an overestimation of the glacial close-off depth by the firn densification model. The glaciological models seem to underestimate the Δ depth during Termination II (we called this the “Termination II Δ depth paradox at EDC”). We have shown that the bipolar seesaw hypothesis confirms that the ice flow is complex in the deep part of the EDC ice core and can help improve the EDC age scale. For the last deglaciation, using $\delta^{15}\text{N}$ data in association with an ice flow model gives Δ depth estimates in agreement with the estimates based on the synchronisation to TALDICE and EDML or based on the seesaw method.

Complete, precise and highly resolved $\delta^{15}\text{N}$ and CH_4 records will be necessary to further improve the EDC gas and ice age scales. An automatic method to synchronise records would both bring rigour and shorten the time to accomplish this tedious task. Further studies will be needed to make the firn densification models more useful for paleoclimatic studies using the Antarctic ice core records. Both the firn modelling and $\delta^{15}\text{N}$ approaches need a precise evaluation of the past surface and lock-in densities and further studies are needed to better constrain them. ^{10}Be measurements are in progress and should allow to extend the Antarctic–Greenland ice synchronisation and thus produce more Δ depth estimates based on this hypothesis-free approach. Studies on the Δ depth comparable to the present one could be applied to other low accumulation Antarctic ice cores such as Vostok and Dome Fuji. This study on the gas/ice depth offset at EDC has important implications on the phasing between CO_2 and Antarctic temperature during climatic changes and consequently on the role of CO_2 during these climatic changes.

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