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## Size Distribution and Optical Properties of Ambient Aerosols during Autumn in Orleans, France

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### ABSTRACT

A new highly sensitive cavity ring-down spectrometer (CRDS) system was designed and assembled to determine the aerosol extinction coefficient ( $b_{\text{ext}}$ ) at 532 nm. The performance of the CRDS was tested by the monodisperse polystyrene latex spheres (PSL) particles with diameters between 200 and 500 nm. By comparing the tested results with Mie theory curve, the uncertainty of the newly developed CRDS system was determined to be < 3%. Simultaneous measurements of the size distribution and extinction coefficient of ambient aerosols were conducted in Orleans, France, from 26<sup>th</sup> October to 21<sup>st</sup> December 2012 by using a scanning mobility particle sizer (SMPS) coupled to the CRDS system. For the non-dehydrated aerosols measured from 26<sup>th</sup> October to 4<sup>th</sup> November, the average  $b_{\text{ext}}$  has been found to be  $41 \pm 35 \text{ Mm}^{-1}$ . For the dehydrated aerosols measured from 7<sup>th</sup> November to 21<sup>st</sup> December,  $b_{\text{ext}}$  expresses a good agreement with the particle number (N) and volume (V) concentration, the average values of  $b_{\text{ext}}$ , N and V are  $36 \pm 31 \text{ Mm}^{-1}$ ,  $3300 \pm 2700 \text{ cm}^{-3}$  and  $3.1 \pm 2.8 \times 10^9 \text{ nm}^3/\text{cm}^3$ , respectively. Further analysis of the particle size distribution reveals that car and boiler emissions maybe the main aerosol sources in Orleans. In addition, back trajectory results indicate that the air parcel transported from Atlantic Ocean may play a role in cleaning up the ambient air in Orleans.

**Keywords:** CRDS; Optical properties; Size distribution; Orleans-France.

### INTRODUCTION

Aerosol particles influence climate both directly, by scattering and absorbing solar radiation (Charlson *et al.*, 1992; Haywood and Shine, 1997), and indirectly, by acting as cloud condensation nuclei (CCN) with subsequent effects on the albedo and lifetime of clouds (Kaufman *et al.*, 2005). In addition, aerosol particles affect visibility (Vis) and air quality, such as the haze and dust storm events (Fu *et al.*, 2010; Han *et al.*, 2012; Huang *et al.*, 2012; Wang *et al.*, 2012; Yang *et al.*, 2012; Zhang *et al.*, 2012; Kang *et al.*, 2013; Park and Cho, 2013; Park *et al.*, 2013). Different investigations suggest that the visibility, as the most readily

perceived impact of air pollution, is highly dependent on the particle size and optical properties, and can be calculated by particle extinction coefficient ( $b_{\text{ext}}$ ), i.e.,  $\text{Vis} = 3.912/b_{\text{ext}}$  (Seinfeld and Pandis, 2006; Li *et al.*, 2011).

Because of the evidence of their role in climate change and visibility, the study of the optical properties of atmospheric aerosols has received an enormous interest in the last years. Recently, cavity ring-down spectroscopy (CRDS) has been introduced as a robust and complimentary method to in situ measurement of the aerosols extinction coefficient (Smith and Atkinson 2001; Thompson *et al.*, 2002; Strawa *et al.*, 2003; Pettersson *et al.*, 2004; Moosmuller *et al.*, 2005; Strawa *et al.*, 2006; Baynard *et al.*, 2007; Riziq *et al.*, 2007; Lang-Yona *et al.*, 2009; Radney *et al.*, 2009; Sakamoto *et al.*, 2009; Li *et al.*, 2011; Qiu *et al.*, 2012; Khalizov *et al.*, 2013). CRDS is not only used in laboratory to measure the optical properties of inorganic aerosols (Bulatov *et al.*, 2002; Riziq *et al.*, 2007), secondary organic aerosols (Nakayama *et al.*, 2010b; Redmond *et al.*, 2011; Nakayama *et al.*, 2012; Zarzana *et al.*, 2012), and soot-reacted particles (Zhang *et al.*

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*al.*, 2008; Xue *et al.*, 2009; Khalizov *et al.*, 2009; Qiu, *et al.*, 2012; Khalizov *et al.*, 2013), but has also been employed for field campaigns in many different places of the world. For example, Nakayama *et al.* (2010a) measured the aerosol optical properties in central Tokyo during summertime using CRDS, a nephelometer and a particle/soot absorption photometer (PSAP). The average  $b_{\text{ext}}$ ,  $b_{\text{sca}}$  (scattering coefficient) and  $b_{\text{abs}}$  (absorption coefficient) at 532 nm were determined to be  $144.7 \pm 85.2$ ,  $130 \pm 81.4$  and  $13.6 \pm 9.2$   $\text{Mm}^{-1}$ , respectively, and  $b_{\text{ext}}$  expressed a good agreement with  $b_{\text{sca}} + b_{\text{abs}}$  values during this observations. In the United States, the same phenomenon, i.e., excellent agreement between  $b_{\text{ext}}$  (CRDS) and that derived from combining nephelometer scattering and PSAP absorption ( $b_{\text{sca}} + b_{\text{abs}}$ ), was also observed in Oklahoma (Strawa *et al.*, 2006). For other cities of the United States, Langridge *et al.* (2011) designed a multi-wavelength 8-channels CRDS system and measured the multi-wavelength aerosol extinction under dry and elevated RH conditions in California. Thompson *et al.* (2002), by measuring the aerosol optical properties in Florida from 4<sup>th</sup> May to 3<sup>rd</sup> July, 2001 have reported the average extinction coefficient of  $80.4 \text{ Mm}^{-1}$  for 510 nm and  $87.0 \text{ Mm}^{-1}$  for 578 nm. In China, the first CRDS system used to investigate the aerosol optical properties was designed and assembled by Li *et al.* (2011). They found that the maximum aerosol extinction coefficient reached up to  $2 \times 10^3 \text{ Mm}^{-1}$  from 28<sup>th</sup> May to 2<sup>nd</sup> June, 2009 in Shanghai and that the visibility derived from CRDS ( $\text{Vis} = 3.912/b_{\text{ext}}$ ) was in good agreement with that reported by Shanghai Meteorological Bureau. Recently, the same authors reported that the particle water-soluble inorganic ions (WSI), especially  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$ ,  $\text{Cl}^-$  and  $\text{NH}_4^+$ , have a good correlation with scattering coefficient in urban Shanghai (Li *et al.*, 2013).

As mentioned above, although the CRDS system has been extensively used in the world, to our knowledge, most of aerosol optical studies conducted in Europe were performed

using the conventional techniques, such as: nephelometer, PSAP and Aethalometer (Vrekoussis *et al.*, 2005; Bryant *et al.*, 2006; Kalivitis *et al.*, 2011; Esteve *et al.*, 2012), hence, up to now, there are no publications about the use of CRDS to measure the extinction coefficient of ambient aerosols in Europe. Since the aerosol extinction coefficient is a function of the particle size, simultaneous observations of the particle size distribution and extinction coefficient will provide a good way for a better understanding of the optical properties of aerosols.

In this study, a pulsed CRDS instrument for measuring extinction coefficient of atmospheric aerosols was developed. First, the instrument was validated using monodisperse laboratory-generated polystyrene latex spheres (PSL) particles. It was then coupled to a SMPS for simultaneous measurements of the extinction coefficient and size distribution at the CNRS campus in Orleans-France during the autumn of 2012.

## EXPERIMENTAL

### Sampling Site and Period

The measurements presented in this study were made during the period 26<sup>th</sup> October–21<sup>st</sup> December, 2012 at the campus of the Centre National de la Recherche Scientifique (CNRS) in Orleans, France ( $47^\circ 50' 16.80''\text{N}$ ,  $1^\circ 56' 39.34''\text{E}$ ). As shown in Fig. 1, Orleans city is located in the central France, about 120 km southwest of Paris. Typical average concentrations in Orleans for  $\text{O}_3$  ( $52.8 \pm 12.7 \mu\text{g}/\text{m}^3$ ),  $\text{NO}_2$  ( $24.0 \pm 14.4 \mu\text{g}/\text{m}^3$ ),  $\text{NO}$  ( $4.3 \pm 3.4 \mu\text{g}/\text{m}^3$ ),  $\text{PM}_{2.5}$  ( $18.0 \pm 5.9 \mu\text{g}/\text{m}^3$ ), and  $\text{PM}_{10}$  ( $13.1 \pm 5.9 \mu\text{g}/\text{m}^3$ ) have been measured by LigAir (<http://www.ligair.fr/>) from May 2012 to March 2013. Our sampling site, positioned at the CNRS campus, about 8 km south of the Orleans city center, is mostly surrounded by residential dwellings. There are no obstructing buildings around the sampling site within 50 m. During the



Fig. 1. Location of the sampling site in Orleans, France.

measurement, CRDS and SMPS sampled the ambient aerosols from an inlet located 1 m above the roof of the site (~3.5 m above the ground).

This study has been conducted during two different periods under different conditions. The first one, the non-dehydrated condition was from 26<sup>th</sup> October to 4<sup>th</sup> November where the extinction coefficient was measured from the ambient air. This represents ‘real’ atmospheric conditions for the aerosols and the resulting extinction coefficient. The data recorded by the weather station at the same sampling site reveal that the RH of the ambient air during this period was in the range 0.54–1.00 with a mean value of 0.86. The second period, the dehydrated condition was from 7<sup>th</sup> November to 21<sup>st</sup> December, where the study was conducted to investigate the particle size distribution effect on the aerosol extinction coefficient after removing the interference from relative humidity. In this study, the sampled ambient aerosols were flowed through a silica drier before entering into the CRDS and SMPS. During this measurement, the dessicant was replaced 2 times per week and the RH of the sample before and after changing the dessicant was monitored for each replacement. The recorded data reveal that the RH of the dehydrated aerosol was less than 15%.

### Experimental setup

The experimental setup consists of a combination of a SMPS and a CRDS. The SMPS has been described in detail elsewhere (Heim *et al.*, 2004). Therefore, only a brief description is presented here. In the SMPS system used in this work (TSI Inc., USA), the particles are firstly neutralized with Kr<sup>85</sup> diffusion charger and size selected using a differential mobility analyzer (DMA, Model 3081). The selected monodisperse particles subsequently goes to a butanol-based condensation particle counter (CPC, Model 3022A), which determines the particle concentration at that size. In this study, the detectable aerosol mobility diameters range from 10.7 to 487.3 nm with a sheath and sample flow rates of 5 LPM and 0.5 LPM, respectively. The total counting time required for each sample is about 130 s, consisting of a 120 s up scan and a 10 s retrace. Before measurements,

SMPS has been calibrated by the PSL particles.

The CRDS system is shown schematically in Fig. 2. The instrument constructed in our group is similar in design to that developed by Li *et al.* (2011). Laser light at 532 nm, generated by a pulsed Nd:YAG laser (EXPL-532-200-E, Spectra-Physics Inc., USA), is firstly introduced into a cavity cell and then bounced numerous times between two highly reflective mirrors (532 nm center wavelength, 99.997% reflectivity, 1m radius of curvature, ATFilms Inc., USA). Each time the pulse interacts with the back mirror, a small amount of light (e.g., 0.003%) leaks out. This light is collected and detected with a photomultiplier tube (PMT, R928, Hamamatsu Inc., Japan). Data acquisition was performed using a 100 MHz Digital card (NI PXIe-1073, USA) and control programs have been written in LabVIEW software. Since the intensity of the transmitted light is proportional to that of the trapped light, the time decay of the light intensity inside the cavity can be determined by monitoring the transmitted signal. The decay time is calculated by using non-linear fitting procedure. The temporal resolution was 1 ms and the data was averaged for every 10 s. After obtaining the decay time of the cavity filled with ( $\tau$ ) and without ( $\tau_0$ ) aerosols, the aerosol extinction coefficient ( $b_{\text{ext}}$ ) can be calculated according to the following equation:

$$b_{\text{ext}} = \frac{L}{cl} \left( \frac{1}{\tau} - \frac{1}{\tau_0} \right) \quad (1)$$

where  $L$  is the length between the mirrors (m),  $l$  is the effective sample path length (m),  $c$  is the speed of light (m/s). To ensure good mixing and even concentration of the particles inside the cavity, the aerosol flow enters the cavity from the two inlets near the mirrors and exits the cavity from the outlet in the middle of the cavity. Meantime, to prevent mirror contamination by deposition of aerosols, a small purge flow of dry particle-free nitrogen (0.03 LPM) is introduced in front of each mirror. In our CRDS system, the distance between the two mirrors is 68 cm and the length of the cavity occupied by particles during the flow is about 63 cm. The minimal detectable extinction coefficient ( $b_{\text{ext min}}$ ) can be calculated by the following equation:

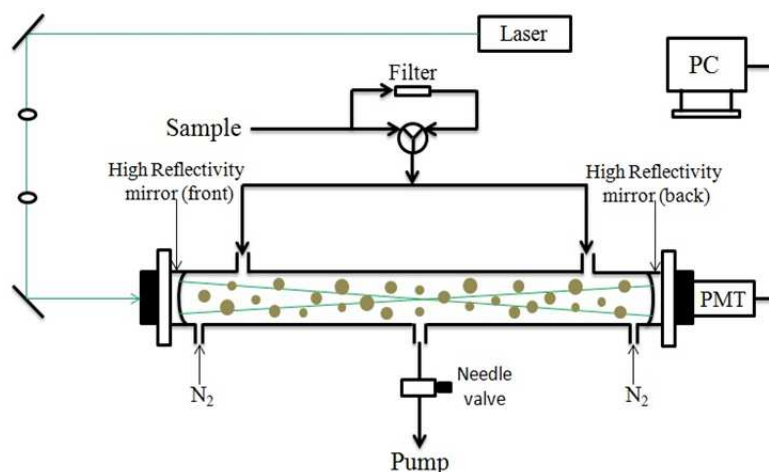


Fig. 2. Schematic diagram of the cavity ring-down spectroscopy system.

$$b_{\text{ext min}} = \frac{L\Delta\tau_{\text{min}}}{c\tau_0^2} \quad (2)$$

where  $\Delta\tau_{\text{min}}$  is the detection limitation of the decay time. During the experiment,  $\tau_0$  is 30  $\mu\text{s}$  and the  $\Delta\tau_{\text{min}}$  is 0.1  $\mu\text{s}$ , resulting in the  $b_{\text{ext min}} = 4.0 \times 10^{-7} \text{ m}^{-1}$ .

During the experiment, the aerosol flow was 2 LPM, resulting in a residence time of around 32 s. In addition, to avoid the interferences from  $\text{NO}_2$  and other absorbing trace gas species, the blank checks of the CRDS system were performed automatically by turning the 3-way electric valves to the filter direction to measure the  $b_{\text{ext}}$  of the particle-free ambient air every 2 h.

### Validation of the CRDS System

To test the performance of the CRDS system developed during the present work, the optical properties of the monodisperse PSL (Duke Inc.) particles with diameters of 200, 269, 300, 400, 450 and 500 nm were monitored. During the calibration experiment, PSL particles produced by the aerosol generator (Model 3076, TSI Inc., USA), are firstly dried to a low RH (RH < 15%) and then fed into the SMPS system in order to select PSL particles of a certain “dry” size. Afterwards the size-selected PSL particles are first drawn into CRDS and then to the CPC (Model 3022A, TSI Inc., USA) to measure the  $b_{\text{ext}}$  and number concentration respectively. To expand the detectable particle size range of SMPS system, the PSL particle sample was diluted with filtered air before entering into the CPC. In addition, for each calibration experiment the PSL number concentration at the CRDS inlet and outlet were measured by CPC to calculate the loss of PSL when they pass through the CRDS cavity. The PSL number concentration used in this paper has been corrected by the dilution rate and the loss of PSL along the cavity.

As shown in Fig. 3(a), the measured extinction coefficients show a linear fitting with the number concentration for each PSL particles. According to Fig. 3(a), the extinction efficiency ( $Q_{\text{ext}}$ ) of PSL particles at 532 nm can be determined by the following equation:

$$Q_{\text{ext}} = \frac{b_{\text{ext}}}{N \times (\pi D^2 / 4)} \quad (3)$$

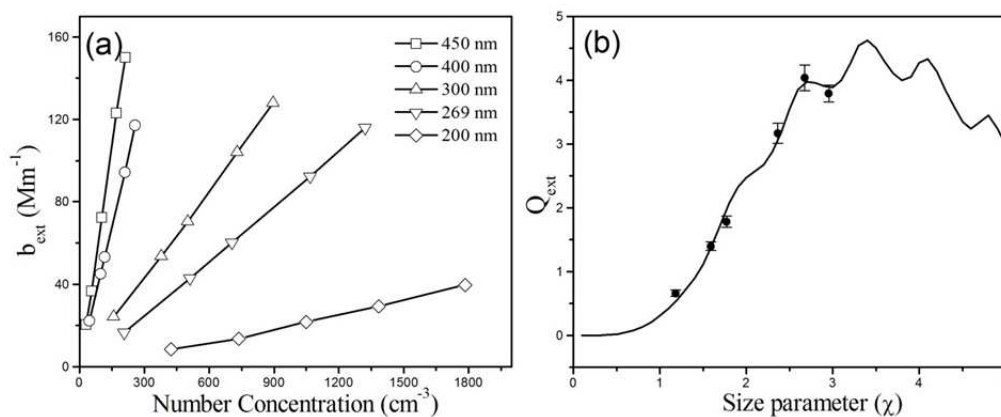
where  $N$  and  $D$  are the particle number concentration and particle diameter, respectively.

Fig. 3(b) shows the  $Q_{\text{ext}}$  of PSL particles as a function of size parameter  $\chi$  ( $\chi = \pi D/\lambda$ ). Points are determined using the Eq. (3) and the curve is obtained from Mie theory using refractive index of  $1.598 + 0.00i$  reported by Boundy and Boyer (1952). As expected, the measured extinction efficiency values are in good agreement with the theoretical curve, the average deviation between the measured and theoretical  $Q_{\text{ext}}$  is within  $\pm 3\%$ . This result indicates that our CRDS system, if we do not consider the uncertainty in the size of particles and the particle number density, could measure an extinction coefficient with errors less than 3%.

## RESULTS AND DISCUSSION

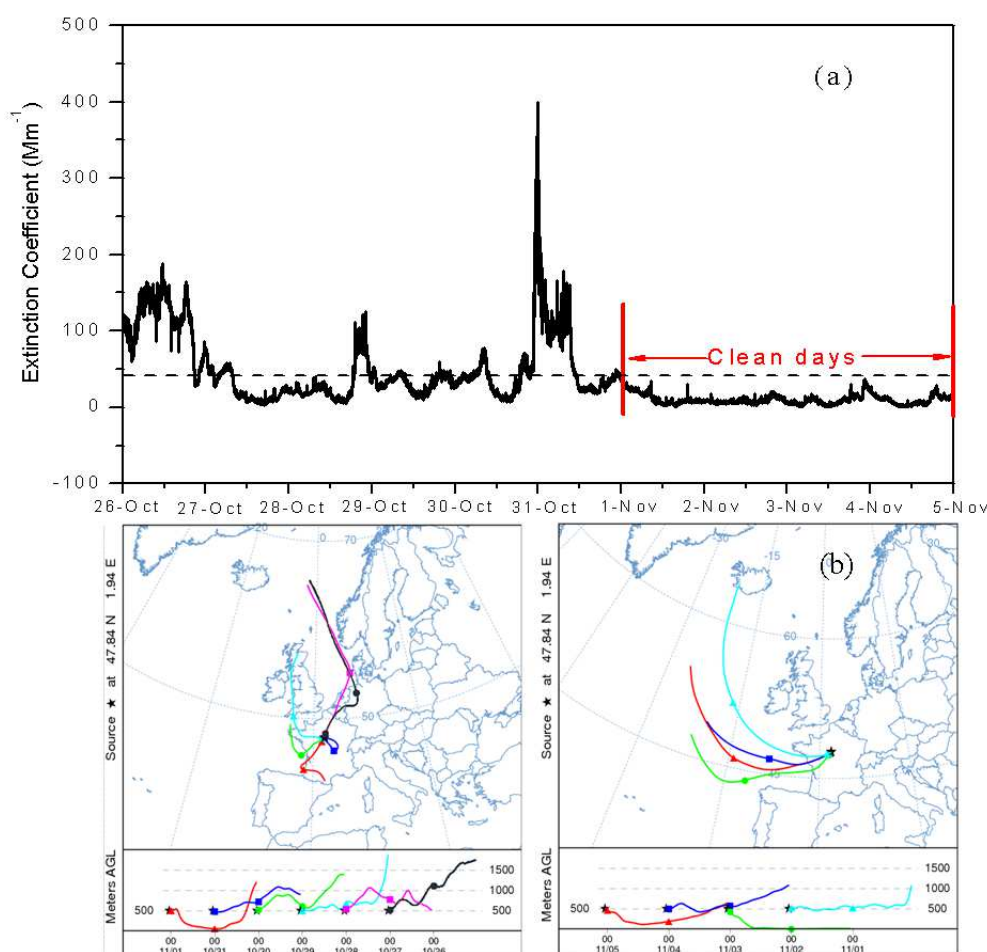
### Extinction Coefficient of the Non-dehydrated Ambient Aerosols

To study the real aerosol optical properties in Orleans, the extinction coefficient of the non-dehydrated ambient aerosols was measured from 26<sup>th</sup> October to 4<sup>th</sup> November 2012. As shown in Fig. 4(a), the average (dashed line) and maximum extinction coefficient values during this period were  $41 \pm 35$  and  $400 \text{ Mm}^{-1}$ . This result is comparable with those observed (Table 1), for example, in Eastern Mediterranean from July 2001 to January 2007 (the average value of  $b_{\text{sca}}$  (543 nm) and  $b_{\text{abs}}$  (565 nm) are  $40.3 \pm 27.1$  and  $5.4 \pm 3.7 \text{ Mm}^{-1}$ , Kalivitis *et al.*, 2011), and Western Mediterranean from March 2006 to December 2010 (the average value of  $b_{\text{sca}}$  at 550 nm is  $80 \pm 50 \text{ Mm}^{-1}$ , Esteve *et al.*, 2012), but much lower than those observed in Tokyo (the average value of  $b_{\text{ext}}$  at 532 nm is  $144.7 \pm 85.2 \text{ Mm}^{-1}$ , Nakayama *et al.*, 2010a) and Shanghai (the average value of  $b_{\text{sca}}$  and  $b_{\text{abs}}$  at 532 nm is  $332 \pm 217$  and  $85 \pm 60 \text{ Mm}^{-1}$ , Li *et al.*, 2011). Furthermore, it is noteworthy that the air quality during 1<sup>st</sup> to 4<sup>th</sup> November period (noted as “clean days”) is much better than the other days. Indeed, further



**Fig. 3.** (a) Extinction coefficient ( $b_{\text{ext}}$ ) at 532 nm measured as a function of particle number concentration of PSL at different sizes; (b) The extinction efficiency ( $Q_{\text{ext}}$ ) as a function of size parameter ( $\chi$ ) of PSL at 532 nm, the curve is obtained from Mie theory using refractive index of  $1.598 + 0.00i$ .





**Fig. 4.** (a) Extinction coefficient ( $b_{\text{ext}}$ ) of the non-dehydrated ambient aerosols. (b) The air parcel backward trajectories arriving at 500 m above ground level at Orleans from 26<sup>th</sup> October to 4<sup>th</sup> November, 2012.

**Table 1.** Summary of the reported atmospheric aerosol optical property measurement studies.

Location	Time	Value ( $\text{Mm}^{-1}$ )	Techniques	Reference
Tokyo Japan	14-Aug. to 2-Sep. 2007	$b_{\text{ext}} (532 \text{ nm}) = 144.7 \pm 85.2$ $b_{\text{sca}} (532 \text{ nm}) = 130 \pm 81.4$ $b_{\text{abs}} (532 \text{ nm}) = 13.6 \pm 9.2$	CRDS Nephelometer (3563) PSAP	Nakayama <i>et al.</i> (2011a)
Florida USA	4-May to 3-Jul. 2001	$b_{\text{ext}} (510 \text{ nm}) = 80.4$ $b_{\text{ext}} (578 \text{ nm}) = 87.0$	CRDS	Thompson <i>et al.</i> (2002)
Shanghai China	27-Nov. to 6 Dec. 2009	$b_{\text{sca}} (532 \text{ nm}) = 332 \pm 217$ $b_{\text{abs}} (532 \text{ nm}) = 85 \pm 60$	CRDS Nephelometer (3563)	Li <i>et al.</i> (2011)
Finokalia Greece	Jul. 2001 to Jan. 2007	$b_{\text{sca}} (543 \text{ nm}) = 40.3 \pm 27.1$ $b_{\text{abs}} (565 \text{ nm}) = 5.4 \pm 3.7$	Nephelometer (903) PSAP Aethalometer (AE21)	Kalivitis <i>et al.</i> (2011)
Valencia Spain	Mar. 2006 to Dec. 2010	$b_{\text{sca}} (550 \text{ nm}) = 80 \pm 50$	Nephelometer (3563)	Esteve <i>et al.</i> (2012)

analysis reveals that the average  $b_{\text{ext}}$  for the “clean days” is only  $10 \pm 7 \text{ Mm}^{-1}$ , about 6 times lower than that observed from 26<sup>th</sup> to 31<sup>st</sup> October 2012.

Fig. 4(b) displays the air parcel backward trajectory curves (calculated by HYSPLIT trajectory model, NOAA, <http://ready.arl.noaa.gov/HYSPLIT.php>) during the sampling time over Orleans. It can be noticed that the air parcel comes from the Atlantic Ocean for the “clean days”, while

it comes from interior regions for the other days. This result suggests that the air parcel transported from Atlantic Ocean may improve the air quality of Orleans through advection or wet deposition.

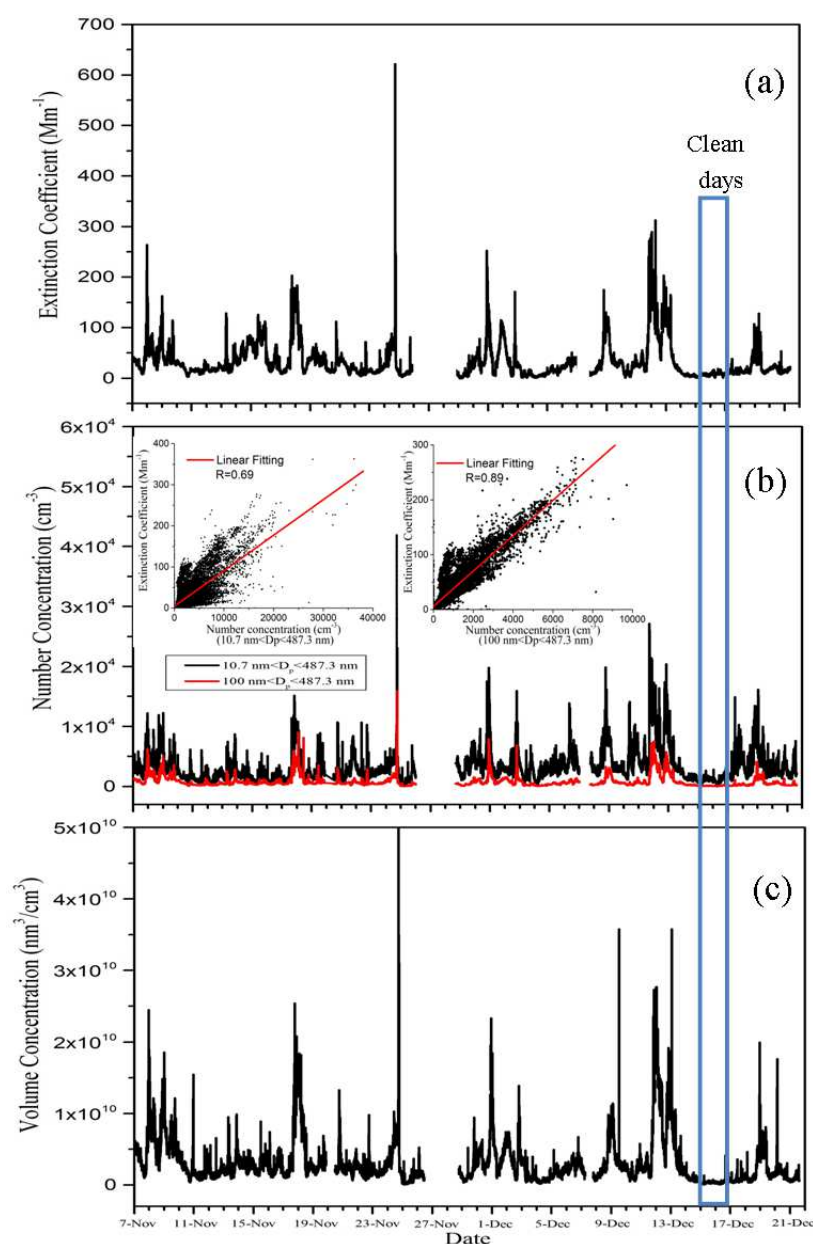
#### **Size Distribution and Extinction Coefficient of the Dehydrated Ambient Aerosols**

Fig. 5 presents the extinction coefficient, number (N)

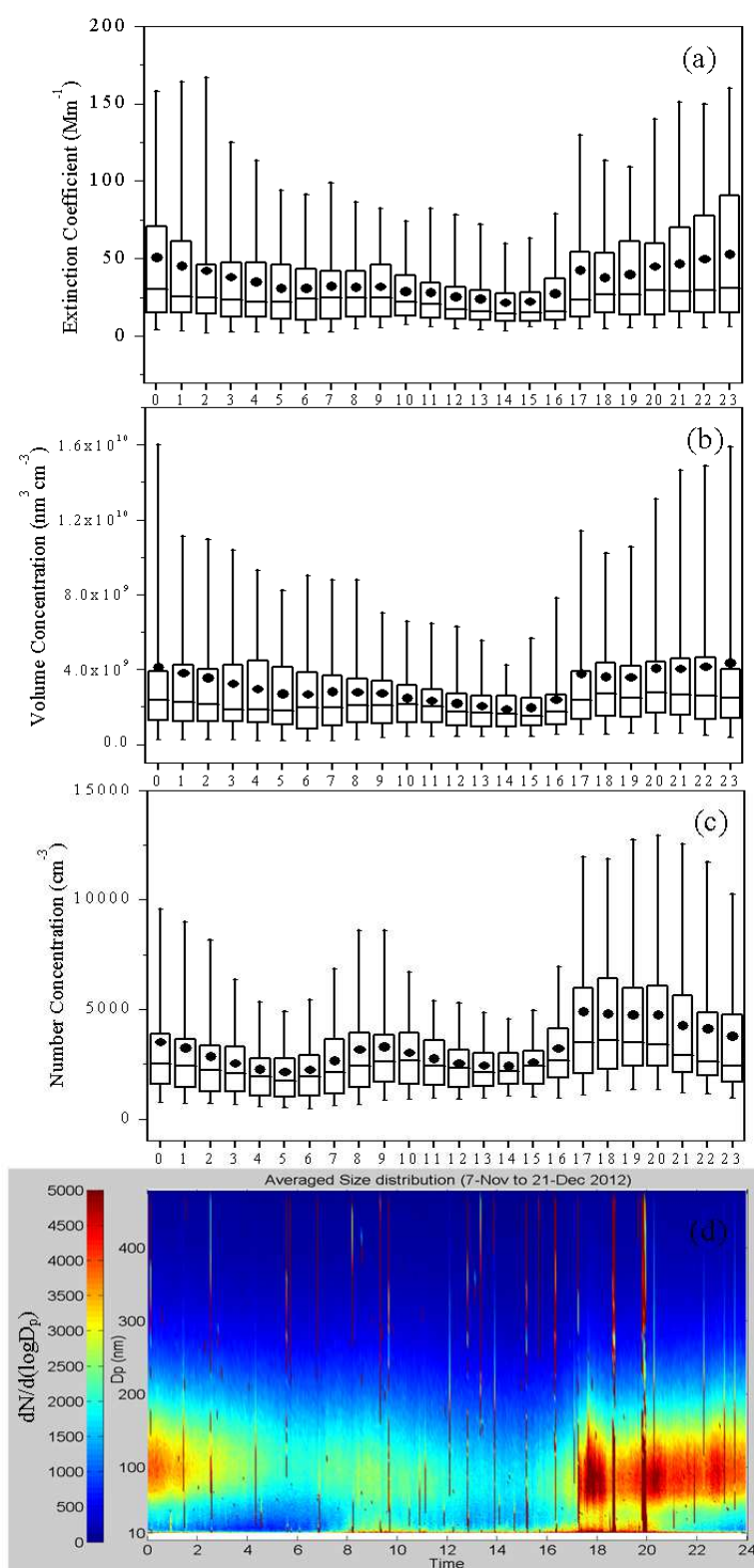
and volume (V) concentration of the dehydrated ambient aerosols from 7<sup>th</sup> November to 21<sup>st</sup> December, 2012 in Orleans. It can be seen that the extinction coefficient exhibit the same trend with the particle number and volume concentration during the whole observation period. The average values of  $b_{\text{ext}}$ , N and V were  $36 \pm 31 \text{ Mm}^{-1}$ ,  $3300 \pm 2700 \text{ cm}^{-3}$  and  $3.1 \pm 2.8 \times 10^9 \text{ nm}^3/\text{cm}^3$ , respectively. Moreover, the correlation plots of  $b_{\text{ext}}$  with particle number concentration ( $10.7 \text{ nm} < D_p < 487.3 \text{ nm}$  and  $100 \text{ nm} < D_p < 487.3 \text{ nm}$ ) were presented as an inset in Fig. 5(b). It reveals that the correlation of  $b_{\text{ext}}$  with the number concentration for particles larger than 100 nm ( $100 \text{ nm} < D_p < 487.3 \text{ nm}$ ,  $R = 0.89$ ) is better than that with the total particle number concentration ( $10.7 \text{ nm} < D_p < 487.3 \text{ nm}$ ,  $R = 0.69$ ). During these measurements, 15<sup>th</sup> and 16<sup>th</sup> December were selected

as the “clean days”, the average  $b_{\text{ext}}$ , N and V for these two days are only  $7 \pm 3 \text{ Mm}^{-1}$ ,  $1100 \pm 500 \text{ cm}^{-3}$  and  $3.8 \pm 2.5 \times 10^8 \text{ nm}^3/\text{cm}^3$ , respectively, about 6, 3 and 9 times lower than that of the other days. However, it can be clearly observed that on 24<sup>th</sup> November the values measured for the extinction coefficient ( $622 \text{ Mm}^{-1}$ ), number concentration ( $38000 \text{ cm}^{-3}$ ) and volume concentration ( $5.2 \times 10^{10} \text{ nm}^3/\text{cm}^3$ ) were higher compared to the other days. This typical event will be discussed in details later.

Figs. 6(a), 6(b) and 6(c) show the statistic diurnal variation of the  $b_{\text{ext}}$ , V and N. In this “box” diagrams, the mean is represented by dots. The dividing segment in the box is the median. The top box limit represents the 75<sup>th</sup> percentile and the bottom box limit represents the 25<sup>th</sup> percentile. The error bars are related to the percentiles 95<sup>th</sup> and 5<sup>th</sup>. In these



**Fig. 5.** (a) Extinction coefficient, (b) number and (c) volume concentration of the dehydrated ambient aerosols in Orleans between 7<sup>th</sup> November to 21<sup>st</sup> December, 2012.



**Fig. 6.** (a) Statistic diurnal variation of the particle extinction coefficient. (b) Statistic diurnal variation of the particle volume concentration. (c) Statistic diurnal variation of the particle number concentration. (d) One-day averaged particle size distribution derived from 7<sup>th</sup> November to 21<sup>st</sup> December, 2012.

figures,  $b_{ext}$  and  $V$  show a similar trend. Furthermore,  $b_{ext}$ ,  $V$  and  $N$  exhibit a clear daily pattern, with two peaks at 9:00 and 17:00 hrs. This daily variability is typical of urban areas,

and is mostly due to the car emissions during the rush hours (Esteve *et al.*, 2012). In addition, it is important to note that  $b_{ext}$  and  $V$  express a clear increase trend even after

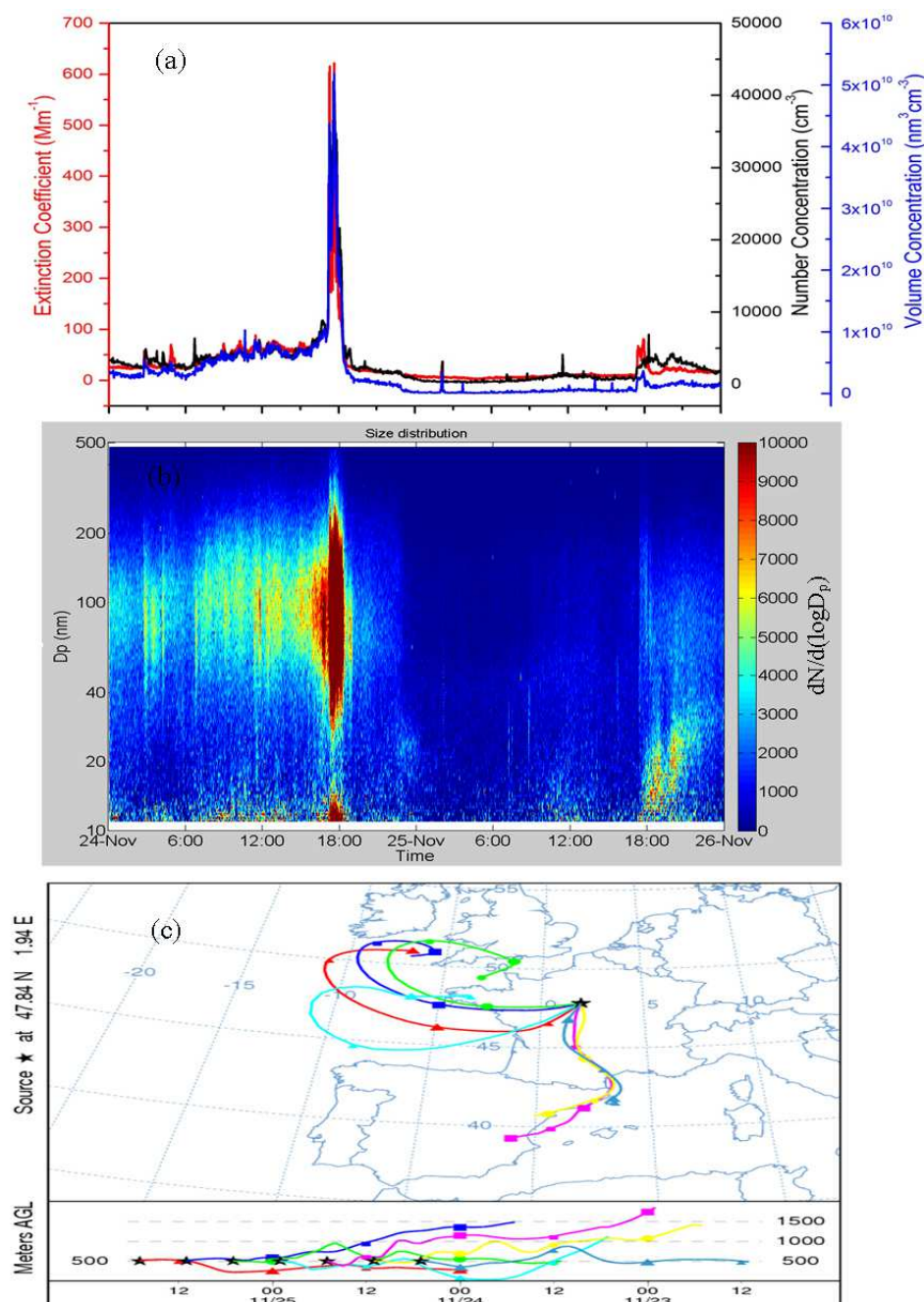


the second rush hour (19:00 to 24:00), while  $N$  decreases during this period. A reasonable explanation for this result is that the particle size is larger during the night time.

To further explain the diurnal variation of  $b_{\text{ext}}$ ,  $V$  and  $N$ , one-day averaged particle size distribution derived from 7<sup>th</sup> November to 21<sup>st</sup> December, 2012 is presented in Fig. 6(d). As expected, two high particle concentration events with particle size less than 40 nm were observed at 9:00 and 18:00, which confirms that, during the day time, car emissions are the main source of ambient aerosols near the measurement site. During the night, the concentrations of

the large particles ( $\approx 100$  nm) increased abruptly at 18:00 and this high concentration persisted until midnight. That is why  $b_{\text{ext}}$  and  $V$  express a clear increase trend even after the second rush hour (19:00 to 24:00). It is believed that the most possible source of these large particles could be the residential heating in the area.

During the measurements, a typical episode (Fig. 5), i.e., a sharp clean process, was observed on 24<sup>th</sup> November (Saturday). As it can be seen in Figs. 7(a) and 7(b),  $b_{\text{ext}}$ ,  $V$  and  $N$  only express one peak during this weekend, and the maximum values reached up to  $622 \text{ Mm}^{-1}$ ,  $5.2 \times 10^{10} \text{ nm}^3/\text{cm}^3$



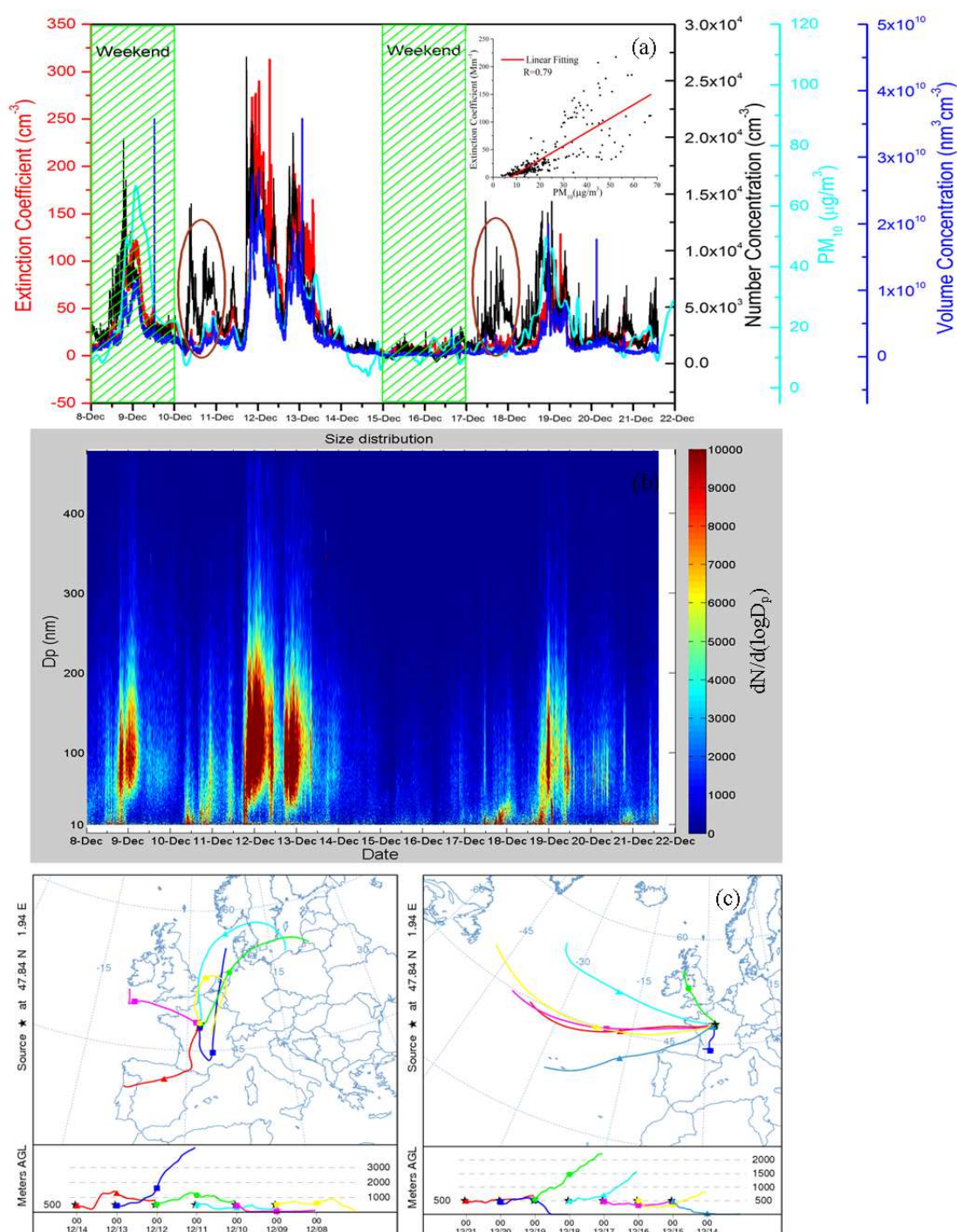
**Fig. 7.** (a) Extinction coefficient, volume and number concentration of the dehydrated ambient aerosols. (b) Size distribution of the dehydrated ambient aerosols. (c) The air parcel backward trajectories arriving at 500 m above ground level at Orleans from 24<sup>th</sup> to 25<sup>th</sup> November, 2012.

and  $38000 \text{ cm}^{-3}$ , respectively, at 17:40. Interestingly, unlike the results expressed in Fig. 6(d), this high pollution event did not persist during the whole night,  $b_{\text{ext}}$ ,  $V$  and  $N$  decreasing dramatically to  $39 \text{ Mm}^{-1}$ ,  $4.7 \times 10^9 \text{ nm}^3/\text{cm}^3$  and  $4700 \text{ cm}^{-3}$  only 50 minutes later. Further analysis of this clean process was conducted by looking at the air parcel backward trajectory curves during this weekend (Fig. 7(c)). This figure shows that the air parcel over Orleans shifts clearly from interior regions to Atlantic oceans after 17:00

(UTC time, corresponding to 18:00 in Orleans). This result further indicates that the air parcel transported from Atlantic Ocean could play a role in “cleaning up” the air over the measurements site area.

#### Typical Analysis the Cleaning-effect of Atlantic Air Parcel

Fig. 8(a) illustrates the aerosol extinction coefficient, particle number and volume concentration and  $\text{PM}_{10}$  concentration measured from 8<sup>th</sup> to 21<sup>st</sup> December 2012. It



**Fig. 8.** (a) Extinction coefficient, number concentration, volume concentration and  $\text{PM}_{10}$  concentration of the dehydrated ambient aerosols. (b) Size distribution of the dehydrated ambient aerosols. (c) The air parcel backward trajectories arriving at 500 m above ground level at Orleans from 8<sup>th</sup> to 21<sup>st</sup> December, 2012.

shows that the aerosol extinction coefficient exhibits the same trend with PM<sub>10</sub> ( $R = 0.79$ , the inset plot in Fig. 8(a)), particle volume and number concentration. However, in some special conditions (brown cycle marked in Fig. 8(a)), no clear change was observed for  $b_{\text{ext}}$  and even particle number concentration exhibits a small peak. Deep analysis reveals that all those peaks occur in the rush hour, particles from car emissions are too small (less than 40 nm, Fig. 8(b)) to increasing the  $b_{\text{ext}}$ .

As shown in Figs. 8(a) and 8(b), there are some clean days during this period, such as 14–17, 20 December, 2012. Fig. 8(c) gives the air parcel backward trajectories curves during this period over Orleans. As expected, the air parcel comes from the Atlantic oceans for all those clean days, while from interior regions for the other days. This confirms the previous hypothesis of the cleaning effect of air parcel transported from Atlantic Ocean.

## CONCLUSIONS

In this study, a pulsed CRDS instrument for measuring extinction coefficient of atmospheric aerosols was developed. The instrument was validated using monodisperse PSL particles. Meantime, size distribution and optical properties of ambient aerosols were measured from 26<sup>th</sup> October to 21<sup>st</sup> December, 2012 in Orleans, France, by using a SMPS coupled to the developed CRDS system. For the non-dehydrated aerosols, the average extinction coefficient at 532 nm is  $41 \pm 35 \text{ Mm}^{-1}$ . For the dried aerosols, extinction coefficient expresses a good agreement with the particle number and volume concentration, and the average values of  $b_{\text{ext}}$ ,  $N$  and  $V$  are  $36 \pm 31 \text{ Mm}^{-1}$ ,  $3300 \pm 2700 \text{ cm}^{-3}$  and  $3.1 \pm 2.8 \times 10^9 \text{ nm}^3/\text{cm}^3$ , respectively. One-day averaged particle size distribution derived from 7<sup>th</sup> November to 21<sup>st</sup> December, 2012 reveals that vehicle and boiler emissions maybe the main ambient aerosol sources near this specific measurement site. In addition, we found that the air parcel was coming from the Atlantic oceans for all clean days, while it was from interior regions for the other days, implying the air parcel transported from Atlantic Ocean could clean the air of Orleans city.

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