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**Northward growth of the Qimen Tagh Range: a new model accounting for the
Late Neogene strike-slip deformation of the SW Qaidam Basin**

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Abstract

Situated along the western termination of the Eastern Kunlun Mountains, the Qimen Tagh Range represents a key area to understand the Cenozoic basin-range interactions between the northeastern Tibetan Plateau and the Qaidam Basin. Within that region, several huge bow-like fault systems such as the Kunbei and Qimen Tagh fault systems accommodate the transpressive deformation but their kinematic evolution is still highly debated. Newly acquired seismic profiles and isopach maps of the Late Eocene sediments strongly suggest that the Kunbei fault system (consisting of the Kunbei, Arlar and Hongliuquan faults) in the southwestern Qaidam Basin was initially a left-lateral strike-slip fault system rather than a thrusting system. Growth strata indicate an Early Miocene onset age for this strike-slip deformation. However, earthquakes focal mechanisms show that the present-day tectonic pattern of this fault system is dominated by NE-SW transpression. As for the Qimen Tagh fault system, numerous linear geomorphic features and fault scarps indicate that it was again a strike-slip fault system. Deformed sediments within the Adatan Valley prove that strike-slip motion prevailed during the Pleistocene, yet the present day deformation is marked by NE-SW transpression. Collectively, the Kunbei and Qimen Tagh fault systems were initially left-lateral strike-slip fault systems that formed during Early Miocene and Pleistocene respectively. Colligating with these southward younging left-lateral strike-slip faulting ages and the fact that these convex-northward structures converge to the center segment of active Kunlun Fault in the east, we thus considered the Kunbei and Qimen Tagh fault systems as former western segments of the Kunlun Fault once located further south in the present-day location of that fault. These faults gradually migrated northward since the Early Miocene while their kinematics changed from left-lateral strike-slip motion to NE-SW transpression.

Keywords: Cenozoic tectonics, North Tibet, Kunlun Fault, Qimen Tagh Range, southwestern

Qaidam Basin.

1. Introduction

The Eastern Kunlun Mountains and adjacent Qaidam Basin present one of the strongest topographic gradients in Eurasia, similar to that of the Longmen Shan Range in East Tibet (Fig. 1). Within that region, the left-lateral strike-slip Kunlun Fault Zone represents one of the key tectonic features controlling the active deformation along the northern margin of the Tibetan Plateau (e.g. Avouac and Tapponnier, 1993; Meyer et al., 1998; Jolivet et al., 2001; Tapponnier et al., 2001; Xia et al., 2001; Yin et al., 2007). Most of the previous studies that dealt with the Tertiary deformation in the East Kunlun ranges focused on the central and eastern segment of the Kunlun Fault (Kirby et al., 2007; Harkins and Kirby, 2008) where the deformation is mostly localized along the major, unique strike-slip fault zone. However, only a limited number of studies describe the more structurally complex western termination of the Kunlun Fault system (Jolivet et al., 2001, 2003).

Although the Cenozoic tectonic evolution of North Tibet has been largely studied, the relationship between the Qaidam Basin and the East Kunlun Range, as well as the tectonic pattern of the Kunlun Fault itself, remains controversial. For example, several contradicting models have been proposed to describe the tectonic structure of the southern edge of the Qaidam Basin: a crustal-scale south-dipping thrust system (Burchfiel et al., 1989); a large transpressional system including the left-lateral strike - slip Kunlun Fault to the south and a north - directed thrust system along the northern flank of the Eastern Kunlun Mountains (Meyer et al., 1998; Tapponnier et al., 2001; Jolivet et al., 2003); or a series of north - dipping thrusts carrying the Qaidam Basin southward onto the Qimen Tagh Range (Yin et al., 2007; Shi et al., 2009; Wang et al., 2011). Within most of those models, the large amount of Cenozoic strike-slip motion along the Kunlun

Fault (Fu and Awata, 2007) is generally considered as secondary in controlling the overall structural setting of the SW Qaidam Basin and East Kunlun Range. However, if the boundary fault between the Qaidam Basin and the East Kunlun Range was indeed accommodating a large amount of lateral displacement, the southern Qaidam Basin should be largely decoupled from the Tibetan Plateau to the south and should not be affected by major N-S compression and shortening since the initiation of the strike-slip Kunlun Fault. Nonetheless, the Qimen Tagh Range, bordering the SW Qaidam Basin to the south forms a convex-northward bow-like tectonic prism clearly accommodating shortening along the southern margin of the basin while strike-slip motion occurs along the Kunlun Fault to the south.

This paper aims to re-evaluate the tectonic pattern along the western portion of the Kunlun Fault and to decipher the Late Cenozoic interaction between the East Kunlun Range and the SW Qaidam Basin in the Qimen Tagh Range. This work is based on new field observation and data, interpretation of newly acquired seismic reflection profiles within the SW Qaidam Basin, remote sensing images, and earthquake focal mechanisms and epicenters distribution. Based on those data we propose that the Late Neogene northward growth of the convex-northward structures in the Qimen Tagh Range and SW Qaidam Basin is driven by left-lateral strike-slip motion along the Kunlun Fault. While migrating northward, the western segment of the Kunlun Fault gradually rotates through time. This northward rotation induces a progressive change from left-lateral strike-slip to NE-SW compressional motion. Continuous activity along the Kunlun Fault finally initiates a new strike-slip segment.

2. Regional Geology

2.1 Eastern Kunlun Mountains

The Eastern Kunlun Mountains, with an average altitude of ~4500m, separates the relatively low-relief Qaidam Basin to the north from the Hoh Xil Basin to the south. The Kunlun Fault Zone itself represents a major tectonic boundary between the Bayan Har-Songpan Garze terrane and the Kunlun-Qaidam basement (e.g. Meyer et al., 1996). The fault extends about 1600 km along-strike, initiating on the eastern margin of the Tibetan Plateau (Kirby et al., 2007; Lin and Guo, 2008) and ending in a splay structure with several curved faults in the west (Xu et al., 2002; Fu and Awata, 2007) (Fig. 1). Intense seismicity attests of the activity of this left-lateral strike-slip fault system (e.g., M_w 7.8 Kunlunshan earthquake in 2001 (Lin et al., 2002; Xu et al., 2002); M_w 7.6 Manyi earthquake in 1997 (Peltzer et al., 1999); M_w 7.5 Manyi earthquake in 1973 (Velasco et al., 2000); M_w 7.5 Tuosu Lake earthquake in 1937 (Guo et al., 2007); M_s 7.0 Alake Lake earthquake in 1963 (Guo et al., 2007)). Both GPS and geological data suggest that the central segment of the Kunlun Fault has a uniform slip rate of ca. 10 mm/yr (Kidd and Molnar, 1988; Van Der Woerd et al., 1998; Wang et al., 2001; Zhang et al., 2004; Fu et al., 2005; Kirby et al., 2007). This rate decreases to ~5 mm/yr along its eastern segment and finally dies out toward the eastern fault tip (Kirby et al., 2007; Harkins and Kirby, 2008). The total displacement along the fault since its initiation is approximately 100 ± 20 km according to the largest cumulative offset of basement rocks along the tectonic lineament (Fu and Awata, 2007).

The Cenozoic deformation and exhumation history of the Eastern Kunlun Mountains is complex and still largely debated. For instance, the onset age of the Cenozoic cooling event in this

area range from 35 to 10 Ma depending on the authors (Mock et al., 1999; Jolivet et al., 2001; Wang et al., 2004a; Liu et al., 2005; Yuan et al., 2006; Clark et al., 2010; Duvall et al., 2013). However, the onset age of left-lateral strike-slip motion along the Kunlun Fault has been estimated to be 15 Ma or earlier (Jolivet et al., 2003). Similarly, based on the growth strata records in the seismic profiles across the SW Qaidam Basin, Yin et al. (2008) asserted that the Cenozoic deformation across the Qimen Tagh Range and in the SW Qaidam Basin, directly related to the tectonic development of the Kunlun Fault Zone, initiated during or after Late Oligocene to Early Miocene times.

2.2 *SW Qaidam Basin and Kumukol Basin*

The ~3500 m high (on average), low relief triangular Qaidam Basin is bounded by the Altyn Tagh Range to the northwest, the Eastern Kunlun Mountains to the south, and the Qilian Shan to the northeast (Fig. 1). Recent geological cross sections within the Qaidam Basin, mostly based on seismic profiles and drill cores data, demonstrate that the upper Cretaceous and Tertiary sediments rest unconformably on the Jurassic to Cretaceous strata. This unconformity indicates that some deformation already took place within the basin during the Cretaceous (Xia et al., 2001; Fang et al., 2007). The Cenozoic sequences mainly consist of a series of continental clastic deposits (Fig. 2). The successive depocenters were consistently located along the long axis of the basin and gradually migrated eastward since the Eocene (Balley, 1986; Song and Wang, 1993; Qiu, 2002; Xia et al., 2001; Sun et al., 2005; Yin et al., 2008).

The chronology of the Cenozoic stratigraphic units has been precisely defined (Fig. 2) mostly using magnetostratigraphy, but also based on palynology and paleontology studies within the entire basin (QBGMR, 1991; Yang et al., 1992; Deng et al., 2004a, 2004b; Sun et al., 2005; Song

et al., 2006; Zhang et al., 2006; Fang et al., 2007; Sun et al., 2007; Wang et al., 2007; Lu and Xiong, 2009; Ke et al., 2013). The series include: (1) the Lulehe Formation, $E_{1+2}l$, >53.5 - 43.8 Ma (Yang et al., 1992; Song et al., 2006; Zhang, 2006; Ke et al., 2013); (2) the lower Xiaganchaigou Formation, E_3^1xg , 43.8 - 37.8 Ma (Song et al., 2006; Zhang, 2006; Sun et al., 2007; Pei et al., 2009); (3) the upper Xiaganchaigou Formation, E_3^2xg , 37.8 - 35.5 Ma (Sun et al., 2005; Sun et al., 2007; Pei et al., 2009); (4) the Shangganchaigou Formation, N_{1sg} , 35.5 - 22.0 Ma (Sun et al., 2005; Lu and Xiong, 2009); (5) the lower Youshashan Formation, N_2^1xy , 22.0 - 15.3 Ma (Fang et al., 2007; Lu and Xiong, 2009); (6) the upper Youshashan Formation, N_2^2sy , 15.3 - 8.1 Ma (Fang et al., 2007); (7) the Shizigou Formation, N_2^3s , 8.1 - 2.5 Ma (Fang et al., 2007); and finally, (8) the Quaternary sediments, including the Qigequan Formation (Q_{1q}) and the Dabuxun - Yanqiao Formation, 2.5 - 0 Ma (Fang et al., 2007; Yin et al., 2008). The boundaries between these formations have been labeled T_R , T_5 , T_4 , T_3 , T_2' , T_2 , T_1 and T_0 (from bottom to top) by petroleum geologists and will be used as it below (Fig. 2).

The Kumukol Basin (Figs. 1 and 2), with an average elevation of ~4000m, is located between the Qimen Tagh Range to the north and the western branches of the Eastern Kunlun Range to the south. Remote sensing image and field evidence indicate that the northern part of the Kumukol Basin is bordered by an active fault which controls the distribution of the present drainage system (IGSSP, 2003). The basin contains an Eocene to Pliocene sequence which has been deformed into a NW-SE oriented syncline (Meng and Fang, 2008). Based on fossils and stratigraphic correlation (Zhang et al., 1996), these Tertiary sequences can be divided into four units (from bottom to top) (Meng and Fang, 2008): the Huatiaoashan Formation (Eocene to Oligocene), the Hongshiliang Formation (Early Miocene), the Fengchenkou Formation (Middle Miocene), and the Jiantuliang

Formation (Pliocene to Pleistocene). The Eocene to Oligocene series in the Kumukol Basin are dominated by lacustrine deposits (IGSSP, 2003) displaying similar facies and depositional environments as the Eocene to Oligocene deposits in the SW Qaidam Basin and Hoh Xil Basin. This similarity of the deposits within the three basins suggests that they were largely connected during the Paleogene (Yin et al., 2008).

3. Active tectonic pattern in the western segment of the East Kunlun Range

The distribution and focal mechanisms of earthquakes reflect the occurrence and kinematics of active faults and allow describing the active tectonic pattern of the study area (e.g. Taylor and Yin, 2009; Jolivet et al., 2013). We analyzed the parameters of all earthquakes in an area between 89°E to 98°E and 35°N to 39°N, for which precisely determined hypocenter depths and focal mechanisms are available. The epicenters and depth - frequency distribution of seismicity used in this study are issued from the China Earthquakes Network Center (CENC). Only events with magnitude $M > 3$ that occurred between 1970 and 2011 have been considered (Fig. 3). Focal mechanisms were obtained from the U.S. Geological Survey and the Global Centroid Moment Tensor Project. The depth - frequency distribution of seismicity is presented along two profiles, A - A' (Fig. 4a) and B - B' (Fig. 4b) (See Fig. 3 for the location of the profiles). The epicenters of earthquakes situated between 89°E to 92°E and 35°N to 39°N are projected on profile A-A'; Those of the earthquakes situated between 90°E to 98°E and 35°N to 37°N are projected on profile B-B'.

Earthquakes magnitude ranges from $M = 3$ (the selected cut-off value) to $M = 7$. Strong tectonic activities (mean M between 5 and 6) are recorded along the Eastern Kunlun Mountains, on the southern edge of the Qaidam Basin. However, diffuse seismicity (mean M between 4 and 5) also occurs inside the basin (Fig. 3 and 4). Surprisingly, the Qimen Tagh Range displays a relative lack of seismicity. Within the Eastern Kunlun Mountains, the distribution and focal mechanisms of the earthquakes are coherent with left-lateral, transpressive strike-slip movement on the Kunlun Fault. The existing few transtensive focal mechanisms can be explained by the occurrence of

extensional relay-zones along the main Kunlun Fault (Jolivet et al., 2003). The focal mechanisms available inside the Qimen Tagh Range and the SW Qaidam Basin all correspond to reverse faults activity indicative of a general compressive stress field with no or extremely limited strike-slip motion.

The depth-frequency distribution of the earthquakes indicates that both in the Eastern Kunlun Mountains and in the Qaidam Basin, seismicity affects the upper 30 to 35 km of the crust (that is thought to be at least 50 km thick in that region (e.g. Jolivet et al., 2003; Zhang et al., 2011)) with most events occurring in the upper 10 to 15 km. This distribution pattern is similar to the depth-frequency distribution of seismicity over the entire Tibetan Plateau (Taylor and Yin, 2009), with exceptions where earthquakes might have occurred within the lower crust or upper lithospheric mantle (Chu et al., 2009). The observed seismicity is thus mostly indicative of the deformation of the upper to middle crust and can be reasonably compared to the fault pattern observed on the surface.

4. Splay Fault System of the Kunlun Fault Zone

Different from the typical horse-tail structure, the western segment of the Kunlun Fault develops several asymmetric northward curved structures topographically characterized by a series of ridges and valleys (Fig. 1). These elongated linear geomorphic landforms are similar to those of the central segment of the Kunlun Fault, such as the Xidatan and Dongdatan valleys (Xu et al., 2002), and indicate the occurrence of left-lateral strike-slip faulting. Meanwhile, the valleys are largely filled with Cenozoic deposits that recorded the history of the strike-slip faulting and especially its onset time. In this study, taking the Qimen Tagh Range as an example, we aim at clarifying the tectonic pattern of the western segment of the Kunlun Fault through time.

4.1 Strike-slip faulting in the Qimen Tagh Range

The Qimen Tagh Range, located in the western part of the Eastern Kunlun Mountains, is over ~4500 m high on average and marks the boundary between the Kumukol Basin to the south and the Qaidam Basin to the north (Fig. 1). The range has a general WNW-ESE elongated shape. To the NW, it bends sharply to a southwestward trend sub-parallel to the Altyn Tagh Range that lies directly to the north. To the SE, the Qimen Tagh Range converges with the Eastern Kunlun Mountains (Fig. 1 and 5). It is separated into two parallel NW-SE oriented sub-ranges by the Adatan Valley (Fig. 2).

To the north, the sharp contact between the Qaidam Basin and the Qimen Tagh Range has been interpreted as evidence for active north-directed thrusting (e.g., Meyer et al., 1998; Jolivet et al., 2003), whereas field examination suggests that this is a steeply north-dipping unconformity between bedrock in the south and north-dipping Cenozoic strata in the north (Yin et al., 2007).

Detailed analysis of both SRTM digital elevation models (DEM) and Landsat 5 satellite images across the Qimen Tagh Range reveals abundant tectonic lineaments (Fig. 5). The extensive Qimen Tagh fault system consists of two major, sub-parallel tectonic lineaments: the Middle Qimen Tagh Fault (MQF) and the South Qimen Tagh Fault (SQF) (Fig. 5).

The MQF affects the northern edge of the Adatan Valley. Based on field measurements and on previous studies (Wang et al., 2007), the MQF dips about 55° to the south. The strike of the MQF changes from ENE in the western part of the range to WNW in the eastern part. The western segment shows a reverse motion associated to a slight left-lateral strike-slip component. Some of the scarps and branched faults are arranged in echelons (Fig. 6). In the eastern part of the MQF the main active fault is obliterated by Quaternary deposits. However, some linear structures are visible on the eastern end of the MQF and in the convergence zone with the Kunlun Fault system (Figs. 5 and 7). This however poorly exposed tectonic pattern suggests a very probable connection between the Qimen Tagh Fault system in the west and the Kunlun Fault system in the east.

The SQF separates the southern sub-range of the Qimen Tagh Range from the Kumukol Basin to the south (Fig. 5). The fault zone is marked by a series of distinctive fault scarps and structural lineaments which are inferred to be reverse faults striking approximately $NE120^{\circ}$ (Fig. 5).

Furthermore, several evidences of recent activity such as river offsets and alluvial fan displacements indicate a left-lateral reverse movement on the fault.

Although the offset channels and geomorphic surfaces indicate a recent, probably still active general left-lateral shearing within the range, the focal mechanism solution mentioned above suggest that this recent left-lateral strike-slip faulting is relatively slight compared to the NE-SW compressive component.

4.2 *Pleistocene series within the Adatan valley*

Fieldwork in the Adatan Valley was conducted to determine the timing of the strike-slip faulting in the south Qimen Tagh sub-range. Detailed analysis of the Cenozoic syntectonic deposits was conducted to provide a better understanding of the deformation in the Qimen Tagh Range (Fig. 8).

The morphology of the Adatan Valley is marked by a strong topographic relief between the valley and the mountains (Fig. 8a). Within the two sub-ranges, the series are mainly composed of Precambrian basement, Paleozoic metasediments and Paleozoic and Mesozoic igneous rocks. The Cenozoic series exposed within the Adatan Valley was assigned a Pleistocene age based on previous geology mapping and regional stratigraphic correlation (IGSSP, 2003; IGSQP, 2004; Wang et al., 2004b; Wang et al., 2007; Fu et al., 2008).

The deposits are deformed into several anticlines cut by modern rivers within the valley (Figs. 8a and 9). Along the northern edge of the valley, the Cenozoic series overlap northward on the Ordovician basement and are dominated by a sequence of conglomerates and sandstones. Along the southern edge of the valley, the Cenozoic series are deformed and unconformably covered by Holocene tillites and alluvial sediments. As in the north, these deposits are composed of coarse-grained conglomerates and sandstones in alluvial fans facies. The petrological composition of the conglomerates corresponds to the whole range of basement rocks exposed in the Qimen Tagh Range, including igneous, metamorphic rocks and marine limestones. Poor sphericity and poor sorting of the grains indicate very proximal deposits. The grain size generally increases upward (Fig. 8b). Similarly to the Quaternary sedimentary within the Xidatan Valley further east (Xu et al., 2002) that record the present active feature of the Kunlun Fault, the Pleistocene

sequences exposed in the Adatan Valley record the topographic growth of both the north and south Qimen Tagh sub-ranges. They indicate a Pleistocene age for the onset of strike-slip faulting in the Qimen Tagh Range.

As a consequence, the linear geomorphic features, occurrence of faults scarps in the Qimen Tagh Range as well as the deformed Pleistocene deposits in the Adatan Valley indicate that the reverse strike-slip deformation in the region has been active since the Pleistocene.

5. Strike-slip faulting in the SW Qaidam Basin

In order to decipher the Cenozoic strike-slip faulting pattern in the SW Qaidam Basin, we choose the isopach map of the upper Xiaganchaigou Fm. (E_3^{2xg}) as an example to analyze the impact of fault motion on the sediment distribution within the basin (Fig. 10). Data were obtained from the Qinghai Oilfield Company (China) who recently compiled information from thousands of drill holes and a dense network of seismic profiles.

The Kunbei fault system was composed of three major WNW trending faults, including the Kunbei Fault, the Arlar Fault and the Hongliuquan Fault. As reported on the isopach map based on seismic data, these three faults controls the general tectonic pattern of the SW Qaidam Basin. The Kunbei Fault and Arlar Fault have been previously interpreted as reverse faults active since the Paleogene (Wang et al., 2010). However, the three major faults mentioned above are arranged into a left-lateral right-step en-échelon pattern (Fig. 10), that does not correspond to the usual imbricated pattern of purely compressive thrust systems such as, for example, in the Appalachian thrust belt (Macedo and Marshak, 1999; Thomas, 2001) or in the Zagros fold and thrust belt (McQuarrie, 2004; Sepehr and Cosgrove, 2004). Meanwhile, the isopach lines of the Late Eocene Upper Xiaganchaigou Fm. are obviously offset across the faults reflecting a left-lateral strike-slip movement along these three faults. According to the well-expressed offset of the 400 meters isopach line, the left-lateral displacement of the Honglingquan Fault, the Arlar Fault and the Kunbei Fault increases from a few kilometers to tens of kilometers (Fig. 10). Therefore, the variations in thickness of the Late Eocene series across these three faults, as well as the en-échelon pattern of the faults, clearly demonstrates the occurrence of left-lateral strike - slip faulting in the

SW Qaidam Basin after the Late Eocene.

To further explore the tectonic pattern of the SW Qaidam Basin and the timing of faulting, three seismic reflection profiles (C-C', D-D', E-E') recently acquired by the Qinghai Oilfield Company have been selected. These three NE-trending seismic lines extend from the northern front of the Qimen Tagh Range towards the center of the SW Qaidam Basin (See location on Figs. 2 and 10).

Seismic Section C-C'

The seismic section C-C' situated in the central part of the SW Qaidam Basin, intersects the Arlar Fault (Fig. 10 and 11). On that section the Arlar Fault is a south dipping high angle fault associated to a few branched faults. The Mesozoic series are not preserved on both sides of the Arlar Fault where only Cenozoic deposits can be identified. Although the strata on each side of the fault consist in the same Cenozoic series, the thickness of each formation is quite different showing obvious stratigraphic throw. This major stratigraphic throw corresponds to the offset observed in the isopach lines of the Late Eocene Upper Xiaganchaigou Fm. on both side of the Arlar Fault (Fig. 10). Combined together, these two sets of data demonstrate a left-lateral reverse motion along the Arlar Fault during the Tertiary.

The thickness of the strata younger than Neogene Xiayoushashan Fm. (N_2^{1xy}) decreases towards the Arlar Fault and growth strata develop on both sides of the fault. Based on these features, previous studies interpreted the Arlar Fault as a purely reverse fault that accommodated rapid uplift in the Eastern Kunlun Mountains since the Late Oligocene - Early Miocene (Song and Wang, 1993; Yin et al., 2007; Wang et al., 2010). Since it is very difficult to identify strike-slip faults on seismic profiles (Durand-Riard et al., 2012), these studies did not take into consideration

the strike-slip component along the major faults within the SW Qaidam Basin. Growth strata are generally used to reveal the initiation time of a reverse fault (e.g. Suppe et al., 1992; Vergés et al., 2002), but strike-slip motion may also induce vertical offset of strata, such as in positive flower structure (e.g. Christophoul et al., 2002). Such motion can also generate growth strata which will then provide an accurate initiation time for the strike-slip faulting (e.g. Hinsch et al., 2005). Consequently, the growth strata associated to the positive flower structure in seismic profile C-C' presented above, as well as the northward thinning of the Early Miocene and younger series in the footwall of the Arlar Fault date the onset time of left-lateral strike-slip faulting along that fault (Fig. 10 and 11).

Seismic Section D-D'

The tectonic pattern of section D-D' is characterized by several high angle faults defining a block-like structure (Fig. 12). The major faults are the Arlar Fault and the Hongliuquan Fault. The sediments within the section are Mesozoic to Cenozoic in age. The cover sequence in the hanging wall of the Arlar Fault is composed of Cenozoic sediments whereas the footwall is dominated by Late Mesozoic, Paleogene and Early Neogene deposits. The Mesozoic sediments are relatively thick within the footwall where the Tertiary series are also obviously thicker than that of the hanging wall. The basement and Cenozoic strata on both sides of the Arlar Fault display obvious stratigraphic throw. Further to the north, the Early to Middle Cenozoic and Late Mesozoic strata are cut by the north dipping high angle Hongliuquan Fault. In the hangingwall of the Hongliuquan Fault, the pre-Early Miocene strata were deformed in an anticline later cut and sealed by the Early Miocene sediments clearly indicating Early Miocene motion of this fault. Like in profile C-C', the thickness of the Cenozoic units is constant within each tectonic compartment but varies sharply

between two adjacent blocks. Combined with the offset isopach lines of the Late Eocene deposits (Fig. 10), this pattern can also be explained by lateral movements along the faults during the Tertiary. Again like in seismic profile C-C', the occurrence of growth strata initiating in the Early Miocene series north of the Arlar and Hongliuquan faults implies that left-lateral strike-slip deformation initiated during the Early Miocene in the SW Qaidam Basin.

Seismic Section E-E'

Seismic section E-E' is located about 20 km east of seismic section C-C' and again runs from the northern front of the Qimen Tagh Range to the western part of the central Qaidam Basin (Figs. 10 and 13). As in section D-D', the tectonic pattern is characterized by several high angle faults but the cover-sequence only consists in Cenozoic strata. To the south, the activity along the Kunbei Fault is marked in the hanging wall by the erosion of the Late Eocene and younger strata and probably by the non-deposition of some of the upper Tertiary and Quaternary series. This erosion also affects the footwall of the Kunbei Fault, though the Early Tertiary series are better preserved with some Late Eocene sediments still present. Like in section C-C' and D-D', the thickness of the Early Tertiary series vary across the fault showing obvious stratigraphic throw. To the north, the XIII fault dips to the north and cuts the Early to Middle Cenozoic sequence, while the large stratigraphic throw on both sides of the fault decreases from basement to Quaternary series. Like in the previous sections, the thickness of the Paleocene to Eocene strata varies sharply from both sides of the Kunbei and XIII faults, again attesting of lateral movements on those structures during the Tertiary. Furthermore, the thickness of the strata younger than the Xiayoushashan Fm. (N_2^1 xy) within the footwall of the Kunbei and XIII faults decreases towards the faults. Growth strata are again initiating in the footwall of both faults during Early Miocene

times dating again the onset of deformation in that region.

From what has been discussed above, we conclude that the tectonic attribute of the Kunbei Fault during the Early Miocene is featured by a left-lateral transpressive system mainly composed of the Kunbei Fault, the Arlar Fault and the Hongliuquan Fault. Among these faults, the Kunbei and Arlar faults are the two major boundary faults separating the Cenozoic series within the SW Qaidam Basin from the Kunlun basement to the south.

6. Discussion

6.1 Available models for the Cenozoic evolution of the SW Qaidam Basin.

In the past decades, several models have been proposed to explain the Cenozoic tectonic and sedimentary evolution of the Qaidam Basin. Most of those concentrated on the tectonic pattern displayed in the Qimen Tagh Range and SW Qaidam Basin, and simply defined the tectonic pattern between the Eastern Kunlun Mountains and the Qaidam Basin through a northward thrusting or southward thrusting model.

The northward thrusting model prevailed in the last two decades, explaining the subsidence of the Qaidam Basin as a result of the basinward propagation of successive thrust sheets along the southern edge of the basin. Burchfiel et al. (1989) proposed that the southern edge of the Qaidam Basin is limited by a south-dipping thrust fault. However, field observation revealed that the contact between the Eastern Kunlun Mountains and the Qaidam Basin is defined by gentle north-dipping Pliocene strata overlapping north-dipping Carboniferous beds (Yin et al., 2007). Based on a synthesis of focal mechanisms and focal-depth distribution combined with geological features, Chen et al. (1999) again speculated the existence of this south-dipping North Kunlun thrust fault along the northern margin of the East Kunlun belt. However, their data projected on the single cross section were collected along the over 1000 km long Eastern Kunlun Mountains and Yin et al. (2007) showed that it was not possible to differentiate between the two possible fault planes solutions (N dipping or S dipping). In addition, many authors indicated that northward basement thrusting across the Eastern Kunlun Mountains was associated to the northward propagation of a crustal-scale accretionary wedge initiating around ca. 30-20 Ma. (Meyer et al.,

1998; Mock et al., 1999; Jolivet et al., 2001; Tapponnier et al., 2001; Jolivet et al., 2003; Wang et al., 2006). These ideas suggest that the Eastern Kunlun Mountains consists in a large transpressional system limited to the south by the left-lateral strike-slip Kunlun Fault and to the north by a series of south-verging basement thrusts affecting the southern edge of the Qaidam Basin. Yet, the observed thickness pattern of the Cenozoic strata across the Qaidam Basin (Yin et al., 2008) seems poorly compatible with a major south-dipping range-bounding thrust along the northern margin of the Eastern Kunlun Mountains. All the individual Cenozoic units in the basin consistently thicken from the margins towards the center of the basin (Bally, 1986; Huang et al., 1996; Sobel et al., 2003; Wang et al., 2006; Yin et al., 2007, 2008). This isopach pattern contradicts the classic tectonically subsiding foreland-basin model that requires the maximum sediment thickness to be localized near the bounding thrust system (Jordan, 1981).

The southward thrusting model has drawn much attention in recent years (Yin et al., 2007; Shi et al., 2009; Wang et al., 2011). Yin and Harrison (2000) suggested that two major thrusts, the Qimen Tagh Thrust and the North Kunlun Thrust mark the southern boundary of the Qaidam Basin. Using seismic profiles within the Qaidam Basin and field observations, Yin et al. (2007) then proposed that the low altitude Qaidam Basin has been thrust onto the Eastern Kunlun Mountains along major north dipping Cenozoic thrusts. This model was later expanded by Wang et al., (2011) based on a high-resolution deep seismic reflection profile across this area. This model indicates that the Eastern Kunlun Mountains should not thrust northward onto the Qaidam Basin and that the initiation of compressive deformation along the southern margin of the Qaidam Basin is much younger than along the northern margin of basin. However, this model does not take any strike-slip component in the Qimen Tagh Range and SW Qaidam Basin into account. In

addition, the convex-northward bow-like tectonic structures in the Qimen Tagh Range are not consistent with southward thrusting of the basin onto the range. Such a deformation pattern would rather generate a southward inflexion of the structures.

Other tectonic models have been proposed aside from the two main ones described above. Xia et al. (2001) proposed that the Qaidam Basin experienced a two stages evolution: an initial extension phase during the Early Tertiary was followed by a compressive phase leading to basin inversion. However, few seismic profiles in the western Qaidam Basin have revealed any extensional faults during the Early Tertiary. Meng and Fang (2008) suggested that the Cenozoic tectonic subsidence of the Qaidam Basin resulted from crustal folding or buckling in response to regional horizontal compression. However, the seismic profiles across the whole Qaidam Basin imply that the magnitude of Cenozoic upper crustal shortening decreases eastward across the basin from >48% in the west to <1% in the east (Yin et al., 2008), whereas the present day basin is wider to the west and relatively narrow to the east. This phenomenon is difficult to explain only by crustal folding or buckling.

To summarize, despite the large number of models that have been proposed, few studies took the left-lateral strike-slip factor into account when considering the relationship between the Qaidam Basin and the Eastern Kunlun Mountains.

6.2 Strike-slip tectonics and northward growth of western segment of Kunlun Fault.

The Qaidam Basin is bordered to the northwest by the Altyn Tagh Fault which has been active since the Early or Middle Eocene (Jolivet et al., 1999, 2001; Yin et al., 2002). In addition, despite that Chen et al. (2002) concluded that the Qaidam Basin experienced an over 20° clockwise rotation during the Neogene resulting in differential shortening in the Nan Shan fold

and thrust belt (Fig. 1), Dupont-Nivet et al. (2003, 2004) suggested that the Nan Shan fold and thrust belt and the Qaidam Basin did not experience any vertical axis rotation during the Neogene. As the limit between the thick, rigid Tarim block and the more deformable Qaidam Basin, the Altyn Tagh Fault is considered stable in time. Consequently, the deformation in the SW Qaidam Basin and the western segment of the Kunlun Fault should be bordered and limited to the NW by the Altyn Tagh Fault. When considering the geometric relation between the Altyn Tagh Fault and Kunlun Fault, the western Kunlun fault segment has to rotate northward: the left-lateral movement on the Altyn Tagh is progressively driving the connection point between that fault and the Kunlun fault towards the east. For the Kunlun Fault to be able to continuously act as a strike-slip fault, it would jump backwards towards its initial position creating a new strike-slip segment. As mentioned above, the distribution and focal mechanisms of earthquakes indicate that the active tectonic pattern along the Kunlun Fault is characterized by left-lateral strike-slip motion. However, in the SW Qaidam Basin and the Qimen Tagh Range the deformation is accommodated by NW-SE directed thrust faults associated to a limited left-lateral strike-slip component (Fig. 3). In addition, the geomorphological WNW trending linear structures in the Qimen Tagh Range (Fig. 5) and deformed Pleistocene sequences in the Adatan Valley (Figs. 8 and 9) record the left-lateral strike-slip movements within the Qimen Tagh Range since the Pleistocene. Finally, the overall tectonic and structural architecture within the SW Qaidam Basin is dominated by a series of high angle faults. The sudden variation in sediment thickness marked by the offset of the isopachs on both side of the faults (Fig. 10) associated to the growth strata and the positive flower structures generated by movements on those faults (Figs. 11, 12 and 13), indicate Early Miocene left-lateral strike-slip faulting in SW Qaidam Basin.

Synthesizes the above analysis, we thus propose that the Late Neogene northward growth of the convex-northward structures in the Qimen Tagh Range and SW Qaidam Basin (northern margin of the Tibetan Plateau) is driven by progressive northward migration of the left-lateral strike-slip Kunlun Fault (Fig. 14).

Around the Early Miocene, the left-lateral strike-slip faulting in the Kunlun Fault Zone initiated. This motion lead to E-W extension along the northern margin of the Tibetan Plateau and the deeply rooted strike-slip faults were acting as pathways for the shoshonitic magmas (Jolivet et al., 2003), leucogranites and rhyolites (Zhang et al., 2012) to reach the surface. During that period, the parallel Kunbei, Arlar and Hongliuquan faults (designated below as the Kunbei fault system) were forming the western segment of the Kunlun Fault, located in the same position as the present-day Kunlun Fault with a nearly E-W striking trend. The left-lateral movement between the Qiangtang-Bayan Har and Qaidam blocks caused the en-échelon arrangement of the faults and generated the growth strata recording the onset time (Early Miocene) of fault activation. Following this initial phase, the Kunbei fault system, progressively rotated clockwise and migrated northwards, forming the successive high-angle left-lateral compressional faults now observed in the SW Qaidam Basin. In order to maintain the strike-slip motion along the western segment of the Kunlun Fault, a new strike-slip fault system, the Qimen Tagh fault system (including the MQF and SQF) developed during the Pleistocene in place of the older western Kunlun Fault segment that migrated northward. This second phase is recorded by the deformed coarse-grain Pleistocene deposits in the Adatan Valley. This newly formed Qimen Tagh strike-slip fault segment again migrated northwards and progressively rotated clockwise forming the present-day Qimen Tagh prism. Another new strike-slip fault segment developed on

the western end of the Kunlun Fault, replacing the previous one. During the process of northward migration of these left-lateral strike-slip faults (the Kunbei Fault, the Arlar Fault, the Hongliuquan Fault, the MQF and the SQF), their motion changed from left-lateral strike-slip to nearly NE-SW compressional motion. The characteristic northward-convex bow-like shape of the Kubei fault system and of the Qimen Tagh prism is governed by the strike-slip motion along the Altyn Tagh Fault to the northwest. The northward migration of the western segment of the Kunlun Fault explains the lack of Early Miocene sediment records at the present day location of this western segment that would relate to the initiation of the Kunlun Fault.

The model proposed above is an evolution of the previously proposed transpressional model (Meyer et al., 1998; Jolivet et al., 2001; Tapponnier et al., 2001). It identifies the Kunbei fault system and Qimen Tagh fault systems as the western segment of the Kunlun Fault Zone and describes their progressive northward migration since Early Miocene. Their initiation as strike-slip faults explains the present-day steepness of the fault planes.

Based on the analysis of geological features in the Jingyu Basin (North Tibet) and the intrusion of Tertiary to Quaternary shoshonites along the western segment of the Eastern Kunlun Mountains, Meyer et al. (1998) and Jolivet et al. (2003) suggested that the strike-slip Kunlun Fault initiated during the Early Neogene. Recently, after detailed geochemistry and geochronology studies on the Miocene leucogranites and rhyolites found in the Hoh Xil Lake area (north of the Kunlun Fault) (Fig. 1), Zhang et al. (2012) suggested that the dehydration melting in this area was probably triggered by localized E - W stretching decompression within the left-lateral strike-slip Kunlun Fault system around 15 Ma. Meanwhile, Duvall et al. (2013) and Yin (2010) also considered that the left-lateral East Kunlun transpressional system initiated around 30 to 20 Ma as

a response to the collision between India and Asia. These ages of initial faulting in the Eastern Kunlun Mountains are roughly consistent with our results.

In addition, from the data described above, the boundary between the SW Qaidam Basin and the Eastern Kunlun Mountains is characterized, since the Early Miocene by a series of northward migrating and eastward rotating left-lateral strike-slip faults rather than by a continuous northward thrusting (e.g. Burchfiel et al., 1989; Meyer et al., 1998) or southward thrusting system (e.g. Yin et al., 2007; Wang et al., 2011). Consequently, the SW Qaidam Basin did not suffer a large amount of N-S compression and shortening before the Early Miocene. This observation is compatible with the Paleo-Qaidam model of Yin et al., (2008) which suggests the existence, during the Paleogene, of a large topographic depression encompassing the Hoh Xil Basin to the south and the Qaidam Basin to the north. This wide basin was partitioned by the uplift of the Eastern Kunlun Mountains during Late Oligocene - Early Miocene times.

7. Conclusions

The Eastern Kunlun Mountains and adjacent Qaidam Basin present one of the largest topographic gradients in the Tibetan Plateau. In that respect, determining the tectonic relations between the Qaidam Basin and the Eastern Kunlun Mountains and deciphering the complex tectonic architecture of the western segment of the Kunlun Fault has major implications in understanding the tectonic, sedimentary and geomorphic Cenozoic history of North Tibet.

The distribution and focal mechanisms of earthquakes presented in this study indicate that the active tectonic pattern in the SW Qaidam Basin and the Qimen Tagh Range is dominated by transpressional structures while that of the Eastern Kunlun Mountains is characterized by left-lateral strike-slip faulting. Geomorphological interpretation of remote sensing images shows that the Qimen Tagh Range is affected by obvious WNW trending linear structures which relate to left-lateral strike-slip motion. The deformed Pleistocene sediments within the Adatan Valley strongly suggest strike-slip faulting along those WNW trending faults during the Pleistocene. To the north, in the SW Qaidam Basin, the offset isopachs of the Late Eocene sediments and the en-échelon arrangement of Kunbei, Arlar and Hongliuquan faults again indicate left-lateral strike-slip faulting. This, associated to growth strata, implies an Early Miocene onset age for the left-lateral strike-slip deformation of the Kunbei fault system in the SW Qaidam Basin. This tectonic pattern and the southwards younging of the left-lateral strike-slip phases lead us to identify the Kunbei Fault system and the Qimen Tagh faults system as former western segments of the Kunlun Fault once located further south in the present-day location of that fault. These faults gradually migrated northward since the Early Miocene while their kinematics changed from

left-lateral strike-slip motion to NE-SW compression.

We conclude that the SW Qaidam Basin has been bordered by a series of strike-slip faults to the south since the Early Miocene, rather than by a continuous northward thrusting or southward thrusting system. The SW Qaidam Basin did not suffer major N-S compression and shortening before the Early Miocene.

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Figure Caption

Fig. 1. Digital topographic map of the western Qaidam Basin and surrounding area. Topography from SRTM data.

Fig. 2. Geological map of the Qimen Tagh Range and SW Qaidam Basin. See Fig. 1 for location and the text for discussion. Modified after the Kulangmiqiti J46C003001 1:250,000 geological map (IGSQP, 2004).

Fig. 3. General topographic map of the Qaidam Basin and Eastern Kunlun Mountains with background seismicity from the China Earthquake Networks Center (CENC) from 1st January 1970 to 1st January 2011 ($M > 3$).

Fig. 4. Depth-frequency distribution of earthquakes within the study area. Data from the China Earthquakes Networks Center (CENC), 1st January 1970 to 1st January 2011 with $M > 3$. Two seismic profiles (AA' (a) and BB' (b) (See location on Fig.3) are investigated with (left) depth distribution of earthquakes and (right) depth - frequency distribution of seismic events. See text for discussion.

Fig. 5. Remote sensing image from Landsat 5 on December 5, 2008 (USGS) (a) and SRTM digital topography map (b) of the Qimen Tagh Range. The red solid lines show the main tectonic structures. See the location in Fig. 1.

Fig. 6. Un-interpreted (a) and interpreted SPOT-5 image (b) (Google Earth) centered on the western part of the Middle Qimen Tagh Fault (MQF). See the location in Fig. 5. The red solid lines indicate the fault segments imaged by fault scarps on the surface. The dashed line represents the blind fault covered by the Quaternary sediments. The fault scarps in en-échelon arrangement indicates a left-lateral strike-slip movement on the MQF.

Fig.7. Un-interpreted (a) and interpreted SPOT-5 image (b) (Google Earth) centered on the eastern part of the Middle Qimen Tagh fault. See the location in Fig. 5. The red dash line indicates the linear fault linking the Qimen Tagh fault system in the west with the Eastern Kunlun Fault system in the east.

Fig. 8. Cenozoic formations within the Adatan Valley. (a) General view of the Qimen Tagh Range.

The picture, taken from ~10 km south of the Atantihan river checkpoint, shows the strong relief between the Adatan Valley and the mountains bordered by a series of triangular facets and fault scarps belonging to the MQF. The foothills domain is represented by Cenozoic strata which are deformed and pushed into gentle folds. See the location in Fig. 9. (b) Uplifted Pleistocene series covered by Holocene tillites (left) and general stratigraphic column (right). See the photographs' location in Fig. 8.

Fig. 9. Geological sketch of the Qimen Tagh Range showing the distribution of the Cenozoic series. Modified after the Kulangmiqiti J46C003001 1:250,000 geological map (IGSQP, 2004) and the Ayakekumu J45C003004 1:250,000 geological map (IGSSP, 2003). See the location in Fig. 2.

Fig. 10. Isopach map of the Upper Xiaganchaigou Formation (E_3^{2xg}). Three major faults (the Kunbei Fault, the Arlar Fault and the Hongliuquan Fault) are also marked on the isopach map based on seismic data. These data were acquired by the Qinghai Oilfield Company (China) who compiled information from thousands of drill holes and a dense network of seismic profiles in recent years. Note the offset of the 400 isopach line controlled by the faults and indicating the left-lateral strike-slip faulting of these faults.

Fig. 11. Seismic profile C - C' in the SW Qaidam Basin. See the location in Figs. 2 and 10. Note that the thickness of the Paleocene to Eocene strata shows a sharp difference between both sides of the Arlar Fault. The growth strata associated to the positive flower structure started to develop during the initial deposition of the Early Miocene unit (N_2^1xy), which suggests that the strike - slip

movement on the Arlar Fault initiated during the deposition of this unit.

Fig. 12. Seismic section D - D' in the SW Qaidam Basin. See Figs. 2 and 10 for location. Note that the strata below the Shangyoushashan Fm. (N_2^1xy) keep a constant thickness in the section, while the strata above it are thinning northward. The thickness of the Paleocene to Eocene strata shows a sharp difference between both sides of the Arlar Fault. Meanwhile, Early Miocene growth strata above the Arlar and Hongliuquan faults indicate the onset of left-lateral strike-slip faulting during the Early Miocene.

Fig. 13. Seismic profile E - E' in the SW Qaidam Basin. See the location in Figs. 2 and 10. Note that the thickness of the Paleocene to Eocene strata shows a sharp difference between both sides of the Kunbei Fault. Meanwhile, Early Miocene growth strata again indicate an onset of left-lateral strike-slip faulting during the Early Miocene.

Fig. 14. Tectonic evolution of the Qimen Tagh Range and the SW Qaidam Basin since the Early Neogene. See the text for discussion.

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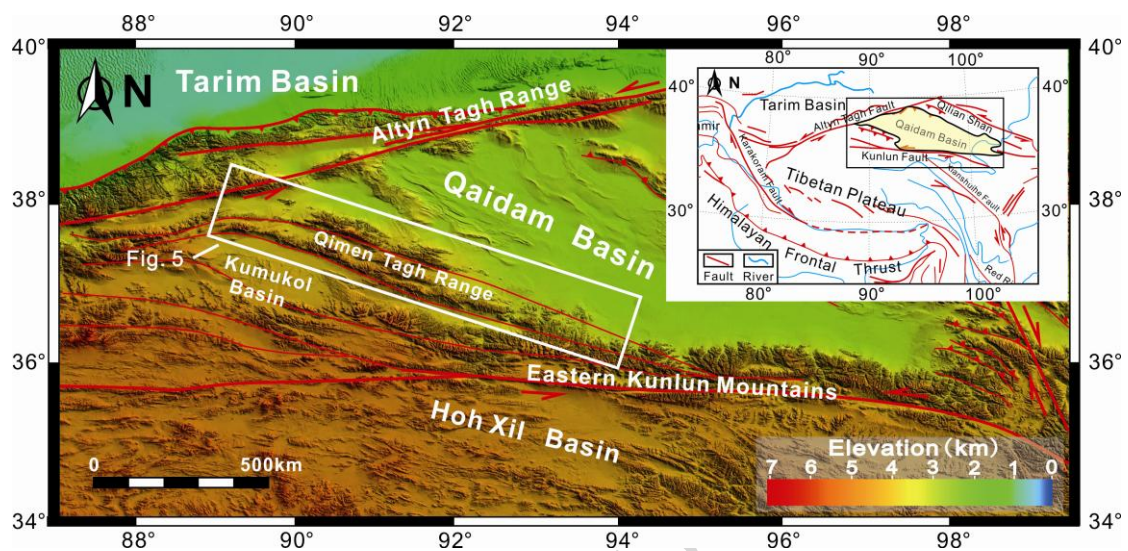


Fig1

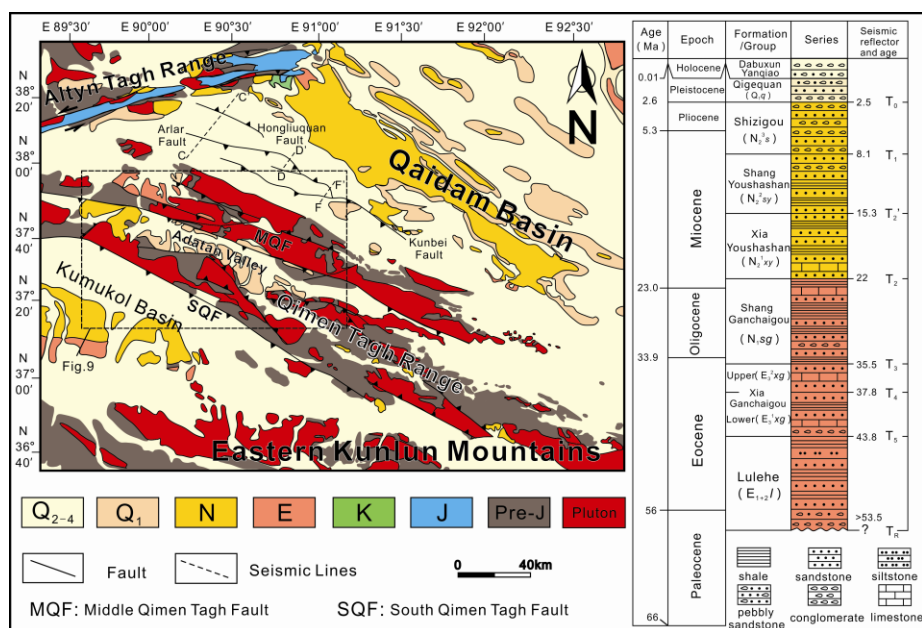


Fig2

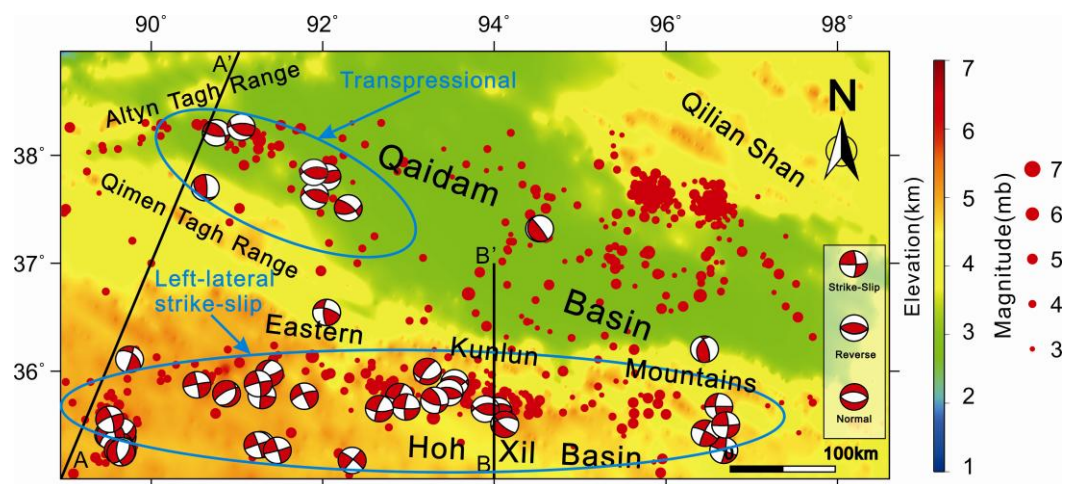


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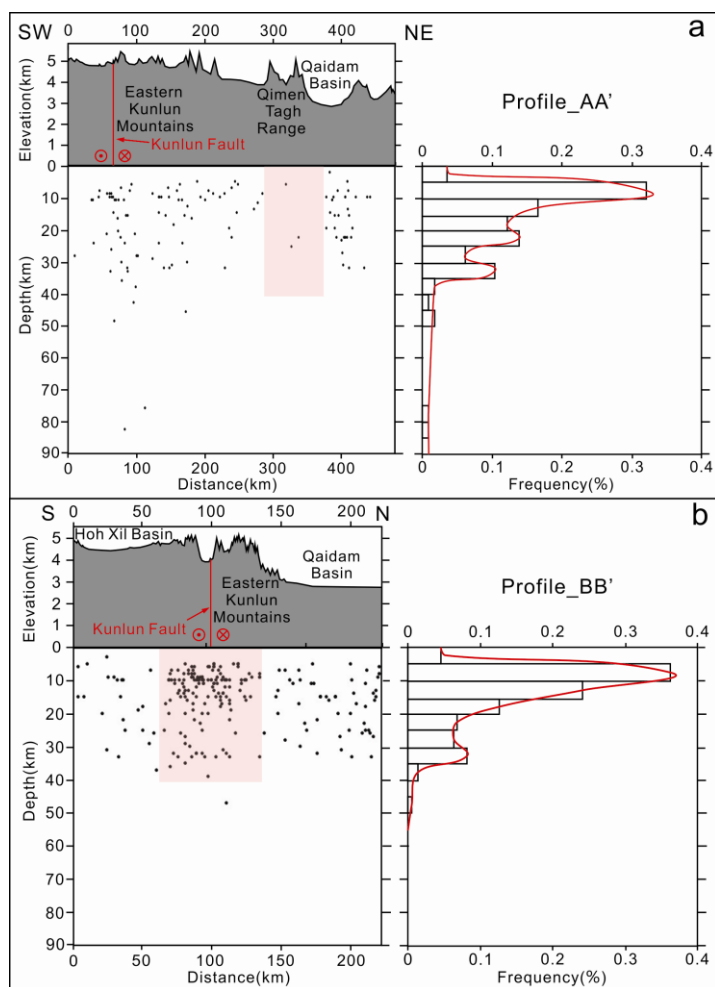


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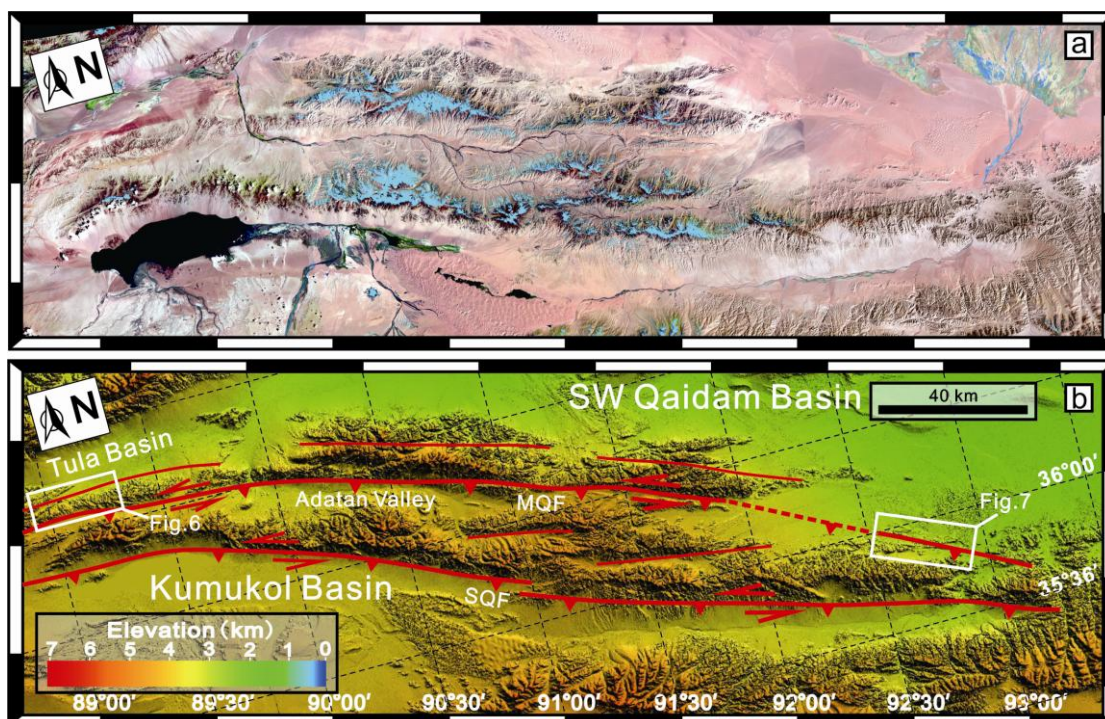


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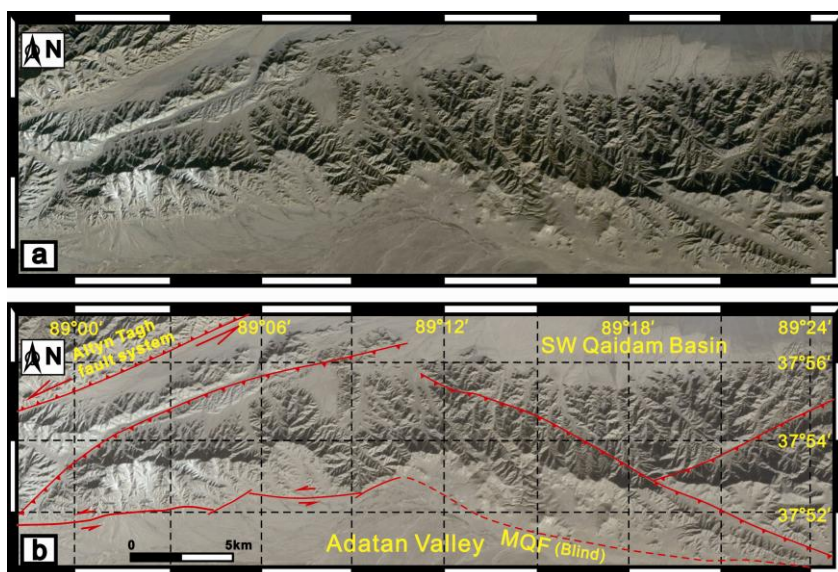


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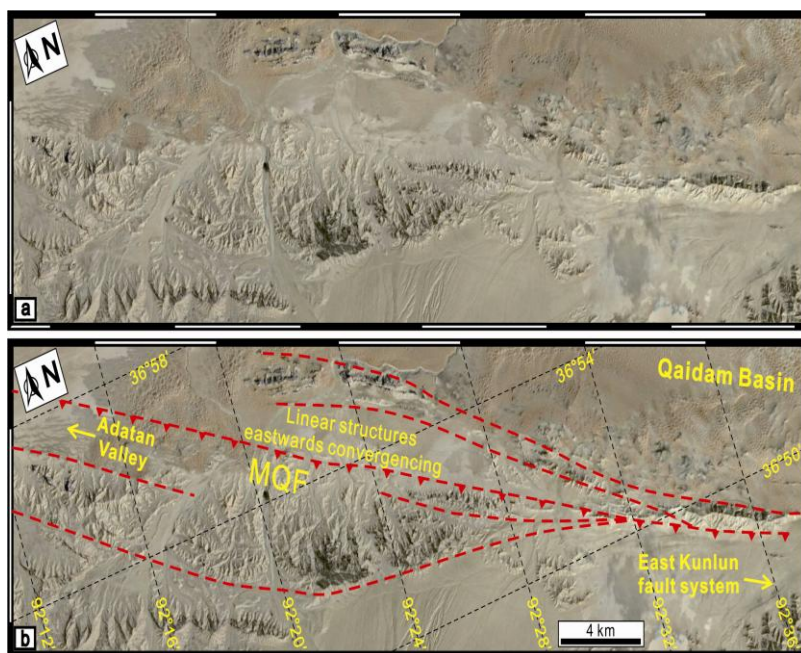


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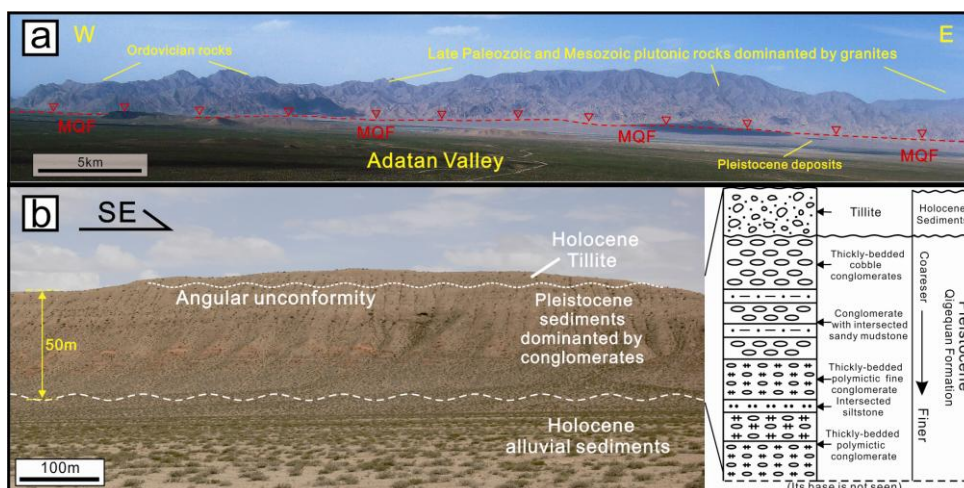


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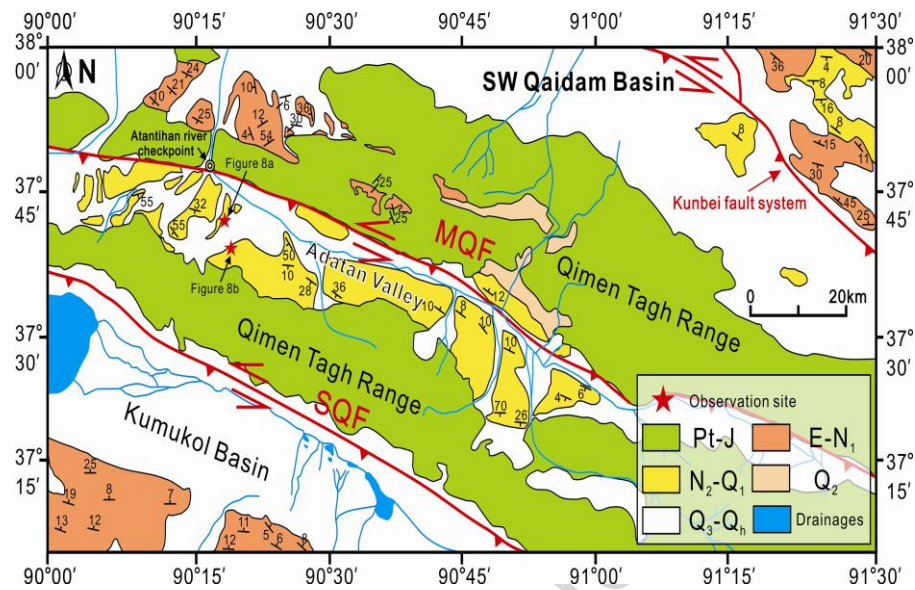


Fig9

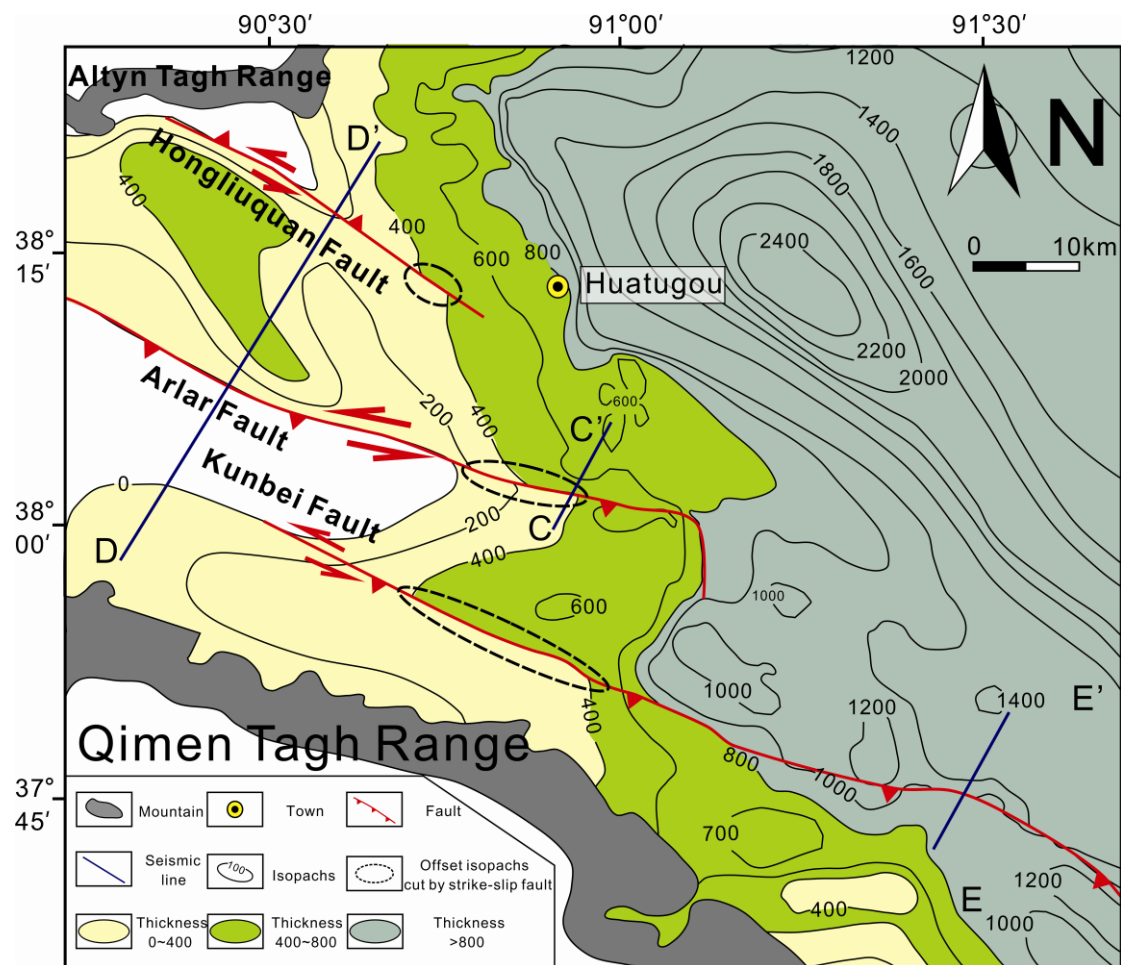


Fig10

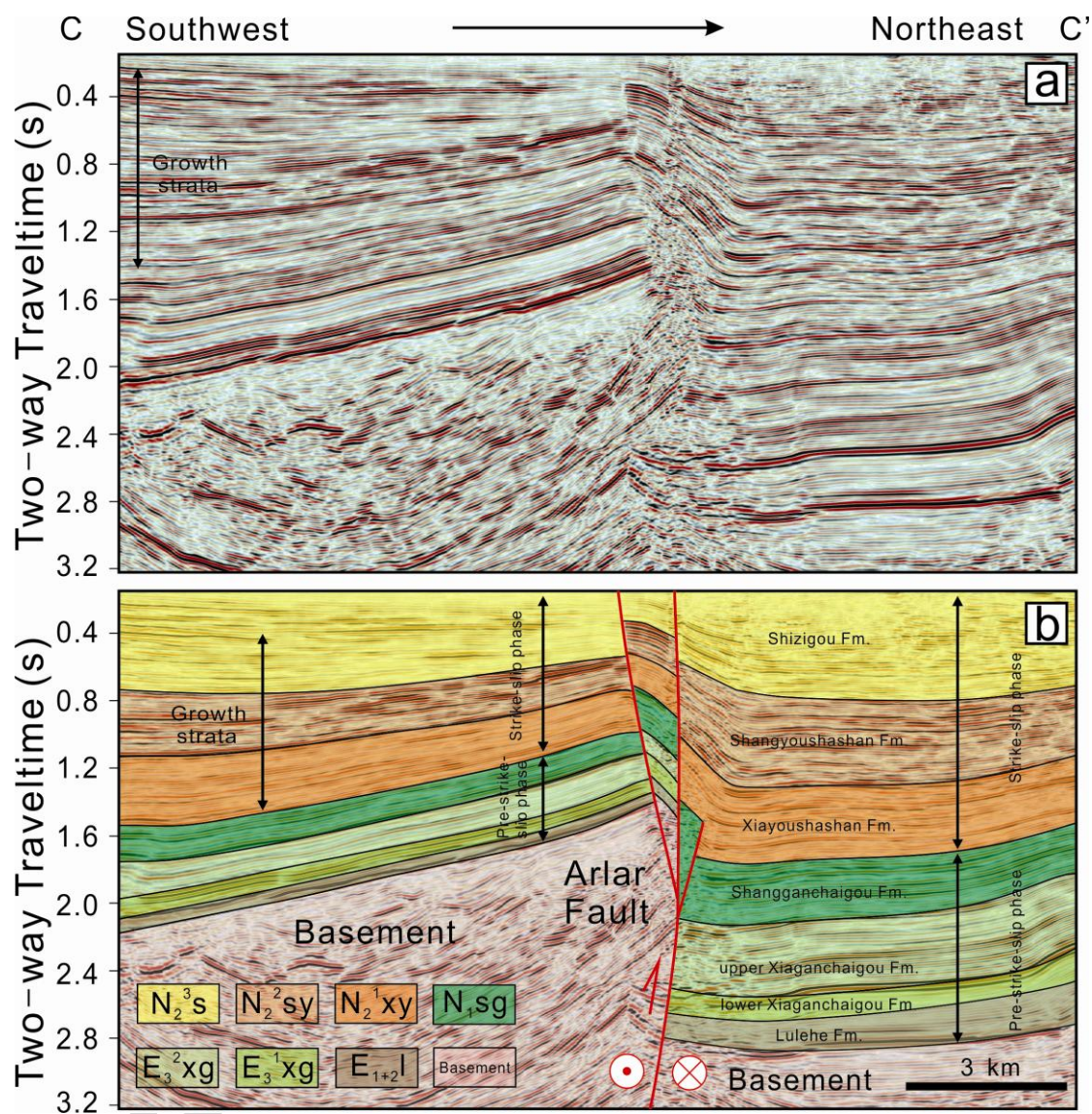


Fig11

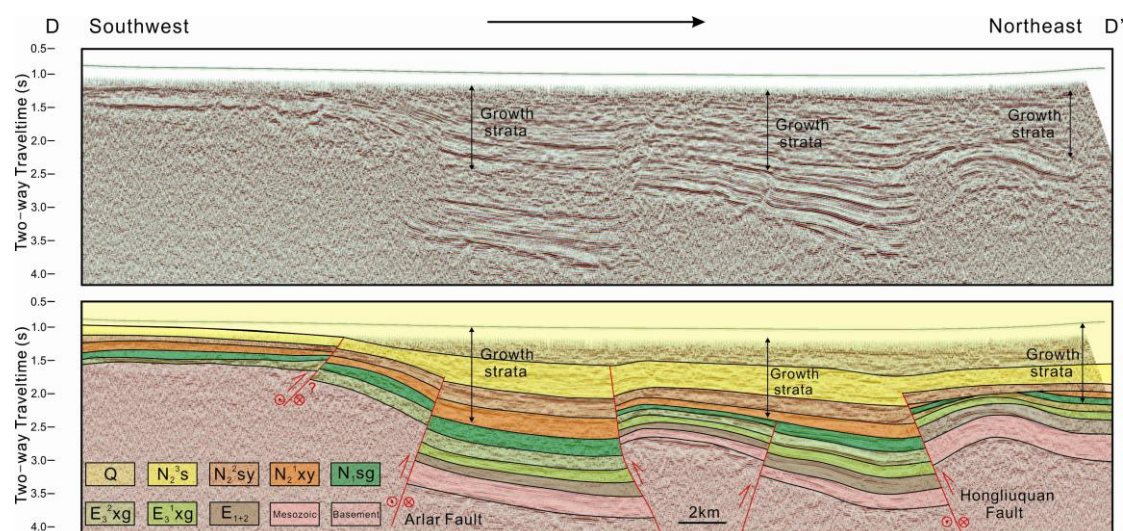


Fig12

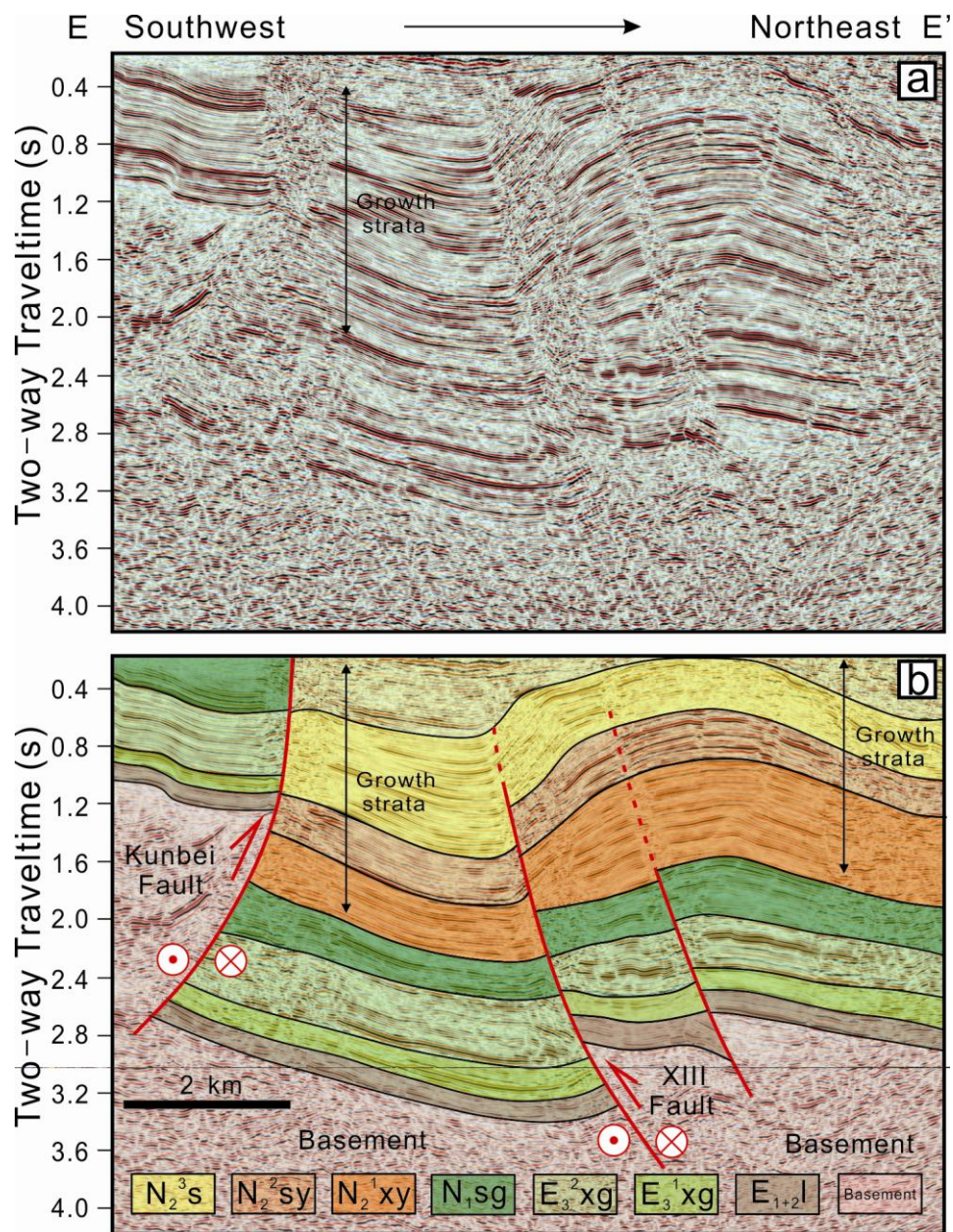


Fig13

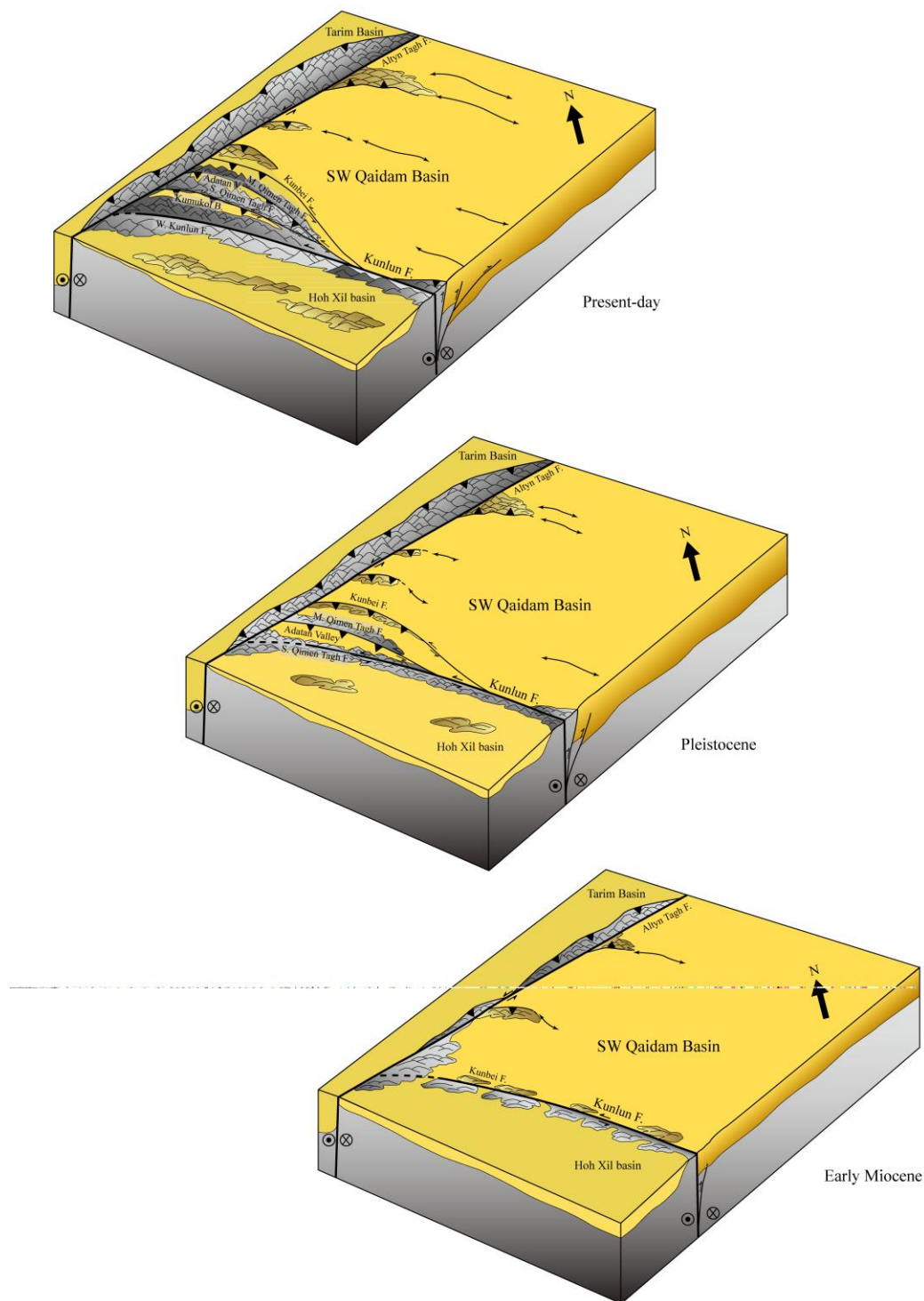
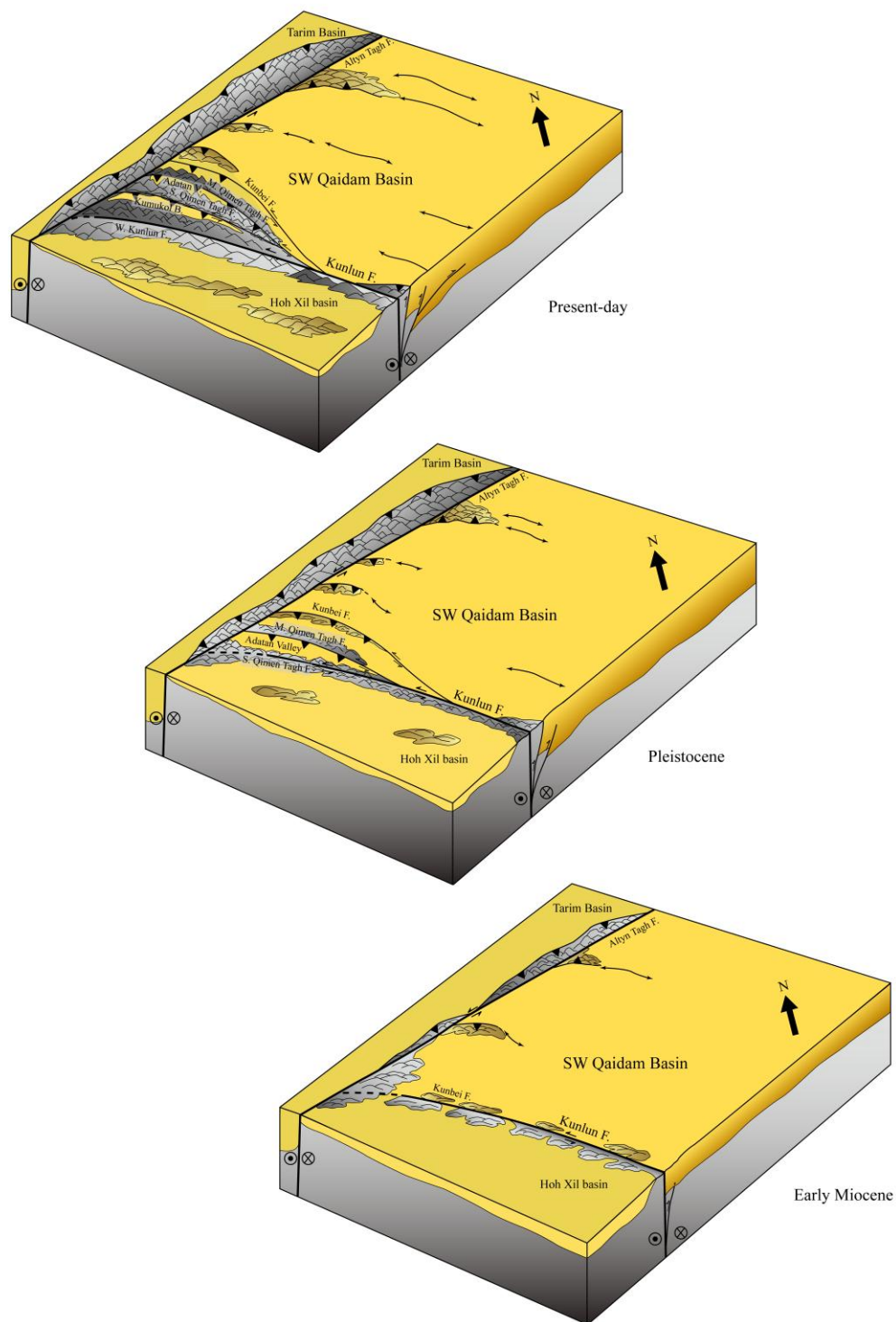


Fig14



Graphical abstract

Highlights

We identify sinistral strike-slip faulting in SW Qaidam Basin and Qimen Tagh Range

Sinistral strike-slip faulting prevailed in SW Qaidam during the Early Miocene

Sinistral strike-slip faulting prevailed in Qimen Tagh Range during the Pleistocene

These strike-slip faults once located further south migrated northwards gradually