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# Impact of Holocene climate variability on lacustrine records and human settlements in South Greenland

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## Abstract

Due to its sensitivity to climate changes, south Greenland is a particularly suitable area to study past global climate changes and their influence on local Human settlements. A paleohydrological investigation was therefore carried out on two river-fed lakes: Lake Qallimiut and Little Kangerluluup, both located close to the Labrador Sea in the historic farming center of Greenland. Two sediment cores (QAL-2011 and LKG-2011), spanning the last four millennia, were retrieved and showed similar thin laminae, described by high magnetic susceptibility and density, high titanium and TOC / TN atomic ratio, and coarse grain size. They are also characterized either by inverse grading followed by normal grading or by normal grading only and a prevalence of red amorphous particles and lignocellulosic fragments, typical of flood deposits.

Flood events showed similar trend in both records: they mainly occurred during cooler and wetter periods characterized by weaker Greenlandic paleo-temperatures, substantial glacier advances, and a high precipitation on the Greenlandic Ice Sheet and North Atlantic ice-rafting events. They can therefore be interpreted as a result of ice and snow-melting episodes. They occurred especially during rapid climate changes (RCC) such as the Middle to Late Holocene transition around 2250 BC, the Sub-boreal/Sub-atlantic transition around 700 BC and the Little Ice Age (LIA) between AD 1300 and AD 1900, separated by cycles of 1500 years and driven by solar forcing. These global RCC revealed by QAL-2011 and LKG-2011 flood events may have influenced Human settlements in south Greenland, especially the paleo-Eskimo cultures and the Norse settlement, and have been mainly responsible for their demise.

## 1 Introduction

The Holocene appears to be a steady climatic period in comparison with older glacial episodes, as climate evolution was not linear but showed several oscillations (O'Brien et al., 1995; Mayewski et al., 2004; D'Andrea et al., 2011; Larsen et al., 2012), result-

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ing from various forcing mechanisms such as solar activity (Bond et al., 2001; Magny, 2013) volcano emissions (Miller et al., 2012; Sigl et al., 2013) and/or ocean/atmosphere interactions (Hurrell and Denser, 2010; Knudsen et al., 2011). It is of prime importance to establish their spatial influence and phasing in order to understand natural climate oscillations and their influence on past Human societies and thus to create better predictive climatic models. Indeed, even if current global warming is mainly the result of an anthropogenic pressure, the contribution of natural climate variability cannot be excluded and may explain the recent rapid rise in temperatures (Chylek et al., 2004; Solanki et al., 2004).

Projected climate scenarios predict a rise in surface temperature associated to an increase in the frequency and intensity of flood events in the coming decades (IPCC, 2013), especially in the Arctic region, which is very sensitive to climate changes (Serreze and Barry, 2011). Due to its proximity to the Greenlandic Ice Sheet (GIS), the polar front and the adjacent North Atlantic oceanic currents, south Greenland is a particularly suitable area to record in natural archives past global climate changes and their associated forcing mechanisms at high resolution (Fig. 1a). While numerous proxy records have been used to infer past temperature in the Arctic (D'Andrea et al., 2011; McKay and Kaufman, 2014; Millet et al., 2014), few records have been used to infer past hydrological activity (Sundqvist et al., 2014) despite the extreme importance of hydrological conditions for Arctic ecosystems and their potential feedback on global climate (MacDonald et al., 2000).

To assess of past hydrological conditions during the Holocene, a sedimentological multiproxy approach was performed, compiling physical (magnetic susceptibility, density and grain size) and geochemical (X-ray fluorescence, X-ray diffraction, elemental analyses and organic geochemistry) measurements, on two lacustrine cores. They were retrieved from two river-fed lakes: Lake Qallimiut and Little Kangerluluup located between the GIS and the Labrador Sea in the Vatnahverfi, known as the archeological farming center of Greenland where two agropastoral phases settled (the Norse period between AD 986 and ca. AD 1450 and over the last century, since the 1920s). Thus,

this study reconstructs at high resolution the global climatic variations and the large scale drivers of Greenlandic climate through hydrological changes in lacustrine catchments that may have influenced past Human settlements (i.e. vulnerability, adaptation strategies, demise, etc. . . ) during the last four millennia.

## 2 Settings

### 2.1 Localization and climate

The study area is located between the GIS and the Labrador Sea (Fig. 1a and b). Due to this geographical position, the local climate is mainly influenced by oceanic currents, such as cold polar waters ( $< 0^{\circ}\text{C}$ ) of the eastern Greenland current and warm waters ( $+ 4^{\circ}\text{C}$ ) of the Irminger current, which mix to the west of Kap Farvel (Bond et al., 1997). The area is also windy with strong and relatively warm foehn winds blowing from the ice cap throughout the year, that are responsible for dryness. Meteorological data of the last 60 years are characteristic of a sub-oceanic climate, with a mean annual temperature of  $0.9^{\circ}\text{C}$ , a yearly precipitation of 907 mm, 211 days with frost and 61 days of snow per year (Cappelen et al., 2001; Cappelen, 2014).

### 2.2 The Qallimiut lake system

Lake Qallimiut ( $60^{\circ}43'27''\text{N}$ ;  $45^{\circ}23'12''\text{E}$ ; 40 m a.s.l.) is a proglacial lake located close to the hamlet of Qallimiut ca. 2 km to the south, near the Alluitsup Kangerlua fjord (Fig. 1c). It is aligned through a north-west/south-east axis with a surface of ca. 82 ha and a maximum depth of 26 m. Two main tributaries feed it (on the north-west and the east shores respectively) while a single outlet drains it directly in the Alluitsup Kangerlua fjord (Fig. 1c). Its catchment, composed of granites (geological survey of Denmark and Greenland; <http://www.geus.dk>), has a surface of  $112\text{ km}^2$  and a catchment to lake ratio around 137. It includes around twenty upstream lakes, making Lake Qallimiut the last sediment trap of this catchment (Fig. 1c). Less than ten ruins, dated from the Norse

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period, were inventoried in the catchment with only two sites directly found around the lake (Madsen et al., 2009). These two Norse archaeological sites have already been surveyed. Ruins group E77, a medium-sized farmstead, is located in the north-west part of the lake and was formerly excavated in AD 1880 (Holm, 1883; Madsen, 2014), and ruins group E77a, near the coast, seems to be a large, highly eroded farmstead (Madsen, 2014). Nowadays, two recent farms can be found on the north and south shores of the lake (the largest one is located near the lake outlet).

**2.3 The Little Kangerluluup lake system**

Lake Little Kangerluluup (60°38'32" N; 45°38'11" E; 10 m.a.s.l.) is also a proglacial lake, located close to Kangerluluup Qaqaa hill and Lake Kangerluluup, near the Torsukattak fjord (Fig. 1c). It is quite circular with a tiny surface of 6 ha and a maximum depth of 6.4 m. Lake Little Kangerluluup is fed by only one major tributary and is drained directly into the Torsukattak fjord by one outlet (Fig. 1c). Like Lake Qallimiut, its catchment is composed of granites (geological survey of Denmark and Greenland; <http://www.geus.dk>). It has a relatively small area (8 km<sup>2</sup>) with a catchment to lake ratio of ca. 133, quite similar that of Lake Qallimiut. The lake catchment is aligned along a north-east/south-west axis with a very steep gradient slope, reaching 900 to 10 m within only 2 km (Fig. 1c). Unlike Lake Qallimiut, only one upstream lake flows into Lake Little Kangerluluup (Fig. 1c). Norse ruins (E178) are located on the isthmus of the far side of the river and were surveyed by Madsen et al. (2009) and Madsen (2014). This large farmstead with barns, enclosures and irrigation systems, attests to agropastoral activities in the catchment during the Middle Ages. Nowadays, there are no farms around the lake, indicating that no human activities took place during the last century around Lake Little Kangerluluup and more generally, at the end of the Torsukattak fjord.

## 2.4 Archaeological context

Human occupation began around 2500 BC with the Saqqaq culture in the southern half of Greenland (Meldgaard, 2004; Jensen, 2009). These small populations of nomad hunters were replaced around 800 BC by the sea-ice hunters of the Greenlandic Dorset culture, who disappeared abruptly around 250 BC (Meldgaard, 2004; Jensen, 2009). Greenland remained unpopulated until the Norse colonization in 986 AD (Jones, 1986). This society, established during the medieval warm period (MWP), was mainly marked by agropastoral activities (Buckland et al., 2009; Arneborg et al., 2012; Guillemot et al., 2015) and ended around AD 1450 (Dugmore et al., 2009, 2012). The ancestors of the modern Greenlandic, the Thule (neo-Eskimos), settled throughout Greenland around AD 1100 and still present today (Gulløv 1983, 2004). Due to recent global warming, sheep farming was established in the area by Danish in the 1920s (Austrheim et al., 2008). Several paleo-environmental studies revealed that farming by the Norse and during the last century, corresponding to the only two agropastoral phases in south Greenland, had a strong impact on the ecosystems through vegetation changes (Gauthier et al., 2010; Schofield et al., 2011), soil erosion (Massa et al., 2012a) and changes in the lake trophic state (Perren et al., 2012; Millet et al., 2014).

## 3 Materials and methods

### 3.1 Fieldwork

In the deepest part of each lake basin (26 m for Lake Qallimiut and 6.4 m for Lake Little Kangerluluup), two cores were retrieved during the summer 2011: QAL-2011 (118 cm long) and LKG-2011 (109.5 cm long), using a UWITEC gravity corer.

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## 3.2 Non-destructive laboratory analyses

Initial core analyses of QAL-2011 and LKG-2011 included macroscopic description, digital photographs and X-ray radiographs (Scopix system, EPOC laboratory, University of Bordeaux, France), providing high-resolution imaging. X-ray radiographs are based on 256 grey levels corresponding in value to X-ray densities (Migeon et al., 1998).  $\gamma$ -ray attenuation density and magnetic susceptibility (MS) measurements were performed with a GEOTEK multi-sensor core logger in a millimeter resolution (Chrono-environnement laboratory, University of Bourgogne/Franche-Comté, France). Elementary analysis was performed with an X-ray microfluorescence core scanner (XRF, AVAATECH system, EDYTEM laboratory, University of Savoie/Mont Blanc, France) at a resolution of 2 mm in two distinct runs. The first one, with a counting time of 20 s and a 10 kV acceleration intensity, was followed by a second with a exposure time of 45 s and a 30 kV acceleration intensity to obtain the relative abundance of elements from Al to Bi. Titanium (Ti), a conservative element and a marker of detrital fluxes (Cohen, 2003; Arnaud et al., 2012), is expressed here in peak area counts. Usually, elements counts are normalized so as to correct drifts coming from variations in the water content or the grain size (Tjallingii et al., 2007; Revel et al., 2010). However, since the normalization of Ti by another chemical compound only smooths the Ti content trends, we used raw Ti counts directly in the present study. To determine the mineralogical composition of the sediment, X-ray diffraction (XRD) analyses were carried out in whole sediment using a D8 Advance Brucker diffractometer equipped with a LinxEye detector (Utinam laboratory, University of Bourgogne/Franche-Comté, France). The crumbled sediment is placed in the specimen holder where a step-scan of  $0.5^\circ$  in the  $2\text{--}52^\circ$  interval during a step time of 0.05 s is realized using a parallel beam geometry and  $\text{CuK}\alpha$  radiation at  $1.54184 \text{ \AA}$ .

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### 3.3 Destructive analyses

QAL-2011 and LKG-2011 were sampled every 0.5 cm slices, based on X-ray images in order to respect lithological boundaries and provide homogenous samples.

84 samples from QAL-2011 and 85 samples from LKG-2011 were therefore used for grain size (GS) analyses, respectively. Around one cm<sup>3</sup> of wet sediment was immersed in a solution of distilled water and hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>, 30 %) to digest organic matter. Residual mineral fractions were then treated in an ultrasonic bath with a solution of sodium-hexametaphosphate to disperse aggregates. Finally, the GS of each sample was determined using a laser diffraction particle analyzer (LS230 Beckman-Coulter) and represented by GS classes, according to granulometric scales (< 4, 4–16, 16–63, > 63 μm). Regions of interest (ROIs) for sedimentary micro-structure identification were selected based on the initial sedimentary description of the split cores as well as on the results of geophysical and geochemical core logging. The selected ROIs cover sediment transitions, including larger laminae (at 32.2 cm in core QAL-2011 and at 41.3 cm in core LKG-2011). Thin sections (one for each core) were sampled from these ROIs and examined using a polarized light microscope.

Total carbon (TC), nitrogen (TN) and sulfur (TS) contents were measured on 161 and 187 selected samples, in QAL-2011 and LKG-2011 respectively, using a vario MAX CNS analyzer (Elementar). As the carbonate fraction was negligible in the sediments because of the Arctic climate and the granitic catchments, the total carbon measured corresponds to the total organic carbon (TOC). It was thus possible to access the relative contribution of lacustrine and terrestrial organic matter (OM) in the sediments via the TOC/TN atomic ratio (Meyers and Ishiwatari, 1993). To identify and quantify the composition of the organic fraction, quantitative organic petrography (QOP), developed by Graz et al. (2010), was performed on the ROIs described above. After elimination of the mineral phases by hydrochloric and hydrofluoric attacks, particles were inventoried and counted according to their shape, color and reflectance using a transmitted and reflected light microscope (Combaz, 1964). Red amorphous particles (rAP), grey

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amorphous particles (gAP) and lignocellulosic fragments (LCF), in particular, were described following the classification frequently used in lake sediments (Simonneau et al., 2013a, b, 2014; Foucher et al., 2014).

### 3.4 Core chronology

The chronology of each core was based on six AMS radiocarbon datings of terrestrial plant macrofossils (Fig. 2a and b; Table 1a and b) conducted in the Lyon radiocarbon laboratory (Lyon, France) and Beta Analytics laboratory (Miami, United States). For the last 150 years, the chronologies of Qal-2011 and LKG-2011 were established thanks to short-lived radio-isotopes ( $^{210}\text{Pb}$ ,  $^{137}\text{Cs}$ ,  $^{241}\text{Am}$ ) measured by gamma spectrometry in the underground laboratory in Modane (LSM-CNRS/CEA, France; Fig. 2a and b). Measurements were made on contiguous 0.5 cm-thick samples of 0.8 to 1.7 g dried sediment, using a high efficiency, well-type Ge detector, during 24 to 72 h (deepest samples with lowest activities). The excess  $^{210}\text{Pb}$  (formed in the atmosphere by the decay of  $^{222}\text{Rn}$ ) was calculated as the difference between the activities of total  $^{210}\text{Pb}$  and  $^{226}\text{Ra}$  (supported  $^{210}\text{Pb}$ ; Appleby, 2001).  $^{210}\text{Pb}$  chronology was derived for both lakes using the Constant Rate of Supply Model (CRS; Appleby and Oldfield, 1978). This type of model was chosen because it allows the sedimentation rate to vary. Results from radiocarbon and short-lived radio-isotopes were then combined with Clam (Blaauw, 2010) to produce two age-depth models using linear interpolation between date points (Fig. 2a and b). The age-depth models were corrected for instantaneous flood event layers by subtracting the thickness of these layers before computation.

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## 4 Results

### 4.1 Chronology

#### Lake Qallimiut

Core QAL-2011 covers the last four millennia. The age-depth model is quite linear (Fig. 2a) and shows steps, especially in the first thirty cm, due to an increased number of instantaneous events (see below, Sect. 5.1). The mean sediment accumulation rate (SAR) is  $0.3 \text{ mm year}^{-1}$  and the average age error is 160 years, not considering the extrapolated section of the model which is unconstrained by a radiocarbon date (below 90 cm, in black dotted lines; Fig. 2a). The last century is contained in the top three cm, as indicated by the peak of  $^{137}\text{Cs}$  between 1.5 and 2.5 cm, corresponding to the maximum fallout consecutive to atmospheric nuclear weapon tests around AD 1963 (Fig. 2a).

#### Lake Little Kangerlulup

Similarly, the age-depth model of LKG-2011 covers the last four millennia. It is quite linear, shows several steps due to instantaneous events and has a mean SAR of  $0.3 \text{ mm year}^{-1}$  (Fig. 2b). A slight rise in the SAR up to  $0.4 \text{ mm year}^{-1}$  is recorded during the last millennium, which is not visible in the QAL-2011 age-depth model. The average age error of LKG-2011 age-depth model is ca. 200 years, not considering the extrapolated section of the model, below 90 cm (in black dotted lines; Fig. 2b). The last century is recorded in the first 4 cm with the AD 1963 peak of  $^{137}\text{Cs}$  located at 2–2.5 cm (Fig. 2b).

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## 4.2 Lithology, physical and chemical properties of the laminated facies

### Lake Qallimiut

Three stratigraphic units (SU1 to SU3) were identified in core QAL-2011 (Fig. 3a).

The first one (SU1, from 118 to 95 cm) is characterized by a brownish silty sediment (with a principal GS mode of ca. 30  $\mu\text{m}$ , Fig. 3a) and interrupted by five thin white laminae closely spaced at 117, 113.9, 113.4, 110.4 and 100.9 cm (Table 2a). The MS is the only parameter that remains constant (around  $100 \times 10^{-5}$  S.I.) all along SU1 (Fig. 3a). The density and the Ti content show a slight reduction at the base (from 1.3 to 1.1  $\text{g cm}^{-3}$  for the density, and from 10 000 to 9 000 peak area for Ti, Fig. 3a) whereas the TOC increases slightly (from 4 to 8 %; Fig. 3). Finally, the mean value of the TOC/TN is 13 (Fig. 3a).

The second stratigraphic unit (SU2, from 95 to 32.5 cm) is the darker part of the QAL-2011 sequence due to its higher organic carbon content (around 8%). It is defined by a dark brown silty sediment (with a principal GS mode of 30  $\mu\text{m}$ , Fig. 3) interrupted by four white laminae at 81, 77.2, 66.8 and 36.7 cm (Table 2a). As previously, MS values remain stable at ca.  $100 \times 10^{-5}$  S.I. except for a peak at 66.8 cm (Fig. 3a). The density shows little variation around 1.1  $\text{g cm}^{-3}$  (Fig. 3a). In spite of slight variations centered on the white laminae, the mean Ti and TOC contents, as well as the mean TOC/TN ratio, 9 000 peak area, 8 % and 13 respectively (Fig. 3a).

The last sedimentary unit (SU3, from 32.5 cm to the top of the sequence) is characterized by a light brown sandy silt sediment (with a principal GS mode twice as high as that of the other two SUs at ca. 60  $\mu\text{m}$ , Fig. 3a). It is interrupted by a maximum of 18 white laminae spread all over SU3. These laminae generate abrupt shifts in all the measured parameters. MS rise to a mean value of  $125 \times 10^{-5}$  S.I., density, Ti content and TOC/TN ratio also increase slightly to values of 1.2  $\text{g cm}^{-3}$ , 10 000 peak area and 13.5, respectively (Fig. 3a), whereas TOC drops to 6 % (Fig. 3).

The white laminae, present all along the QAL-2011 record, are globally infra-centimetric, making naked-eye identification difficult in most cases (Table 2a). X-ray

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radiography is very helpful because the laminae are particularly well-defined in the high-resolution image, appearing in intense black, corresponding to high X-ray density values, whereas the background sediment is represented by light-grey color. Regardless of the SU, white laminae are characterized by higher MS (with a maximum value of  $450 \times 10^{-5}$  S.I. at 5 cm, Fig. 3a), density (maximum value of  $1.7 \text{ g cm}^{-3}$ , Fig. 3a) and TOC/TN ratio (maximum value of 16, Fig. 3a). They are also rich in Ti (maximum value of 20 000 peak area, Fig. 3a) and in coarse fraction (maximum value of 25 % in the  $> 63 \mu\text{m}$  massic class, Fig. 3a). Inversely, they are distinguished by much lower TOC (minimum value of 1.5 %, Fig. 3a). XRD tests indicated a similar mineralogical composition, composed of quartz, feldspaths and phyllosilicates, not different from the background sedimentation. The thicker lamina (1.8 cm thick, between 32.2 and 34 cm) showed a normal grading (fining upward, Fig. 3b) and was mainly composed of rAP and LCF, coming from soil and upper vegetation debris (Simonneau et al., 2013a), as part of other organic particles (Fig. 3b). On the basis of their similar physical and geochemical properties, the results obtained on the thicker lamina were then extrapolated to the other laminae.

### Lake Little Kangerlulup

Core LKG-2011 is also divided into three distinct stratigraphic units (SU1 to SU3, Fig. 4a).

The first one (SU1, from 108 to 74 cm) is characterized by a light brown sandy silt background sediment (with a principal GS mode of  $30 \mu\text{m}$ , Fig. 4a), interrupted by three thick brownish laminae (at 94.8, 81.6 and 79.7 cm) and by two thin white laminae (at 107.7 and 104.9 cm; Table 2b). MS values are much lower than in core QAL-2011, with background values of  $10 \times 10^{-5}$  S.I. (10 times less than in SU1 in core QAL-2011, Fig. 4a). Density, Ti content, TOC and TOC/TN ratio are quite similar to those observed in QAL-2011 SU1 (mean values of  $1.2 \text{ g cm}^{-3}$ , 6000 counts, 7 % and 14, respectively, Fig. 4a).

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The second stratigraphic unit (SU2, from 74 to 51 cm) is composed of a dark brown silty sediment with the same principal GS mode as SU1 (30  $\mu\text{m}$ , Fig. 4a), and is not interrupted by any laminae. All the measured parameters are particularly steady: MS has a value of  $5 \times 10^{-5}$  S.I.; density, a value around  $1.1 \text{ g cm}^{-3}$ ; Ti content, a value of 6000 peak area; TOC, a mean value of 10% and TOC/TN ratio, a value around 14 (Fig. 4a).

The last sedimentary unit (SU3, from 51 cm to the top of the core) corresponds to the darker and coarse part of core LKG-2011 with a black sandy silt sediment and a principal GS mode of 50  $\mu\text{m}$  (Fig. 4a). Five thick dark grey laminae, only visible in the X-ray radiographs, were identified at 49.2, 41.3, 29.2, 17, and 4 cm (Table 2b). The background MS value is  $5 \times 10^{-5}$  S.I. and density is  $1.2 \text{ g cm}^{-3}$  (Fig. 4a). There is also a pronounced rise in Ti content to 8000 peak area in SU3 (Fig. 4a). TOC has a background value around 9% and the TOC/TN ratio, a value of 14 (Fig. 4a).

Globally, the different types of laminae present within SU1 and SU3 (Table 2b), have similar physical and geochemical characteristics (Fig. 4a). The brownish, light-grey and dark-grey laminae reach the highest values in MS (maximum value of  $110 \times 10^{-5}$  S.I., Fig. 4a), density (maximum value of  $1.7 \text{ g cm}^{-3}$ , Fig. 4a), Ti content (maximum values of 15 000 peak area, Fig. 4a), coarse fraction (maximum of 60% for the massic class  $> 63 \mu\text{m}$ , Fig. 4a) and TOC/TN ratio (maximum of 19, Fig. 4a). Laminae are also systematically marked by a strong decrease in TOC (minimum value of 3%, Fig. 4a). XRD analyses conducted on these laminae showed a similar mineralogical composition composed of quartz, feldspaths and phyllosilicates, comparable to the background sedimentation. These physical and geochemical characteristics are similar to those observed within the laminae in core QAL-2011 (Fig. 3a), even if LKG-2011 laminae are larger ( $> 1 \text{ cm}$ , Table 2b) than in QAL-2011. Moreover, the observations of the thin-sections and the QOP analysis made on the brownish and dark-grey laminae from core LKG-2011, revealed the presence of an inversely graded bed followed by a normal one and the predominance of rAP and LCF in the OM content (Fig. 4b). On the basis of the

similar physical and geochemical properties of LKG-2011 laminae, these results were extrapolated to other laminae.

## 5 Discussion

### 5.1 Sedimentary events (SEs)

5 The white laminae observed within core QAL-2011, as well as the brownish, light-grey and dark-grey laminae within LKG-2011, showed similar sedimentological properties, clearly contrasting with the background sedimentation except for the mineralogical composition. Generally, they are characterized by high values of MS and density, higher Ti content and a rise in the massic class over 63  $\mu\text{m}$  (Figs. 3a–4a). All these  
10 properties are characteristic of terrestrial inputs (Chapron et al., 2005; Arnaud et al., 2012). Moreover, the laminae composition from core QAL-2011 is also characterized by normal grading, high TOC/TN ratio, a predominance of rAP and LCF particles produced in soils (Fig. 3b; Di-Giovanni et al., 1998) and a mineralogical composition of granite, both resulting from the reworking of detrital inputs from the catchment area in  
15 lake sediments (Simonneau et al., 2013a, b; Foucher et al., 2014). This organic signature therefore confirms the terrestrial origin of these laminae. Associated with the fining upward, both organic and granulometric results from core QAL-2011 are typical of hyperpycnal flood deposits (Mulder and Alexander, 2001; Simonneau et al., 2013a). The fining upward sequence can therefore be correlated to the falling limb of the flood hydrograph (Simonneau et al., 2013a). However, because Lake Qallimiut is the last  
20 sediment trap of the catchment with many lakes upstream, flood event records are thin and do not show the coarsening upward of the deposit, characteristic of the rising limb of the flood hydrograph (St-Onge et al., 2004). This suggests that the 27 white laminae described within core QAL-2011 can be related to exceptional hyperpycnal flood events  
25 associated with heavy rainfall and/or snowmelt events occurring all over the drainage basin and essentially reworking terrestrial materials.

Within core LKG-2011, thin-section and QOP analyses of the dark-grey and brownish laminae showed that the deposits are successively inversely and normally graded grain size, mainly constituted of rAP and LCF organic particles and composed of minerals present in granites (Fig. 4b). All these characteristics indicate that, as for core QAL-2011, the 10 laminae described within core LKG-2011 (Table 2b) also correspond to hyperpycnal floods deposits too. Because there is only one lake upstream Lake Little Kangerluluup, and because of the elongated shape and the high gradient slope of its catchment with increases detrital inputs, the LKG-2011 flood event deposits are thicker and coarser than those identified within core QAL-2011.

## 5.2 Flood frequencies during the Late Holocene

Flood event records from both sites were smoothed over a 250 year moving window (K1D software, Gavin et al., 2006). QAL-2011 and LKG-2011 flood frequencies are globally synchronous during the last four millennia, considering age-depth models uncertainties (Fig. 5).

In more detail, a total of five high flood frequency periods were identified (Fig. 5; Table 2a and b):

- between ca. 2300 BC and ca. 1600 BC;
- between ca. 1100 BC and ca. 900 BC;
- between ca. 600 BC and ca. 200 BC;
- between ca. AD 1050 and ca. AD 1200;
- between ca. AD 1350 and ca. AD 1650.

These episodes of increased flood events are synchronous with low Greenlandic lake temperatures (correlated to the air temperatures; D'Andrea et al., 2011), Arctic glacier advances (Kelly et al., 2008; Briner et al., 2009, 2011; Larsen et al., 2011, 2012; Winsor et al., 2014) and North Atlantic ice-rafting events (Bond et al., 2001; Fig. 5). Thus, over



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the second half of the Holocene, QAL-2011 and LKG-2011 flood deposits seemed to occur especially during cooler and wetter periods, associated to increased summer ice and snow-melting episodes. During colder and wetter periods, intense rainfall episodes created a sufficient stock of snow and ice. After short warming episodes in spring and summer (such as several accumulated positive degree days or hot spell melt events caused by foehn winds), ice and snow melting events increased river discharge, sometimes inducing sporadic flood events (Mernild and Hasholt, 2009). In glacial areas, ice-melting episodes might also be sometimes associated to the rupture of a glacier ice dam (Mayer and Schuler, 2005; Carrivick et al., 2013). In addition, high flood frequency also appeared synchronously with weak solar irradiance phases (Bond et al., 1997, 2001; Steinhilber et al., 2009), probably suggesting a solar forcing for these events.

Four of the five Greenlandic flood-event phases were already defined at a global scale as rapid climate change periods (RCC) by Mayewski et al. (2004), separated by ca. 1500 year cycles and likely driven by the solar activity too (Bond et al., 1997, 2001; Mayewski et al., 1997, 2004). For example, the period of high flood frequency between ca. 2300 BC and ca. 1600 BC may correspond to the Middle to Late Holocene transition (ca. 2250 BC), defined as a cooler period by Walker et al. (2012). Then, the period between ca. 1100 BC and ca. 900 BC, as well as the one between 600 BC and 200 BC, could match to the bipartite cooler Sub-boreal/Sub-atlantic transition (Van Geel et al., 1996; Martin-Puertas et al., 2012; Magny, 2013). Surprisingly, no flood events were recorded during the RCC corresponding to the Dark-Ages (DA) cooling (between ca. AD 200 and ca. AD 900; Bianchi and McCave, 1999; McDermott et al., 2001; Sicre et al., 2008). This could be explained by the fact that the DA was a drier climatic episode and/or it did not really impact the study area, especially the hydrological systems of Lake Qallimiut and Little Kangerluluup. Finally, the period between ca. AD 1000 and ca. AD 1650 revealed a maximum of flood deposits over the last four millennia, partly corresponding to the cooling of the Little Ice Age (LIA). With a maximum of flood events, the LIA cooling, occurring between AD 1300 and AD 1900, can be defined as the

main episode of climatic pejouration recorded in south-west Greenland over the Late Holocene.

### 5.3 High-resolution flood frequency during the last 1000 years

During the last millennium, flood events are especially recorded between ca. AD 1050 and AD 1200, and between AD 1350 and AD 1650 (Fig. 6).

The time interval between ca. AD 1350 and AD 1650 corresponds to the climatic pejouration of the LIA. A major decrease in Arctic temperatures (Sicre et al., 2008; Andrea et al., 2011; McKay and Kaufman, 2014; Millet et al., 2014; Figs. 5–6), a maximum of Arctic glacier advances (Kelly et al., 2008; Briner et al., 2009, 2011; Larsen et al., 2011, 2012; Winsor et al., 2014; Fig. 6), a maximum rate of ice accumulation (Andersen et al., 2006; Fig. 6) and a solar minimum (Spörer minimum; Bard et al., 2000; Fig. 6) were identified during this period, confirming that these flood events likely resulted from ice and snow-melting episodes and were potentially driven by solar forcing.

The highest flood frequency occurred between ca. AD 1050 and ca. AD 1200 during the Oort solar minimum (Bard et al., 2000) and the Norse settlement (Arneborg et al., 2012; Fig. 6). It was the only time over four millennia that flood events took place during a warmer and drier period: the Medieval Warm Period (MWP; Andersen et al., 2006; McKay and Kaufman, 2014; Millet et al., 2014; Fig. 6). During this phase, flood deposits are thick (Fig. 6), attesting to the increase of detrital inputs probably induced by Norse farming activities in the lake watersheds. Indeed, around lakes Qallimiut and Little Kangerluluup, archaeological sites attested to agropastoral practices during the Norse settlement (Madsen et al., 2009; Madsen, 2014). They are likely responsible for enhancing the detrital inputs due to the reduction in the vegetation cover and to live-stock trampling (Gauthier et al., 2010; Massa et al., 2012a), making the remobilization of soil material easier during flood events and less intense hydrological events. Indeed a similar sedimentological response was observed, regardless of the intensity of the hydrological events, because of soil remobilization due to agropastoral activities in the catchment. The highest flood deposit thickness was recorded in Lake Little Kangerlu-

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luup (Fig. 6), potentially explained by a larger number of archaeological sites identified around this lake (Madsen et al., 2009).

In south-west Greenland, there were two main agropastoral phases: the Norse settlement and the last century with higher impacts on ecosystems in recent decades (Massa et al., 2012a). Thus, we can expect enhanced flood deposits in the last decades. However, thicker flood deposits occurred before the 1930s when the anthropogenic impacts were low (Massa et al., 2012a). It should be pointed out, though, that Lake Qallimiut and Little Kangerluluup are not proper lakes to trace recent agropastoral impacts since around Lake Qallimiut, few herds are grazing and the recent farm is located at its outlet probably rejecting its effluents directly into the fjord. Concerning Lake Little Kangerluluup, there has been no human occupation or pastoral activities in the catchment during the last century (Madsen et al., 2009).

#### 5.4 Climatic influences on Human settlements

A compilation of ca. 200 calibrated radiocarbon dates ( $2\sigma$  ranges) from south Greenland Norse sites was made (Arneborg, 2007; Edvardsson et al., 2007; Schofield et al., 2007; Arneborg et al., 2012; Walker et al., 2012; Ledger et al., 2013, 2014) to illustrate the patterns of occupation across time and especially its relative magnitude (Rick, 1987; Fig. 6). For example, a maximum of radiocarbon dates could reveal a period of maximum occupation (Kuzmin and Keates, 2005; Rick, 1987; Kuper and Kröppelin, 2006). Here, two periods of maximum occupation were therefore inventoried between AD 1000-AD 1100 and between AD 1250-AD 1350 (Fig. 6). From AD 850, the radiocarbon dates increased progressively until ca. AD 1050 because of the establishment of the settlement since AD 986. After ca. AD 1050, synchronously with a drop in radiocarbon dates, flood frequency rose, potentially indicating some short cooler and wetter episodes during the MWP in spite of a global warm and dry trend. These periods are associated to a short demise of the Norse, recorded by a drop of radiocarbon dates, probably linked to these climatic perturbations. Radiocarbon age frequency rose again after ca. AD 1200 when flood events ceased, indicating repopulation by the Norse or

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a renewal of their activities, possibly linked to the improvement in the climate. Around AD 1300, the maximum frequency of radiocarbon dates was synchronous with the peak of the estimated Norse population (Madsen, 2014). Radiocarbon dates fell to zero in AD 1450, with the progressive and definitive demise of the Norse settlement in Greenland, approximately dating from the mid to late 15th century (Dugmore et al., 2009). This abandon is still a controversial issue (Barlow et al., 1997; Dugmore et al., 2012). However, it is generally assumed that the climatic deterioration played an important role in their abandon (D'Andrea et al., 2011; Massa et al., 2012a). Here, the radiocarbon dates decreased around AD 1300, synchronous with the beginning of the cooling of the LIA and the second major rise in flood frequencies, supporting the idea of a progressive Norse demise mainly linked to the climatic deterioration of the LIA.

Global cooler and wetter episodes, revealed by the flood frequency within core QAL-2011 and LKG-2011, revealed a potential impact on the Norse settlement. However, they were not the only society present in south Greenland during the last four millennia. The paleo-Eskimo culture named the Saqqaq, established from ca. 2500 BC to 850 BC, experienced cold and warm episodes during their settlement (D'Andrea et al., 2011; Fig. 5), showing a global adaptation to climate changes. For example, technological tool changes and inventions were observed around 1450 BC (Møbjerg, 1999), possibly revealing an adaptation to climate change (Sørensen, 2010), as this period corresponds to a solar minima phase (Steinhilber et al., 2009), characterized by cold temperatures (D'Andrea et al., 2011), an increase in North-Atlantic ice-rafted debris (Bond et al., 2001) and high flood frequencies (Fig. 5), potentially related to the beginning of the Sub-boreal/Sub-atlantic transition. Globally, the Saqqaq were defined as open water hunters (Meldgaard, 2004) in response to the relatively warm temperature trend occurring during their settlement (Moros et al., 2006; D'Andrea et al., 2011; Fig. 5). Thus, when the greatest decrease in temperatures (D'Andrea et al., 2011), increase in North-Atlantic ice-rafted debris (Bond et al., 2001) and south Greenlandic glacier advances (Kelly et al., 2008; Larsen et al., 2011, 2012; Winsor et al., 2014; Fig. 5) occurred at the end of the Sub-boreal/Sub-atlantic transition around 850 BC,

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they disappeared abruptly. Contemporary to their abandon, the Greenlandic Dorset, defined as sea-ice hunters because of their equipment (sledge shoes, soapstone vessels for burning seal fat and snow knives; Jensen, 2009) settled in Greenland. At the beginning of the Roman climate warming (RCW) when sea-ice cover diminished (Moros et al., 2006; Sicre et al., 2008; D'Andrea et al., 2011), the Dorset, who were ill-adjusted to warm conditions, left Greenland around 50 BC (Jensen, 2009). Finally, the Thule culture, present from the 12th century in Greenland, did not seem to be especially affected by the LIA. They just migrated in response to this cooler event and did not modify their diet or their tools, allowing them a long-term persistence in Greenland (Sørensen, 2010).

Thus, the RCC, occurring during the second half of the Holocene and recorded through QAL-2011 and LKG-2011 flood deposits, seem to have influenced Greenlandic Human occupation since the first colonization in ca. 2500 BC.

## 6 Conclusion

To reconstruct past Holocene climatic variations in south-west Greenland, a multiproxy study was performed on two lacustrine cores retrieved from Lake Qallimiut and Lake Little Kangerluluup, covering the last four millennia (QAL-2011 and LKG-2011). This study, including physical (magnetic susceptibility, density and grain size) and geochemical (X-ray microfluorescence, elemental analyses and organic petrography) analysis, has evidenced a total of 37 flood events in both sequences. They fall into five distinct phases: between 2300 BC and 1600 BC, 1100 BC–900 BC; 600 BC–200 BC; AD 1050–AD 1200; AD 1350–AD 1900 synchronous with high negative temperature anomalies, global glacier advances and an increase North-Atlantic in ice-rafted debris. High flood frequencies probably resulting from ice and snow-melting, coincide with solar minima and could be matched with global RCC such as the Middle to Late Holocene transition (ca. 2250 BC), the Sub-boreal/Sub-atlantic transition (ca. 700 BC) and the LIA (from ca. AD 1300 to ca. AD 1900). The RCC with a maximum of flood events

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recorded was the LIA, suggesting major climate deterioration in south Greenland at this time. Climatic deterioration revealed by QAL-2011 and LKG-2011 flood deposits may have potentially impacted local Human societies. Indeed, the Sub-boreal/Sub-atlantic transition might be responsible for the demise of the Saqqaq and the colonization of the Dorset. Moreover, the climatic deterioration of the LIA, marked by two major successive rises in flood frequencies, seems to be one of the main reasons for the Norse abandon, illustrated by the radiocarbon dates frequency measured in archaeological sites. Lakes Qallimiut and Little Kangerluluup are proper sites to reconstruct past climate changes during the last four millennia and their impact on past Human societies except during the Middle Ages, when Norse agropastoral activities enhanced the detrital inputs and disturbed their sedimentological responses.

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**Table 1.** Radiocarbon dates made on terrestrial remains and the age calibrations obtained from a linear model, using R program and the calibration curve of IntCal13 (Reimer, 2013). (a) Radiocarbon dates on Lake Qallimiut core. (b) Radiocarbon dates on Lake Little Kangerluluup core.

(a) Qallimiut	Depth cm	Material	Radiocarbon ages AD/BC	Calibrated ages AD/BC	Lab code
	19.5	wood	1550 ± 30	1480 ± 85	Beta-395 567
	51.25	wood	240 ± 30	355 ± 85	Beta-363 187
	71.1	wood	−530 ± 30	−630 ± 290	Lyon-9888
	88.4	wood	−1150 ± 30	−1355 ± 145	Lyon-9889
	93.25	wood	−1425 ± 30	−1660 ± 90	Lyon-9890
	95.75	wood	−1480 ± 30	−1720 ± 120	Lyon-9891
(b) Little Kangerluluup	Depth cm	Material	Radiocarbon ages AD/BC	Age AD calibr� AD/BC	Lab code
	20.55	wood	1705 ± 30	1655 ± 50	Lyon-9885
	31.9	wood	1445 ± 30	1420 ± 50	Lyon-9882
	34.5	wood	1345 ± 30	1335 ± 75	Lyon-9886
	52.65	wood	735 ± 30	825 ± 125	Lyon-9887
	77.75	wood	−165 ± 30	−125 ± 155	Lyon-9883
	93.75	wood	−940 ± 30	−1060 ± 140	Beta-386 534

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**Table 2.** Depth, thickness and age of the flood deposits. (a) Lake Qallimiut flood deposits description. (b) Lake Little Kangerluluup flood deposits description. NA indicates that the age error is not available because the end of the age-depth models is not constrained with a radiocarbon date.

(a) Qallimiut	Top cm	Bottom cm	Thickness cm	Age AD/BC	Age error
	1.4	1.9	0.5	1965	35
	3.2	5.7	2.5	1904	150
	7.5	8	0.5	1835	130
	10.2	10.95	0.75	1755	125
	15	15.5	0.5	1605	145
	16.6	17	0.4	1565	155
	17.5	17.75	0.25	1545	160
	18.5	19.1	0.6	1515	170
	21.8	21.95	0.15	1400	165
	22.2	22.35	0.15	1390	160
	23.3	23.9	0.6	1350	155
	27.4	28	0.6	1195	150
	29.5	29.6	0.1	1130	145
	29.9	30.1	0.2	1125	145
	30.2	30.4	0.2	1115	140
	30.6	30.85	0.25	1110	140
	31.8	32.05	0.25	1065	140
	32.2	34	1.8	1055	140
	36.7	37.2	0.5	940	135
	66.8	67.8	1	-465	255
	77.2	77.7	0.5	-905	200
	81	81.5	0.5	-1050	160
	100.9	101.3	0.4	-1885	NA
	110.4	110.6	0.2	-2145	NA
	113.4	113.7	0.3	-2225	NA
	113.9	114.2	0.3	-2230	NA
	117	117.5	0.5	-2310	NA
(b) Little Kangerluluup	Top cm	Bottom cm	Thickness cm	Age AD/BC	Age error
	4	6	2	1930	210
	17	19	2	1725	370
	29.2	30.7	1.5	1455	115
	41.3	44.1	2.8	1083	120
	49.2	51	1.8	880	160
	79.7	81.1	1.4	-275	235
	81.6	82.7	1.1	-310	225
	94.8	98.5	3.7	-1140	NA
	104.9	105.8	0.9	-1580	NA
	107.7	109	1.3	-1715	NA

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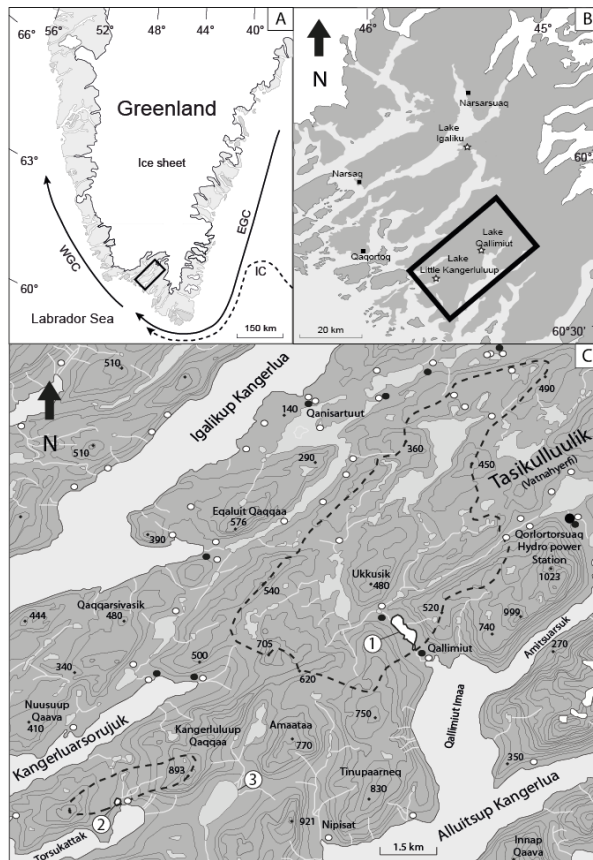
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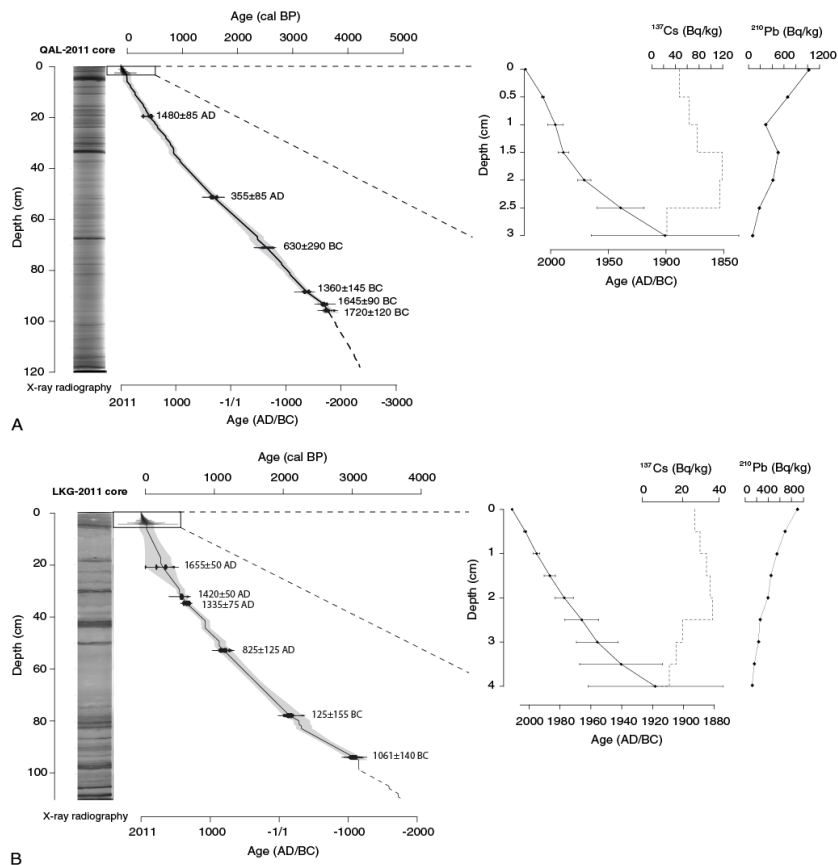
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**Figure 1.** Localization of Lake Qallimiut and Little Kangerluluup. **(a)** Map of Greenland. **(b)** Map of South Greenland. **(c)** Catchments outlines (dotted black lines). White points represent Norse archeological sites and the black ones, recent farms. Number 1 indicates the location of Lake Qallimiut, number 2 matches Lake Little Kangerluluup and number 3, Lake Kangerluluup.

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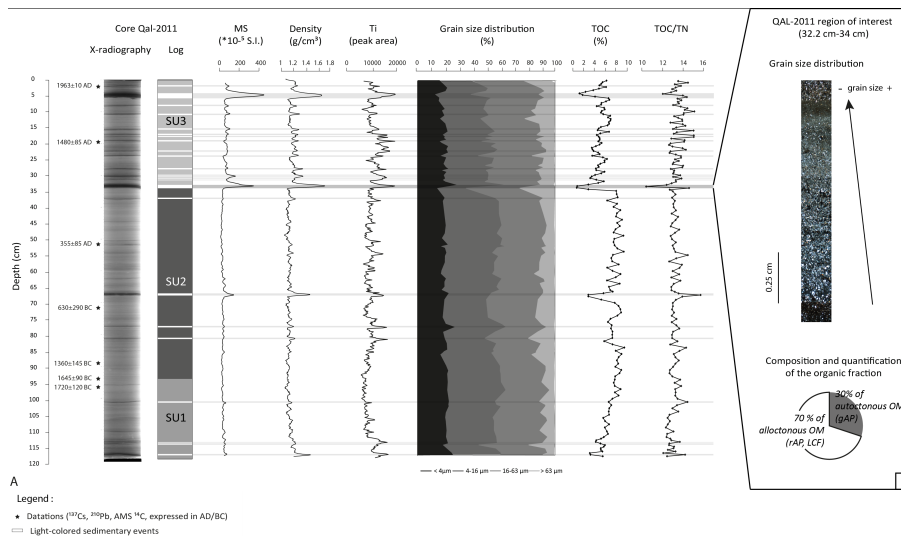
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**Figure 2.** Age-depth models based on radiocarbon dates and on short-lived radio-isotopes measurements. **(a)** Lake Qallimiut age-depth model. **(b)** Lake Little Kangerluluup age-depth model.

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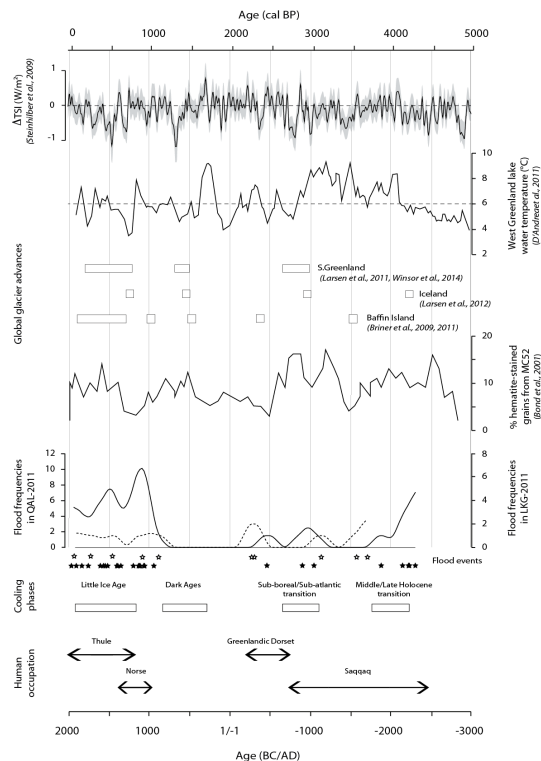


**Figure 3.** (a) X-radiographs and log representing the Lake Qallimiut core associated to magnetic susceptibility,  $\gamma$ -density, XRF, grain size, TOC and TOC/TN ratio measurements. Grey zones highlight the physical and geochemical responses of the white laminae. (b) The frame contains photographs of the thin section in polarized light and the piechart shows the composition of the OM.



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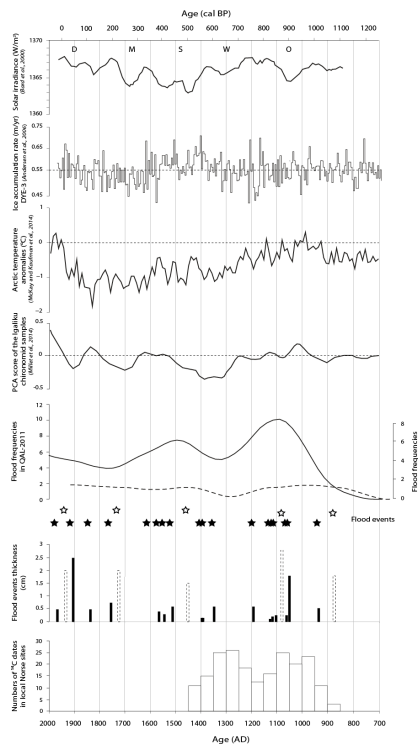


**Figure 5.** Flood frequencies expressed in a time step of 250 years in the last four millennia. In black, flood frequency measured in Qallimiut sequence and in dotted line, flood frequency observed in Little Kangerluluup core. They are compared to solar irradiance (Steinhilber et al., 2009), west Greenland lake water temperatures (D’Andrea et al., 2011), global glacier advances (Briner et al., 2011, 2009; Kelly et al., 2008; Larsen et al., 2011; Winsor et al., 2014), the main rapid climate changes (Mayewski et al., 2004) and periods of Human occupations (Dugmore et al., 2009; Gulløv, 2004; Jones, 1986; Meldgaard, 2004).



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**Figure 6.** Focus on Qallimiut and Little Kangerluluup flood frequencies in the last millennium. They are associated to their thickness (in black, Qallimiut flood deposit thickness and in dotted line, Little Kangerluluup flood events thickness). They are also compared to solar irradiance (Bard et al., 2000), ice accumulation rate (Andersen et al., 2006), Arctic temperature anomalies (McKay and Kaufman, 2014), PCA score of the Igaliku chironomid samples (Millet et al., 2014) and the number of radiocarbon dates measured in local archaeological sites (Arneborg, 2007; Edvardsson et al., 2007; Madsen et al., 2009; Arneborg et al., 2012).