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2 **Archaeointensity study of five Late Bronze Age fireplaces from Corent (Auvergne,**
3 **France)**

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13

14

15 **Abstract**

16 Recent excavations at Coirent (France) unearthed a vast Late Bronze Age settlement. The
17 high density of fireplaces especially highlights it. The present study focuses on the
18 archaeomagnetic study of five fireplaces. These ones were dated between 950 and 800
19 BC by cross-dating of metallic and ceramic artefacts and by radiocarbon. The main
20 objective of our study is to increase the archaeointensity database in Western Europe at
21 the beginning of the first millennium BC. The sampling was conducted on 64 fragments
22 of baked clay and sherds from the fireplaces floor. The classical Thellier-Thellier
23 protocol provides 48 successful archaeointensity results, yielding to five mean values
24 between 58 and 69 μT at the site. Together with previously published results, our new
25 data point out two successive maxima of the intensity of the geomagnetic field. The first
26 maximum $\sim 70 \mu\text{T}$ in the ninth century BC and the second $\sim 90 \mu\text{T}$ in ~ 700 BC are
27 separated by a $\sim 45\text{-}50 \mu\text{T}$ minimum at $\sim 800\text{-}750$ BC. The resulting fast variation of the
28 field intensity will be very useful for archaeomagnetic dating purposes. As the direction
29 of the geomagnetic field has also a strong variation during this period (Hervé *et al.*,
30 2013a), archaeomagnetism promises to be a powerful dating tool to recover the
31 historical processes at the transition between the Bronze and Iron Ages in Western
32 Europe.

33

34 **Keywords**

35 archaeomagnetism; archaeointensity; France; Late Bronze Age; Puy de Coirent

36

37 **1. Introduction**

38 The number of archaeomagnetic intensity results considerably grew in Western Europe
39 during the last few years (e.g. Genevey et al., 2009, 2013; Gómez-Paccard et al., 2008,
40 2012; Hervé et al., 2013b; Schnepf et al., 2009; Tema et al., 2013). Most of them cover
41 the past 2500 years and only few have been published for older periods (Aidona et al.,
42 2006; Gallet et al., 2009; Hervé et al., 2011; Hill et al., 2008; Kapper et al., 2015;
43 Kovacheva et al., 2009). The latter highlight a fast secular variation in intensity,
44 especially between 1000 and 500 BC that is at the Late Bronze Age and the Early Iron
45 Age. This fast changing of intensity was also recovered in the Middle East (e.g. Ertepinar
46 et al., 2012; Gallet and Le Goff, 2006; Gallet et al., 2015; Kovacheva et al., 2014; Shaar et
47 al., 2011). A better constraint of the secular variation during this period in Western
48 Europe will allow to better understand the geomagnetic field behaviour at the regional
49 and global scale (Hong et al., 2013).

50 By the other hand, this fast secular variation lets also expect a great potential for the
51 archaeomagnetic dating technique. A directional (inclination and declination) curve is
52 already available for Western Europe (Hervé et al., 2013a). However, Western Europe
53 intensity data for Late Bronze and Early Iron Age are still too few to build a precise and
54 accurate regional secular variation curve. Adding the intensity to the direction will
55 provide a more efficient chronological tool for archaeologists. The five new data from
56 the Late Bronze Age settlement of Corent presented in this study are a new step to
57 better recover the intensity secular variation in Western Europe and to improve the
58 dating method for this period.

59

60 2. Archaeological context

61 The Puy de Corent is located on a plateau overlooking the Grande Limagne plain, 19 km
62 away from Clermont-Ferrand in Auvergne (Latitude: 45.665°N; Longitude: 3.189°E).
63 Since 2001, two teams of researchers have excavated this site, one from Université
64 Lumière Lyon II conducted by Matthieu Poux and another one from Université Toulouse
65 – Jean Jaurès conducted by Pierre-Yves Milcent. This location is very famous for its
66 *oppidum* of the Late La Tène period, but is also characterized by earlier important
67 agglomerations (Milcent *et al.*, 2014a and 2014b).

68 One of these important occupations of Puy de Corent's is dated at the end of the Bronze
69 Age (from the end of the 11th to the end of 9th century BC), during which a vast and
70 dense settlement developed on the lower part of the plateau and covered a minimum
71 surface of 15 ha (Figure 1a). Its limits have not yet been reached and we have now some
72 evidence that the site could be one of the first proto-urban settlements in Western
73 Europe (Ledger *et al.*, 2015). Three successive phases of occupation and development of
74 the agglomeration were recognized: "Bronze Final 2 récent" (~1050 – ~950 BC),
75 "Bronze Final 3 ancien" (~950 – ~900 BC) and "Bronze Final 3 récent" (~900 – ~800
76 BC). These phases are determined by stratigraphy. They are dated by few radiocarbon
77 dates and by comparison of the abundant ceramics and metallic artefacts with similar
78 objects coming from accurately dated alpine lake's palafittes. The various occupation
79 levels display a high density of fireplaces, also dated by their relative positions in the
80 stratigraphic sequence and according to their close relationships with the ceramic and
81 metallic material. Radiocarbon dating (Lyon-11289, 2785±35 BP) of the occupancy level
82 numbered [20450] related to the fireplace FY20462, assigned to "Bronze Final 3 ancien"
83 by ceramics, confirms the archaeological dating ([950; 900] BC) with the dating interval

84 [1012; 839] BC at 95 per cent of confidence and [979; 899] BC at 68 per cent of
85 confidence.

86 The 64 fireplaces of the Late Bronze Age (1 per 50 m² in average) discovered since 2001
87 display, whatever their phase, some recurrent features in their shape and their
88 construction type. Most of the time, fireplaces are built on simple or mixed raft
89 foundation of small pebbles, basalt blocks, re-used fragments of stone macro-equipment
90 (grindstones and granite thumb-wheels) or ceramic sherds (Figure 1b-c). They support
91 screeds with thickness generally varying between 1 and 3 cm made of mixed clay and
92 sand. The best-preserved fireplaces are either circular or rectangular with round angles,
93 and measure between 1.00 and 1.60 m of diameter for the first ones, and 0.95 m x 0.70
94 m for the latter. Repeatedly, we observed, under the raft foundation and in the center of
95 the fireplace, a little *locus* with a depth of 5 to 10 cm and with a diameter varying from
96 18 to 28 cm. Although they seemed sealed by the fireplace, some *loci* sheltered another
97 deliberate deposition. The deposition are composed of bone, bronze objects (pin, ring,
98 metal droplet) and even exceptionally a fig seed, whose the growing was limited to the
99 Mediterranean regions at the Late Bronze Age. Positioning exactly the fireplaces in
100 relation to the constructions on standing posts of the site remains difficult: while some
101 were clearly inside the buildings, others seem to have been outside.

102

103 **3. Archaeomagnetic analyses**

104 3.1 Sampling

105 We sampled five fireplaces. The best-preserved fireplaces FY20462 and FY22783
106 (Figure 1b) were sampled *in-situ* using plaster cap method. Respectively 9 and 13 blocks

107 of baked clay were surrounded with plaster, levelled horizontally using a bubble and
108 oriented using a magnetic compass. In the laboratory of Rennes, the baked clay
109 fragments were prepared in 8 cm³ cubic specimen after consolidation using sodium
110 silicate. In the case of the disturbed fireplaces FY22705, FY22798 and FY22842 (Figure
111 1c), we collected without orientation between 12 and 16 baked clay fragments and
112 pottery sherds per structure. Those and the pottery sherds of the fireplaces FY20462
113 and FY22783 were divided in ~1 cm³ chips with the same orientation. The cutting
114 reference was a flat side of the fragment or the sherd, which corresponded or was
115 parallel to the surface of the fireplace. This would help to identify the component of
116 remanent magnetization acquired *in situ*. Each chip was then packed into cylindrical
117 quartz holder filled by quartz wool.

118

119 3.2 Rock magnetism

120 To investigate the ferromagnetic mineralogy, thermomagnetic curves were measured on
121 small chips of 34 samples using a KLY3-CS3 susceptibility meter with a fitted furnace.
122 The variation of the susceptibility was measured during heating to 400 and 600°C and
123 during the subsequent cooling. In all baked clay fragments and pottery sherds,
124 thermomagnetic curves reveal a dominant ferromagnetic phase with Curie
125 temperatures between 550 and 580°C identified as titanium-poor titanomagnetite
126 (Figure 2a-b). All heating-cooling cycles (up to 600°C) of pottery sherds are reversible
127 (Figure 2b). On some fragments of baked clay the slight irreversibility suggests
128 mineralogical evolutions at high temperature (Figure 2a). None samples were
129 nevertheless rejected for archaeointensity experiments, because all cycles up to 400°C
130 were fully reversible. Isothermal remanent magnetization (IRM) acquisition curves were

131 acquired on 21 specimens using an ASC impulse magnetizer. Saturation occurred at low
132 magnetic fields (~300 mT), indicating the lack of any high-coercivity ferromagnetic
133 phase (Figure 2c).

134

135 3.3 Thermal demagnetization

136 Prior to the archaeointensity experiment, one specimen per sample was thermally
137 demagnetized in a Magnetic Measurement Thermal Demagnetizer (MMTD) oven, in
138 order to identify the component of thermoremanent magnetization (TRM) acquired *in*
139 *situ*. Sample's positions in the field suggest that the expected TRM should have an
140 inclination of circa $\pm 60-70^\circ$, as based on the data from the neighbour and
141 contemporaneous site of Lignat (Gallet et al., 2002; Moutmir, 1995).

142 Almost all (50/54) fragments of backed clay from the upper layer of the fireplaces carry
143 a single TRM component with the expected inclination. Three pottery sherds carry two
144 clear components of magnetization. The low-temperature component between 100 and
145 400-500°C has a $\sim 60-70^\circ$ inclination and was therefore acquired *in situ*. Directions of
146 the high-temperature component are totally dispersed and they are probably associated
147 to the initial firing of the pottery. As these three sherds were found below the baked clay
148 layer, they reached a lower temperature during the last heating of the fireplace and
149 therefore they carry two components of magnetization. Six pottery sherds (three from
150 FY22798 and three from FY22842) and four baked clay samples (from FY 22798) have a
151 multiple component magnetization, none of them close to the expected inclination. We
152 did not perform archaeointensity experiment on these ten samples.

153 Oriented block samples from FY20462 and FY22783 carry a TRM with an easterly
154 declination and an inclination in the range of the expected values for the Late Bronze
155 Age (Hervé *et al.*, 2013a) (Figure 3). However the scatter between directions indicates
156 slight displacements of the baked clay fragments since the last high-temperature heating
157 and prevents the calculation of a mean direction of magnetization.

158

159 3.4 Archaeointensity study

160 Archaeointensity experiments were performed using the classical Thellier-Thellier
161 method (Thellier and Thellier, 1959) with partial thermoremanent magnetization
162 (pTRM) checks on 56 specimens (50 baked clay fragments and 6 pottery sherds). At
163 each temperature step, specimens were heated and cooled twice, first in a laboratory
164 field $+F_{lab}$ and secondly in the opposite field $-F_{lab}$ of 60 μ T. The protocol was performed
165 using 13 temperature steps up to 570 or 580°C in a Pyrox amagnetic oven. All remanent
166 magnetization were measured with a 2G cryogenic magnetometer. The anisotropy of
167 TRM was determined at 510 or 540°C using 6 successive heating and one stability check
168 (Chauvin *et al.*, 2000). The cooling rate effect on TRM intensity was also corrected at 535
169 or 555°C using the procedure in four heating steps of Gómez-Paccard *et al.* (2006). The
170 slow cooling rate was fixed to 8 hours.

171 We used the following criteria to select the specimens for the mean archaeointensity
172 calculation: NRM fraction factor higher than 0.5, maximum angular deviation Mad lower
173 than 5°, deviation angle $Dang$ lower than 5° and ratio of the standard error of the slope
174 to the absolute value of the slope β lower than 0.05. A pTRM-check was said positive if
175 its difference with the original pTRM was lower than 10%. Samples showing a concave-

176 up NRM-TRM diagram indicating some mineralogical changes during heating were
177 rejected from the analysis (Figure 4a). All the accepted specimens have a linear NRM-
178 TRM diagram (Figure 4b). In the case of pottery sherds, the archaeointensity was
179 computed using the secondary component of magnetization after correction of the NRM-
180 TRM diagram (Hervé et al., 2013b, Figure 4c). The selection procedure yields an
181 acceptance rate of 86% (Supplementary material).

182 The archaeointensity values were corrected for TRM anisotropy if the alteration factor
183 inferred from the stability check was lower than 10%. The anisotropy degree varied
184 from 3 to 27%. The cooling rate correction was applied only when the absolute value of
185 the correction factor was higher than the alteration factor (Gómez-Paccard et al., 2006).
186 Otherwise, that is for ten specimens, the cooling rate correction was not accounted for.
187 The correction factor was usually lower than 6%, except for specimens from FY20462
188 with values between 10 and 15%. The mean archaeointensity per fireplace was
189 computed using the weighting method of Prévot et al. (1985) (Table 1). The five mean
190 values agree well with each other.

191

192 **4. Discussion**

193 Our new mean archaeointensities were relocated to Paris using the Virtual Axial Dipole
194 Moment (VADM) correction. On Figure 5, they are compared to published Western
195 European data for Late Bronze and Iron Ages. These data include the selected data set
196 described in Hervé et al., (2013b) completed by new Swiss data (in grey on Figure 5,
197 Kapper et al., 2015). Given the small number of published results, the five fireplaces of
198 Corent represent a significant step to recover the secular variation of the geomagnetic

199 field intensity between 1500 and 600 BC. The large range of archaeointensities (~50-
200 90 μ T) between 1000 and 600 BC points out high and fast variations of the geomagnetic
201 field strength.

202 An increase of the field intensity is observed during the tenth century BC up to ~65-70
203 μ T, followed by a possible decrease during the ninth century with values close to 45 μ T
204 at ~800-750 BC. After that, the intensity would increase up to ~90 μ T in ~700-600 BC, as
205 supported by Gallet et al. (2009), Hervé et al. (2011) and Hill et al. (2008) data.

206 From ~800 to ~700-600 BC, the secular variation rate would have been around ~3
207 μ T/decade. This rate is higher than the typical one (1 μ T/decade) observed over the last
208 two millennia in Western Europe (Genevey et al., 2013) but is similar to the rate during
209 the early Middle Age (Gómez-Paccard et al., 2012). The sharp secular variation at the
210 beginning of the Early Iron Age may be recorded in the Swiss mean data with an unusual
211 standard deviation of 15.5 μ T (Kapper et al., 2015, grey triangle data on Figure 5). This
212 average archaeointensity was computed from two pottery sherds coming from the same
213 archaeological layer. They provide archaeointensities of 45 and 75 μ T. We suggest that
214 these potteries were magnetized at slightly different times during a period of fast
215 changes of the field intensity.

216 The Corent data do not record the geomagnetic spikes (short-lived high field anomalies),
217 highlighted in the Levantine area during the ninth century BC (Ben-Yosef et al., 2009;
218 Shaar et al., 2011). Other new reference data are needed to investigate the presence of
219 such events in Western Europe and to better estimate the secular variation rate during
220 the Late Bronze Age.

221 Finally, our new data are compared to the prediction at Paris of the geomagnetic model
222 SHA.DIF.14k that is valid in the Northern hemisphere (Figure 5)(Pavón-Carrasco et al.,
223 2014a). This model is developed by inversion of archaeomagnetic and volcanic results
224 using spherical harmonic analysis in space and penalised cubic B-splines in time. For the
225 1500-500 BC time period the data set used to build the model includes a large amount of
226 results obtained on Eastern Europe sites. All the data in black on Figure 5 are also used
227 to build the SHA.DIF.14k model but not the more recently published Swiss data in grey
228 (Kapper et al., 2015). The SHA.DIF.14k model's prediction does not fit well most of the
229 data between 1000 and 500 BC. Maxima of intensity are observed but with shifts in time
230 and amplitude compared to Western Europe archaeointensity results. The
231 archaeointensity values obtained at Corent are higher than the model's prediction and
232 differ up to 15 μ T. Inhomogeneous quality of the archaeointensities and inhomogeneous
233 geographical distribution of sites in the global database probably explains these
234 inconsistencies (Pavón-Carrasco et al., 2014b). Finally this comparison indicates that
235 reliable intensity data are still needed in order to better constrain global models of the
236 past geomagnetic field.

237

238 **5. Conclusion**

239 The study of fireplaces from the archaeological site Puy de Corent provides five new
240 high quality archaeointensities at the Late Bronze Age in Western Europe. This
241 represents a new step to increase the amount of reliable data set and to build a
242 reference curve of the secular variation of the geomagnetic field intensity. The large and
243 fast secular variation at the Late Bronze Age and the Early Iron Age lets expect that this
244 reference curve will give precise archaeomagnetic dating both for in place and displaced

245 objects. Together with directional data of the geomagnetic field, which also shows large
246 variations, archaeomagnetic dating technique will provide a very valuable alternative to
247 radiocarbon, especially problematic during this period due to plateau effects.
248 Archaeomagnetism will efficiently contribute to the refinement of our knowledge of the
249 evolutions of the societies in Western Europe at the transition from the Bronze Age to
250 the Iron Age.

251

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257

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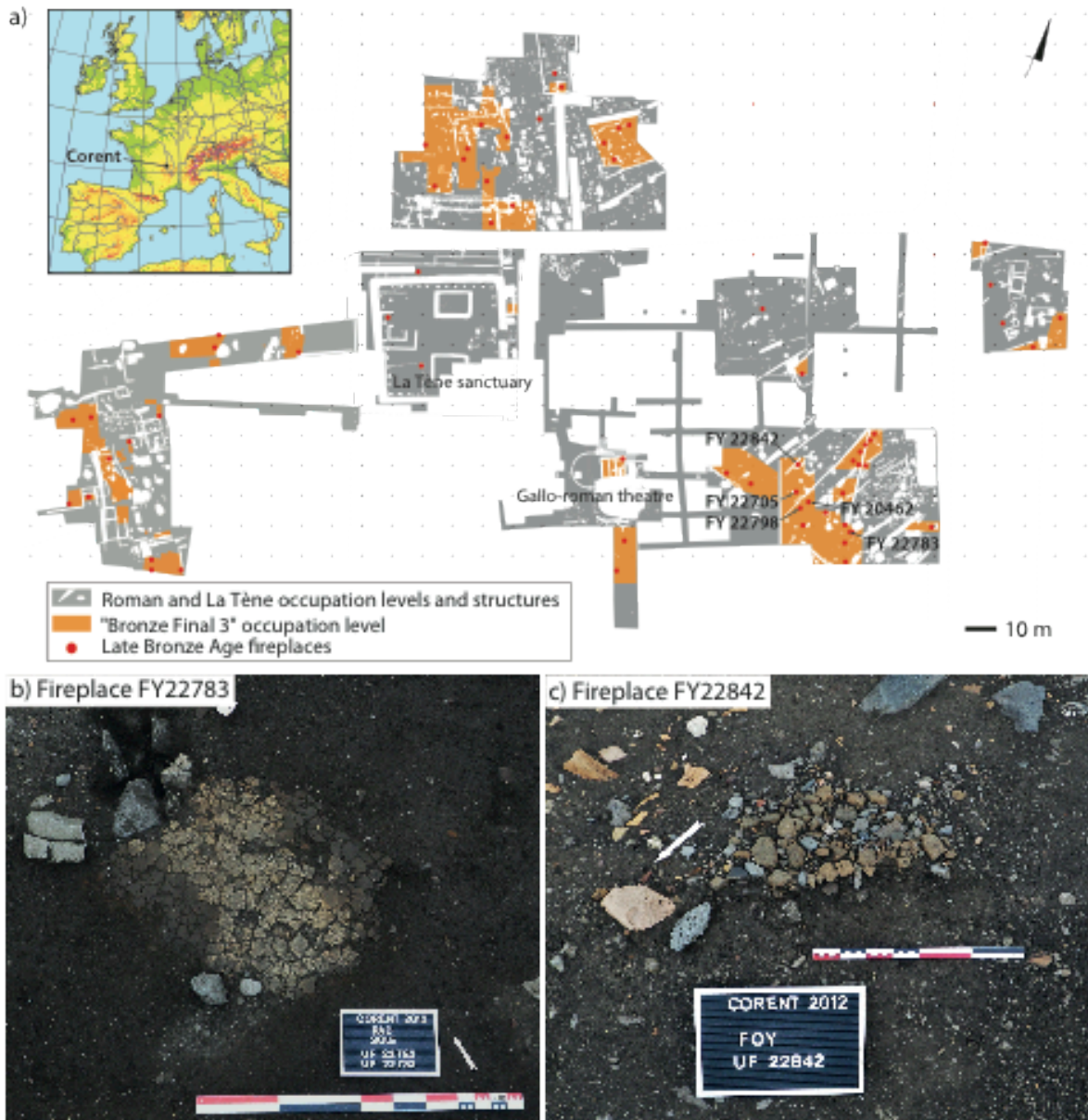
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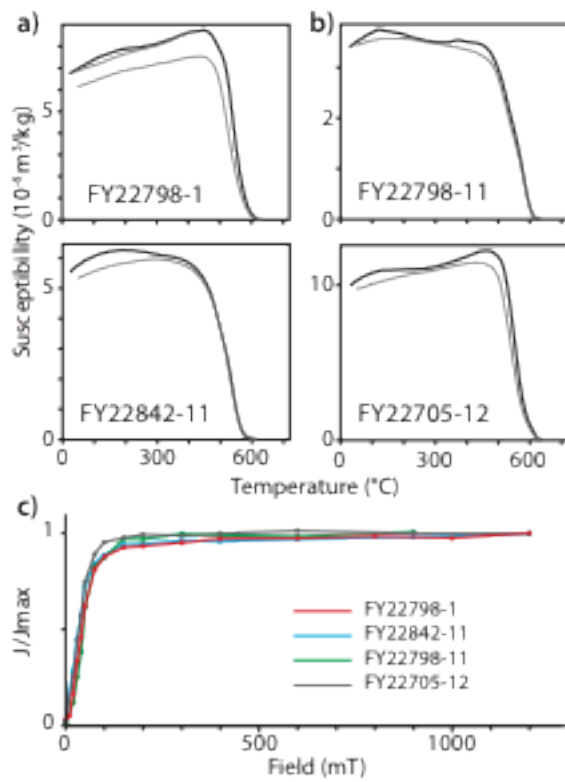
364 Figure 1: Map of the central area of Corent archaeological site (2001 to 2015
365 excavations) emphasizing the levels from the Late Bronze Age ("Bronze Final 3") and
366 their associated fireplaces (a) and pictures of two sampled fireplaces (b-c).



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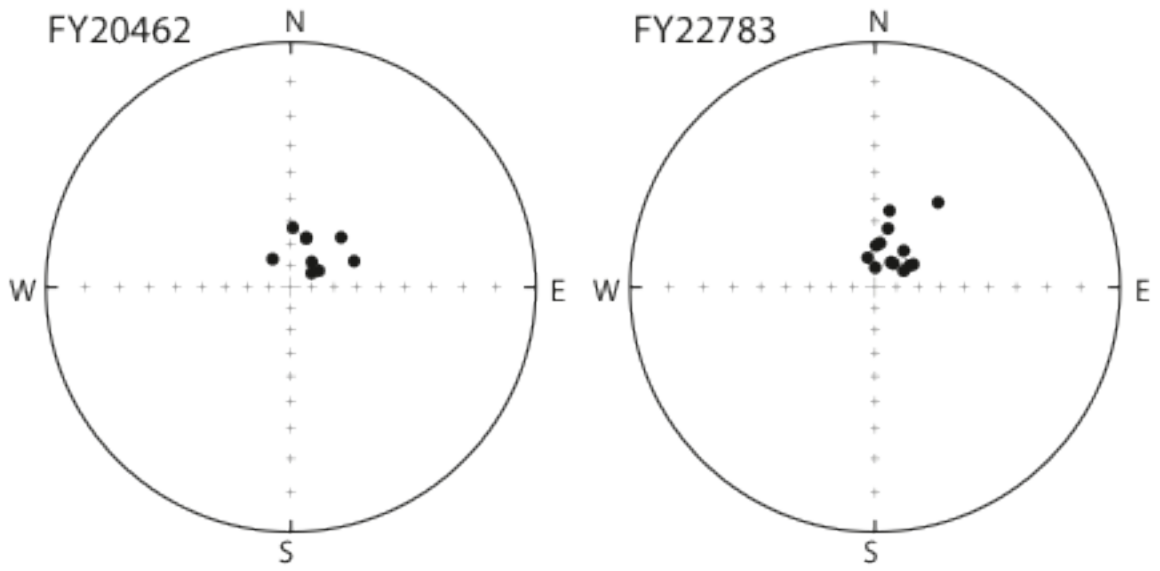
369 Figure 2: Representative magnetic mineralogy results with thermomagnetic curves of
370 baked clay fragments (a) and of pottery sherds (b) together with acquisition curves of
371 isothermal remanent magnetization (c). In thermomagnetic curves, the black curve is
372 the variation of susceptibility during the heating and the grey curve during the cooling.



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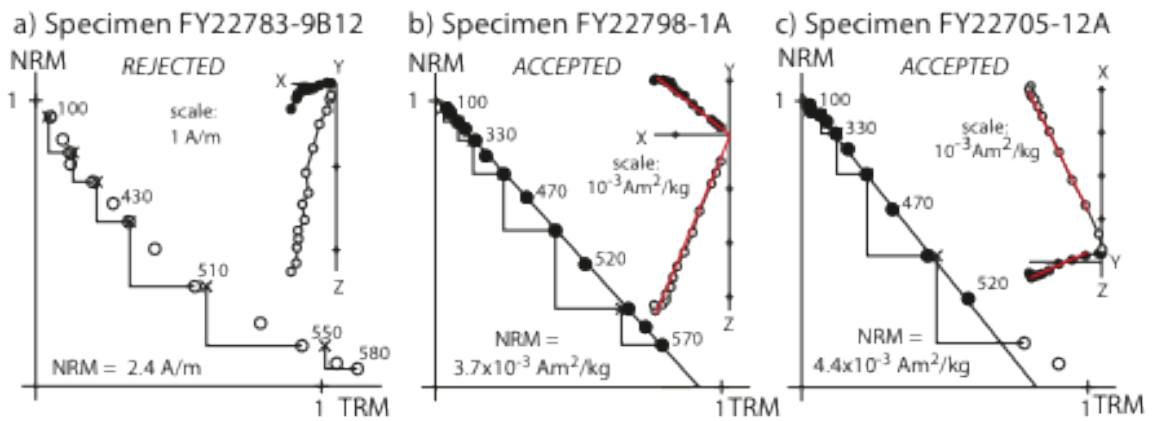
375 Figure 3: Stereographic plot of the TRM directions obtained on oriented block samples
376 from FY20462 and FY22783 fireplaces.



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379 Figure 4: Archaeointensity results of baked clay fragments (a-b) and of pottery sherds
 380 (c). Solid circles on NRM-TRM diagrams indicate the temperature steps used in the
 381 intensity determination. Corresponding demagnetization directions are shown in
 382 sample coordinates in the orthogonal diagrams. Open (solid) circles denote the
 383 projection on the vertical (horizontal) plane.

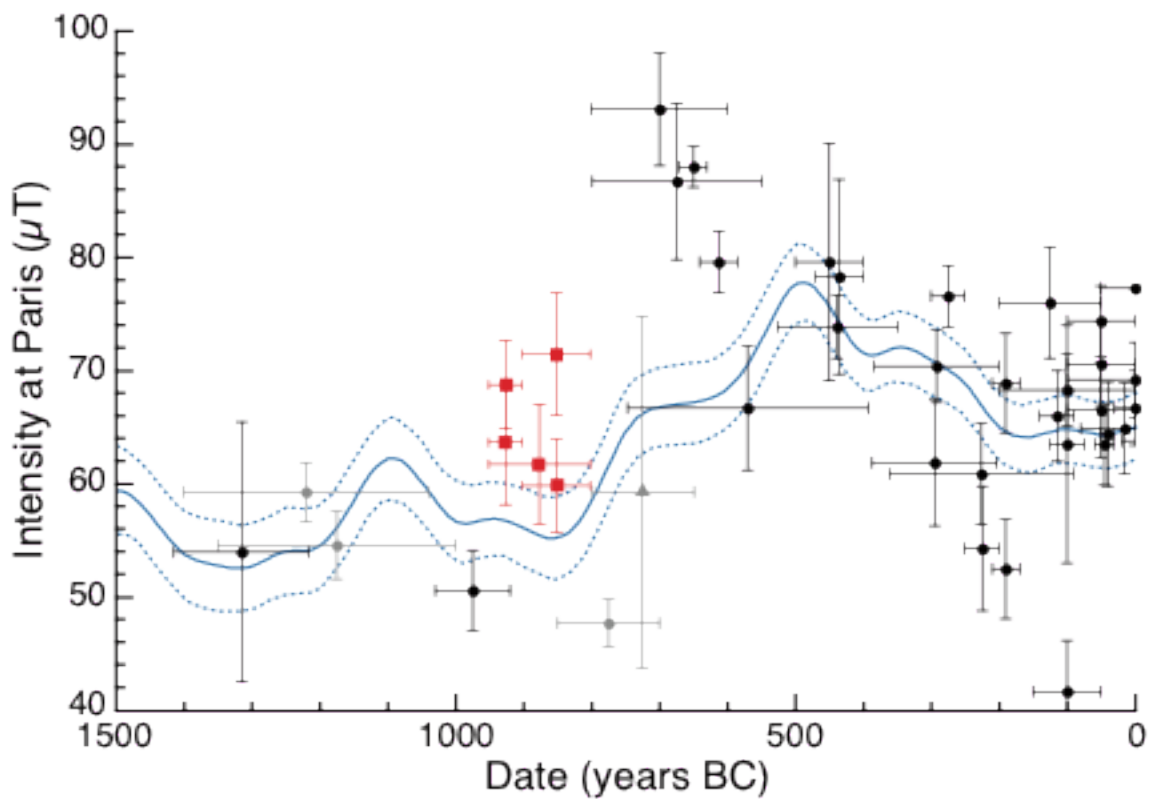


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386 Figure 5: Secular variation of the geomagnetic intensity at the Late Bronze and Iron Ages
387 in Western Europe. Corent data (red squares) are plotted with other published data
388 (black squares represent the selection of Hervé et al. 2013b, whereas gray circles and
389 triangles are the Swiss data of Kapper et al. 2015). The gray triangle indicates the Swiss
390 data with an unusual standard deviation discussed in the text. All data are relocated to
391 Paris. The blue curve is the mean intensity with its 95 per cent confidence envelop
392 predicted by the geomagnetic model SHA.DIF.14k (Pavón-Carrasco et al., 2014). The
393 Swiss data in grey are not included in this model.

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397 Table 1: Archaeological dating and mean archaeointensities of studied fireplaces.
 398 Fireplace, name of the sampled structure; Age, dating interval in years BC of the last use
 399 of the fireplace; N, number of specimens used in the calculation of the structure mean
 400 archaeointensity; $F \pm SD$, mean raw archaeointensity and standard deviation; $F_a \pm SD$,
 401 mean archaeointensity and standard deviation corrected for TRM anisotropy; $F_{a+c} \pm SD$,
 402 mean archaeointensity and standard deviation corrected for TRM anisotropy and
 403 cooling rate; F_{Paris} , mean archaeointensity relocated to Paris using Virtual Axial Dipole
 404 Moment correction.

Fireplace	Age (BC)	N	$F \pm SD$ (μT)	$F_a \pm SD$ (μT)	$F_{a+c} \pm SD$ (μT)	F_{Paris} (μT)
FY20462	[950; 900]	8	69.7 ± 5.6	70.3 ± 5.7	61.6 ± 5.5	63.6
FY22783	[950; 900]	12	65.2 ± 3.1	69.0 ± 3.8	66.5 ± 3.9	68.7
FY22798	[900; 800]	6	62.5 ± 5.5	60.7 ± 4.0	57.9 ± 4.1	59.8
FY22705	[900; 800]	12	71.5 ± 4.5	71.6 ± 4.8	69.2 ± 5.4	71.4
FY22842	[950; 800]	10	60.5 ± 4.8	60.8 ± 5.0	59.7 ± 5.3	61.7

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