

Vanadium

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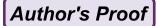
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Vanadium

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Element Data
6
        Atomic Symbol: V
        Atomic Number: 23
        Atomic Weight: 50.9415 g/mol
        Isotopes and Abundances: <sup>50</sup>V 0.25 %, <sup>51</sup>V 99.75 %
10
         1 Atm Melting Point: 1910 °C
11
         1 Atm Boiling Point: 3407 °C
12
         Common Valences: 5+, 4+, 3+, 2+
13
        Ionic Radii: 64 pm (in 3+, 6-fold coordination)
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        Pauling Electronegativity: 1.63
        First Ionization Energy: 6.74 eV
16
         Chondritic (CI) Abundance: 54.6 ppm<sup>a</sup>
17
        Silicate Earth Abundance: 86 ppm<sup>a</sup>
18
         Crustal Abundance: 138 ppm<sup>b</sup>
19
        Seawater Abundance: ~30-37 nmol/kg<sup>c</sup>
20
         Core Abundance: ~150 ppm<sup>d</sup>
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         <sup>a</sup>Palme et al. (2014)
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        <sup>b</sup>Rudnick and Gao (2014)
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         <sup>c</sup>Bruland et al. (2014)
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        <sup>d</sup>McDonough (2014)
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Properties

Vanadium (chemical symbol, V) is a d-block transition metal, silver in color, appearing in the first long period of the periodic table between titanium and chromium. Vanadium has two stable isotopes: ⁵⁰V and ⁵¹V, with atomic abundance of 0.25 % and 99.75 %, respectively. Vanadium has several 31 oxidation forms (between 2+ and 5+). In the lithosphere, 32 V occurs as reducing V(III) form, whereas in oxidizing conditions V prevails under V(IV) form. Vanadium(II) is partic- 34 ularly unstable in the environment. Vanadium(III) is more 35 stable than V(II), but it is also gradually oxidized by the air 36 or dissolved oxygen. Vanadium(V) is expected to be the 37 prevailing form in waters exposed to atmospheric oxygen, 38 whereas V(IV) may be present in reducing environments. 39 Depending upon geometry and environment, V ionic radii 40 vary between 36 pm and 79 pm. Vanadium has a high melting 41 point of 1910 °C and is a mildly incompatible, refractory, 42 lithophile (siderophile in the iron core and chondrites) ele- 43 ment. Vanadium has an electronegativity of 1.63 on the Pau- 44 ling scale and displays a first ionization potential of 45 6.74 eV. More details can be found in Richards (2006) and 46 Haynes (2015).

History and Use

Vanadium was first discovered by del Rio in 1801 and was 49 incorrectly considered as impure chromium. Vanadium was 50 rediscovered and named by Sefström in 1830 after the Scan- 51 dinavian goddess Vanadis. In 1867, Roscoe isolated V in 52 nearly pure form by reducing V chloride with hydrogen. 53 Vanadium of high purity (99.3-99.8 %) was not produced 54 until 1927. The V redox chemistry and its specific application 55 to renewable energy technologies (e.g., vanadium redox bat- 56 tery; a type of rechargeable flow battery using V in different 57 oxidation states to store chemical potential energy) have 58 already add to the demand for this element and will definitely 59 continue to increase in the near future. Vanadium demand for 60 construction materials is the greatest consumer of mined V - 61 (Moskalyk and Alfantazi, 2003). Indeed, vanadium is the 62 most widely used alloying element for strengthening steels 63 employed in buildings and bridges. Vanadium volume that is 64 cycled through terrestrial and aquatic reservoirs can thus be 65

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2 Vanadium

expected to further growth. It results in uncontrolled V releases in the environment, and environmental issues appear in developing and developed countries (Imtiaz et al., 2015).

Geochemical Behavior

Widely and sparsely distributed, V is not found as the free metal in nature. It originates from primary sources such as ores (i.e., iron oxides deposits), metallurgical slags (i.e., processing of U and Ti ores), and petroleum residues. Other sources include vanadiferous sandstones, bauxite, coals, or oil

Vanadium is widely distributed in igneous and sedimentary rocks and minerals. Among more than fifty minerals, carnotite $[K_2(UO_2)_2(VO_4)_2.3H_2O]$, roscoelite $[KV^{3+}_2(Si_3A])$ O₁₀(OH)₂], vanadinite [Pb₅(VO₄)₃Cl], mottramite [PbCu (VO₄)(OH)], and patrónite (VS₄) are the most important V-carriers. It is also present in some crude oils as organic complexes. Its average content in the Earth's crust is approximately 138 ppm and is similar to that of Zn and Ni. It is, however, more dispersed in the crust than either of those elements, and concentrated mineral deposits are consequently rare. The average V contents of CI chondrites range from 54.6 ppm to 56.5 ppm. The silicate Earth has a V content of 86 ppm, the bulk Earth of 105 ppm, whereas its content in the metallic core is slightly higher (150 ppm). Vanadium contents in the bulk continental crust vary significantly from 96 ppm to 230 ppm. These aspects have been further reviewed by McDonough (2014), Palme et al. (2014), and Rudnick and Gao (2014).

Historically, geochemists' interest in vanadium has derived from the occurrence of variable redox states in Earth surface environments, +3, +4, +5, and as a consequence its application as a redox indicator (Huang et al., 2015). Indeed, V chemical speciation and solubility are strong functions of pH and Eh conditions. Incorporation of reduced V species (III, IV) in clays and other secondary minerals provides a useful redox indicator for past anoxic conditions. Moreover, V behaves as an incompatible element because V(III) is preferentially incorporated into crystals: V-bearing minerals which crystallized in equilibrium with a low fO2 magma have a higher V content than crystals from the same magma with high fO₂ (Mallmann and O'Neill, 2009). It results in an increasing interest in V partitioning during temperature – high pressure magma differentiation processes. Studies on V partitioning among Earth's crustal reservoirs have provided critical insights into the interaction and evolution of these reservoirs. Vanadium depletion in the silicate Earth is explained by it preferential partitioning into the metallic core during "deep magma ocean" process (at high pressure and low fO₂; Wood et al., 2006). The core could

contain half of the total V budget of the bulk Earth 116 (McDonough, 2014). Perspectives to exploit V stable isotope 117 chemistry are considered: Wu et al. (2015) proposed first- 118 principles calculations that predict V isotopes fractionation 119 among V species with different valences in aqueous systems 120 and during sorption of V(V) to goethite.

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Vanadium is also a naturally occurring element in air, soil, 122 plants, and water. The concentration of dissolved V in natural 123 freshwater and seawater is generally several mg/L or lower 124 (Bruland et al., 2014; Gaillardet et al., 2014). Vanadium 125 concentrations in marine and nonmarine sediments are gen- 126 erally below 100 ppm, reflecting a net accumulation from the 127 aqueous phase. Vanadium in trace amounts represents an 128 essential element for normal cell growth, but it may cause 129 adverse effects when its concentration is much greater than a 130 few tenths of µg/L (Chatterjee, 2009). Most data on the 131 release of V into the environment have been related to industrial activities, especially from oil refineries and power plants 133 using V-rich fuel oil, and coal crude oil is enriched in V with 134 respect to many other trace elements, with concentrations 135 occasionally exceeding 1 mg/L. Thus, the fraction of 136 dissolved V in surface waters might be an environmental 137 indicator of oil combustion or pollution (Hope, 2008). Such 138 pollution sources may be responsible for appreciable amounts 139 of V into the environment, well above the natural background 140 levels associated with rock weathering and sediment 141 leaching. Fluvial dissolved V concentrations might also be 142 indicative of the types of rocks being weathered or of the 143 nature of the involved weathering processes. Weathering rate 144 and source rock type, rather than solution chemistry or anthropogenic influences, appeared to be the important controlling 146 factors on fluvial dissolved V concentrations (Hope, 2008).

The oxidation rate of V(IV) to V(V) and the equilibrium 148 between these two species in aqueous solution will regulate 149 V prevalence in water (Pourret et al., 2012). It further depends 150 on several factors, such as pH, V concentration, redox poten- 151 tial, ionic strength of the aqueous system, and biological 152 activity. In water, V(IV) is commonly present as a vanadyl 153 cation $[VO^{2+}, VO(OH)^{+}]$, whereas V(V) exists as a vanadate 154 oxyanion (H₂VO₄⁻, HVO₄²⁻). VO²⁺ is strongly adsorbed 155 onto solid phases, including organic and oxyhydroxide 156 phases. Adsorption of anionic V (H₂VO₄⁻, HVO₄²⁻) is 157 much lower than the cations; however, VO²⁺ solubility may 158 be greatly increased through complexation with organic matter. Whereas V(IV) is not thermodynamically stable at 160 pH > 7, complexation by various organic and inorganic species may considerably increase its stability. Eventually, V (V) oxidation state ion is more toxic than V(IV) ion one.



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Biological Utilization and Toxicity

Whereas low concentrations of V are required for plants to 165 positively influence chlorophyll synthesis, K consumption, or 166 N assimilation, higher V concentrations are toxic as they 167 cause chlorosis and limit growth. At trace amounts V is essen-168 tial for normal cell growth, whereas possibly toxic at high 169 concentrations, especially when occurring as pentoxid 170 Chronic exposure may result in inflammation of bronchi and 171 trachea, eyes and skin irritation, pulmonary edema, and systemic poisoning 173

Summary

Vanadium's occurrence receives more attention as the global 175 demand for V increases. However, its prevalence in environment becomes a new critical issue. Vanadium has been suggested to be a potentially dangerous pollutant, and the 178 United States Environmental Protection Agency classifies 179 V in the priority list of environmental risk elements. Fixing 180 the environmental issue of V is constrained by the limited 181 understanding of V biogeochemistry relative to other d-block 182 transition metals. better understanding 183 V biogeochemical behavior may support to assess the risk to 184 the environment and to human health and to assist in devel-185 oping new remediation tools. 186

Cross-References 187

- ► Complexation 188
- ► Lithophile Elements 189
- Oxidation-Reduction 190
- **▶** Transition Elements

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