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Origin of the Late Jurassic to Early Cretaceous peraluminous granitoids in the northeastern Hunan province (middle Yangtze region), South China: Geodynamic implications for the Paleo-Pacific subduction

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Abstract

The Late Mesozoic granitic belt in the northeastern Hunan province (situated in the south of the middle Yangtze region) represents the western front of the large magmatic province of SE China. In order to determine their ages and petrogenesis, we carried out zircon U–Pb dating, Hf isotope and whole-rock geochemical analyses for four granitic plutons, namely Taohuashan, Dayunshan-Mufushan, Wangxiang and Lianyunshan. Our SIMS zircon U–Pb ages, together with previously published data, reveal that the magmatic activities in this area can be roughly subdivided into three phases at 151–146 Ma, 132–127 Ma and ca. 117 Ma, and the Dayunshan-Mufushan batholith therein is a composite pluton. These four plutons are mainly composed of weakly to strongly peraluminous biotite or two-mica monzogranites, with a minor amount of biotite granodiorites. Their geochemical features are similar to S-type as well as fractionated S-type granites, with enrichment in LREEs and negative Ba, Sr, Nb, P and Ti anomalies. All samples show negative zircon $\epsilon_{\text{Hf}}(t)$ values ranging from -12.5 to -3.6, corresponding to crustal Hf model (T_{DM}^{C}) ages of 1.4–2.0 Ga. It is inferred that these granitoids were derived from partial melting of metasedimentary rocks analogous to the Neoproterozoic Lengjiayi Group, predominantly with psammitic component. Fractional crystallization probably played an important role in the magma evolution, while input of mantle-derived magma was insignificant. Combined with other geological evidence, our new data allow us to propose that the Cretaceous (132–127 Ma and ca. 117 Ma) magmatism might be response to episodic slab rollback of the

Paleo-Pacific plate, while the early-stage (151–146 Ma) magmatism that overlapped the epilogue of Jurassic magmatic flare-up and subsequent magmatic quiescence probably foreshadowed the transformation from foundering of a subducted flat-slab to slab rollback. Alternatively, slab foundering after a SE-directed intracontinental subduction in the central SCB cannot be ruled out for geodynamic interpretation of the Jurassic magmatism in SE China.

Keywords: South China; Late Mesozoic tectono-magmatism; Peraluminous granitoids; Geochronology and petrogenesis; Paleo-Pacific subduction

1. Introduction

Southeast China is an important part of the circum-Pacific magmatic and metallogenic belt (e.g., Zhou and Li, 2000; Zhou et al., 2006). Jurassic-Cretaceous (referred to as the “Yanshanian” period in Chinese literature) intrusive and volcanic rocks occupy a vast area of SE China (mostly in Anhui, Zhejiang, Jiangxi, Fujian and Guangdong provinces), forming a ca. 600 km wide magmatic belt parallel to the present coastline (Fig. 1). Despite intensive research in the past decades, origin and evolution of this large magmatic province are still hotly debated. For instance, its temporal-spatial configuration and geodynamic setting have not been well understood yet. The Jurassic (ca. 195–150 Ma) igneous rocks are widespread in the hinterland (e.g., the Nanling Range), while the Cretaceous (ca. 140–90 Ma) intrusive rocks coeval with voluminous volcanic rocks are mainly distributed in the coastal area and the middle-lower reaches of the Yangtze River. An oceanward younging trend for the Late Mesozoic magmatism resulted from northwestward subduction of the Paleo-Pacific plate with an increasing subduction angle had been proposed (Zhou and Li, 2000). Instead, Wang et al. (2011) recently suggested that the 180 to 125 Ma igneous rocks experienced a northeastward migration due to southwestward Paleo-Pacific subduction and corresponding slab rollback. Furthermore, it is noteworthy that more Jurassic granites have been identified in the coastal area in recent years (Cui et al., 2012; Liu et al., 2012; Zhang et al., 2015). Thus temporal-spatial distribution of the Jurassic-Cretaceous magmatic rocks deserves a reappraisal. Geodynamic trigger of the Mesozoic tectono-magmatism in eastern China has been generally ascribed to the subduction of the Paleo-Pacific plate beneath the Eurasian continent. However, the subduction pattern is still an open question,

diverse models have been proposed for SE China, such as shallow subduction and subsequent slab rollback (Zhou and Li, 2000; Jiang et al., 2009, 2011), break-up and foundering of a subducted flat-slab (Li and Li, 2007; X.H. Li et al., 2007, 2013b; Z.X. Li et al., 2012), ridge subduction (Ling et al., 2009, Sun et al., 2010; H. Li et al., 2012) and repeated slab advance-retreat (Jiang et al., 2015; Wang et al., 2016). Particularly, the Paleo-Pacific subduction beneath SE China was supposed to initiate during the Permian (Li and Li, 2007), Late Triassic (Jiang et al., 2015), Early Jurassic (Zhou et al., 2006) or Early Cretaceous (Chen et al., 2008). Other outdated or minority models include Alpine-type continental collision (Hsü et al., 1990), wrench faulting (Xu, 1993), strike-slip faulting plus concomitant rifting (Gilder et al., 1996), and lithospheric extension maybe unrelated to the Paleo-Pacific subduction (Li, 2000).

More importantly, numerous polymetallic deposits in SE China have been inferred to be closely related to the Mesozoic granitoids (Wang et al., 2011; Mao et al., 2013 and references therein). In addition to the geodynamic setting, controversial issues on the granitoids themselves involve their petrogenetic types (such as I-type, S-type or A-type granites, as well as fractionated granites) and the nature of source rocks, especially the degree of mantle contribution to the magma generation (e.g., X.H. Li et al., 2007; Wong et al., 2009; Jiang et al., 2009, 2011; Huang et al., 2013, 2015; Huang and Jiang, 2014; Zhang et al., 2015; Wang et al., 2016). Several Late Mesozoic granitic plutons crop out in the northeastern part of Hunan province, situated in the south of the middle Yangtze region (middle reach of the Yangtze River; Figs. 1 and 2). These granitic plutons represent the western front of the large magmatic province of SE China. However, due to the lack of detailed investigations, the geochronology and petrogenesis of these granitic rocks remain in dispute (Wang and Deng, 2004; Li et al., 2005; L.X. Wang et al., 2008, 2014; Xu et al., 2009). In this paper, we present zircon U–Pb ages and Hf isotopic data, major and trace element geochemical data for four granitic plutons (namely Taohuashan, Dayunshan-Mufushan, Wangxiang and Lianyunshan) in this area. This provides us an opportunity to decipher their origin and tectonic implications for the Late Mesozoic tectono-magmatic history of SE China.

2. Geological setting

2.1. Tectonic framework

Southeast China belongs to the eastern South China block (SCB) that consists of the Yangtze craton in the northwest and the Cathaysia block in the southeast (Fig. 1). It is generally accepted that the amalgamation of the Yangtze and Cathaysia blocks formed the Neoproterozoic Jiangnan orogen, with subduction-collision process before ca. 820 Ma (e.g., Charvet et al., 1996; Li et al., 2009a; Zhao et al., 2011; X.L. Wang et al., 2014; Zhao, 2015 and references therein). Subsequently, the SCB underwent continental rifting after ca. 820 Ma, with development of the Nanhua failed rift basin followed by platform sedimentation (Wang and Li, 2003). During the Phanerozoic, the SCB experienced three main tectono-thermal events (for a review, see Y.J. Wang et al., 2013 and references therein), including: (1) Early Paleozoic intracontinental orogeny mainly in the Wuyi-Yunkai domain of the Cathaysia block (e.g., Lin et al., 2008; Faure et al., 2009; Charvet et al., 2010; Z.X. Li et al., 2010; Shu et al., 2014, 2015); (2) Triassic orogeny around the SCB such as the northward continental deep-subduction along the Dabie-Sulu belt to the north and the South China-Indochina collision to the south, as well as in interior of the SCB (i.e., the intracontinental Jiuling-Xuefengshan belt superposed upon the pre-existing Jiangnan orogen) (e.g., Chu et al., 2012a, 2012b; Chu and Lin, 2014; Faure et al., 2008, 2014, 2016); (3) Late Mesozoic intracontinental deformation and magmatism, especially the development of the so-called “Southeast China basin and range tectonics” under an overall extensional regime (e.g., Gilder et al., 1996; Li, 2000; Lin et al., 2000; Shu et al., 2009; Ji et al., 2014; J.H. Li et al., 2014).

The Jurassic-Cretaceous magmatic rocks are mostly distributed in the Cathaysia block, whereas a small amount of contemporaneous plutons crop out in the Yangtze craton. In the northeastern part of Hunan province, Late Mesozoic granitoids intruded into the middle segment of the Jiuling-Xuefengshan belt (Figs. 1 and 2). To the northwest, the Cretaceous-Cenozoic Jiangnan basin covers most area of the middle Yangtze craton. The Precambrian basement exposed in the study area is a pile of weakly metamorphosed (greenschist facies) flysch sequences, i.e., the Lengjiaxi Group, mainly consisting of metasandstone, metasilstone and slate (BGMRHN, 1988; Gu et al., 2002; Xu et al., 2007). As in other regions, the Lengjiaxi Group and its equivalents were intruded by ca. 820 Ma strongly peraluminous (S-type) granitoids (such as the largest Jiuling batholith to the southeast, northwestern Jiangxi province), and then unconformably covered by rifting-related clastic sedimentary rocks represented by the Neoproterozoic (Nanhuan period) Banxi Group and its equivalents (W. Wang et al., 2010, 2012, 2013; Gao et al., 2011; Meng et al., 2013; Zhao et al., 2011, 2013; X.L.

Wang et al., 2013, 2014; Yan et al., 2015). This angular unconformity is widely considered to be the response to final amalgamation of the Yangtze and Cathaysia blocks. The late Neoproterozoic (Sinian period) to Middle Triassic strata mostly distributed in the northeastern part of the study area are roughly continuous, except for a distinct hiatus between the Silurian and Devonian (BGMRHN, 1988). They mainly consist of carbonate and siliciclastic rocks in a shallow-marine environment. Two regional unconformities, respectively below the Middle Devonian conglomerate and the sporadically Late Triassic to Jurassic terrigenous clastic rocks, correspond to the Early Paleozoic and Triassic intracontinental orogeny of the SCB. All the rocks mentioned above were strongly folded, and unconformably overlain by NE-trending small fault depressions filled with Cretaceous-Paleogene alluvial and lacustrine deposits (red beds).

2.1 Geology of the granitoids

Four granitic plutons (namely Taohuashan, Dayunshan-Mufushan, Wangxiang and Lianyunshan) in the northeastern Hunan province and adjacent areas have been investigated in this work (Fig. 2). The dominant rock type of these plutons is biotite or two-mica monzogranite, with a minor amount of biotite granodiorite (BGMRHN, 1988; Wang and Deng, 2004; L.X. Wang et al., 2008, 2014; Xu et al., 2009). The main minerals of the monzogranite are quartz, alkali feldspar, plagioclase and biotite, with or without muscovite (Fig. 3a, b, c). Primary muscovite is commonly present in the two-mica monzogranite or leucogranite. Occasionally, minor muscovite (generally <1 %) is also visible in the biotite monzogranite. Accessory minerals in the monzogranite mainly include apatite, zircon and monazite (Fig. 3d). The biotite granodiorite has a mineral assemblage of quartz, plagioclase, alkali feldspar, biotite and minor hornblende, with accessory minerals such as apatite, zircon and titanite as well as epidote and allanite (Fig. 3e, f). Country rocks of these granitoids are predominantly Neoproterozoic slate, as well as locally Early Paleozoic limestone and shale. About 200 m to 2 km wide contact aureoles are usually developed around these plutons, where hornfels and schist are common.

The Taohuashan pluton crops out over an area of about 150 km², despite its considerable part covered by Quaternary deposits. It is mainly composed of fine- to medium-grained two-mica monzogranite with subordinate porphyritic biotite

monzogranite in the north. Contacts between the two-mica and biotite phases are gradational, and the former usually shows weak gneissic structure.

The Dayunshan-Mufushan batholith is a huge composite pluton, with a total outcrop area of about 2440 km². It consists of early-stage biotite-rich intrusion and late-stage two-mica intrusion. The former is mainly composed of porphyritic biotite monzogranite, as well as subordinate fine- to medium-grained biotite granodiorite in the northeastern part; while the latter is composed of fine- to medium-grained two-mica monzogranite, with minor phenocrysts sometimes. The early-stage intrusion was dismembered into three larger parts due to the disturbance of the late-stage intrusion. Finally, fine-grained stocks of two-mica or biotite monzogranite intruded the previous facies. As shown later, there is an interval of about 20 Myr between the early and late intrusions. Except for the common andalusite hornfels, garnet- and/or staurolite-bearing schist can also be locally found in the contact zone. In addition, a detachment fault with half-graben basin on the hanging wall developed along the western tongue-shaped margin of the batholith (Dayunshan area), and the late-stage intrusion was involved in mylonitization (Yu et al., 1998).

The Wangxiang batholith is located in the Wangcheng-Xiangyin area, and shows an elliptical shape with an outcrop area of about 1480 km². The main phase of this batholith is fine- to medium-grained or porphyritic two-mica monzogranite with minor fine- to medium-grained biotite granodiorite as marginal facies in the northern part. Moreover, supplementary pulse of two-mica monzogranite that occurs as stocks or apophysises can be distinguished from the main phase according to field occurrence and petrography.

The Lianyunshan pluton was cut and deformed by a NE-trending fault zone that also constitutes the boundary fault of a narrow and long rift basin to the west. This pluton has an outcrop area about 130 km², and is lentoid in shape with a bulgy body in the northeast and forked tails in the southwest. It is mainly composed of muscovite-bearing monzogranite with fine- to medium-grained granitic texture or porphyritic texture, usually showing deformation features. Garnet, sillimanite and staurolite are visible in schist of the contact metamorphic zone, indicating a relatively high-temperature metamorphism. Particularly, the thermal metamorphism extends up to 10 km wide to the NE, locally accompanied by weakly migmatization.

3. Analytical methods

Fifteen representative samples were collected from the four plutons mentioned above, and zircons from eleven samples therein were determined for SIMS U–Pb dating and LA-MC-ICPMS Hf isotope analyses. In addition, ten samples were selected for whole-rock geochemical analyses.

Zircon grains were separated using standard density and magnetic separation techniques. They were mounted in epoxy mounts together with zircon standards, and then polished to section the crystals in half for analysis. In order to investigate the morphology and internal textures of the zircons, which were documented with transmitted and reflected light micrographs, as well as cathodoluminescence (CL) images (Fig. 4). Measurements of U, Th and Pb isotopes were conducted using a Cameca IMS-1280 at the Institute of Geology and Geophysics, Chinese Academy of Sciences (IGGCAS). The instrument description and analytical procedure can be found in Li et al. (2009b). The beam spot was about $20 \times 30 \mu\text{m}$ in size, and positive secondary ions were extracted with a 10 KV potential. Standard zircons Plešovice (Sláma et al., 2008) and Qinghu (Li et al., 2013a) were alternately analyzed with unknown zircons from the samples. Measured compositions were corrected for common Pb using the non-radiogenic ^{204}Pb . Data reduction was carried out using the Isoplot program (Ludwing, 2003). Uncertainties of individual analyses are reported at 1σ in Supplementary Table S1. The concordia diagrams of U–Pb data are shown in Figure 5, where weighted mean $^{206}\text{Pb}/^{238}\text{U}$ ages or concordia ages are quoted at either 95% confidence interval or 2σ , as summarized in Table 1.

In situ Hf isotope analyses were conducted mostly on the dated zircons with concordant ages by using a Neptune MC-ICPMS, equipped with a Geolas 193 nm laser-ablation system at IGGCAS. The detailed analytical technique and calibration methods are described in Wu et al. (2006). During analyses, zircon standards GJ-1 and Mud Tank were used as reference materials (Woodhead and Hergt, 2005; Morel et al., 2008). The spot size was either 40 or 60 μm depending on grain sizes, with a laser repetition rate of 8 Hz and laser beam energy density of 10 J/cm^2 . The chondritic Hf-isotope values of Blichert-Toft and Albarède (1997) were used for calculation of $\epsilon\text{Hf}(t)$ values. Depleted-mantle Hf model ages (T_{DM}) were calculated using the measured $^{176}\text{Lu}/^{177}\text{Hf}$ ratios, in reference to the depleted mantle reservoir with a present-day $^{176}\text{Hf}/^{177}\text{Hf} = 0.28325$ and $^{176}\text{Lu}/^{177}\text{Hf} = 0.0384$ (Griffin et al., 2000). We also present crustal Hf model ages (T_{DM}^{C}) for crustal sources that were originally derived from the depleted mantle, assuming that the parental magma was produced from an average continental

crust with $^{176}\text{Lu}/^{177}\text{Hf} = 0.015$ (Griffin et al. 2002). The Hf isotope data with uncertainties at 2σ are reported in Supplementary Table S2, and also summarized in Table 1. For comparison, histograms of the $\epsilon\text{Hf}(t)$ values of each samples are shown in Figure 6.

Whole-rock geochemical analyses were carried out at IGGCAS. Major elements were determined by XRF (X-ray fluorescence) using fused glass disks. Trace elements (including rare earth elements, REEs) were determined for the same samples by Agilent 7500a ICP-MS after acid digestion. Analytical uncertainties are generally less than 5% for major elements and 10% for trace elements. The geochemical data are presented in Table 2.

4. Results

4.1. Zircon U–Pb ages and Hf isotopic data

4.1.1. Taohuashan pluton

Zircons from a biotite monzogranite (JH155) of the Taohuashan pluton are mostly euhedral, colorless and transparent, about 60–200 μm in length. Most grains are characterized by magmatic growth zonation under CL (Fig. 4). Nineteen analyses of zircons were obtained during a single analytical session. Their U and Th contents respectively vary in the ranges of 276–2159 ppm and 235–1194 ppm, with Th/U ratios between 0.28 and 0.95 (Table S1). Excluding one discordant analysis (spot #12), the remaining analyses are concordant in $^{206}\text{Pb}/^{238}\text{U}$ and $^{207}\text{Pb}/^{235}\text{Pb}$ within analytical errors and yield a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 128.0 ± 1.0 Ma (MSWD = 0.79) (Fig. 5a). This age is interpreted as the best estimate of the crystallization age for the Taohuashan biotite monzogranite.

Eighteen spots on the dated zircons were further analyzed for Hf isotopic composition, their $^{176}\text{Hf}/^{177}\text{Hf}$ ratios range from 0.282409 to 0.282500, corresponding to $\epsilon\text{Hf}(t)$ values of -10.1 to -6.9 and T_{DM}^{C} ages of 1.62–1.82 Ga (Fig. 6a; Table S2).

4.1.2. Dayunshan-Mufushan batholith

Six samples were collected from the Dayunshan-Mufushan batholith for U–Pb dating and Hf isotope analyses, including DM35 and QD69 respectively from the

northern and southern parts of the early-stage biotite monzogranite, DYS14 from the early-stage biotite granodiorite, QD49 and DM80 respectively from the western and eastern parts of the late-stage two-mica monzogranite, as well as QD62 from a later stock of two-mica monzogranite.

Zircons from sample DM35 are commonly euhedral, ranging from 100 to 300 μm in length with aspect ratios of 1.5:1 to 3:1. They show conspicuous oscillatory zoning and rare inherited cores under CL (Fig. 4). Eighteen zircons were analyzed, their U and Th contents respectively range from 517 to 2406 ppm and from 256 to 793 ppm, with Th/U ratios of 0.21–0.65 (Table S1). Two discordant analyses indicate their partially inherited origin, the remaining analyses have similar $^{206}\text{Pb}/^{238}\text{U}$ ages with a weighted mean age of 151.2 ± 1.1 Ma (MSWD = 0.26) (Fig. 5b).

Zircons from sample QD69 are mostly euhedral, with lengths of 100–250 μm and aspect ratios of 2:1 to 3:1. Most grains are nearly black and display faint zoning in CL images, a few of which contain bright inherited cores (Fig. 4). Nineteen analyses were conducted, which are all concordant. The only spot on an inherited core obtained a $^{206}\text{Pb}/^{238}\text{U}$ age of 837 ± 12 Ma. Sixteen out of the remaining eighteen analyses yielded a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 151.4 ± 1.1 Ma (MSWD = 0.43), which agrees with the concordia age of 151.3 ± 1.1 Ma (MSWD of concordance = 0.53) (Fig. 5c). The two rejected analyses with older ages (159.6 and 160.1 Ma) are due to very high U concentrations (up to 5370 and 7007 ppm; Table S1).

Accordingly, ca. 151 Ma is the best estimate of crystallization age for the Dayunshan-Mufushan biotite monzogranite. The zircons with ages around ca. 151 Ma from sample DM35 and sample QD69 have similar Hf isotopic compositions. A total of thirty-three measured $^{176}\text{Hf}/^{177}\text{Hf}$ ratios vary from 0.282440 to 0.282564, corresponding to $\epsilon\text{Hf}(t)$ values of -8.5 to -4.3 and T_{DM}^{C} ages of 1.47–1.74 Ga (Fig. 6b; Table S2). In addition, the ca. 837 Ma inherited core from sample QD69 shows positive $\epsilon\text{Hf}(t)$ value of 4.3.

Zircons from sample DYS14 are mostly euhedral, about 100–250 μm in length with aspect ratios of 1.5:1 to 3:1. Oscillatory zoning of magmatic origin is common in CL images (Fig. 4). Twenty-one spots on twenty zircons were analyzed, all of which obtained concordant analyses. They have moderate U (261–948 ppm) and Th (74–620 ppm) contents, with Th/U ratios of 0.08–0.98 (Table S1). One analysis of inherited core (spot #15) yielded a $^{206}\text{Pb}/^{238}\text{U}$ age of 819 ± 12 Ma, nineteen out of the remaining twenty analyses yielded a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 149.0 ± 1.0 Ma (MSWD =

0.71) (Fig. 5d). This age is interpreted as crystallization age of the Dayunshan-Mufushan biotite granodiorite. In addition, sixteen Hf isotope analyses of the ca. 149 Ma zircons have $^{176}\text{Hf}/^{177}\text{Hf}$ ratios of 0.282438–0.282524. Their $\epsilon\text{Hf}(t)$ values vary from -8.6 to -5.6, and T_{DM}^{C} ages from 1.55 to 1.75 Ga (Fig. 6c; Table S2). By contrast, the ca. 819 Ma inherited core shows positive $\epsilon\text{Hf}(t)$ value of 3.2.

Zircons from sample QD49 are commonly euhedral, about 80–300 μm in length with aspect ratios varying from 1.5:1 to 4:1. A majority of the grains show magmatic growth zonation under CL, but considerable cores are present with various zoning patterns (Fig. 4). Twenty-one analyses show U contents of 88–3685 ppm and Th contents of 81–748 ppm respectively (Table S1). Three spots on the inherited cores obtained concordant analyses respectively with Proterozoic ages of 1246 ± 5 Ma ($^{207}\text{Pb}/^{206}\text{Pb}$ age, spot #18), 831 ± 12 Ma and 848 ± 12 Ma ($^{206}\text{Pb}/^{238}\text{U}$ age, spot #10 and #13) (Fig. 5e; Table S1). Apart from two discordant analyses, the remaining fifteen analyses forms a cluster, yielding a concordia age of 131.8 ± 1.5 Ma (MSWD of concordance = 0.21) (Fig. 5e).

Zircons from sample DM80 are predominantly elongated grains (about 80–200 μm in length and 20–50 μm in width) with dark CL images, some of which contains inherited cores (Fig. 4). Besides, there are also several tubby grains with bright zonation but no core-rim textures, which are likely xenolithic zircons. A total of thirty-three analyses were conducted. Nine spots on the xenocrysts or inherited cores obtained informative results, among which several concordant or near-concordant analyses yielded $^{206}\text{Pb}/^{238}\text{U}$ ages of 850–756 Ma as well as ca. 417 Ma (Fig. 5f). We speculated that these xenocrysts with old ages were captured from the wall-rocks during magma ascent. Twenty-four analyses of the dark crystals and rims show high U contents, mostly higher than 2000 ppm (Table S1). These spots are plotted as a cluster on or near the concordia curve, with apparent $^{206}\text{Pb}/^{238}\text{U}$ ages varying from 137.0 to 153.4 Ma. Despite a weighted mean age of 143.5 ± 1.8 Ma with a large MSWD value of 3.6 can be concluded, there is a distinct correlation that the apparent ages become older while the U concentrations increasing (Fig. 5f). Such a phenomenon is evident for the present of “high-U matrix effect” (White and Ireland, 2012; Gao et al., 2014). Therefore, the crystallization age of sample DM80 is theoretically younger than ca. 143.5 Ma, and probably approximates to the minimum $^{206}\text{Pb}/^{238}\text{U}$ age of 137.0 ± 2.1 Ma therein. Considering that sample DM80 was taken from the same intrusion as sample QD49, we suggest that the age of sample QD49 (ca. 132 Ma) is a reliable estimate of crystallization

age for the main body of the Dayunshan-Mufushan two-mica monzogranite.

Zircons from sample QD62 are euhedral to subhedral, about 50–250 μm in length with aspect ratios between 1:1 and 4:1. They are either dark with faint zonation or bright with distinct zonation under CL, and quite a few inherited cores can also be observed (Fig. 4). Twenty-three analyses were performed on twenty-one zircons, with Th/U ratios range from 0.08 to 0.42 (Table S1). Except for two discordant analyses resulted from inherited cores, the other analyses form a slightly scattered population with apparent $^{206}\text{Pb}/^{238}\text{U}$ ages ranging from 118.8 Ma to 140.4 Ma. Seven analyses therein with a bit older or younger ages were rejected, a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 127.0 ± 1.4 Ma (MSWD = 1.7) was concluded using the remaining fourteen analyses (Fig. 5i). We consider this age to be the crystallization age for the later stocks within the Dayunshan-Mufushan batholith.

The dated zircons from sample QD49 and sample QD 62 were chosen for *in situ* Hf isotope analyses (Fig. 6d; Table S2). Although crystallization ages of these two samples show a discrepancy about 5 Myr, their Hf isotopic compositions largely overlap with $^{176}\text{Hf}/^{177}\text{Hf}$ ratios varying from 0.282424 to 0.282589. All the analyses show $\epsilon\text{Hf}(t)$ values of -3.6 to -9.6 and T_{DM}^{C} ages of 1.42–1.79 Ga. However, the three concordant cores with Proterozoic ages from sample QD49 have positive $\epsilon\text{Hf}(t)$ values of 2.0 to 5.7 and T_{DM}^{C} ages of 1.50–1.69 Ga.

4.1.3. Wangxiang batholith

Two samples of two-mica monzogranite DYS01 and DYS03 from the main and supplementary phases of the Wangxiang batholith were dated by SIMS method. Zircons from sample DYS01 are relatively small, mostly less than 150 μm in length. CL images reveal that more than one third of them contain inherited cores (Fig. 4). We analyzed five spots on cores, and three analyses therein are concordant with a weighted $^{206}\text{Pb}/^{238}\text{U}$ mean age of 837 ± 13 Ma (Fig. 5j). Furthermore, the oldest discordant core with $^{207}\text{Pb}/^{206}\text{Pb}$ age of ca.1715 Ma indicates the presence of potential Paleoproterozoic material. Other eighteen analyses were conducted on rims or magmatic crystals without core. Apart from two analyses possibly due to high U content or Pb loss, the remaining sixteen analyses yielded a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 147.3 ± 1.5 Ma (MSWD = 1.7) (Fig. 5j).

Zircons from sample DYS03 are either elongated or stubby grains, the latter

usually contain inherited cores surrounded by dark rims (Fig. 4). A total of twenty-five analyses were performed on twenty-three zircons. Four analyses on zircon cores are discordant possibly due to partial mixing of core-rim domains; the remaining analyses plotted on concordia curve show a relatively wide range of apparent $^{206}\text{Pb}/^{238}\text{U}$ ages from 135.2 to 155.8 Ma. Excluding six analyses with slightly either older or younger ages therein, the other fifteen analyses yielded a weighted $^{206}\text{Pb}/^{238}\text{U}$ mean age of 146.0 ± 1.4 Ma (MSWD = 1.3) (Fig. 5k). This age is slightly younger but indistinguishable within error from the age of sample DYS01.

Twenty-one dated zircons from sample DYS01 and seventeen dated zircons from sample DYS03 were measured for Hf isotope determination (Fig. 6e; Table S2). The three inherited cores with ages of ca. 837 Ma from sample DYS01 have variable $\epsilon\text{Hf}(t)$ values (-4.1, 3.9 and 11.7). The remaining thirty-five zircons have similar Hf isotopic compositions with $^{176}\text{Hf}/^{177}\text{Hf}$ ratios of 0.282327–0.282517, corresponding to $\epsilon\text{Hf}(t)$ values of -12.5 to -5.9 and T_{DM}^{C} ages of 1.57–1.99 Ga.

4.1.4. Lianyunshan pluton

We collected two samples DYS06 (two-mica monzogranite) and DYS07 (biotite monzogranite) from the Lianyunshan pluton for age determinations. Zircons from sample DYS06 are euhedral, commonly 60–150 μm in length. Most grains show magmatic oscillatory zoning in CL images, a few inherited cores are also observed (Fig. 4). Twenty-four analyses were conducted, with Th/U ratios between 0.26 and 3.31 (Table S1). One of the two spots on inherited cores obtained a concordant analysis with $^{206}\text{Pb}/^{238}\text{U}$ age of 825.8 ± 11.6 Ma. Twenty-one analyses of magmatic crystals yielded a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 150.4 ± 1.7 Ma (MSWD = 2.3) that represents the crystallization age of sample DYS06 (Fig. 5).

Zircons from sample DYS07 are mostly prismatic, 80–200 μm in length with aspect ratios of 2:1 to 3:1. Although most grains display low luminance under CL, oscillatory zoning is more or less visible (Fig. 4). Furthermore, a few of the zircons also show clear core-rim textures, but most of the cores are homogeneous without zonation. A total of twenty-four spots on twenty-two zircons were analyzed, including five bright cores with low U contents (145–726 ppm) and nineteen dark crystals or rims with high U contents (1529–6341 ppm) (Table S1). Except for one discordant core, all the other analyses with similar apparent $^{206}\text{Pb}/^{238}\text{U}$ ages are plotted on the concordia curve. Apart

from two relatively discrete analyses, the remaining twenty-one analyses yielded a weighted mean $^{206}\text{Pb}/^{238}\text{U}$ age of 149.0 ± 1.4 (MSWD = 1.8) (Fig. 5). This age is interpreted as the crystallization age of sample DYS07, which is coincident within error with that of sample DYS06.

Twenty-two dated zircons from sample DYS06 were further analyzed for Hf isotopic compositions (Fig. 6f; Table S2). The ca. 826 Ma inherited core has a negative $\epsilon\text{Hf}(t)$ value of -3.6. The $^{176}\text{Hf}/^{177}\text{Hf}$ ratios of the other magmatic zircons ranges from 0.282330 to 0.282440, corresponding to $\epsilon\text{Hf}(t)$ values of -12.4 to -8.5 and T_{DM}^{C} ages of 1.74–1.98 Ga. By comparison, twenty-one Hf isotope analyses of the dated zircons from sample DYS07 have higher $^{176}\text{Hf}/^{177}\text{Hf}$ ratios of 0.282486–0.282579, corresponding to $\epsilon\text{Hf}(t)$ values of -6.9 to -3.7 and T_{DM}^{C} ages of 1.44–1.64 Ga (Fig. 6f; Table S2).

4.2. Whole-rock major and trace elements

Except for the major and trace element data of ten granitic samples in this study (Table 2), the data from previous literature for the studied plutons have also been selected for comparison. Together with the data presented by L.X. Wang et al. (2008, 2014), the two-mica and biotite monzogranite from the four plutons show high SiO_2 contents mostly varying from 70.2 to 74.0 wt.% (up to 76.6 wt.% for sample DYS07), while about 65.5 wt.% for the biotite granodiorite (Fig. 7a). All the granitic rocks (monzogranites and granodiorites) have moderate total alkali $\text{K}_2\text{O}+\text{Na}_2\text{O}$ contents of 5.8–8.6 wt.% with $\text{K}_2\text{O}/\text{Na}_2\text{O}$ ratios varying from 0.8 to 1.7, mostly belonging to high-K calc-alkaline series (Fig. 7a). On the A/CNK vs. A/NK diagram, they are weakly to strongly peraluminous, with aluminum saturation index ($\text{ASI} = \text{A}/\text{CNK}$) mostly between 0.99 and 1.17 (Fig. 7b).

The monzogranites as well as granodiorites have similar chondrite-normalized REE patterns, characterized by enrichment of LREEs relative to HREEs and variable degree of negative Eu anomalies (Fig. 8; Xu et al., 2009; L.X. Wang et al., 2008, 2014). Our samples have total REEs contents of 92–232 ppm, with $(\text{La}/\text{Yb})_{\text{N}}$ ratios of 20.6–97.8 and Eu/Eu^* ratios of 0.35–0.77 (Table 2). By comparison, the negative Eu anomalies of the 151–149 Ma granitoids (i.e., the early-stage intrusion of the Dayunshan-Mufushan batholith, Lianyunshan pluton) are overall less notable than the other younger granitoids. In the primitive mantle-normalized spidergrams, all the

samples are relatively enriched in large ion lithophile elements (LILEs) such as Rb, Th and K, and show distinct negative Ba, Nb, Sr, P and Ti anomalies with a few exceptions (Fig. 8; Xu et al., 2009; L.X. Wang et al., 2008, 2014). The Nb/Ta ratios of our samples mostly range from 5.9 to 11.7 (average at 9.6, except for the unusual sample DYS07), in accord with crustal-derived granites (Table 2).

5. Discussion

5.1. Emplacement ages of the granitoids in the northeastern Hunan province

As compiled in Table 1 of Shen et al. (2012), numerous geochronological data for the Late Mesozoic granitoids in the study area has been obtained by different methods. However, we notice that most of the data are whole-rock Rb–Sr isochron and biotite or muscovite K–Ar ages, as well as TIMS zircon U–Pb ages. These ages are highly variable ranging from 170 to 108 Ma, their reliabilities are hard to appraise. The Rb–Sr and K–Ar systems are easy to be reset by diffusion and hydrothermal alteration during post-magmatic process, while traditional TIMS method may give mixing age due to possible involvement of inherited cores in magmatic zircons. In contrast, *in situ* zircon U–Pb dating (LA–ICPMS or SIMS) is capable of providing reliable age for pluton emplacement.

Wang et al. (2008) firstly conducted LA–ICPMS zircon U–Pb dating for the Taohuashan two-mica monzogranite, and the age of 129.4 ± 1.1 Ma was considered as the crystallization age. Shen et al. (2012) also dated the Taohuashan pluton by LA–ICPMS method. They obtained a lower intercept age of 118.0 ± 3.7 Ma calculated mostly using discordant spots, and also a concordia age of 127.4 ± 0.6 Ma from four concordant analyses. However, the former was interpreted as the crystallization age, whereas the latter was considered to be inherited from an earlier intrusion phase. Our SIMS zircon U–Pb dating of the Taohuashan biotite monzogranite yielded a concordia age of 128.0 ± 1.0 Ma, which is in agreement with the age of the Taohuashan two-mica monzogranite reported by Wang et al. (2008). Therefore, we suggest that the emplacement age of Taohuashan pluton occurred during 127–129 Ma, and yet the lower intercept age of Shen et al. (2012) likely holds no significance. Besides, there is a small pluton named Xiaomoshan to the north, which yielded two identical LA–ICPMS zircon U–Pb ages of 117.4 ± 1.0 Ma and 116.6 ± 0.5 Ma (Wang et al., 2008; Shen et al., 2012).

This pluton is mainly composed of biotite monzogranite, representing the latest magmatic activity in the study area.

Recently, L.X. Wang et al. (2014) obtained several zircon U–Pb ages for the Dayunshan-Mufushan batholith by LA–ICPMS method, including 151.5 ± 1.3 Ma, 148.3 ± 1.4 Ma and 145.8 ± 0.9 Ma for the biotite granodiorite, biotite monzogranite and two-mica monzogranite, respectively. Moreover, a small dioritic stock in the northern part of the batholith yielded an older age of 154.0 ± 1.9 Ma. They suggested that the Dayunshan-Mufushan batholith intruded episodically from 154 to 146 Ma, with a magmatic evolution from mafic to felsic. However, such a prolonged (about 8 Myr) process of fractional crystallization proposed by L.X. Wang et al. (2014) is not supported by our new SIMS zircon U–Pb dating results. The two biotite monzogranite samples (DM35 and QD69) in this study have identical ages of 151.2 ± 1.1 Ma and 151.3 ± 1.1 Ma, while the age of biotite granodiorite sample (DYS14, 149.0 ± 1.0 Ma) is about 2 Myr younger than the former. By contrast, the two samples of two-mica monzogranite (QD49 and QD62) from the main body of the batholith and the later stock present Cretaceous ages of 131.9 ± 1.9 Ma and 127.0 ± 1.4 Ma, respectively. Thus we argued that the Dayunshan-Mufushan batholith is a composite pluton, which was products of multistage magmatic activities during the Late Jurassic (151–149 Ma) and Early Cretaceous (132–127 Ma).

Our SIMS zircon U–Pb dating also shows that the main and supplementary phases of the Wangxiang batholith crystallized in a very short period during 147–146 Ma, while the Lianyunshan pluton was emplaced during 150–149 Ma. So far, no *in situ* zircon U–Pb data was previously available for the Miaoshan, Changlejie, Cenchuan and Jinjing plutons in the study area (Fig. 2). Their rock associations are similar to the early-stage intrusion of the Dayunshan-Mufushan batholith, Wangxiang batholith and Lianyunshan pluton (BGMRHN, 1988; Wang and Deng, 2004; Li et al., 2005). Combining with the unreliable radiometric results in the report of regional geological surveys, we infer that they were likely also emplaced during the Late Jurassic (151–146 Ma). Nevertheless, precise dating of these plutons is expected in the future.

5.2. Petrogenetic type

Chappell and White (1974, 2001) initially proposed the classification of I- and S-type granites. The I-type granites are mainly metaluminous ($ASI < 1.0$) and have higher

Na₂O contents (normally >3.2 wt.%), derived largely by partial melting of metagneous rocks; while the S-type granites are usually strongly peraluminous (ASI >1.1) and have lower Na₂O contents (normally <3.2 wt.% with about 5 wt.% K₂O), derived largely by partial melting of metasedimentary rocks. In view of the data from both this study and the literature (L.X. Wang et al., 2008, 2014), the majority of the Late Mesozoic granitoids in the northeastern Hunan province is weakly to strongly peraluminous, and have moderate Na₂O contents of 2.8–3.9 wt.% (average at 3.3 wt.%). In this regard, it seems difficult to attribute these granitoids to typical I-type or S-type granites, implying that source heterogeneity or magma mixing may contribute to their origin (Clemens, 2003). On the ACF diagram, however, all the granitic samples plot within the field of S-type granites, in agreement with their muscovite-bearing or biotite-rich peraluminous compositions (Fig. 9a). Furthermore, there is no obvious decrease of P₂O₅ when SiO₂ increases (without respect to the granodiorites; Fig. 9b), and increasing of Y and Th against Rb (Fig. 9c, d). These trends are inconsistent with the features of the I-type granites in the Lachlan fold belt (Chappell, 1999), and also the fractionated I-type granites in SE China (Li et al., 2007; Zhang et al., 2015). In addition, most of the granitic samples have relatively low 10000×Ga/Al (<2.6) and Zr+Nb+Y+Ce (<350 ppm), low to moderate FeO*/MgO and (K₂O+Na₂O)/CaO, falling into the fields of unfractionated and minor fractionated granites (Fig. 9e, f). The above criteria indicate that these rocks are significantly different from the A-type granites (Whalen et al., 1987). Thus, the Late Mesozoic granitoids in the study area, especially the dominant monzogranites, are more similar to S-type as well as fractionated S-type granites. Such a deduction is further supported by the fact that metasedimentary xenoliths or restites are largely present within these granitic rocks. Moreover, considerable zircons in our dated samples contain old inherited cores, indicative of partial melting of metasedimentary rocks.

Barbarin (1996) once further distinguished two main types of peraluminous granitoids, namely muscovite-bearing granitoids (MPGs, unusually two-mica monzogranite to leucogranite) and biotite-rich, cordierite-bearing granitoids (CPGs, unusually tonalite, granodiorite and monzogranite). It has been suggested that the MPGs are produced through ‘wet’ anatexis of crustal rocks and crystal fractionation of the magmas, while the CPGs are produced through ‘dry’ anatexis of crustal rocks enhanced by underplating or injection of hot mantle-derived magmas. According to the rock types and mineral assemblages, both the MPGs and CPGs occur in the study area.

For example, the late-stage intrusion (two-mica monzogranite) of the Dayunshan-Mufushan batholith, and the Wangxiang batholith are dominantly MPGs, while the early-stage intrusion (biotite monzogranite and biotite granodiorite) of the Dayunshan-Mufushan batholith is similar to CPGs but without cordierite.

5.3. Origin of the granitoids

Previous studies on experimental petrology and geochemistry have documented that peraluminous granitoids (including strongly peraluminous or S-type granites) are mainly generated by partial melting of metasedimentary rocks (metapelite, metagreywacke), possibly added with metaigneous rocks (orthogneiss, amphibolite) (e.g., Miller, 1985; Patiño-Douce and Johnston, 1991; Sylvester, 1998). As demonstrated by Sylvester (1998), strongly peraluminous granites can inherit different CaO/Na₂O ratios by partial melting of different sources, pelite-derived melts tend to have lower CaO/Na₂O ratios (<0.3) than psammite-derived counterparts. The majority of granitic rocks in the study area have high CaO/Na₂O ratios (mostly >0.3), indicating that they mainly derived from partial melting of psammitic rocks (Fig. 10a). On the Rb/Sr vs. Rb/Ba plot, these samples mainly fall into the plagioclase-rich, clay-poor sources field (Fig. 10b). Therefore, we suggest that the source rocks are mostly metagreywacke with minor metaigneous rocks, as shown on the molar Al₂O₃/(MgO+FeO*) vs. molar CaO/(MgO+FeO*) diagram (Altherr et al., 2000; Fig. 10c). However, minor mixture of the two end-members consisting of crust-derived melt and basaltic magmas cannot be ruled out, especially for the granodiorites with rather higher CaO/Na₂O, lower Rb/Sr and Rb/Ba ratios. Indeed, L.X. Wang et al. (2014) identified a small stock of mantle-derived diorite within the biotite monzogranite of the Dayunshan-Mufuhshan batholith. Underplating of basaltic magma acting as a heat source that triggered partial melting of the middle to lower crust and formation of the felsic magmas is very likely. Nevertheless, a direct input of mantle-derived magma in genesis of these granitoids appears to be insignificant since mafic enclaves are rarely observed in the field. In addition, the Al₂O₃, Fe₂O₃*, CaO, MgO and TiO₂ contents decrease with increasing SiO₂, indicative of fractionation of feldspar, biotite and Fe-Ti oxides. The negative Eu anomaly, depletion in Sr and Ba, and linear trends of Sr vs. Rb and Ba further support the removal of plagioclase and K-feldspar from the original magma (Figs. 8 and 11). Dykes or small stocks of garnet-bearing leucogranite are

particularly abundant in the Taohuashan pluton and Dayunshan-Mufushan batholith. These rocks represent highly-fractionated products, showing geochemical features similar to the Himalaya leucogranite (L.X. Wang et al., 2008, 2014). Thus fractional crystallization probably played an important role during the magmatic evolution, which seems likely more effective for generation of the MPGs.

Zircon as early crystallization mineral from the magma can retain its original isotopic composition, and thus was widely used to trace petrogenetic process (Griffin et al., 2002). Zircons ($n = 178$) with crystallization ages from the studied granitoids show variable $\epsilon_{\text{Hf}}(t)$ values ranging from -12.5 to -3.6 (Table 1; Fig. 6). The corresponding T_{DM}^{C} ages vary from 1.42 to 1.99 Ga, suggesting that the source rocks were chemically weathered Paleo- to Mesoproterozoic crustal materials of the SCB. As mentioned before, the exposed basement in the study area is predominantly the Lengjiayi Group, which was traditionally regarded as slightly metamorphosed Mesoproterozoic strata (BGMRHN, 1988; Gu et al., 2002; Xu et al., 2007). Recently, U–Pb dating of detrital zircons as well as tuff beds have demonstrated that the deposition of the Lengjiayi Group occurred during the early Neoproterozoic (ca. 860–820 Ma; W. Wang et al., 2010, 2012, 2013; Gao et al., 2011; Meng et al., 2013; X.L. Wang et al., 2014; Yan et al., 2015). In our dated samples, there are quite a few Neoproterozoic inherited cores (Table S1), overlapping the major age population of detrital zircons from the Lengjiayi Group. These relict zircons show $\epsilon_{\text{Hf}}(t)$ values mostly between -4.1 and 4.3 (Table S2), which also fall into the wide range of $\epsilon_{\text{Hf}}(t)$ values of the Neoproterozoic detrital zircons. Moreover, the Lengjiayi Group also contains older detrital zircons with Paleo- to Mesoproterozoic ages equal to the crustal Hf model ages of these studied granitoids. Accordingly, we infer that the Late Mesozoic peraluminous granitoids in the study area were derived from partial melting of metasedimentary basement analogous to the Neoproterozoic Lengjiayi Group. The slight variations of Hf isotopic compositions among the different plutons, even different samples from the same pluton (e.g., Liyunshan), indicate isotope disequilibrium of the magma probably inherited from heterogeneous metasedimentary source or due to wall-rock assimilation.

5.4. Geodynamic implications

Subduction of the Paleo-Pacific plate beneath the Eurasia continent is the most

popular model to account for the origin and evolution of the Jurassic-Cretaceous magmatism in SE China so far. For example, Zhou and Li (2000) proposed a comprehensive model of a combination of lithospheric subduction and underplating of mafic magmas, with an increasing subduction angle of the Paleo-Pacific plate to explain the oceanward younging trend of the magmatic activity. Alternatively, Li and Li (2007) presented a flat-slab subduction model to account for the Mesozoic migration of intracontinental orogeny and magmatism in SE China. They interpreted the Jurassic magmatism in the interior as anorogenic products formed in response to the upwelling of asthenosphere mantle after the break-up and foundering of an Early Mesozoic subducted flat-slab, whereas the Cretaceous magmatism with both extensional and arc-related features in the coastal area was regarded as a retreating arc system. Similarly, based on detailed compilations of available geochronological, geochemical and geological data on the Cretaceous plutonism in the coastal Fujian province, a recurrence of slab break-off and rollback of the Paleo-Pacific plate has been suggested (Z. Li et al., 2014). More recently, Jiang et al. (2015) put forward a repeated slab advance-retreat model in view of the identification of several Mesozoic NE-trending A-type granite belts and the Early Jurassic mafic rocks in SE China. In addition, a ridge subduction model has been proposed for genesis of the Cretaceous magmatism (especially “adakitic rocks” and A-type granites) and related polymetallic mineralization in the lower Yangtze River belt (Ling et al., 2009; Sun et al., 2010; H. Li et al., 2012). However, the so-called “adakitic rocks” penetrated perpendicularly into the continental interior more than 1000 km away from the present continental margin, such a spatial distribution would require an unlikely ridge subduction with very shallow angle (X.H. Li et al., 2010, 2013b). Moreover, if there was a slab window formed by ridge subduction, the related magmatism would show a younging trend toward the continental interior, which is contrary to the present observation. Considering the subduction angle variations, Wu et al. (2012) argued that slab tearing instead of ridge subduction might account for the eastward younging trend of the 150-136 Ma magmatism in this area. Nevertheless, it is beyond the scope of this paper to reconstruct the history of the Paleo-Pacific plate, since available data are rare (Engebretson et al., 1985; Maruyama et al., 1997; Sun et al., 2007).

The Jurassic-Cretaceous magmatism in SE China took place in two distinct stages. The Jurassic magmatism includes an subordinate phase of ca. 190 Ma A-type granites, mafic intrusions and bimodal volcanic rocks mainly in the E-W trending Nanling Range

(Chen et al., 2008; He et al., 2010; Jiang et al., 2015), and a major magmatic flare-up at ca. 165–150 Ma throughout the Cathaysia interior dominated by I-type or fractionated I-type granites along with minor A-type granites and syenites (Li et al., 2007; Jiang et al., 2009; Huang et al., 2013, 2015; Zhang et al., 2015; Wang et al., 2016). The Cretaceous intrusive and volcanic rocks are confined mostly to the coastal region, with two major phases at ca. 140–125 Ma and ca. 115–90 Ma. They consist predominantly of calc-alkaline granites and felsic volcanic rocks, and subordinate A-type granites and basaltic rocks (Li, 2000; Wong et al., 2009; Jiang et al., 2011; Guo et al., 2012; Liu et al., 2012, 2014; Z. Li et al., 2014). In view of geochronological statistics, ca. 150–140 Ma is a relatively “magmatic quiescence” in SE China (X.H. Li et al., 2010; Y.J. Wang et al., 2013). As discussed before, the Late Mesozoic magmatic activities in the northeastern Hunan province can be roughly subdivided into three phases, namely 151–146 Ma, 132–127 Ma and ca. 117 Ma. The latter two phases of magmatism are in agreement with the Cretaceous magmatism in SE China. Most of the researchers favor slab rollback of the Paleo-Pacific plate as the geodynamic trigger for the Cretaceous magmatism in SE China (e.g., Zhou and Li, 2000; Li and Li, 2007; Wong et al., 2009; Jiang et al., 2011, 2015; Liu et al., 2012, 2014; Wang et al., 2016). Such a setting is evidently supported by the Cretaceous lithospheric extension in the eastern SCB, as indicated by the development of extensional domes and rift basins (Gilder et al., 1996; Li, 2000; Lin et al., 2000; Shu et al., 2009; Ji et al., 2014; J.H. Li et al., 2014). In the study area, the Dayunshan syntectonic pluton was emplaced under a NW-SE extension according to our unpublished structural and AMS (anisotropy of magnetic susceptibility) data. Moreover, a NW-directed back-thrusting event along the Changle-Nan'ao belt and associated metamorphism occurred between 130 and 110 Ma, as a response to collision of the West Philippines microcontinent with the SCB (Li et al., 2015; Wei et al., 2015). Thus we infer that there might be two episodes of slab rollback during the Cretaceous, respectively corresponding to the two main phases of Cretaceous magmatism (Fig. 12A). Particularly, the early-stage (151–146 Ma) magmatism identified in this study, which overlapped the epilogue of Jurassic magmatic flare-up and the subsequent “magmatic quiescence”, is rarely developed in SE China. Nevertheless, a minor amount of 150–140 Ma igneous rocks can also be found in the lower Yangtze River belt and the Changle-Nan'ao belt (e.g., X.H. Li et al., 2010; Cui et al., 2012; Liu et al., 2012; Wu et al., 2012). Such a configuration agrees with the broad outward-younging trend for the Jurassic-Cretaceous magmatism as

suggested by Li et al. (2013a). Thus we support the interpretation that the Jurassic magmatic flare-up in SE China was mainly induced by break-up and foundering of a subducted oceanic flat-slab in an intraplate setting, and the apparent 150–140 Ma magmatic gap probably represented a transition period prior to slab rollback of the Paleo-Pacific plate (Fig. 12A).

The effect of Paleo-Pacific subduction on the Mesozoic geological evolution of SE China is generally acknowledged, and a northwestward polarity is usually assumed (e.g., Zhou et al., 2000; Li and Li, 2007; Jiang et al., 2015). However, when and how far the Paleo-Pacific plate subducted is still highly speculative. It is worth nothing that the tectonic nature and polarity of the Jurassic flat-slab beneath SE China remain an open question. In the present state of knowledge, a continental slab with southeastward polarity cannot be ruled out. Indeed, southward or southeastward continental subduction accounts well for the geodynamic scenarios of the Triassic orogeny along the southern margin of the SCB (from Yunnan province to Hainan Island, possibly farther to the east; see Faure et al., 2014, 2016), and especially in interior of the SCB along the NE-SW trending Jiuling-Xuefengshan belt. (Chu et al., 2012a, 2012b; Chu and Lin, 2014). The pre-existing lithospheric weak zone in the central SCB (e.g., the Jiangshan-Shaoxing fault and the Chenzhou-Linwu fault) was reactivated by the Triassic intracontinental orogeny. At variance to the Paleo-Pacific subduction models, the Jurassic magmatism in SE China might be a protracted consequence of slab foundering after a SE-directed intracontinental subduction (Fig. 12B). However, future work is required to test this new geodynamic interpretation.

6. Conclusions

A comprehensive study of zircon U–Pb dating, Hf isotope and geochemical data of the Late Mesozoic granitoids in the northeastern Hunan province allows us to draw the following conclusions:

(1) The Late Mesozoic magmatic activities in this area occurred at 151–146 Ma, 132–127 Ma and ca. 117 Ma;

(2) These high-K calc-alkaline, weakly to strongly peraluminous granitoids are mainly represented by two-mica or biotite monzogranites with minor biotite granodiorites, showing geochemical features similar to S-type as well as fractionated S-type granites;

(3) The source rocks of these granitic rocks were mainly metasedimentary rocks predominantly with psammitic component, without significant involvement of mantle material. Fractional crystallization played an important role in the magma evolution especially for the dominant monzogranites;

(4) The Cretaceous (132–127 Ma and ca. 117 Ma) magmatism might be response to episodic slab rollback of the Paleo-Pacific plate; while the early-stage (151–146 Ma) magmatism probably foreshadowed the transformation from foundering of a subducted flat-slab to slab rollback.

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Appendix A. Supplementary data

Supplementary data related to this article can be found online.

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Figure captions

Fig. 1. Schematic map showing distribution of the Jurassic-Cretaceous magmatic rocks in SE China (modified after Zhou et al., 2006) and location of the study area. The dashed line indicates the Jiangshan-Shaoxing fault at the northeastern segment (representing the reworked Neoproterozoic suture zone) and the Chenzhou-Linwu fault at the southwestern segment (probably as a scar of the Triassic intracontinental orogeny). Inset shows the tectonic sketch of the SCB, with the Jiuling-Xuefengshan belt in the central part.

Fig. 2. Geological map of the northeastern Hunan province and adjacent areas showing the main Late Jurassic to Early Cretaceous granitic plutons including those studied in this paper.

Fig. 3. Microphotographs of the representative granitic samples. (a) Plagioclase with compositional zoning in biotite monzogranite (sample JH155). (b) Euhedral to subhedral muscovite, along with biotite and K-feldspar in two-mica monzogranite (sample DYS01). (c) Secondary muscovite replacing biotite and plagioclase in biotite

monzogranite, note also myrmekite developed in feldspar (sample QD69). (d) Monazite enclosed within biotite in biotite monzogranite (sample QD69). (e) Biotite mostly altered to chlorite in granodiorite (sample DYS14). (f) Magmatic epidote including allanite in granodiorite (sample DYS14). Mineral abbreviations: Bt, biotite; Mus, muscovite; Kfs, K-feldspar; Pl, plagioclase; Qtz, quartz; Mnz, monazite; Ep, epidote; Aln, allanite; Myr, myrmekite.

Fig. 4. Cathodoluminescence images of representative zircons from the granitic samples. White ellipses indicate the analytical spots with measured ages.

Fig. 5. Concordia diagrams of the analyzed zircons, sample locations are shown in Figure 2.

Fig. 6. Histograms of the zircon $\epsilon_{\text{Hf}}(t)$ values, n is the number of analyzed grains.

Fig. 7. (a) K_2O vs. SiO_2 diagram (after Peccerillo and Taylor, 1976); (b) A/NK vs. A/CNK diagram: A/NK = molar $\text{Al}_2\text{O}_3/(\text{Na}_2\text{O}+\text{K}_2\text{O})$, A/CNK (alumina saturation index, ASI) = molar $\text{Al}_2\text{O}_3/(\text{CaO}+\text{Na}_2\text{O}+\text{K}_2\text{O})$ (after Maniar and Piccoli, 1989).

Fig. 8. Chondrite-normalized REE patterns (left column), and primitive mantle-normalized trace element spidergrams (right column). The normalization values are from McDonough and Sun (1989).

Fig. 9. (a) ACF diagram: A = $\text{Al}_2\text{O}_3-\text{Na}_2\text{O}-\text{K}_2\text{O}$, C = CaO, F = $\text{MgO}+\text{FeO}^*$ (after Chappell and White, 1992); (b) variation of SiO_2 vs. P_2O_5 , (c) and (d) plots of Y and Th against Rb (for the I- and S-type granites trends, cf. Chappell, 1999); (e) and (f) classification diagrams of $10000 \times \text{Ga}/\text{Al}$ vs. FeO^*/MgO and $\text{Zr}+\text{Nb}+\text{Ce}+\text{Y}$ vs. $(\text{K}_2\text{O}+\text{Na}_2\text{O})/\text{CaO}$, respectively (after Whalen et al., 1987). Symbols are the same as those in Figure 7.

Fig. 10. (a) $\text{Al}_2\text{O}_3/\text{TiO}_2$ vs. $\text{CaO}/\text{Na}_2\text{O}$ and (b) Rb/Sr vs. Rb/Ba (after Sylvester, 1998); (c) molar $\text{CaO}/(\text{MgO}+\text{FeO}^*)$ vs. molar $\text{Al}_2\text{O}_3/(\text{MgO}+\text{FeO}^*)$ (after Altherr et al., 2000). Symbols are the same as those in Figure 7.

Fig. 11. Sr vs. Ba (a) and Rb (b) plots. Labeled vectors correspond to up to 50% fractionation crystallization of the main rock-forming minerals (after Janoušek et al.,

2004). Symbols are the same as those in Figure 7.

Fig. 12. Tectonic model for the Mesozoic tectono-magmatic evolution in SE China. (A) Northwestward subduction of the Paleo-Pacific plate (modified after Li and Li, 2007; Li et al., 2007); (B) Jurassic slab foundering after a SE-directed intracontinental subduction (modified after Faure et al., 2016).

Table captions

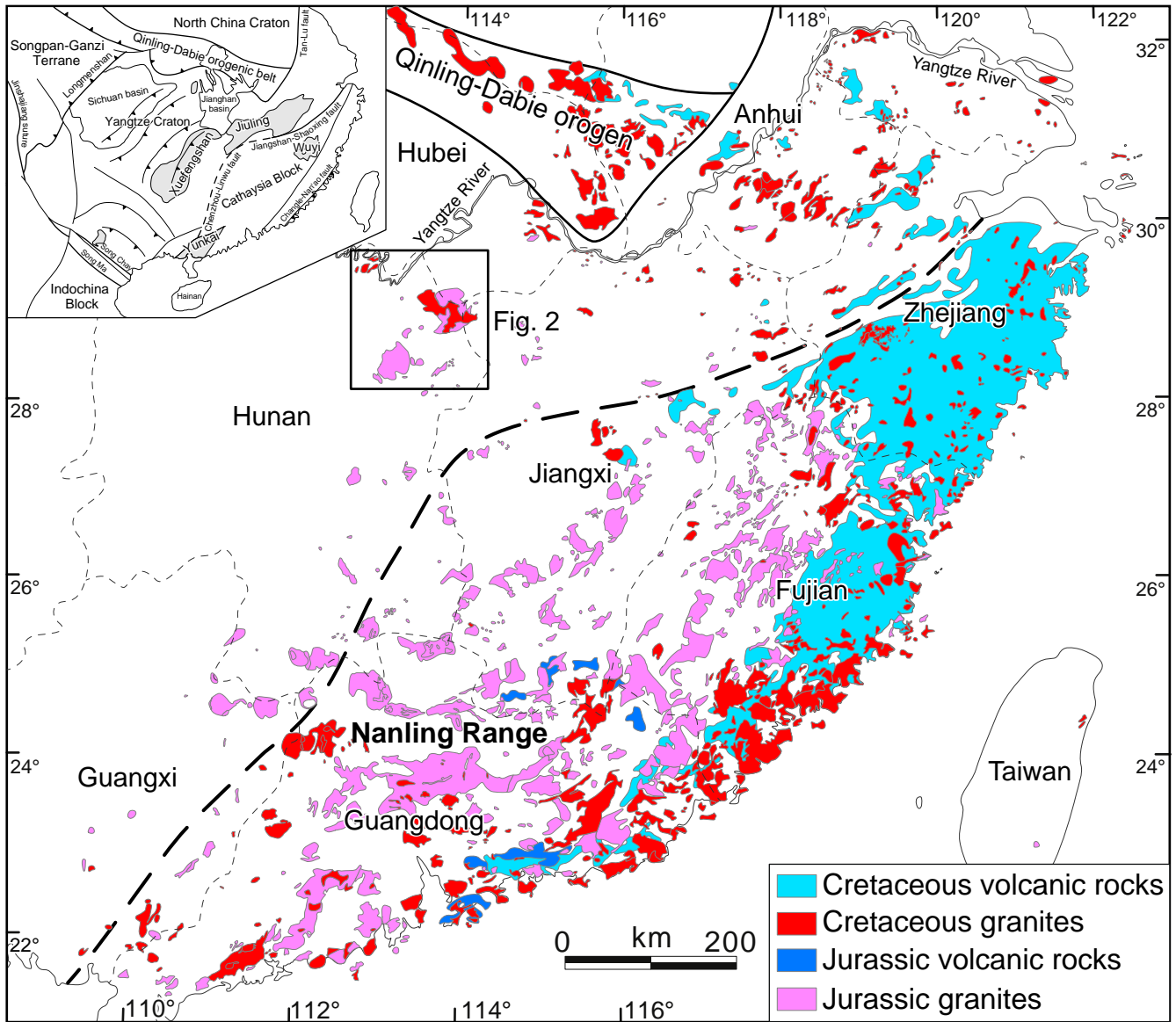
Table 1 Summary of zircon U–Pb ages and Hf isotopic data for the granitic samples

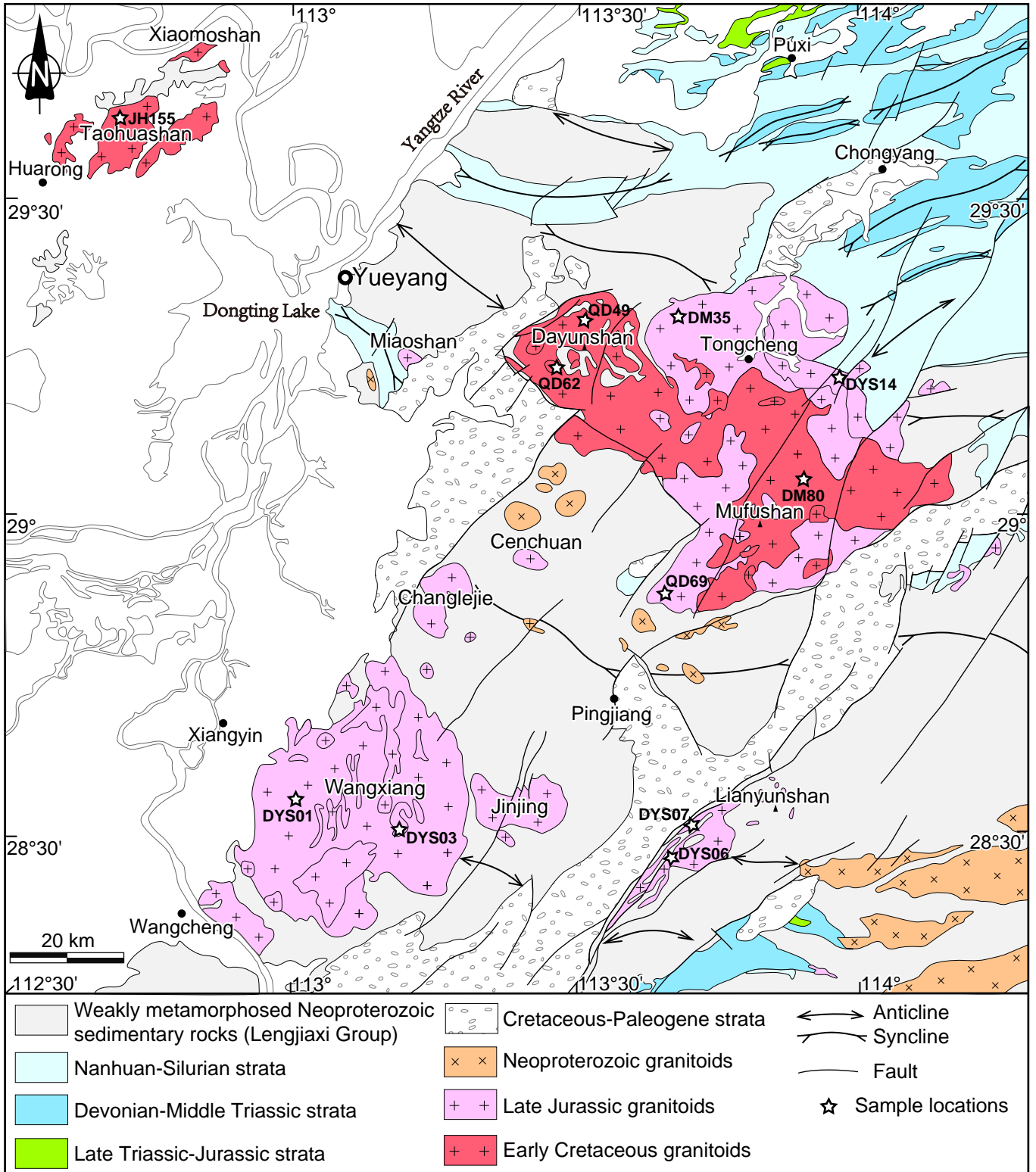
Table 2 Major (wt.%) and trace element (ppm) compositions of the granitic samples

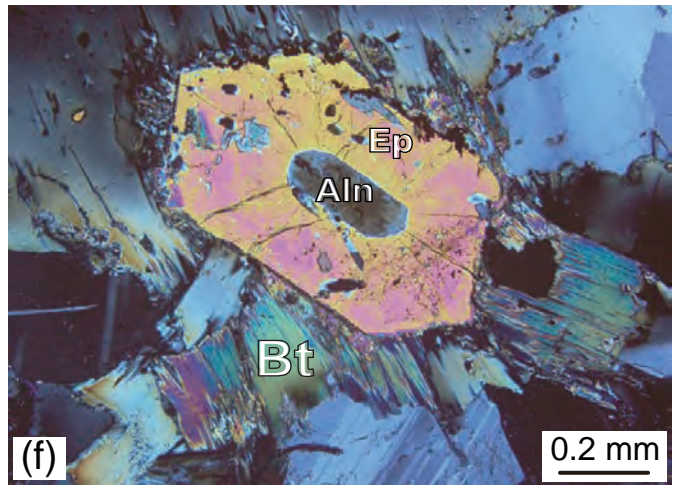
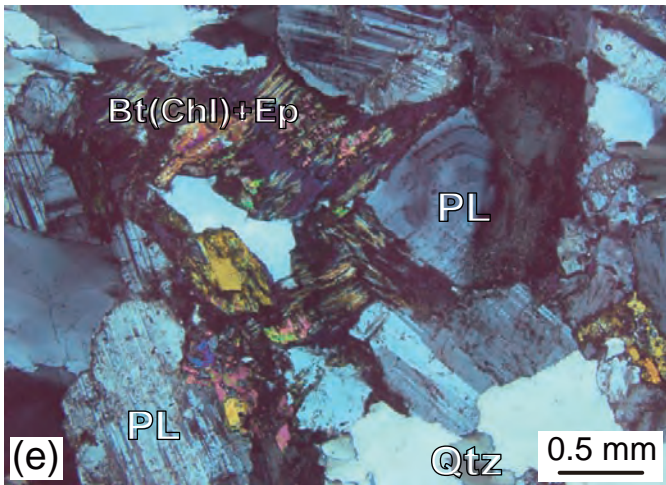
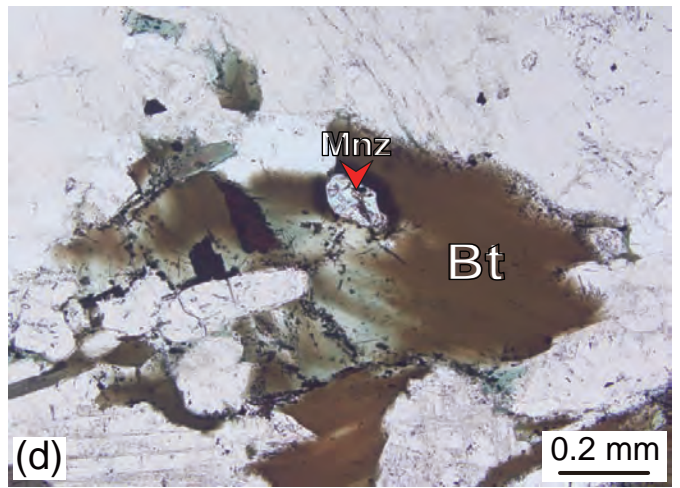
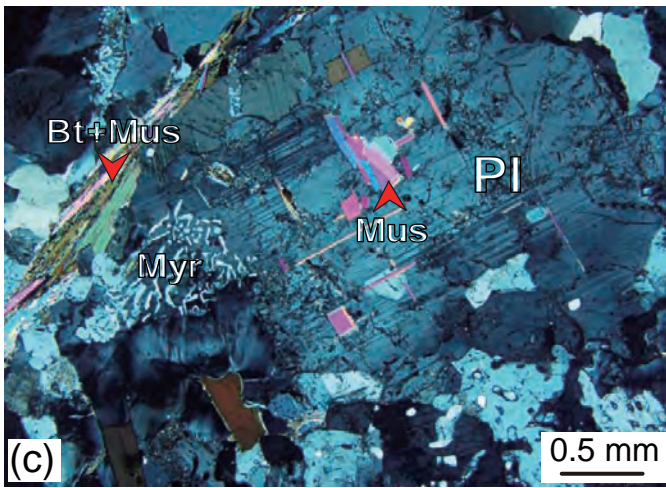
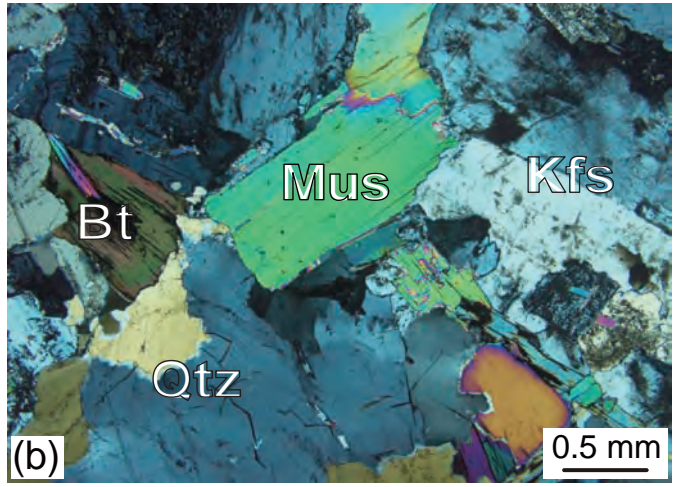
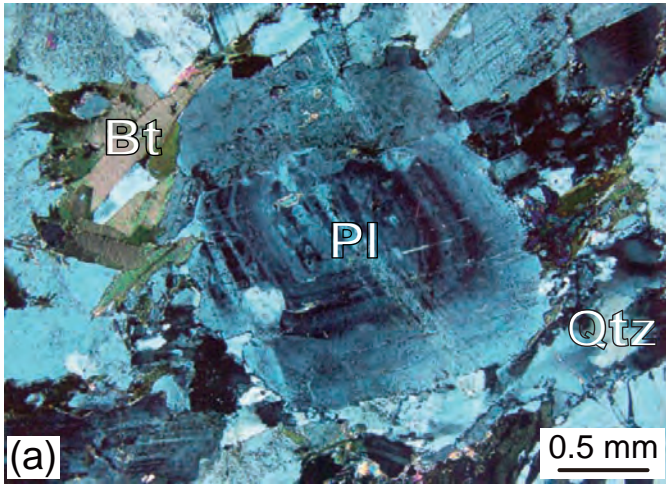
Supplementary data

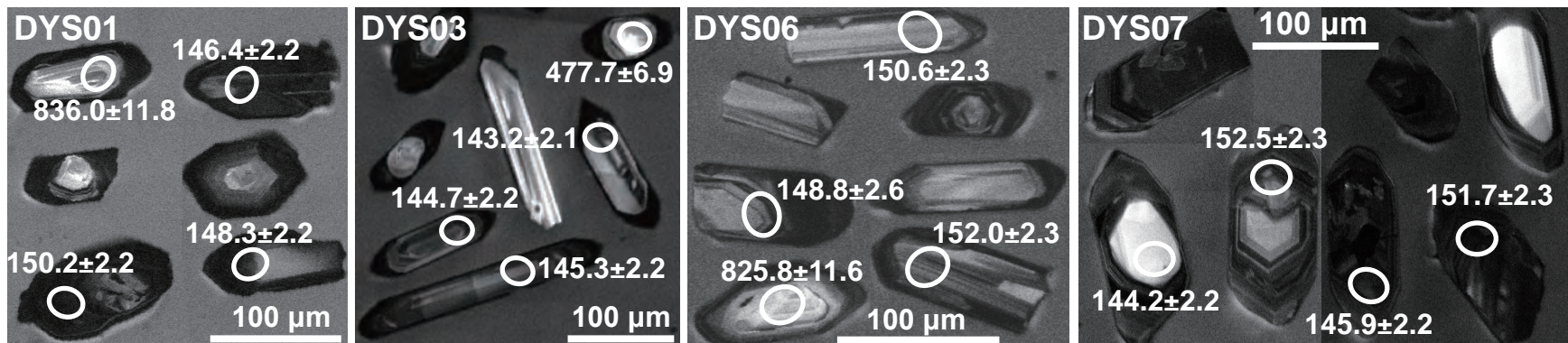
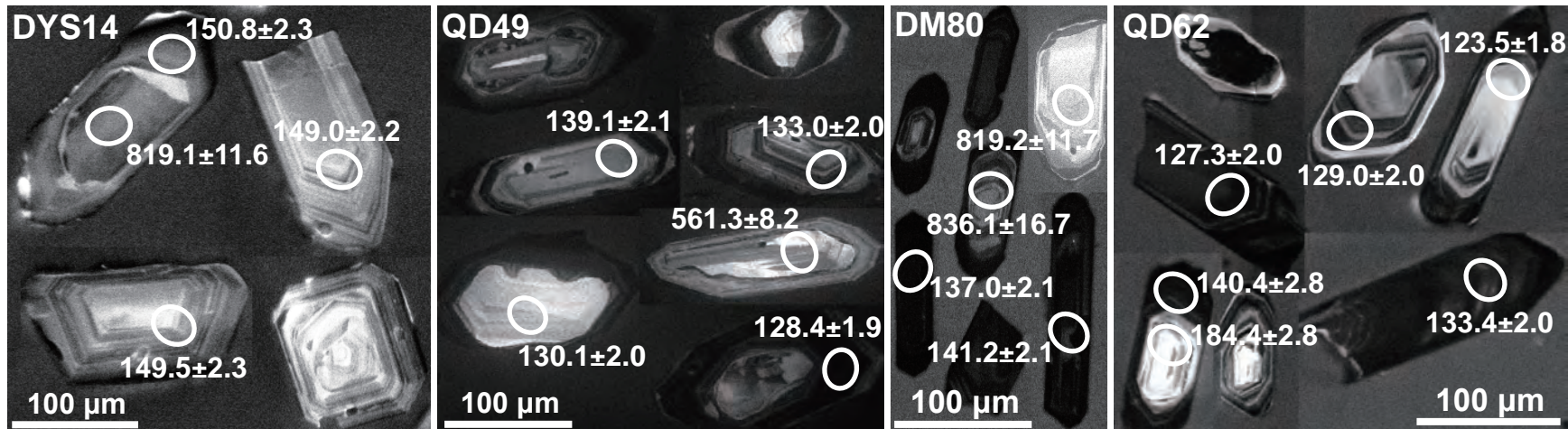
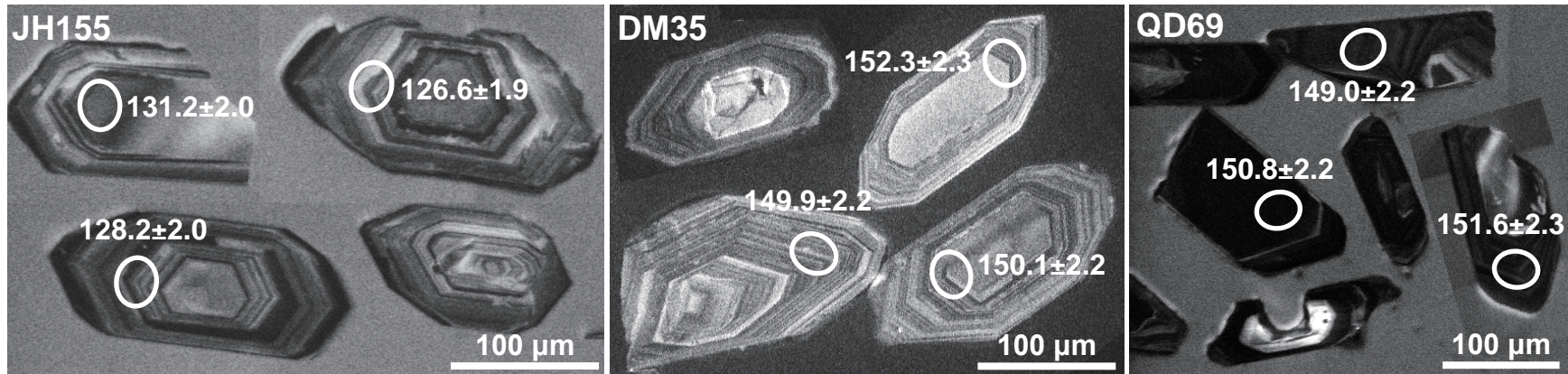
Table S1 SIMS zircon U–Pb data of the granitic samples

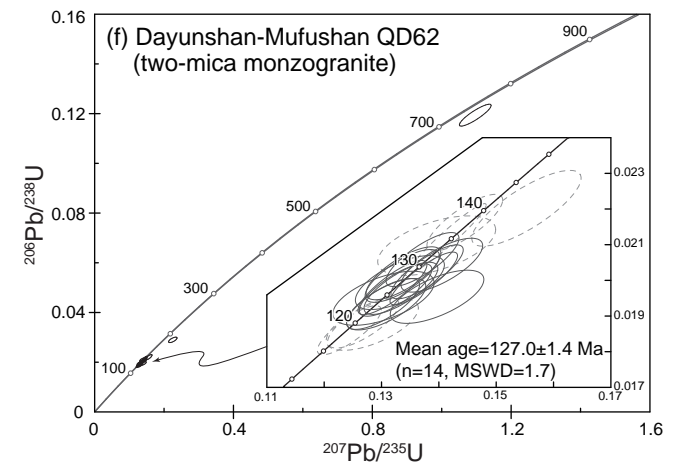
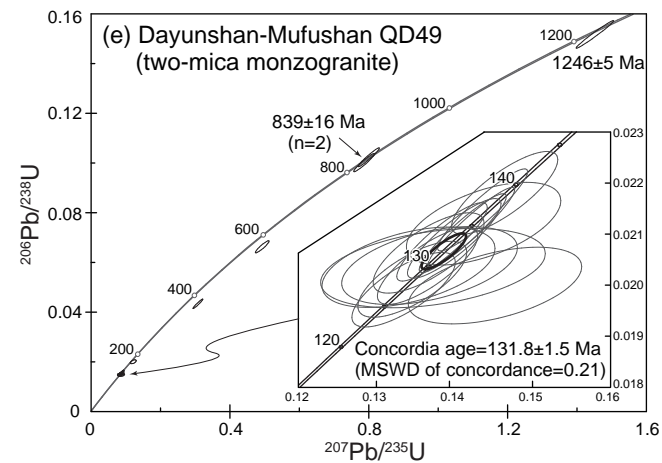
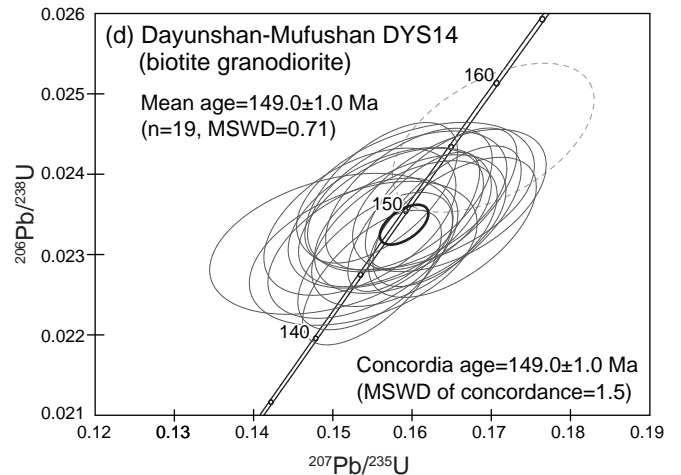
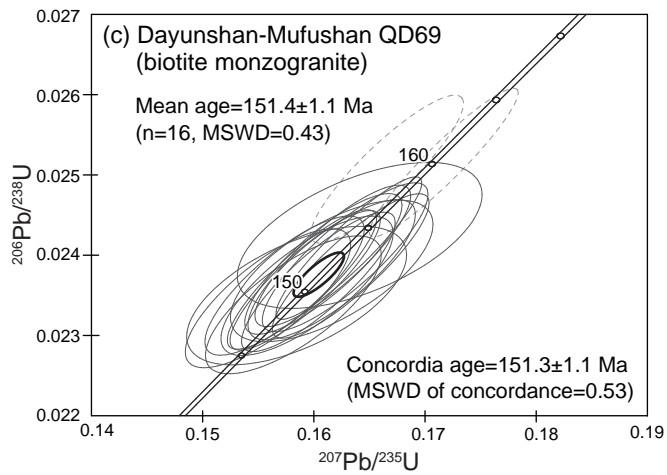
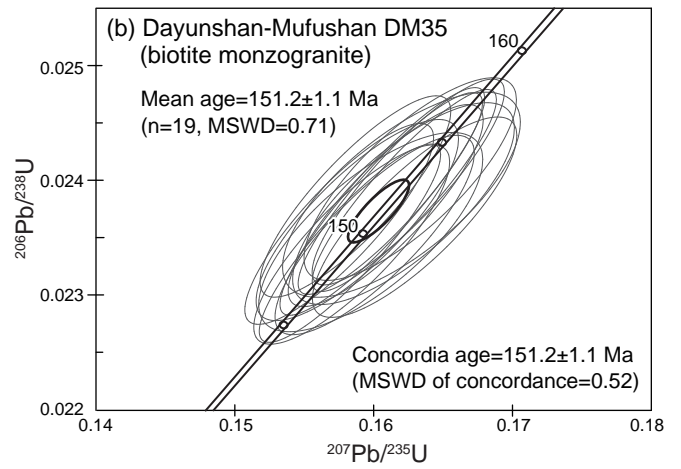
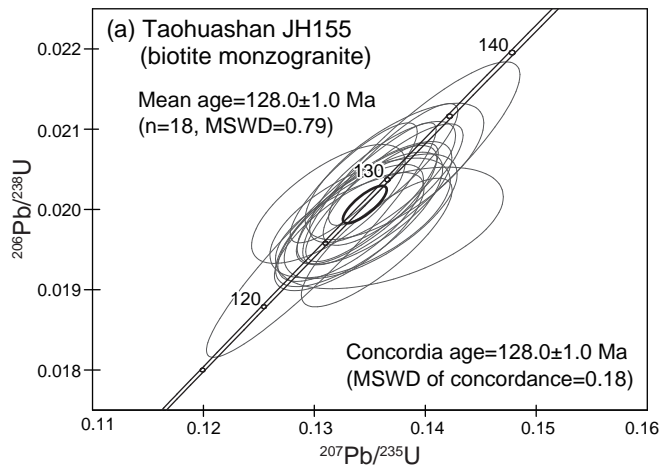
Table S2 Zircon Hf isotopic data of the granitic samples

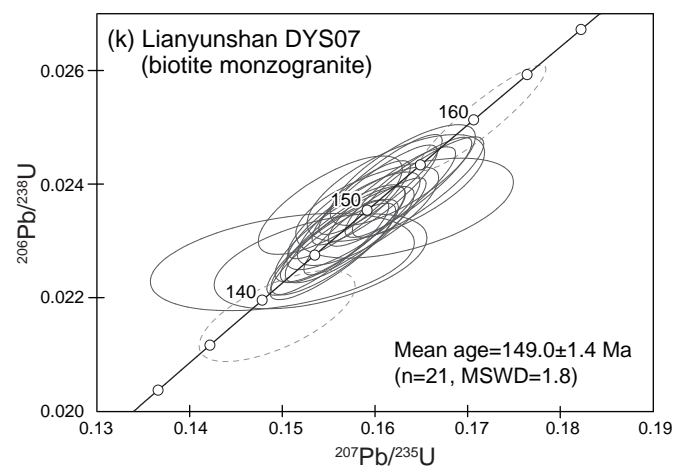
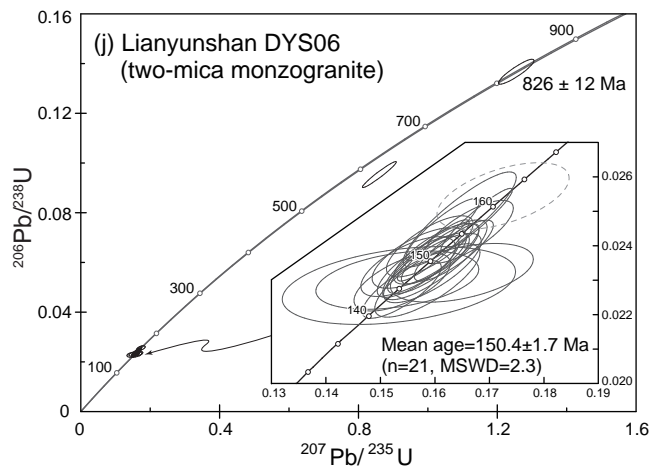
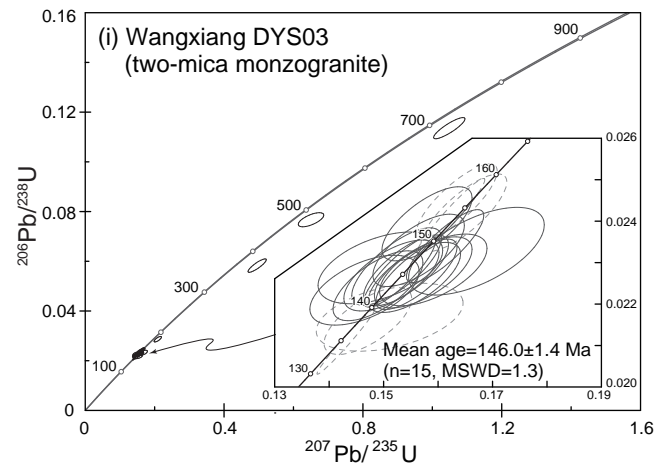
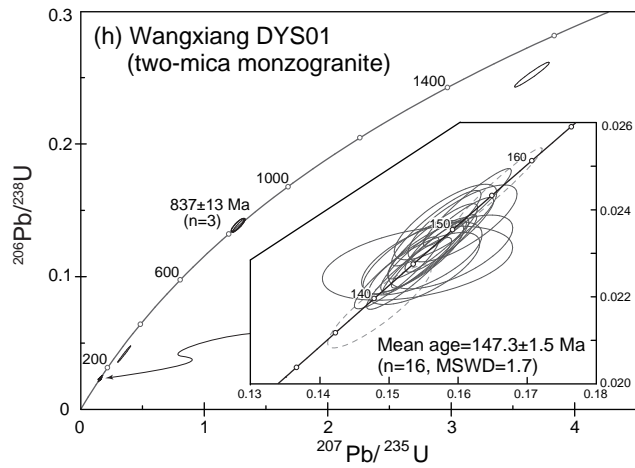
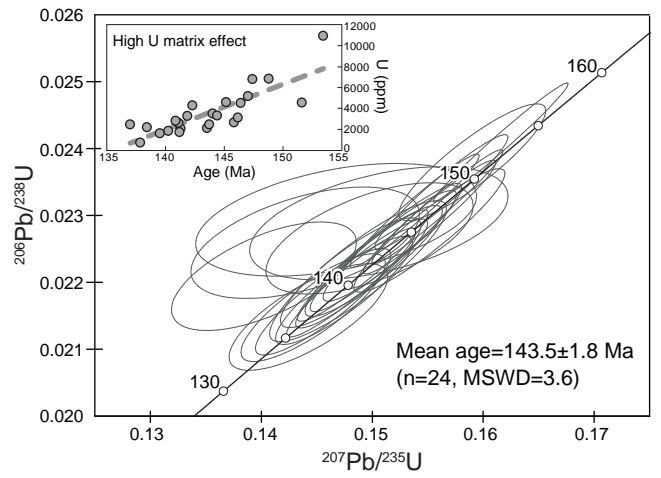
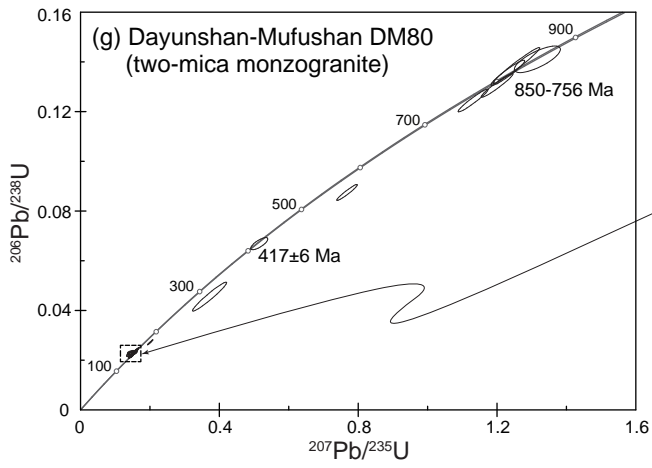


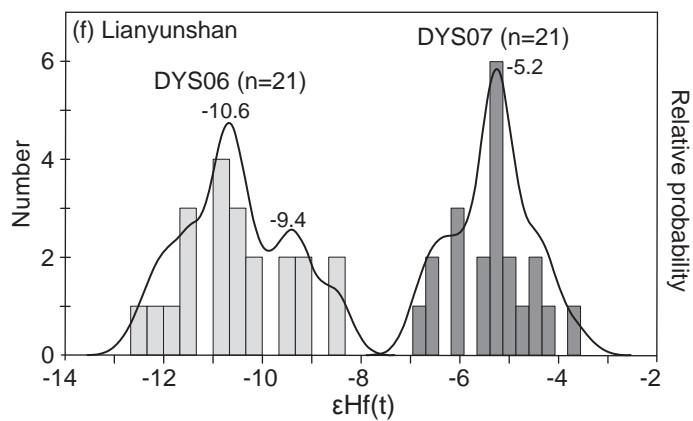
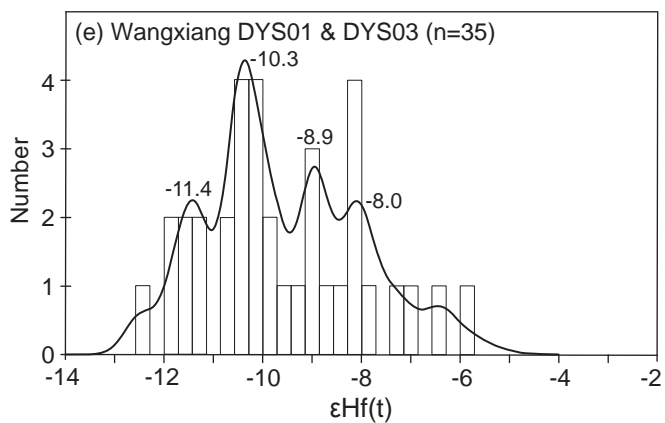
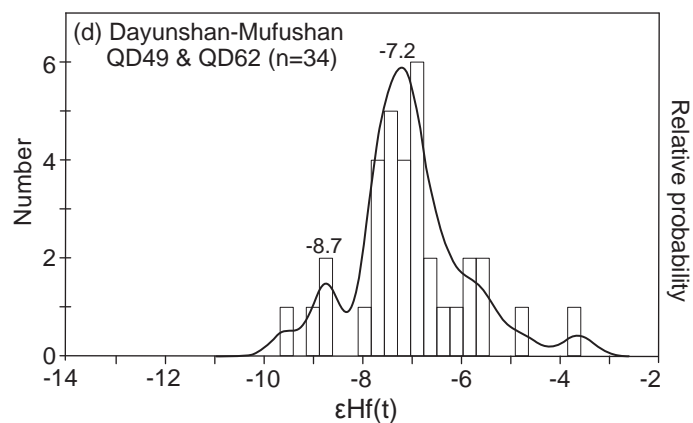
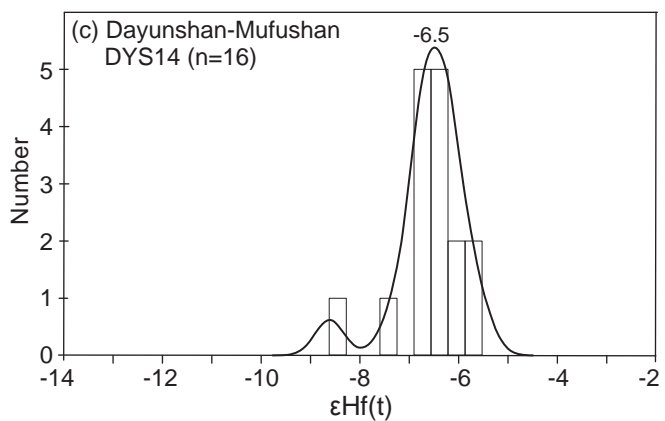
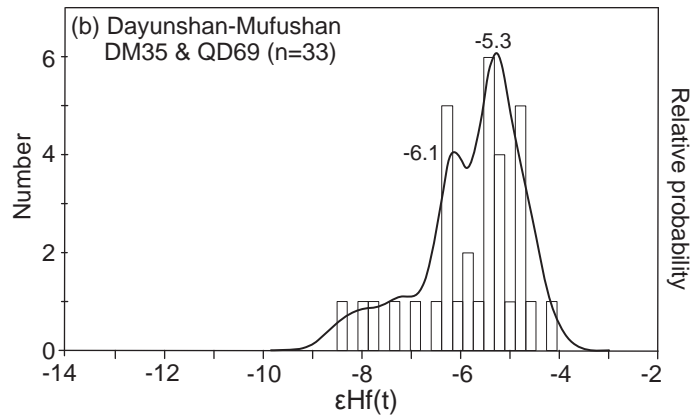
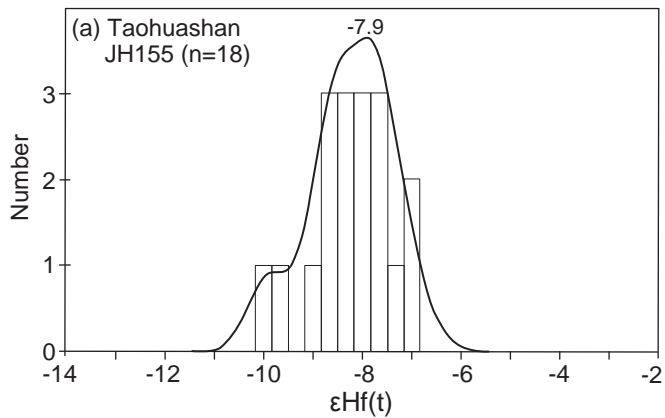


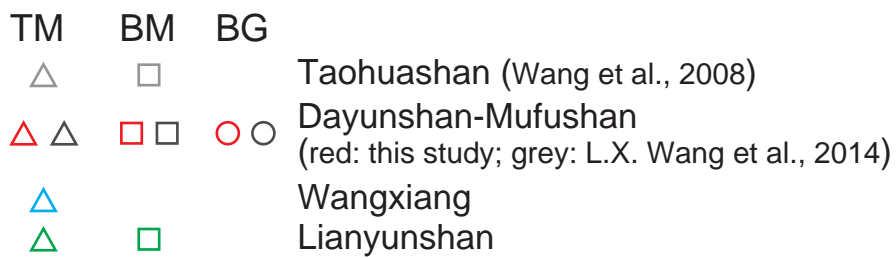
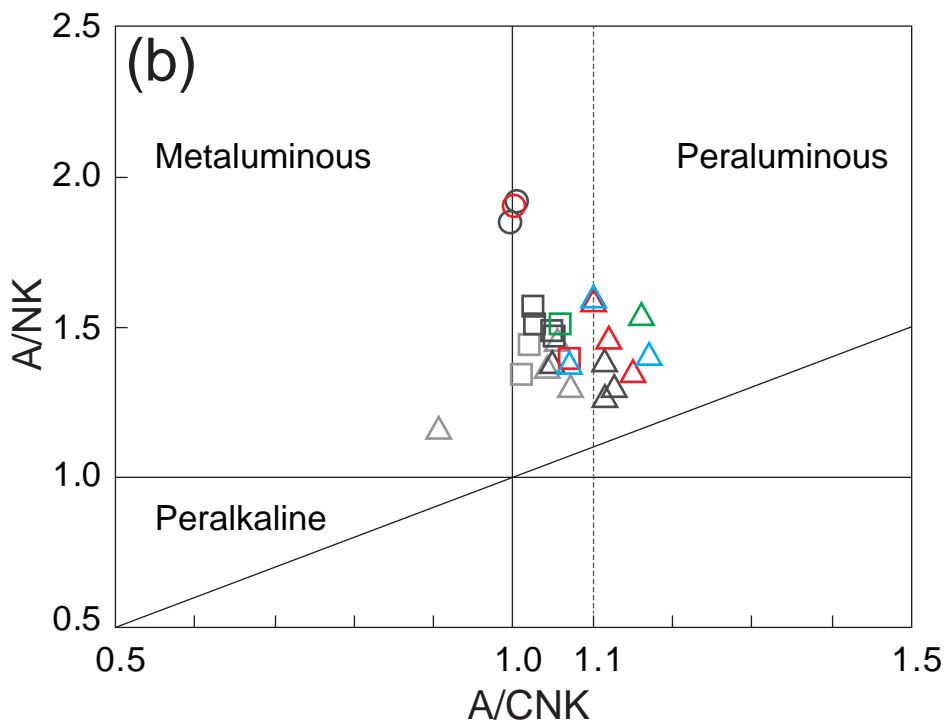
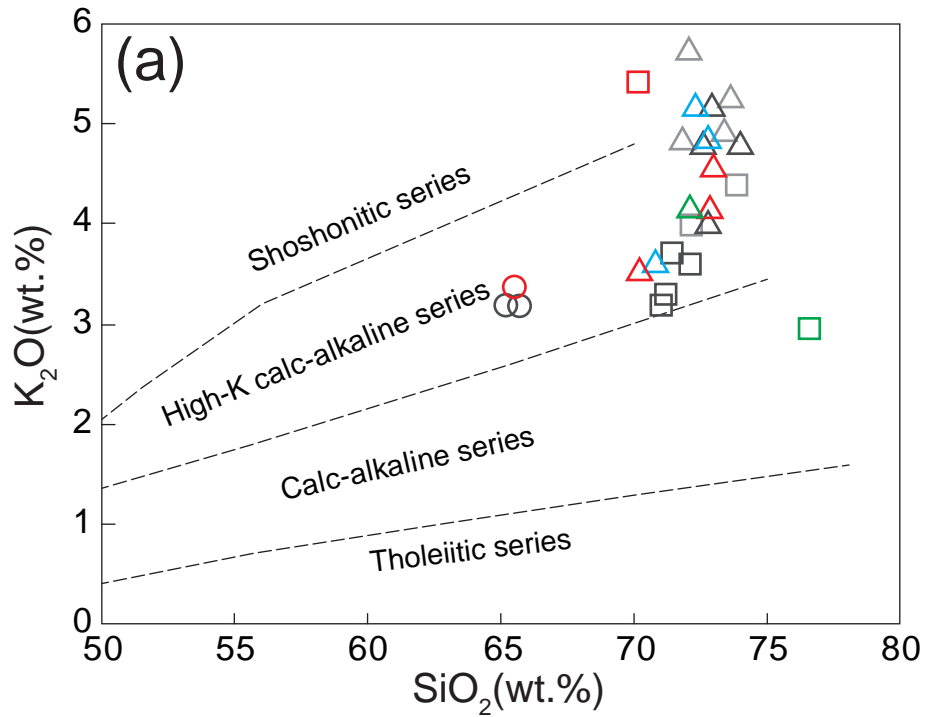




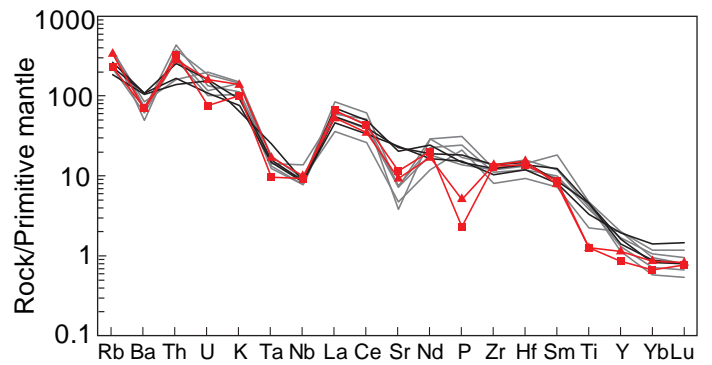
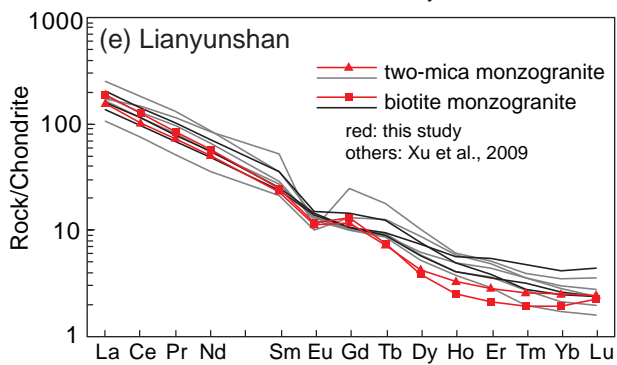
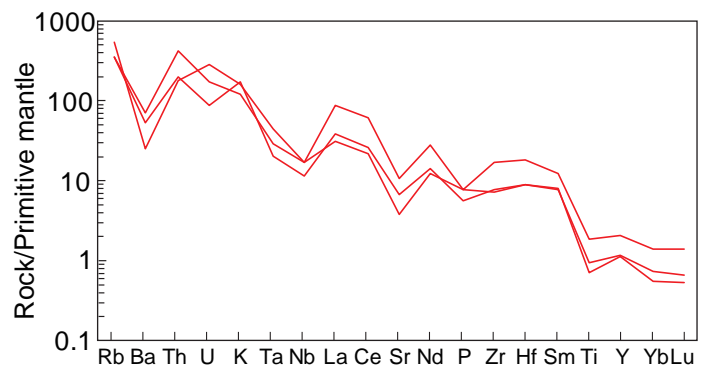
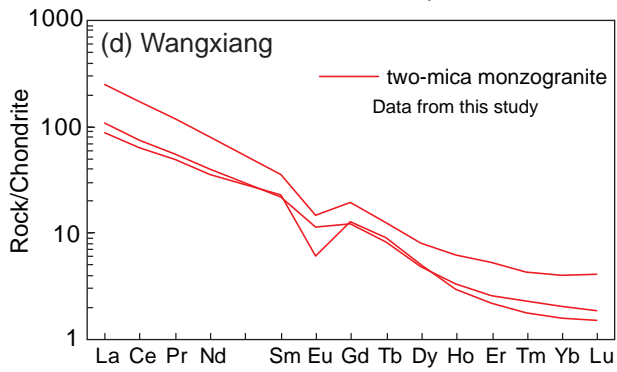
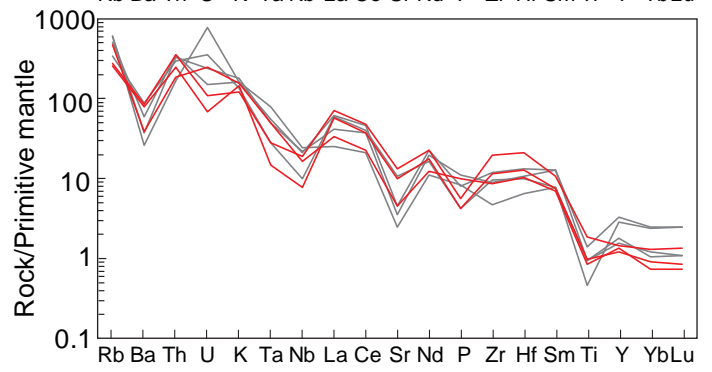
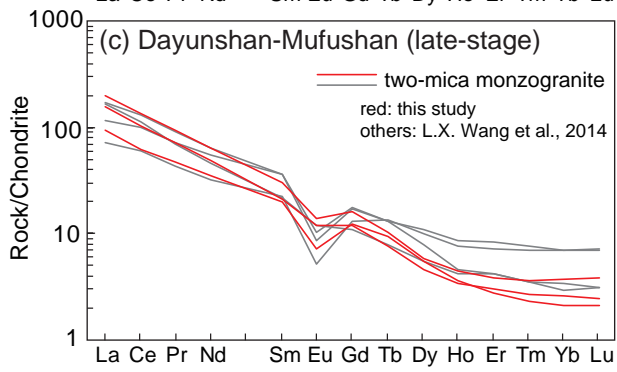
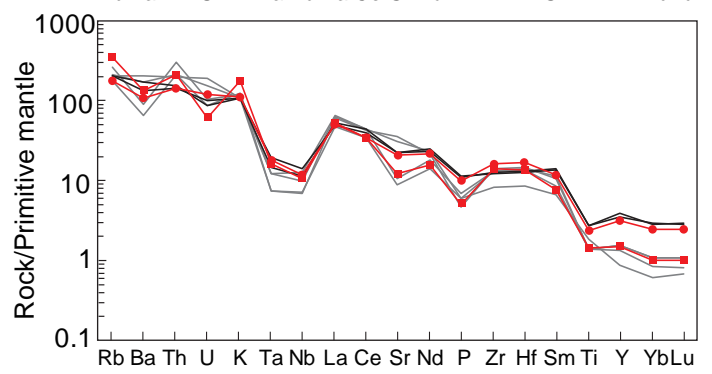
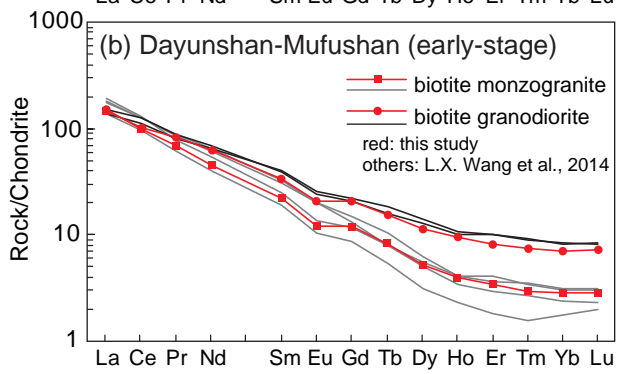
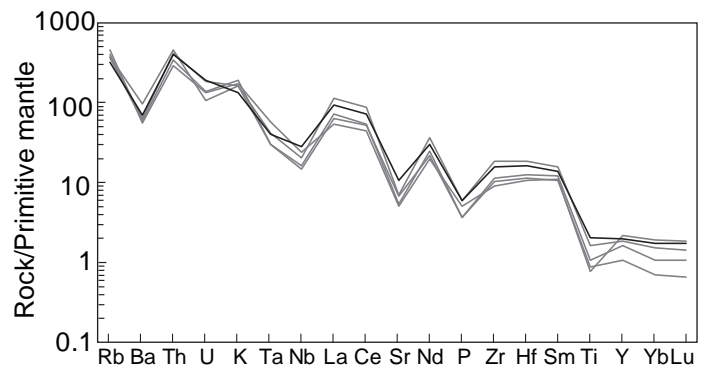
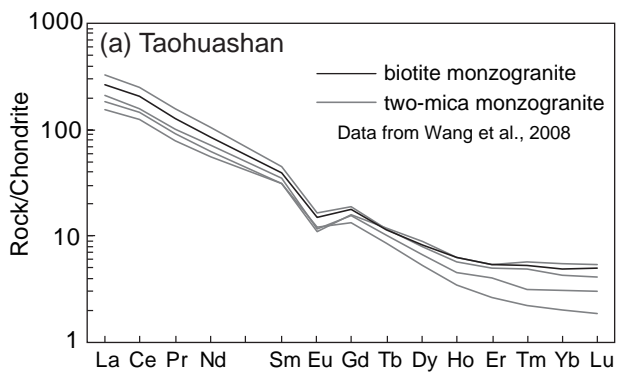


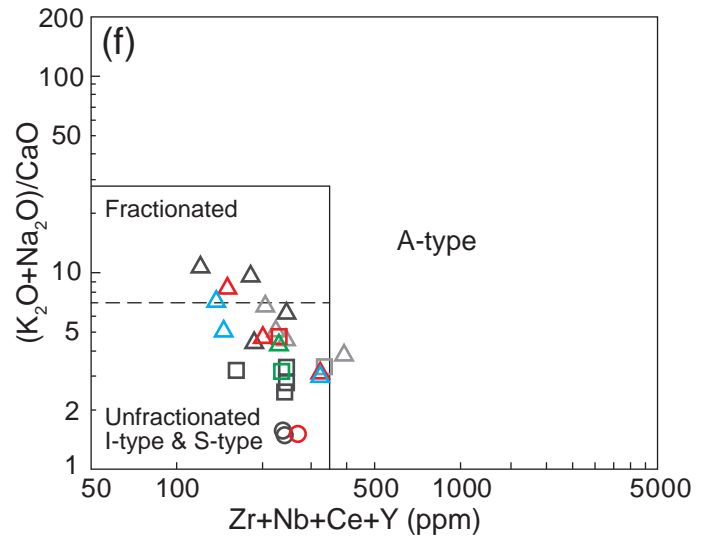
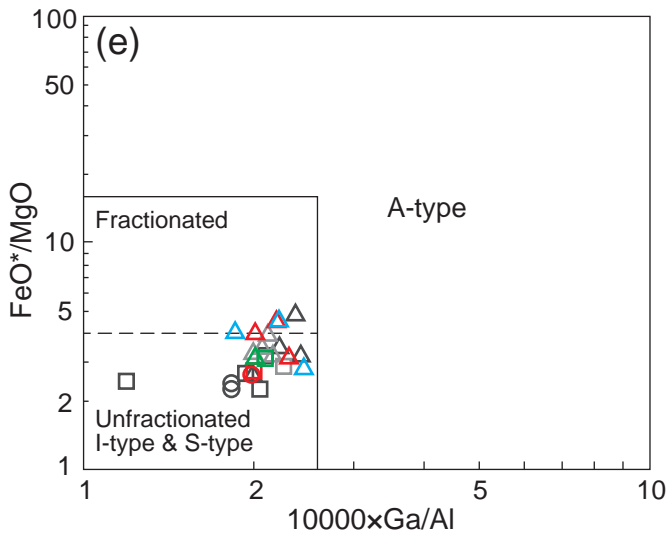
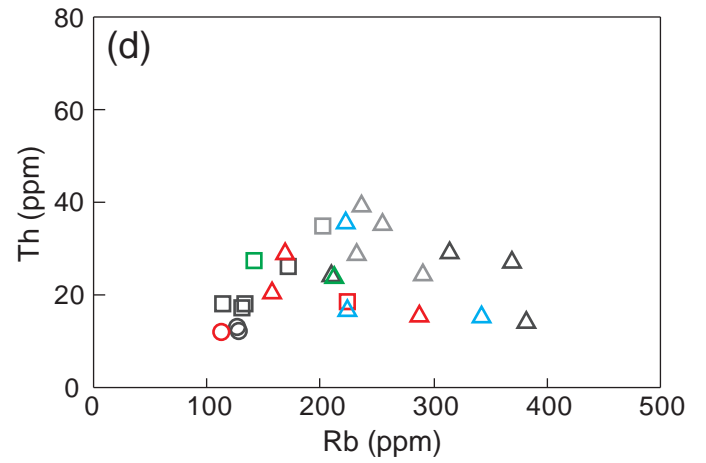
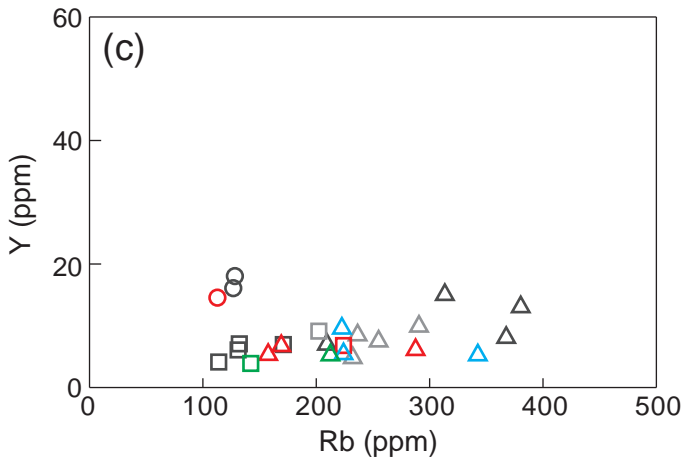
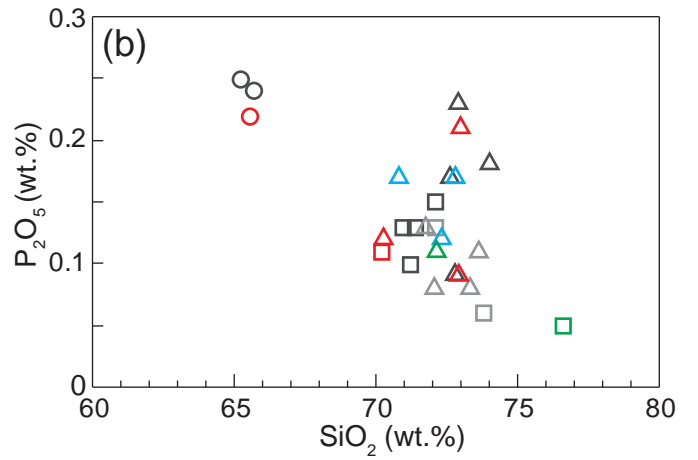
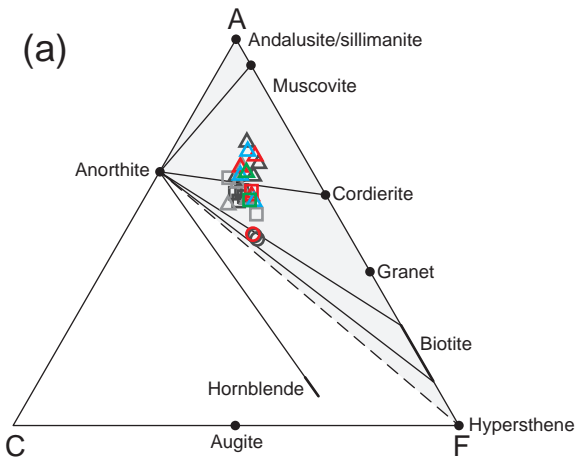


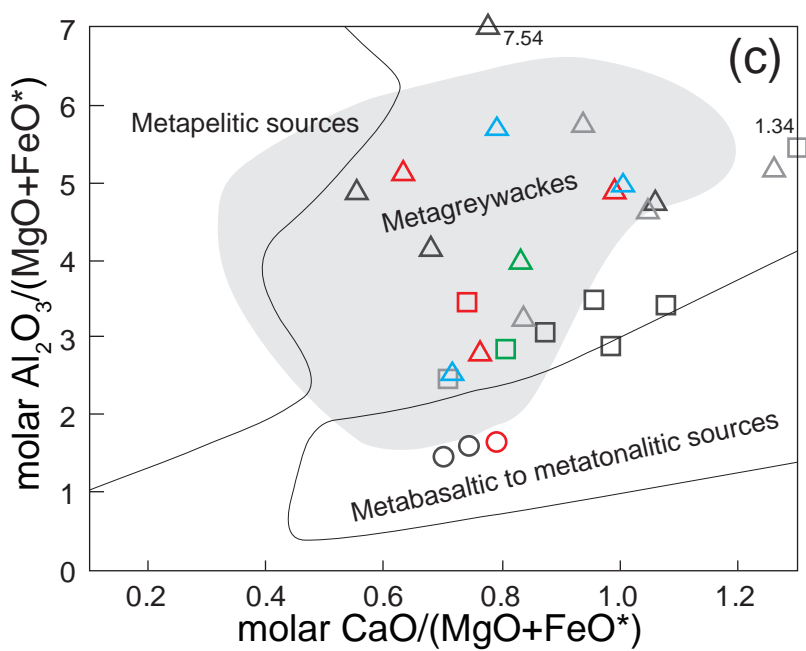
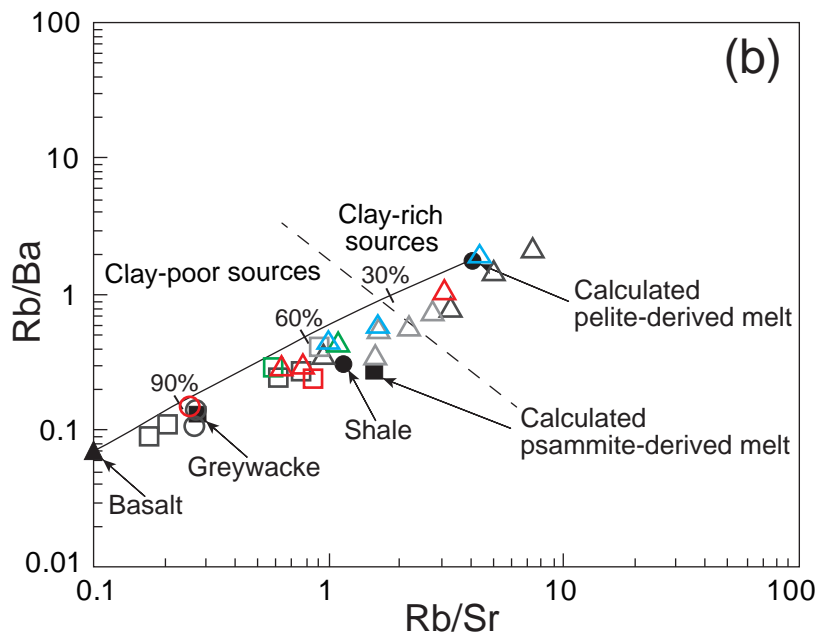
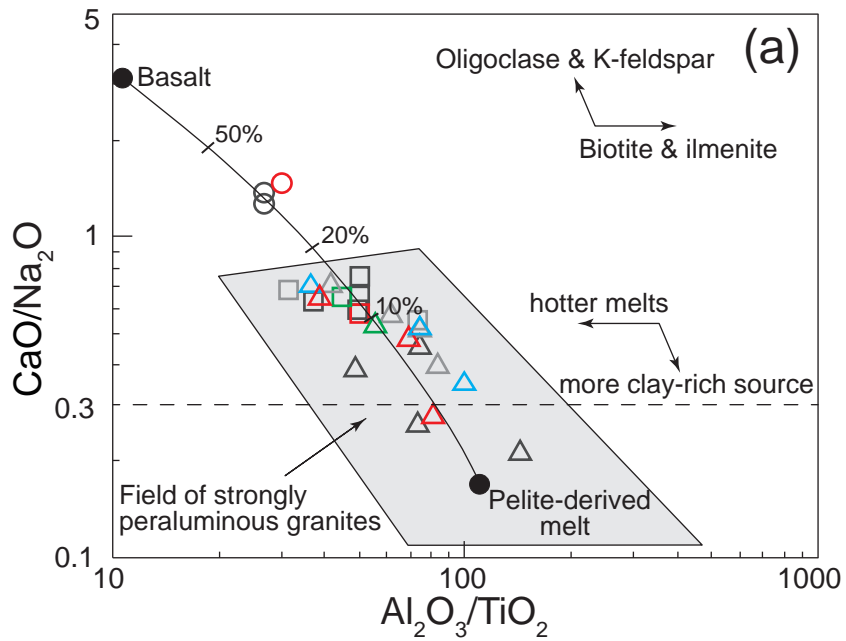


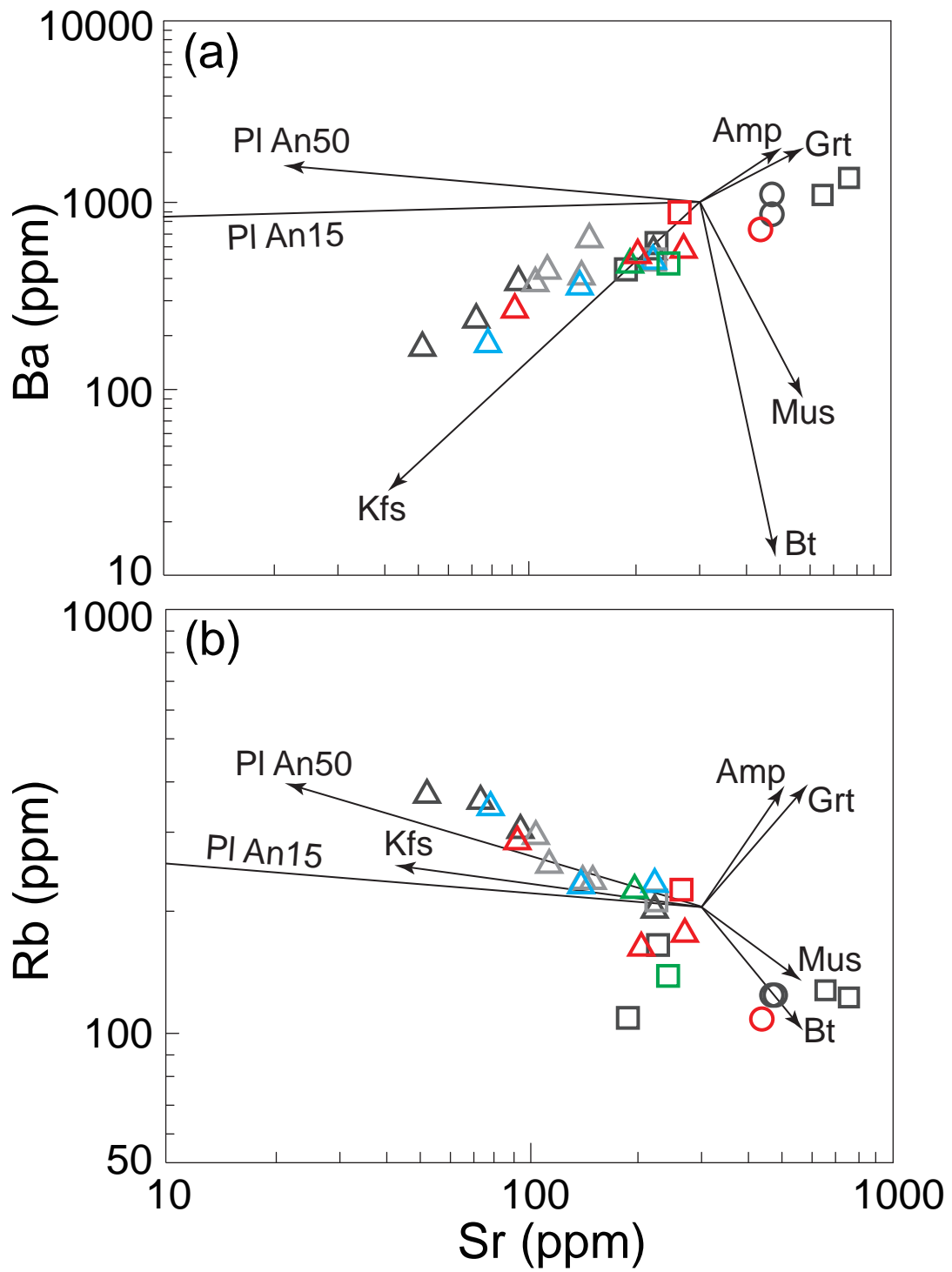


TM: two-mica monzogranite
 BM: biotite monzogranite
 BG: biotite granodiorite



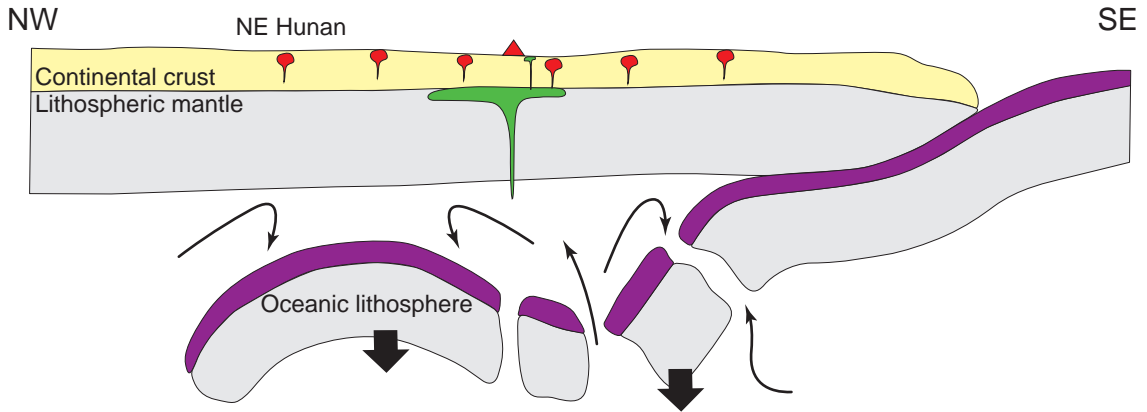




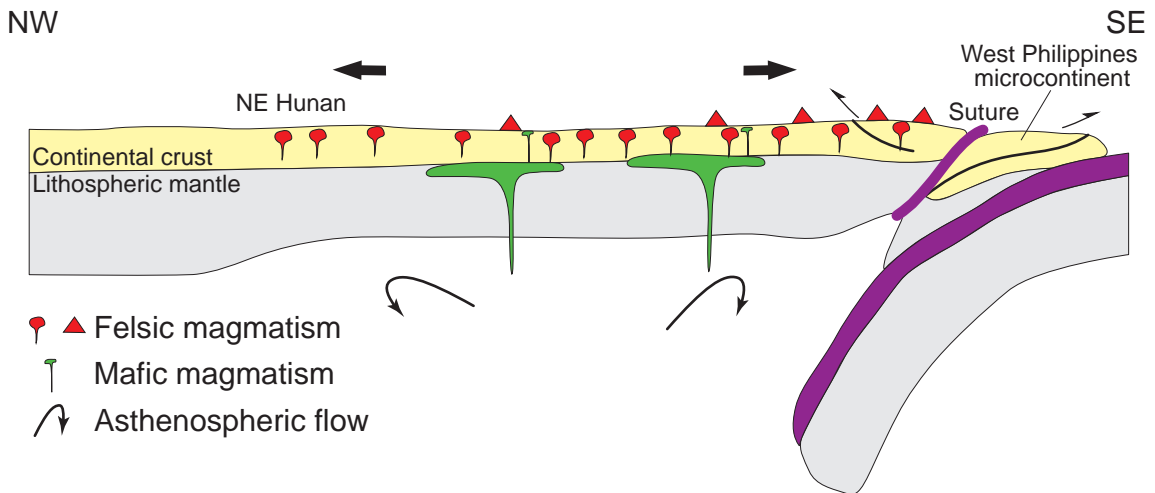


A) Northwestward subduction of the Paleo-Pacific plate

(a) Jurassic: foundering of a subducted oceanic flat-slab



(b) 140-90 Ma: episodic slab rollback with collision during the interval



B) Jurassic slab foundering after a SE-directed intracontinental subduction

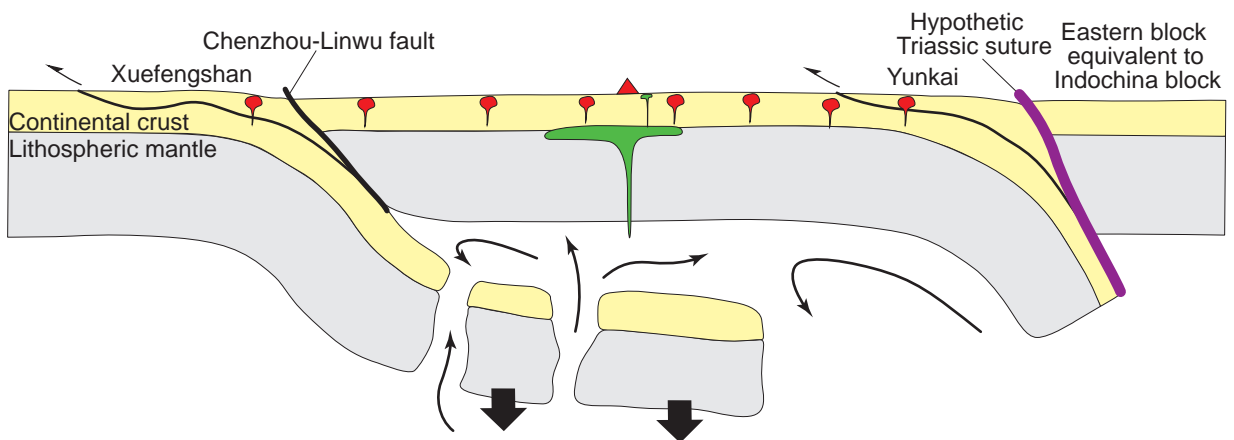


Table 1

Summary of zircon U-Pb ages and Hf isotopic data for the granitic samples

Sample	Site	Lithology	GPS location	Zircon U-Pb age	$\epsilon\text{Hf}(t)$	T_{DM}^{C}
JH155	Taohuashan	biotite monzogranite	N29°37'59", E112°40'40"	128.0 ± 1.0 Ma	-10.1 to -6.9	1.62–1.82 Ga
DM35	Dayunshan-Mufushan	biotite monzogranite	N29°19'35", E113°40'12"	151.2 ± 1.1 Ma	-8.5 to -4.7	1.50–1.74 Ga
QD69	Dayunshan-Mufushan	biotite monzogranite	N28°51'56", E113°37'21"	151.3 ± 1.1 Ma	-7.3 to -4.3	1.47–1.66 Ga
DYS14	Dayunshan-Mufushan	biotite granodiorite	N29°13'19", E113°56'48"	149.0 ± 1.0 Ma	-8.6 to -5.6	1.55–1.75 Ga
QD49	Dayunshan-Mufushan	two-mica monzogranite	N29°17'58", E113°30'52"	131.8 ± 1.5 Ma	-7.9 to -3.6	1.42–1.69 Ga
DM80	Dayunshan-Mufushan	two-mica monzogranite	N29°03'02", E113°53'44"			
QD62	Dayunshan-Mufushan	two-mica monzogranite	N29°13'52", E113°27'56"	127.0 ± 1.4 Ma	-9.6 to -6.2	1.58–1.79 Ga
DYS01	Wangxiang	two-mica monzogranite	N28°28'49", E112°57'56"	147.3 ± 1.5 Ma	-12.5 to -6.4	1.61–1.99 Ga
DYS03	Wangxiang	two-mica monzogranite	N28°30'08", E113°10'28"	146.0 ± 1.4 Ma	-11.6 to -5.9	1.57–1.93 Ga
DYS06	Lianyunshan	two-mica monzogranite	N28°28'16", E113°40'12"	150.4 ± 1.7 Ma	-12.4 to -8.5	1.74–1.98 Ga
DYS07	Lianyunshan	biotite monzogranite	N28°30'34", E113°42'11"	149.0 ± 1.4 Ma	-6.9 to -3.7	1.44–1.64 Ga

Table 2

Major (wt.%) and trace element (ppm) compositions of the granitic samples

Site	Dayunshan-Mufushan					Wangxiang			Lianyunshan	
Sample	QD70	DYS14	QD49	DYS10	DYS15	DYS01	DYS02	DYS03	DYS06	DYS07
Lithology	BM	BG	TM	TM	TM	TM	TM	TM	TM	BM
Major elements (wt.%)										
SiO ₂	70.2	65.5	72.9	73.0	70.3	72.3	70.8	72.8	72.1	76.6
TiO ₂	0.31	0.52	0.21	0.18	0.40	0.20	0.41	0.15	0.27	0.27
Al ₂ O ₃	15.5	15.7	14.5	14.5	15.3	14.8	14.8	14.8	15.0	12.0
Fe ₂ O ₃ *	2.13	4.39	1.60	1.58	2.72	1.61	2.75	1.44	1.88	2.09
FeO*	1.92	3.95	1.44	1.42	2.45	1.45	2.48	1.30	1.69	1.88
MnO	0.03	0.07	0.02	0.03	0.03	0.02	0.04	0.03	0.03	0.02
MgO	0.70	1.51	0.36	0.32	0.79	0.36	0.89	0.29	0.54	0.61
CaO	1.83	4.10	1.61	0.98	2.30	1.64	2.27	1.12	1.72	1.87
Na ₂ O	3.20	2.79	3.37	3.57	3.57	3.15	3.25	3.21	3.23	2.86
K ₂ O	5.43	3.38	4.15	4.58	3.54	5.19	3.63	4.87	4.17	2.96
P ₂ O ₅	0.11	0.22	0.09	0.21	0.12	0.12	0.17	0.17	0.11	0.05
LOI	0.44	1.52	0.44	0.70	0.50	0.46	0.54	0.96	0.66	0.54
Total	99.9	99.7	99.3	99.7	99.6	99.9	99.6	99.8	99.8	99.9
K ₂ O+Na ₂ O	8.6	6.2	7.5	8.2	7.1	8.3	6.9	8.1	7.4	5.8
K ₂ O/Na ₂ O	1.7	1.2	1.2	1.3	1.0	1.6	1.1	1.5	1.3	1.0
ASI	1.07	1.00	1.12	1.15	1.10	1.07	1.10	1.17	1.16	1.06
Trace elements (ppm)										
Sc	4.21	8.75	3.86	2.38	4.46	2.83	4.94	2.34	3.65	3.70
V	32.2	82.7	13.9	17.5	31.3	11.9	42.0	7.6	21.8	28.3
Cr	201	241	218	221	229	262	234	232	172	358
Co	5.76	11.0	4.26	3.61	6.43	4.60	7.70	3.42	4.56	6.30
Ni	10.3	12.6	12.7	15.5	13.7	40.6	17.1	9.39	11.3	22.9
Ga	21.3	21.2	20.0	21.7	24.2	18.8	24.8	22.3	20.6	17.2
Rb	224	112	157	287	169	224	222	343	213	141
Sr	261	437	202	92.9	270	139	222	78.2	193	242
Y	6.77	14.5	5.49	6.03	6.60	5.32	9.45	5.09	5.09	3.94
Zr	155	183	126	93.7	215	87.2	191	81.2	155	146
Nb	7.44	8.56	5.49	11.6	13.15	8.04	12.1	12.1	7.10	6.59
Cs	10.9	3.90	6.61	44.9	8.83	9.13	13.8	36.4	15.0	5.61
Ba	929	740	527	267	584	376	495	177	479	484
La	36.0	34.7	37.8	22.8	48.1	26.1	60.2	20.9	36.6	44.7
Ce	60.7	62.0	64.0	38.9	83.6	46.1	107.1	39.1	61.4	76.8
Pr	6.59	7.96	6.87	4.63	9.13	5.38	11.4	4.67	6.85	8.05
Nd	21.4	29.3	23.0	16.5	30.1	18.8	38.5	16.8	23.2	27.0
Sm	3.45	5.21	3.30	3.09	4.66	3.39	5.51	3.50	3.44	3.80

Eu	0.71	1.23	0.70	0.43	0.82	0.68	0.87	0.36	0.65	0.68
Gd	2.53	4.28	2.44	2.55	3.41	2.53	4.07	2.68	2.36	2.70
Tb	0.31	0.59	0.28	0.35	0.38	0.31	0.48	0.35	0.27	0.28
Dy	1.33	2.90	1.16	1.43	1.49	1.24	2.09	1.30	1.07	0.97
Ho	0.23	0.55	0.19	0.21	0.25	0.19	0.36	0.17	0.19	0.14
Er	0.57	1.35	0.50	0.46	0.65	0.44	0.88	0.37	0.46	0.35
Tm	0.08	0.19	0.07	0.06	0.09	0.06	0.11	0.05	0.07	0.05
Yb	0.49	1.21	0.44	0.36	0.64	0.35	0.70	0.27	0.42	0.33
Lu	0.07	0.19	0.06	0.05	0.10	0.05	0.11	0.04	0.06	0.06
Hf	4.21	5.24	3.84	3.14	6.40	2.72	5.59	2.71	4.68	4.23
Ta	0.67	0.74	0.59	1.98	1.12	0.84	1.19	1.79	0.69	0.39
Th	18.4	12.1	20.6	15.5	29.0	16.7	35.7	15.2	23.9	27.4
U	1.30	2.55	1.41	5.07	2.20	1.86	3.60	6.02	3.39	1.59
ΣREEs	134	152	141	91.8	183	106	232	90.6	137	166
(La/Yb) _N	53.4	62.1	55.4	62.3	97.8	44.9	20.6	54.0	61.7	52.6
Eu/Eu*	0.68	0.54	0.35	0.66	0.62	0.45	0.77	0.60	0.72	0.70
Nb/Ta	11.1	11.5	9.3	5.9	11.7	9.6	10.2	6.8	10.3	16.9
Zr/Hf	36.9	34.9	32.9	29.9	33.6	32.1	34.1	30.0	33.2	34.5

Supplementary Table S1

SIMS zircon U-Pb data of the granitic samples

Sample#spot	U ppm	Th ppm	Th/U	f ₂₀₆ %	$\frac{^{207}\text{Pb}}{^{235}\text{U}}$	±σ (%)	$\frac{^{206}\text{Pb}}{^{238}\text{U}}$	±σ (%)	t _{207/206} (Ma)	±σ (Ma)	t _{207/235} (Ma)	±σ (Ma)	t _{206/238} (Ma)	±σ (Ma)
JH155														
JH155#1	829	621	0.75	0.08	0.13126	2.16	0.0201	1.51	71.2	36.2	125.2	2.5	128.1	1.9
JH155#2	2159	1194	0.55	0.08	0.13520	1.81	0.0201	1.50	138.8	23.4	128.8	2.2	128.2	1.9
JH155#3	651	583	0.90	0.07	0.13328	2.01	0.0198	1.50	135.3	31.3	127.0	2.4	126.6	1.9
JH155#4	652	452	0.69	0.03	0.13542	2.10	0.0201	1.56	143.4	33.0	129.0	2.6	128.2	2.0
JH155#5	1510	1053	0.70	0.00	0.13712	1.74	0.0206	1.51	117.3	20.2	130.5	2.1	131.2	2.0
JH155#6	1345	372	0.28	0.01	0.13563	2.05	0.0196	1.67	205.1	27.2	129.1	2.5	125.1	2.1
JH155#7	748	502	0.67	0.06	0.13508	2.36	0.0199	1.59	157.0	40.2	128.7	2.9	127.1	2.0
JH155#8	1176	662	0.56	0.00	0.13710	1.88	0.0203	1.52	145.8	25.7	130.5	2.3	129.6	2.0
JH155#9	803	463	0.58	0.06	0.13471	1.92	0.0201	1.50	133.8	27.9	128.3	2.3	128.0	1.9
JH155#10	494	374	0.76	0.00	0.13404	2.30	0.0199	1.72	143.5	35.2	127.7	2.8	126.9	2.2
JH155#11	276	261	0.95	0.00	0.13677	3.07	0.0198	1.57	203.1	60.2	130.2	3.8	126.2	2.0
JH155#12	623	235	0.38	0.11	0.42838	1.90	0.0513	1.67	622.6	19.6	362.0	5.8	322.7	5.2
JH155#13	1160	703	0.61	0.09	0.13615	1.98	0.0203	1.52	125.8	29.6	129.6	2.4	129.8	1.9
JH155#14	1401	472	0.34	0.02	0.13355	1.75	0.0203	1.50	81.5	21.3	127.3	2.1	129.8	1.9
JH155#15	668	593	0.89	0.04	0.13315	2.07	0.0197	1.52	145.8	32.5	126.9	2.5	125.9	1.9
JH155#16	1224	1166	0.95	0.09	0.13875	2.88	0.0206	2.51	143.5	32.9	131.9	3.6	131.3	3.3
JH155#17	751	477	0.63	0.06	0.13413	2.48	0.0199	1.53	142.9	45.1	127.8	3.0	127.0	1.9
JH155#18	934	547	0.59	0.05	0.13063	3.24	0.0196	3.04	113.4	26.0	124.7	3.8	125.2	3.8
JH155#19	722	546	0.76	0.00	0.13722	2.40	0.0205	1.70	128.9	39.5	130.6	2.9	130.7	2.2
DM35														
DM35#1	874	289	0.33	0.07	0.16198	2.16	0.0239	1.53	157.5	35.4	152.4	3.1	152.1	2.3
DM35#2	917	392	0.43	0.10	0.15749	1.77	0.0236	1.51	116.6	21.4	148.5	2.4	150.5	2.2
DM35#3	1121	585	0.52	0.04	0.15880	1.88	0.0237	1.51	132.7	26.0	149.7	2.6	150.7	2.3
DM35#4	1288	751	0.58	0.06	0.15827	1.69	0.0235	1.50	144.9	17.9	149.2	2.3	149.5	2.2
DM35#5	1195	583	0.49	0.03	0.16061	1.88	0.0238	1.70	143.3	18.6	151.2	2.6	151.7	2.5

DM35#6	1448	793	0.55	0.07	0.16219	1.67	0.0239	1.50	156.9	17.1	152.6	2.4	152.3	2.3
DM35#7	942	419	0.45	0.02	0.15977	2.07	0.0236	1.50	157.4	32.8	150.5	2.9	150.1	2.2
DM35#8	709	256	0.36	0.00	0.16005	1.83	0.0235	1.51	164.0	24.3	150.7	2.6	149.9	2.2
DM35#9	517	334	0.65	0.03	0.16261	1.92	0.0240	1.50	153.3	28.0	153.0	2.7	153.0	2.3
DM35#10	1292	725	0.56	0.04	0.16118	1.68	0.0239	1.50	144.8	17.8	151.7	2.4	152.2	2.3
DM35#11	836	334	0.40	0.07	0.16293	1.85	0.0237	1.50	190.4	24.8	153.3	2.6	150.9	2.2
DM35#12	2406	497	0.21	0.08	0.43199	7.53	0.0427	6.80	1022.7	64.2	364.6	23.3	269.7	18.0
DM35#13	917	510	0.56	0.03	0.16085	1.79	0.0238	1.51	150.3	22.2	151.4	2.5	151.5	2.3
DM35#14	882	394	0.45	0.17	0.15866	1.77	0.0239	1.50	110.4	22.1	149.5	2.5	152.0	2.3
DM35#15	893	420	0.47	0.06	0.16213	2.02	0.0238	1.51	167.2	30.9	152.6	2.9	151.6	2.3
DM35#16	1227	694	0.57	0.07	0.15839	1.82	0.0234	1.51	148.0	23.7	149.3	2.5	149.4	2.2
DM35#17	887	287	0.32	0.22	1.07765	1.61	0.1156	1.50	856.9	11.7	742.5	8.5	705.1	10.0
DM35#18	1075	584	0.54	0.15	0.16247	1.74	0.0240	1.53	153.8	19.6	152.9	2.5	152.8	2.3

QD69

QD69#1	1839	345	0.19	0.05	0.15737	1.86	0.0236	1.50	115.3	25.8	148.4	2.6	150.5	2.2
QD69#2	3341	422	0.13	0.00	0.16283	1.87	0.0240	1.58	160.5	23.3	153.2	2.7	152.7	2.4
QD69#3	2516	271	0.11	0.00	0.16042	1.76	0.0237	1.50	155.8	21.5	151.1	2.5	150.8	2.2
QD69#4	755	27	0.04	0.00	0.16162	2.31	0.0238	1.52	159.8	40.2	152.1	3.3	151.6	2.3
QD69#5	7007	979	0.14	0.00	0.17165	1.60	0.0251	1.50	172.0	12.7	160.8	2.4	160.1	2.4
QD69#6	4952	771	0.16	0.34	0.16161	1.91	0.0238	1.50	158.3	27.3	152.1	2.7	151.7	2.3
QD69#7	4816	700	0.15	0.01	0.16316	1.64	0.0241	1.51	154.2	15.4	153.5	2.3	153.4	2.3
QD69#8	3176	612	0.19	0.06	0.16086	1.80	0.0239	1.50	140.2	23.2	151.5	2.5	152.2	2.3
QD69#9	2483	405	0.16	0.01	0.16172	1.77	0.0240	1.50	142.8	21.9	152.2	2.5	152.8	2.3
QD69#10	126	85	0.68	0.08	1.26671	2.11	0.1387	1.50	814.3	30.7	830.9	12.1	837.1	11.8
QD69#11	521	397	0.76	0.15	0.15846	2.56	0.0235	1.50	144.0	47.8	149.4	3.6	149.7	2.2
QD69#12	1466	106	0.07	0.00	0.15587	1.95	0.0235	1.51	108.6	29.0	147.1	2.7	149.5	2.2
QD69#13	3703	727	0.20	0.07	0.15923	1.74	0.0237	1.50	136.3	20.4	150.0	2.4	150.9	2.2
QD69#14	5370	652	0.12	0.23	0.16664	1.69	0.0251	1.50	110.0	18.5	156.5	2.5	159.6	2.4
QD69#15	2129	217	0.10	0.12	0.15816	2.05	0.0234	1.50	151.2	32.3	149.1	2.8	149.0	2.2
QD69#16	2873	246	0.09	0.01	0.16333	1.77	0.0239	1.54	178.3	20.0	153.6	2.5	152.0	2.3

QD69#17	2996	340	0.11	0.04	0.15894	1.73	0.0237	1.50	131.1	20.1	149.8	2.4	151.0	2.2
QD69#18	427	250	0.59	0.14	0.16414	2.74	0.0242	1.53	152.8	52.5	154.3	3.9	154.4	2.3
QD69#19	1295	1783	1.38	0.06	0.15940	1.99	0.0236	1.50	145.7	30.2	150.2	2.8	150.5	2.2
DYS14														
DYS14#1	303	185	0.61	0.31	0.17024	3.05	0.0245	1.55	217.8	59.8	159.6	4.5	155.7	2.4
DYS14#2	381	298	0.78	0.15	0.16478	3.04	0.0237	1.53	217.4	59.7	154.9	4.4	150.8	2.3
DYS14#3	439	226	0.52	0.26	0.15967	2.70	0.0237	1.52	142.2	51.5	150.4	3.8	150.9	2.3
DYS14#4	330	258	0.78	0.29	0.15909	3.02	0.0233	1.59	176.6	58.7	149.9	4.2	148.2	2.3
DYS14#5	457	216	0.47	0.02	0.16518	2.63	0.0233	1.53	255.7	48.4	155.2	3.8	148.7	2.2
DYS14#6	378	372	0.98	0.40	0.16049	3.83	0.0237	1.58	150.7	79.8	151.1	5.4	151.2	2.4
DYS14#7	517	168	0.32	0.15	0.15539	2.58	0.0229	1.52	156.1	47.9	146.7	3.5	146.1	2.2
DYS14#8	532	301	0.57	0.11	0.15615	2.53	0.0238	1.50	81.9	47.6	147.3	3.5	151.4	2.2
DYS14#9	627	281	0.45	0.13	0.16374	2.39	0.0237	1.52	200.7	42.4	154.0	3.4	151.0	2.3
DYS14#10	942	529	0.56	0.06	0.16027	2.13	0.0234	1.51	179.0	34.5	150.9	3.0	149.2	2.2
DYS14#11	414	277	0.67	0.38	0.15388	3.83	0.0230	1.59	123.3	80.1	145.3	5.2	146.7	2.3
DYS14#12	326	162	0.50	0.12	0.15723	3.34	0.0235	1.52	129.2	68.6	148.3	4.6	149.5	2.3
DYS14#13	294	258	0.88	0.11	0.15983	3.10	0.0234	1.51	175.5	62.0	150.6	4.3	149.0	2.2
DYS14#14r	694	117	0.17	0.14	0.15945	3.23	0.0237	1.51	141.8	65.8	150.2	4.5	150.8	2.3
DYS14#14c	948	74	0.08	0.01	1.24630	1.61	0.1355	1.51	829.0	11.6	821.7	9.1	819.1	11.6
DYS14#15	699	620	0.89	0.34	0.15739	3.04	0.0235	1.50	131.3	61.0	148.4	4.2	149.5	2.2
DYS14#16	261	166	0.64	0.72	0.14985	4.20	0.0231	1.55	48.2	90.8	141.8	5.6	147.4	2.3
DYS14#17	676	370	0.55	0.17	0.16449	2.33	0.0235	1.52	232.5	40.3	154.6	3.4	149.6	2.3
DYS14#18	629	381	0.61	0.47	0.15384	3.19	0.0231	1.54	116.1	64.7	145.3	4.3	147.1	2.2
DYS14#19	726	561	0.77	0.00	0.15499	2.32	0.0227	1.51	171.0	40.6	146.3	3.2	144.8	2.2
DYS14#20	674	504	0.75	0.23	0.15346	3.24	0.0234	1.60	76.8	65.6	145.0	4.4	149.2	2.4
QD49														
QD49#1	5166	109	0.02	0.03	0.14105	1.62	0.0210	1.50	132.1	14.1	134.0	2.0	134.1	2.0
QD49#2	2432	358	0.15	0.04	0.13996	1.75	0.0209	1.53	125.5	20.2	133.0	2.2	133.4	2.0
QD49#3	3685	438	0.12	0.04	0.14215	1.60	0.0211	1.50	137.3	12.6	135.0	2.0	134.8	2.0
QD49#4	88	81	0.92	0.72	0.13724	4.29	0.0203	1.56	145.7	91.2	130.6	5.3	129.8	2.0

QD49#5	831	697	0.84	0.07	0.14644	1.90	0.0218	1.52	133.5	26.5	138.8	2.5	139.1	2.1
QD49#6	162	154	0.95	0.15	0.13679	3.37	0.0204	1.53	132.3	69.0	130.2	4.1	130.1	2.0
QD49#7	157	164	1.04	0.13	0.14519	3.13	0.0200	1.54	315.0	60.9	137.7	4.0	127.6	1.9
QD49#8	2981	675	0.23	0.02	0.14360	1.64	0.0215	1.53	123.8	14.0	136.2	2.1	137.0	2.1
QD49#9	938	325	0.35	0.00	0.13668	2.13	0.0206	1.56	101.9	33.8	130.1	2.6	131.6	2.0
QD49#10	525	334	0.64	0.01	1.27868	1.56	0.1377	1.50	849.3	9.1	836.3	8.9	831.4	11.7
QD49#11	161	92	0.57	0.16	0.19698	2.96	0.0273	1.64	299.6	55.4	182.6	5.0	173.6	2.8
QD49#12	367	278	0.76	0.16	0.14373	2.82	0.0212	1.56	155.7	53.9	136.4	3.6	135.3	2.1
QD49#13	618	416	0.67	0.01	1.29991	1.56	0.1406	1.51	839.8	8.3	845.7	9.0	847.9	12.0
QD49#14	706	748	1.06	0.13	0.13346	2.03	0.0199	1.50	129.6	31.9	127.2	2.4	127.1	1.9
QD49#15	427	188	0.44	0.02	0.80096	1.63	0.0910	1.51	736.9	13.0	597.4	7.4	561.3	8.1
QD49#16	252	222	0.88	0.36	0.13409	3.38	0.0203	1.58	89.0	69.4	127.8	4.1	129.9	2.0
QD49#17	1186	442	0.37	0.03	0.49973	1.86	0.0594	1.80	639.8	10.3	411.5	6.3	372.0	6.5
QD49#18	413	119	0.29	0.08	0.13458	2.12	0.0201	1.51	125.3	34.6	128.2	2.6	128.4	1.9
QD49#19	1272	375	0.29	0.04	0.13839	1.71	0.0208	1.51	106.6	19.2	131.6	2.1	133.0	2.0
QD49#20	632	148	0.23	0.08	0.13654	1.93	0.0204	1.51	129.0	27.9	130.0	2.4	130.0	1.9
QD49#21	975	525	0.54	0.01	2.35906	1.53	0.2087	1.50	1245.5	5.4	1230.3	10.9	1221.6	16.7
QD62														
DQ62#1	3724	483	0.13	0.06	0.13365	1.99	0.0201	1.50	107.4	30.4	127.4	2.4	128.4	1.9
DQ62#2	1266	474	0.37	0.02	1.09491	1.71	0.1195	1.51	820.2	16.6	750.9	9.1	727.8	10.4
DQ62#3	537	50	0.09	0.18	0.13033	2.77	0.0195	1.58	122.4	52.7	124.4	3.2	124.5	2.0
DQ62#4	9826	1410	0.14	0.04	0.14160	1.93	0.0212	1.85	121.6	12.6	134.5	2.4	135.2	2.5
DQ62#5r	1170	245	0.21	0.04	0.13080	2.16	0.0193	1.53	151.2	35.3	124.8	2.5	123.4	1.9
DQ62#5c	3033	397	0.13	0.32	0.15528	2.50	0.0220	1.98	247.2	34.6	146.6	3.4	140.4	2.8
DQ62#6	1719	321	0.19	0.32	0.22448	2.22	0.0290	1.53	456.4	35.1	205.6	4.1	184.4	2.8
DQ62#7	2839	425	0.15	0.33	0.13321	2.37	0.0199	1.55	121.3	41.8	127.0	2.8	127.3	2.0
DQ62#8	1708	363	0.21	0.07	0.13338	1.97	0.0198	1.51	143.4	29.5	127.1	2.4	126.3	1.9
DQ62#9	2516	533	0.21	0.16	0.12682	2.01	0.0188	1.52	143.4	30.6	121.2	2.3	120.1	1.8
DQ62#10	1260	526	0.42	0.00	0.14053	2.07	0.0204	1.51	190.0	32.8	133.5	2.6	130.4	1.9
DQ62#11	602	96	0.16	0.08	0.13963	3.02	0.0210	1.63	113.2	58.8	132.7	3.8	133.8	2.2

DQ62#12	2099	541	0.26	0.01	0.14456	1.87	0.0216	1.50	123.4	26.2	137.1	2.4	137.9	2.0
DQ62#13	2693	376	0.14	0.08	0.12438	1.83	0.0186	1.51	123.3	24.3	119.0	2.1	118.8	1.8
DQ62#14	1634	186	0.11	0.12	0.13385	1.99	0.0202	1.50	102.3	30.7	127.5	2.4	128.9	1.9
DQ62#15	2636	509	0.19	0.06	0.13191	1.87	0.0198	1.57	114.8	23.7	125.8	2.2	126.4	2.0
DQ62#16	1802	321	0.18	0.02	0.13246	1.94	0.0197	1.52	132.6	28.2	126.3	2.3	126.0	1.9
DQ62#17	1967	244	0.12	0.06	0.13716	1.90	0.0206	1.50	115.1	27.1	130.5	2.3	131.4	2.0
DQ62#18-1	1619	291	0.18	0.04	0.13025	2.00	0.0193	1.51	140.3	30.6	124.3	2.3	123.5	1.8
DQ62#18-2	1702	130	0.08	0.01	0.12839	2.55	0.0188	1.59	176.6	45.8	122.7	3.0	119.9	1.9
DQ62#19	2239	227	0.10	0.01	0.13796	1.87	0.0202	1.53	168.8	25.1	131.2	2.3	129.2	2.0
DQ62#20	3402	420	0.12	0.02	0.13739	1.75	0.0201	1.50	180.0	20.5	130.7	2.1	128.0	1.9
DQ62#21	851	106	0.12	0.06	0.13970	2.39	0.0195	1.56	278.8	41.0	132.8	3.0	124.8	1.9
DM80														
DM80#1	4582	500	0.11	0.02	0.15123	1.85	0.0225	1.74	167.3	16.0	146.5	2.4	145.2	2.4
DM80#2	5175	1926	0.37	0.01	0.15517	1.57	0.0230	1.51	116.7	8.9	145.3	2.1	147.0	2.2
DM80#3	2666	3332	1.25	0.12	0.15158	1.62	0.0225	1.50	86.1	64.5	142.4	4.2	145.8	2.2
DM80#4	2093	623	0.30	0.04	0.14557	1.71	0.0217	1.50	131.6	23.6	140.8	2.9	141.3	2.7
DM80#5	2611	941	0.36	2.31	0.14396	2.92	0.0226	1.51	140.4	14.9	141.1	2.2	141.1	2.1
DM80#6	3496	563	0.16	0.14	0.15515	1.79	0.0228	1.66	0.0	68.1	136.6	3.7	144.0	2.1
DM80#7	2806	366	0.13	0.00	0.15902	1.70	0.0238	1.50	-0.1	45.4	133.3	3.0	140.8	2.1
DM80#8	1861	331	0.18	0.58	0.14024	2.43	0.0221	1.50	140.0	12.9	140.2	2.1	140.3	2.1
DM80#9	726	264	0.36	0.08	0.14425	2.03	0.0215	1.52	125.4	24.1	137.2	2.4	137.9	2.1
DM80#10	2183	654	0.30	0.03	0.14872	2.19	0.0222	1.94	130.3	19.3	138.0	2.2	138.4	2.1
DM80#11	4500	2228	0.50	0.02	0.14916	1.62	0.0222	1.53	147.1	10.2	146.5	2.1	146.4	2.2
DM80#12	607	97	0.16	0.05	1.25570	2.19	0.1385	2.13	798.9	10.0	826.0	12.4	836.1	16.7
DM80#13	6792	217	0.03	0.02	0.15381	1.55	0.0231	1.50	142.2	8.5	147.1	2.1	147.4	2.2
DM80#14	2125	692	0.33	0.09	0.14805	1.75	0.0221	1.51	133.8	14.5	143.0	2.5	143.6	2.5
DM80#15	4550	1187	0.26	0.02	0.15624	1.57	0.0233	1.50	122.3	18.8	149.8	2.4	151.6	2.2
DM80#16	3131	761	0.24	1.79	0.14674	3.53	0.0229	1.50	18.9	74.9	139.0	4.6	146.2	2.2
DM80#17	2475	875	0.35	0.61	0.15060	3.16	0.0229	1.51	133.6	31.3	136.8	2.6	137.0	2.1
DM80#18	1743	710	0.41	0.03	0.14746	1.71	0.0219	1.51	123.4	20.4	140.2	2.3	141.2	2.1

DM80#19	3337	1885	0.56	0.07	0.14905	1.63	0.0221	1.50	96.4	48.0	141.7	3.4	144.4	2.1
DM80#20	4268	1384	0.32	0.85	0.14976	2.55	0.0227	1.50	137.8	17.0	142.0	2.2	142.3	2.1
DM80#21	3295	499	0.15	0.04	0.14811	1.61	0.0220	1.51	129.6	12.4	141.2	2.1	141.9	2.1
DM80#22	810	113	0.14	0.02	0.16159	1.54	0.0241	1.51	804.4	9.1	768.1	8.5	755.7	10.8
DM80#23	405	70	0.17	0.03	0.19986	1.54	0.0270	1.51	802.7	10.7	814.8	9.1	819.2	11.7
DM80#24	68	43	0.63	0.07	0.15588	1.55	0.0231	1.51	860.0	27.9	852.8	11.9	850.1	12.3
DM80#25	690	535	0.78	0.04	1.20012	1.56	0.1305	1.50	828.5	9.0	800.6	8.7	790.7	11.2
DM80#26	260	255	0.98	0.06	1.13078	1.57	0.1244	1.51	447.6	27.8	421.3	6.8	416.5	6.1
DM80#27	1452	262	0.18	0.04	0.76740	1.57	0.0874	1.50	575.7	30.4	320.4	15.0	286.3	14.7
DM80#28	1633	595	0.36	0.06	1.23105	1.61	0.1355	1.53	142.3	18.8	139.7	2.2	139.5	2.1
DM80#29	2488	784	0.32	0.02	0.15014	1.68	0.0223	1.52	136.6	14.3	143.3	2.2	143.7	2.1
DM80#30	10911	941	0.09	0.02	0.14468	1.84	0.0216	1.52	131.1	6.3	152.1	2.2	153.4	2.3
DM80#31	1329	584	0.44	0.05	0.51429	1.96	0.0668	1.50	731.9	9.6	578.3	6.9	539.9	7.8
DM80#32	7413	115	0.02	0.01	1.31618	2.06	0.1410	1.55	355.1	7.4	185.0	2.6	172.0	2.6
DM80#33	6842	244	0.04	0.03	0.37095	5.43	0.0454	5.24	125.3	10.8	147.4	2.2	148.8	2.2
DYS01														
DYS01#1	2497	18	0.01	0.03	0.15107	1.66	0.0225	1.50	128.1	16.3	142.9	2.2	143.7	2.1
DYS01#2	6697	339	0.05	0.01	0.16590	1.55	0.0245	1.50	150.0	8.6	155.9	2.2	156.2	2.3
DYS01#3	533	39	0.07	0.09	0.15502	2.13	0.0226	1.50	178.3	34.9	146.3	2.9	144.4	2.1
DYS01#4	3374	41	0.01	0.04	0.15511	1.60	0.0229	1.51	153.3	12.6	146.4	2.2	146.0	2.2
DYS01#5	656	206	0.31	0.00	0.15348	1.94	0.0229	1.50	128.0	28.6	145.0	2.6	146.0	2.2
DYS01#6	12087	689	0.06	3.39	0.15409	3.65	0.0228	1.53	151.5	75.9	145.5	5.0	145.1	2.2
DYS01#7	7008	369	0.05	0.02	0.16130	1.62	0.0241	1.52	129.9	12.8	151.8	2.3	153.3	2.3
DYS01#8	373	202	0.54	0.02	3.65417	1.57	0.2523	1.52	1715.1	7.4	1561.4	12.6	1450.3	19.7
DYS01#9	3687	269	0.07	0.10	0.15628	1.70	0.0233	1.52	133.2	17.8	147.4	2.3	148.3	2.2
DYS01#10	6866	1318	0.19	0.59	0.15681	1.69	0.0236	1.51	110.8	17.8	147.9	2.3	150.2	2.2
DYS01#11	1757	1615	0.92	0.01	0.15514	1.68	0.0230	1.51	147.2	17.2	146.4	2.3	146.4	2.2
DYS01#12	1396	285	0.20	0.05	0.16061	2.01	0.0237	1.53	156.0	30.1	151.2	2.8	150.9	2.3
DYS01#13	351	419	1.20	0.01	1.28026	1.61	0.1385	1.50	839.4	12.2	837.0	9.2	836.0	11.8
DYS01#14	227	79	0.35	0.03	1.26987	1.69	0.1388	1.50	817.6	16.0	832.3	9.6	837.9	11.8

DYS01#15	1878	245	0.13	0.32	0.35173	6.10	0.0416	6.06	649.5	16.1	306.0	16.3	262.9	15.6
DYS01#16	3882	55	0.01	0.01	0.15649	1.61	0.0230	1.50	159.8	13.7	147.6	2.2	146.9	2.2
DYS01#17	122	104	0.85	0.05	1.27447	1.80	0.1389	1.50	823.5	20.6	834.4	10.3	838.5	11.8
DYS01#18	2085	464	0.22	1.29	0.15886	2.32	0.0230	1.51	196.2	40.4	149.7	3.2	146.8	2.2
DYS01#19	2960	22	0.01	0.01	0.15287	1.64	0.0225	1.50	157.3	15.5	144.4	2.2	143.7	2.1
DYS01#20	1975	238	0.12	0.01	0.15870	1.65	0.0234	1.50	154.3	16.0	149.6	2.3	149.3	2.2
DYS01#21	1516	734	0.48	0.04	0.14858	2.05	0.0218	1.81	170.9	22.6	140.7	2.7	138.9	2.5
DYS01#22	319	245	0.77	0.00	0.15471	2.30	0.0229	1.51	149.6	40.4	146.1	3.1	145.8	2.2
DYS01#23	658	690	1.05	0.07	0.15814	1.92	0.0237	1.50	116.5	27.9	149.1	2.7	151.1	2.2
DYS03														
DYS03#1	710	356	0.50	0.03	1.05139	1.76	0.1138	1.53	837.9	17.8	729.6	9.2	694.8	10.1
DYS03#2	519	120	0.23	0.10	0.65209	2.30	0.0769	1.50	656.5	37.0	509.8	9.3	477.7	6.9
DYS03#3	1980	86	0.04	0.57	0.14649	2.80	0.0224	1.52	75.7	55.1	138.8	3.6	142.5	2.1
DYS03#4	2124	867	0.41	0.55	0.14896	2.27	0.0225	1.51	104.3	39.5	141.0	3.0	143.2	2.1
DYS03#5	857	748	0.87	0.05	0.15144	2.36	0.0227	1.52	117.8	42.2	143.2	3.2	144.7	2.2
DYS03#6	3677	1087	0.30	0.11	0.15293	1.76	0.0228	1.51	131.3	21.2	144.5	2.4	145.3	2.2
DYS03#7	1651	336	0.20	0.15	0.15799	2.14	0.0239	1.50	93.3	35.8	148.9	3.0	152.5	2.3
DYS03#8c	756	135	0.18	0.09	0.49649	2.17	0.0584	1.89	661.1	22.5	409.3	7.3	366.1	6.7
DYS03#8r	12950	115	0.01	0.00	0.16242	1.57	0.0241	1.51	139.5	10.0	152.8	2.2	153.7	2.3
DYS03#9	4548	728	0.16	1.75	0.16132	1.80	0.0234	1.51	194.7	22.6	151.9	2.5	149.1	2.2
DYS03#10	1959	1694	0.86	0.04	0.15565	2.23	0.0227	1.50	186.8	38.0	146.9	3.1	144.4	2.1
DYS03#11c	1754	121	0.07	0.22	0.20955	2.05	0.0288	1.57	321.2	29.7	193.2	3.6	182.9	2.8
DYS03#11r	38596	271	0.01	0.01	0.14310	1.54	0.0212	1.51	146.3	6.5	135.8	2.0	135.2	2.0
DYS03#12	884	466	0.53	0.09	0.15729	2.33	0.0228	1.51	198.9	40.7	148.3	3.2	145.2	2.2
DYS03#13	1579	108	0.07	0.01	0.15413	2.07	0.0228	1.51	146.7	33.0	145.5	2.8	145.5	2.2
DYS03#14	592	786	1.33	0.34	0.15113	3.64	0.0217	1.53	217.1	74.8	142.9	4.9	138.5	2.1
DYS03#15	361	188	0.52	0.61	0.16708	3.05	0.0233	1.83	287.6	54.9	156.9	4.4	148.4	2.7
DYS03#16	608	16	0.03	0.10	0.15914	2.54	0.0228	1.52	220.0	46.3	150.0	3.5	145.6	2.2
DYS03#17	1540	722	0.47	0.04	0.15604	1.95	0.0227	1.50	185.0	28.8	147.2	2.7	144.9	2.2
DYS03#18	2625	1269	0.48	0.17	0.16346	1.85	0.0245	1.50	122.5	25.4	153.7	2.6	155.8	2.3

DYS03#19	3181	949	0.30	0.50	0.15240	2.21	0.0229	1.64	110.8	34.7	144.0	3.0	146.1	2.4
DYS03#20	6468	1230	0.19	0.59	0.15644	2.20	0.0233	1.51	133.2	37.2	147.6	3.0	148.5	2.2
DYS03#21	2076	402	0.19	0.06	0.16477	1.95	0.0242	1.55	166.3	27.5	154.9	2.8	154.1	2.4
DYS03#22	867	171	0.20	0.00	0.14711	2.24	0.0216	1.50	164.6	38.3	139.4	2.9	137.9	2.0
DYS03#23	584	329	0.56	0.07	0.14941	3.14	0.0229	1.52	67.0	64.0	141.4	4.1	145.9	2.2
DYS06														
DYS06#1	1784	775	0.43	0.00	0.15970	1.68	0.0239	1.52	125.9	16.9	150.4	2.4	152.0	2.3
DYS06#2	611	450	0.74	0.04	0.15862	2.49	0.0230	1.89	197.1	37.1	149.5	3.5	146.5	2.7
DYS06#3	2394	7919	3.31	1.57	0.15341	5.70	0.0228	2.02	134.0	120.8	144.9	7.7	145.6	2.9
DYS06#4	1028	1036	1.01	0.03	0.16456	2.58	0.0248	2.27	109.3	28.4	154.7	3.7	157.7	3.5
DYS06#5	1371	910	0.66	0.26	0.15392	1.92	0.0227	1.50	153.8	28.0	145.4	2.6	144.9	2.1
DYS06#6	478	877	1.83	0.78	0.16495	3.25	0.0232	1.50	265.4	64.8	155.0	4.7	147.9	2.2
DYS06#7	660	1346	2.04	0.02	0.15882	1.93	0.0235	1.53	150.5	27.4	149.7	2.7	149.6	2.3
DYS06#8	721	350	0.49	0.04	0.15692	1.96	0.0234	1.55	131.8	27.7	148.0	2.7	149.0	2.3
DYS06#9	1420	376	0.26	0.05	0.16141	1.70	0.0245	1.50	91.7	18.7	151.9	2.4	155.8	2.3
DYS06#10	551	808	1.47	0.10	0.16026	2.00	0.0238	1.52	143.8	30.2	150.9	2.8	151.4	2.3
DYS06#11	487	575	1.18	0.08	0.15906	2.04	0.0240	1.51	104.9	32.2	149.9	2.9	152.7	2.3
DYS06#12	499	788	1.58	0.14	0.15510	2.29	0.0234	1.51	103.9	40.1	146.4	3.1	149.0	2.2
DYS06#13	1107	982	0.89	0.05	0.16099	1.81	0.0238	1.51	154.5	23.2	151.6	2.6	151.4	2.3
DYS06#14	908	653	0.72	0.03	0.15714	2.25	0.0236	1.52	109.6	38.7	148.2	3.1	150.6	2.3
DYS06#15	571	931	1.63	0.07	0.15955	2.30	0.0233	1.76	174.6	34.1	150.3	3.2	148.8	2.6
DYS06#16	628	1052	1.67	1.19	0.15311	3.99	0.0230	1.50	113.4	85.0	144.7	5.4	146.6	2.2
DYS06#17	975	279	0.29	0.21	0.85995	2.32	0.0953	2.24	789.1	12.8	630.1	11.0	586.7	12.6
DYS06#18	4980	1882	0.38	0.82	0.15399	2.35	0.0229	1.50	132.5	41.9	145.4	3.2	146.2	2.2
DYS06#19	920	775	0.84	0.00	0.16346	1.82	0.0239	1.52	181.2	23.2	153.7	2.6	152.0	2.3
DYS06#20	230	124	0.54	0.01	1.25457	1.67	0.1367	1.50	824.7	15.1	825.5	9.5	825.8	11.6
DYS06#21	4792	1949	0.41	0.01	0.16767	1.59	0.0248	1.50	154.1	12.6	157.4	2.3	157.6	2.3
DYS06#22	1404	1519	1.08	0.94	0.17251	2.86	0.0254	1.57	157.1	55.1	161.6	4.3	161.9	2.5
DYS06#23	880	992	1.13	0.04	0.16616	2.03	0.0244	1.53	166.8	30.7	156.1	2.9	155.4	2.4
DYS06#24	3364	1432	0.43	0.07	0.16000	2.45	0.0236	2.00	157.4	32.6	150.7	3.4	150.3	3.0

DYS07

DYS07#1	1529	33	0.02	0.01	0.15926	1.96	0.0234	1.52	161.6	28.6	150.1	2.7	149.3	2.2
DYS07#2	3377	537	0.16	0.04	0.15724	1.72	0.0235	1.50	120.4	19.5	148.3	2.4	150.0	2.2
DYS07#3	5009	929	0.19	0.33	0.16301	1.96	0.0242	1.50	144.9	29.4	153.3	2.8	153.9	2.3
DYS07#4c	282	197	0.70	0.15	0.15002	3.88	0.0226	1.55	103.7	82.1	141.9	5.2	144.2	2.2
DYS07#5c	381	65	0.17	0.11	0.16321	2.94	0.0236	1.53	204.7	57.3	153.5	4.2	150.2	2.3
DYS07#6	2291	466	0.20	0.06	0.15768	1.86	0.0234	1.51	146.6	25.5	148.7	2.6	148.8	2.2
DYS07#7	3255	142	0.04	0.01	0.15493	1.73	0.0229	1.50	153.2	20.3	146.3	2.4	145.8	2.2
DYS07#8	2586	173	0.07	0.04	0.15546	1.78	0.0228	1.52	168.2	21.5	146.7	2.4	145.4	2.2
DYS07#9	2868	135	0.05	0.01	0.15566	1.80	0.0229	1.50	163.4	23.1	146.9	2.5	145.9	2.2
DYS07#10	3319	774	0.23	0.04	0.15999	1.73	0.0238	1.51	135.2	19.7	150.7	2.4	151.7	2.3
DYS07#11	4070	263	0.06	0.06	0.16061	1.69	0.0240	1.50	126.2	18.4	151.2	2.4	152.8	2.3
DYS07#12	3025	239	0.08	0.01	0.16184	1.74	0.0238	1.51	165.0	20.1	152.3	2.5	151.5	2.3
DYS07#13	2690	296	0.11	0.04	0.16467	1.76	0.0239	1.51	189.0	20.9	154.8	2.5	152.6	2.3
DYS07#14	2715	427	0.16	0.39	0.14944	2.30	0.0217	1.51	195.1	39.9	141.4	3.0	138.2	2.1
DYS07#15c	339	129	0.38	0.00	0.15414	3.04	0.0226	1.51	166.2	60.6	145.6	4.1	144.3	2.2
DYS07#16r	3354	364	0.11	0.03	0.16363	1.71	0.0239	1.51	174.7	18.6	153.9	2.4	152.5	2.3
DYS07#16c	726	287	0.40	0.03	0.16227	2.41	0.0239	1.61	156.2	41.6	152.7	3.4	152.5	2.4
DYS07#17	2776	88	0.03	0.02	0.15623	1.83	0.0231	1.51	147.4	24.0	147.4	2.5	147.4	2.2
DYS07#18	2683	105	0.04	0.02	0.15656	1.76	0.0231	1.50	154.1	21.5	147.7	2.4	147.3	2.2
DYS07#19r	6341	793	0.13	0.03	0.17159	1.63	0.0252	1.51	169.4	14.8	160.8	2.4	160.2	2.4
DYS07#19c	145	95	0.66	0.20	0.95262	2.27	0.1060	1.51	779.2	35.2	679.5	11.3	649.8	9.4
DYS07#20	5820	518	0.09	0.01	0.15894	1.69	0.0233	1.50	175.0	18.2	149.8	2.4	148.2	2.2
DYS07#21	5384	916	0.17	1.24	0.15518	2.05	0.0236	1.50	81.2	32.7	146.5	2.8	150.5	2.2
DYS07#22	2926	94	0.03	0.13	0.15728	1.81	0.0231	1.50	165.1	23.4	148.3	2.5	147.3	2.2

Supplementary Table S2

Zircon Hf isotopic data of the granitic samples

Sample#spot	$^{176}\text{Yb}/^{177}\text{Hf}$	$^{176}\text{Lu}/^{177}\text{Hf}$	$^{176}\text{Hf}/^{177}\text{Hf}$	2σ	$^{176}\text{Hf}/^{177}\text{Hf}_i$	Age (Ma)	$\varepsilon_{\text{Hf}}(t)$	2σ	T_{DM} (Ma)	T_{DM}^{C} (Ma)	$f_{\text{Lu/Hf}}$
JH155											
JH155#01	0.0224	0.0008	0.282409	0.000021	0.282407	128	-10.1	0.7	1187	1823	-0.97
JH155#02	0.0233	0.0009	0.282481	0.000019	0.282478	128	-7.6	0.7	1089	1663	-0.97
JH155#03	0.0130	0.0005	0.282450	0.000021	0.282449	128	-8.6	0.7	1119	1728	-0.99
JH155#04	0.0137	0.0005	0.282480	0.000020	0.282479	128	-7.6	0.7	1078	1662	-0.98
JH155#05	0.0246	0.0009	0.282420	0.000020	0.282418	128	-9.7	0.7	1173	1798	-0.97
JH155#06	0.0227	0.0009	0.282465	0.000020	0.282463	128	-8.1	0.7	1110	1698	-0.97
JH155#07	0.0148	0.0005	0.282482	0.000021	0.282481	128	-7.5	0.7	1076	1658	-0.98
JH155#08	0.0211	0.0008	0.282475	0.000021	0.282473	128	-7.8	0.7	1094	1676	-0.98
JH155#09	0.0182	0.0007	0.282437	0.000019	0.282436	128	-9.1	0.7	1144	1758	-0.98
JH155#10	0.0238	0.0009	0.282473	0.000018	0.282471	128	-7.8	0.6	1099	1680	-0.97
JH155#11	0.0129	0.0005	0.282500	0.000024	0.282499	128	-6.9	0.9	1051	1619	-0.98
JH155#12	0.0165	0.0006	0.282445	0.000019	0.282444	128	-8.8	0.7	1130	1740	-0.98
JH155#13	0.0169	0.0007	0.282454	0.000018	0.282453	128	-8.5	0.6	1119	1721	-0.98
JH155#14	0.0170	0.0006	0.282461	0.000019	0.282460	128	-8.2	0.7	1107	1704	-0.98
JH155#15	0.0231	0.0008	0.282472	0.000018	0.282470	128	-7.9	0.6	1100	1683	-0.97
JH155#16	0.0203	0.0008	0.282456	0.000019	0.282454	128	-8.4	0.7	1119	1717	-0.98
JH155#17	0.0148	0.0006	0.282492	0.000019	0.282490	128	-7.2	0.7	1064	1637	-0.98
JH155#18	0.0166	0.0006	0.282447	0.000022	0.282446	128	-8.7	0.8	1128	1737	-0.98
DM35											
DM35#01	0.0393	0.0012	0.282463	0.000021	0.282460	151	-7.7	0.8	1122	1692	-0.97
DM35#02	0.0279	0.0009	0.282541	0.000020	0.282539	151	-4.9	0.7	1003	1515	-0.97
DM35#03	0.0293	0.0009	0.282529	0.000019	0.282527	151	-5.4	0.7	1021	1542	-0.97
DM35#04	0.0270	0.0008	0.282517	0.000018	0.282515	151	-5.8	0.7	1036	1569	-0.97
DM35#05	0.0316	0.0010	0.282548	0.000021	0.282545	151	-4.7	0.7	997	1501	-0.97
DM35#06	0.0266	0.0008	0.282483	0.000018	0.282481	151	-7.0	0.6	1083	1645	-0.98
DM35#07	0.0226	0.0007	0.282440	0.000019	0.282438	151	-8.5	0.7	1139	1739	-0.98

DM35#08	0.0220	0.0007	0.282533	0.000018	0.282531	151	-5.2	0.6	1011	1533	-0.98
DM35#09	0.0304	0.0009	0.282547	0.000019	0.282544	151	-4.7	0.7	997	1503	-0.97
DM35#10	0.0249	0.0008	0.282543	0.000022	0.282540	151	-4.9	0.8	1000	1511	-0.98
DM35#11	0.0268	0.0009	0.282549	0.000018	0.282547	151	-4.6	0.6	992	1497	-0.97
DM35#12	0.0231	0.0008	0.282456	0.000020	0.282454	151	-7.9	0.7	1120	1705	-0.98
DM35#13	0.0261	0.0009	0.282544	0.000019	0.282542	151	-4.8	0.7	999	1508	-0.97
DM35#14	0.0303	0.0010	0.282504	0.000021	0.282501	151	-6.3	0.7	1059	1599	-0.97
DM35#15	0.0278	0.0009	0.282528	0.000020	0.282525	151	-5.4	0.7	1023	1546	-0.97

QD69

QD69#01	0.0355	0.0013	0.282506	0.000015	0.282503	151	-6.2	0.5	1063	1595	-0.96
QD69#02	0.0319	0.0011	0.282533	0.000015	0.282530	151	-5.2	0.5	1020	1534	-0.97
QD69#03	0.0288	0.0011	0.282514	0.000015	0.282511	151	-5.9	0.5	1048	1578	-0.97
QD69#04	0.0308	0.0011	0.282533	0.000012	0.282530	151	-5.3	0.4	1021	1536	-0.97
QD69#05	0.0295	0.0010	0.282503	0.000015	0.282500	151	-6.3	0.5	1060	1601	-0.97
QD69#06	0.0402	0.0014	0.282548	0.000014	0.282544	151	-4.8	0.5	1007	1504	-0.96
QD69#07	0.0490	0.0018	0.282530	0.000016	0.282525	151	-5.4	0.6	1043	1545	-0.95
QD69#08	0.0229	0.0008	0.282534	0.000012	0.282532	151	-5.2	0.4	1011	1530	-0.98
QD69#09	0.0355	0.0013	0.282477	0.000015	0.282473	151	-7.3	0.5	1106	1662	-0.96
QD69#10	0.0415	0.0015	0.282394	0.000014	0.282370	837	4.3	0.5	1230	1465	-0.95
QD69#11	0.0520	0.0016	0.282532	0.000016	0.282527	151	-5.3	0.5	1037	1541	-0.95
QD69#12	0.0366	0.0013	0.282511	0.000014	0.282507	151	-6.1	0.5	1056	1585	-0.96
QD69#13	0.0595	0.0020	0.282564	0.000019	0.282558	151	-4.3	0.7	1001	1472	-0.94
QD69#14	0.0418	0.0014	0.282521	0.000014	0.282517	151	-5.7	0.5	1046	1564	-0.96
QD69#15	0.0251	0.0008	0.282506	0.000011	0.282503	151	-6.2	0.4	1052	1594	-0.98
QD69#16	0.0102	0.0004	0.282525	0.000013	0.282524	151	-5.5	0.5	1013	1549	-0.99
QD69#17	0.0434	0.0014	0.282497	0.000015	0.282493	151	-6.6	0.5	1082	1618	-0.96
QD69#18	0.0369	0.0012	0.282528	0.000014	0.282524	151	-5.5	0.5	1031	1548	-0.96
QD69#19	0.0424	0.0014	0.282504	0.000015	0.282500	151	-6.3	0.5	1072	1602	-0.96

DYS14

DYS14#01	0.0183	0.0007	0.282488	0.000015	0.282485	149	-6.9	0.5	1075	1635	-0.98
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DYS14#02	0.0223	0.0009	0.282497	0.000016	0.282495	149	-6.5	0.6	1066	1615	-0.97
DYS14#03	0.0187	0.0008	0.282491	0.000018	0.282489	149	-6.7	0.6	1071	1627	-0.98
DYS14#04	0.0199	0.0008	0.282492	0.000016	0.282489	149	-6.7	0.6	1070	1626	-0.98
DYS14#05	0.0220	0.0009	0.282490	0.000017	0.282488	149	-6.8	0.6	1076	1630	-0.97
DYS14#06	0.0179	0.0007	0.282510	0.000016	0.282508	149	-6.1	0.6	1043	1585	-0.98
DYS14#07	0.0158	0.0006	0.282438	0.000017	0.282436	149	-8.6	0.6	1141	1745	-0.98
DYS14#08	0.0210	0.0009	0.282517	0.000016	0.282515	149	-5.8	0.6	1037	1570	-0.97
DYS14#09	0.0183	0.0007	0.282524	0.000017	0.282522	149	-5.6	0.6	1023	1553	-0.98
DYS14#10	0.0164	0.0007	0.282490	0.000016	0.282489	149	-6.8	0.6	1068	1628	-0.98
DYS14#11	0.0187	0.0008	0.282501	0.000016	0.282499	149	-6.4	0.6	1057	1606	-0.98
DYS14#12	0.0567	0.0020	0.282381	0.000016	0.282350	819	3.2	0.6	1265	1522	-0.94
DYS14#13	0.0210	0.0008	0.282505	0.000015	0.282503	149	-6.3	0.5	1052	1596	-0.98
DYS14#14	0.0259	0.0010	0.282475	0.000017	0.282472	149	-7.3	0.6	1101	1665	-0.97
DYS14#15	0.0201	0.0008	0.282505	0.000016	0.282503	149	-6.3	0.6	1053	1597	-0.98
DYS14#16	0.0204	0.0008	0.282508	0.000015	0.282506	149	-6.2	0.5	1048	1590	-0.98
DYS14#17	0.0184	0.0008	0.282500	0.000017	0.282498	149	-6.4	0.6	1058	1607	-0.98

QD49

QD49#01	0.0120	0.0004	0.282507	0.000015	0.282506	132	-6.5	0.5	1038	1600	-0.99
QD49#02	0.0196	0.0006	0.282526	0.000020	0.282525	132	-5.9	0.7	1017	1558	-0.98
QD49#03	0.0262	0.0008	0.282495	0.000017	0.282493	132	-7.0	0.6	1066	1629	-0.98
QD49#04	0.0089	0.0003	0.282487	0.000018	0.282486	132	-7.2	0.6	1063	1645	-0.99
QD49#05	0.0223	0.0008	0.282501	0.000019	0.282499	132	-6.8	0.7	1056	1615	-0.98
QD49#06	0.0078	0.0003	0.282466	0.000021	0.282466	132	-7.9	0.7	1090	1689	-0.99
QD49#07	0.0167	0.0005	0.282499	0.000020	0.282498	132	-6.8	0.7	1053	1618	-0.98
QD49#08	0.0124	0.0004	0.282497	0.000019	0.282496	132	-6.9	0.7	1052	1621	-0.99
QD49#09	0.0186	0.0006	0.282589	0.000018	0.282587	132	-3.6	0.7	930	1418	-0.98
QD49#10	0.0344	0.0012	0.282376	0.000020	0.282358	831	3.7	0.7	1244	1497	-0.97
QD49#11	0.0095	0.0003	0.282524	0.000020	0.282523	132	-5.9	0.7	1013	1561	-0.99
QD49#12	0.0150	0.0005	0.282533	0.000020	0.282532	132	-5.6	0.7	1004	1542	-0.99
QD49#13	0.0376	0.0012	0.282320	0.000024	0.282300	848	2.0	0.8	1325	1616	-0.96

QD49#14	0.0208	0.0007	0.282536	0.000020	0.282535	132	-5.5	0.7	1005	1535	-0.98
QD49#15	0.0099	0.0003	0.282489	0.000025	0.282488	132	-7.2	0.9	1062	1640	-0.99
QD49#16	0.0507	0.0018	0.282559	0.000020	0.282555	132	-4.8	0.7	1001	1490	-0.95
QD49#17	0.0115	0.0004	0.282477	0.000021	0.282476	132	-7.6	0.8	1079	1667	-0.99
QD49#18	0.0404	0.0014	0.282186	0.000032	0.282152	1246	5.7	1.1	1522	1691	-0.96

DYS62

QD62#01	0.0512	0.0018	0.282446	0.000016	0.282442	127	-8.9	0.6	1164	1746	-0.95
QD62#02	0.0184	0.0007	0.282484	0.000015	0.282482	127	-7.5	0.5	1078	1656	-0.98
QD62#03	0.0581	0.0020	0.282481	0.000017	0.282476	127	-7.7	0.6	1122	1670	-0.94
QD62#04	0.0078	0.0003	0.282473	0.000012	0.282472	127	-7.8	0.4	1081	1678	-0.99
QD62#05	0.0457	0.0017	0.282498	0.000015	0.282494	127	-7.1	0.5	1087	1630	-0.95
QD62#06	0.0513	0.0018	0.282450	0.000015	0.282446	127	-8.8	0.5	1159	1736	-0.95
QD62#07	0.0452	0.0016	0.282485	0.000018	0.282481	127	-7.5	0.6	1103	1658	-0.95
QD62#08	0.0412	0.0015	0.282521	0.000017	0.282517	127	-6.2	0.6	1048	1577	-0.96
QD62#09	0.0121	0.0004	0.282496	0.000016	0.282495	127	-7.0	0.6	1055	1628	-0.99
QD62#10	0.0124	0.0004	0.282481	0.000014	0.282480	127	-7.6	0.5	1075	1661	-0.99
QD62#11	0.0119	0.0004	0.282482	0.000014	0.282481	127	-7.5	0.5	1073	1658	-0.99
QD62#12	0.0215	0.0008	0.282488	0.000015	0.282486	127	-7.3	0.5	1076	1648	-0.98
QD62#13	0.0501	0.0018	0.282500	0.000015	0.282496	127	-7.0	0.5	1087	1626	-0.95
QD62#14	0.0381	0.0014	0.282492	0.000014	0.282489	127	-7.2	0.5	1086	1641	-0.96
QD62#15	0.0260	0.0009	0.282424	0.000016	0.282422	127	-9.6	0.6	1167	1789	-0.97
QD62#16	0.0382	0.0014	0.282513	0.000016	0.282510	127	-6.5	0.6	1056	1594	-0.96
QD62#17	0.0204	0.0008	0.282450	0.000014	0.282448	127	-8.7	0.5	1128	1731	-0.98
QD62#18	0.0171	0.0006	0.282497	0.000014	0.282495	127	-7.0	0.5	1059	1627	-0.98
QD62#19	0.0383	0.0014	0.282476	0.000016	0.282473	127	-7.8	0.6	1109	1677	-0.96

DYS01

DYS01#01	0.0133	0.0005	0.282399	0.000016	0.282397	147	-10.0	0.6	1191	1833	-0.98
DYS01#02	0.0182	0.0007	0.282442	0.000019	0.282441	147	-8.5	0.7	1136	1736	-0.98
DYS01#03	0.0019	0.0001	0.282431	0.000015	0.282431	147	-8.8	0.5	1133	1758	-1.00
DYS01#04	0.0175	0.0007	0.282389	0.000013	0.282387	147	-10.4	0.5	1210	1855	-0.98

DYS01#05	0.0039	0.0001	0.282327	0.000017	0.282327	147	-12.5	0.6	1277	1988	-1.00
DYS01#06	0.0319	0.0011	0.282380	0.000014	0.282377	147	-10.7	0.5	1236	1877	-0.97
DYS01#07	0.0207	0.0007	0.282365	0.000014	0.282363	147	-11.2	0.5	1244	1908	-0.98
DYS01#08	0.0190	0.0007	0.282384	0.000016	0.282382	147	-10.6	0.6	1217	1867	-0.98
DYS01#09	0.0288	0.0010	0.282409	0.000014	0.282406	147	-9.7	0.5	1193	1813	-0.97
DYS01#10	0.0397	0.0013	0.282392	0.000021	0.282388	147	-10.3	0.7	1227	1852	-0.96
DYS01#11	0.0076	0.0003	0.282474	0.000018	0.282473	147	-7.3	0.6	1080	1663	-0.99
DYS01#12	0.0450	0.0016	0.282387	0.000020	0.282362	836	3.9	0.7	1243	1485	-0.95
DYS01#13	0.0444	0.0016	0.282158	0.000020	0.282133	838	-4.1	0.7	1569	2000	-0.95
DYS01#14	0.0155	0.0006	0.282424	0.000014	0.282422	147	-9.1	0.5	1159	1777	-0.98
DYS01#15	0.0200	0.0007	0.282393	0.000018	0.282391	147	-10.2	0.7	1206	1846	-0.98
DYS01#16	0.0493	0.0020	0.282610	0.000024	0.282579	839	11.7	0.9	934	989	-0.94
DYS01#17	0.0437	0.0016	0.282504	0.000018	0.282499	147	-6.4	0.6	1078	1606	-0.95
DYS01#18	0.0159	0.0006	0.282351	0.000017	0.282350	147	-11.7	0.6	1260	1938	-0.98
DYS01#19	0.0156	0.0006	0.282348	0.000025	0.282346	147	-11.8	0.9	1264	1946	-0.98
DYS01#20	0.0192	0.0007	0.282365	0.000018	0.282363	147	-11.2	0.6	1244	1908	-0.98
DYS01#21	0.0196	0.0007	0.282429	0.000014	0.282427	147	-9.0	0.5	1155	1767	-0.98

DYS03

DYS03#01	0.0142	0.0005	0.282388	0.000014	0.282386	146	-10.4	0.5	1206	1857	-0.99
DYS03#02	0.0163	0.0005	0.282406	0.000025	0.282404	146	-9.8	0.9	1183	1818	-0.98
DYS03#03	0.0347	0.0013	0.282462	0.000017	0.282459	146	-7.9	0.6	1126	1696	-0.96
DYS03#04	0.0269	0.0009	0.282421	0.000025	0.282418	146	-9.3	0.9	1172	1786	-0.97
DYS03#05	0.0377	0.0013	0.282484	0.000024	0.282480	146	-7.1	0.9	1097	1649	-0.96
DYS03#06	0.0207	0.0007	0.282453	0.000024	0.282451	146	-8.2	0.9	1122	1714	-0.98
DYS03#07	0.0316	0.0011	0.282454	0.000024	0.282451	146	-8.1	0.9	1133	1713	-0.97
DYS03#08	0.0296	0.0011	0.282517	0.000029	0.282514	146	-5.9	1.0	1042	1573	-0.97
DYS03#09	0.0115	0.0004	0.282410	0.000024	0.282409	146	-9.6	0.8	1172	1808	-0.99
DYS03#10	0.0051	0.0002	0.282354	0.000017	0.282354	146	-11.6	0.6	1242	1929	-0.99
DYS03#11	0.0156	0.0005	0.282384	0.000015	0.282383	146	-10.6	0.5	1212	1865	-0.98
DYS03#12	0.0299	0.0011	0.282396	0.000016	0.282393	146	-10.2	0.6	1213	1843	-0.97

DYS03#13	0.0252	0.0009	0.282432	0.000014	0.282430	146	-8.9	0.5	1157	1761	-0.97
DYS03#14	0.0367	0.0014	0.282452	0.000018	0.282448	146	-8.2	0.7	1143	1719	-0.96
DYS03#15	0.0184	0.0006	0.282396	0.000017	0.282394	146	-10.2	0.6	1199	1840	-0.98
DYS03#16	0.0229	0.0008	0.282456	0.000015	0.282454	146	-8.0	0.5	1120	1707	-0.98
DYS03#17	0.0255	0.0011	0.282361	0.000015	0.282358	146	-11.4	0.5	1262	1921	-0.97

DYS06

DYS06#01	0.0171	0.0006	0.282359	0.000020	0.282357	150	-11.4	0.7	1250	1920	-0.98
DYS06#02	0.0133	0.0005	0.282342	0.000016	0.282340	150	-12.0	0.6	1268	1957	-0.99
DYS06#03	0.0150	0.0005	0.282356	0.000014	0.282355	150	-11.5	0.5	1249	1925	-0.99
DYS06#04	0.0193	0.0007	0.282440	0.000017	0.282439	150	-8.5	0.6	1138	1739	-0.98
DYS06#05	0.0254	0.0009	0.282377	0.000021	0.282375	150	-10.8	0.7	1234	1881	-0.97
DYS06#06	0.0143	0.0005	0.282375	0.000016	0.282373	150	-10.8	0.6	1225	1884	-0.98
DYS06#07	0.0140	0.0005	0.282381	0.000018	0.282380	150	-10.6	0.7	1215	1869	-0.99
DYS06#08	0.0126	0.0004	0.282330	0.000020	0.282329	150	-12.4	0.7	1284	1982	-0.99
DYS06#09	0.0178	0.0006	0.282409	0.000021	0.282407	150	-9.6	0.8	1180	1808	-0.98
DYS06#10	0.0108	0.0004	0.282376	0.000018	0.282375	150	-10.8	0.7	1219	1881	-0.99
DYS06#11	0.0213	0.0008	0.282419	0.000018	0.282417	150	-9.3	0.6	1171	1787	-0.98
DYS06#12	0.0140	0.0005	0.282380	0.000018	0.282379	150	-10.6	0.6	1216	1871	-0.99
DYS06#13	0.0137	0.0005	0.282375	0.000021	0.282373	150	-10.8	0.8	1223	1884	-0.99
DYS06#14	0.0582	0.0022	0.282421	0.000020	0.282415	150	-9.3	0.7	1212	1791	-0.94
DYS06#15	0.0187	0.0007	0.282395	0.000021	0.282393	150	-10.1	0.7	1202	1840	-0.98
DYS06#16	0.0123	0.0004	0.282437	0.000022	0.282436	150	-8.6	0.8	1134	1745	-0.99
DYS06#17	0.0041	0.0001	0.282157	0.000020	0.282155	826	-3.6	0.7	1509	1957	-1.00
DYS06#18	0.0319	0.0011	0.282396	0.000022	0.282392	150	-10.1	0.8	1214	1842	-0.97
DYS06#19	0.0142	0.0005	0.282354	0.000024	0.282352	150	-11.6	0.9	1254	1931	-0.98
DYS06#20	0.0138	0.0005	0.282339	0.000023	0.282338	150	-12.1	0.8	1273	1963	-0.99
DYS06#21	0.0161	0.0006	0.282382	0.000020	0.282380	150	-10.6	0.7	1216	1869	-0.98
DYS06#22	0.0322	0.0012	0.282416	0.000019	0.282413	150	-9.4	0.7	1188	1796	-0.96

DYS07

DYS07#01	0.0281	0.0010	0.282524	0.000013	0.282522	149	-5.6	0.5	1029	1554	-0.97
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DYS07#02	0.0269	0.0009	0.282533	0.000013	0.282530	149	-5.3	0.5	1017	1535	-0.97
DYS07#03	0.0374	0.0014	0.282558	0.000016	0.282554	149	-4.4	0.6	993	1481	-0.96
DYS07#04	0.0518	0.0018	0.282514	0.000020	0.282509	149	-6.0	0.7	1067	1582	-0.95
DYS07#05	0.0228	0.0008	0.282533	0.000014	0.282530	149	-5.3	0.5	1014	1534	-0.97
DYS07#06	0.0205	0.0007	0.282541	0.000013	0.282539	149	-4.9	0.5	1000	1515	-0.98
DYS07#07	0.0085	0.0003	0.282532	0.000018	0.282531	149	-5.2	0.6	1001	1533	-0.99
DYS07#08	0.0368	0.0013	0.282512	0.000018	0.282508	149	-6.0	0.6	1056	1584	-0.96
DYS07#09	0.0058	0.0002	0.282486	0.000016	0.282485	149	-6.9	0.6	1062	1636	-0.99
DYS07#10	0.0110	0.0004	0.282536	0.000020	0.282535	149	-5.1	0.7	998	1524	-0.99
DYS07#11	0.0349	0.0012	0.282495	0.000018	0.282492	149	-6.6	0.6	1077	1620	-0.96
DYS07#12	0.0523	0.0017	0.282579	0.000019	0.282574	149	-3.7	0.7	972	1436	-0.95
DYS07#13	0.0350	0.0013	0.282564	0.000017	0.282560	149	-4.2	0.6	983	1468	-0.96
DYS07#14	0.0377	0.0013	0.282535	0.000017	0.282532	149	-5.2	0.6	1023	1532	-0.96
DYS07#15	0.0150	0.0006	0.282524	0.000016	0.282522	149	-5.5	0.6	1021	1553	-0.98
DYS07#16	0.0227	0.0008	0.282509	0.000016	0.282507	149	-6.1	0.6	1047	1587	-0.98
DYS07#17	0.0297	0.0010	0.282499	0.000014	0.282496	149	-6.5	0.5	1067	1612	-0.97
DYS07#18	0.0050	0.0002	0.282530	0.000018	0.282529	149	-5.3	0.7	1001	1537	-0.99
DYS07#19	0.0217	0.0008	0.282534	0.000015	0.282532	149	-5.2	0.5	1010	1531	-0.98
DYS07#20	0.0238	0.0008	0.282558	0.000016	0.282555	149	-4.4	0.6	979	1479	-0.98
DYS07#21	0.0165	0.0006	0.282547	0.000016	0.282546	149	-4.7	0.5	988	1500	-0.98
