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Tracing interactions between natural argillites and hyper-alkaline fluids from engineered cement paste and concrete: Chemical and isotopic monitoring of a 15-years old deep-disposal analogue

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ABSTRACT

Samples of Toarcian argillite were collected both next to and far from a CEM II cement paste and a CEM II concrete, within the specific context of a 15-a old borehole located in the Tournemire Experimental Platform (Aveyron, France). The objectives were evaluation of the mineralogical and geochemical changes of the claystone at the contact with the cementitious materials and determination of the spatial extent of the interactions. The approach includes the examination of the mineralogical (XRD, SEM, TEM), chemical (major, trace, rare earth elements) and isotopic (Sr, C, O) compositions of argillite whole-rocks and of various soluble phases, at two scales: in the rock matrix (P1 scale) and along micro-cracks (P2 scale). The two study scales outline nearly similar mineralogical modifications, shown by the presence of Ca silicate hydrates (C-S-H) and newly-formed CaCO₃ within 10–15 mm of the cement paste and concrete. Chemical data from whole-rock argillites indicate few changes in a slightly thicker zone (18–20 mm), mainly consisting of an increase in the CaO wt.%, and a decrease in Sr contents. The other elementary contents remained quite constant except for MgO, which suggests redistribution with precipitation of a Mg-rich mineral phase at 20 mm from cement paste/concrete interface. Acetic acid leachates had more pronounced variations, including a decrease of the total elementary content in the same 'geochemical disturbed zone' (GDZ), together with a significant increase of the Sr isotopic ratios. A combination of Sr and C/O isotopic patterns was used to distinguish the behavior of secondary cementitious phases in the clay-rich rock: (i) calcite dissolution and re-precipitation is supported by C/O isotopic data and (ii) C-S-H neoformation is evidenced by the ⁸⁷Sr/⁸⁶Sr ratios; this tool also contributes to determine the origin of the fluids. The proportion of newly-formed C-S-H in the matrix and in the micro-cracks of the argillite is modeled.

1. Introduction

Deep geological disposal of high-activity and long-period radioactive wastes in France is designed with a confinement system based on the multiplication of argillaceous and cement-bearing barriers called 'engineered barriers'. The role of these barriers is to avoid the release of radio-elements into the biosphere, as well as to prevent the potential addition of external fluids to the waste materials. In a deep clay-rich medium, cementitious materials are expected to account for most of the building structures, the packages containing waste, as well as the cell grouts or plugs. These cementitious materials will be emplaced at the immediate contact with the natural argillaceous formation and/or the bentonite plugs.

Hydrolysis of cementitious phases is known to produce hyper-alkaline pore fluids with pHs ranging from 10 to 13.5 (Atkinson et al., 1985; Berner, 1992; Hewlett, 1998), whereas natural argillaceous settings are characterized by a nearly neutral pH. Numerous studies have been performed in the last decades to assess interactions in such a chemical disequilibrium context, with the specific aim of identifying their effect on the confinement material properties. These studies have been based mainly on evaluation of mineralogical and petrophysical changes during propagation of high-pH fluids in claystones, either by laboratory experiments (Ramirez et al., 2002, 2005; Fernandez et al., 2009b; Dauzères et al., 2009, 2010a), computer modeling (Savage et al., 2002; De Windt et al., 2004; Gaucher and Blanc, 2006; Fernandez et al., 2009a), or characterization of natural and engineered analogues (Khoury et al., 1985; Linklater et al., 1996; Milodowski et al., 2001; Techer et al., 2006; Tinseau et al., 2006; Fourcade et al., 2007; Elie et al., 2007). They

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have allowed identifying the mineral transformations that have occurred in claystones and the cementitious materials during interaction with alkaline fluids over various time sequences.

In order to better constrain the reaction processes induced by alkaline fluids/argillite interaction and to specify the geometry of the system, this study will focus on a combined acquisition of mineralogical and geochemical data from contact zones located between a cement paste or a concrete and a natural argillite after a 15-a interaction. This approach is based on the characterization of an engineered analogue that was described preliminarily by Tinseau et al. (2006) and modeled by De Windt et al. (2008). This engineered analogue, called the DM borehole, is located in the Tournemire Experimental Platform (Aveyron, France) and is to a 15-a old borehole crosscutting the Toarcian argillite formation and sealed with concrete. The over-coring of the borehole exhibits the natural claystone in contact with a CEM II hardened cement paste or with a CEM II concrete. A mineralogical and geochemical characterization of samples collected either next to or away from the cement-bearing materials, at various depths in the borehole is presented. In combination with SEM and TEM investigations, Sr, O and C isotope systems are used to constrain the propagation of alkaline fluids generated by the cement into the clay medium: limited alkaline fluid/rock interactions that cannot be identified

by mineralogical observations due to the relative high detection limit will be discussed on the basis of isotope modeling based on distinguishable C and Sr isotope compositions in the clays and in the cement paste and concrete (Techer et al., 2006; Fourcade et al., 2007).

2. Description of the sampling site and samples

The study area is located in the Tournemire Experimental Platform of the French Institute for Radioprotection and Nuclear Safety (IRSN) located in Aveyron (SE France). The platform is devoted to multi-disciplinary research on the feasibility of deep geological disposal of radioactive wastes. The site was initially constructed between 1882 and 1888 as a train tunnel of several hundred meters beneath the 'Plateau des Causses'. Secondary galleries were then excavated horizontally from the tunnel between 1996 and 2008. The underground structures crosscut a 200-m thick argillite Toarcian formation from South to North which is composed mainly of phyllosilicates (40–50 wt.% of illite, chlorite, kaolinite and illite/smectite mixed-layers), calcite, dolomite, K and Na–Ca feldspars, pyrite and quartz (Boisson et al., 2001; Tinseau et al., 2006). In 1990, exploration boreholes were drilled from the tunnel for

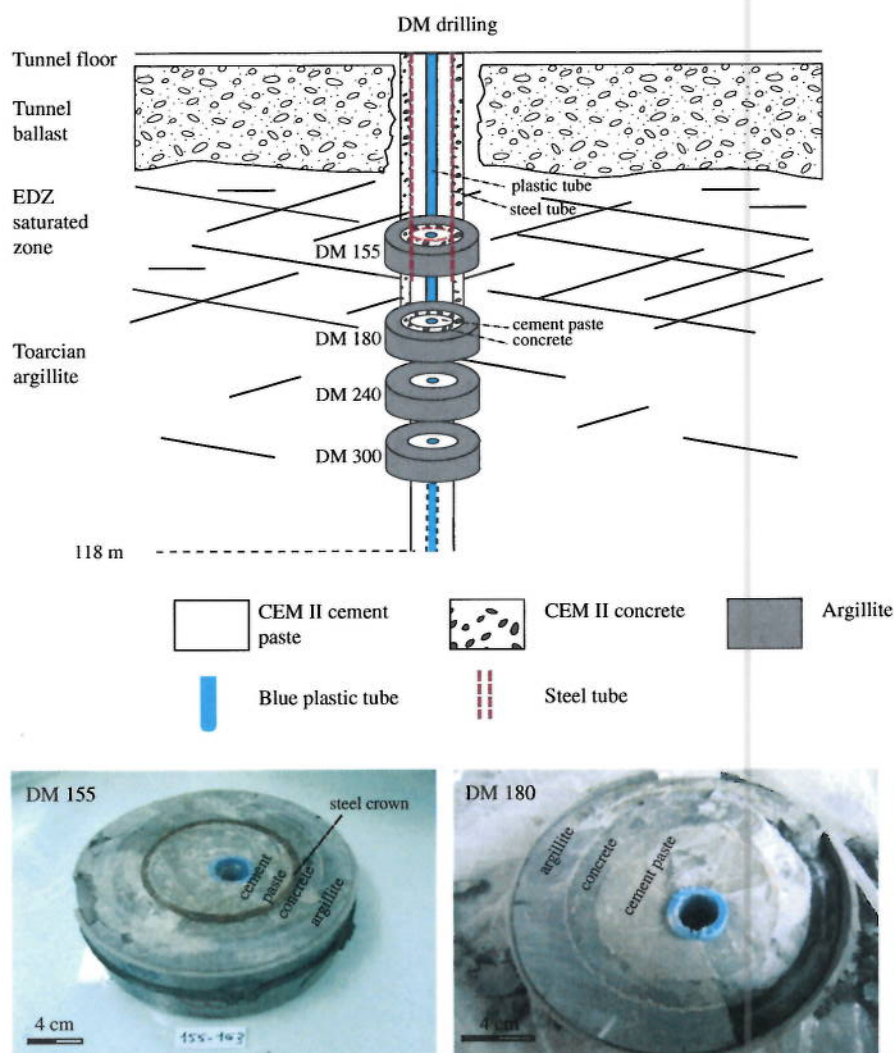


Fig. 1. Scheme of the DM borehole in the Tournemire Experimental Platform. Location of the studied layers along the borehole: DM155, DM180, DM240 and DM300 and representation of the cementitious materials (CEM II concrete or cement paste) in contact with the argillite. The images are from DM155 and DM180 sections.

hydro-geological characterization of the basement formations, in some cases downwards into the Toarcian argillite over a depth interval of several hundred meters. Soon after these hydro-geological measurements, the boreholes were filled with CEM II cement paste and concrete. The deeper parts of the boreholes were entirely filled with cement paste, which was mixed with concrete in the upper meters below the tunnel floor (Fig. 1). Over-coring of these boreholes results in contacts between argillite and cementitious materials that can be expected to be similar in a deep geological waste disposal. Of these boreholes, Tinseau et al. (2006) focused on the DM-borehole that was over-drilled by 3 m, for a mineralogical characterization of the cementitious materials and the claystone in contact. The samples of the present study come from four layers located between 1.55 m and 3.00 m of this over-core (Fig. 1). The four sets of samples are labeled as a function of their depth location: 1.55 m (DM155), 1.80 m (DM180), 2.40 m (DM240) and 3.00 m (DM300). The layers are located in the excavation disturbed zone (EDZ) of the tunnel that can be considered to be pore-fluid saturated as it is drained by the Cernon fault (Fig. 1). The deepest layer (DM300) is, however, expected to be less affected by the EDZ, and thus less saturated.

Each sample set consists of argillite and cement paste or concrete, which were collected along a profile perpendicular to the axis of the core. In sample DM155, the center of the core corresponds to a white cement paste (CEM II-A-LL, Portland cement base material with calcareous wt.% between 6 and 20) surrounded

by a steel crown and a ~1.5 cm thick concrete crown (CEM II) in contact with the argillite (Fig. 1). This sample was considered by De Windt et al. (2008) for reactive transport modeling using the HYTEC code. In sample DM180, the same configuration occurs but without the steel crown (Fig. 1). A 2.5–2.8 cm thick crown of argillite was observed surrounding the concrete in these two samples (Fig. 1). Samples DM240 and DM300 show a core consisting of a fine-grained cement paste (CEM II) surrounded by 5–6 cm of argillite. The cement paste is white colored, shows a fine-grain chalky texture in sample DM240, and is more indurated with inclusion of gray zones in sample DM300. The argillite crowns are thicker in these layers (5.5–6 cm). Argillites collected in the deep layers (DM240, DM300), and far from cement paste/concrete interface, will be considered as undisturbed reference materials.

Mineralogical and geochemical characteristics of argillite located away from or next to the cement paste/concrete interface were investigated at two scales: (1) in the rock matrix (labeled P1) and (2) in micro-cracks crosscutting the claystone perpendicular to the cement paste/concrete interface (labeled P2). These micro-cracks were more numerous in the upper layers of the borehole and very discrete in the deeper ones. Along P1, the argillite was sampled continuously in sections perpendicular to the central axis of the core, every 2–4 mm over about 20 mm. Some samples were collected at the external limit of the profiles as undisturbed references. Although no secondary mineral recrystallization was detected to the naked eye along P2 (even if a bright

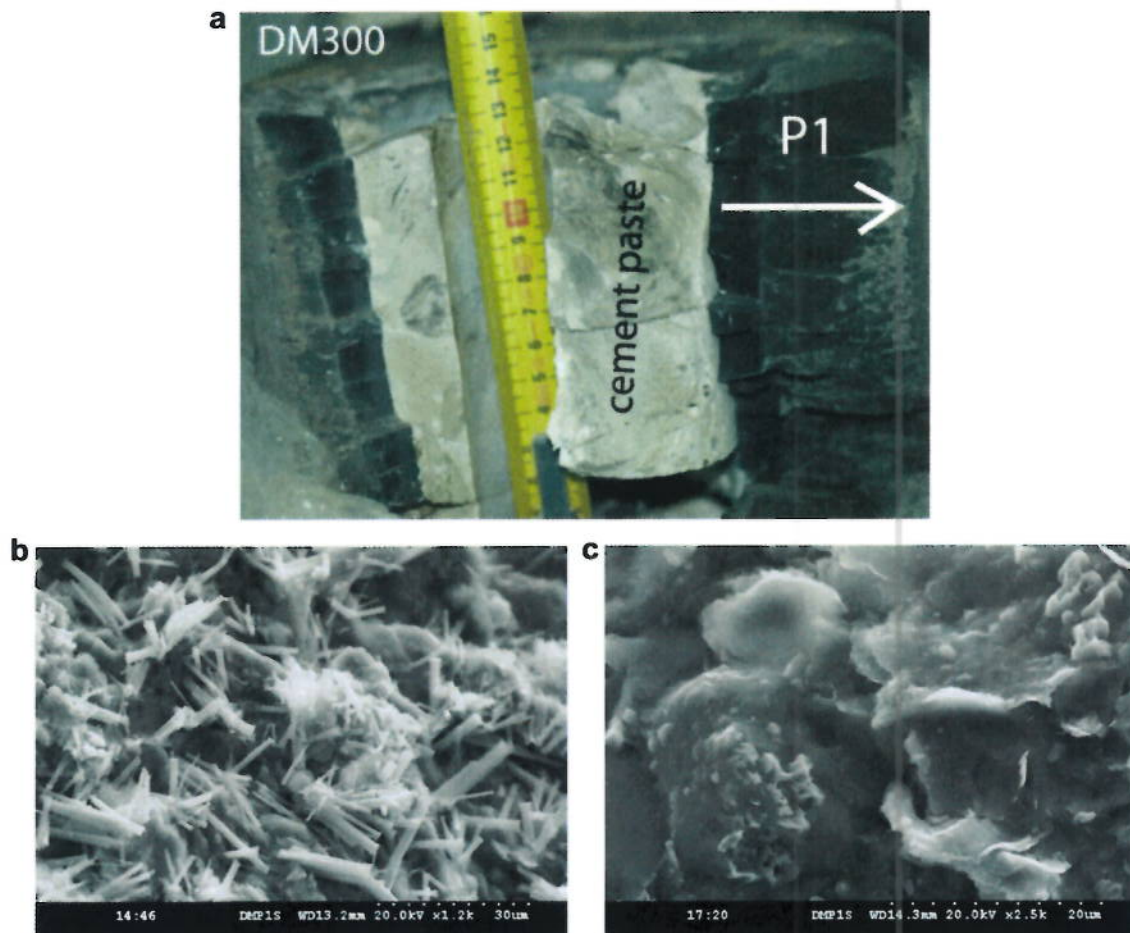
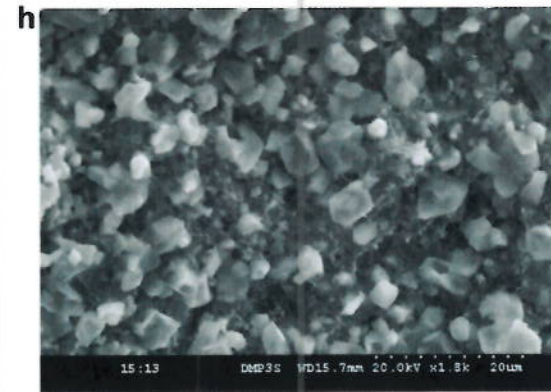
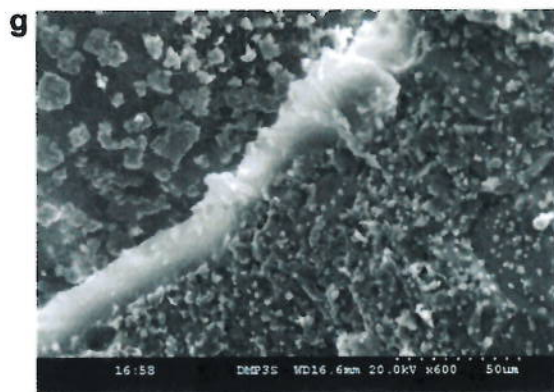
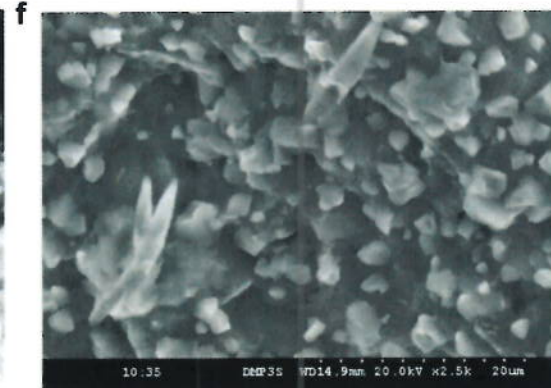
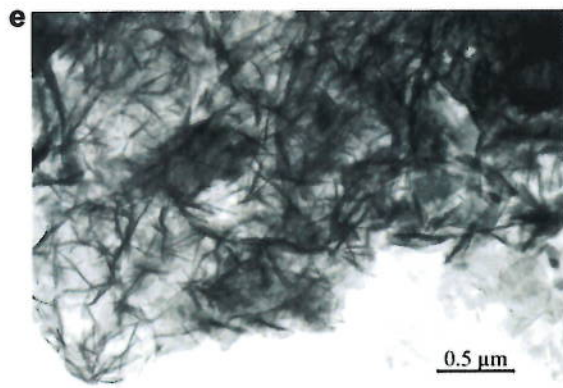
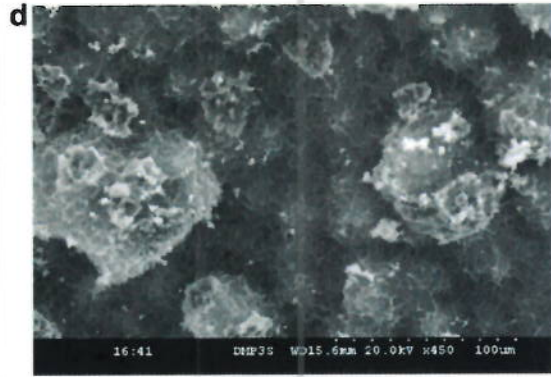
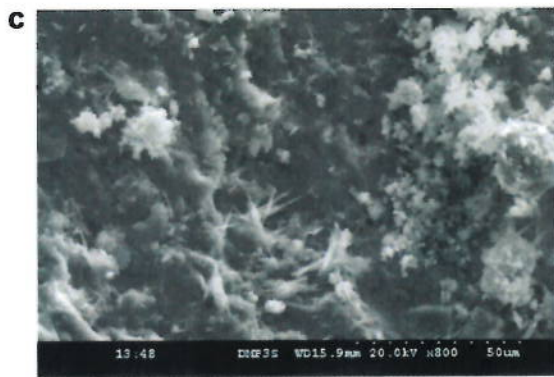
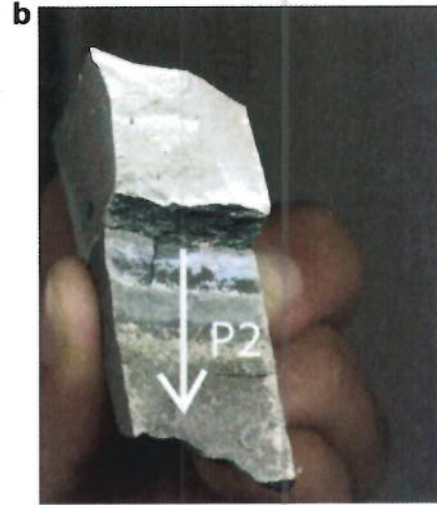
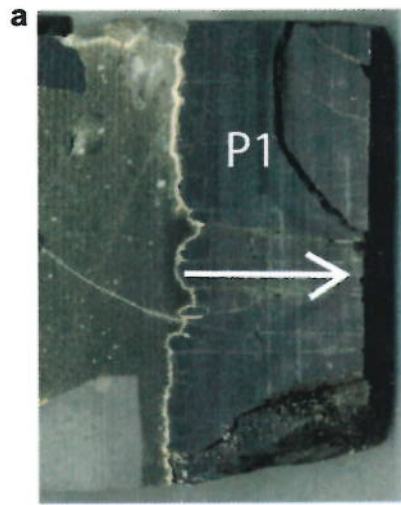


Fig. 2. (a) Macroscopic observation of DM300 section and representation of the P1 studied profile crosscutting the Toarcian argillite in contact with the cement paste; the 15 mm thick dark band is well outlined at the interface of the two materials. (b and c) SEM images showing 'needle-like' ettringite crystals (b) and a Si-Ca rich gel in the dark band matrix (c).



aspect and specific blue colors were noticeable in some cases), the contact surfaces were carefully scraped to collect any discrete secondary minerals. Despite all precautions, the possibility of contamination by host argillite material cannot be ruled out for these samples, especially considering the small amount of infilling materials.

3. Analytical procedures

3.1. Mineralogical characterization

The argillite samples were crushed and sieved for mineral analysis of the whole-rocks and of the <2 μm clay fractions by X-ray diffraction (XRD). Measurements were performed using a Siemens D500 diffractometer operating at 40 kV and 30 mA with a Cu K α radiation (MNHN, Paris, France). Some rock fragments were kept, C-coated and observed by scanning electron microscopy (SEM). The SEM was equipped with an energy-dispersive spectrometry system (Hitachi S-3500, Oxford-Isis EDS) (IRSN, Paris, France) for qualitative overall or localized analyses. Transmission electron microscopy (TEM) was also used to examine the morphology of mineral grains that were added to alcohol suspensions and settled onto perforated carbon grids. The TEM was coupled with an energy dispersive X-ray analyser (TEM-EDX) to determine semi-quantitative chemical compositions (Philips CM20) (SCMEM, Nancy, France). Analyses were carried out in nanoprobe mode with a probe diameter of 10 nm and 40 s counting time. Values for the constant k were determined using standards.

3.2. Whole-rock chemical analyses

The argillite samples collected along the P1 profile of samples DM240 and DM300 were crushed for chemical analyses. The DM155 and DM180 profiles were not investigated at this scale due to their short length. For the whole-rock, the compositions of major, trace and rare-earth elements (REE) were determined by ICP-AES and ICP-MS (SARM, Nancy, France). Approximately 100 mg of ignited rock powders were fused with Li₂BO₄ in an Ar-atmosphere and dissolved in 25 mL of HNO₃-glycerine solution, for analysis. The relative standard deviations for the major, trace and rare-earth elements were 3% and 5% and at 8%, respectively.

3.3. Extraction of soluble phases: chemical and isotopic analyses

In most cases, 30–60 mg of rock or mineral powder were used to determine the total elementary, Sr, O and C isotope compositions. However, in some cases the powder amount was not sufficient for a complete analysis, as for sampling in micro-cracks (P2). For these samples, only Sr isotopic measurements were performed.

For the C and O isotopic procedure, the carbonate fraction (1–10 mg) was allowed to react with 100% orthophosphoric acid at 50 °C (McCrea, 1950) using $\alpha(\text{O}) \text{CO}_2/\text{CaCO}_3$ (extraction) = 1.009306 considering that the carbonate phase consisted entirely of calcite despite identification of some dolomite in the Toarcian argillite matrix (Tinseau et al., 2006). The carbonate

content was determined by measuring the CO₂ partial pressure during the reaction, with an uncertainty at 0.5 wt.%. The CO₂ was then purified from H₂S released during interaction with orthophosphoric acid and the sulfides disseminated in the argillite, by reaction with Ag₃PO₄ for 5 min at ca. 50 °C. The isotopic analyses were carried out on a VG SIRA 10 triple collector mass spectrometer (Géosciences Department, University of Rennes, France), and are reported using the conventional δ notation vs. SMOW (for O) and PDB (for C). Analyses were made at distinct periods so that specific corrections were applied to the results according to replicate analysis of the internal carbonate standard Prolabo Rennes and NBS 19 reference material. The average uncertainties were estimated at <0.1‰ for $\delta^{13}\text{C}$ and at ca. 0.1‰ for $\delta^{18}\text{O}$.

For the Sr isotopic determinations, the soluble material, also referred to as leachate, was extracted by reacting the mineral powders with 1 N acetic acid, with an acid/rock ratio of 0.02 and a leaching time of about 5 min at room temperature. The leachates were then separated by centrifugation and dried before dissolution in HNO₃. Strontium separation was performed on the Eichrom Sr-resin, following the Pin et al. (2003) method. Strontium was loaded on a Ta-filament and the ⁸⁷Sr/⁸⁶Sr isotopic ratio was measured by solid-source thermal ionization mass spectrometry (Finnigan TRITON TI) (University of Nîmes, France). Approximately 90 ratios were measured for each aliquot to ensure an analytical error (2σ) of about 5×10^{-6} . The external reproducibility of the isotopic measurements was controlled by repetitive analysis of the NBS987 standard providing a mean value of $0.710258 \pm 2 \times 10^{-6}$ (2σ).

A second aliquot of the mineral powders was treated by an identical acetic acid leach for the determination of the major elements and Sr. The same 'acid/powder' ratio and reaction time were used for all samples as before. The leachates were analyzed by ICP-AES and ICP-MS following a procedure described by Samuel et al. (1985) (LhyGeS, University of Strasbourg, France). Reproducibility and accuracy tests were performed routinely by a systematic analysis of the BE-N and GL-O standards for all elements of each data set. The relative standard deviation for Ca was 3% for Ca, and 5% for the other elements.

4. Results

The macroscopic observation of the 'cementitious material/argillite' system collected along the DM borehole provides determining information about the scales of transformations. The first optical change observed in the system corresponds to the occurrence of a thin white crust at the concrete/argillite interface (in samples DM155 and DM180), whereas no such crust is detected at the cement paste contact (samples DM240 and DM300). Next to the cement paste and concrete, peculiar features are observed in the argillite in a 10–15-mm thick zone. There, the claystone is characterized by a dark color and an isotropic appearance (Fig. 2a). The dark zone is thicker in the deep layers (15 mm in DM300) than in the upper ones (10 mm in DM180). It is less evident in sample DM155 at the contact with the concrete (DM155) in which 2-mm thick dark bands parallel the concrete interface in the clay-rich matrix; they delineate a supposed disturbed zone about 10 mm thick (Fig. 3a). This upper layer of the borehole also

Fig. 3. (a and b) Macroscopic observations of the DM155 section and representation of the P1 and P2 studied profiles crosscutting the Toarcian argillite in contact with concrete: 2-mm thick dark bands parallel the concrete interface in the clay-rich matrix (a-P1 scale) and a zonation is clearly visible along 10 mm open discontinuities underlined by a specific bluish color (b-P2 scale). (c–h) SEM and TEM images obtained in the various zones of the P2 profile: (c) acicular ettringite crystals in the first zone, (d) C–S–H honeycomb structures in the second zone that was identified by the more pronounced bluish color, (e) these C–S–H phases were characterized by a foil-like morphology by TEM, (f) sub-automorphous calcite and celestine were also observed in this second zone, (g) a smooth Ca–Si rich white band marked the boundary between the second and the third zone, (h) calcite and C–S–H neoformations in the fourth zone. (For interpretation of the references to color in this figure legend, the reader is referred to the web version of this article.)

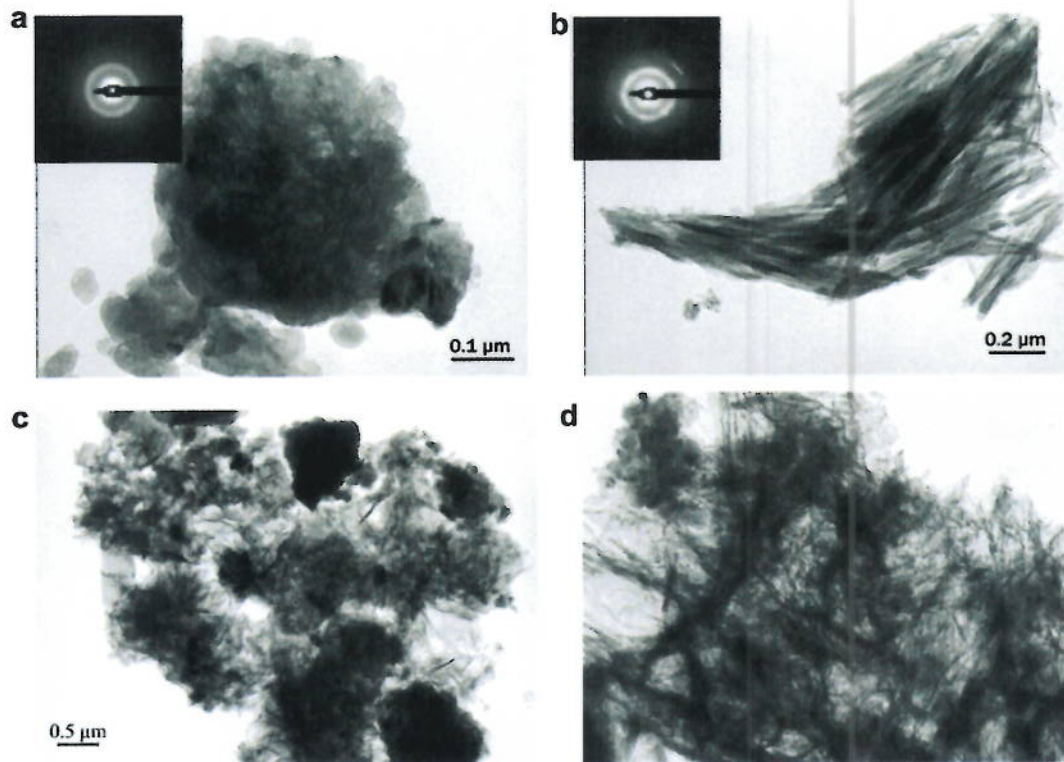


Fig. 4. (a–c) TEM images of C–S–H phases observed in the concrete and cement paste materials: (a) identified with a fluffy appearance (flakes) and diffuse SAED rings in the concrete, (b) characterized either with fiber morphology and a SAED showing a secondary discontinuous ring with a periodicity of 0.18 nm in the DM300 and DM240 cement paste, (c) or with foil-like structures. (d) TEM image of the C–S–H phases observed in the thin crust at the concrete/argillite interface with morphology mixing thin fibers, foil-like and granular forms.

outlines many open discontinuity surfaces corresponding to a decompression process following the claystone coring. Along these plan surfaces (P2 scale), a zonation is clearly visible over about 10 mm (Fig. 3b), underlined by a specific bluish color. The mineralogical and chemical characteristics of so-defined 'cementitious material/argillite' system is described below.

4.1. Cement paste and concrete

The CEM II cement paste and the CEM II concrete display similar XRD patterns and are mainly composed of portlandite, calcite and ettringite. Calcium silicate hydrates (C–S–H), which are typical hydration products of cement, are not detected by XRD in the studied materials. Instead, they are observed by TEM as non-diffracting phases, mainly with 'fluffy flakes' and 'foil-like' morphologies (Fig. 4a–c). Such non-diffracting phases are also identified in the thin crust sampled between the concrete and the argillite, together with ettringite and CaCO_3 . The C–S–H phases appear in this crust with morphologies that mix thin fibers, foil-like forms and granular forms (Fig. 4d). The three isomorphs of CaCO_3 are also detected by XRD (calcite 55%, aragonite 30% and vaterite 15%) in the rim.

The cement paste and the concrete are distinguishable on the basis of their mineral contents and their geochemical characteristics. The CEM II concrete is identified by carbonate contents ranging from 19 to 33 wt.% that are related to the intrinsic properties of the material and its aggregate heterogeneity (Fig. 5). The cement paste is more homogeneous with a mean CaCO_3 content of 39 wt.%, which is expected for a CEM II-A-LL cement base material (Fig. 5). However, one sample of this cement paste from the DM300 layer yields an unexpected composition with a low CaCO_3 value of 26 wt.%. This heterogeneity may result from addition of gray and

indurated materials observed in the layer, probably induced by the infilling of the borehole by the cement.

The $\delta^{13}\text{C}$ values are low in all cementitious materials, ranging from -8.3‰ to -13.2‰ (Table 1). Such data belong to the domain defined for various mortars and cements (Letolle et al., 1992). The DM240 cement paste yields the lowest values (-11.2‰ and -13.2‰), whereas the highest (-8.3‰) is obtained on the gray and indurated part of the DM300 cement paste (Fig. 6). In contrast, the $\delta^{18}\text{O}$ values are identical in the concrete and cement paste, with a mean value of 19.4‰ , except the DM300 cement paste characterized by a low $\delta^{18}\text{O}$ value of 16.5‰ (Table 1, Fig. 6c).

The dilute acid leachates extracted from cement paste and concrete are mainly calcitic, and also contain significant amounts of Si, Al and Fe (Table 3). Concrete and cement paste leachates show distinct compositions: in comparison to the concrete leachates, those of the cement paste are enriched in Ca, Si and Al, yet they are depleted in Sr contents ($230\text{--}340\ \mu\text{g L}^{-1}$ relative to $380\text{--}400\ \mu\text{g L}^{-1}$). As a result, the leachates of the cement paste are more concentrated than those from concrete ($760\text{--}1360\ \text{mg L}^{-1}$ and $635\ \text{mg L}^{-1}$, respectively). The previously mentioned DM300 gray cement paste is quite specific with a leachate enriched in all elements. Similarly, variable Sr isotope ratios are obtained regardless of the nature of the cementitious material (Table 1; Fig. 7): the concrete is identified by high $^{87}\text{Sr}/^{86}\text{Sr}$ ratios ranging between 0.708352 and 0.708380 ($\pm 6 \times 10^{-6}$), whereas the cement paste displays relatively lower ratios (from 0.708122 to 0.708308). The ratios are scattered in the DM240 cement paste and are fairly homogeneous in the DM300 cement paste with a mean value of 0.708260 ($\pm 5 \times 10^{-6}$) (Table 1). No specific composition is seen in the special DM300 gray cement paste sample mentioned above.

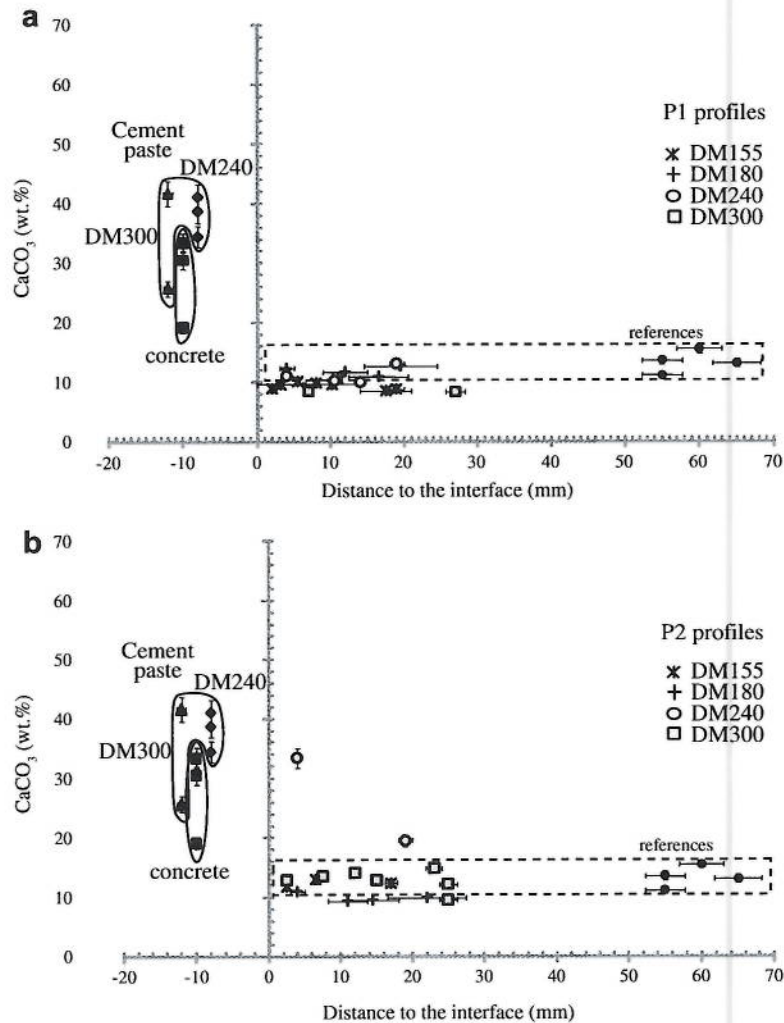


Fig. 5. Calcite content (CaCO₃ wt.%) of the CEM II cement pastes and concrete and of the argillite in the matrix along P1 profiles (a) and in the micro-cracks along P2 profiles (b).

4.2. The Toarcian argillite

4.2.1. Variations in the matrix (P1 scale)

No significant changes are shown by XRD in the mineral composition of the Toarcian argillite along the P1 profiles. The claystone is characterized next to or away from cementitious materials by the same mineral sequence with predominant phyllosilicates (illite, micas, kaolinite, and chlorite), carbonates (calcite, dolomite), K and Na–Ca feldspars and pyrite. The clay-fraction assemblage determined after air-drying and ethylene-glycol treatment is also similar in all samples, consisting of chlorite, non-expandable illite- or mica-type dioctahedral layers and kaolinite. A broad diffraction domain observed between 11 and 12 Å on the air-dried samples is assigned to an illite/smectite mixed-layer mineral. Alternatively, distinct mineralogical features are visible by SEM in the 10–15-mm thick dark zone when compared to the reference claystone. Next to the contact, large amounts of ettringite crystals are identified as more or less isolated acicular or ‘needle-like’ crystals (Fig. 2b). Throughout the entire zone, the structure of argillite appears denser and is less recognizable due to the occurrence of a Si–Ca rich gel, celestine, neofomed calcite and C–S–H phases (Fig. 2c). These latter phases are detected mainly by TEM based on their diffraction patterns and their foil-like morphology. These changes are close to those expected from the modeling (De Windt et al., 2008).

The chemical composition of the argillite whole-rock appears quite constant along the two longer profiles (DM240 and DM300) (Table 2). Only the 18–20-mm thick zone next to the interface with the cement paste shows slight changes: a clear and continuous increase in the CaO content progressing towards the cement paste, which is negatively correlated with a slight decrease in the K₂O content. The behavior of MgO is specific with a maximum value measured systematically at the external limit of the zone (18–20 mm) (Table 2; Fig. 8). This chemically modified zone is also enriched in Cs, whereas it is depleted in Sr and to a lesser extent in REEs (Table 2; Fig. 8).

Major changes are observed in the chemical composition of the argillite leachates (Table 3; Fig. 9). Reference samples located far from cementitious material are characterized by calcitic leachates with minor amounts of Si, Al, Mg and K, Sr contents of about 500 µg L⁻¹ and total elementary contents in the solution of about 200 mg L⁻¹. Relative to this reference, samples collected close to the concrete along the DM180 P1 profile show a decrease in the total elemental content of their leachates at the immediate contact with the cementitious material (0–10 mm; 180 mg L⁻¹). This decrease is mainly linked to a significant decrease in the Ca and Sr contents. A peak in Mg and Al amounts occurs roughly in the middle of the studied zone (Table 3; Fig. 9). The Sr isotopic results for the leachates of the reference samples are quite homogeneous with

Table 1

Location of the samples along the DM over-core corresponding to -155, -180, -240 and -300 cm layers with distances to the cement paste/concrete interface. Strontium chemical and isotopic compositions of the leachates; C and O isotope compositions and CaCO₃ wt.% of the extracted carbonate phases.

DM sample	Profile type	Sample ref.	Location (mm from the cement/concrete interface)	⁸⁷ Sr/ ⁸⁶ Sr	2σ	δ ¹⁸ O	δ ¹³ C	CaCO ₃ wt.%	
DM155	Argillite along a P1 profile	S1	0-4	0.708094	9E-06	25.7	-1.93	8.96	
		S2	4-7	0.708002	7E-06	25.2	-1.93	10.2	
		S3	7-9	0.708013	5E-06	25.4	-1.95	9.82	
		S4	9-13	0.707892	6E-06	25.7	-1.44	10.6	
		S5	13-25	0.707858	4E-06	26.1	-1.21	8.76	
	Argillite along a P1 profile	A2	0-6.5	0.707903	6E-06	24.6	-2.08	9.62	
		B2	6.5-14	0.707883	8E-06	25.6	-1.58	9.66	
		C2	14-21	0.707820	7E-06	26.4	-1.37	8.48	
	Material scraped along a P2 profile	A (blue z.)	0-5	0.708379	4E-06	24.5	-2.80	11.8	
		B	5-6.5	0.708142	5E-06	25.0	-2.39	13.0	
		D	8-14	0.708023	4E-06	nd	nd	nd	
		E	14-20	0.708001	4E-06	25.2	-1.58	12.4	
		Concrete	S6	-10	0.70838	7E-06	20.1	-9.67	19.2
			S7	-14	0.708364	4E-06	18.7	-11.7	33.5
	DM180	Argillite along a P1 profile	T1 (blue z.)	0-8	0.707949	5E-06	25.0	2.01	12.4
			T2	8-16	0.707917	6E-06	25.1	-1.27	11.7
			T3	16-17	0.707897	5E-06	24.9	-1.14	10.9
T4			17-22	0.707870	5E-06	25.2	-1.15	12.7	
Material scraped along a P2 profile		E3 (blue z.)	0-8	0.708050	6E-06	24.1	-1.96	11.0	
		E4	8-14	0.707993	6E-06	25.4	-1.40	9.38	
		E6	14-15	0.708019	8E-06	26.1	-0.85	9.52	
		E5	15-28	0.707916	7E-06	25.3	-0.95	9.95	
		Concrete	C_DM180	-10	0.708352	6E-06	19.2	-10.6	30.6
DM240		Argillite along a P1 profile	1/C2	0-8	0.707918	4E-06	24.4	-2.56	11.1
			1/D2	8-13	0.707960	4E-06	24.9	-2.49	10.3
	1/E2		13-15	0.707918	5E-06	25.0	-2.27	10.1	
	1/F2		15-23	0.707862	5E-06	25.6	-1.67	13.1	
	ARG REF2		55	0.707789	3E-06	25.6	-1.40	13.6	
	ARG REF3		65	nd	nd	25.8	-1.25	13.1	
	Material scraped along a P2 profile		1/C (blue z.)	0-8	0.708350	5E-06	12.9	-14.2	33.4
		1/D	8-13	nd	nd	nd	nd	nd	
		1/E	13-15	nd	nd	nd	nd	nd	
		1/F	15-23	0.708018	5E-06	24.6	-3.68	19.6	
	Cement	Pa1	-2	0.708166	6E-06	29.4	-12.5	38.8	
		Pa2	-6	0.708122	6E-06	20.6	-11.2	34.5	
		C_DM240	-10	0.708308	6E-06	19.0	-13.2	41.2	
	DM300	Argillite along a P1 profile	300/K	0-14	0.707919	6E-06	25.1	-1.97	8.5
300/j			13-41	0.707832	4E-06	26.6	-0.92	8.3	
300/N			60	0.707791	4E-06	25.0	-0.65	15.5	
300/O			55	0.707792	5E-06	26.5	-0.64	11.1	
Material scraped along a P2 profile			300/G	0-5	0.707885	4E-06	25.0	-1.97	13.0
		300/F	5-10	0.707942	3E-06	24.8	-1.77	13.6	
		300/E	10-13	0.707939	4E-06	24.9	-1.28	14.2	
		300/C	13-17	0.707939	5E-06	25.4	-0.98	12.9	
		300/A	15-20	0.707877	5E-06	nd	nd	nd	
		300/B	14-30	0.707921	3E-06	25.5	-0.79	14.9	
		300/I	20-30	0.707825	7E-06	26.3	-0.76	9.61	
		300/H	30-41	0.707799	4E-06	26.1	-0.74	12.2	
Cement		300/BC	-10	0.708267	4E-06	16.5	-10.7	41.7	
		300/BS	-10	0.708252	6E-06	18.9	-8.33	25.8	

a mean ⁸⁷Sr/⁸⁶Sr ratio of 0.707791 ± 4 × 10⁻⁶ (Fig. 7). This value agrees with previous measurements obtained on carbonates extracted from Toarcian argillite (Michel, 1999). Progressing towards the cement paste/concrete, the ⁸⁷Sr/⁸⁶Sr ratio of the leachates increases to a maximum value of 0.708094 (±9 × 10⁻⁶). This increase defines a 20-25-mm thick transition zone between the interface of the cement paste and the concrete and the reference claystone (Fig. 7a).

The carbonate content of the reference argillite samples is between 11 and 16 wt.% with a mean value of 13 wt.% (±2.2) (Fig. 5b). The δ¹³C and δ¹⁸O values for these carbonates are typical for marine materials (between -0.64‰ and -1.40‰ and

between 25.0‰ and 26.5‰ respectively) at more than 20 mm from the cement paste/concrete interface (Fig. 10). When compared to these reference compositions, a slight depletion in the CaCO₃ content is visible in the argillite within the 20 mm next to the cement paste/concrete interface, the values ranging around 10 wt.%. These carbonates are also characterized by a progressive decrease of the δ¹³C towards the cement paste/concrete interface (from -1.40‰ to -2.01‰). The P1 profile from sample DM240 shows the lowest δ¹³C value (-2.6‰) at the immediate contact with the cement paste (Fig. 6a). The δ¹⁸O remains constant along the profiles, ranging between 25‰ and 26.5‰.

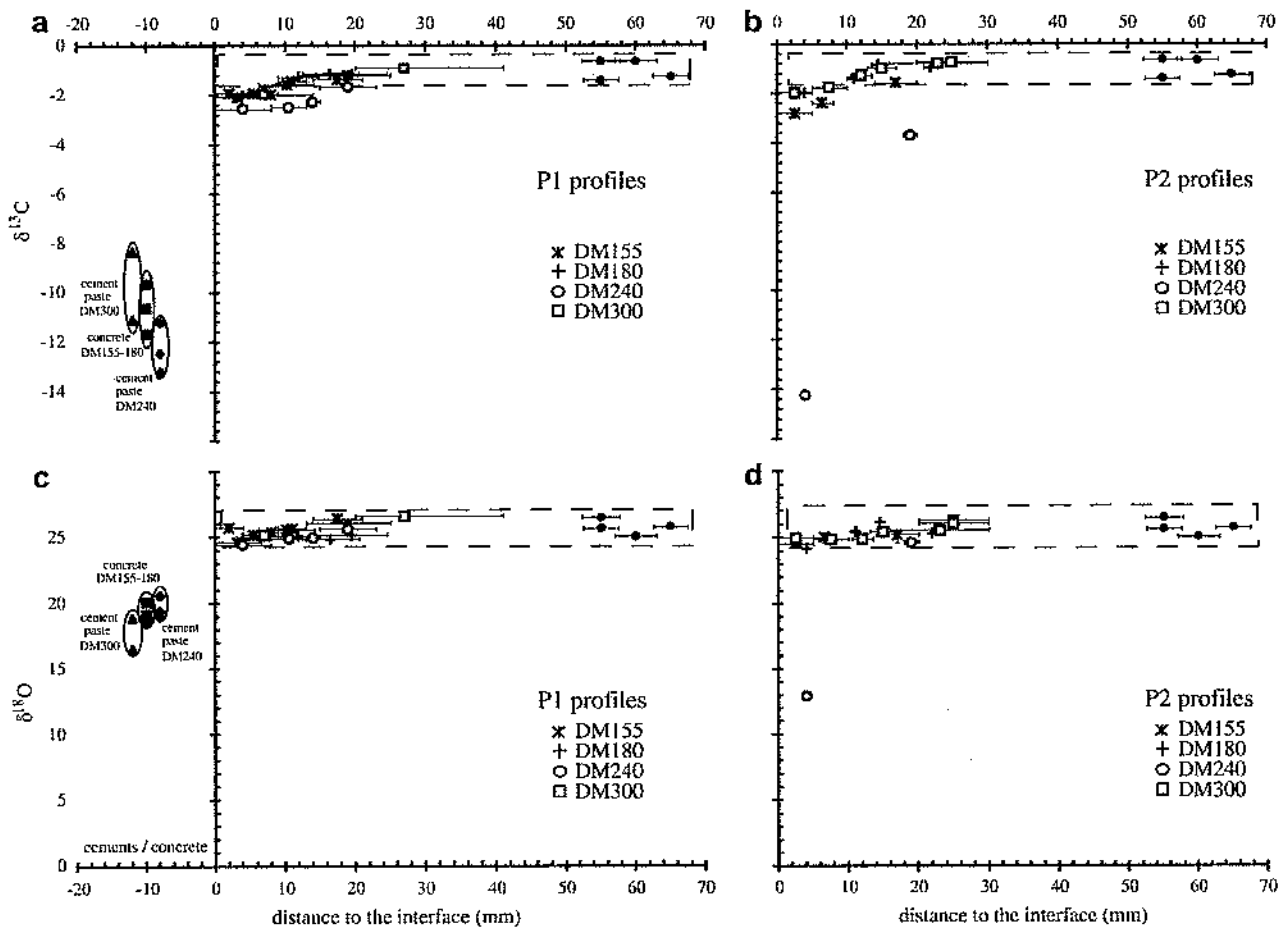


Fig. 6. Evolution of the C and O isotopic data ($\delta^{13}\text{C}/\delta^{18}\text{O}$) in the argillite matrix (P1 profiles; a and c) and in the argillite micro-cracks (P2 profiles; b and d). The data are compared with values measured in the cement paste and in the concrete.

4.2.2. Variation at the micro-cracks (P2 scale)

Along the P2 surfaces, the 10-mm thick zone identified by a zonation and a blue color yields the same mineralogical assemblage as that defined in the reference claystone by XRD. Only gypsum is observed as a specific phase. However, significant modifications become apparent when the samples are analyzed by SEM and TEM. These changes define five zones in the claystone. The first zone about 10- μm thick is at the immediate contact with the cement paste or concrete. This zone is enriched in acicular ettringite crystals (Fig. 3c) that appear to decrease in number away from the contact. A second zone (500- μm thick) is characterized by the occurrence of C-S-H crystals, which can be identified by a foil-like morphology by TEM and absence of diffraction (Fig. 3d and e). The blue color of the argillite appears to be related with the amount of C-S-H; sub-automorphic calcite and celestine are also observed in this zone (Fig. 3f). A third zone (500- μm thick) is identified by a higher amount of calcite, and minor amounts of celestine and C-S-H. The boundary between the second and third zones is marked by a smooth Ca-Si rich white band (Fig. 3g). A fourth zone is defined over 500 μm by significant calcite neoformation and a minor amount of C-S-H grains (Fig. 3h); the last 300 μm of this zone contains calcite crystals oriented towards the fifth zone. In this fifth zone, the texture of the Toarcian argillite is still recognizable and is slightly modified by some disoriented newly-formed calcite and some angular K-feldspar suggesting an overgrowth.

Changes are also observed in the chemical composition of the leachates of powders scraped from argillite P2 plans. Two distinguishable trends are visible according to the location along the DM over-core. In the upper layer (DM155), the total elemental content of the leachates decreases at the immediate contact with the concrete (0–10 mm; 153 mg L^{-1}). This decrease is mainly represented by a drop in Ca and Sr contents. Peaks in the amount of Mg and Al are also observed near the middle of the profile; these values being close to those observed along the P1 profiles. In the deeper layer (DM300 sample) close to the cement paste (<10 mm), the total elemental content of the leachates is higher than in the reference samples and increases towards the interface (finally reaching 289 mg L^{-1} relative to 200 mg L^{-1} in the reference argillite); this increase relates to an increase in nearly all the elements (Table 3; Fig. 9). Along each studied P2 profile, the $^{87}\text{Sr}/^{86}\text{Sr}$ isotopic ratio of the leachates is higher than in the marine-like distal Toarcian samples (Fig. 7b), even at 25 mm from the cement paste/concrete interface. At this interface, the $^{87}\text{Sr}/^{86}\text{Sr}$ ratio increases from $0.707799 (\pm 5 \times 10^{-6})$ to $0.708379 (\pm 4 \times 10^{-6})$.

No significant variation is visible in the carbonates when comparing the values of the scraped powder with those of the reference matrix (Fig. 5b). However, low $\delta^{13}\text{C}$ values (-1.8‰ to -2.8‰) are measured in the carbonates mainly located within 11 mm of the cement/concrete interface. One exception is in the DM240 profile, where a clear increase in the CaCO_3 content is found in micro-cracks (20–33 wt.%). These carbonates are also

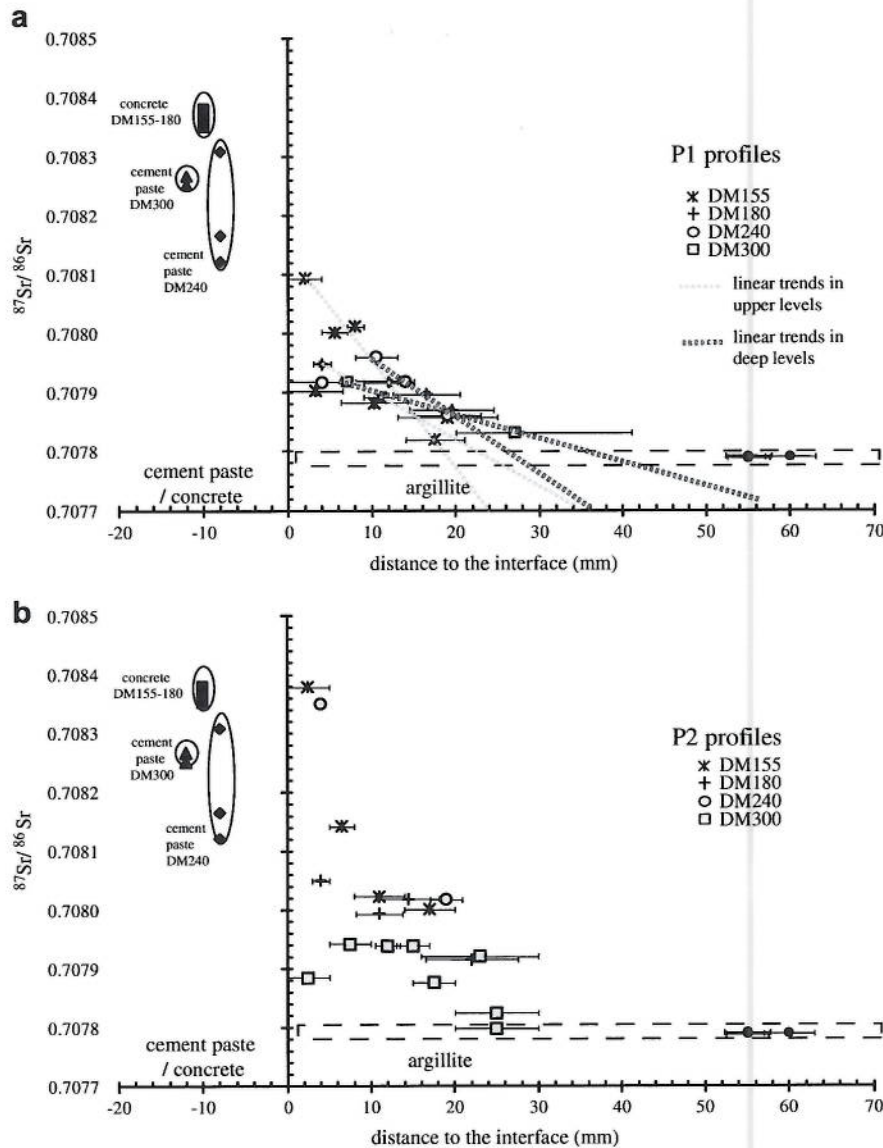


Fig. 7. Evolution of the Sr isotopic data ($^{87}\text{Sr}/^{86}\text{Sr}$) in the argillite matrix (P1 profiles, a) and in the argillite micro-cracks (P2 profiles, b). The data are compared with values measured in the cement paste and in the concrete.

strongly ^{13}C depleted with $\delta^{13}\text{C}$ values as low as -14.2‰ (Fig. 6b). The $\delta^{18}\text{O}$ composition is also different in these newly-formed minerals, yielding lower values (12.9‰) than those in the other samples ($25\text{--}26.5\text{‰}$) (Fig. 6d).

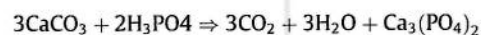
5. Discussion

5.1. On the interest of decoupling the data of carbonates and other soluble phases

In addition to the whole-rock analyses that are commonly performed in geochemical studies, the chemical and isotopic characteristics of selective extracted phases, carbonates and other soluble phases were examined. In fact, these phases are expected to be very sensitive to water/rock interaction, and even to record slight changes due to mineral dissolution and/or precipitation. They can, therefore, potentially record subtle and tiny disturbances

(Techer et al., 2009). The isotopic data (C, O, Sr) obtained on the cement paste/concrete and on the argillite outline distinct behaviors along the profiles. Before the processes responsible for these changes will be discussed, it is probably appropriate to detail the specifics of the two types of measurements. In fact, the C/O and Sr isotopic data were obtained on two different fractions.

The $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ values were measured on the carbonate fraction of the rock powders that were extracted by reaction with orthophosphoric acid according to the equation:



In argillite samples, this reaction results in the release of CO_2 extracted mainly from calcite (CaCO_3) and dolomite ($(\text{Ca}, \text{Mg})\text{CO}_3$). Numerous and variable carbonate minerals are typically encountered in the specific environments of cement paste and concrete, predominantly represented by calcite (Atkins et al., 1991; Vieillard and Rassineux, 1992) and minor phases such as spurrite ($\text{Ca}_5\text{Si}_2\text{O}_8\text{CO}_3$) and hydrocalcite ($\text{Mg}_{0.75}\text{Al}_{0.125}(\text{OH})_2(\text{CO}_3)_{0.125}(\text{H}_2\text{O})_{0.5}$)

Table 2
Argillite whole rock chemical composition: Panel (a) Major elements (in wt.%), Panel (b) Trace elements (mg kg⁻¹), Panel (c) In rare earth elements (mg kg⁻¹).

DM depth level	Sample ref.	Location (mm from the cement/concrete interface)	SiO ₂	Al ₂ O ₃	Fe ₂ O ₃	MnO	MgO	CaO	Na ₂ O	K ₂ O	TiO ₂	P ₂ O ₅	PF																
Panel (a) DM240	1P1-1	7.00	47.5	15.8	5.11	0.05	1.90	9.54	0.22	2.90	0.82	0.20	15.4																
	1P1-2	15.0	47.1	15.3	5.03	0.05	3.01	8.80	0.22	2.84	0.81	0.22	15.7																
	1P1-3	17.0	48.0	15.8	5.27	0.05	1.73	8.83	0.22	2.87	0.82	0.21	15.5																
	1P1-4	21.5	48.4	16.2	5.16	0.05	1.69	8.34	0.23	2.92	0.84	0.20	14.8																
	1P1-5	28.5	48.8	16.3	5.30	0.05	1.73	8.18	0.22	2.95	0.84	0.20	14.8																
	1P1-6	36.5	48.3	16.0	5.30	0.05	1.71	8.57	0.23	2.92	0.83	0.20	14.9																
	1P1-7	46.0	48.7	15.9	5.17	0.05	1.73	8.44	0.23	2.89	0.83	0.21	15.0																
	1P1-8	56.0	48.9	15.8	5.07	0.05	1.72	8.72	0.23	2.85	0.82	0.21	14.7																
DM300	2P1-1	1.25	47.2	16.1	5.07	0.05	1.79	9.25	0.21	2.84	0.84	0.22	15.9																
	2P1-2	3.5	47.7	16.1	5.00	0.05	1.81	9.25	0.21	2.87	0.83	0.22	15.2																
	2P1-3	9.0	48.1	16.1	5.10	0.05	1.98	8.98	0.22	2.89	0.84	0.24	14.7																
	2P1-4	16.0	47.9	16.3	5.19	0.05	2.83	8.27	0.21	2.88	0.84	0.24	14.6																
	2P1-5	20.0	48.3	16.5	5.18	0.05	1.68	7.99	0.22	3.12	0.83	0.22	14.9																
	2P1-6	24.5	48.5	16.7	5.16	0.05	1.69	7.57	0.22	3.15	0.83	0.23	14.7																
	2P1-7	31.5	48.7	16.6	5.27	0.05	1.74	7.88	0.23	3.19	0.85	0.24	14.5																
	2P1-8	40.0	48.8	16.7	5.28	0.05	1.76	7.72	0.22	3.06	0.85	0.24	14.6																
2P1-9	49.5	48.7	16.7	5.40	0.05	1.76	7.56	0.23	3.21	0.86	0.25	14.6																	
	58.5	48.9	16.9	5.31	0.05	1.78	7.44	0.23	3.25	0.86	0.24	14.3																	
	64.5	49.1	16.8	5.35	0.05	1.76	7.52	0.23	3.23	0.86	0.24	14.3																	
Panel (b) DM240	1P1-1	7.00	326	2.06	0.22	14.3	108	9.08	15.4	20.0	1.68	4.79	<L.D.	0.52	13.9	46.3	19.1	11.4	0.40	2.47	215	1.17	10.8	2.40	101	1.40	22.6	63.6	180
	1P1-2	15.0	328	1.90	0.23	14.2	108	7.60	14.2	19.3	1.73	4.64	<L.D.	0.65	13.7	49.2	17.7	108	0.40	2.41	239	1.15	10.5	2.35	97.1	1.32	23.1	55.1	173
	1P1-3	17.0	320	1.90	0.26	15.1	111	7.83	15.6	20.6	1.81	5.24	<L.D.	0.79	14.4	52.4	19.5	113	0.44	2.55	228	1.22	11.4	2.58	99.3	1.50	24.0	65.0	194
	1P1-4	21.5	325	1.90	0.25	14.2	112	8.15	15.4	21.6	1.82	5.27	<L.D.	0.61	14.9	49.3	19.9	118	0.42	2.70	230	1.26	11.7	2.67	104	1.47	24.9	59.9	193
	1P1-5	28.5	329	1.78	0.27	13.8	109	8.36	15.0	21.4	1.95	5.35	<L.D.	0.72	14.9	47.9	20.5	118	0.45	2.78	230	1.33	12.2	2.77	99.7	1.57	24.7	59.5	191
	1P1-6	36.5	328	1.86	0.27	14.3	104	8.21	15.0	21.1	1.82	5.17	<L.D.	0.58	14.6	46.1	20.7	116	0.48	2.74	232	1.31	12.3	2.74	97.7	1.58	24.7	61.9	186
	1P1-7	46.0	304	2.18	0.25	15.1	110	7.98	16.0	20.4	1.81	4.91	<L.D.	0.58	14.0	48.7	18.7	110	0.38	2.56	217	1.19	11.0	2.48	103	1.42	23.6	58.3	182
	1P1-8	56.0	309	2.16	0.24	14.5	109	7.88	15.3	20.1	1.79	5.05	<L.D.	0.55	14.4	45.2	17.7	110	0.42	2.49	220	1.18	10.9	2.49	104	1.49	24.5	59.9	190
	2P1-1	1.25	319	2.11	0.24	15.0	130	13.8	18.1	20.9	1.88	4.87	<L.D.	0.77	14.6	54.1	26.6	117	0.38	2.50	198	1.24	11.2	2.51	105	1.57	24.8	61.6	177
	2P1-2	3.5	315	2.11	0.24	13.8	112	9.78	16.6	20.4	1.86	4.75	<L.D.	0.82	14.0	49.1	18.7	114	0.37	2.62	199	1.20	11.0	2.41	101	1.44	23.4	56.4	174
	2P1-3	9.0	315	2.11	0.25	14.5	107	8.18	15.7	20.6	1.78	5.00	<L.D.	0.61	14.4	47.9	19.5	118	0.40	2.60	228	1.22	11.6	2.53	102	1.47	24.4	60.1	182
2P1-4	16.0	315	2.07	0.24	14.8	110	7.84	15.6	20.2	1.81	4.80	<L.D.	0.61	14.2	50.2	19.1	114	0.36	2.60	229	1.20	11.1	2.43	101	1.42	23.8	55.3	173	
2P1-5	20.0	312	2.36	0.26	15.0	118	8.64	17.1	21.6	2.12	4.90	<L.D.	0.73	15.3	52.2	19.6	120	0.41	2.89	237	1.30	12.1	2.63	109	1.54	25.0	53.1	187	
2P1-6	24.5	321	2.36	0.27	15.1	121	8.90	18.3	22.1	2.16	5.25	0.11	0.77	15.6	52.0	20.5	122	0.44	3.00	226	1.35	12.5	2.75	112	1.57	25.9	361.0	197	
2P1-7	31.5	316	2.37	0.26	15.1	116	8.68	17.3	22.0	2.09	5.10	<L.D.	0.98	15.2	50.3	20.3	119	0.41	2.96	228	1.31	12.6	2.76	109	1.54	25.9	59.0	189	
2P1-8	40.0	304	2.29	0.26	14.7	113	8.37	16.7	21.2	1.93	5.02	<L.D.	0.87	14.8	49.4	20.3	114	0.43	2.78	216	1.28	12.0	2.62	106	1.48	25.1	54.0	186	
2P1-9	49.5	603	2.31	0.26	14.8	151	8.31	16.6	21.2	2.00	5.07	<L.D.	0.57	14.8	49.5	19.7	114	0.40	2.78	222	1.29	12.2	2.69	107	1.50	25.6	50.3	188	
2P1-10	58.5	316	2.36	0.26	15.4	119	8.88	17.3	22.2	2.11	5.13	<L.D.	0.65	15.8	51.4	20.2	121	0.47	2.89	223	1.31	12.6	2.80	112	1.55	26.8	58.1	191	
2P1-11	64.5	324	2.37	0.27	15.4	127	8.69	17.2	22.1	2.06	5.15	<L.D.	0.90	15.5	56.1	20.4	119	0.44	2.82	226	1.33	12.5	2.72	110	1.59	25.8	53.8	188	

(continued on next page)

Table 2 (continued)

DM depth level	Sample ref.	Location (mm from the cement/concrete interface)	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm	Yb	Lu
Panel (C) DM240	1P1-1	7.00	35.9	72.6	8.33	30.7	5.73	1.22	4.45	0.69	4.06	0.79	2.22	0.34	2.33	0.35
	1P1-2	15.0	35.0	71.9	8.37	31.5	6.00	1.24	4.61	0.72	4.13	0.79	2.26	0.34	2.30	0.35
	1P1-3	17.0	36.7	74.8	8.72	32.2	6.09	1.25	4.59	0.74	4.34	0.84	2.38	0.36	2.46	0.38
	1P1-4	21.5	37.7	77.4	8.97	33.1	6.25	1.28	4.72	0.75	4.39	0.87	2.44	0.38	2.54	0.39
	1P1-5	28.5	38.1	77.4	8.94	33.0	6.09	1.26	4.52	0.75	4.33	0.86	2.47	0.37	2.62	0.40
	1P1-6	36.5	38.2	78.4	9.06	33.4	6.17	1.29	4.71	0.75	4.42	0.86	2.47	0.37	2.57	0.40
	1P1-7	46.0	36.8	75.6	8.63	31.8	6.04	1.27	4.88	0.73	4.29	0.82	2.36	0.36	2.43	0.38
	1P1-8	56.0	36.4	75.5	8.77	32.2	6.05	1.29	4.90	0.76	4.44	0.86	2.45	0.37	2.56	0.39
DM300	2P1-1	1.25	37.0	76.1	8.85	33.0	6.24	1.30	5.03	0.76	4.42	0.86	2.44	0.37	2.48	0.39
	2P1-2	3.5	36.4	74.9	8.64	32.3	6.11	1.28	4.89	0.76	4.41	0.83	2.34	0.36	2.40	0.37
	2P1-3	9.0	37.6	77.4	9.07	34.0	6.47	1.37	5.14	0.80	4.63	0.87	2.42	0.36	2.50	0.38
	2P1-4	16.0	36.6	74.8	8.77	32.4	6.23	1.31	5.00	0.76	4.43	0.84	2.36	0.37	2.43	0.38
	2P1-5	20.0	38.4	79.3	9.12	33.4	6.44	1.36	5.09	0.79	4.60	0.88	2.44	0.37	2.52	0.39
	2P1-6	24.5	40.1	82.5	9.54	35.0	6.72	1.39	5.24	0.83	4.81	0.91	2.58	0.40	2.64	0.43
	2P1-7	31.5	39.3	81.7	9.49	34.9	6.76	1.43	5.29	0.83	4.80	0.92	2.61	0.39	2.65	0.42
	2P1-8	40.0	37.9	78.4	9.13	33.2	6.45	1.38	5.18	0.80	4.68	0.89	2.53	0.38	2.58	0.41
	2P1-9	49.5	38.1	79.6	9.29	34.4	6.71	1.41	5.30	0.82	4.83	0.90	2.54	0.38	2.60	0.41
	2P1-10	58.5	39.8	82.2	9.56	35.7	6.77	1.42	5.47	0.84	4.93	0.93	2.68	0.41	2.70	0.41
	2P1-11	64.5	39.2	81.1	9.49	35.1	6.67	1.42	5.42	0.82	4.69	0.91	2.62	0.40	2.63	0.41

(Taylor, 1997). The alteration of these cementitious materials also leads to the formation of secondary carbonate phases such as new calcite and thaumasite ($\text{Ca}_6[\text{Si}(\text{OH})_3]_2(\text{CO}_3)_2(\text{SO}_4)_2 \cdot \text{H}_2\text{O}$) that can be generated by extensive replacement of the C-S-H during coupled dissolution of sulfate and carbonate (Iden and Hagelia, 2003) at low temperature ($<18^\circ\text{C}$) for thaumasite. In the case of the DM study, the main components of the cement paste and concrete are portlandite ($\text{Ca}(\text{OH})_2$), calcite (CaCO_3), ettringite ($3\text{CaO} \cdot \text{Al}_2\text{O}_3 \cdot 3\text{CaSO}_4 \cdot 32\text{H}_2\text{O}$) and amorphous C-S-H. C-S-H and calcite were also observed as secondary minerals in the argillite close to the cement paste/concrete interface. Thus, the reported C/O isotope signatures obtained by extraction of CO_2 from the various samples essentially represent calcite.

The $^{87}\text{Sr}/^{86}\text{Sr}$ ratios were measured on whole-rock powders that were leached with a 1 N acetic acid solution. Such a procedure is expected to dissolve calcite, but also salts, sulfates and oxides, together with removal of pore water in most of the geological media (Clauer and Chaudhuri, 1995; Techer et al., 2009). In the case of the cement paste and concrete environments, other highly soluble phases can be extracted during this leaching procedure. Among the typical cementitious phases formed by hydration of cement paste and/or by alteration, C-S-H and portlandite – together with calcite – are known to be highly soluble and unstable under acid conditions (Hewlett, 1998). Decalcification of C-S-H and dissolution of portlandite are commonly observed when in contact with deionized water in ambient conditions (Planel et al., 2006). The two phases are known to be unstable in low pH environments (Berner, 1992; Vieillard and Rassinieux, 1992; Adenot and Buil, 1992; Faucon et al., 1998; Le Bescop and Solet, 2006) and in clay-rich environments (Dauzères et al., 2009, 2010a,b). Thus, it can be expected that leachates may not only carry calcite but also C-S-H and portlandite. The complementary analysis of carbonates and leachates therefore represents an interesting issue to follow the occurrence of secondary cementitious phases and, as a consequence, the imprints of high-pH fluids.

5.2. Alteration of the Toarcian argillite close to the cement paste/concrete

Thin rims of Toarcian argillite samples were cut parallel to the cement paste/concrete interface or scraped off discontinuity surfaces. For each analyzed zone, aliquots between 100 and 500 mg were separated. Studying such small amounts of representative material increases the possibilities of technical and analytical artifacts. In order to overcome this issue, experiments were performed in duplicate, two sampling profiles being studied of the same borehole layer whenever possible.

In the deep layers of the DM borehole, at more than 5 cm from the interface with the cement, argillite measurements reveal CaCO_3 contents (11–15.5 wt.%), $\delta^{13}\text{C}$ (-0.64‰ , -1.40‰), $\delta^{18}\text{O}$ (25‰, 26.5‰) and $^{87}\text{Sr}/^{86}\text{Sr}$ (0.70779) data that are typical of a marine undisturbed Toarcian claystone (Michel, 1999). These characteristics do not remain constant along the matrix profiles closer towards the cement paste or the concrete, and are different from those measured in the material scraped from micro-cracks close to the cementitious materials. Several changes are observed within ca. 20 mm of the argillite matrix (P1) and within ca. 25 mm of the micro-cracks (P2); these zones are referred to as the 'geochemical disturbed zone' (GDZ). The thickness of the GDZ is larger than that observed by mineralogical methods, which is limited to ca. 10–15 mm along the P1 scale and to 10 mm along the P2 scale. The extent of mineralogical transformations in the matrix (P1) increases with borehole depth and is more pronounced in layers that are less influenced by the train tunnel EDZ. When discontinuity plans (P2) do not occur in the clay-rich matrix, transformations seem to imprint the matrix (P1) more significantly.

Table 3
Chemical composition of the argillite, cement paste and concrete leachates (in mg L⁻¹; 2σ is ± 10% for both).

DM depth level	Profile type	Sample ref.	Location (mm from the cement/concrete interface)	Si	Al	Mg	Ca	Fe	Na	K	P	Sr 10 ⁻³	Sum
DM155	Material scraped along a P2 profile	A (blue z.)	0–5	2.82	1.34	3.07	139	2.49	0.53	3.16	0.34	348	153
		B	5–6.5	2.59	2.67	4.34	173	2.77	0.83	5.51	0.39	503	193
		D	8–14	4.19	4.42	6.52	204	4.24	0.88	7.62	0.44	534	233
		E	14–20	1.80	1.90	3.08	208	3.03	0.79	5.29	0.34	488	225
	Concrete	S7	–10	60.8	9.42	3.01	551	8.94	0.12	0.45	–	377	635
DM180	Argillite along a P1 profile	T1 (blue z.)	0–8	1.67	1.21	3.43	165	2.56	0.60	4.40	0.31	498	180
		T2	8–16	0.71	0.75	5.49	196	2.67	0.65	4.31	0.45	562	212
		T3	16–17	1.14	0.85	4.15	189	3.19	0.61	3.95	0.25	539	204
		T4	17–22	0.93	0.91	3.53	193	3.11	0.60	4.07	0.23	538	208
	Concrete	C_DM180	–10	59.8	10.2	2.80	553	9.48	0.11	0.55	–	408	637
DM240	Cement	C_DM240	–10	42.9	9.16	1.96	698	8.18	–	–	0.15	231	761
DM300	Argillite along a P1 profile	300 N	60	0.87	0.80	3.51	182	3.26	0.61	2.66	0.20	499	195
		300O	55	0.89	0.88	3.93	195	3.50	0.72	2.89	0.17	554	210
	Material scraped along a P2 profile	300F	5–10	1.44	1.54	5.65	270	4.35	0.90	4.04	–	717	289
		300E	10–13	0.36	0.86	5.15	210	3.35	0.62	4.34	–	644	226
		300C	13–17	0.28	0.66	5.14	191	2.92	0.70	2.34	–	593	205
		300H	30–41	0.42	0.52	3.79	195	2.64	0.48	3.28	0.24	538	207
	Cement	C_DM300	–10	48.1	24.8	4.48	1259.9	22.8	0.31	0.71	0.36	339	1363

5.2.1. In the matrix (P1 scale)

The mineral transformation of the claystone within the 10–15-mm thick dark zone corresponds mainly to neoformation of calcite, ettringite and C–S–H phases. The proportion of these neoformed crystals is, however, very limited as they cannot be detected by XRD means. Less than 5 wt.% of each secondary mineral is expected in the disturbed matrix of the claystone. Despite this low variation intensity, the petrophysical characteristics of the rocks appears to have changed upon visual inspection. Gaboreau et al. (submitted for publication) investigated these changes by autoradiography and petrophysical measurements. They demonstrated a clogging of the claystone porosity in the dark zone that is likely linked to the mineralogical changes as defined here. On the basis of chemical and isotopic changes, a thicker disturbed zone is identified, reaching 20 mm in thickness. This GDZ is mainly characterized by higher CaO contents together with slightly lower carbonate proportions (10–12 wt.%), in comparison to the reference state. This indicates slight carbonate dissolution together with neoformation of a calcitic phase. Carbonate dissolution can be expected in the clay-rich matrix at the contact with the cement paste if considering the Ca contents in the two materials at an initial state. In the presence of KOH and NaOH, Ca²⁺ contents can be lower in the cement paste, than in the argillite (CHESS calculation) (Van der Lee, 1998). This difference can induce a Ca²⁺ flux from claystone to the cementitious material. If this dissolution of carbonates cannot be detected by means of the mineralogical techniques due to the low extent of the variation, the neoformation of the calcitic phase is supported by XRD, SEM and TEM, as well as by the occurrence of a Si–Ca rich gel. In the GDZ, the remaining carbonate phase does not yield marine isotopic signatures, but shows instead low δ¹³C values (–2‰ to –2.56‰) that are intermediate between those of marine carbonates (–0.64‰, –1.40‰) and those of cement paste and concrete (δ¹³C = –8‰ to –13‰). Fourcade et al. (2007) demonstrated that in alkaline conditions, such as those produced during cement paste and concrete alteration, very low δ¹³C carbonates (labeled hereafter as ‘cementitious carbonates’) can precipitate. A minimum δ¹³C-value of about –30‰ was measured by the authors in such secondary phases by *in situ* SIMS analyses. The relatively low δ¹³C-values measured in the GDZ of the Toarcian argillite next to the cement paste and concrete could illustrate a mixing between

a pristine (mean δ¹³C-value of –1‰) and a cementitious carbonate (mean δ¹³C-value of –30‰) crystallized in high-pH conditions. Calculations indicate that the lowest δ¹³C composition of –2.56‰ in the GDZ could result from dissolution of authigenic carbonates and crystallization of about 0.60 wt.% of cementitious carbonates. Such a minute carbonate re-precipitation is certainly insufficient for detection by classical mineralogical methods. This hypothesis of a combined ‘dissolution–precipitation’ process of carbonate species may result in the disappearance of authigenic dolomite and neoformation of new calcite and other Mg-rich phases. The changing MgO contents in the whole-rock, points to the redistribution of Mg in the GDZ with a clear Mg-enriched phase that precipitated at the external limit of the zone (Fig. 8). The mobility of this element was discussed by Fernandez et al. (2009b), a clear exchange between alkaline cations and Mg²⁺ from a compacted bentonite being observed during diffusion experiments. The newly-formed calcite carries a mean homogeneous δ¹⁸O composition of about 26‰. This could correspond to an isotopic re-equilibration between the mineral and local fracture-water at a surface temperature of 12 °C, which is the mean temperature in the train tunnel. In support of this, a calcite in equilibrium with present-day fracture water (δ¹⁸O_{fracture water} = –8‰ to –7‰) (Patriarche, 2001) should be enriched by about 33‰ relative to water (fractionation parameter of Zheng, 1994) and provide δ¹⁸O values near 25–26‰. Combined with the carbonate dissolution–precipitation process, neoformation of a calcitic phase is evidenced in the GDZ. Part of this calcitic phase is poorly soluble as indicated by the leachates that are mainly depleted in Ca and Sr (Fig. 9); it may correspond to ettringite. This phase is known to occur in alkaline environments as a secondary mineral of low solubility (Perkins and Palmer, 1999). The soluble fraction or leachate, even if Ca-depleted (Fig. 9), is significantly enriched in radiogenic ⁸⁷Sr suggesting derivation of the element with a higher ⁸⁷Sr/⁸⁶Sr ratio from cement paste and/or concrete. As for the C isotopes, the ⁸⁷Sr/⁸⁶Sr increase in the argillite matrix can result from a mixing of a pristine Toarcian with a newly-formed cementitious soluble phase. To test this hypothesis, a mixing model has been computed (Fig. 11), according to the equations given by Faure (1986). The Sr contents and ⁸⁷Sr/⁸⁶Sr ratios that characterized the argillite pristine phase are, respectively, 500–550 μg L⁻¹ and about 0.70779. The cementitious soluble

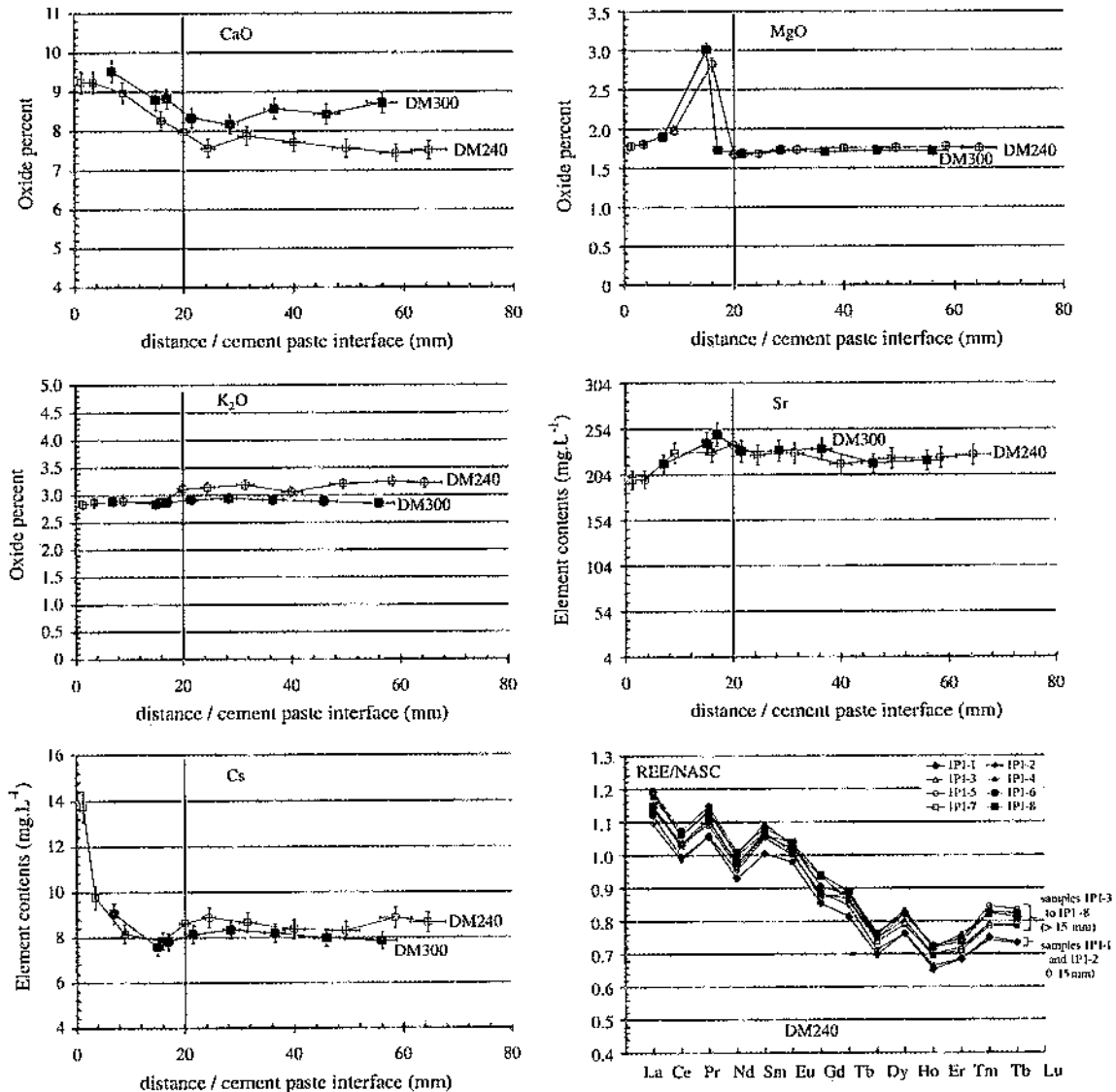


Fig. 8. Chemical composition for argillite whole-rock in major, trace and rare-earth elements according to the distance to the cement paste interface at the DM240 and DM300 layers.

phase is more radiogenic ($^{87}\text{Sr}/^{86}\text{Sr} = 0.70835$) with heterogeneous Sr contents (from 230 to 400 $\mu\text{g L}^{-1}$). Sample values from the DM180 P1 profile define an intermediate domain between concrete and argillite end-members, which is consistent with a proportion of 12–33 wt.% of cementitious phases that have precipitated in the GDZ of the Toarcian argillite (33 wt.% at the immediate contact and decreasing when moving away). The main part likely corresponds to C–S–H and/or portlandite phases and only 0.6% of these phases correspond to carbonates, as discussed above. Such high contents of secondary phases can explain the clogging of the claystone porosity (total porosity of the argillite decreasing from 9% in the reference samples to $1.9 \pm 0.3\%$ in the 15–20 mm thick zone close to the concrete) (Gaboreau et al., submitted for publication). This imprint of cementitious fluids in the Toarcian argillite matrix appears to be recorded in more than 20 mm of argillite from the interface (Fig. 7a), which is more than the dark zone where C–S–H were observed by mineralogical techniques (10–15 mm). At the DM155 and DM180 layers, no stable reference state is reached in the last sample of the P1 profile

(20–25 mm), suggesting an imprint of cementitious fluids beyond this limit distance from concrete (Fig. 7). Unfortunately, no additional samples were available to constrain this hypothesis. When considering the trends of the $^{87}\text{Sr}/^{86}\text{Sr}$ data as a function of distance to the cement paste/concrete in the GDZ, linear functions can be fitted with correlation factors that range from 0.86 to 0.96 (Fig. 7a). This relationship suggests a diffusion process for the propagation of the cementitious fluids into the argillite matrix. This propagation expands more greatly into the matrix of the deeper borehole layers than into that of the upper layers. When discontinuity surfaces are visible, for instance in the EDZ of the train tunnel, the fluid propagation appears to be less developed in the claystone matrix.

5.2.2. In the micro-cracks (P2 scale)

Along the few micro-cracks that developed perpendicularly to the cement paste/concrete interface (P2 profiles), no secondary minerals were unequivocally detected by the naked eye, even when suspected due to a bright crystalline aspect and a blue

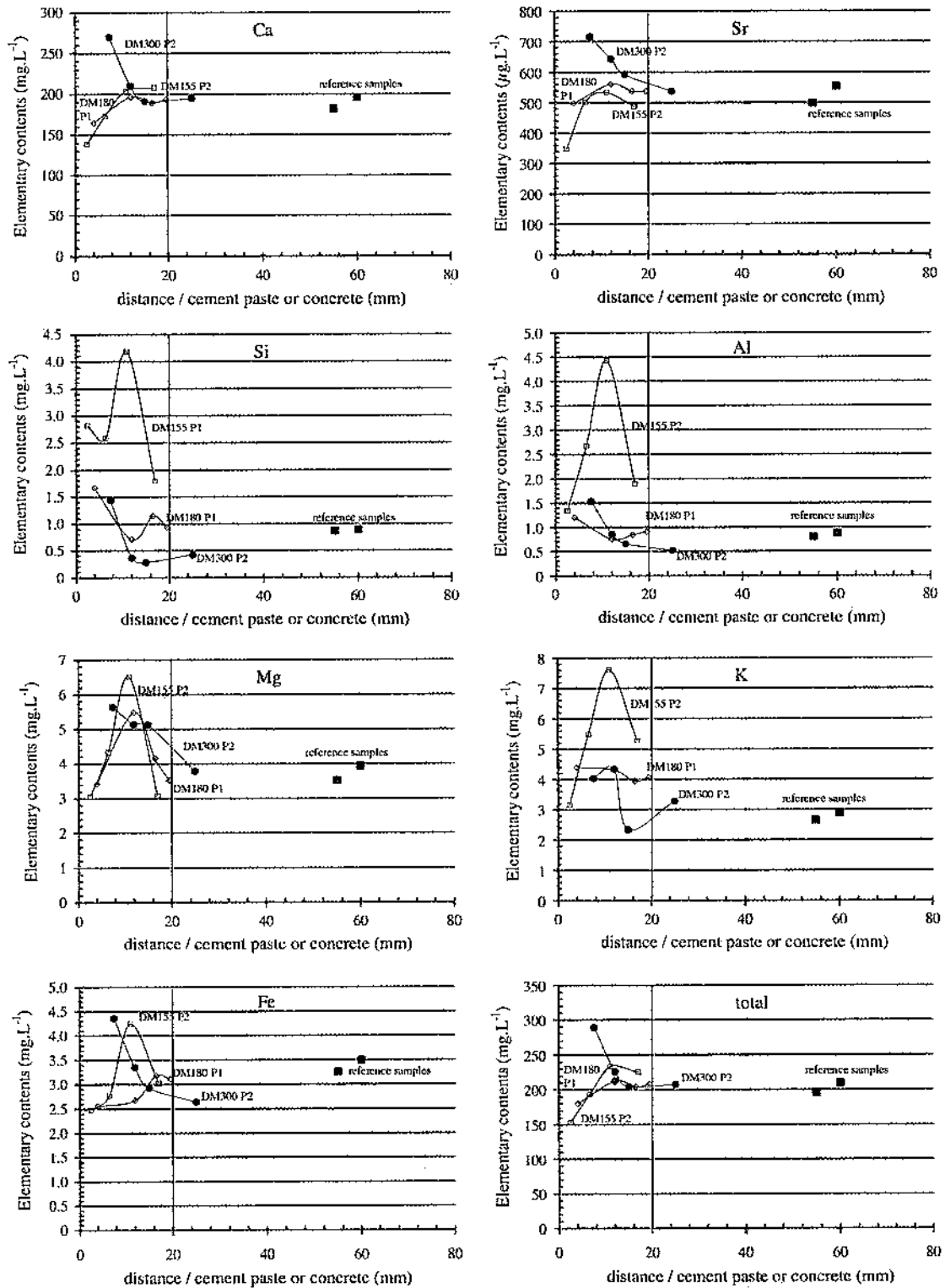


Fig. 9. Chemical composition for argillite leachates away from and next to the cementitious materials from DM155, DM180 and DM300 layers. Evolution along P1 (open symbols) and P2 profiles (black circles).

color. SEM and TEM observations clearly indicate mineralogical transformations, mainly with neoformation of ettringite, but also of calcite and C-S-H distributed in various rims over ca. 15 mm. Measurement of carbonate contents reveals significantly higher

values than those obtained in the argillite matrix (P1 scale) at a similar distance from the cement paste/concrete interface (Fig. 5b). The C isotopic signatures of the carbonates, even if negative (which suggests an alkaline origin) are not as low as

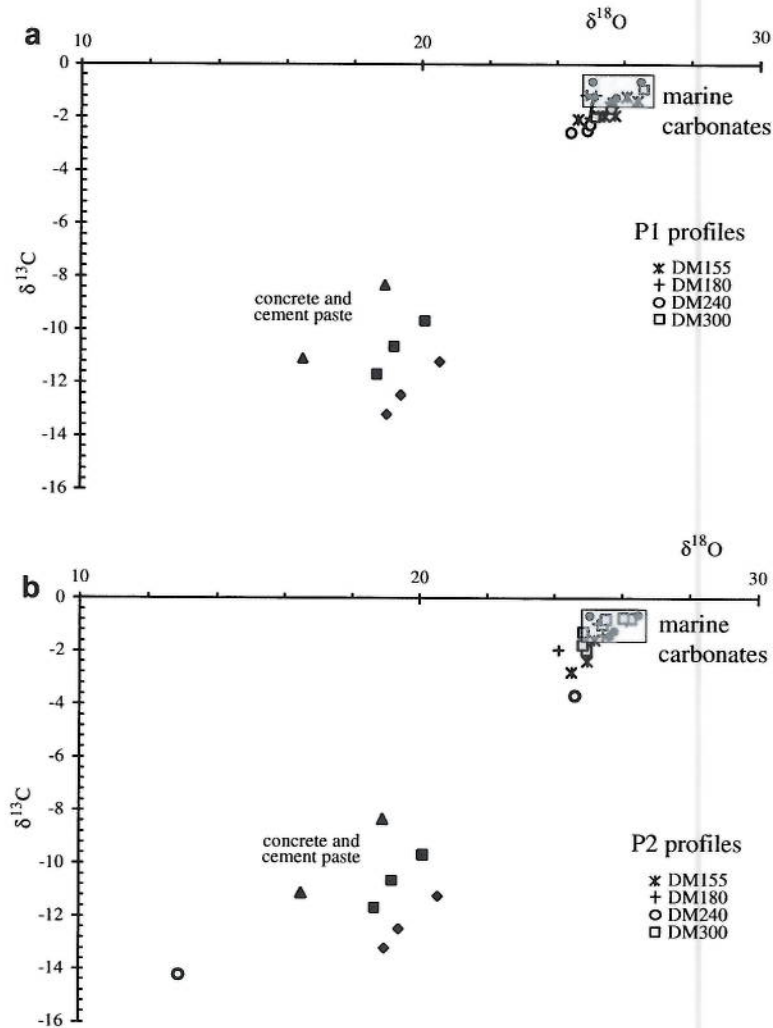


Fig. 10. $\delta^{13}\text{C}$ versus $\delta^{18}\text{O}$ diagram for the cementitious carbonates (cement paste and concrete), for the claystone carbonates and for carbonates extracted along the P1 and P2 profiles.

those expected for carbonates precipitating from alkaline fluids (Fourcade et al., 2007). This could result from a sampling artifact: the scraping could have removed some argillite matrix containing authigenic calcite. The P2 data thus reflect a mixing of secondary calcite precipitated on the P2 surfaces and marine calcite of the Toarcian matrix. Despite this potential artifact, it appears clearly that a P2 alteration superimposes that of the P1 matrix scale with lower $\delta^{13}\text{C}$ values reported in the micro-cracks (Fig. 6a and b). Similarly, the chemical and isotopic Sr data obtained on the soluble fractions extracted from P2 material outline a trend of disturbance close to that obtained for P1 samples, but significantly higher. For instance, when considering the Sr mixing line (Fig. 11), the results obtained on the DM155 profile are located on the same mixing line as that obtained for the results from the P1 profile calling for mixing fractions of 45–100 wt.% of newly formed minerals through percolation of altering concrete fluids. Once again, the main part of these secondary phases likely corresponds to C–S–H and/or portlandite phases. A precise identification of the percolation process along the discontinuities (diffusion or any other mechanism) is challenging due to a possible sampling artifact (i.e., mixing between P1 and P2).

The two P2 profiles studied in the deep layers (DM240, DM300) also show specific behaviors. The micro-crack sampled at the

DM240 layer shows a very well expressed carbonation (20–32 wt.% of carbonate minerals), with low $\delta^{13}\text{C}$ values (-3.68‰ to -14.2‰) that are even lower than those of the cement paste. DM240 is the sole sample showing such high secondary calcite precipitation in the studied micro-cracks. In addition, the O isotope composition ($\delta^{18}\text{O} = 12.6\text{‰}$) of these newly formed minerals is significantly abnormal when compared to that of the matrix or the other cementitious carbonates ($\delta^{18}\text{O}$ of about 26‰). The lower $\delta^{18}\text{O}$ value in this layer is close to what is predicted for an equilibrium at higher temperature or in alkaline condition for pH higher than 10 (Fourcade et al., 2007). The low $\delta^{13}\text{C}$ values are also compatible with precipitation from high-pH solutions. Therefore, it can be suggested that the changes observed in this layer of the DM-borehole result from interactions that occurred in the first stage of the borehole infilling (soon after injection of the cement paste) and not from 15-a interaction.

Specific chemical and Sr isotopic data for the micro-crack sampled at the DM300 layer show: low $^{87}\text{Sr}/^{86}\text{Sr}$ ratios, high Ca and Sr contents, and high total elemental contents (Fig. 9). As a result, a single two end-member mixing cannot explain the patterns of the powders scraped from these micro-cracks (Fig. 11). In the Sr mixing diagram, the data define a range of compositions and contents that do not fit a mixing line, neither when considering the

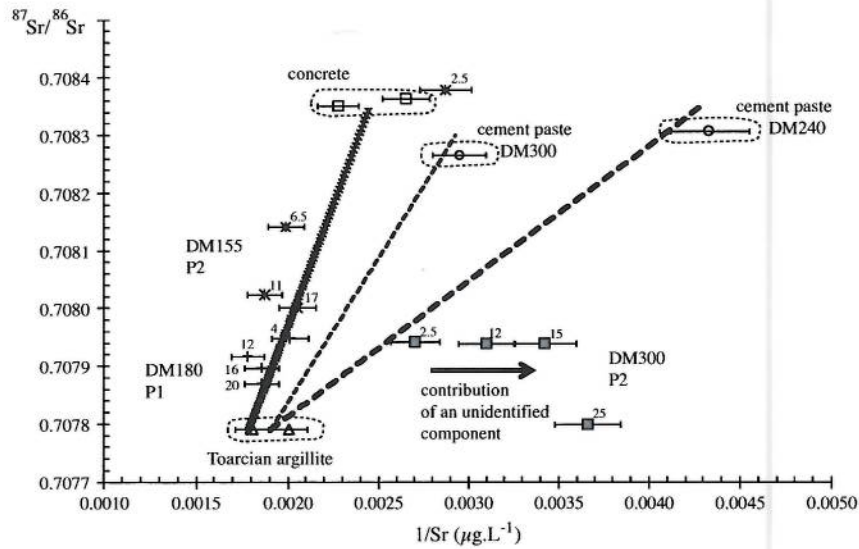


Fig. 11. Evolution of the Sr isotopic ratios ($^{87}\text{Sr}/^{86}\text{Sr}$) as a function of the Sr elementary contents ($1/\text{Sr}$ in $\mu\text{g.L}^{-1}$). Theoretical mixing between carbonates from Toarcian argillite and carbonates from cement paste or concrete are represented by the dotted lines connecting the two end-members (equation from Faure (1986)). Values measured in samples collected along three studied profiles (DM155 P2, DM180 P1, DM300 P2) are reported in this diagram with the reference of their sampling location (number above the symbols referring to the distance to the cement paste/concrete interface). The DM115 P2 and DM180 P1 data can be explained by a two-end member mixing as defined above. Values reported along the DM300 P2 profile cannot be interpreted as a consequence of cementitious material/argillite interaction (see text for explanations).

characteristics of the cement paste, nor when considering those of the concrete (Fig. 11). It can be predicted that the percolation of high-alkaline fluids is not the only process responsible for the crystallization of secondary minerals observed in the micro-cracks of the deep layer of the DM borehole. In the hypothesis of a mixing model, the data call for a third mixing component that could have had a greater imprint than the alkaline one and may be in equilibrium with the Tournemire present-day fracture water, if considering the $\delta^{18}\text{O}$ patterns. The imprint of this unidentified component is especially important, as the sample is located far from the interface with the cement paste, and, therefore, next to the over-coring pipe. Such data could call for a contamination process with a solution depleted in Sr at surface temperature, for instance with water pumped into the Tournemire canal during the over-drilling. This water corresponding to that of the Cernon fault could have stayed in the depth of the borehole and impacted the external samples.

5.3. The isotopes as tools for tracking imprints of high-pH fluids

The combination of C and Sr isotopes allows discrimination between carbonates and other soluble phases that are the main markers of high-alkaline fluid percolations in a clay-rich medium. This isotopic approach appears, therefore, very appropriate for tracking the occurrence of very discrete percolations of high-alkaline fluids, to evaluate their interactions with the natural host rocks, to discuss their spatial extents and to demonstrate their origin. The use of isotopic markers is extremely important in contexts where mineralogical changes are tiny and not sufficient enough to be detected directly by means of routine methods. For instance, $\delta^{13}\text{C}$ and $\delta^{18}\text{O}$ data allow distinction and identification of newly-formed cementitious carbonates. These carbonates can be calcite, hydroxalite or thaumasite in the context studied here. Discrimination between these various types of minerals is difficult but can be evaluated by considering the chemical compositions of the whole rocks. For instance, in the matrix scale (P1 scale) of the external zone of the GDZ, an increase in MgO and Al_2O_3 contents suggests

the precipitation of hydroxalite. The Sr isotope data also give complementary information on other soluble phases and in their proportions, as for C-S-H. As a result, combined C/O and Sr isotopic data clearly identify an alkaline fluid imprint. This imprint appears to be more extended when considering isotopic patterns than on the basis of the sole mineralogical results, which reflects the difference in sensitivity of the two tracers. Alternatively, isotopic methods are much more sensitive and extremely tenuous variations in mineral content can induce significant changes in the signal. On the basis of C isotope measurements, it is possible to detect for instance the neoformation of less than 1% of very low- $\delta^{13}\text{C}$ cementitious carbonate in samples of marine origin. Furthermore, using XRD or SEM to detect and identify cementitious phases such as C-S-H, is often challenging in low intensity processes, but this can be improved with combined isotopic approaches. Variations between the GDZ and the mineralogical altered zone can thus be viewed as a difference in the detection limits of the analytical methods used.

Although the sensitivity of the C and Sr isotopic methods is clearly well-suited for tracking most of the minerals formed during high-pH fluids interactions in the studied DM-borehole, it should be noted that ettringite and/or C-A-S-H can also form (Gaucher and Blanc, 2006; Fernandez et al., 2009a; Dauzères et al., 2010a). There are still uncertainties about the solubility of these minerals (Myneni et al., 1998; Perkins and Palmer, 1999), but they appear to be more stable in acetic acid leaching conditions, and are probably not extractable by the HAc protocol used in this study. A selective extraction method for the more stable mineral phases needs to be promoted to improve the evaluation of the imprint of the alkaline fluids into clay-rich media.

Last but not least, the isotopes offer the opportunity to define the origin of fluids. In the case of the DM borehole, the observed mineralogical transformations were initially attributed to high-pH fluid percolations. Importantly, changes defined along the discontinuities in the deepest layer of the core cannot be explained by a single cement/argillite interaction, and thus cannot be considered as an appropriate analogue for studies on interactions among clay/cement paste or concrete.

6. Conclusions

The interactions between natural clay and cementitious materials were assessed in a 15-a old engineered analogue for deep geological storage, using a combined mineralogical and chemical/isotopic study. This exceptional analogue was observed in the IRSN Tournemire experimental tunnel in Aveyron, France, along a 3 m long borehole (DM borehole). To assess the behavior of the Toarcian argillite that was in direct contact with a CEM II cement paste or a CEM II concrete for 15 a the variations of the argillite whole-rocks were tracked as well as of selective extracted mineral phases. The whole rock CO₂ extraction and O/C analysis were aimed at tracking the occurrence of cementitious carbonates, whereas leachates obtained by dilute acid reactions at a low acid/rock ratio consisted mostly of soluble mineral phases such as carbonates, C-S-H and portlandite.

The study also emphasizes how the isotopic variations mimic the mineralogical changes with a somewhat extended volume of alteration. The thickness of disturbance and the relative proportion of newly-formed minerals can be estimated in the argillite matrix and in the secondary micro-cracks; the origin of the fluids responsible for their crystallization was also identified. By this combined mineralogical and chemical/isotopic approach, the role of cementitious materials can be better evaluated and validated in engineered analogues. Furthermore, the presence of cementitious secondary minerals and their relative contents could be tracked in the 'geochemical disturbed zones' (GDZ) over greater distances by these isotope methods, than by the classical mineralogical techniques. This difference emphasizes the potential of isotopes as appropriate tools for the characterization of alteration processes among cement materials and natural clays after long-lasting interactions. Quantitative and predictive modeling may now be determining in the interpretation of the described disturbances.

The study also outlines the scale and the type of disturbances that may occur in claystones when interacting with cement paste or concrete, which is expected in a deep disposal storage. After 15 a of interaction, the observed changes are extremely limited in the clay-rich matrix as the alteration extends over a few centimeters, thus resulting from slow and discrete processes. The changes are of the same order of magnitude in micro-cracks induced for instance in an EDZ.

The data indicate that disturbances in a deep storage will mainly affect the near field and will not extend into the geological barrier. Alternatively, they could contribute to the insulation of the waste package as they are mainly characterized by the neoformation of amorphous phases (C-S-H) and to a lesser extent of calcite, intimately linked to the clay matrix and to a clogging of the rock porosity.

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References

- Adenot, F., Buil, M., 1992. Modelling of the corrosion of cement paste by de-ionized water. *Cem. Concr. Res.* 22, 489–496.
- Atkins, M., Glasser, F.P., Moroni, L.P., 1991. The long term properties of cement and concretes. *Mater. Res. Soc. Symp. Proc.* 212, 373–386.
- Atkinson, A., Harris, A.W., Hearne, J.A., 1985. Hydrothermal Alteration and Ageing of Synthetic Calcium Silicate Hydrate Gels. Safety Series Reports NSS/R374. UK Nirex Limited, Harwell, UK.
- Berner, U.R., 1992. Evolution of pore water chemistry during degradation of cement in a radioactive waste repository environment. *Waste Manage.* 12, 201–219.
- Boisson, J.Y., Bertrand, L., Heitz, J.F., Moreau-Le Golvan, Y., 2001. In situ and laboratory investigations of fluid flow through an argillaceous formation at different scales of space and time, Tournemire tunnel, southern France. *Hydrogeol. J.* 9, 108–123.
- Clauer, N., Chaudhuri, S., 1995. *Clays in Crustal Environments. Isotope Dating and Tracing*. Springer Verlag, Heidelberg, Berlin, New York.
- Dauzères, A., Le Bescop, P., Sardini, P., 2009. Physico-chemical investigation of cement paste degradation in a clayey environment: experimental approach and preliminary modelling. In: *Concrete in Aggressive Aqueous Environment, Performance, Testing and Modelling*, 3–5 June 2009, Toulouse, France, RILEM Proc. PRO 63, 1, pp. 228–239.
- Dauzères, A., Le Bescop, P., Sardini, P., Cau Dit Coumes, C., 2010a. Physico-chemical investigation of clayey/cement-based materials interaction in the context of geological waste disposal, experimental approach and results. *Cem. Concr. Res.* 40, 1327–1340.
- Dauzères, A., Le Bescop, P., Sardini, P., Cau Dit Coumes, C., 2010b. Study of CEM I and low pH cement paste leaching in multi-ionic underground water. In: *Concrete Under Severe Conditions-environment and Loading*, CONSEC 10, Merida (Mexico) 1, pp. 495–503.
- De Windt, L., Pellegrini, D., Van Der Lee, J., 2004. Coupled modelling of cement/claystone interactions and radionuclide migration. *J. Contam. Hydrol.* 68, 165–182.
- De Windt, L., Marsal, F., Tinseau, E., Pellegrini, D., 2008. Reactive transport modeling of geochemical interactions at a concrete/argillite interface, Tournemire site (France). *Phys. Chem. Earth* 33, 295–305.
- Elie, M., Techer, I., Trotignon, L., Khoury, H., Salameh, E., Vandamme, D., Bouvais, P., Fourcade, S., 2007. Cementation of kerogen-rich marls by alkaline fluids released during weathering of thermally metamorphosed marly sediments. Part II: Organic matter evolution, magnetic susceptibility and metals (Ti, Cr, Fe) at the Khushaym Matruk natural analogue (Central Jordan). *Appl. Geochem.* 22, 1311–1328.
- Faucou, P., Adenot, F., Jacquinet, J.F., Petit, J.C., Cabrilac, R., Jorda, M., 1998. Long-term behaviour of cement pastes used for nuclear waste disposal: review of physico-chemical mechanisms of water degradation. *Cem. Concr. Res.* 28, 847–857.
- Faure, G., 1986. *Principles of Isotope Geology*, second ed. John Wiley & Sons.
- Fernandez, R., Cuevas, J., Mader, U., 2009a. Modelling concrete interaction with a bentonite barrier. *Eur. J. Mineral.* 21, 177–191.
- Fernandez, R., Mader, U., Rodriguez, M., Vigil De La Villa, R., Cuevas, J., 2009b. Alteration of compacted bentonite by diffusion of highly alkaline solutions. *Eur. J. Mineral.* 21, 725–735.
- Fourcade, S., Techer, I., Bouvais, P., Elie, M., Trotignon, L., Vandamme, D., 2007. Cementation of a clayey sediment by alkaline fluids released by combustion-produced cements: I. Isotopic (C, O) and paleomagnetic study of the Khushaym Matruk natural analogue (Central Jordan). *Appl. Geochem.* 22, 1293–1310.
- Gaboreau, S., Prêt, D., Tinseau, E., Claret, F., Pellegrini, D., Stammose, D., in press. 15 Years of in situ cement-argillite interaction from Tournemire URL: characterisation of the multi-scale spatial heterogeneities of pore space evolution. *Appl. Geochem.* doi:10.1016/j.apgeochem.2011.07.013.
- Gaucher, E., Blanc, P., 2006. Cement/clay interactions – a review: experiments, natural analogues, and modeling. *Waste Manage.* 26, 776–788.
- Hewlett, P.C., 1998. *Leas's Chemistry of Cement and Concrete*, fourth ed. Elsevier, Butterworth-Heinemann.
- Iden, I.K., Hagelii, P., 2003. C, O and S isotopic signatures in concrete which have suffered thaumasite formation and limited thaumasite form of sulfate attack. *Cem. Concr. Compos.* 25, 839–846.
- Khoury, H., Salameh, E., Abdul-Jaber, Q., 1985. Characteristics of an unusual highly alkaline water from the Maqarin area, Northern Jordan. *J. Hydrol.* 81, 79–91.
- Le Bescop, P., Sotet, C., 2006. External Sulphate Attack by Ground Water. Revue Européenne de Génie Civil.
- Letolle, R., Gegout, P., Moranville-Regourd, M., 1992. Stable isotopes of carbon and oxygen for the study of carbonation/decarbonation processes in concretes. *Cem. Concr. Res.* 22 (2–3), 235–240.
- Linklater, C.M., Albinsson, Y., Alexander, W.R., Casa, I., McKinley, I.G., Sellin, P., 1996. A natural analogue of high-pH cement pore waters from the Maqarin area of northern Jordan: comparison of predicted and observed trace-element chemistry of uranium and selenium. *J. Contam. Hydrol.* 21, 59–69.
- McCrea, J.M., 1950. On the isotope chemistry of carbonates and a paleotemperature scale. *J. Chem. Phys.* 18, 849–857.
- Michel, O., 1999. Caractérisation isotopique Rb/Sr et Pb/Pb des roches totales, des minéraux de remplissage de fracture et des eaux des formations sédimentaires jurassiques de Tournemire (Aveyron, France). Implications sur les interactions Eau/Roches passées et actuelles. Thèse de doctorat de l'Université de Montpellier II.
- Mitodowski, A.E., Hyslop, E.K., Khoury, H., Hugues, C., Mader, U.K., Griffault, L., Trotignon, L., 2001. Mineralogical alteration by hyper-alkaline groundwaters in northern Jordan. In: Cidu, R. (Ed.), *Proc. 10th Internat. Symp. Water-Rock Interaction - WRI 10*, Balkema, pp. 1347–1350.
- Myneni, S.C.B., Traina, S.T., Logan, T.J., 1998. Ettringite solubility and geochemistry of the Ca(OH)₂-Al₂(SO₄)-H₂O system at 1 atm pressure and 298 K. *Chem. Geol.* 148, 1–19.
- Patriarche, D., 2001. Caractérisation et modélisation des transferts de traceurs naturels dans les argillites de Tournemire. Thèse de doctorat Hydrologie et

- hydrogéologie quantitatives, CIG - Centre d'Informatique Géologique, Paristech, ENSMP.
- Perkins, R.P., Palmer, C.D., 1999. Solubility of ettringite ($\text{Ca}_3\text{Al}(\text{OH})_6(\text{SO}_4)_2 \cdot 26\text{H}_2\text{O}$) at 75 °C. *Geochim. Cosmochim. Acta* 63, 1969–1980.
- Pin, C., Joannon, S., Bosq, C., Le Fèvre, B., Gauthier, P.J., 2003. Precise determination of K, Sr, Ba and Pb in geological materials by isotope dilution and ICP-quadrupole mass spectrometry following selective separation of the analytes. *J. Anal. At. Spectrom.* 18, 135–141.
- Planel, D., Sercombe, J., Le Bescop, P., Adenot, F., Torrenti, J.M., 2006. Long-term performance of cement paste during combined calcium leaching-sulfate attack: kinetics and size effect. *Cem. Concr. Res.* 36, 137–143.
- Ramirez, S.R., Cuevas, J., Vigil, R., Leguey, S., 2002. Hydrothermal alteration of "La Serrata" bentonite (Almeria, Spain) by alkaline solutions. *Appl. Clay Sci.* 21, 257–269.
- Ramirez, S.R., Vieillard, P., Bouchet, A., Cassagnabere, A., Meunier, A., Jacquot, E., 2005. Alteration of the Callovo-Oxfordian clay from Meuse-Haute-Marne underground laboratory (France) by alkaline solutions. I. A XRD and CEC study. *Appl. Geochem.* 20, 89–99.
- Samuel, J., Rouault, K., Besnus, Y., 1985. Analyse multi-élémentaire standardisée des matériaux géologiques en spectrométrie d'émission par plasma à couplage inductif. *Analysis* 13, 312–317.
- Savage, D., Noy, D., Mihara, M., 2002. Modelling the interaction of bentonite with hyperalkaline fluids. *Appl. Geochem.* 17, 207–223.
- Taylor, H.F.W., 1997. *Cement Chemistry*. Academic Press, London, UK.
- Techer, I., Khoury, H.N., Rassineux, F., Salameh, E., Claude, C., Clauer, N., Pagel, M., Lanceiot, J., Griffault, L., Jacquot, E., 2006. Propagation of alkaline fluids in an argillaceous formation: study of the Khushaym Matruk natural analogue (Central Jordan). *J. Geochem. Explor.* 90, 53–67.
- Techer, I., Clauer, N., Liewig, N., 2009. Ageing effect on the mineral and chemical composition of Opalinus Clays (Mont Terri, Switzerland) after excavation and surface storage. *Appl. Geochem.* 24, 2000–2014.
- Tiaseau, E., Bartier, D., Hassouta, L., Devol-Brown, I., Stammose, D., 2006. Mineralogical characterization of the Tournemire argillite after in situ interaction with concretes. *Waste Manage.* 26, 789–800.
- Van der Lee, Y., 1998. *Thermodynamic and Mathematical Concepts of CHES*. Technical Report Nr LHM/RD/98/39. Ecole des Mines de Paris, Fontainebleau, France.
- Vieillard, P., Rassineux, F., 1992. Thermodynamic and geochemical modeling of the alteration of two cement matrices. *Appl. Geochem.* 7 (Suppl. 1), 125–136.
- Zheng, Y.F., 1994. Oxygen isotope fractionation in metal monoxides. *Mineral. Mag.* 58A, 1000–1001.