



HAL
open science

Pentevrian and brioverian in the North armorican massif: characteristics and geodynamic evolution

Bernard Auvray

► **To cite this version:**

Bernard Auvray. Pentevrian and brioverian in the North armorican massif: characteristics and geodynamic evolution . 1989, pp.9 + figures. insu-01576103

HAL Id: insu-01576103

<https://insu.hal.science/insu-01576103v1>

Submitted on 22 Aug 2017

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

"PENTEVRIAN AND BRIOVERIAN IN THE NORTH ARMORICAN MASSIF :
CHARACTERISTICS AND GEODYNAMIC EVOLUTION".

Bernard AUVRAY

Abstract :

The North Armorican Massif is one of the rare areas within the Hercynian belt of Western Europe where the Precambrian basement outcrops. Within this basement, two main cycles may be distinguished (1) The Pentevrian or Icartian (2.2 - 1.8 Ga) composed of metamorphic rocks mainly derived from magmatic protoliths. (2) The Brioverian which includes several sedimentary and volcanic units among which stratigraphic correlations are not perfectly established. The numerous plutons emplaced during this period allow a chronology of the main tectonometamorphic events to be established.

Nowadays, the Late Precambrian evolution (660 Ma - 540 Ma) of the North Armorican domain is interpreted as resulting from a subduction geodynamic process, but the vergence of the subduction remains a matter for discussion. Nevertheless, most of the major structures now visible in this domain are related to this Precambrian history.

AUVRAY Bernard
Université de Rennes I
Institut de Géologie
Campus de Beaulieu
35042 - RENNES Cédex

Poste de Professeur
Université de Rennes I
Section CNU : 35ème
J.O. n°16 du 19/01/89

The first comprehensive synthesis concerning the Precambrian terrains in the North Armorican Massif has been proposed by Cogné, in 1962, mainly based on the cartographic data of Barrois, Pruvost, Graindor and Cogné himself. As it is possible to observe on fig.1, the Precambrian in this area was subdivided into two main cycles (Cogné 1962).

(1) A highly metamorphosed basement, granodioritic in composition in evidence at Jospinet on the eastern flank of the Baie de St Brieuc and called by Cogné the Pentevrian (Cogné 1959).

(2) This gneissic basement was unconformably overlain by the Brioverian composed of two stages separated by an orogenic phase : at the base, a volcano-sedimentary sequence; at the upper part, and exclusively sedimentary series. The main phase of metamorphism, deformation and granitisation occurred at the end of the first stage and was achieved at the end of the second stage with the intrusion of the Mancellian granites.

Fig.1 →

The first questions concerning the validity of this scheme appeared in 1971 when a new datation of the Erquy volcanic formation, considered as a reference level for the base of the Brioverian, yielded a Lower Paleozoic age of about 480 Ma (Vidal et al. 1971). One of the main consequences of this new radiometric data was to introduce some doubt concerning the existence of Pentevrian at least in the area where it was defined and also to introduce confusion in the Brioverian stratigraphy.

An investigation among the voluminous bibliography concerning the Northern part of the Armorican Massif demonstrates :

(1) That a large amount of various available data have been obtained during the last twenty years.

(2) That the precambrian stratigraphy, especially the Brioverian stratigraphy, has been strongly modified.

(3) That all the authors try to explain the geodynamic evolution during the Precambrian by using a plate tectonic model and more precisely a subduction model. But there is no complete agreement concerning the direction of this subduction.

Consequently, it seems now necessary to separate the facts which seem well established from interpretations and hypothesis; this is the main purpose of this paper.

II - MAIN CHARACTERISTICS OF THE PRECAMBRIAN FORMATION IN THE NORTH ARMORICAN DOMAIN :

Let us define the main characteristics of the Precambrian in the North Armorican Massif and first of all the Pentevrian.

(A) The Pentevrian :

The only area where the Pentevrian formations are well-defined is located at the northern margin of the North Armorican Massif, along a line starting from the Tregor area to Cap de la Hague in the Cotentin Peninsula (figure 2).

Fig. 2 → The radiometric data obtained by Adams (1967, 1976), Calvez and Vidal (1978), Auvray *et al.* (1980), Piton (1985), yield a Lower Proterozoic age ranging from 2,2 by to 1,8 by. The Pentevrian, particularly well exposed in Guernsey and in the Trebeurden area near to Lannion is mainly composed of metavolcanics, metasediments and metagranitoids. Metavolcanics, varying from basic to acidic types (amphibolites to acidic gneisses) are largely dominant; the only well developed sedimentary sequences being the Jerbourg and Pleinmount metasediments in Guernsey (Roach 1966). Granitoids, actually transformed into augen-gneisses, are the most typical facies with compositions varying from granodiorite to quartz-monzonite. All the magmatic rocks (including volcanics and plutonites) show a calc-alkaline evolution in the AFM diagram (fig. 3a) and in the Tarney diagram (fig. 3b). All the formations have been metamorphosed in the amphibolite facies and in places, it is possible to observe the first stages of a migmatisation. Because of the scattering and small extent of the outcrops, it is impossible to define a large scale metamorphic zoneography and structure in this Lower Proterozoic basement. The ages have been determined on the augen-gneisses and consequently are a minimum age for the volcanics and the sediments intruded by these granitoids.

Fig. 3 → Nowhere else in the Armorican Massif has this basement been recognized with certainty, especially in the Jospinet area where it was defined for the first time. At that place, there is no radiometric data and no geological evidence, the main argument (that is the Brioverian age of the volcanic series of Erquy) being highly doubtful.

Finally, concerning the name of this period, it is proposed to use the term Pentevrian as equivalent of Icartian, considering that in its original definition Pentevrian included all the terrains older than the Brioverian.

4

(B) The Brioverian : Many attempts have been made to establish a comprehensive stratigraphy among the various lithologic units belonging to this stage. Actually, these correlations are not well defined.

At present, with the exception of the Perelle Gneisses in Guernsey which have yielded an age around 700 Ma (Wilson 1986, unpublished data), the oldest age in the Brioverian has been obtained on a granitic pebble of the Cesson Conglomerate and is around 660 Ma (Guerrot et al. 1986). Consequently, this conglomerate, which is considered as the lower part of the Brioverian, has a maximum age of 660 Ma.

Fig. 4 → All over the Baie de St Briec area, from the Tregor to St Malo, are exposed numerous lithological units composed of volcanics and sedimentary rocks. The correlations are difficult to establish between these units since they actually belong to tectonic blocks joined side by side during the Cadomian tectogenesis. The most recent syntheses clearly show that there is a tendency to consider that the variations observed in the different units result from lateral facies variations more than vertical ones. This concept is clearly expressed for instance in the stratigraphic table proposed by Chantraine et al. (1988) and Dissler et al. (1988) (fig.4) and also by Balé (1986) to explain the correlations between the various units outcropping in the South part of the Baie de St Briec (fig.5).

Fig. 5 → These units have suffered various degrees of metamorphism from greenschist facies (in the Tregor or Lamballe series) to amphibolite facies in the Lanvollon or Yffiniac formations.

The facts which seem well established are as follows (fig.4) :

(1) The volcanics are very abundant in the northern part of the domain and decrease towards the South-South East in the St Cast, St Malo and Lamballe series.

(2) Among the sedimentary rocks, the black cherts or "phtanites" are considered as a good stratigraphic marker. Consequently, Le Corre (1977) has proposed to use the term "Phtanitic Brioverian" to designate all these units which belong to the Lower Brioverian.

(3) During this first stage of Brioverian, and with the exception of the 660 Ma age of a granodioritic pebble in the Cesson conglomerate, radiometric data are the following :

(a) The 615 Ma age of magmatism in the Tregor area (Graviou 1974).

(b) The 600 Ma age of the basic-ultrabasic sequence of Belle-Isle-en-Terre - Yffiniac (Peucat et al. 1981 ; Guerrot & Peucat this volume).

(4) Except the St Malo - St Cast and Lamballe units, the tectonic and metamorphic evolution ended around 590 Ma, the age of the syntectonic emplacement of the dioritic massifs like that of Coëtmieux - Fort-la-Latte (Vidal et al. 1974 ; Genestier 1984 ; Balé 1986).

(5) Eastwards, in Normandy, appears the same phtanitic Brioverian including large amounts of volcanics (Le Vast, Montsurvent, La Terrette and les Mortes Femmes séries, Dupret 1986 ; Dissler 1987) metamorphosed in the greenschist facies; in this area, the minimum age for the "Phtanitic Brioverian", is given by the emplacement of the Coutances diorite at 585 Ma (Guerrot et al., 1986).

After the Phtanitic cycle, occurs à flysch-type sedimentation with rare volcanic intercalations like the Vassy spilites in Normandy (Dupret et al. 1985). This post-phtanitic cycle is well developed in Normandy (Granville and La Laize series) but is only known in Brittany in the small area of Le Minard near to Paimpol (Denis 1987). The only appreciable metamorphism in this formation is the metamorphism developed at the contact of the Mancellian granites which were emplaced at 540 Ma (Graviou et al. 1988). This age indicates the end of the Cadomian evolution. During this post-phtanitic stage, the St-Malo - Lamballe and Rance series suffered a high temperature - low pressure metamorphism (Brun & Martin 1978) characterized by the appearance of migmatites in the St Malo area, the age of the metamorphism being 540 Ma (Peucat 1982, 1986 ; Peucat & Martin 1985).

The post-phtanitic Brioverian occurs, then, between 590 Ma and 540 Ma. This time range includes the magmatism dated around 550 Ma in the Northern part of the Armorican Massif (Auvray 1975). The ignimbritic volcanics are characteristic of this magmatism and in Normandy, at St Germain Le Gaillard, are directly overlain by the Lower Cambrian formations.

In short, according to the available radiometric data, the Brioverian cycle covers a period of about 120 Ma, the major tectonic events occurring since 590 Ma.

Consequently, the geodynamic models are constrained by the main facts summarized in the table I.

Tabl. I →

III - GEODYNAMIC MODELS - DISCUSSION :

It was at the end of the seventies that the first plate tectonic models were proposed to explain the Precambrian evolution in the North Armorican Massif. Fig.6 shows an example of this modelisation proposed by Auvray (1979) with an oceanic plate plunging under a continental margin constituted by the North Armorican Domain.

Fig. 6 →

The models were based on the following main arguments :

- (1) The existence of an oceanic domain outlined by a well-marked magnetic anomaly, trending SW-NE in the middle part of the Channel (Lefort 1975, 1977).
- (2) The presence of an old sialic basement at the south margin of this oceanic domain.
- (3) Emplacement in this continental margin of volcanics and plutonics of calc-alkaline affinity largely developed in the Tregor area (fig. 7).

Fig. 7 →

This calc-alkaline affinity was the strongest argument favouring the subduction model, with a subduction plane dipping towards S-SE, the melting of the plate itself or of the above mantle wedge giving rise to the calc-alkaline magmatism.

At the same time, the age 600 Ma for the Belle-Isle-en-Terre Complex of oceanic affinity (Peucat et al. 1981), appeared to confirm the opening of a marginal basin south of the active margin, as was previously proposed by Hirbec (1979) and Auvray (1979).

Since this first proposal, numerous data concerning the characteristics of the sedimentation, the geochemistry of the volcanics, the geochemical and chronological zonations in the plutonism have allowed improvements to the model (see Chantraine et al. 1988, for references). Fig. 8, for instance, shows the regional zonation of the plutonites as given by Graviou and Auvray (1985 ; this volume).

Fig. 8 →

More recently, Balé and Brun (1983) and Balé (1986), using structural arguments have demonstrated in the Baie de St Brieuc area :

- (1) That it is possible to correlate the Belle-Isle-en-Terre and Yffiniac formations, correlation confirmed by datation, (Guerrot and Peucat, this volume) or the migmatitic belt of St Malo with the migmatitic series of Guingamp as shown on fig. 9.

Fig. 9 →

- (2) That the major Cadomian structures may be interpreted as interactions between imbricate thrust zones towards the S-SW and ductile sinistral wrench zones trending N 50°. This type of deformation has been related to an oblique

subduction followed by collision and obduction which developed continuously from 600 Ma until 540 Ma. In fig. 10, it is possible to observe the resulting structures in the Brioverian formations in the St Brieuc area.

Fig. 10 →

The result of this new data, is that in more recent geodynamic models two domains are distinguished (fig. 11) :

(a) in the North, a volcanogenic domain characterized by its calc-alkaline magmatism and its M-type plutonisms with a mean age of emplacement around 600 Ma but in fact scattered between 620 Ma and 580 Ma.

(b) to the South a terrigenous domain characterized by the lack (except the Vassy spilites in Normandy, fig.4) of volcanism and the nature and age of its plutonism (C-type and 540 Ma in age).

(c) between these two domains, a suture zone represents the closure of the Belle-Isle-en-Terre - Yffiniac back-arc basin.

Fig. 11 →

All these concepts are more or less taken in account in the various models proposed by Dissler et al. (1988) to explain the Precambrian evolution in Normandy (fig. 12) and by Chantraine et al. (1988) (fig. 13) or Graviou et al. (1988) (fig. 14) to explain the same evolution all over the North Armorican Domain. In the three models, the subduction plane is dipping towards the South-East. One of the main differences is the more complicated arc-system in the model of Dissler et al. (1988 ; fig. 12). The existence of a fore-arc, an intra-arc and a back-arc basin results from the various geochemical characteristics of numerous volcanic series (Le Vast, Montsurvent, la Terrette, Les Mortes Femmes series). In fact, some of the chemical data are conflicting and need further determinations; consequently, this arc-model has perhaps to be simplified in the future. Another important difference is that in the model for Normandy (Dissler et al., 1988), the vergence of the deformations is towards the North; it means that all the structures are synthetic, with regard to the direction of the subduction. In opposition, in the models of Chantraine et al. (1988) and Graviou et al. (1988), the vergence of the deformations is systematically towards the South, all the structures being antithetic compared to the subduction dipping.

Fig. 12, 13, 14 →

However, all these models with a subduction plane dipping towards the South become questionable if we consider the structural arguments of Balé and Brun (1983) and Balé (1986) :

(1) In the hypothesis of a subduction dipping towards the South, the systematic antithetic character of the structures, almost in the St Malo and St Brieuc areas, is surprising and has never been observed in modern subduction and

collision systems.

(2) If we agree with the translation hypothesis of the Leon block proposed by Balé and Brun (1986) during Paleozoic times (fig. 15), the Channel ocean is not in its original position and must be translated towards the SW, in the prolongation of the Belle-Isle - Yffiniac suture, (see Brun & Balé, this volume, for more details).

Fig. 15 →

Considering these two strong structural arguments and taking into account all the other constraints, it is possible to propose an alternative geodynamic model with a subduction dipping towards the North (fig. 16).

- To the North, a continental domain southern margin of which is represented by the Dòmenean domain but which also includes the Precambrian formations of Wales, South England and South-East Ireland; this Precambrian block (Northern continent in fig. 16) may be a part of the Armorica plate defined by the geophysicists (Van der Voo 1979).

- Then an oceanic domain of limited extension, as is proved by the lack of true ophiolitic assemblage, the non typical oceanic characters of the volcanic series, the marginal type sedimentation and the short time for the closure of this ocean, about 10 Ma.

- Finally, a Southern continent, the Mancellian domain the southern extension of which is ignored in the paper.

After the opening of the ocean before 620 Ma the closure phase begins by a subduction of the oceanic plate under the Northern continent. Then follows the collision around 590 Ma and the underplating of the southern continent under the northern one until 540 Ma. The various stages of this geodynamic evolution are schematically shown in fig. 16.

Fig. 16 →

But this new model is not, again, entirely satisfying, mainly because there is a problem of timing between the opening of the ocean and the subduction process. The subduction has a minimum age of 620 Ma, which corresponds to the age of calc-alkaline magmatism in the Tregor; this means that, at that time, the ocean must have been open, but the only radiometric of the supposed oceanic domain (Belle-Isle - Yffiniac) is at 600 Ma.

Another question concerns the vergence of the deformations (see above); it seems clearly established that, in the Baie de St-Brieuc domain, the vergence is towards the South (Balé and Brun 1983). In contrast, in Normandy, the vergence of the Cadomian deformations has been described as being towards the North. (Dissler *et al.* 1988; fig. 12). Moreover, Lefort and Bardy (1987), on the base of geophysical arguments yielded by the SWAT-ECORS

seismic profiles conclude that the main reflectors dipping towards the South observed in the Channel, are the traces of synthetic deformations related to Cadomian subduction of the Channel Ocean. This is an important problem which needs to be clarified.

Fig. 17 →

The non-existence of the Channel ocean which is a consequence of the Balé and Brun hypothesis, does not have a dramatic effect on south dipping subduction models because it is always possible to find another ocean, for instance, the Celtic or Longmyndian ocean as has been previously proposed by Cogné and Wright in 1980 (fig. 17). In that hypothesis, Wales, South England, Southeastern Ireland and North Armorican Massif compose a single continental margin (Northern continent, fig. 17); this would explain the similar geological and geochronological characters of these regions (Shackleton 1975 ; Dunning 1975 ; Max 1975 ; Cogné and Wright 1980 ; Thorpe et al. 1984) . The Belle-Isle - Yffiniac-Coutances suture zone becomes again the trace of a marginal basin developed during the subduction. But remains the structural problem of synthetic deformations; to my knowledge, no characteristic Cadomian structure with a northward vergence has been described in the northern margin of this Northern continent, perhaps mainly because of the strong influence of Caledonian and Hercynian tectonics in that area.

In conclusion, the problem of the vergence of the subduction is not solved for the moment. The main fact enhanced in this revue is that there is a great paradox between the large amount of various geological information acquired concerning the Precambrian domain in the North Armorican Massif and the difficulties encountered in integrating all these data in a global evolution. This is partly because, today, we observe only some fragments of a puzzle; and also perhaps, because the scale of observation is very limited and does not correspond to the continental scale required to apply global tectonic models.

REFERENCES

ADAMS, C.J.D. 1967 - A geochronological and related isotopic study of rocks from north-western France and the Channel Islands (United Kingdom). Ph. D. Thesis, Oxford, 259 p.

ADAMS, C.J.D. 1976 - Geochronology of the Channel Islands and adjacent french mainland. *Journal of Geological Society, London*, 132, 233-250.

AUVRAY, B. 1975 - Relations entre plutonisme acide et volcanisme ignimbritique : exemple des manifestations magmatiques cambriennes du Nord de la Bretagne. *Pétrologie*, 1, 125-138.

AUVRAY, B. 1979 - Genèse et évolution de la croûte continentale dans le Nord du Massif Armoricaïn. Thesis, Rennes, 680 p.

AUVRAY, B., CHARLOT, R., VIDAL, Ph. 1980 - Données nouvelles sur le Protérozoïque inférieur du domaine nord-armoricain (France) : âge et signification. *Canadian Journal of Earth Sciences*, 17, 532-538.

BALE, P. 1986 - Tectonique cadomienne en Bretagne Nord. Interaction décrochement-chevauchement : champs de déformation et modélisations expérimentales. Thesis, Rennes, 361 p.

BALE, P., BRUN, J.P. 1983 - Les chevauchements cadomiens de la Baie de St Brieuc (Massif Armoricaïn). *Comptes Rendus de l'Académie des Sciences, Paris*, 297, (II), 359-362.

BALE, P., BRUN, J.P. 1986 - Les complexes métamorphiques du Léon (NW Bretagne) : un segment du domaine éo-hercynien sud-armoricain translaté au Dévonien. *Bulletin de la Société Géologique de France*, (8), II, 3, 471-477.

BRUN, J.P., BALE, P. - Cadomian tectonics (North Brittany). *Journal of Geological Society, London*, this volume.

BRUN, J.P., MARTIN, H. 1978 - Relations métamorphisme-déformation au cours de l'évolution dynamique d'un dôme migmatitique : le massif de St Malo,

France. Bulletin de la Société Géologique de France, 7, 91-101.

CALVEZ, J.Y., VIDAL, Ph. 1978 - Two billion years old relicts in the Hercynian belt of Western Europe. Contribution to Mineralogy and Petrology, 65, 395-399.

CHANTRAINE, J., CHAUVEL, J.J., BALE, P., DENIS, E., RABU, D. 1988 - Le Briovérien (Protérozoïque supérieur à terminal) et l'orogénèse cadomienne en Bretagne (France). Bulletin de la Société Géologique de France, IV, 5, 815-829.

COGNE, J. 1959 - Données nouvelles sur l'Antécambrien dans l'Ouest de la France : Pentévrien et Briovérien en Baie de St Briec. Bulletin de la Société Géologique de France, (7), I, 112-118.

COGNE, J. 1962 - Esquisse des caractères stratigraphiques, métamorphiques, structuraux et paléontologiques de l'Antécambrien récent dans le Massif Armoricaïn. Bulletin de la Société Géologique de France, (7), IV, 413-430.

COGNE, J., WRIGHT, A.E. 1980 - L'orogénèse cadomien. Vers un essai d'interprétation paléogéodynamique... in "Géologie de l'Europe du Précambrien aux bassins sédimentaires post-hercyniens". Cogné et Slansky (ed.), 26ème Congrès Géologique International, Colloque C6, Mémoire du B.R.G.M., 108, 29-55.

DENIS, E. 1987 - Les sédiments briovériens (Protérozoïque supérieur) de Bretagne septentrionale et occidentale : nature, mise en place et évolution. Thesis, Rennes, 202 p.

DISSLER, E. 1987 - Evolution géodynamique cadomienne du Nord-Cotentin (Massif Armoricaïn). Thesis, Caen, 255 p.

DISSLER, E., DORE, F., DUPRET, L., GRESSELIN, F., LE GALL, J. 1988 - L'évolution géodynamique cadomienne du Nord-Est du Massif Armoricaïn. Bulletin de la Société Géologique de France, IV, 5, 801-814.

DUNNING, F.W. 1975 - Precambrian craton of central England and the Welsh Borders. in "Precambrian - A correlation of Precambrian rocks in the British Isles". Geological Society, London, Special Reports. N°6, 83-95.

DUPRET, L. 1986 - Le Protérozoïque du Massif Armoricaïn nord-oriental (Normandie et Maine). Bulletin de la Société Linnéenne de Normandie. (in press).

DUPRET, L., LE GALL, J., DORE, F., GATINOT, F., DISSLER, E. 1985 - les spilites de Vassy (Calvados), témoin d'un volcanisme tholéiitique et distensif, dans la sédimentation du Briovérien supérieur de Normandie. Comptes Rendus de l'Académie des Sciences, Paris, II, 300, 687-692.

GENESTIER, V. 1984 - Pétrologie d'une intrusion calco-alkaline et de ses enclaves : le massif cadomien de "Fort La Latte" (Bretagne Nord). DEA, Rennes, 52 p.

GRAVIOU, P. 1984 - Pétrogenèse des magmas calco-alkalins : exemples des granitoïdes cadomiens de la région trégorroise (Massif Armoricaïn). 3rd cycle thesis, Rennes, 236 p.

GRAVIOU, P., AUVRAY, B. 1985 - Caractérisation pétrographique et géochimique des granitoïdes cadomiens du domaine nord-armoricaïn : implications géodynamiques. Comptes Rendus de l'Académie des Sciences, Paris, 301, (II), 315-318.

GRAVIOU, P., AUVRAY, B. - Late Precambrian M-type granitoïd genesis in the Cadomian belt of N.W. France. Journal of Geological Society, London, this volume.

GRAVIOU, P., PEUCAT, J.J., AUVRAY, B., VIDAL, Ph. 1988 - L'orogénèse cadomienne dans le Nord du Massif Armoricaïn : Pétrologie et géochronologie. Hercynica, (in press).

GUERROT, C., PEUCAT, J.J., DUPRET, L. 1986 - Age du Précambrien sédimentaire (Briovérien) dans le Massif Armoricaïn. 11^{ème} Réunion des Sciences de la Terre, Clermont-Ferrand, abstract, p.86.

GUERROT, C., PEUCAT, J.J. - U-Pb geochronology in the Late Precambrian of North Brittany, France : the Cadomian orogeny. Journal of Geological Society, London, this volume.

HIRBEC, Y. 1979 - Le complexe basique de Belle-Isle-en-Terre (Côtes du Nord). Sa place dans l'évolution géodynamique du Nord du Massif Armoricaïn. 3rd cycle thesis, Rennes, 146 p.

LE CORRE, Cl. 1977 - Le Briovérien de Bretagne Centrale : essai de synthèse lithologique et structurale. Bulletin du BRGM, section I, 3, 219-254.

LEFORT, J.P. 1975 - Le socle péri-armoricain : étude géologique et géophysique du socle submergé à l'Ouest de la France. Thesis, Rennes, 217 p.

LEFORT, J.P. 1977 - Possible "Caledonian" subduction under the Domnonian domain North Armorican area. *Geology*, 5, 523-526.

LEFORT, J.P., BARDY, Ph. 1987 - Mise en évidence de chevauchements cadomiens sur le profil sismique SWAT n°10 levé en Manche. Bulletin de la Société Géologique de France, (8), III, 6, 1189-1194.

MAX, M.D. 1975 - Precambrian rocks of South-East Ireland in "Precambrian - A correlation of Precambrian rocks in the British Isles". Geological Society (London), Special Reports n°6, 96-101.

PEUCAT, J.J. 1982 - Géochronologie des roches métamorphiques (Rb-Sr et U-Pb). Exemples choisis au Groenland, en Laponie, dans le Massif armoricaïn et en Grande Kabylie. Thesis, Rennes, 172 p.

PEUCAT, J.J. 1986 - Behaviour of Rb-Sr whole rock and U-Pb zircon systems during partial melting as shown in migmatitic gneisses from the St Malo Massif, NE Brittany, France. *Journal of Geological Society*, London, 143, 875-885.

PEUCAT, J.J., HIRBEC, Y., AUVRAY, B., COGNE, J., CORNICHE, J. 1981 - Late Proterozoic zircon ages from a basic-ultrabasic complex : a possible cadomian orogenic complex in the Hercynian belt of Western Europe. *Geology*, 9, 169-173.

PEUCAT, J.J., MARTIN, H. 1985 - Are Rb-Sr thin slabs migmatite ages always meaningful? *Neues Jahrbuch für Mineralogie, Abh.*, 152, 1, 1-21.

PITON, Ph. 1985 - Etude géochimique et géochronologique de la croûte

continentale protérozoïque (1,8-2,2) au Nord du Massif Armoricaïn. DEA, Rennes, 53 p.

ROACH, R.A. 1966 - Outline and guide to the geology of Guernesey. Report and Transactions, La Société Guernesaise. XVIII, 751-776.

SHACKLETON, R.M. 1975 - Precambrian rocks of Wales in "Precambrian - A correlation of Precambrian rocks in the British Isles". Geological Society (London), Special Reports. n°6, 76-82.

THORPE, R.S., BECKINSALE, R.D., PATCHETT, P.J., PIPER, J.D.A., DAVIES, G.R., EVANS, J.A. 1984 - Crustal growth and late precambrian - early paleozoic plate tectonic evolution of England and Wales. Journal of Geological Society, London, 141, 521-536.

VAN DER VOO, R. 1979 - Paleozoic assembly of Pangea : a new plate tectonic model for the Taconic, Caledonian and hercynian orogenesis. Eos Transactions. AGU, 60, p.241.

VIDAL, Ph., AUVRAY, B., COGNE, J., HAMEURT, J., JEANNETTE, D. 1971 - Données géochronologiques sur la série spilitique d'Erquy : problèmes nouveaux à propos du Briovérien de Bretagne septentrionale. Comptes Rendus de l'Académie des Sciences, Paris, (D), 273, 132-135.

VIDAL, Ph., DEUTSCH, S., MARTINEAU, F., COGNE, J. 1974 - Nouvelles données radiométriques en Baie de St Brieuc. Le problème d'un socle antécadomien nord-armoricaïn. Comptes Rendus de l'Académie des Sciences, Paris, (D), 279, 631-634.

B. AUVRAY
Institut de Géologie
LP CNRS 4661
Université de Rennes I
Avenue du Général Leclerc
35042 - RENNES Cédex (France).

Figure Captions

Fig. 1 - Precambrian stratigraphy in the North Armorican domain according to Cogné (1962).

Fig. 2 - Location (stars) of the outcrops of the Lower Proterozoic formations at the northern margin of the North Armorican Massif (after Auvray 1979) (1) Moulin de la Rive, (2) Trébeurden, (3) Port-Béni, (4) Guernsey, (5) Sark, (6) Alderney, (7) La Hague.

Fig. 3 - Calc-alkaline affinity of igneous rocks (volcanics and plutonites) constituting the Pentevrian or Icartian basement in the Tregor area (North Brittany). (a) AFM diagram, (b) Tarney's diagram (1976) (after Auvray 1979).

Fig. 4 - Stratigraphic correlations in the Brioverian formations of the North Armorican Massif (after Chantraine et al. 1988 and Dissler et al. 1987). (1) Port-Béni, (2) Tréguier, (3) Trédrez, (4) Paimpol, (5) La Roche Derrien, (6) Le Minard, (7) Le Moulin de la Rive - Trébeurden, (8) Locquirec, (9) L'Armorique, (10) St-Michel-en-Grêves, (11) Yffiniac, (12) Cesson, (13) Lanvollon, (14) Binic, (15) St Malo, (16) Lamballe, (17) La Rance, (18) La Hague, (19) Le Vast, Monsurvent, La Terrette, Les Mortes Femmes, (20) Phtanitic Brioverian of Normandy, (21) Granville and La laize, (22) Vassy.

Fig. 5 - Correlations between the various Brioverian lithological units outcropping in the Baie de St Brieuc (from Balé 1986).

Fig. 6 - Subduction geodynamic model proposed by Auvray (1979) for the Brioverian evolution in the North Armorican domain.

Fig. 7 - Calc-alkaline character of the brioverian magmatism in the Tregor area, North Brittany (after Auvray 1979).

Fig. 8 - Geochemical and chronological zonation within the brioverian magmatism (from 620 Ma to 540 Ma) in the North Armorican Massif (after Graviou & Auvray 1985 ; Graviou & Auvray, this volume). **M** = mantle type magmatic rocks; **C** = crustal type magmatic rocks and migmatites.

Fig. 9 - Geological sketch-map of the Baie de St Brieuc area according to Balé & Brun 1983 and Balé 1986 : (1) Metamorphic and migmatitic series of Guingamp and St Malo, (2) Belle-Isle-en-Terre and Yffiniac basic-ultrabasic formations.

Fig. 10 - Tectonic interpretation of the western part of the Baie de St Brieuc area proposed by Balé & Brun 1983. There is a general overthrusting of all the lithological units towards the S-SW.

Fig. 11 - During the Brioverian evolution, two domains separated by a suture zone are distinguished in the North Armorican Massif : a volcanologic domain to the North, a terrigenous domain to the South (after Chantraine et al. 1988).

Fig. 12 - Geodynamic model proposed by Dissler et al. 1988. (1) Pentevrian basement, (2) Active volcanism, (3) Remnant volcanism, (4) Early Cadomian Granitoids, (5) Dioritic intrusions (Coutances type), (6) Late Cadomian Gabbroic intrusions, (7) Late Cadomian Granitoids, (8) Phtanites (black-shirts) formations, (9) Upper Brioverian sediments.

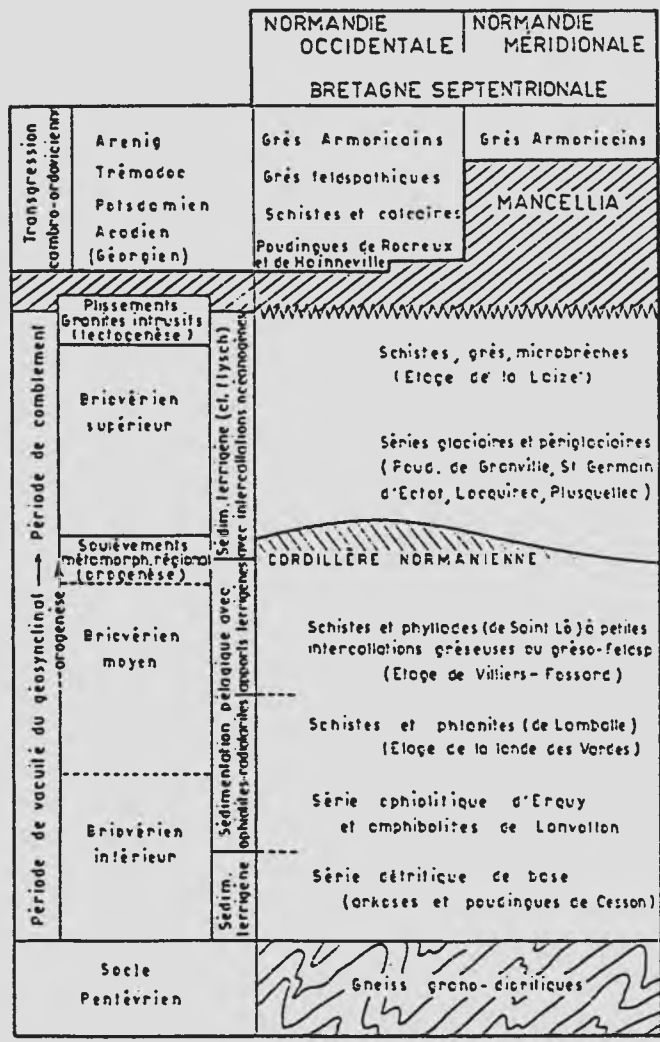
Fig. 13 - Geodynamic model proposed by Chantraine et al. 1988.

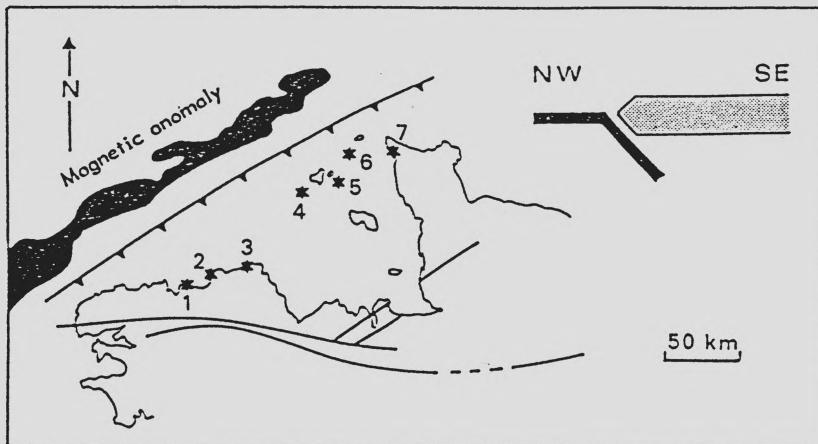
Fig. 14 - Geodynamic model proposed by Graviou et al. 1988. T = Tregor massif ; LP = Lannion-Paimpol volcanics, MF = Les Mortes Femmes volcanics; M = Montsurvent volcanics; Te = La Terrette volcanics; BI = Belle-Isle-en-Terre series; BS = Baie de St Brieuc series; SQ = St Quay diorite massif; C-FI = Coëtmieux - Fort-La-Latte diorite massif; SM = St Malo migmatite belt.

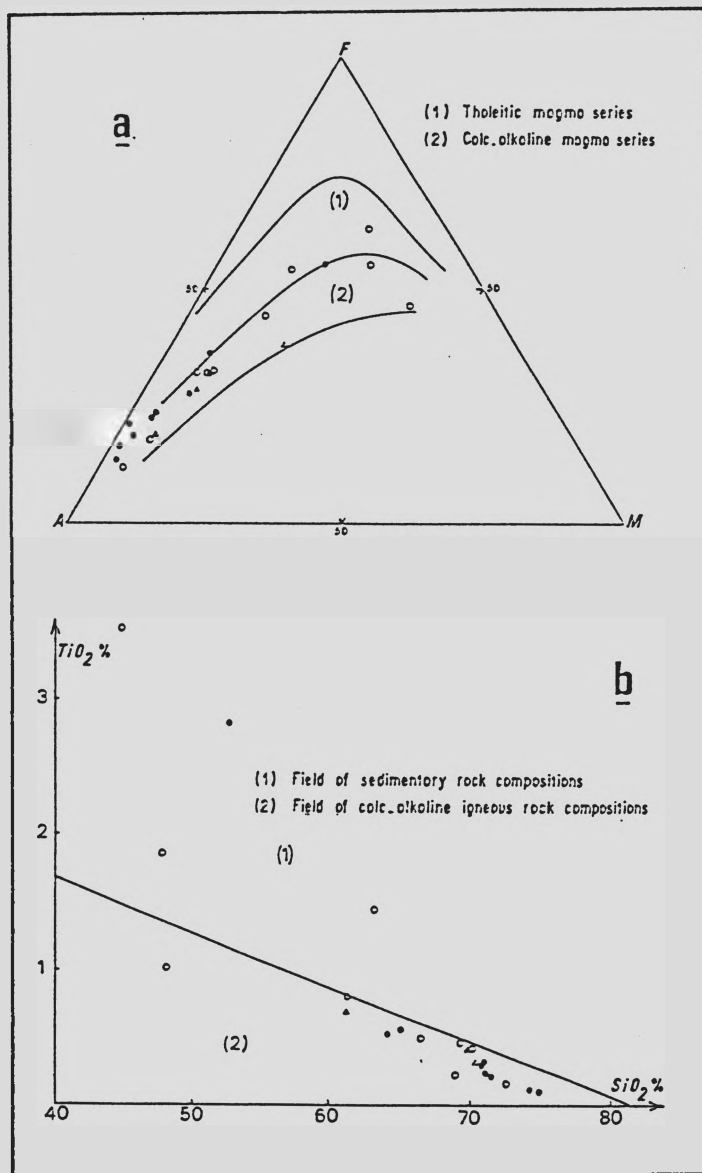
Fig. 15 - Translation of the Leon block from South to North during the Devonian (from Balé & Brun 1986).

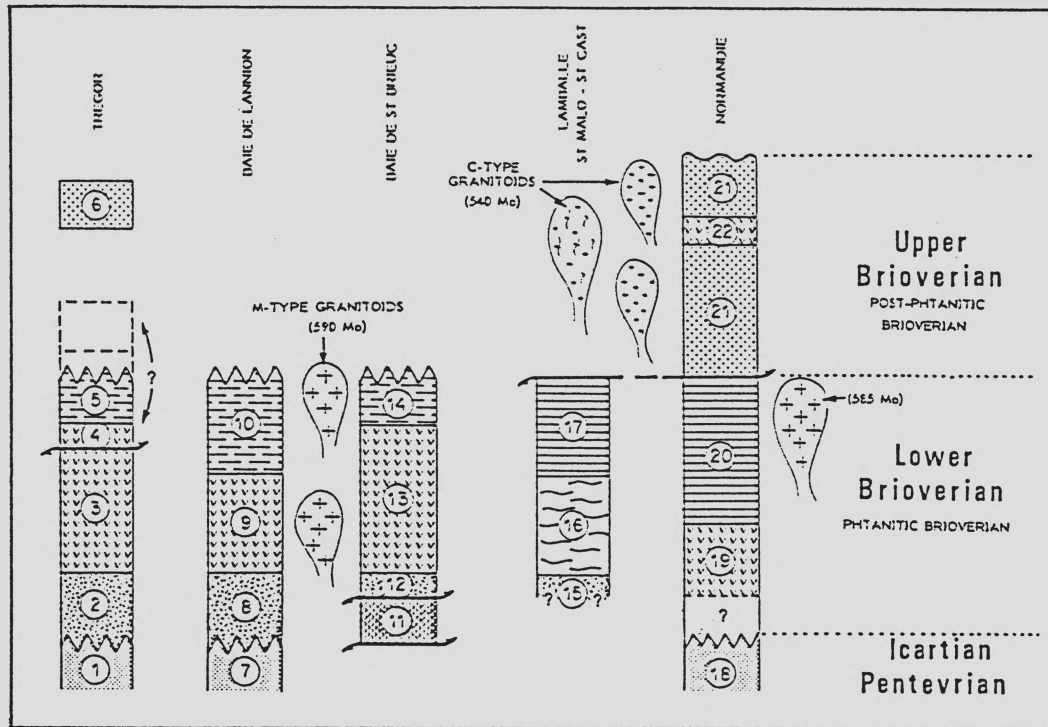
Fig. 16 - Geodynamic evolution of the North Armorican Massif explained by a subduction process with a subduction plane dipping towards N-NW.

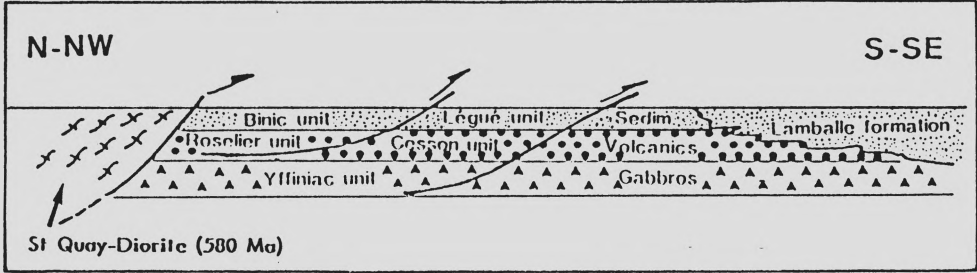
Fig. 17 - Subduction model proposed by Cogné & Wright 1980, to explain the Brioverian evolution. The closure of the Celtic Ocean induced a subduction towards S-SE under the Pentevrian micro-plate (Wales, South England, Southern Ireland and North Armorican Massif).

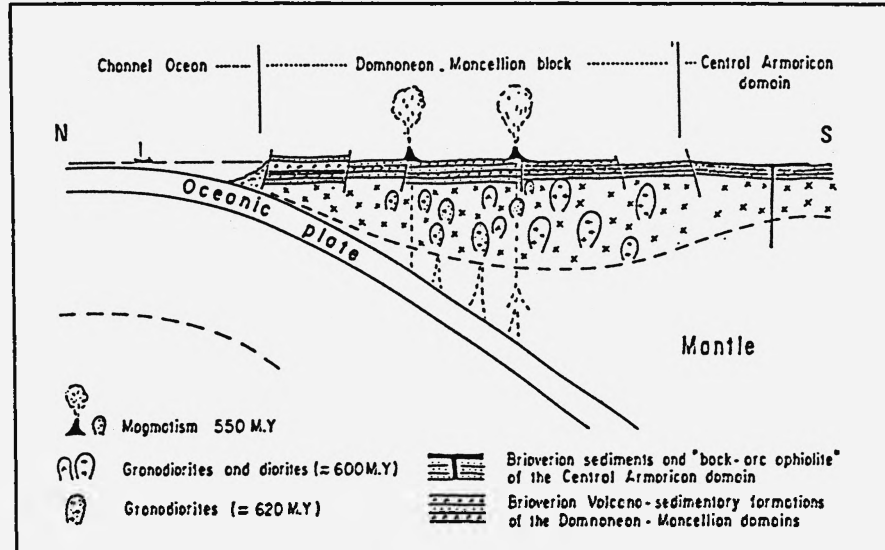


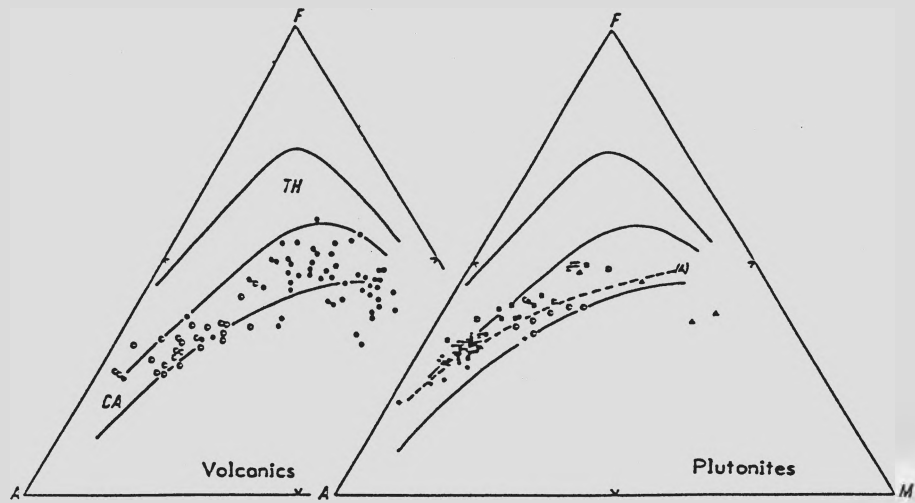


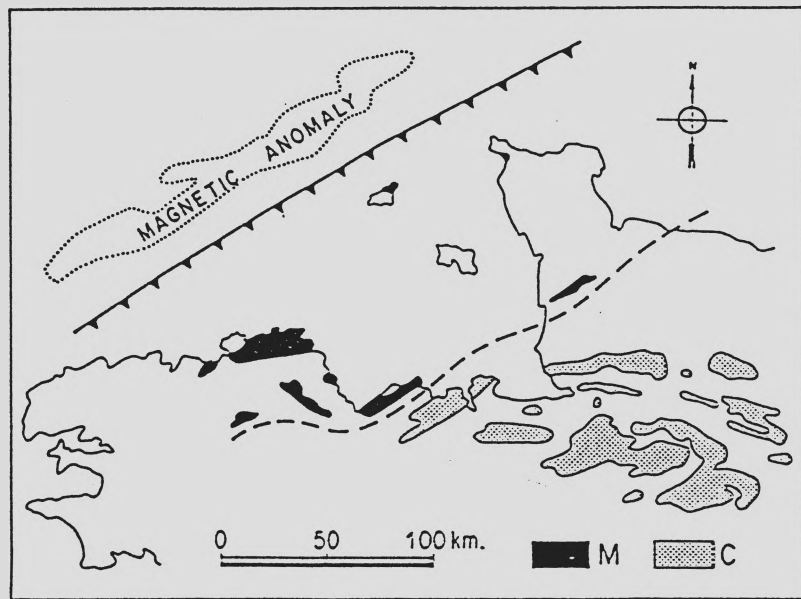


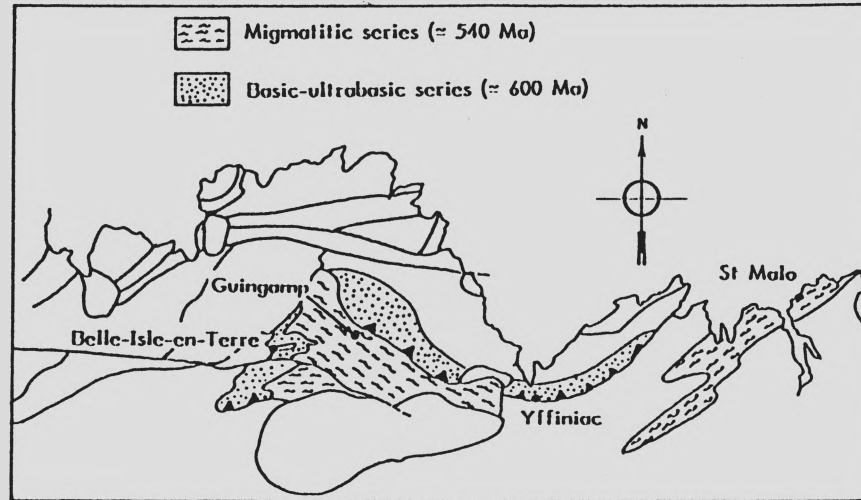


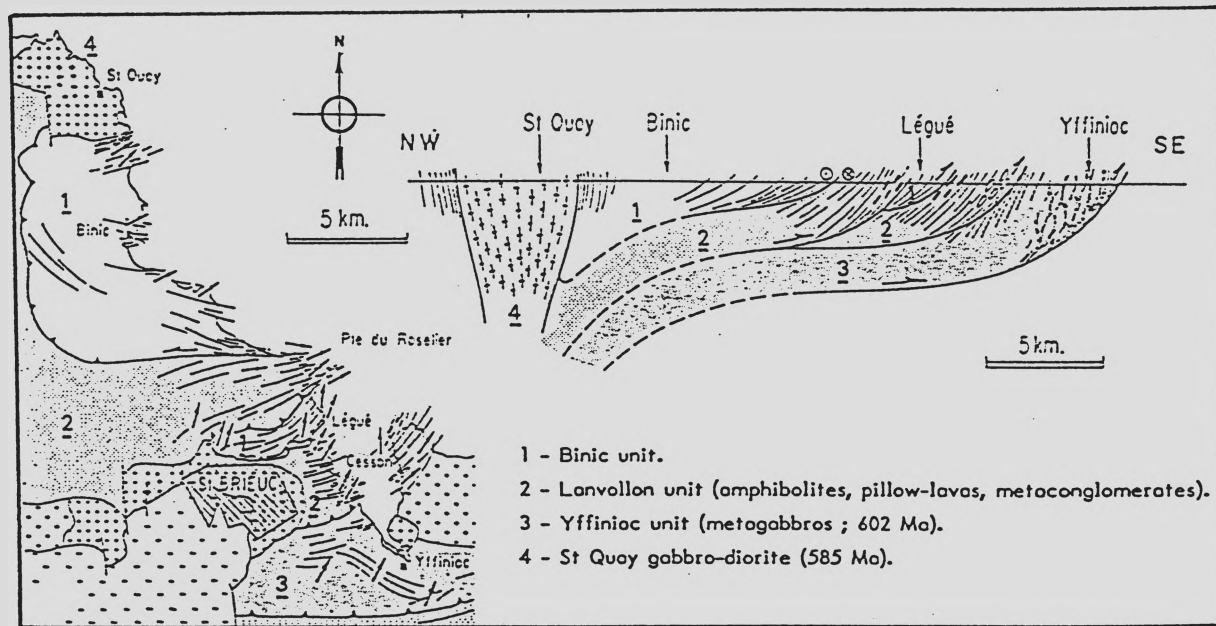


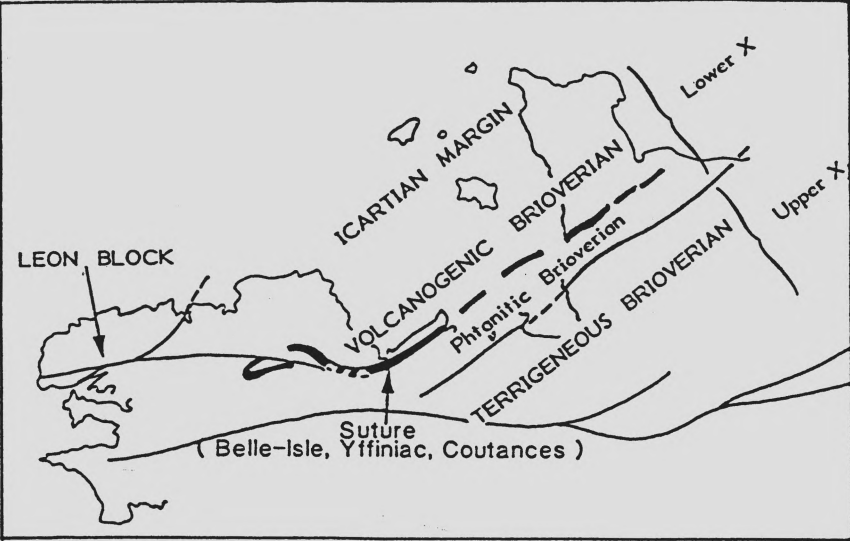


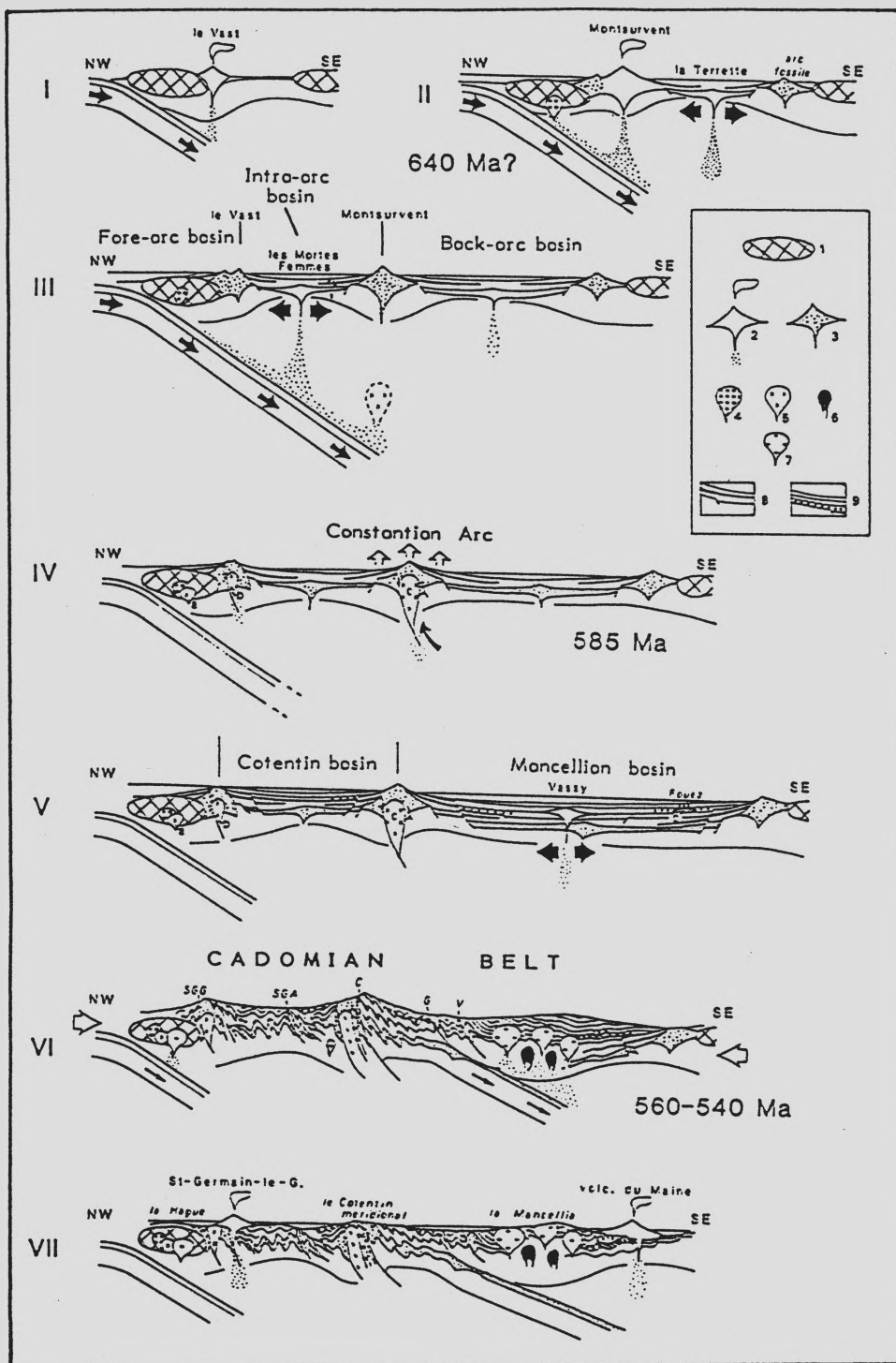


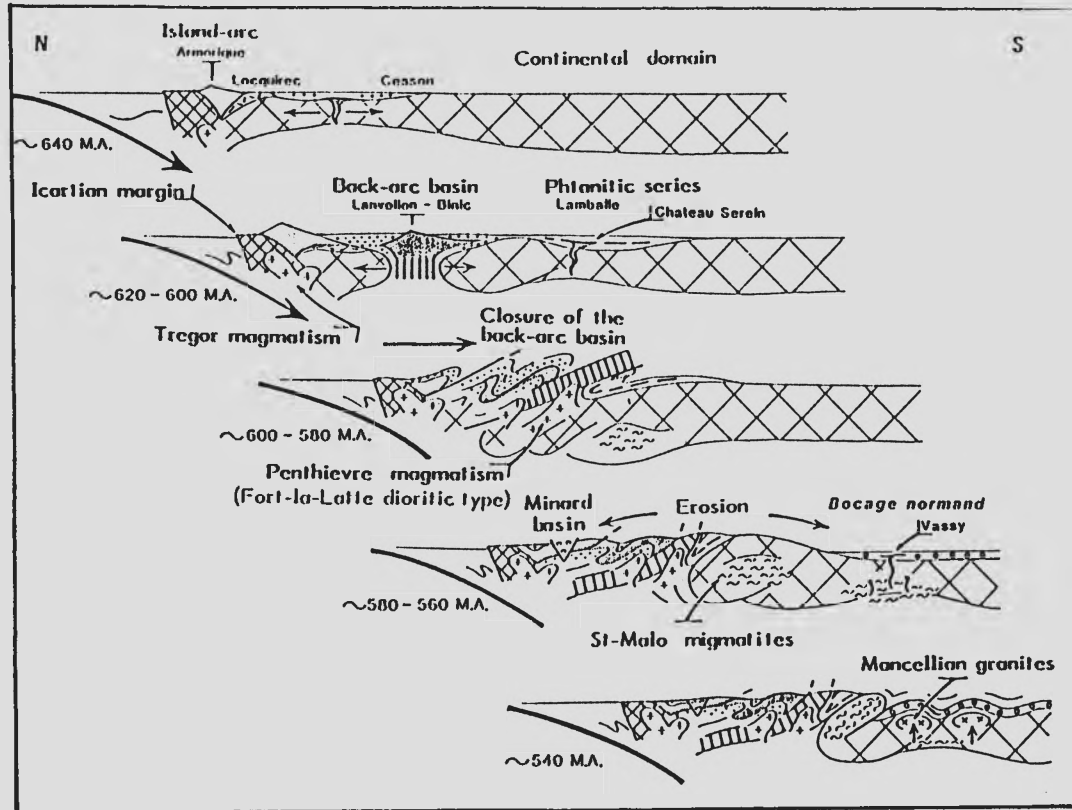


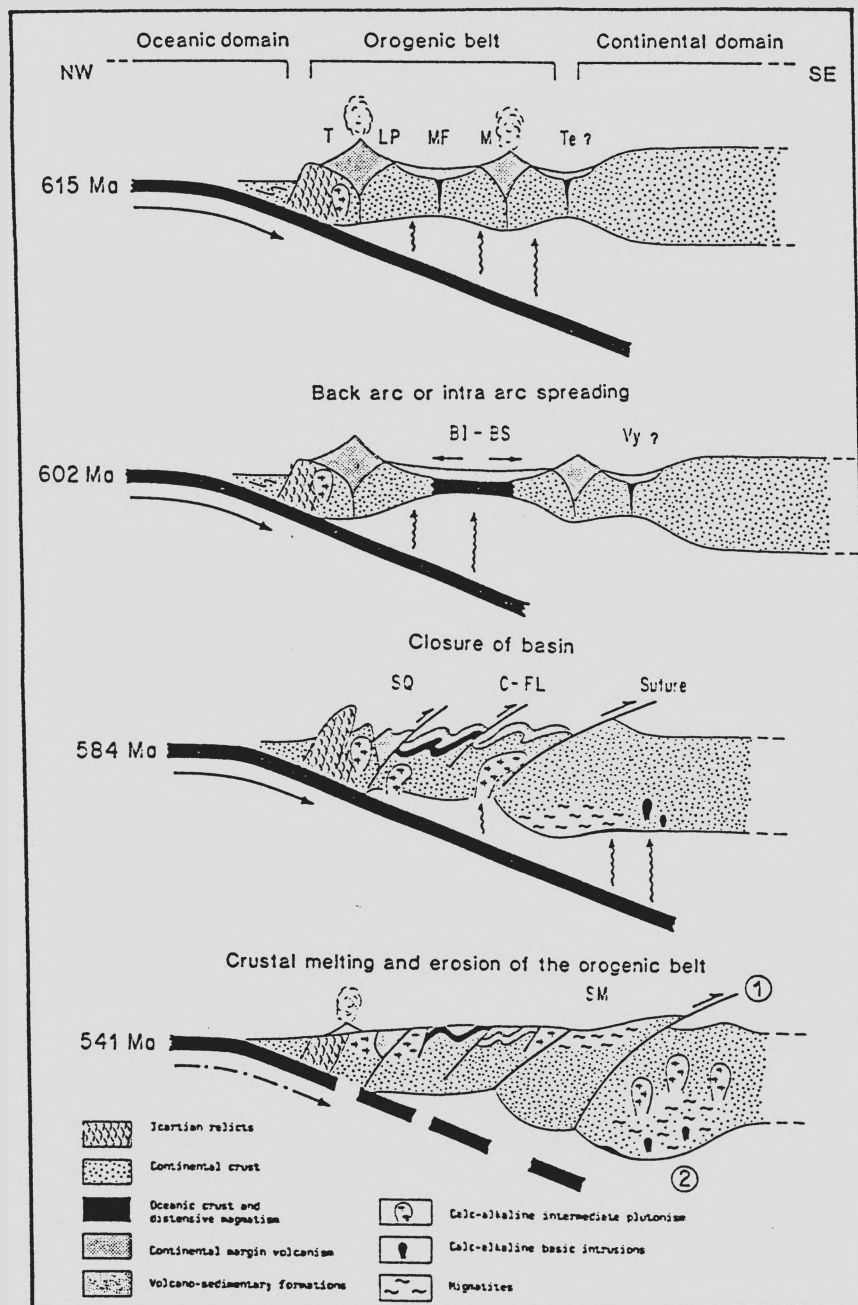


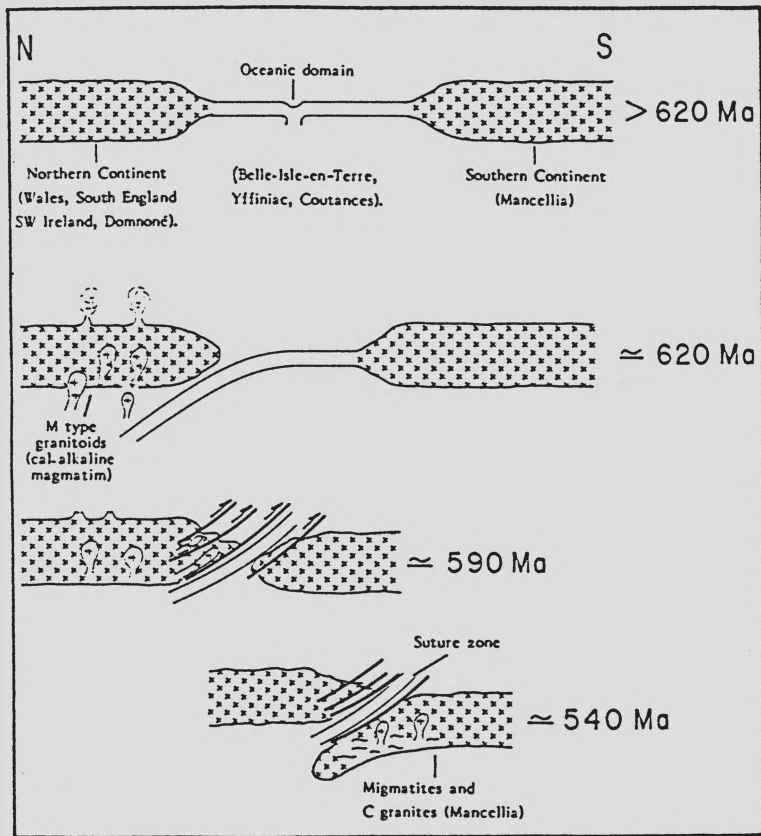


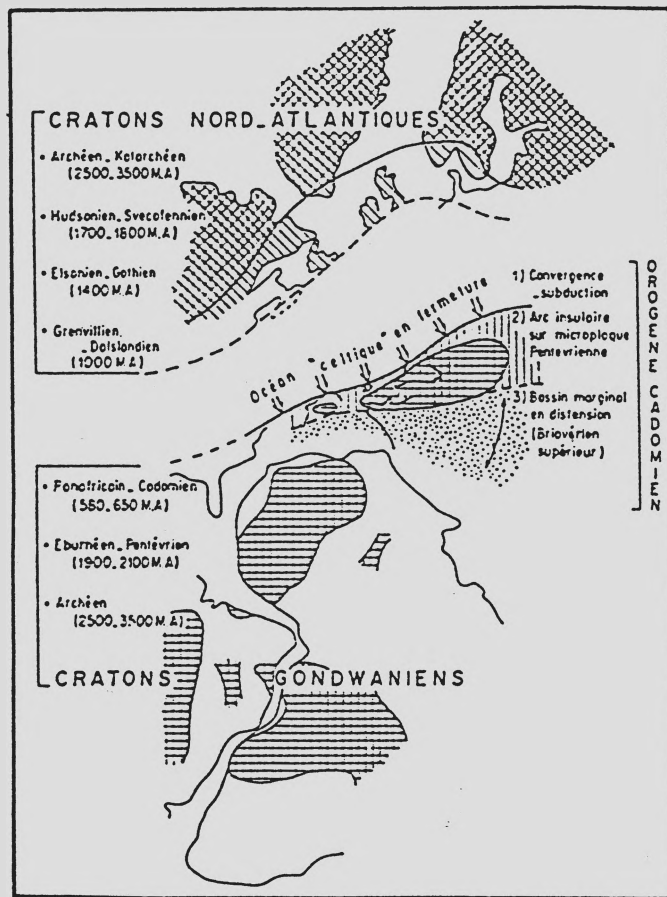












MAIN CHARACTERISTICS OF PRECAMBRIAN IN THE NORTH ARMORICAN MASSIF

- 1° - Occurrence of a gneissic basement - Lower Proterozoic age (2.0 - 1.8 Ga) - in the northern part of the area.
- 2° - Development, between 660 Ma. (perhaps 700 Ma) and 540 Ma. of a two-stage Drioverian cycle :
 - (a) - The phanitic stage (650 Ma - 590 Ma) with large amounts of volcanics at the base and sedimentary rocks (black cherts = phanites) towards the upper part.
Emplacement of plutonic rocks of calc-alkaline affinities all during this stage.
 - (b) - The post-phanitic stage (590 Ma - 540 Ma) characterized by flysch type sedimentation with rare basic volcanics (Normandy).
- 3° - Metamorphism and deformation occur during ≈ 60 Ma (600 Ma to 540 Ma)
They mainly, but irregularly, affect the formations which belong to the phanitic stage.
- 4° - Finally, the Precambrian evolution ends with the emplacement at about 550 Ma - 540 Ma.
of magmatic rocks : mainly volcanics in the North (ignimbrites)
and plutonites in the South-East (Manceillon granites).

Table 1