

Neptunium sorption and redox speciation at the illite surface under highly saline conditions

Nidhu Lal Banik, Remi Marsac, Johannes L Lützenkirchen, Christian Michael Maquardt, Kathy Dardenne, Joerg Rothe, Kerstin Bender, Horst Geckeis

▶ To cite this version:

Nidhu Lal Banik, Remi Marsac, Johannes L Lützenkirchen, Christian Michael Maquardt, Kathy Dardenne, et al.. Neptunium sorption and redox speciation at the illite surface under highly saline conditions. Geochimica et Cosmochimica Acta, 2017, 215, pp.421-431. 10.1016/j.gca.2017.08.008 . insu-01588250

HAL Id: insu-01588250 https://insu.hal.science/insu-01588250

Submitted on 15 Sep 2017

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

Neptunium sorption and redox speciation at the illite surface under highly saline conditions Nidhu lal Banik^{a,b}, Rémi Marsac^{a,c,*}, Johannes Lützenkirchen^a, Christian Michael Marquardt^a, Kathy Dardenne^a, Joerg Rothe^a, Kerstin Bender^a and Horst Geckeis^a Address: ^a Institut für Nukleare Entsorgung, Karlsruhe Institute of Technology, P.O. Box 3640, D-76021 Karlsruhe, Germany. Present addresses: ^b JRC-KARLSRUHE, G.II.6 - Nuclear Safeguards and Forensics, European Commission, P.O.Box 2340, D-76125 Karlsruhe ^c Géosciences Rennes UMR 6118, Université Rennes 1, CNRS, 35042 Rennes cedex, France. *Corresponding author: e-mail address: remi.marsac@univ-rennes1.fr Tel: +332 23 23 53 56. Fax: +332 23 23 60 90

ABSTRACT. Neptunium (Np) uptake on illite is investigated in 1 and 3.2 molal (m) NaCl 26 solutions under inert (Ar) atmosphere for 4 $< \rm pH_m < 10~(\rm pH_m$ = - log $m_{H^+})$ and 5×10^-8 <27 $[Np(V)]_{tot} < 3 \times 10^{-4}$ M. In agreement with a previous study in 0.1 m NaCl solutions [Marsac et 28 al. (2015) Geochim. Cosmochim. Acta 152, 39-51], Np(V) is the prevailing oxidation state in 29 the aqueous solution, but Np uptake by illite is affected by surface induced reduction. The 30 extent of Np(V) reduction to Np(IV) follows the measured redox potential (or the $pe = -\log \frac{1}{2}$ 31 ae.), which is influenced by the introduced Np(V) amount, because of the low redox capacity 32 of the illite. The presence of Np(IV) on the solid phase is verified by X-ray Absorption Near 33 34 Edge Spectroscopy (XANES). We can conclude that Np uptake by illite is not significantly affected by the variation of m_{NaCl} from 0.1 to 3.2 m and thus is in agreement with reports on 35 tetravalent actinide and Np(V) sorption to clays at high ionic strength. The combination of (i) 36 37 the two site protolysis non-electrostatic surface complexation and cation exchange model, (ii) the specific ion interaction theory to calculate activity coefficients for dissolved species and 38 (iii) by accounting for redox equilibria and the stability of surface Np species, the overall Np 39 uptake by illite can be simulated as a function of pH_m , pe and m_{NaCl} using a single set of 40 parameters. The present experimental and modeling results are particularly important in the 41 42 context of deep geological nuclear waste disposal since many sedimentary rocks or clay formations that are deemed suitable for this purpose exhibit highly saline porewaters. 43

Keywords: Neptunium, illite, redox, saline, high ionic strength, surface complexation,spectroscopy, geochemical modeling.

1. INTRODUCTION

The final disposal in deep geologic formations is considered the most appropriate 47 strategy to isolate high level nuclear waste (HLW) from the biosphere (e.g. OECD/NEA, 48 2008). Sedimentary clay rock formations are potential host rocks for HLW in several 49 European countries encompassing Opalinus Clay (OPA) in Switzerland (Nagra, 2002), Boom 50 and Ypresian clays in Belgium (Ondraf, 2001), Callovo-Oxfordian (COx) clays in France 51 (Andra, 2005) or Boda Claystone in Hungary (Lázár and Máthé, 2012). Indeed, clays exhibit 52 low permeability, large surface areas and pronounced sorption capacity for many relevant 53 radionuclides via ion exchange reactions as well as surface complexation to silanol and/or 54 55 aluminol groups at the edges of clay particles. Illite and smectite are the most important components of the proposed claystone host rocks and may amount to about 50 wt.% of the 56 material (Lauber et al., 2000; Gaucher et al., 2004; Bradbury and Baeyens, 2011; Pearson et 57 al., 2011; Chen et al., 2014). Although actinide sorption to clays has been extensively studied 58 59 (e.g. Bradbury and Baeyens, 2005; 2009a,b; and references therein), the underlying mechanisms were rarely investigated for ionic strengths above 0.1 M. While OPA and COx 60 porewaters exhibit ionic strengths of 0.3-0.4 M, clay rock pore waters as e.g. in the Jurassic 61 and lower Cretaceaous clay rock in Northern Germany, which are also discussed as 62 potentially appropriate host rock formations, may contain salt contents as high as about 5 M 63 (Brewitz, 1980). Sedimentary rocks currently investigated in Canada are even in contact with 64 brine solutions up to 6.5 M (Fritz and Frape, 1982). Therefore, detailed sorption investigations 65 and the development of geochemical models that can predict actinide sorption and speciation 66 67 in concentrated brine solutions are required (Vilks, 2011).

68 Neptunium (Np) is a minor constituent of high level radioactive waste. Yet, its 69 environmental chemistry is of considerable interest due to the long half-life of its main 70 isotope 237 Np (t_{1/2} = 2×10⁶ a) and the high solubility and mobility of its pentavalent state

 (NpO_2^+) under oxidizing conditions. Np is redox sensitive and its most relevant oxidation 71 72 states in the geosphere are the penta- and tetravalent states (Kim, 1986). Both redox states show drastically different geochemical behavior concerning, for instance, complexation by 73 organic and inorganic ligands, solubility, as well as sorption to minerals (e.g. Altmaier et al., 74 2013). Under reducing conditions, Np(IV) prevails and is considered rather immobile because 75 of its strong sorption to minerals and its low solubility, as other tetravalent actinides (An; e.g. 76 Th(IV)) or other tetravalent elements (e.g. Sn(IV)) (Bradbury and Baeyens, 2009a,b). Under 77 oxidizing conditions, Np(V) prevails, which weakly sorbs to minerals and is regarded as 78 rather mobile (e.g. Geckeis et al., 2013). However, partial reduction of Np(V) to Np(IV) when 79 in contact with Opalinus Clay was suggested by comparing sorption data obtained under 80 aerobic and anaerobic conditions: stronger Np sorption (Fröhlich et al., 2011) was found 81 under anaerobic conditions and significant reduction to Np(IV) was observed using 82 83 synchrotron based spectroscopic techniques (Fröhlich et al., 2012). A more recent Np-illite batch sorption study under oxygen-free argon atmosphere in 0.1 M NaCl showed that the 84 strongly adsorbing Np(IV) is thermodynamically favored at a mineral surface compared to 85 Np(V) (Marsac et al., 2015a). Although Np(V) prevailed in aqueous solution at pH \approx 7, as 86 evidenced by liquid-liquid extraction method and by capillary electrophoresis hyphenated to 87 88 inductively coupled plasma mass spectrometry (Graser et al., 2015), Np(IV) was observed at the illite surface by X-ray absorption near edge structure spectroscopy (XANES). When 89 accounting for the uptake of both Np(V) and Np(IV) on illite and the redox potential of the 90 91 illite suspension, Marsac et al. (2015a) were able to describe the overall Np uptake on illite for various redox conditions. The modeling approach was subsequently extended to the 92 sorption behavior of plutonium, which exhibits an even more complex redox chemistry than 93 Np. Pu can be found in the redox states +III to +VI under environmental conditions. The 94 experimentally determined overall Pu uptake on kaolinite (Marsac et al., 2015b) and illite 95 (Banik et al., 2016) in 0.1 M NaCl/ClO₄ was successfully reproduced. 96

The above cited studies focused on aqueous solutions of low ionic strength. At high 97 98 salt concentrations, activity coefficients of aqueous species change dramatically, whereas the effect of high ionic strengths on the surfaces properties of adsorbents and surface species is 99 100 elusive. Previous work showed that non-electrostatic models are quite suitable to predict proton and metal ion sorption at high ionic strength to naturally occurring matrices that bear 101 102 surface functional groups that are affected by pH and can be treated by the same formalism 103 and with similar numerical models. This includes marine microalgae (Schjif and Herbling, 2010; Zoll and Schjif, 2012) or bacteria (Ams et al., 2013). Eu(III) sorption to illite and 104 smectite was investigated experimentally in $0.1 < m_{NaCl} < 3.9$ molal (m) (Schnurr et al., 105 106 2015). The results obtained in the latter study showed that the 2 site protolysis non-107 electrostatic surface complexation and cation exchange (2 SPNE SC/CE) model coupled to either the specific ion interaction theory (SIT; Ciavatta, 1980) or the Pitzer formalism (Pitzer, 108 1991) to account for activity coefficients in concentrated media was able to reproduce 109 sorption data also at elevated ionic strength. By using the modeling approach of Schnurr et al. 110 (2015), the Pu-illite study in 0.1 M NaCl of Banik et al. (2016) was extended to high ionic 111 strength (0.1 < m_{NaCl} < 3.2 m; Marsac et al., 2017). In the latter studies, the Pu(IV)/Pu(III) is 112 involved. The present study is a similar extension of the work of Marsac et al. (2015a) on Np 113 uptake on illite in 0.1 M NaCl to highly saline conditions (up to $m_{NaCl} = 3.2$ m), but here a 114 An(V)/An(IV) couple is involved. Together, the work of Marsac et al. (2017) and the present 115 116 one should elucidate the actinide(III,IV,V) sorption and redox speciation at the illite surface under highly saline conditions. Classical batch experiments at various pH, total Np 117 concentrations and $m_{NaCl} = 1.0$ and 3.2 m were performed, and the redox potential was 118 monitored. XANES was applied to determine the Np redox state at the illite surface. The 119 120 coupling of the 2 SPNE SC/CE model with SIT was used to describe Np sorption and redox 121 speciation on illite under saline conditions.

2. MATERIALS AND METHODS

All chemicals were of pro analytical quality or better and were obtained from Merck 124 (Darmstadt, Germany) or Riedel de Haen (Seelze, Germany). Experiments were conducted 125 with de-ionized "MilliO" water (specific resistivity, 18.2 M Ω cm). The purified Na-illite was 126 provided within the EC project CP CatClay. The source material derives from lacustrine 127 continental sediments deposited at the Upper Eocene (~ -35 Ma) in the basin of Le Puy en 128 Velay (Massif Central, France). The purification procedures and the characterization of the 129 purified illite (<63 µm) were previously described in detail (Marsac et al., 2015). Concerning 130 redox processes, it is noted that the iron content is 0.88 mol Fe per kg of illite (Marsac et al., 131 2015). The aqueous concentration of 237 Np in sorption experiments after phase separation was 132 determined by liquid scintillation counting (LSC) using the scintillation cocktail Ultima Gold 133 XR (Packard Instruments Co., USA) and the liquid scintillation counter (LSC) Tri-Carb 134 (Packard Instruments Co., USA). Results for ²³⁷Np-α-activities were corrected for 135 interferences by the β -radiation from the daughter nuclide protactinium (²³³Pa) by applying 136 the α/β discrimination method. The procedure is described in detail by Marsac et al. (2015a) 137 and will not be repeated here. ²³³Pa is a short-lived isotope ($t_{1/2} = 27$ days) unlike ²³⁷Np ($t_{1/2} =$ 138 2.14 10⁶ years). Since the ²³³Pa mass concentration is much smaller than that of ²³⁷Np, its 139 140 competitive effect on Np sorption to illite can be neglected.

141

142 **2.1. Neptunium stock solution**

The initial ²³⁷Np solution was evaporated to dryness and dissolved in 8 M HCl. The solution was purified via anion exchange (BIO RAD[®] AG 1-X8). After a washing step with several milliliters of 8 M HCl, Pu contaminations were removed with a fresh solution of 150 mg NH₄I in 5 mL 8 M HCl. Subsequently, Np was eluted from the ion exchanger with 4 M HCl/0.05 M HF and evaporated to dryness in a PTFE beaker. The sample was fumed twice with 1 M HClO₄ and the residue was dissolved in deionized water to obtain a solution containing only NpO_2^+ , as required for all our experiments, and its purity was verified by UV-Vis/NIR spectroscopy (Sjoblom and Hindman 1951). The final Np concentration was determined by LSC.

152 **2.2. Determination of pH and Eh**

pH in the clay suspensions was measured by an Orion 525A pH meter and a Ross electrode calibrated with 4 standard buffers (pH 3, 5, 7 and 9; Merck). The uncertainty of pH measurements is \pm 0.05. For converting the pH value measured in saline solutions (I > 0.1 m) (pH_{exp}) into a chemically meaningful quantity, a correction is applied in order to obtain the molal proton concentration, i.e. $-\log m_{H^+}$ (pH_m). An empirical correction coefficient (A) which has been accurately determined by Altmaier et al. (2003) for NaCl solutions for our setup with known -log m_{H^+} was added to pH_{exp} values according to equations 1 and 2:

$$160 \quad pH_m = pH_{exp} + A_{NaCl} \tag{1}$$

161
$$A_{NaCl} = 0.0013 \times m_{NaCl}^2 + 0.1715 \times m_{NaCl} - 0.0988$$
 (2)

162 where m_{NaCl} is the molality (mol kg⁻¹) of the background electrolyte.

The redox potentials in the clay suspensions were measured using an Orion 525A (E_h 163 meter) and a Pt combined electrode with a Ag/AgCl reference (Metrohm). They were 164 converted into Eh vs. standard hydrogen electrode (SHE) by correcting for the potential of the 165 reference electrode. A commercial redox-buffer (220 mV, Schott instruments) was used for 166 calibration. An equilibration time of 15 min was allowed for all Eh measurements, after 167 having stirred the suspension. Uncertainties in Eh measurements are \pm 50 mV (Altmaier et al., 168 169 2010; Kirsch et al., 2011). Eh was converted to the logarithm of the apparent electron activity, pe = $-\log a_{e} = 16.9 \times Eh(V)$ at 25°C. To our best knowledge, no ionic strength dependent 170 171 correction needs to be performed for experimentally measured Eh values (as the one that has 172 been described for pH measurements in brines above).

173 **2.3. Batch sorption experiments**

All sorption experiments were performed as batch type experiments with initial Np(V)174 concentrations in the range of 5×10^{-8} - 3×10^{-6} M. The batch experiments were carried out in 40 175 mL polypropylene centrifuge tubes at room temperature in an argon glove box (< 1 ppm O_2 , 176 absence of CO₂). The sample volume was 25 mL. At a solid-to-liquid ratio of 2 g L⁻¹, the 177 suspensions were preconditioned at a given m_{NaCl} under continuous shaking for 4-5 days to 178 achieve a given target pH_m value by adding 0.1 M HCl or 0.1 M NaOH. After mixing the 179 NpO_2^+ solution with the preconditioned illite suspension, pH_m was readjusted by adding acid 180 or base (HCl/NaOH). The samples were shaken end-over-end. According to Marsac et al. 181 (2015a), Np(V) adsorption or reduction to Np(IV) does not significantly change with time 182 after 7 days. Therefore, after 7 days, pH_m and Eh were measured in the suspension and an 183 aliquot of each sample was centrifuged in a Beckman L7 Ultracentrifuge at 90000 rpm for 1 h. 184 The supernatant was analyzed for dissolved Np by LSC. Results obtained in batch 185 experiments will be expressed throughout as distribution coefficients (R_d in L kg⁻¹), calculated 186 by the following equation: 187

188
$$R_{d} = \left([Np]_{tot} / [Np]_{ag} - 1 \right) \times V/m$$
(3)

189 where $[Np]_{aq}$ and $[Np]_{tot}$ (mol L⁻¹) are the dissolved (final) equilibrium and total (initial) 190 concentrations of Np in solution, respectively. The term V/m corresponds to the aqueous 191 solution volume to illite mass ratio (L kg⁻¹). An uncertainty of \pm 0.3 is commonly associated 192 with log R_d determined for radionuclide sorption to clay minerals (Bradbury and Baeyens, 193 2009a), although for low [Np]_{tot} and high uptake, the uncertainty on log R_d might be larger.

194 2.4. X-ray absorption near-edge spectroscopy (XANES)

One sample was prepared as described in the previous section for Np L₃-XANES measurements for $[NpO_2^+]_{tot} = 3 \times 10^{-4} \text{ M}, \text{ m/V} = 20 \text{ g L}^{-1}, \textbf{m}_{NaCl} = 3.2 \text{ and } \text{pH}_m = 7.9 \text{ under}$

inert (Ar) atmosphere. The Eh measured after one week equilibration time was 0.43 ± 0.05 V 197 (pe = 7.3 ± 0.8). A similar experiment is described in Marsac et al. (2015a), however, 198 performed in 0.1 m NaCl and at Eh = 0.40 ± 0.05 V (pe = 6.8 ± 0.8). Conditions for XANES 199 measurements and the data treatment were identical to what has been described in detail in 200 Marsac et al. (2015a). Briefly, an aliquot was centrifuged at 3000 g in a polyethylene vial 201 (500 µL). The supernatant was discarded, but an aliquot was taken for LSC. The wet illite 202 203 paste was transferred into a vial, mounted inside an Ar-flushed cell, which acts as a second containment and keeps the sample vials in an anaerobic atmosphere (Brendebach et al., 2009). 204 Measurements were performed at the 2.5 GeV synchrotron light source ANKA, KIT, 205 206 Germany, at the INE-Beamline for actinide research (Rothe et al., 2006; 2012). A pair of Ge(422) crystals was used in the Lemonnier-type double crystal monochromator (DCM). The 207 monochromatic radiation is focused by a Rh-coated toroidal mirror to a spot of $< 1 \text{ mm} \times 1$ 208 mm at the sample position. Higher harmonic radiation in the incident beam was suppressed by 209 detuning the parallel alignment of the DCM crystals to 70 % of photon flux peak intensity in 210 the middle of the spectral range at the beginning of each scan. The incident flux was measured 211 by an Ar-filled ionization chamber at ambient pressure and held constant by a digital 212 MOSTAB feedback system. The Np L₃-edge spectra ($E(Np^0 2p_{3/2}) = 17.61$ keV) were 213 calibrated against the first derivative XANES spectrum of a Zr foil (energy of the first 214 inflection point set to $E(Zr^0 1s) = 17.998$ keV). All Np L₃-XANES spectra were recorded in 215 fluorescence detection mode by registering the Np L_{α} fluorescence yield at ~13.95 keV as a 216 function of the incident photon energy with a 5-pixel Ge solid state detector (Canberra LEGe). 217 Up to 5 scans were collected at room temperature and averaged for each sample. The 218 illuminated area of the sample was changed between each scan to minimize potential Np 219 220 redox reactions induced by the beam. XANES spectra were extracted following subtraction of a linear pre-edge background function and normalization of the edge jump to unity using the 221 Demeter Athena software (Ravel and Newville, 2005). The same energy range was 222

considered for the pre-edge and for the normalization of all XANES spectra (pre-edge: 17.3017.55 keV; post-edge: 17.77-18.20 keV). Preliminary tests ensured that the differences in the
XANES were not due to the normalization procedure.

226

227 **2.5. Thermodynamic modeling**

pe-pH diagrams for Np were obtained using PhreePlot (Kinniburgh and Cooper, 228 2009), which contains an embedded version of the geochemical speciation program 229 PHREEQC (Parkhurst and Appelo, 1999). PhreePlot also includes a parameter optimization 230 procedure, which allows fitting a model to experimental data by minimizing the weighted sum 231 of squares of the residuals. Thermodynamic constants for Np aqueous speciation were taken 232 from the NEA thermodynamic database (Guillaumont et al., 2003). The specific ion 233 interaction theory (SIT; Ciavatta, 1980) was used for the calculation of the activity 234 coefficients of aqueous species. The 2 SPNE SC/CE model was used to simulate Np sorption 235 236 to illite (Bradbury and Baeyens, 2009b). Cation exchange reactions of a metal ion with a charge +n (Mⁿ⁺) on Na-illite can be written as follows (Gaines and Thomas, 1953): 237

238
$$M^{n+} + n \text{ Na-Illite} \rightleftharpoons n \text{ Na}^+ + M \text{-Illite}$$
 (4)

The cation exchange capacity (CEC) of the illite was set to 0.225 eq/kg. Surface complexation
reactions can be written as follows:

241
$$\equiv$$
 SOH + Mⁿ⁺ + (x-1) H₂O = \equiv SOM(OH)_{(x-1})^(n-x) + x H⁺; ^SK_(x-1) (5)

where M^{n+} is a metal ion with a charge of +n and \equiv SOH is a protonated surface site.

Only the least abundant, i.e. strong, sites of the 2 SPNE SC/CE model were considered in the adsorption calculations with a site density of 2×10^{-3} mol kg⁻¹, since the weak sites are irrelevant at trace levels of Np. Bradbury and Baeyens (2009b) determined a NpO₂⁺-Na⁺ selectivity coefficient for illite as well as Np(V) surface complexation constants for illite within the 2 SPNE SC/CE model. Marsac et al. (2015a) determined surface complexation

constants for Np(IV). However, the fitted constants relied on the experimental Eh and, 248 therefore, are subject to large uncertainties. Np and Pu are neighbors in the actinide series and 249 the hydrolysis constants of Np(IV) and Pu(IV) are not significantly different (Guillaumont et 250 al., 2003). On the basis of the linear free energy relationship determined for illite by Bradbury 251 and Baeyens (2009b), the surface complexation constant for Np(IV) and Pu(IV) should not 252 significantly differ. Recently, the surface complexation constants for Pu(IV) were refined for 253 illite by Banik et al. (2016). The constants from the latter study were used here for Np(IV). 254 No constants are available for Np(V) and Np(IV) surface complexation to the more abundant 255 but weaker sites defined in the 2 SPNE SC/CE model. The complete aqueous thermodynamic 256 database, SIT coefficients and the parameters for the 2SPNE SC/CE model are given in the 257 supporting information. 258

259

260

3. RESULTS

261 3.1. Experimental batch results

Figure 1 presents experimental log R_d (R_d in L kg⁻¹) values versus pH_m for initial 262 Np(V) concentrations of 10^{-6} M (Fig. 1a) and 5×10^{-8} M (Fig. 1b), and $m_{NaCl} = 1.0$ and 3.2 m. 263 The results of Marsac et al. (2015a) in 0.1 m NaCl obtained after 1 week contact time are 264 included for $[Np(V)]_{tot} = 10^{-6} \text{ M}$ (Fig. 1a) and $3 \times 10^{-8} \text{ M}$ (Fig. 1b). No data are available for 265 $[Np(V)]_{tot} = 5 \times 10^{-8}$ M in 0.1 m NaCl. For both values of $[Np(V)]_{tot}$, no clear effect of the ionic 266 267 strength on Np sorption is observed. This is consistent with the results of experiments related to the sorption of Np(V) onto illite, shale and bentonite (MX80) in NaCl/CaCl₂ solutions up to 268 I~4.7 M (Nagasaki et al., 2016). For $[Np(V)]_{tot} = 10^{-6}$ M, log R_d slightly increases with pH_m 269 whereas, for $[Np(V)]_{tot} = 3$ or 5×10^{-8} M, the increase in log R_d with pH_m is more pronounced. 270 As observed by Marsac et al. (2015a), Np sorption to illite is stronger for low [Np(V)]_{tot}. By 271 comparison with the uptake of other cations on illite, it was demonstrated that this behavior 272

cannot be due to site saturation effects (see Figure 2 and corresponding text in Marsac et al., 273 2015a). In 0.1 m NaCl, the observations for Np were explained by the relatively low pe 274 measured in the suspension and the formation of Np(IV) at the illite surface, with Np(V)275 prevailing in the aqueous solution. By addition of 10^{-6} M Np(V) the overall concentration of 276 oxidants in the suspension increases and apparently overcompensates the relatively low redox 277 capacity of the illite suspension. As a consequence, pe increases and a smaller fraction of 278 Np(V) is reduced to Np(IV), which explains the observed reduced Np uptake at high 279 $[Np(V)]_{tot}$. Our results must be interpreted by taking into account the actual redox conditions 280 in the system. Figure 2 shows the complete datasets for various [Np(V)]_{tot} (0.05, 0.09, 0.3, 0.5, 281 282 0.9 and 3 µM) in 1.0 m (Fig. 2a,b) and 3.2 m NaCl (Fig. 2c,d). On the left side (Fig. 2a,c), uptake data are shown as log R_d versus pH_m whereas, on the right side (Fig. 2b,d), the 283 corresponding pe versus pH_m data are given. As expected, the samples with the highest 284 285 $[Np(V)]_{tot}$ show the highest pe and the lowest log R_d , whereas the highest log R_d values are measured at the lowest pe and the lowest [Np(V)]tot. Although a small amount of Fe(II) 286 287 contained in the illite was supposed to be responsible for the reduction of Np(V) to Np(IV)(Marsac et al. (2015a)), this could not be monitored experimentally. As calculated in the latter 288 study, the amount of Fe(II) required for the reduction of Np(V) in the most concentrated 289 sample (i.e. XANES sample: $[Np(V)]_{tot} = 3 \times 10^{-4} \text{ M}$; 20 g/L illite) corresponds to about 0.08 290 % of the total amount of Fe contained in the illite. Such small fractions are very likely below 291 the detection limit of available experimental techniques. 292



Figure 1. Np sorption to illite (log R_d in L kg⁻¹) after 1 week contact time in 1.0 and 3.2 m NaCl versus pH_m for [Np(V)]_{tot} = 10⁻⁶ M (a) and 5×10⁻⁸ M (b). The results obtained in 0.1 m NaCl are taken from Marsac et al. (2015a) after 1 week contact time, [Np(V)]_{tot} = 10⁻⁶ M (a) and 3×10⁻⁸ M (b).

297

298 3.2. Simulation of Np(V) and Np(IV) uptake

Calculations are made using the 2 SPNE SC/CE model in combination with SIT and the surface complexation constants for Np(V) and Pu(IV) determined by Bradbury and Baeyens (2009b) and Banik et al. (2016), respectively, for [NaCl/ClO₄] \leq 0.1 m. Np(IV) sorption is considered to be identical to that of Pu(IV). The model is used to calculate the uptake of Np(V) and Np(IV) without considering redox transitions. Results are shown in figure 2a (for 1 m NaCl) and 2c (for 3.2 m NaCl). As for Pu(IV) (Marsac et al., 2017), no significant ionic strength effect on Np(IV) sorption onto illite is observed for 0.1 < m_{NaCl} <

3.2 m. This is also consistent with the outcome of a previous study on Th(IV) sorption to 306 montmorillonite between 0.1 and 1 m NaClO₄ (Bradbury and Baeyens, 2005). Also the Np(V) 307 surface complexation to illite is found to be weakly affected by ionic strength for $0.1 < m_{NaCl}$ 308 < 3.2 m (see also Nagasaki et al., 2016). As all metal ions, Np(V) is known to sorb to illite via 309 cation exchange at low pH (pH < 6; for Np(V)) and at pH > 6 via inner sphere complexation. 310 However, due to the small charge of the $\mathrm{NpO_2^+}$ ion, and the multiple excess of $\mathrm{Na^+}$ 311 concentration, cation exchange becomes negligible and hence ionic strengths effects are not 312 pronounced. In figures 2a (for 1 m NaCl) and 2c (for 3.2 m NaCl), our experimentally 313 determined log R_d values lie between those calculated for Np(V) and Np(IV) sorption data. 314



316

317 Figure 2. On the left: experimental log R_d data measured after 1 week versus the pH_m for the various [Np]_{tot} investigated (0.05, 0.09, 0.3, 0.5, 0.9 and 3 μ M; shown as different symbols), for $m_{NaCl} = 1.0$ m (a) and 3.2 m 318 319 (c). Calculated pH-edges for sorption of Np(V) and Np(IV) (no redox transition considered), respectively, are 320 shown as lines. On the right: corresponding pH_m-pe values measured in $m_{NaCl} = 1.0$ m (b) and 3.2 m (d) for the 321 different [Np]_{tot} are shown with the same symbols as on the left. The calculated Np(V)/Np(IV) redox borderlines 322 in solution $({Np(V)/Np(IV)}_{aq}; black line)$ and at the illite surface $({Np(V)/Np(IV)}_{surf}; grey line)$ are also shown. 323 Error bars are not shown for clarity but uncertainties for log R_d (a,c) lie in the range of ± 0.3 and for pe (b,d) in 324 the range of \pm 0.8. The pH-pe value for the sample analyzed by XANES (m/V = 20 g L⁻¹; 3.2 m NaCl; [Np]_{tot} = 3×10^{-4} M) is also shown (d). 325

327 3.3. pH-pe diagrams in solution and at the illite surface

Redox borderlines between the predominance fields of Np(V) and Np(IV) in solution (noted as $\{Np(V)/Np(IV)\}_{aq}$) are shown in figures 2b (for 1.0 m NaCl) and 2d (for 3.2 m NaCl) as black lines. The borderline between the predominance fields of Np(V) and Np(IV) species adsorbed at the surface of illite (noted as $\{Np(V)/Np(IV)\}_{surf}$) is described by:

332
$$\{Np(V)/Np(IV)\}_{surf} = \{Np(V)/Np(IV)\}_{aq} + (\log R_d(Np(IV)) - \log R_d(Np(V)))$$
 (6)

and is plotted as well in figures 2b and 2d (grey lines). In eq. 6, $\{Np(V)/Np(IV)\}_{aq}$ and $\{Np(V)/Np(IV)\}_{surf}$ correspond to pe values, where $[Np(V)]_{aq}/[Np(IV)]_{aq}$ and $[Np(V)]_{surf}/[Np(IV)]_{surf}$ equal 1. They can vary with physico-chemical conditions (pH_m, m_{NaCl}, T, etc) in a similar way as log R_d(Np(IV)) and log R_d(Np(V)).

337 As shown in previous studies (Marsac et al., 2015a,b; Banik et al., 2016; Marsac et al., 2017), the stability fields of surface sorbed Np(IV) and Pu(IV) species are extended relative to 338 those for dissolved species because of their stronger sorption relative to that of other redox 339 states (i.e. Pu(III, V, VI); Np(V)) under the conditions investigated in those studies. For pH < 6, 340 at elevated ionic strength (i.e. 1.0 and 3.2 m), we find a shift of $\{Np(V)/Np(IV)\}_{surf}$ to 341 significantly higher values because Np(V) sorption to illite is decreased. Np(V) and Np(IV) 342 343 inner-sphere surface complexation to illite is only weakly affected by ionic strength so that 344 $\{Np(V)/Np(IV)\}_{surf}$ hardly varies between 0.1 and 3.2 m NaCl for pH > 6.

By comparing the experimental pH_m -pe values and the calculated values for {Np(V)/Np(IV)}_{aq} and {Np(V)/Np(IV)}_{surf}, we expect Np(V) to prevail in solution while Np(IV) predominantly exists as surface species. This qualitatively explains why the measured Np sorption data lie between those calculated for pentavalent and tetravalent Np species.

349

350 **3.4.** Spectroscopic results

Np redox speciation at the illite surface was further investigated by XANES for $pH_m =$ 351 7.9 (pe = 7.3 ± 0.8), m/V = 20 g L⁻¹, 3.2 m NaCl and 3×10^{-4} M Np (introduced as NpO₂⁺). In 352 figure 3, the XANES spectrum recorded for the Np-illite sample prepared in 3.2 m NaCl is 353 compared to the spectrum recorded in $m_{NaCl} = 0.1$ m by Marsac et al. (2015a) under similar 354 355 conditions (pH_m = 7.4; pe = 6.8 ± 0.8). As explained before, phase separation was performed to remove the solution from the sample. Therefore, the following analysis of spectra pertains 356 only to Np species associated to the solid phase. The Np-illite XANES spectra are compared 357 to those reported for Np^{4+} and NpO_2^+ aquo ions in 0.1 M HClO₄ (Gaona et al., 2012), 358

measured at the same beamline. The presence of a shoulder at 17625 eV indicates multiple 359 scattering of the outgoing Np $2p_{3/2}$ electrons along the axis of the NpO₂⁺ dioxo cation. This 360 indicates the prevalence of the +V redox state in both Np-illite samples. The "white line" (WL) 361 position as well as the inflection point of Np redox states +IV and +V are similar and cannot 362 be used as a conclusive probe for Np redox speciation. However, Np(V) and Np(IV) XANES 363 show significant differences in the intensity of the WL. The WL intensity increases from the 364 NpO_2^+ aquo ion to the Np-illite sample in 0.1 m NaCl. The sample prepared in 3.2 m NaCl 365 shows intermediate intensity. As discussed in more detail in Marsac et al. (2015a) and 366 Denecke et al. (2005), the WL of Np(V) sorbed to a mineral surface will be less intense than 367 368 that of the Np(V) aquo-ion. Therefore, the present results indicate the presence of Np(IV). By fitting the 0.1 m NaCl Np-illite XANES spectrum with a linear combination of Np(V) and 369 Np(IV) reference samples (i.e. aquo-ions), Marsac et al. (2015a) found that 14±1% (error 370 371 provided by the Demeter Athena software; Ravel and Newville, 2005) of the adsorbed Np was Np(IV) (i.e. Np(V)-to-Np(IV) concentration ratio at the illite surface $[Np(V)]_{surf}/[Np(IV)]_{surf}$ 372 6.1). Repeating the same exercise for the 3.2 m NaCl sample yields the presence of at least 373 8±1% of the Np existing as tetravalent species at the illite surface (i.e. [Np(V)]_{surf}/[Np(IV)]_{surf} 374 = 11.5). Although this value is small, it must be considered as a minimum value since the 375 376 aquo-ions exhibit the most intense WLs. The same exercise with either a Np(V) or a Np(IV)reference, where Np is in a more condensed state (i.e. with a less intense WL), would lead to a 377 larger fraction of Np(IV) (Marsac et al., 2015a). XANES analysis, therefore, yields a similar 378 379 result compared to what has been obtained for the experiment at low ionic strength and provides a clear indication for partial reduction of Np(V), also at elevated ionic strength. 380 Using experimental pe and spectroscopically determined [Np(V)]_{surf}/[Np(IV)]_{surf} values, we 381 find that $\{Np(V)/Np(IV)\}_{surf} = 6.0$ (for $m_{NaCl} = 0.1$) and 6.2 (for $m_{NaCl} = 3.2$). These values 382 compare very well with $\{Np(V)/Np(IV)\}_{surf} = 5.9$ (for both 0.1 and 3.2 m NaCl), as 383 determined by modeling (see e.g. the grey lines in Fig. 2b,d). This further confirms that the 384

Np(V)-Np(IV) redox transition at the illite surface is not substantially affected by ionicstrength.



387

Figure 3. (a) Np L₃-XANES measured for Np-illite samples ($[Np]_{tot} = 3 \times 10^{-4} \text{ M}$, m/V = 20 g L⁻¹) prepared in 0.1 m NaCl (pH_m = 7.4; pe = 6.8; Marsac et al., 2015a) and 3.2 m NaCl (pH_m = 7.9; pe = 7.3; present study). Reference XANES of aqueous Np(IV) and Np(V) in 1 M HClO₄ from Gaona et al. (2012) are shown for comparison. Arrows highlight the partial Np(V) reduction to Np(IV) in the present study. The area between 17610 and 17650 eV is enlarged in (b). Results of the linear combination fit are shown in the supporting information (Figure S1).

394

4. DISCUSSION

397 4.1. Effects of pH_m , pe and m_{NaCl} on Np sorption to illite

Because Np uptake on illite is sensitive to pH_m and pe, it appears difficult to study in 398 detail the effect of a third parameter, namely m_{NaCl} . Figure 4 shows data for experimentally 399 determined log R_d plotted versus pe for $m_{NaCl} = 0.1$, 1.0 and 3.2. It is obvious that there is a 400 strong correlation between R_d and pe showing that increasing the pe from 2 to 6 reduces R_d by 401 \sim 3 orders of magnitude. In so far "pe edges" (log R_d plotted vs. pe) are more informative than 402 the "pH edges" in Figures 2a and 2c (log R_d vs. pH). It appears that pe is the more relevant 403 parameter determining Np retention under given conditions. A few experimental data do not 404 405 fit into the trend and have been excluded from figure 4: (i) sorption data obtained at $pH_m < 6$ 406 show a significant pH dependence because the Np(V)/Np(IV) borderlines in this range clearly evolve with pH (see Fig. 2b,d); (ii) at quite high pe values, where Np(V) prevails in solution 407 408 and at the illite surface, the influence of pH on Np sorption is of course dominant (see the pHedges of Np(V) in Fig. 2a,c). The exclusion of data points amount to ~20% of the complete 409 data set, and make sure that, for the selected data, pH has no significant effect. 410

Figure 4 seems to suggest that Np sorption to illite is stronger in 0.1 than in 1.0 and 411 3.2 m NaCl at least for pe < 4.5, an issue that does not occur on Figure 1. Recent experimental 412 413 investigations did not show significant influence of ionic strength neither on tetravalent actinide ion (I = 0.1 - 3.2 m; Bradbury and Baeyens, 2005; Marsac et al., 2017; results of both 414 studies on Th(IV) and Pu(IV) sorption are shown in the supporting information) nor on Np(V) 415 416 sorption to illite via inner-sphere surface complexation (I = 0.1 to 4.6 M; Nagasaki et al., 2016). Moreover, our XANES results suggest no ionic strength effect on Np(V)-Np(IV) redox 417 transition at the illite surface. In fact, the interpretation of the Np-illite results is strongly 418 affected by the experimental pe. Unlike Np, Th(IV) is not a redox sensitive metal ion and its 419 420 sorption on illite is independent of the ionic strength within the experimental uncertainties. The same independency of the ionic strength is observed for Pu(IV) illite uptake. From these results it would appear difficult to assume, that the ionic strength has no significant influence on the Np(IV) uptake, although trends in average values could suggest such effects. Therefore, it is more probable that the variations of Np-illite uptake data are due to the uncertainties of experimental pe measurements.

To highlight the consistency between Np-illite and Pu-illite sorption data in 0.1, 1.0 426 427 and 3.2 m NaCl (Banik et al., 2016; Marsac et al., 2017; present study), simulations are made with the 2 SPNE SC/CE model for $pH_m = 7$ and the results are plotted in Figure 4. This pH_m 428 value is arbitrarily chosen because insignificant pH_m effect is expected for the selected dataset, 429 430 as explained above. The surface complexation constant for Np(IV) given by Marsac et al. (2015a) (Table S3) describes sorption data determined for low ionic strength accurately, but 431 overestimates sorption when comparing with experimental data determined in this study. In 432 433 this case, $\log R_d(Np(IV)) = 6.1$ (i.e. $\log R_d$ values extrapolated to pe values where all Np(V) in the system is reduced to Np(IV) in solution and at the illite surface). Adjusting model 434 435 parameters to fit our data in 1.0 and 3.2 m NaCl using PhreePlot leads to $\log R_d(Np(IV)) = 4.7$ 436 (corresponding surface complexation constants are given in Table S3). In average, the model would give log $R_d(Np(IV)) = 5.4 \pm 0.7$ (which is also determined by Phreeplot when using the 437 complete dataset in 0.1, 1.0 and 3.2 m NaCl; Table S3). This average value is in excellent 438 agreement with log $R_d(Pu(IV)) = 5.3 \pm 0.3 L \text{ kg}^{-1}$ determined by Marsac et al. (2017) for 0.1 < 439 m_{NaCl} < 3.2 m. Figure 4 shows that calculated log R_d evolves linearly with pe for 440 intermediate pe values (for $\{Np(V)/Np(IV)\}_{aq} < pe < \{Np(V)/Np(IV)\}_{surf}$, all the selected data 441 are within this range in Figure 4). Therefore, an error of ± 0.8 on pe leads to an error of ± 0.8 442 on the calculated log R_d value, which compares well with the error on determined log 443 R_d(Np(IV)) using Np-illite data at 0.1, 1 and 3.2 m NaCl (±0.7). These results suggest that it 444 is reasonable to impose surface complexation constants for Np(IV) equal to those of Pu(IV) 445

and to attribute the trend of average Np-illite uptake with ionic strength effects touncertainties in pe.

448



449

Figure 4. Np-illite sorption datasets from Marsac et al. (2015a; $m_{NaCl} = 0.1$ m) and the present study ($m_{NaCl} =$ 1.0 and 3.2 m) are plotted as log R_d versus pe. Lines are calculations made with the 2 SPNE SC/CE model for pH_m = 7 using the surface complexation constants determined by Marsac et al. for Np(IV) (2015a; top dotted line), Banik et al. for Pu(IV) (2016; bold line) or fitting these constants to the 1.0 and 3.2 m NaCl data (bottom dotted curve). The corresponding log R_d(Np(IV)) is given on the y-axis.

455

456 4.2. Modeling An(III,IV,V) sorption to illite as a function of pH_m , pe and m_{NaCl}

Values of log R_d are calculated for all the presently investigated conditions (including 457 the conditions of Marsac et al., 2015a) using the experimental pH_m and pe values. The model 458 results are plotted versus the experimental log R_d in Figure 5. The results of Banik et al. (2016; 459 0.1 m NaCl) and Marsac et al. (2017; 1.0 and 3.2 m NaCl), involving the Pu(IV)/Pu(III) 460 couple, are also included. The 1:1 line shows the good agreement between model and 461 462 experiment. Deviations of $\pm 1.1 \log R_d$ unit are shown as dotted lines. This error for the calculated log R_d accounts for both the error on pe (±0.8) and on experimental log R_d (±0.3). 463 For a wide range of pH_m , $[Np/Pu]_{tot}$, m_{NaCl} and pe investigated in these different studies (i.e. 464 179 experimental log R_d values) the discrepancies between model and the experimental 465

results rarely exceed 1.1 log R_d unit. The mean value of the deviation between experimental and calculated log R_d values is 0.42. It is important to note that sorption data are highly sensitive to the redox conditions and pe measurements are associated with considerable uncertainties. We nevertheless conclude that the combination of (i) the 2 SPNE SC/CE model, (ii) SIT and (iii) the approach developed by Marsac et al. (2015a,b), which accounts for redox equilibria and the stability of surface species, allows to reasonably model actinide(III,IV,V) uptake by illite as a function of pH_m, pe and m_{NaCl} up to highly saline conditions.

473



474

Figure 5. Modelled versus experimental log R_d values for Np (data from this study and Marsac et al., 2015a) and Pu in $0.1 < m_{NaCl} < 3.2$ m (data from Banik et al., 2016; Marsac et al., 2017). The 1:1 (solid) line is associated with an uncertainty of \pm 1.1 log R_d units (dashed lines).

- 478
- 479

4. CONCLUSIONS

In agreement with previous studies performed at low ionic strength (Marsac et al., 2015a), Np sorption increases with decreasing redox potential also at high m_{NaCl} (up to 3.2 m). Using a combined batch sorption, spectroscopic and geochemical modeling approach, the strong interaction between Np and illite is attributed to the partial reduction of Np(V) to 484 Np(IV) at the illite surface driven by the higher thermodynamic stability of Np(IV) surface 485 complexes as compared to Np(V) surface species. By the combination of the 2 SPNE SC/CE 486 model with SIT, it is possible to reproduce experimentally obtained overall Np uptake data for 487 illite as a function of pH_m, pe and m_{NaCl} using a single set of parameters. The ionic strength 488 has a relatively small impact on sorption of both Np(V) and Np(IV). Main parameters driving 489 the sorption behavior of redox sensitive actinides are pH_m and pe.

Notably the role of the redox potential on the uptake of redox sensitive actinides on 490 minerals is to be considered when applying laboratory derived conditional R_d values for 491 performance safety assessment purposes. The applicability of such data is only justified, if 492 redox conditions in laboratory experiments and those expected in the real system are identical. 493 This is not always the case as extraction of natural materials from deep environments, 494 transport to the laboratory and experimental handling are technically challenging and redox 495 496 conditions may change due to e.g. access of oxygen. However, if the pH and redox potentials of in situ and laboratory conditions are known, mechanistic models as proposed here, are 497 498 capable of assessing the overall uptake of redox sensitive actinides by natural adsorbents with 499 reasonable accuracy. In agreement with recent studies, the use of non-electrostatic models 500 appears suitable for the prediction of metal ion sorption to various types of surfaces even at elevated ionic strength, e.g. marine microalgae (Zoll and Schjif, 2012), bacteria (Ams et al., 501 502 2013), illite (Schnurr et al., 2015; Marsac et al., 2017) or smectite (Schnurr et al., 2015). Accordingly, the 2 SPNE SC/CE model is a reliable predictive tool for performance safety 503 504 assessment for clays even under elevated saline conditions.

505

506 ACKNOWLEDGEMENTS

507 This work was financed by the Federal Ministry of Economic Affairs and Energy 508 (Germany) under contracts No. 02E10206 and 02E10961. The research leading to these results has received funding from the European Union's European Atomic Energy
Community's (Euratom) Seventh Framework Program FP7/2007-2011 under grant agreement
n° 249624 (CATCLAY project).

512

513 Supporting Information Available

514 Additional material is provided in the supporting information section.

516 **REFERENCES**

- Altmaier M., Gaona X. and Fanghänel T. (2013) Recent advances in aqueous actinide
 chemistry and thermodynamics. *Chem. Rev.* 113, 901-943.
- Altmaier M., Gaona X., Fellhauer D. and Buckau G. (2010) Intercomparison of redox
 determination methods on designed and near-neutral aqueous systems. KIT-SR 7572,
 Karlsruhe Institute of Technology, Karlsruhe.
- Altmaier M., Metz V., Neck V., Müller R. and Fanghänel Th. (2003) Solid-liquid equilibria of
 Mg(OH)_{2(cr)} and Mg₂(OH)₃Cl•4H₂O_(cr) in the system Mg-Na-H-OH-Cl-H₂O at 25°C.
 Geochim. Cosmochim. Acta 67, 3595–3601.
- Ams D. A., Swanson J. S., Szymanowski J. E. S., Fein J. B., Richmann M. and Reed D. T.
 (2013) The effect of high ionic strength on neptunium (V) adsorption to a halophilic
 bacterium. *Geochim. Cosmochim. Acta* 110, 45–57.
- Andra, Dossier 2005 Synthèse Évaluation de la faisabilité du stockage géologique en
 formation argileuse. 2005, Agence nationale pour la gestion des déchets radioactifs
 (Andra): Châtenay-Malabry, France. p. 241.
- Banik N. L., Marsac R., Lützenkirchen J., Diascorn A., Bender K. Marquardt C. M. and
 Geckeis H. (2016) Sorption and redox speciation of plutonium on illite. *Environ. Sci. Technol.* 50, 2092–2098.
- Bradbury M. H. and Baeyens B. (2005) Modelling the sorption of Mn(II), Co(II), Ni(II),
 Zn(II), Cd(II), Eu(III), Am(III), Sn(IV), Th(IV), Np(V) and U(VI) on montmorillonite:
 Linear free energy relationships and estimates of surface binding constants for some
 selected heavy metals and actinides. *Geochim. Cosmochim. Acta* 69, 875–892.
- Bradbury M. H. and Baeyens B. (2009a) Sorption modelling on illite Part I: Titration
 measurements and the sorption of Ni, Co, Eu and Sn. *Geochim. Cosmochim. Acta.* 73,
 990–1003.
- Bradbury M. H. and Baeyens B. (2009b) Sorption modeling on illite Part II: actinide sorption
 and linear free energy relationships. *Geochim. Cosmochim. Acta.* 73, 1004–1013.
- Bradbury M. H. and Baeyens B. (2011) Predictive sorption modelling of Ni(II), Co(II),
 Eu(IIII), Th(IV) and U(VI) on MX-80 bentonite and Opalinus Clay: A "bottom-up"
 approach, *Appl. Clay Sci.* 52, 27–33.
- 546 Brendebach B., Banik N. L., Marquardt C. M., Rothe J., Denecke M. and Geckeis H. (2009)
 547 X- ray absorption spectroscopic study of trivalent and tetravalent actinides in solution at
 548 varying pH value. *Radiochim. Acta* 97, 701–708.
- 549 Brewitz W. (1980) Zusammenfassender Zwischenbericht, GSF T 114.
- Chen Z., Montavona G., Guo Z., Wang X., Razafindratsima S., Robinet J. C., Landesmana C.
 (2014) Approaches to surface complexation modeling of Ni(II) on Callovo-Oxfordian
 clayrock. *Appl. Clay Sci.* 101, 369–380.
- Ciavatta L. (1980). The specific interaction theory in the evaluating ionic equilibria. *Ann. Chim.* 70, 551–562.

- Denecke M. A., Dardenne K. and Marquardt C. M. (2005) Np(IV)/Np(V) valence
 determinations from Np L3 edge XANES/EXAFS. *Talanta* 65, 1008–1014.
- Fritz P. and Frape S. K. (1982). Saline groundwaters in the Canadian shield a first overview.
 Chem. Geol. 36, 179–190.
- Fröhlich D. R., Amayri S., Drebert J. and Reich T. (2011) Sorption of neptunium(V) on
 Opalinus Clay under aerobic/anaerobic conditions. *Radiochim. Acta* 99, 71–77.
- Fröhlich D. R., Amayri S., Drebert J., Grolimund D., Huth J., Kaplan U., Krause J. and Reich
 T. (2012) Speciation of Np(V) uptake by Opalinus Clay using synchrotron microbeam
 techniques. *Anal. Bioanal. Chem.* 404, 2151–2162.
- Gaines G. I. and Thomas H. C. (1953) Adsorption studies on clay minerals. II. A formulation
 of the thermodynamics of exchange adsorption. *J. Phys. Chem.* 21, 714-718.
- Gaona X., Tits J., Dardenne K., Liu X., Rothe J., Denecke M. A., Wieland E. and Altmaier M.
 (2012) Spectroscopic investigations of Np(V/VI) redox speciation in hyperalkaline
 TMA-(OH, Cl) solutions. *Radiochim. Acta* 100, 759–770.
- Gaucher E., Robelin C., Matray J. M., Négrel G., Gros Y., Heitz J. F., Vinsot A., Rebours H.,
 Cassagnabère A. and Bouchet A. (2004) ANDRA underground research laboratory:
 interpretation of the mineralogical and geochemical data acquired in the Callovian–
 Oxfordian formation by investigative drilling. *Phys. Chem. Earth* 29, 55–77.
- Geckeis H., Lützenkirchen J., Polly R., Rabung T. and Schmidt M. (2013) Mineral-water
 interface reactions of actinides. *Chem. Rev.* 13 (2), 1016–1062.
- Graser C.-H., Banik N. L., Bender K. A., Marquardt C. M., Marsac R., Monytoya, V. and
 Geckeis H. (2015) Sensitive Redox speciation of elements relevant to nuclear waste
 disposal by Capillary Electrophoresis hyphenated to inductively coupled plasma sector
 field mass spectrometry. *Anal. Chem.* 87, 9786–94.
- Guillaumont R., Fanghänel Th., Fuger J., Grenthe I., Neck V., Palmer D. A. and Rand M. H.
 (2003) Update on the chemical thermodynamics of Uranium, Neptunium, Plutonium,
 Americium and Technetium, Mompean, F.J., Domenech-Orti, C., Ben-Said, K.,
 OECD/NEA Data Bank, Eds., vol. 5 of Chemical Thermodynamics, Elsevier,
 Amsterdam.
- Kim J. I. (1986) Chemical behaviour of transuranic elements in natural aquatic systems. In:
 Handbook on the Physics and Chemistry of the Actinides. (Freeman, A. J., ed.) Elsevier
 Science Publishers, B. V., Amsterdam p. 413.
- 587 Kinniburgh D. G. and Cooper D. M. (2009) PhreePlot: Creating graphical output with
 588 PHREEQC. http://www.phreeplot.org>.
- Kirsch R., Fellhauer D., Altmaier M., Neck V., Rossberg A., Fanghänel Th., Charlet L. and
 Scheinost A. C. (2011) Oxidation state and local structure of plutonium reacted with
 magnetite, mackinawite, and chukanovite. *Environ. Sci. Technol.* 45, 7267–7274
- Lauber M., Baeyens B. and Bradbury M. H. (2000) Physico-chemical characterisation and
 sorption measurements of Cs, Sr, Ni, Eu, Th, Sn and Se on Opalinus Clay from Mont
 Terri. PSI Technical Report 00-10, Paul Scherrer Institut, Villigen/Switzerland

- Lázár K. and Máthé Z. (2012) Claystone as a potential hostr for nuclear waste storage, in
 "Clay minerals in nature Their characterization, modification and application.", Marta
 Valaskova (Ed.), InTech, ISBN 978-953-51-0738-5.
- Marsac R., Banik N. L., Diascorn A., Kupcik T., Lützenkirchen J., Marquardt C. M., Schäfer
 T., Schild D., Rothe J., Dardenne K. and Geckeis H. (2015a) Neptunium redox
 speciation at the illite surface. *Geochim. Cosmochim. Acta* 152, 39–51.
- Marsac R., Banik N. L., Lützenkirchen J., Buda R. A., Kratz J. V. and Marquardt C. M.
 (2015b) Modeling plutonium sorption to kaolinite: accounting for redox equilibria and the stability of surface species. *Chem. Geol.* 400, 1–10.
- Marsac R., Banik N. L., Lützenkirchen J., Diascorn A., Bender K., Marquardt C. M. and
 Geckeis H. (2017) Sorption and redox speciation of plutonium on illite under highly
 saline conditions. J. Coll. Int. Sci. 485, 59–64.
- Nagasaki S., Saito T. and Yang T. T. (2016) Sorption behavior of Np(V) on illite, shale and
 MX-80 in high ionic strength solutions. *J. Radioanal. Nucl. Chem.*, **308**, 143–153.
- Nagra (2002) Project opalinus clay: safety report. Demonstration of disposal feasibility
 (Entsorgungsnachweis) for spent fuel, vitrified high-level waste and long-lived
 intermediate-level waste. Nagra Technical Report NTB 02-05, Nagra, Wettingen,
 Switzerland.
- OECD/NEA (2008) Radioactive waste management committee: "Collective statement on moving forward to geological disposal of radioactive waste", ISBN 978-92-64-99057 9.
- Ondraf (2001) SAFIR 2: safety assessment and feasibility interim report 2. NIROND-2001-06
 E. Ondraf, Brussels.
- Parkhurst D. L. and Appelo C. A. J. (1999) User's guide to PHREEQC (Version 2) a
 computer program for speciation, batch reaction, one-dimensional transport and inverse
 geochemical calculation. Water-resources Investigation Report 99-4259, USGS, Denver,
 Colorado, p. 312.
- Pearson F. J., Tournassat C. and Gaucher E. C. (2011) Biogeochemical processes in a clay
 formation in situ experiment: Part E Equilibrium controls on chemistry of pore water
 from the Opalinus Clay, Mont Terri Underground Research Laboratory, Switzerland. *Appl. Geochem.* 26, 990–1008.
- 626 Pitzer K. S. (1991). Activity Coefficients in Electrolyte Solutions, CRC Press, Boca Raton.
- Ravel B. and Newville M. (2005) ATHENA and ARTEMIS: data analysis for X-ray absorption spectroscopy using IFEFFIT. J. Synchrotron Radiat. 12, 537–541.
- Rothe J., Butorin S., Dardenne K., Denecke M. A., Kienzler B., Löble M., Metz V., Seibert A.,
 Steppert M., Vitova T., Walther C. and Geckeis H. (2012) The INE-Beamline for
 actinide science at ANKA. *Rev. Sci. Instrum.* 83, 043105.
- Rothe J., Denecke M. A., Dardenne K. and Fanghänel T. (2006) The INE-beamline for
 actinide research at ANKA. *Radiochim. Acta.* 94, 691–696.

- Schijf J. and Ebling A. M. (2010) Investigation of the ionic strength dependence of Ulva
 lactuca acid functional group pK_{as} by manual alkalimetric titrations. *Environ. Sci. Technol.* 44, 1644–1649
- Schnurr A., Marsac R., Kupcik T., Rabung T., Lützenkirchen J. and Geckeis H. (2015)
 Sorption of Cm(III) and Eu(III) onto clay minerals under saline conditions: Batch adsorption, Laser-fluorescence spectroscopy and modeling. *Geochim. Cosmochim. Acta* 151, 192–202.
- Sjoblom R. and Hindman J.C. (1951) Spectrophotometry of neptunium in perchloric acid
 solutions. J. Am. Chem. Soc. 73(4), 1744–1751.
- Vilks P. (2011) Sorption of selected radionuclides on sedimentary rocks in saline conditions –
 Literature review. NWMO TR-2011-12, Atomic Energy of Canada Limited, Toronto,
 Ontario, Canada.
- Zoll A. M. and Schjif J. (2012) A surface complexation model of YREE sorption on *Ulva lactuca* in 0.05–5.0 M NaCl solutions. *Geochim. Cosmochim. Acta* 97, 183–199.

649 **Figure caption**

Figure 1. Np sorption to illite (log R_d in L kg⁻¹) after 1 week contact time in 1.0 and 3.2 m NaCl versus pH_m for $[Np(V)]_{tot} = 10^{-6}$ M (a) and 5×10^{-8} M (b). The results obtained in 0.1 m NaCl are taken from Marsac et al. (2015a) after 1 week contact time, $[Np(V)]_{tot} = 10^{-6}$ M (a) and 3×10^{-8} M (b).

Figure 2. On the left: experimental $\log R_d$ data measured after 1 week versus the pH_m for the 654 various $[Np]_{tot}$ investigated (0.05, 0.09, 0.3, 0.5, 0.9 and 3 μ M; shown as different symbols), 655 for $m_{NaCl} = 1.0$ m (a) and 3.2 m (c). Calculated pH-edges for sorption of Np(V) and Np(IV) 656 (no redox transition considered), respectively, are shown as lines. On the right: corresponding 657 pH_m-pe values measured in $m_{NaCl} = 1.0$ m (b) and 3.2 m (d) for the different [Np]_{tot} are 658 shown with the same symbols as on the left. The calculated Np(V)/Np(IV) redox borderlines 659 in solution $({Np(V)/Np(IV)}_{aq}; black line)$ and at the illite surface $({Np(V)/Np(IV)}_{surf}; grey$ 660 line) are also shown. Error bars are not shown for clarity but uncertainties for log R_d (a,c) lie 661 in the range of ± 0.3 and for pe (b,d) in the range of ± 0.8 . The pH-pe value for the sample 662 analyzed by XANES (m/V = 20 g L⁻¹; 3.2 m NaCl; $[Np]_{tot} = 3 \times 10^{-4} M$) is also shown (d). 663

Figure 3. (a) Np L₃-XANES measured for Np-illite samples ($[Np]_{tot} = 3 \times 10^{-4}$ M, m/V = 20 g L⁻¹) prepared in 0.1 m NaCl (pH_m = 7.4; pe = 6.8; Marsac et al., 2015a) and 3.2 m NaCl (pH_m = 7.9; pe = 7.3; present study). Reference XANES of aqueous Np(IV) and Np(V) in 1 M HClO₄ from Gaona et al. (2012) are shown for comparison. Arrows highlight the partial Np(V) reduction to Np(IV) in the present study. The area between 17610 and 17650 eV is enlarged in (b). Results of the linear combination fit are shown in the supporting information (Figure S1).

Figure 4. Np-illite sorption datasets from Marsac et al. (2015a; $m_{NaCl} = 0.1$ m) and the present study ($m_{NaCl} = 1.0$ and 3.2 m) are plotted as log R_d versus pe. Lines are calculations made with the 2 SPNE SC/CE model for pH_m = 7 using the surface complexation constants determined by Marsac et al. for Np(IV) (2015a; top dotted line), Banik et al. for Pu(IV) (2016; bold line) or fitting these constants to the 1.0 and 3.2 m NaCl data (bottom dotted curve). The corresponding log R_d(Np(IV)) is given on the y-axis.

Figure 5. Modelled versus experimental log R_d values for Np (data from this study and Marsac et al., 2015a) and Pu in $0.1 < m_{NaCl} < 3.2$ m (data from Banik et al., 2016; Marsac et al., 2017). The 1:1 (solid) line is associated with an uncertainty of \pm 1.1 log R_d units (dashed lines).







Figure 3



Figure 5