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Organic chemistry in a CO₂ rich early Earth atmosphere

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Abstract

The emergence of life on the Earth has required a prior organic chemistry leading to the formation of prebiotic molecules. The origin and the evolution of the organic matter on the early Earth is not yet firmly understood. Several hypotheses, possibly complementary, are considered. They can be divided in two categories: endogenous and exogenous sources. In this work we investigate the contribution of a specific endogenous source: the organic chemistry occurring in the ionosphere of the early Earth where the significant VUV contribution of the young Sun involved an efficient formation of reactive species. We address the issue whether this chemistry can lead to the formation of complex organic compounds with CO₂ as only source of carbon in an early atmosphere made of N₂, CO₂ and H₂, by mimicking experimentally this type of chemistry using a low pressure plasma reactor. By analyzing the gaseous phase composition, we strictly identified the formation of H₂O, NH₃, N₂O and C₂N₂. The formation of a solid organic phase is also observed, confirming the possibility to trigger organic chemistry in the upper atmosphere of the early Earth. The identification of Nitrogen-bearing chemical functions in the solid highlights the possibility for an efficient ionospheric chemistry to provide prebiotic material on the early Earth.

Key words: Atmospheres, chemistry; Organic chemistry; Prebiotic chemistry; Earth

1. Introduction

Life is supposed to have appeared on Earth before 3.5 Ga during the Archean and maybe before the Late Heavy Bombardment (LHB) during the Hadean (Nisbet and Sleep, 2001). The presence of liquid water and organic matter is now widely accepted as conditions for the apparition of life (Cottin et al., 2015). Understanding the origins of these organic molecules, which are implicated in the apparition of life is of prime interest for astrobiology. Different origins are proposed for the organic matter on the early Earth at this period: exogenous delivery by meteorites and comets, or endogenous formation in hydrothermal vents or in the primitive atmosphere (Chyba and Sagan, 1992).

Reduced atmospheres are known to be an important source of organic matter as pointed out by the observation of Titan the largest satellite of Saturn (Tomasko and West, 2010; Waite et al., 2010) or by experiments realized by Miller about the reactivity of the early Earth atmosphere (Miller, 1953). However, if the question of the degree of oxidation of the early Earth atmosphere and so its composition is not completely solved, there is evidence that the upper mantle of the Earth was at the present redox state since 3.9 Ga (Delano, 2001) and probably 4.4 Ga (Trail, 2011) resulting in a relatively oxidant primitive atmosphere dominated by molecular nitrogen N_2 and carbon dioxide CO_2 (Kasting, 1993). The composition of the early Earth's atmosphere is an important parameter to consider for the production of organic matter. Different experiments highlighting the fact that chemistry in oxidant atmospheres produces less organic compounds than in reduced atmospheres (Schlesinger and Miller, 1983a, b).

Recent experimental works showed that organic aerosols can be produced in an atmosphere dominated by CO_2 with a minor amount of methane CH_4 (Trainer et al., 2004; Trainer et al., 2006). But it remains uncertain whether CH_4 was present in the primitive atmosphere of the

Earth. Indeed, today 70% of the methane emissions have biogenic sources and for an important part from ecosystems, which do not exist during the Hadean and the Archean eons (Denman et al., 2007). The two principal abiotic sources of methane considered for the early Earth are delivery by meteoritic and cometary impacts and production by serpentinization process (Kasting, 2005). These sources could be responsible for a methane amount of only few ppm_v in the primitive atmosphere of the Earth (Emmanuel and Ague, 2007; Feulner, 2012; Guzmán-Marmolejo et al., 2013; Kasting, 2005), which is much lower than the amount considered in experimental simulations (Trainer et al., 2006) or climate model (Charnay et al., 2013). For these reasons, we have chosen to explore a potential organic chemistry in an early Earth's atmosphere with only CO₂ as source of carbon.

Moreover, if CH₄ concentration should be low in the early Earth's atmosphere, molecular hydrogen H₂ could be present in a larger amount. Indeed, assuming that the loss of hydrogen to space was limited by diffusion, a mixing ratio of 10⁻³ has been established for the early Earth's atmosphere (Kasting, 1993). But, assuming a lower Jeans escape than today, Tian, 2005 has proposed that H₂ mixing ratio could be 100 times higher, reaching 10% and up to 30%. According to the authors, this lower efficiency of the Jeans escape for H₂ would be explained by a lower temperature of the exobase at this period. Indeed, in one hand the higher CO₂ concentration at this period would result to a higher radiations backed to space and in the other hand, the lower oxygen concentration in the atmosphere (O₂) would result to a lower heating of the exobase by the atomic oxygen UV absorbance (Martin et al., 2006; Tian et al., 2005). This new calculation has been discussed since, notably because of the limited amount of CO₂ present in the atmosphere to cool the exobase (Catling, 2006; Tian et al., 2006). A more recent model has determined a hydrogen mixing ratio of 1% at the homopause. It results that the exact concentration of H₂ during the Hadean and the Archean remains difficult to

determine but could be important. For this reason, the influence of H₂ on the primitive atmospheric reactivity needs to be studied.

We present in this paper an experimental study of the organic growth occurring in an early Earth's atmosphere made of N₂, CO₂ and H₂. A first study performed in similar conditions highlighted an important formation of water vapor in the stratosphere and ionosphere of the early Earth (Fleury et al., 2015). We are now interested in investigating the concomitant formation of organic molecules, as both gaseous and solid products. The composition of the Earth atmosphere has evolved over geological time, notably CO₂ whose concentration during the Hadean and the Archean is not precisely known (Feulner, 2012). For this reason, we also take into account the influence of the CO₂ initial amount on the atmospheric chemistry of the early Earth simulated in our experiments.

2. Experimental setup and analytical protocols

2.1. Experimental simulation

We used here the PAMPRE experimental setup, which has been described in details in previous publication (Szopa et al., 2006). It is a Radio Frequency Capacitively Coupled Plasma (RF CCP) at low pressure. In this experiment, a discharge is generated between a polarized electrode and a cylindrical grid grounded electrode confining the plasma. Before each experiment, the reactor is heated and pumped down to 2×10^{-6} mbar. The gaseous flow is then injected continuously and pumped through a rotary vane vacuum pump. Three gas bottles are used in the experiments to generate the reactive gaseous mixture: one with high-purity of N₂ (99.999%), one containing a N₂-H₂ mixture with 5% of H₂ and one with CO₂ (> 99.995%).

The generator RF power is set to 30 W and the total gas flow rate to 55 sccm resulting in a 0.9 mbar pressure in the reactor. The modeling of the used plasma has been done in pure N₂ (Alves et al., 2012). The electrons energy distribution function (EEDF) presents a maximum at 2 eV and a relatively populated tail for electron energy above 4 eV and up to 14 eV mimicking the young solar spectrum, which presents a higher UV and X-ray emission flux and a lower visible and infrared emission flux than today (Claire et al., 2012). The electrons present in our plasma have sufficient energy to dissociate and ionize N₂ and activate nitrogen chemistry as well as CO₂ and H₂. The PAMPRE experiment can be so used to simulate the reactivity of the upper layers (above the troposphere) of the early Earth, where N₂, CO₂ and H₂ can be dissociated and ionized.

The hydrogen mixing ratio is kept constant at 4% for all experiments. This concentration is chosen in agreement with recent modeling studies of the H₂ mixing ratio, giving an upper limit about 1% for the Archean atmosphere (Kuramoto et al., 2013) and up to 30% (Tian et al., 2005). The gas flow is adjusted among N₂ and CO₂ from an experiment to another to introduce CO₂ at different mixing ratios: 1%, 5% and 10%.

2.2. Cryogenic trapping

In order to detect and identify gas species produced in low quantity, we have trapped the volatile species and accumulated them by cooling the plasma box. For that, a part of the experiments presented below is performed with a continuous injection of liquid nitrogen (L_{N2}) inside the stainless-steel block supporting the grounded electrode. The plasma box is cooled by thermal conduction. The temperature is fixed at 173 K to prevent N₂, H₂ and CO₂ condensation. However, the trapping of CO₂ has been observed in the results presented in the section 3.4. The trapping of the reactants is discussed in the same section and has been taken into account for the interpretation of the results. Products are accumulated during 4 or 8 hours

of plasma duration. After plasma turning off, the reactor is pumped to eliminate N₂, H₂ and CO₂. The reactor is isolated and the plasma box is warmed up to room temperature. Trapped gases are released and analyzed by mass spectrometry, infrared spectroscopy and by Gas Chromatography coupled to Mass Spectrometry (GC-MS).

2.3. Mass spectrometry on gaseous phase

Measurements of the gas phase trapped during a part of the experiments, as described in section 2.1, are achieved with a Pfeiffer QME 200 quadrupole mass spectrometer (QMS). In the spectrometer, neutral species are ionized by electron impact at 70 eV. Gases are transferred to the spectrometer through a capillary tube, which is long enough to keep the pressure below 10⁻⁵ mbar in the spectrometer when the pressure inside the reactor is 0.9 mbar. Its resolution is 100 at m/z 100 and it covers 1-100 u mass range. Gas trapped are analyzed with a long scanning acquisition of 1 s for each mass between m/z 1 and 60.

2.4. Infrared spectroscopy on gaseous phase

Trapped molecules are also analyzed with a Thermo Scientific Nicolet 6700 Fourier Transform Infrared (FTIR) spectrometer. A schema of the FTIR setup on the PAMPRE reactor, seen from above, is presented in Figure 1. The infrared beam is emitted by the FTIR source and passes through the reactor via two KBr windows. Then, the beam is collected by a Mercury Cadmium Telluride (MCT) detector cooled by liquid nitrogen.

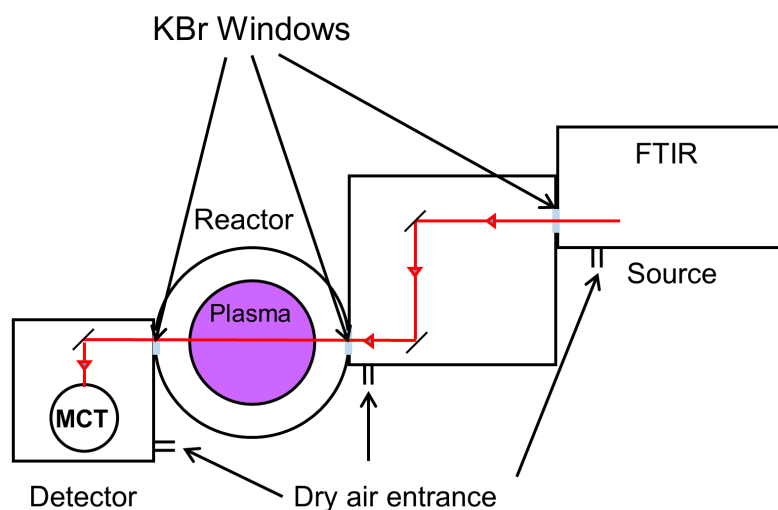


Figure 1: Schema of the FTIR setup on the PAMPRE reactor.

In the results presented below, IR spectra are recorded in the $650\text{-}4500\text{ cm}^{-1}$ range with a resolution of 1 cm^{-1} after a co-addition of 500 scans. With only one passage of the beam through the reactor, the corresponding path length is $50.8 \pm 0.2\text{ cm}$.

2.5. GC-MS analysis of the gaseous phase

To analyze the gaseous phase *ex situ* by GC-MS, we transfer the trapped gases in the reactor into another external cold trap. It is a cylindrical glass coil immersed in liquid nitrogen and connected to the reactor. Before transferring gases, the cold trap is pumped down to 5×10^{-5} mbar. Then, the valve isolating the trap from the reactor is opened and the gaseous species are transferred into the external cold trap. No formation of a solid residue is observed through the warming process here, contrarily to an experiment made in Titan-like conditions (Gautier et al., 2011).

The GC-MS analyses of the gas trapped are achieved using a Thermo Scientific trace GC ultra with an ITQ Thermo Scientific mass spectrometer. Gases are injected through a six port gas valve. The mass spectrometer is composed by a quadrupole using a 70 eV electron ionization system. For the gas separation, the column is a MXT-QPLOT (Restek, 30m long, 0.25 mm internal diameter and $10\text{ }\mu\text{m}$ stationary phase thickness). The column temperature is set at 40°C during 5 min, then the temperature is increased with a gradient of $5\text{ }^\circ\text{C}/\text{min}$ up to

190°C and kept at this final temperature for 5 min. Helium is used as the carrier gas (>99.9995% purity) at a constant 1.5 ml/min flow rate. A blank is done before each sample analysis.

2.6. Solid phase collection and infrared analysis of thin films

CaF₂ substrates are placed on the grounded electrode to collect thin films possibly produced during the experiments. These potential films are deposited at room temperature during 40 hours of plasma duration. Then, they are analyzed by infrared spectroscopy (Attenuated Total Reflectance technique). Samples are placed on the surface of a prism with a high refraction index (ATR crystal). The infrared signal is collected by a Deuterium TriGlycine Sulfate (DTGS) detector in the 1200-4000 cm⁻¹ range with a resolution of 4 cm⁻¹ after a co-addition of 500 scans.

3. Results

3.1. CO₂ and H₂ consumption

The first aspect of the gas phase reactivity involves the consumption of the three initial species: N₂, CO₂ and H₂. The N₂ consumption is low, with a dissociation of about 4%, theoretically evaluated in (Alves et al., 2012). The consumptions of H₂ and CO₂ are more important and are monitored at room temperature by *in situ* mass spectrometry using the time-tracking of CO₂⁺ at *m/z* 44 and H₂⁺ at *m/z* 2 at a time resolution of 0.5 s. The time-evolution of H₂ is given on Figure 2, and the time-evolution of CO₂ has been previously studied (Fleury et al., 2015).

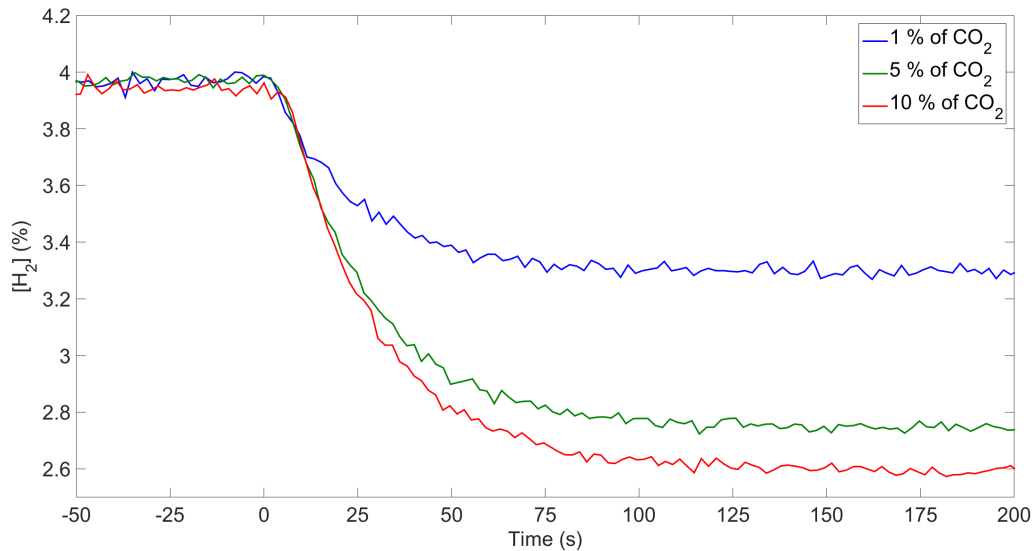


Figure 2: Evolution of the H₂ mixing ratio in the gaseous reactive medium with the plasma duration. Origin of the time is set as the moment when the plasma is turned on.

In (Fleury et al., 2015), the CO₂ consumption efficiency e_{CO_2} is defined according to the following equation (1):

$$e_{CO_2} = \frac{\Delta CO_2}{[CO_2]_0} = \frac{[CO_2]_0 - [CO_2]_{ss}}{[CO_2]_0} \quad (1),$$

where $[CO_2]_0$ and $[CO_2]_{ss}$ represent the initial and steady-state concentrations of carbon dioxide and Δ_{CO_2} represents the consumption of CO₂.

Similarly we defined the hydrogen consumption efficiency e_{H_2} with:

$$e_{H_2} = \frac{\Delta_{H_2}}{[H_2]_0} = \frac{[H_2]_0 - [H_2]_{ss}}{[H_2]_0} \quad (2),$$

where $[H_2]_0$ and $[H_2]_{ss}$ represent the initial and steady-state percentages of molecular hydrogen respectively and Δ_{H_2} the consumption of H_2 .

We defined also the elemental ratio in the fractions of gases consumed C/H:

$$\frac{C}{H} = \frac{\Delta_{CO_2}}{2\Delta_{H_2}} \quad (3),$$

Where Δ_{CO_2} represents the absolute consumption of CO_2 and Δ_{H_2} the absolute consumption of H_2 .

Similarly we defined the elemental ratio in the fractions of gases consumed O/H:

$$\frac{O}{H} = \frac{\Delta_{CO_2}}{\Delta_{H_2}} \quad (4),$$

Where Δ_{CO_2} represents the absolute consumption of CO_2 and Δ_{H_2} the absolute consumption of H_2 .

Table 1: Evolutions as a function of $[CO_2]_0$ of: Δ_{CO_2} , e_{CO_2} , Δ_{H_2} , and C/H and O/H ratios in the fraction of gases consumed. The uncertainties are given as 2σ (standard deviation) and are calculated from the standard fluctuation of the mass spectrometry measurements.

$[CO_2]_0$ (%)	Δ_{CO_2}	Δ_{H_2}	e_{CO_2} (%)	e_{H_2} (%)	C/H	O/H
1	0.24	0.72	24 ± 2	18 ± 1	0.2	0.3
5	1.2	1.3	24 ± 1	32 ± 1	0.5	0.9
10	2	1.4	20 ± 1	36 ± 1	0.7	1.4

Consumption efficiencies are reported in Table 1. A first important observation is that CO_2 and H_2 are efficiently consumed whatever the experimental conditions with a relative decrease of about 20-30%. This first result is important meaning that new molecules are

produced in the discharge from the reactions of these two species, involving the formation of possible organic molecules. Organic growth seems possible even in the absence of methane, with CO₂ as the sole carrier for organic chemistry.

The consumption efficiencies of H₂ and CO₂ are moreover rather different according to the initial CO₂ concentration. On the one hand the CO₂ relative consumption is stable at about 25%, so that when the initial CO₂ amount is multiplied by a factor of 10, the CO₂ consumed is multiplied by about the same factor. Considering the whole carbon balance among the reactants and the products in the discharge, we therefore expect a C-content similarly extended in the budget of the products. On the other hand the H₂ consumption strongly increases with the initial CO₂ concentration, but it does not compensate the larger absolute CO₂ consumption. This leads to an important increase, from 0.2 up to 0.7, of the total C/H ratio provided to the product budget by the reactant consumption. We therefore expect a production of organic products whose hydrogenation rate will decrease when the CO₂ initial concentration increases, limited by the constant 4 % H₂ initial concentration in the gas mixture. To study this possible change of the chemistry as a function of the initial amount of CO₂, we will now study the resulting products of this chemistry in a gaseous mixture made of N₂, CO₂ and H₂.

3.2. Solid phase production

A solid thin film is only observed in experiments realized with 10% of CO₂ on the CaF₂ substrates disposed on the grounded electrode. The film is organic as confirmed by its mid-IR signature presented in Figure 3.

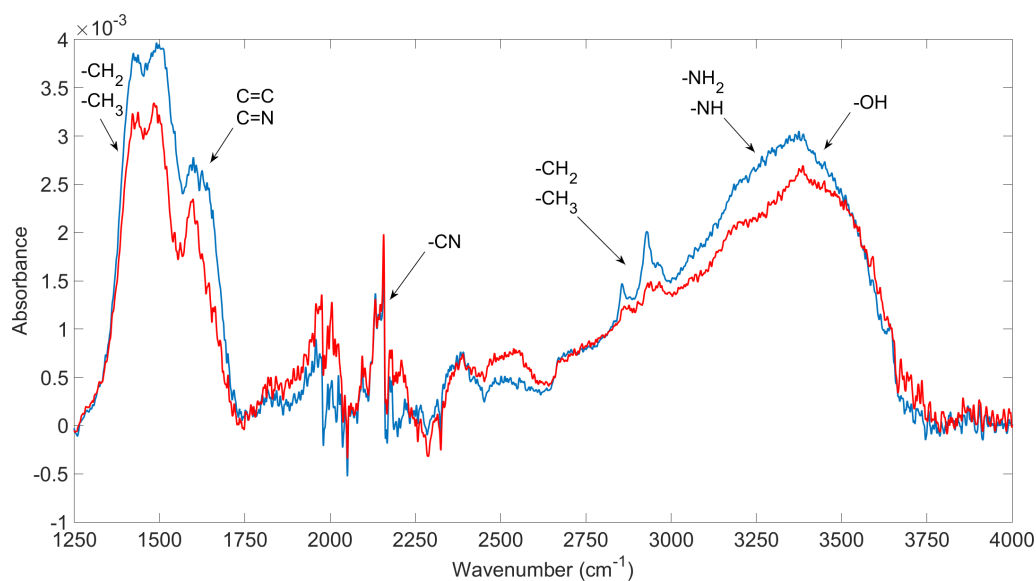


Figure 3: Infrared absorption spectra of the films deposited on two substrates during the same experiments in a $N_2/CO_2/H_2$ (86/10/4) plasma. The films were produced during 40 hours of plasma duration.

Different absorption bands characteristic of solid organics are observed in the spectrum.

First, two broad bands are observed at lower wavenumbers, centered at 1465 cm^{-1} and 1630 cm^{-1} . The 1465 cm^{-1} band can be a contribution of the asymmetric bending mode of $-CH_3$ and a scissor in plane bending mode of $-CH_2$. The second band at 1630 cm^{-1} corresponds to several possible functional groups as $C=N$ and $C=C$ double bonds, aliphatic and aromatic $-NH_2$ or aromatic and hetero-aromatic functions.

Another small absorption band is visible at 2140 cm^{-1} , which corresponds to nitrile bonds $-C\equiv N$ or isocyanides $-N\equiv C$.

The 2860 cm^{-1} band is attributed to $-CH_2$ symmetric stretching mode. The 2930 cm^{-1} and 2960 cm^{-1} bands are attributed respectively to the $-CH_2$ asymmetric stretching mode and the $-CH_3$ asymmetric stretching mode.

Finally, a broad band is observed at 3380 cm^{-1} . This band is consistent with N-H amine bonds.

No obvious C=O carbonyl bond can be detected at $\approx 1700\text{ cm}^{-1}$ in spite of the CO₂ gas reactant used in the experiment. The incorporation of oxygen in the solid material is not possible to confirm by mid-IR spectroscopy. There may be some hydroxyl bonds but those are ambiguously overlapping with amine functions. Nitrogen-bearing chemical functions are present involving at least nitrile or isocyanide functions. The formation of a N-rich solid highlights the possibility for an efficient ionospheric chemistry to provide prebiotic material on the early Earth.

3.3. Identification of the gas-phase products

We will now study the gaseous intermediate species formed during the experiments. In order to detect and identify gas species produced in low quantity, we have trapped the volatile species and accumulated them by cooling the plasma box as described in section 2.2. The plasma box is then warmed up to room temperature. A pressure increase is observed associated to the sublimation of the trapped species.

Ammonia: a precursor of the solid phase

The broad band observed at 3380 cm^{-1} suggests an important amine content in the solid phase produced with 10% of CO₂. A possible precursor in the gas phase for this chemical signature could be ammonia. To confirm this, we realized an analysis of the gaseous phase by infrared spectroscopy. Figure 4 presents an infrared spectrum in the $700\text{-}1200\text{ cm}^{-1}$ range recorded at room temperature after the release of the trapped gases.

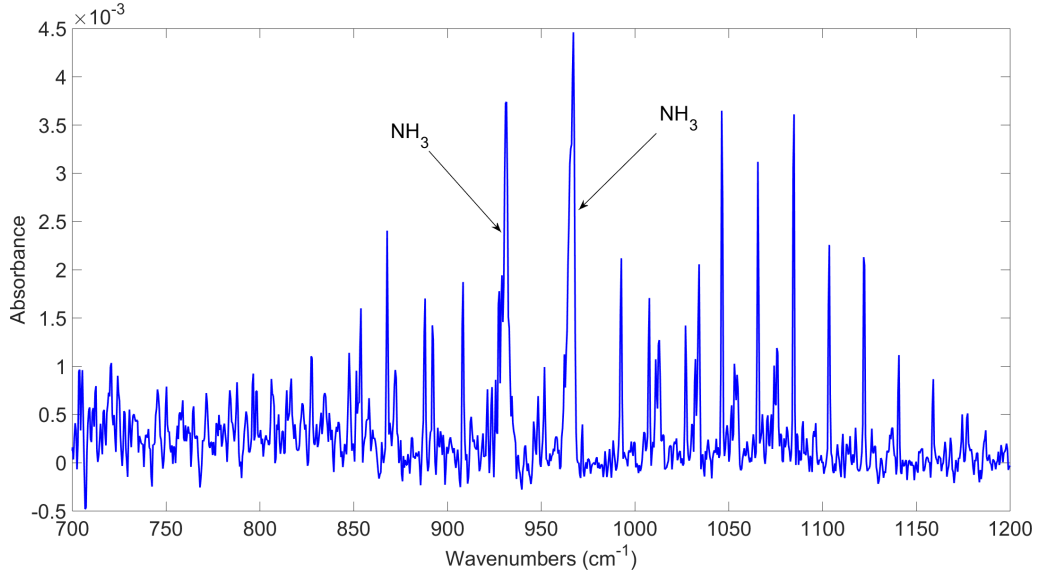


Figure 4: Infrared spectra in the 700-1200 cm^{-1} range recorded at 300 K after 4 hours of cryogenic trapping in a $\text{N}_2\text{-CO}_2\text{-H}_2$ (91/5/4) plasma.

This spectrum presents two strong absorption bands at 930 cm^{-1} and 967 cm^{-1} . They corresponds to the two Q branches of the ν_2 vibrational transition of NH_3 . The existence of these two bands, characteristic of the ammonia molecule, is caused by the inversion of the molecule by quantum tunneling effect. This results in the splitting of the low energy levels in a pair with slightly different energy values. In addition, some rotational lines of the P and R branches are also observed on this spectrum. Those bands allow its identification unambiguously. Since NH_3 has absorption bands with no overlap over other species, it is possible to estimate its concentration using the Beer-Lambert law. The absorbance $A(\lambda)$ at a given wavelength is defined by:

$$A(\lambda) = \varepsilon(\lambda) \times l \times [C] \quad (5),$$

where $\varepsilon(\lambda)$ is the absorption cross-section of the molecule at a given wavelength, l is the path length of the beam through the gas cell and $[C]$ is the concentration of absorbing molecules in the reactor.

$$[C] = \frac{A(\lambda)}{\varepsilon(\lambda) \times l} \quad (6)$$

The ammonia absorption cross-sections have been calculated from the line by line parameters provided by the HITRAN 2012 database (Rothman et al., 2013), using the HITRAN Application Programming Interface (HAPI) (Kochanov et al., 2016). In order to overcome the difference of resolution between the database and laboratory data, we calculate the concentration from the area of the bands. For NH₃, integration is performed both in the 920-940 cm⁻¹ and in the 950-970 cm⁻¹ ranges.

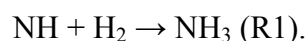
Knowing the volume of the reactor we calculate the number of molecules of ammonia N_{NH3} trapped during four hours. Table 2 presents the evolution of N_{NH3} as a function of the CO₂ initial amount. A good agreement is found for N_{NH3} calculated with the two bands.

Table 2: Evolution as a function of [CO₂]₀ of the number of NH₃ molecules formed in 4 hours of plasma duration. The uncertainties are given as 2σ (standard deviation) and are calculated from the standard fluctuations of the infrared spectroscopy measurements.

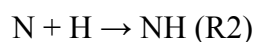
[CO ₂] ₀ (%)	N _{NH3} (molecules) 920-940 cm ⁻¹	N _{NH3} (molecules) 950-970 cm ⁻¹
1	(5.3 ± 0.1) × 10 ¹⁶	(6.4 ± 0.5) × 10 ¹⁶
5	(6.7 ± 0.4) × 10 ¹⁶	(6.9 ± 0.5) × 10 ¹⁶
10	-	(4.1 ± 0.2) × 10 ¹⁵

The number of ammonia molecules trapped in the 1% and 5% conditions is similar, but this number decreases by one order of magnitude for the 10% case. This result comes in front of the solid organic phase production, which is observed in the only 10% case. So compared to the lower CO₂ initial concentrations, in the 10% condition NH₃ disappears from the gas phase whereas an N-H bond signature is observed in the solid. These observations suggest that in all likelihood ammonia is a precursor for the solid phase with 10% of CO₂ and leads to the amine signature of the solid.

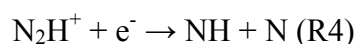
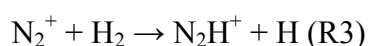
The formation of ammonia in our CO₂-N₂-H₂ plasma requires similar pathways as discussed in (Carrasco et al., 2012) for the case of a N₂-CH₄ plasma. Those involve NH radicals reacting with molecular hydrogen (Carrasco et al., 2012):



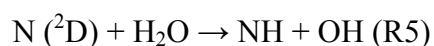
NH radicals are produced at least through three pathways considering the literature and our gaseous mixture. The first one involves radical chemistry (Carrasco et al., 2012; Mutsukura, 2001):



The second one involves ion chemistry (Green and Caledonia, 1982):



The last one involves radical chemistry (Dobrijevic et al., 2014) and is related to the production of water in the CO₂-N₂-H₂ discharge (Fleury et al., 2015):



No other molecule than water and ammonia can be detected by IR spectroscopy by lack of sensitivity. We therefore complete the gas phase analysis, first with ex-situ GC-MS and secondly by MS analysis.

Ex situ analysis by GC-MS of the gaseous products after cryogenic trapping

The cryogenic trapping is performed at 173 K for 8 hours of plasma duration. After the warming of the plasma box, the pressure in the reactor is about 4.1 mbar. Then the gases released are transferred into an external cold trap, resulting in a pressure of about 50 mbar in

the external trap before the injection in the GC-MS. The chromatogram obtained for an initial gaseous mixture containing 5% of CO₂ is presented in the top of the Figure 5. All the peaks have been identified using their retention time and their mass spectra. Four species are detected.

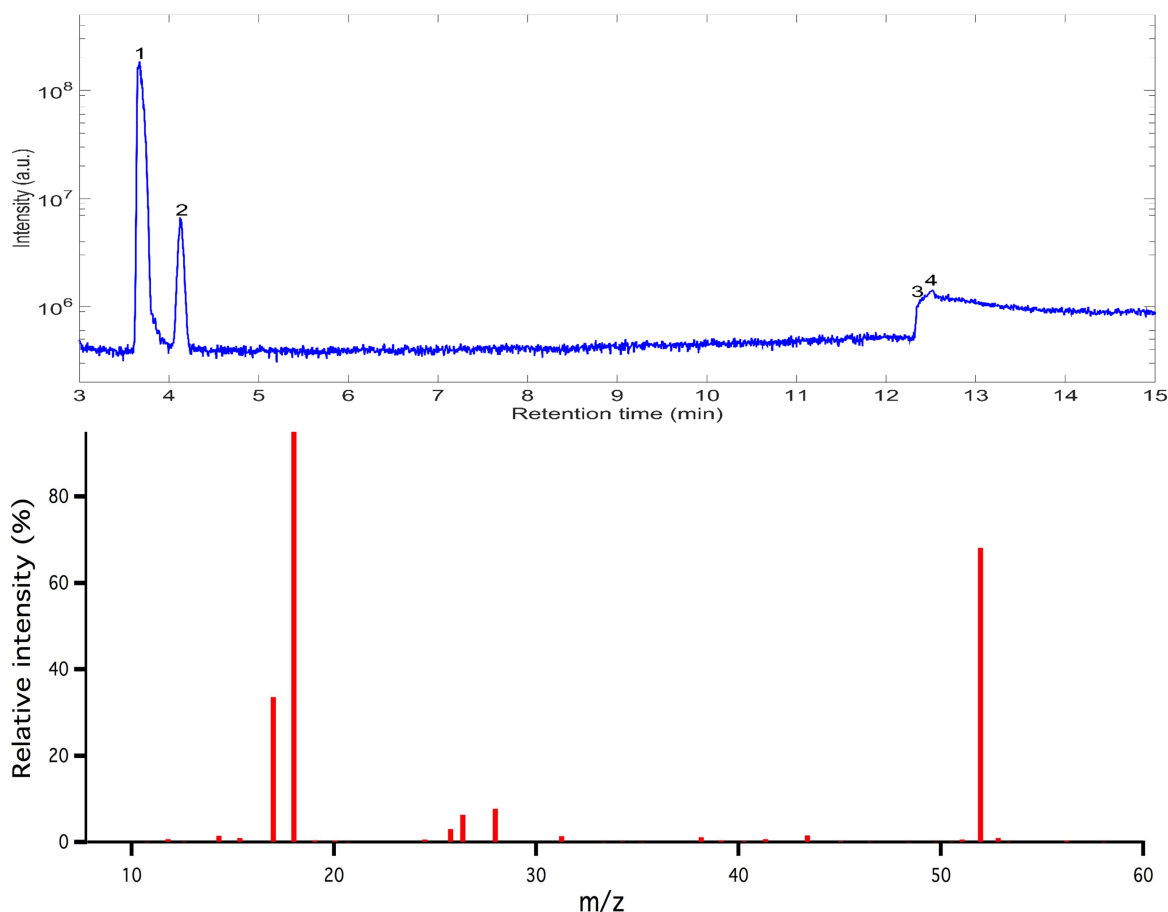


Figure 5: Top: Chromatogram of the gases trapped during an experiment with 5% of CO₂. Bottom: mass spectrum corresponding of the peak 4 of the chromatogram. This peak is attributed to C₂N₂ in agreement with the ion detected at m/z 52. The presence of the mass peak at m/z 17 and 18 is explained by the presence of water co-eluted with C₂N₂.

The first molecule identified is carbon dioxide, with a retention time of 3.42 min, and corresponds to a major peak in the chromatogram. This detection is surprising. Indeed, CO₂ is not expected to condense in these conditions of pressure and temperature. Most likely explanation, an important formation of water ice is observed during the experiment, which can trap small molecules depending on the temperature (Bar-nun et al., 1985; Notesco and Bar-Nun, 1997). This unexpected trapping of CO₂ has to be taken into account in the

following and possibly extended to other gas products, which are not expected to condense such as hydrocarbons with 2 carbon atoms.

The second species detected is nitrous oxide N_2O with a retention time of 4.06 min. N_2O has previously been detected in the gaseous products in a simulation realized with the PAMPRE experiment and using an initial gaseous mixture made of N_2 , CO and CH_4 (Fleury et al., 2014).

The third species detected is water with a retention time of 12.34 min. because of the saturation of the chromatographic column. This detection of water is in agreement with its previous detection (Fleury et al. 2015).

And the last detected species, co-eluted with water, is ethanedinitrile (C_2N_2) with a retention time of 12.52 min. The mass spectrum corresponding to this peak is presented in the bottom of the Figure 5. The important peak at m/z 52, with a relative intensity of 70 %, allows an unambiguous identification of C_2N_2 . The presence of two other important peaks at m/z 17 and 18 is explained by the co-elution of water with C_2N_2 and, which is responsible for these two peaks in the mass spectrum.

[Analysis of the trapped gaseous products by mass spectrometry](#)

The identification of the gaseous products is completed by a direct analysis of the gas trapped released in the reactor by using mass spectrometry. Mass spectra are found similar for the 5 and the 10% CO_2 experiments, in agreement with the total pressure of the released products. Intensities are found lower in the 1% CO_2 experiment, also in agreement with the total gas pressure collected at the end of the warming. No qualitative difference is noted according to the initial CO_2 concentration. The further analysis is therefore focused on one condition, the 5% CO_2 experiment.

Figure 6 presents two mass spectra obtained with an initial CO₂ amount of 5%: the first spectrum is recorded at 173 K and the second after the warming of the plasma box to room temperature (294 K). As water and ammonia were previously identified and quantified by FT-IR spectroscopy, the interpretation of the spectra is achieved only for $m/z > 20$.

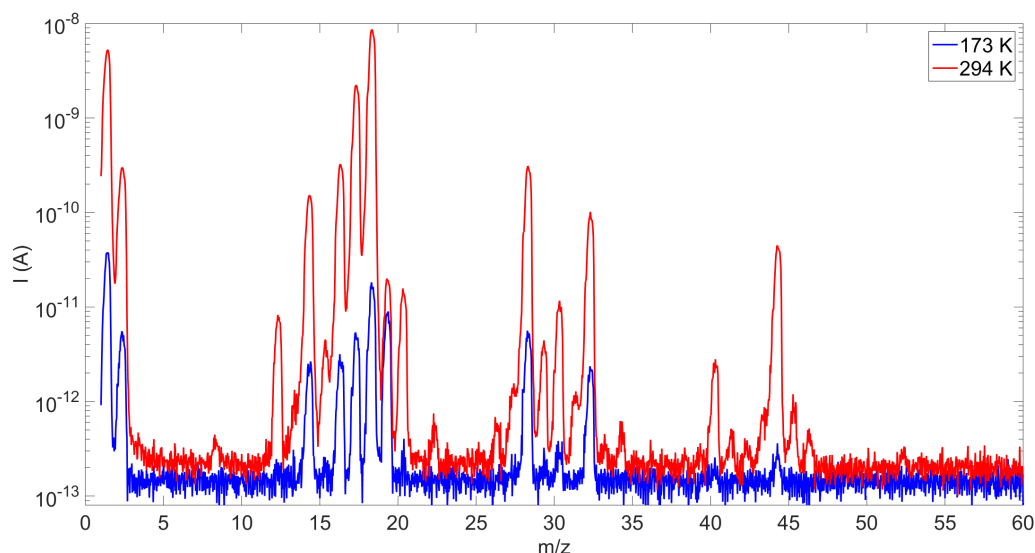


Figure 6: Mass spectra recorded at 173 K (blue) and 294 K (red) after 4 hours of gaseous trapping in N₂-CO₂-H₂ (91/5/4) plasma.

First, the signature of the residual air in the mass spectrometer (contributions of N₂, O₂, CO₂ and Ar) is identified in the mass spectrum recorded at 173 K. It corresponds to the vacuum limit of the MS pumping, at 3×10^{-8} mbar.

At room temperature, Figure 6 shows an evolution of the mass spectrum with the one recorded at 173 K. We observe the increase of the intensity of some peaks and the apparition of new ones with masses up to 60 u. This reflects a release of gaseous species in agreement with the increase of the pressure measured in the reactor.

The low resolution of our mass spectrometer does not allow differentiating molecules with close masses. However, few molecules can be tentatively identified. As discussed previously in the GC-MS section with the detection of CO₂, the important formation of water ice on the walls of the plasma box impacts the species simultaneously trapped. This water trapping

explains the four highest intensities for $m/z > 20$ observed in Figure 6 at 294 K: m/z 28, 32, 40 and 44 for N_2 and CO, O_2 , Ar and CO_2 respectively. Indeed these molecules do not condense in the conditions of pressure and temperature of the experiment. Given the detection of N_2O by GC-MS analysis, a contribution of this molecule is expected at m/z 44 in addition to the main CO_2 signature. Indeed, in mass spectrometry, N_2O has a principal fragment at m/z 44, which is mixed with the CO_2 signature according to the analysis of the gaseous phase achieved by GC-MS and presented above. The second fragment is at m/z 30 and represents 31% of m/z 44 intensity. This pattern is compatible with the mass spectrum presented in Figure 6.

A first important precursor for the formation of complex organic molecules is hydrogen cyanide HCN. Its signature in mass spectrometry is a major peak at m/z 27 with a relative contribution at m/z 26 (20%). This signature is visible on the mass spectrum in Figure 6, but is low, involving a low concentration of HCN in the gas phase. This concentration is even under the detection limit of our FTIR diagnosis so that it cannot be definitely quantified in the experiment. A pathway for HCN formation is well established in the case of the chemistry of N_2 and CH_4 in Titan's atmosphere (Hébrard et al., 2012; Krasnopolsky, 2009; Loison et al., 2015; Wilson and Atreya, 2003). It involves the key species CH_3 produced by the photolysis of CH_4 in Titan's atmosphere. This mechanism is less obvious in a $N_2/CO_2/H_2$ plasma discharge, in agreement with the low concentration of HCN observed.

At m/z 26 the ratio between the peak at 27 and 26 is larger than the ratio expected from the NIST database for HCN. This more important intensity of the m/z 26 peak can be explained by the contribution of all abundant nitriles with their CN^+ fragments even if the presence of acetylene C_2H_2 cannot be discarded. Indeed, if C_2H_2 does not condense in these conditions of pressure and temperature but could be trapped in water ice.

Given the trapping of N₂ or CO revealed at m/z 28, the peak at m/z 29 could be mainly attributed to the isotopologues ¹⁴N¹⁵N or ¹³C¹⁶O. In our experiment, the intensity ratios I₂₉:I₂₈ are 100:1.13, while the expected I₂₉:I₂₈ ratios for natural abundance isotopologues of N₂ and CO are 100:0.75 and 100:1.12 respectively as given by the NIST database. The ratio measured in our experiment is in good agreement with the one expected for CO, which is so the most likely candidate contributing to these peaks. The presence of CO in the mass spectrum can result of the dissociation of CO₂ in the mass spectrometer or of its trapping during the experiment, CO being a major product of the CO₂ dissociation in the plasma.

The peak at m/z 30 is important in the mass spectrum observed at 294 K. There are two possible species at this mass: formaldehyde H₂CO and ethane C₂H₆ with a similar explanation for ethane trapping as for acetylene. Given the abundance of the reactants CO₂ and H₂ in the gas mixture, formaldehyde would be expected as an important contributor. H₂CO are at 29 and 30 amu with a 29:30 ratio of 100:58. This is in disagreement with the ratio observed in the mass spectrum of the Figure 6, where the main contribution to the peak at 29 amu is attributed to ¹³CO. Our results point out a possible production of ethane even in the oxidant conditions used in the present work.

At m/z 41, one possible species is acetonitrile CH₃CN, in agreement with the formation of HCN, which is a precursor of CH₃CN (Dobrijevic and Dutour, 2006; Gautier et al., 2011).

The intensities of the peaks at m/z 45 and 46 are consistent with isotopes of CO₂.

And a last peak is observed at m/z 52 in agreement with the detection of ethanedinitrile C₂N₂ by GC-MS.

In conclusion in situ mass spectrometry completes our overview of the gas products by pointing out nitrile molecules. Among them hydrogen cyanide HCN is an important

molecules for prebiotic chemistry, for example involved with ammonia NH_3 in the Strecker amino acid synthesis (Strecker, 1854).

4. Discussion

The possible formation of organic aerosols in the atmosphere of the early Earth is of the prime interest for the comprehension of the environment of the Earth during the Hadean and the Archean eons. First of all, the formation of fractal organic hazes can provide an ultraviolet shield for the early Earth (Wolf and Toon, 2010) against the high UV flux of the young Sun (Claire et al., 2012). This shield affects the composition of the atmosphere protecting some constituents of the atmosphere such as ammonia from photochemical destruction (Sagan and Chyba, 1997). Furthermore, the formation of haze would affect the climate of the early Earth providing an antigreenhouse effect (Haqq-Misra et al., 2008; Hasenkopf et al., 2011; McKay et al., 1999).

In our gaseous mixture, solid organic products are only observed as deposited on a substrate. The solid formation seems to be promoted on the substrate with the formation of an organic thin film, which has grown on a substrate placed on the grounded electrode of the plasma device. On the contrary, the formation of spherical shaped individual grains is not observed. The mechanism of formation of these two types of solid in the plasma is not fully known. Both grow from reactive gas species present in the plasma. For the formation of grains in the plasma, the process starts with the formation of nanometer-size monomers, which aggregate to form larger particles with a negative charge when the density of the monomers in the plasma reaches a critical value. Then the particles grow by deposition of species present in the gaseous phase (Wattiaux et al., 2015). The hypothesis, which can be done to explain the absence of these grains in our gaseous mixture compared to CH_4 -rich experiments (Szopa et

al. 2006) is that the lower reactivity of CO_2 does not allow to reach the critical value of monomers in the plasma necessary to start the growth of larger particles.

With such a mechanism, we expect that the water formation observed in this gaseous mixture would sustain the formation of solid organic in the case of the atmosphere of the early Earth. Indeed, as discussed in Fleury et al. 2015, water formed above the troposphere leads to the formation of high altitude clouds analogues to the present PSC and PMC. These clouds are composed of water ice, which would play the role of nucleus cores, promoting a heterogeneous nucleation of organic molecules at the surface of ice grains.

Moreover, the formation of organic volatiles in interaction with water ice can also be important in the perspective of a potential enrichment of the lower atmospheric layers of the early Earth by prebiotic molecules formed in the upper atmosphere. Indeed, these molecules could be trapped in the water ice clouds and transport to the lower atmospheric layers involving an important source of nitrogenized molecules for the lower atmospheric layers.

Conclusion

We have shown that an organic growth is possible in an oxidized atmosphere made of N_2 , CO_2 and H_2 simulating the early Earth atmosphere. This suggests that an organic growth can be sustained only by carbon composing CO_2 in the primitive atmosphere of the Earth without methane. Firstly this results in the formation of volatiles molecules with the identification of H_2O , NH_3 , C_2N_2 , N_2O and the possible detection of HCN. The detections of these molecules are interesting in a prebiotic chemistry point of view. Indeed, they are reactive molecules, which conduct to the formation of more complex molecules and HCN and NH_3 are involved in the formation of amino acids and nitrogenous bases. Secondly we have shown that in the absence of methane CO_2 also plays a role in the formation of organic compounds. Indeed, if the composition of the gaseous phase does not change as a function of

the CO₂ initial amount, an increase of the CO₂ content promotes the formation of organic compounds highlighting that a high level of CO₂ in the primitive atmosphere of the Earth can promote organic chemistry in such an atmosphere. And finally, an organic thin film formation is observed in the experiments requiring a threshold concentration of CO₂. Ammonia has also been shown to be precursor for this solid formation. This highlights the important complexity of the organic chemistry, which could be initiated in the early Earth atmosphere from N₂, CO₂ and H₂.

Another result, from this work is that in our gaseous mixture, solid organic products are only observed as thin films deposited on a substrate and a low quantity. In such mechanism, the water formation observed in our experiment can play a role in the formation of solid organic in the atmosphere of the early Earth. Indeed, water formed above the troposphere would conduct to the formation of high altitude clouds analogues to the present PSC and PMC. These clouds could then play the role on nucleus cores, promoting a heterogeneous nucleation of organic at the surface of ice grains.

In this work we have investigated the influence of the CO₂ concentration on the chemistry but the hydrogen concentration in the primitive atmosphere of the Earth is also debated. Further work will be necessary to understand the influence of the hydrogen concentration of the early Earth's atmospheric reactivity for different concentrations of carbon dioxide.

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References

- Alves, L.L., Marques, L., Pintassilgo, C.D., Wattieaux, G., Et, E.-s., Berndt, J., Kovacević, E., Carrasco, N., Boufendi, L., Cernogora, G., 2012. Capacitively coupled radio-frequency discharges in nitrogen at low pressures. *Plasma Sources Science and Technology* 21, 045008.
- Bar-nun, A., Herman, G., Laufer, D., Rappaport, M.L., 1985. Trapping and release of gases by water ice and implications for icy bodies. *Icarus* 63, 317-332.
- Carrasco, N., Gautier, T., Es-sebbar, E.-t., Pernot, P., Cernogora, G., 2012. Volatile products controlling Titan's tholins production. *Icarus* 219, 230-240.
- Catling, D.C., 2006. Comment on "A Hydrogen-Rich Early Earth Atmosphere". *Science* 311, 38.
- Charnay, B., Forget, F., Wordsworth, R., Leconte, J., Millour, E., Codron, F., Spiga, A., 2013. Exploring the faint young Sun problem and the possible climates of the Archean Earth with a 3-D GCM. *Journal of Geophysical Research: Atmospheres* 118, 4104-410431.
- Chyba, C., Sagan, C., 1992. Endogenous production, exogenous delivery and impact-shock synthesis of organic molecules: an inventory for the origins of life. *Nature* 355, 125-132.
- Claire, M.W., Sheets, J., Cohen, M., Ribas, I., Meadows, V.S., Catling, D.C., 2012. The Evolution of Solar Flux from 0.1 nm to 160 μ m: Quantitative Estimates for Planetary Studies. *The Astrophysical Journal* 757, 95.
- Cottin, H., Kotler, J.M., Bartik, K., Cleaves, H.J., Cockell, C.S., Vera, J.-P.P., Ehrenfreund, P., Leuko, S., Ten Kate, I.L., Martins, Z., Pascal, R., Quinn, R., Rettberg, P., Westall, F., 2015. Astrobiology and the Possibility of Life on Earth and Elsewhere.... *Space Sci Rev*, 1-42.
- Delano, J., 2001. Redox History of the Earth's Interior since \sim 3900 Ma: Implications for Prebiotic Molecules. *Orig Life Evol Biosph* 31, 311-341.
- Denman, K.L., G. Brasseur, A. Chidthaisong, Ciais, P., Cox, P.M., Dickinson, R.E., Hauglustaine, D., Heinze, C., Holland, E., Jacob, D., Lohmann, U., Ramachandran, S., Dias, P.L.d.S., Wofsy, S.C., X. Zhang, 2007. Couplings Between Changes in the Climate System and Biogeochemistry, in: Solomon, S., Qin, D., Manning, M., Chen, Z., Marquis, M., Avery, K.B., M.Tignor, Miller, H.L. (Eds.), *Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change*. Cambridge University Press.
- Dobrijevic, M., Dutour, I., 2006. A random graphs model for the study of chemical complexity in planetary atmospheres. *Planetary and Space Science* 54, 287-295.

Dobrijevic, M., Hébrard, E., Loison, J.C., Hickson, K.M., 2014. Coupling of oxygen, nitrogen, and hydrocarbon species in the photochemistry of Titan's atmosphere. *Icarus* 228, 324-346.

Emmanuel, S., Ague, J.J., 2007. Implications of present-day abiogenic methane fluxes for the early Archean atmosphere. *Geophysical Research Letters* 34, L15810.

Feulner, G., 2012. The faint young Sun problem. *Reviews of Geophysics* 50, RG2006.

Fleury, B., Carrasco, N., Gautier, T., Mahjoub, A., He, J., Szopa, C., Hadamcik, E., Buch, A., Cernogora, G., 2014. Influence of CO on Titan atmospheric reactivity. *Icarus* 238, 221-229.

Fleury, B., Carrasco, N., Marcq, E., Vettier, L., Määttänen, A., 2015. Water Formation in the Upper Atmosphere of the Early Earth. *The Astrophysical Journal Letters* 807, L29.

Gautier, T., Carrasco, N., Buch, A., Szopa, C., Sciamma-O'Brien, E., Cernogora, G., 2011. Nitrile gas chemistry in Titan's atmosphere. *Icarus* 213, 625-635.

Green, B.D., Caledonia, G.E., 1982. Observations of NH infrared chemiluminescence. *The Journal of Chemical Physics* 77, 3821-3823.

Guzmán-Marmolejo, A., Segura, A., Escobar-Briones, E., 2013. Abiotic Production of Methane in Terrestrial Planets. *Astrobiology* 13, 550-559.

Haqq-Misra, J.D., Domagal-Goldman, S.D., Kasting, P.J., Kasting, J.F., 2008. A Revised, Hazy Methane Greenhouse for the Archean Earth. *Astrobiology* 8, 1127-1137.

Hasenkopf, C.A., Freedman, M.A., Beaver, M.R., Toon, O.B., Tolbert, M.A., 2011. Potential Climatic Impact of Organic Haze on Early Earth. *Astrobiology* 11, 135-149.

Hébrard, E., Dobrijevic, M., Loison, J.C., Bergeat, A., Hickson, K.M., 2012. Neutral production of hydrogen isocyanide (HNC) and hydrogen cyanide (HCN) in Titan's upper atmosphere. *A&A* 541, A21.

Kasting, J., 1993. Earth's early atmosphere. *Science* 259, 920-926.

Kasting, J.F., 2005. Methane and climate during the Precambrian era. *Precambrian Research* 137, 119-129.

Kochanov, R.V., Gordon, I.E., Rothman, L.S., Wcisło, P., Hill, C., Wilzewski, J.S., 2016. HITRAN Application Programming Interface (HAPI): A comprehensive approach to working with spectroscopic data. *Journal of Quantitative Spectroscopy and Radiative Transfer* 177, 15-30.

Krasnopolsky, V.A., 2009. A photochemical model of Titan's atmosphere and ionosphere. *Icarus* 201, 226-256.

Kuramoto, K., Umemoto, T., Ishiwatari, M., 2013. Effective hydrodynamic hydrogen escape from an early Earth atmosphere inferred from high-accuracy numerical simulation. *Earth and Planetary Science Letters* 375, 312-318.

Loison, J.C., Hébrard, E., Dobrijevic, M., Hickson, K.M., Caralp, F., Hue, V., Gronoff, G., Venot, O., Bénilan, Y., 2015. The neutral photochemistry of nitriles, amines and imines in the atmosphere of Titan. *Icarus* 247, 218-247.

Martin, H., Claeys, P., Gargaud, M., Pinti, D., Selsis, F., 2006. 6. Environmental Context. *Earth Moon Planet* 98, 205-245.

McKay, C.P., Lorenz, R.D., Lunine, J.I., 1999. Analytic Solutions for the Antigrreenhouse Effect: Titan and the Early Earth. *Icarus* 137, 56-61.

Miller, S.L., 1953. A Production of Amino Acids Under Possible Primitive Earth Conditions. *Science* 117, 528-529.

Mutsukura, N., 2001. Deposition of Diamondlike Carbon Film and Mass Spectrometry Measurement in CH₄/N₂ RF Plasma. *Plasma Chemistry and Plasma Processing* 21, 265-277.

Nisbet, E.G., Sleep, N.H., 2001. The habitat and nature of early life. *Nature* 409, 1083-1091.

Notesco, G., Bar-Nun, A., 1997. Trapping of Methanol, Hydrogen Cyanide, and n-Hexane in Water Ice, above Its Transformation Temperature to the Crystalline Form. *Icarus* 126, 336-341.

Rothman, L.S., Gordon, I.E., Babikov, Y., Barbe, A., Chris Benner, D., Bernath, P.F., Birk, M., Bizzocchi, L., Boudon, V., Brown, L.R., Campargue, A., Chance, K., Cohen, E.A., Coudert, L.H., Devi, V.M., Drouin, B.J., Fayt, A., Flaud, J.M., Gamache, R.R., Harrison, J.J., Hartmann, J.M., Hill, C., Hodges, J.T., Jacquemart, D., Jolly, A., Lamouroux, J., Le Roy, R.J., Li, G., Long, D.A., Lyulin, O.M., Mackie, C.J., Massie, S.T., Mikhailenko, S., Müller, H.S.P., Naumenko, O.V., Nikitin, A.V., Orphal, J., Perevalov, V., Perrin, A., Polovtseva, E.R., Richard, C., Smith, M.A.H., Starikova, E., Sung, K., Tashkun, S., Tennyson, J., Toon, G.C., Tyuterev, V.G., Wagner, G., 2013. The HITRAN2012 molecular spectroscopic database. *Journal of Quantitative Spectroscopy and Radiative Transfer* 130, 4-50.

Sagan, C., Chyba, C., 1997. The Early Faint Sun Paradox: Organic Shielding of Ultraviolet-Labile Greenhouse Gases. *Science* 276, 1217-1221.

Schlesinger, G., Miller, S., 1983a. Prebiotic synthesis in atmospheres containing CH₄, CO, and CO₂. *J Mol Evol* 19, 383-390.

Schlesinger, G., Miller, S., 1983b. Prebiotic synthesis in atmospheres containing CH₄, CO, and CO₂. *J Mol Evol* 19, 376-382.

Strecker, A., 1854. Ueber einen neuen aus Aldehyd - Ammoniak und Blausäure entstehenden Körper. *Justus Liebigs Annalen der Chemie* 91, 349-351.

Szopa, C., Cernogora, G., Boufendi, L., Correia, J.J., Coll, P., 2006. PAMPRE: A dusty plasma experiment for Titan's tholins production and study. *Planetary and Space Science* 54, 394-404.

Tian, F., Toon, O.B., Pavlov, A.A., 2006. Response to Comment on "A Hydrogen-Rich Early Earth Atmosphere". *Science* 311, 38.

- Tian, F., Toon, O.B., Pavlov, A.A., De Sterck, H., 2005. A Hydrogen-Rich Early Earth Atmosphere. *Science* 308, 1014-1017.
- Tomasko, M., West, R., 2010. Aerosols in Titan's Atmosphere, in: Brown, R., Lebreton, J.-P., Waite, J.H. (Eds.), *Titan from Cassini-Huygens*. Springer Netherlands, pp. 297-321.
- Trail, D.E.B.N.D., 2011. The oxidation state of Hadean magmas and implications for early Earth's atmosphere. *Nature* 480, 79-82.
- Trainer, M.G., Pavlov, A.A., Curtis, D.B., McKay, C.P., Worsnop, D.R., Delia, A.E., Toohey, D.W., Toon, O.B., Tolbert, M.A., 2004. Haze Aerosols in the Atmosphere of Early Earth: Manna from Heaven. *Astrobiology* 4, 409-419.
- Trainer, M.G., Pavlov, A.A., DeWitt, H.L., Jimenez, J.L., McKay, C.P., Toon, O.B., Tolbert, M.A., 2006. Organic haze on Titan and the early Earth. *Proceedings of the National Academy of Sciences* 103, 18035-18042.
- Waite, J.H., Jr., Young, D.T., Westlake, J.H., Lunine, J.I., McKay, C.P., Lewis, W.S., 2010. High-Altitude Production of Titan's Aerosols, in: Brown, R., Lebreton, J.-P., Waite, J.H. (Eds.), *Titan from Cassini-Huygens*. Springer Netherlands, pp. 201-214.
- Wattiaux, G., Carrasco, N., Henault, M., Boufendi, L., Cernogora, G., 2015. Transient phenomena during dust formation in a N_2 - CH_4 capacitively coupled plasma. *Plasma Sources Science and Technology* 24, 015028.
- Wilson, E.H., Atreya, S.K., 2003. Chemical sources of haze formation in Titan's atmosphere. *Planetary and Space Science* 51, 1017-1033.
- Wolf, E.T., Toon, O.B., 2010. Fractal Organic Hazes Provided an Ultraviolet Shield for Early Earth. *Science* 328, 1266-1268.