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1 **Intrusion-Related Gold Deposits: new insights from gravity and hydrothermal integrated 3D**
2 **modeling applied to the Tighza gold mineralization (Central Morocco)**

3
4 Khalifa Eldursi ^{a,*(1)}, Yannick Branquet^{a,b,c}, Laurent Guillou-Frottier^c, Guillaume Martelet^c,
5 Philippe Calcagno^c

6 ^a Université d'Orléans, ISTO, UMR 7327, 1A rue de la Férellerie, 45071 Orléans, Cedex 2, France

7 ^b CNRS/INSU, ISTO, UMR 7327, 45071 Orléans, France

8 ^c BRGM, ISTO, UMR 7327, 3 av. Claude Guillemin, BP 36009, 45060 Orléans, Cedex 2, France

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17 * Corresponding author: Khalifa Eldursi

18 eldursik@uregina.ca

19 Phone: +1 639-999-7464

20 Fax: +1 306-585-5433

(1) Present address: Department of Geology, University of Regina, 3737 Wascana Parkway, Regina,
SK, S4S 0A2, Canada

21 Abstract

22 The Tighza (or Jebel Aouam) district is one of the most important polymetallic districts in Morocco. It
23 belongs to the Variscan Belt of Central Meseta, and includes W-Au, Pb-Zn-Ag, and Sb-Ba
24 mineralization types that are spatially related to late-Carboniferous granitic stocks. One of the
25 proposed hypotheses suggests that these granitic stocks are connected to a large intrusive body lying
26 beneath them and that W-Au mineralization is directly related to this magmatism during a 287-285 Ma
27 time span. A more recent model argues for a disconnection between the older barren outcropping
28 magmatic stocks and a younger hidden magmatic complex responsible for the W-Au mineralization.
29 Independently of the magmatic scenario, the W-Au mineralization is consensually recognized as of
30 intrusion-related gold deposit (IRGD) type, W-rich. In addition to discrepancies between magmatic
31 sceneries, the IRGD model does not account for published older age corresponding to a high-
32 temperature hydrothermal event at ca. 291 Ma. Our study is based on gravity data inversion and hydro-
33 thermal modeling, and aims to test this model of IRGD and its related magmatic geometries, with
34 respect to subsurface geometries, favorable physical conditions for deposition and time record of
35 hydrothermal processes. Combined inversion of geology and gravity data suggests that an intrusive
36 body is rooted mainly at the Tighza fault in the north and that it spreads horizontally toward the south
37 during a trans-tensional event (D2). Based on the numerical results, two types of mineralization can be
38 distinguished: 1) the “Pre-Main” type appears during the emplacement of the magmatic body, and 2)
39 the “Main” type appears during magma crystallization and the cooling phase. The time-lag between
40 the two mineralization types depends on the cooling rate of magma. Although our numerical model of
41 thermally-driven fluid flow around the Tighza pluton is simplified, as it does not take into account the
42 chemical and deformation contributions, it provides evidence for abandoning the time-lag as an
43 argument for invalidating the existence of a genetic link between older mineralizations and/or
44 hydrothermal events spatially related to younger parent plutons.

45 **Keywords:** Jebel Aouam, intrusion-related gold deposits, hydrothermal fluid flow, 3D geological and
46 gravity modeling, combined inversion, rock alteration index

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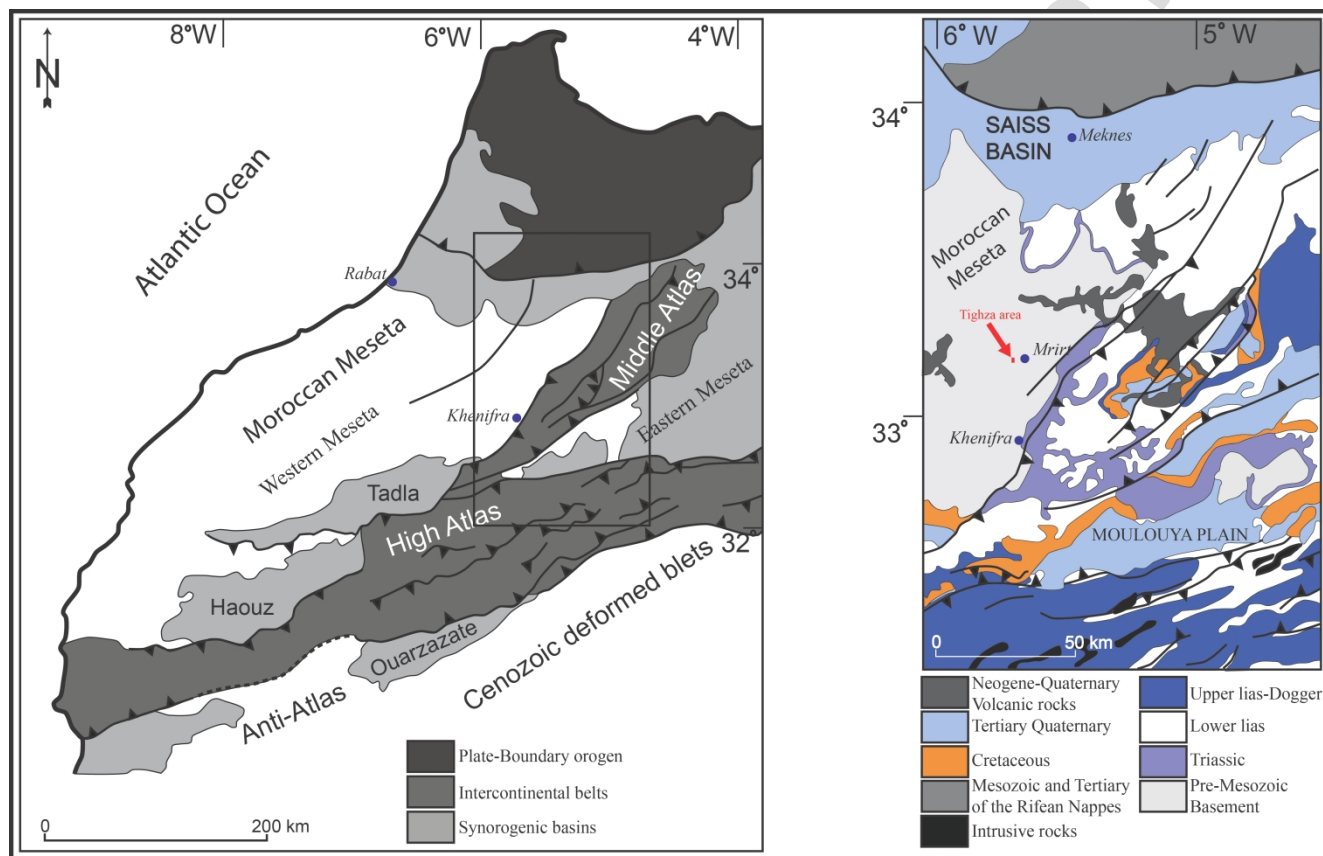
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69 1. Introduction

70 Intrusion-Related Gold Deposits (IRGD) form a major class of ore deposits, including various deposit
71 styles as diverse as skarns, mineralized breccias, sheeted veins, and disseminations within or peripheral
72 to the intrusion (Thompson et al. 1999; Baker and Lang 2001; Lang and Baker 2001; Hart 2005). They
73 form a distinct class from gold-rich porphyry deposits (Sillitoe 2000) because: i) they are associated
74 with moderately oxidized to reduced small intrusions (hence the name “Reduced Intrusion-Related
75 Gold Deposits” preferred by some authors, such as Hart 2007) which may be emplaced at depths down
76 to 8 km; ii) the metal association mostly includes tin, tungsten, molybdenum, bismuth, tellurium, and
77 arsenic, rather than copper and silver; iii) the sulfide mineral content is low in veins; and iv) the
78 hydrothermal fluids are CO₂-rich. The IRGD mineralization is located within the intrusion and/or its
79 fractured thermal aureole. The intrusions are weakly deformed as they postdate regional ductile
80 shearing and peak metamorphism. The exposed intrusions form one to two kilometer-wide stocks
81 which may correspond to apices or cupolas of an underlying larger plutonic body.

82 IRGD are generally assumed to be genetically linked to, and coeval with, the emplacement and
83 cooling of these felsic intrusions (e.g. Baker and Lang 2001; Gloaguen et al. 2014). However, little is
84 known about the 3D geometry of the magmatic plumbing system at depth, although it is well-known
85 that pluton morphology can have a complex effect on the formation of related mineralization (e.g.
86 White 1981; Carten et al. 1988; Seedorff 1988; Wallace 1991; Guillou-Frottier et al. 2000;
87 Guillou-Frottier and Burov 2003; Eldursi et al. 2009; Gloaguen et al. 2014), with the general
88 mineralization morphology and location being dependent on the morphology/shape of the associated
89 intrusions. Knowledge of the 3D shape of the intrusion is thus crucial in the exploration for
90 mineralized apices or cupolas. In addition, exploration models can be greatly improved if a detailed
91 history of mineralization processes can be obtained. However, little is known about detailed timing of
92 ore deposition during the emplacement-crystallization-cooling cycles of an intrusive. In particular,

93 during magma emplacement, crystallization, and cooling, several phases of advective processes (e.g.
 94 localized convection cells, hot fluid channels) may disturb local isotherms and thus mineral closure
 95 temperatures, which makes obtained mineralization ages very difficult to interpret relative to magmatic
 96 cooling ages.



97
 98 *Figure 1 Main geological features of Morocco and location of the Tighza area in the Variscan*
 99 *basement of the eastern Moroccan Meseta (modified after Arboleya et al. 2004).*

100
 101 The Tighza mining district is located in the Moroccan Meseta which is the largest outcrop
 102 region of the Variscan Belt rocks in Morocco (Figure 1). It produces 52% of the Pb, 26% of the Ag,
 103 and 3.2% of the Zn, of the national production of Morocco at the end of the last century (Wadjinny
 104 1998). The annual production is about 25600 concentrated tones of Pb-Ag, and 4300 concentrated

105 tones of Zn in 2013 (CMT website⁽²⁾). The Tighza district contains three main granitic stocks, with
106 other scattered small granitic plugs and dykes (Figure 2). The three main stocks are, from north to
107 south, the Mispickel Granite, Mine granite, and the Kaolin granite. Because these granitic bodies share
108 similar compositional, textural, and mineralogical features, many authors assume that they are part of
109 one single large granitic body at depth to (e.g. Agard et al. 1958; Marcoux et al. 2015).

110 In addition to other economic resources, such as Pb-Zn vein deposits, the area hosts W-Au
111 mineralization with features typical of an IRGD system (Marcoux et al. 2015). This W-Au
112 mineralization and its magmatic context have been extensively studied, particularly dealing with
113 geochronological constraints (Cheilletz and Zimmermann, 1982; Cheilletz, 1984, Watanabe, 2002;
114 Nerci, 2006; Marcoux et al., 2015; Rossi et al., 2016). Even though a general IRGD model is accepted,
115 geochronological data from granitic rocks have led authors to suggest various genetic models for this
116 W-Au mineralization.

117 Two major points will be addressed in this contribution: i) the role of 3D pluton geometry in
118 the mineralization process and localization through 3D gravity data modeling, and ii) the timing of the
119 mineralization, through 3D coupled hydro-thermal numerical simulations. The combination of well-
120 exposed rocks and numerous geochronological data relevant to magma emplacement and associated
121 hydrothermal activities makes the Tighza mining district and its IRGD a particularly suitable candidate
122 for addressing these points.

123

124 **2. The Tighza W-Au mineralization: previous models**

125 **2.1 Geological setting**

126 The Tighza region (or Jebel Aouam region) is located in the Paleozoic Moroccan Meseta,

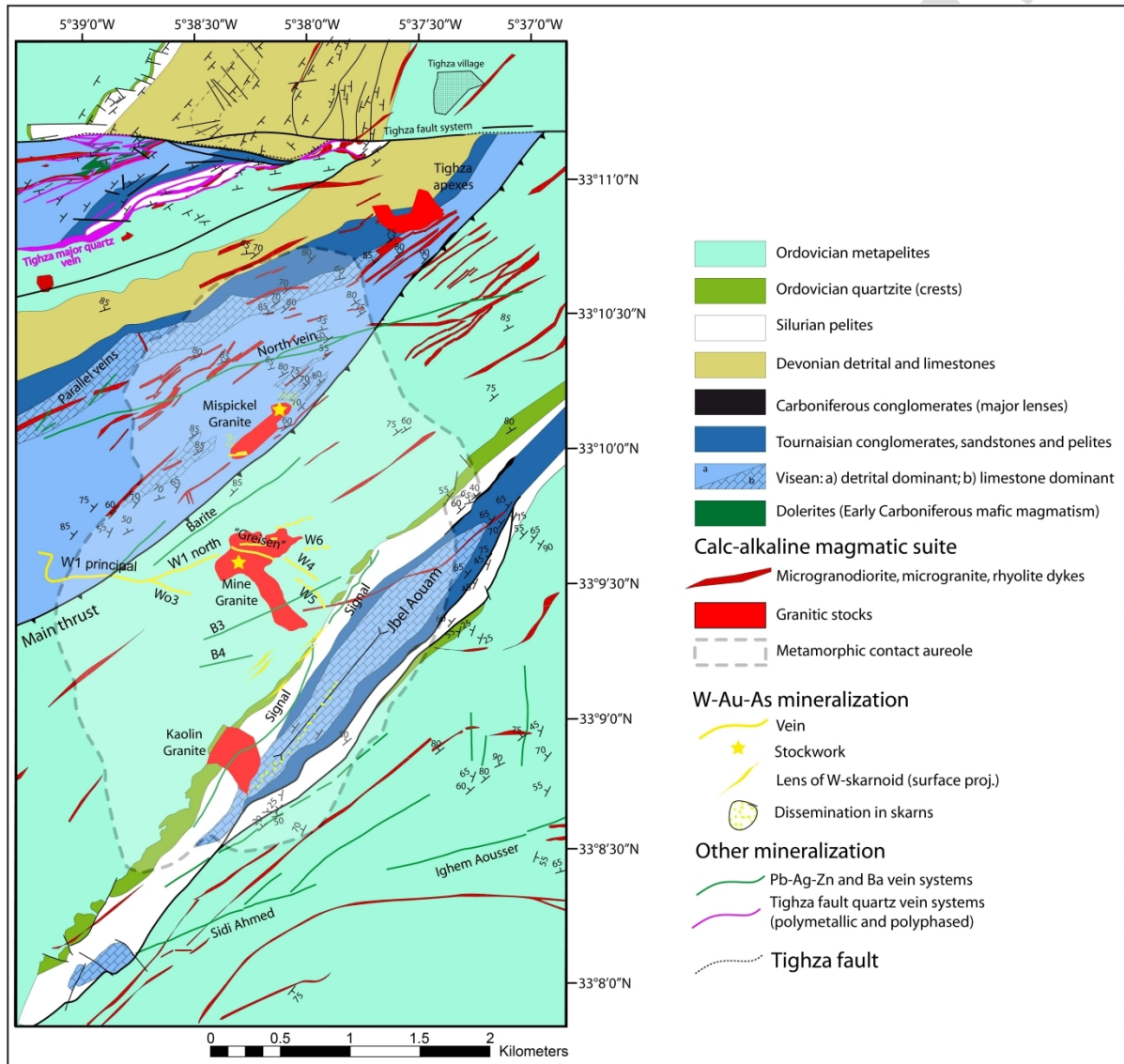
⁽²⁾ <http://www.miniere-touissit.com/cmt-tighza.html>

127 in the eastern part of central Morocco, near the border with the Middle Atlas. It is around thirty
128 kilometers north-west of the town of Khenifra and 7 km west of the city of Mrirt (Figure 1). The region
129 is an area of high to moderate relief (1220 m mean altitude) in the center of which Jbel Aouam (1496
130 m) emerges. The study area (Figure 2) constitutes a low-grade metamorphic Paleozoic (Ordovician to
131 Viséan) series that has been intruded by numerous stocks and dykes. The southeastern portion of the
132 region corresponds to an allochthonous unit that was thrust over a northwestern autochthonous unit
133 (Termier 1936; Ribeyrolles 1972).

134 After the D1 gravity-driven nappe event (late Viséan-Early Namurian, Faïk, 1988), the entire
135 area underwent NE-SW cylindrical folding, resulting from NW-SE horizontal shortening, during the
136 Variscan D2 event (Late Carboniferous - Early Permian, Destreucq, 1974; Marcoux et al., 2015 and
137 references therein). As a result, both bedding planes and weakly developed axial plane cleavages are
138 NE-trending with a dominant steep dip to the NW. The E-W-trending Tighza wrench fault system
139 (Figure 2) is a branch of the regional-scale Aguelmous-Mrirt shear zone (Cheilletz, 1983, 1984). In the
140 study area, it separates a northern from a southern domain. This strike-slip fault has a complex
141 polyphase history. It was first a transtensional right-lateral wrench zone during D2, while the last
142 movements correspond to left-lateral wrenching during a D3 brittle event (Destreucq 1974; Cheilletz
143 1983, 1984; Marcoux et al. 2015) of Late Permian-Middle Triassic in age (Rossi et al., 2016, 2017).

144 With the exception of early Carboniferous dolerite sills emplaced during sediment deposition,
145 all magmatic rocks are late to post- Variscan in age. Most of the intrusives belong to an early calc-
146 alkaline magmatic suite and correspond to: i) the three main biotite granite stocks (Mispickel, Mine
147 and Kaolin granites), as well as a small granodioritic stock located in the northeastern part of the study
148 area (Figure 2), and (ii) a well-developed dyke swarm, generally parallel to S0/S1, that are mainly
149 composed of microgranodiorite, microgranite, and aplite (Figure 2). Based on previous studies
150 (Cheilletz 1984; Cheilletz and Isnard 1985; Nerci 2006; Marcoux et al. 2015; Tarrieu, 2014; Rossi et

151 al. 2016), the four granitic/granodioritic stocks are homogeneous in texture, mineralogy, and
 152 geochemistry, suggesting that they are probably derived from a single magmatic source. Other
 153 intrusives rocks in the region (not represented in Figure 2) are tholeiitic microdioritic and sparse
 154 leucogranitic dykes.



155
 156 *Figure 2 Geological map of the Tighza region compiled from CMT maps and the current field study.*
 157 *The extent of the contact metamorphic aureole is from Cheilletz and Iznard (1985). Only silicic dykes*
 158 *of the calc-alkaline magmatic suite have been represented. W-Au-As mineralization types are from*
 159 *Marcoux et al. (2015).*

160 The Tighza region is a polymetallic district containing: i) Pb-Zn-Ag mineralization in thick
161 veins that make the Tighza district the second largest Pb producer in Morocco (Wadjinny 1998;
162 Compagnie Minière de Touissit (“CMT”) - Pers. Comm. 2012), ii) W-Au granite-related
163 mineralization (IRGDs), and iii) Ba-Sb vein mineralization. There is a consensus that the Pb-Zn-Ag
164 mineralization formed in epithermal conditions during a late Permian-Triassic magmatic-hydrothermal
165 event disconnected from an older W-Au mineralization (Tarrieu 2014; Cheilletz et al. 2015; Marcoux
166 et al. 2015; Rossi et al. 2016; 2017). Pb-Zn-Ag mineralization developed during the D3 left-lateral
167 wrench event that had a transtensional component. A low-temperature Ba-Sb mineralization in veins is
168 also present in the district (“Barite”, B3, B4, and within the Tighza fault quartz vein system; Figure 2).
169 Although no strong evidence has been observed in the Tighza area, based on other neighboring
170 mineralization, the Ba-Sb hydrothermal system is assumed to be older than the Pb-Zn-Ag mineralizing
171 event (Agard et al. 1958).
172 For additional detailed geological data and descriptions concerning the geological setting, refer to
173 Marcoux et al. (2015) and references therein.

174 **2.2 The Tighza W-Au IRGD mineralization: a magmatic-hydrothermal model**

175 The Tighza W-Au mineralization occurs in various styles, such as veins around the Mine granite (“W”
176 in Figure 2), greisen bands crosscutting the Mine granite, stockworks, disseminations in skarns, and
177 lenses of W-rich skarnoids (Figure 2). Three main hydrothermal stages have been proposed (Nerci
178 2006; Marcoux et al. 2015): (i) a primary W-rich skarn stage formed under lithostatic fluid pressure
179 (ca 1.5 kbar) and temperatures up to 580°C, and associated with intrusion of calc-alkaline suite granitic
180 magma, (ii) a second stage responsible for the deposition of Au-As-Bi-Te, and (iii) a third stage that
181 includes base metal (Zn-Cu) mineralization. The three stages occurred during the emplacement-
182 cooling cycle of granitic intrusive activity. The pressure conditions associated with stages 2 and 3 are
183 close to the hydrostatic regime (ca. 550 bars) with temperature around 300°C. Stages 1 and 2 are

184 characterized by CO₂-rich and saline aqueous fluids expelled from the granitic magma. Structurally,
185 the emplacements of the granitic intrusions and the mineralized veins were controlled and coeval with
186 the D2 right-lateral transtensional deformation. The W-Au mineralization is spatially associated with
187 the main granitic intrusions (Figure 2), however, W-Au mineralization is also found far from the four
188 outcropping granitic stocks (Figure 2). This spatial relationship suggests that the mineralizing system
189 is not restricted only to the close vicinity of outcropping stocks and dykes. In addition, Cheillettz and
190 Isnard (1985) identified a hydrothermal metamorphic aureole, centered on the Mine/Mispickel/Kaolin
191 granitic stocks, that encompasses the W-Au occurrences (dashed contour in Figure 2). As the extent of
192 this metamorphic aureole is large relative to the outcrop extent of the granitic stocks and dykes, the
193 presence of a larger plutonic body, hidden at depth, was proposed. This hypothesis has been recently
194 re-proposed by Marcoux et al. (2015) and Rossi et al. (2016, 2017). Because of this relationship to the
195 igneous intrusive bodies, igneous geochemistry, and the origin/nature of the associated fluids, a model
196 of Intrusion-Related Gold Deposits (IRGD) has been proposed for the Tighza W-Au mineralization
197 (Marcoux et al. 2015). Subtle distinctions led Rossi et al. (2016, 2017) to classify this W-Au
198 mineralization as “porphyry-style”, a deposit group that includes and/or overlaps the IRGD definition
199 according to different authors and classifications.

200 The main difference between the various Tighza W-Au metallogenic models is the age of the
201 granitic magmatism associated with the mineralization. K/Ar and Ar/Ar dating on biotite and
202 muscovite from the W-Au hydrothermal materials yields a coherent and unique age of ca. 285 Ma
203 (Cheillettz 1984; Nerci 2006, Marcoux et al. 2015). Only one Ar/Ar age (291.8 ± 0.3 Ma), from a
204 sample of phlogopite associated with quartz from the W1 north veins (Figure 2), suggests the
205 possibility of an earlier phase of high-temperature hydrothermal deposition (Marcoux et al. 2015). In
206 this sample, no ore mineral is directly associated with the phlogopite, however, the quartz and
207 phlogopite assemblage is the classical gangue material to the Tighza W ore. This suggests that the

208 earlier high-temperature hydrothermal event indicated by this age date might not be barren. The
209 significance of this older age and the timing of this hydrothermal event will be discussed later in the
210 paper.

211 In contrast, radiometric ages obtained for the granitic magmatism are more widely distributed.
212 The K/Ar and Ar/Ar ages of igneous micas from the Mine, Kaolin, and Mispickel granitic stocks
213 cluster around 286 Ma (Cheilletz 1984; Watanabe 2002; Nerci 2006; Marcoux et al. 2015), but U-
214 Th/Pb zircon intercept ages yield 309 ± 10 Ma for the Mine granite and 295 ± 9 Ma for the Kaolin
215 granite (Rossi et al. 2016). Based on these ages, Rossi et al (2016, 2017) proposed that the outcropping
216 granitic stocks of the Tighza area (Figure 2) were emplaced during an old magmatic event (between
217 320 and 300 Ma), and thus they could not be responsible for the W-Au hydrothermal event and are
218 disconnected from the IRGD genetic model. This hydrothermal event would then be related to a
219 younger hidden laccolith emplaced between 300 and 280 Ma with a large associated hydrothermal
220 contact metamorphic halo. The Rossi et al. (2016, 2017) model thus implies the non-coeval intrusion
221 of at least two plutonic complexes at depth. Alternatively, Marcoux et al. (2015) proposed that the W-
222 Au mineralization is genetically related to the emplacement of the outcropping granitic stocks, with the
223 stocks being interpreted to be plutonic apices (cupolas) related to a single underlying batholith (e.g. the
224 Yerrington batholith in Nevada, USA, which has several apices that are separated by ~1-2 km; Dilles
225 and Profett 1995).

226 **3. Coupled gravity and geological study**

227 Modeling of gravity data has long provided information valuable in defining pluton shapes at depth,
228 pluton thickness variations, and locations of possible root zones. In addition, gravity and 3D geological
229 models aid in understanding pluton emplacement mechanisms, regional kinematics, and strain fields
230 (e.g. Vigneresse 1990; Lyons et al. 1996; Martelet et al. 2004; Talbot et al. 2004; Joly et al. 2008; Wei
231 et al. 2016).

232 **3.1 Gravity data acquisition**

233 A detailed gravity survey of the Tighza region was acquired using a grid with minimum average
234 sampling spacing of about 250 m. The survey was supported by: i) a detailed (1:5,000), digital, and
235 georeferenced mining company (CMT) topographic map, ii) a 30 m Digital Elevation Model (DEM)
236 used in areas without available topographic mapping, and iii) a detailed geological map from the CMT
237 improved by new field data collected during this study (Figure 2), and iv) field structural
238 measurements (bedding, faults, fold axes, etc.) as point data in a GIS database. Gravity data from 320
239 gravity stations were obtained using a Scintrex Model-CG3-M gravimeter. Station altitudes were
240 determined through plotting GPS station location data on the detailed topographic map and the DEM.
241 The gravity data were fully processed to produce the complete Bouguer anomaly, following reduction
242 of gravity earth-tides, instrumental drift, and free air, plateau, and terrain corrections (using three high,
243 intermediate, and low resolution DEMs, following the method of Martelet et al. 2002). We used the
244 Hayford (1930) reference ellipsoid and the Bouguer reduction density was chosen to be 2600 kg/m^3 ; a
245 value close to the average density of the rocks outcropping in the area. Given the rather gentle
246 topography in the study area, the error due to the terrain corrections is estimated to not exceed
247 0.2 mGal. Taking into account all other sources of error, in particular the vertical positioning accuracy,
248 the overall accuracy of the dataset is estimated to be around $\pm 1 \text{ mGal}$. Finally, the data were
249 interpolated onto a 100 m grid using a standard minimum curvature algorithm. As the new gravity
250 survey was not tied to the existing Moroccan gravity station network, these new data cannot be merged
251 into the regional gravity database. At the scale of the study area, no regional trend in the gravity data
252 was observed, therefore the data interpretation and modeling was based on the complete Bouguer
253 anomaly.

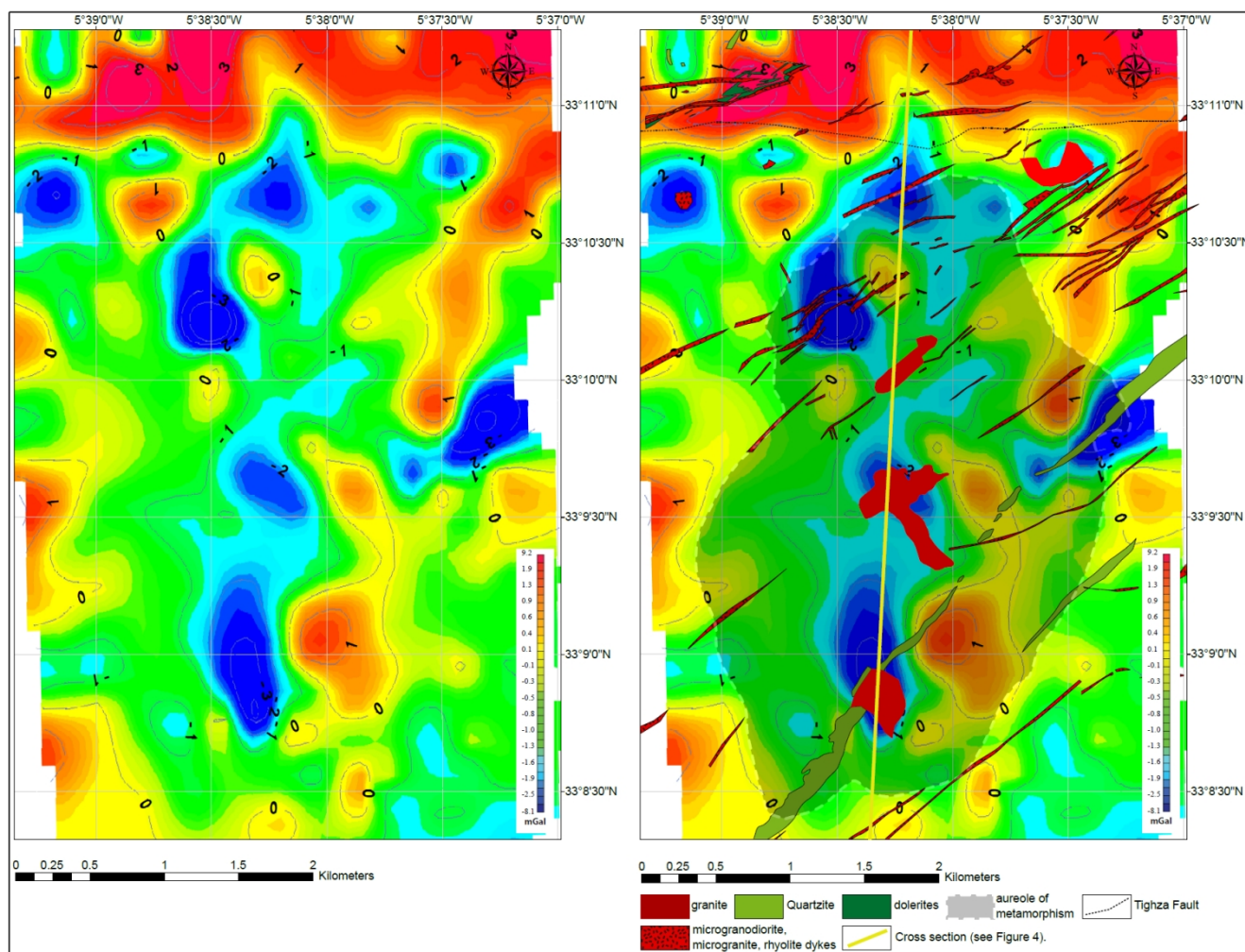
254 **3.2 Gravity data results and interpretation**

255 The Bouguer gravity data are presented in Figure 3. Three different ranges of Bouguer anomalies are

256 present in the region: i) strong negative values (dark blue colours) that correspond, in part, to the
257 outcropping granitic stocks and granitic dyke swarms, while others are likely related to thicker zones
258 of a probable underlying pluton; ii) strong positive values (red colours) that correspond to high-
259 density rocks, which are normally associated with the host Paleozoic sediments; and iii) moderate
260 anomalies (light blue to yellow colours) that are not easily related to a specific outcropping geology.

261 In the northern part of the study area, the Tighza fault system is clearly recognizable in the
262 gravity data as a sharp E-W-trending gravity contrast corresponding to the contact between a lower-
263 density southern domain and a higher-density northern domain (Figure 3). The southern domain is
264 characterized by numerous and voluminous outcropping lower-density granitic rocks and a large
265 associated contact metamorphic thermal aureole, while very few granitic rocks are found within the
266 northern domain.

267 From these gravity data, it is reasonable to suggest that a large-scale plutonic complex (referred
268 to herein as the “Tighza pluton”) is present at depth, south of the Tighza fault system. Because the
269 Tighza fault system was active during successive D2 and D3 events, it is difficult to identify a
270 potential fault control during the emplacement of the Tighza pluton. However, as the W-series
271 structures controlled the emplacement of the Mine granite apex during D2 (Figure 2; Marcoux et al.
272 2015), it cannot be ruled out that the Tighza fault may have controlled spatial emplacement of the
273 proposed pluton. In this case, a major feeder zone could be located in the vicinity of the fault, with the
274 magma ascending along the fault and moving southward to form a laccolith (as proposed by Rossi et
275 al. 2017 for their “hidden pluton”). In support of this hypothesis, several gravity lows demarcate the
276 southern side of Tighza fault. In addition, other local strong negative Bouguer anomalies suggest other
277 possible roots or apices, in particular in the eastern and southern parts of the study area. In contrast, the
278 western region is associated with slightly positive anomalies, suggesting few or no granitic intrusives
279 in this area.



280

281 *Figure 3. Left: Bouguer anomaly map of the Tighza region. Right: Same map with draping of the main*
 282 *geological features. The Tighza fault system corresponds to a strong contrast between a positive*
 283 *anomaly in the northern domain and a southern domain displaying low to strongly negative*
 284 *anomalies.*

285

286 3.3 2D Gravity modeling

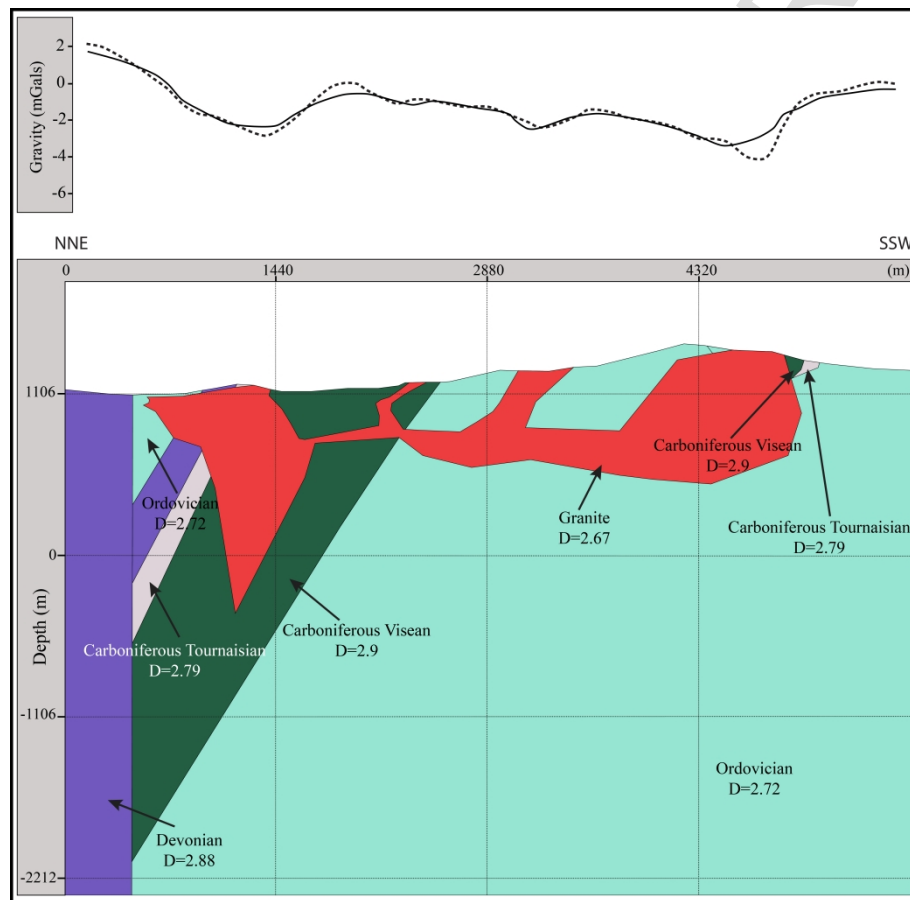
287 Mapping of the gravity data alone does not provide accurate constraints on the geometry and
 288 thickness of the proposed underlying Tighza pluton. Localized gravity lows suggest the presence of
 289 thicker granitic bodies and indistinct low-intermediate values that occupy a large part of the area could
 290 attest to the presence of granitic material at depth, south of the Tighza fault. Without field indications

291 of these suggested features, the hypothesis to be tested was that the localized gravity lows, parts of
292 which correspond to outcropping granitic stocks, could be apices of a batholith lying beneath them
293 (Figure 3). Noteworthy, a multi-phase magmatic history as suggested by Rossi et al. (2016, 2017)
294 cannot be easily depicted by gravity modeling in first order approximation, since gravity data cannot
295 solve time-dependence of magmatic history. Deciphering such complex multi-stage magmatism by
296 gravity modeling would be enhanced by (i) high density contrast between the different magmatic
297 injections; (ii) non-telescoped plutonic complexes with a sufficient thickness of host rocks in between.
298 Since both conditions are not clearly established, the gravity modeling might be inefficient in imaging
299 such complex magmatic scenario. Several 2D cross-sections were modeled to construct the pluton
300 geometry at depth, using the Geosoft GM-SYS software. A cross-section oriented in the NNE-SSW
301 direction, cross-cutting the three outcropping granitic stocks (Mispickel, Mine, and Kaolin granites;
302 Figure 3) is presented in Figure 4. The 2D modeling was constrained by: 1) surface geological and
303 structural data, 2) geometrical constraints derived from indirect interpretation of the measured gravity
304 data, and 3) densities of the lithological units. The subsurface extension of the granitic stocks at depth
305 (geometries) and surrounding rocks were adjusted until the gravity effect of the model reproduced the
306 profile of the field-measured Bouguer gravity data. Despite the non-uniqueness of this solution, the
307 modeled geometry is consistent with the available gravity data. The obtained geometry (Figure 4)
308 suggests that the granitic stocks merge at a shallow depth into a single plutonic body with a thickness
309 ranging between 0.3 and 1 km.

310 **3.4 Coupled gravimetry and geology 3D model**

311 The coupled gravimetry and geology 3D model was built in a georeferenced system and takes into
312 account: i) the digital elevation model (DEM), (ii) surface geological mapping, and (iii) 2D

313 geophysical cross-sections. We used 3D GeoModeller⁽³⁾ modeling software to integrate the available
 314 information, such as geophysical cross-sections and geological observations (lithological and structural
 315 measurements), into a common 3D geometrical framework. All these data were jointly used to produce
 316 a 3D geostatistical interpolation with resulting 3D geological geometries (Lajaunie et al. 1997;
 317 Calcagno et al. 2002; Calcagno et al. 2008).
 318



319
 320 *Figure 4 Bottom: 2D gravity model of the proposed Tighza pluton; top: gravity effect of the model*
 321 *(solid line) fitted to the measured Bouguer anomaly (dotted line). D is the rock density used for*
 322 *modeling.*

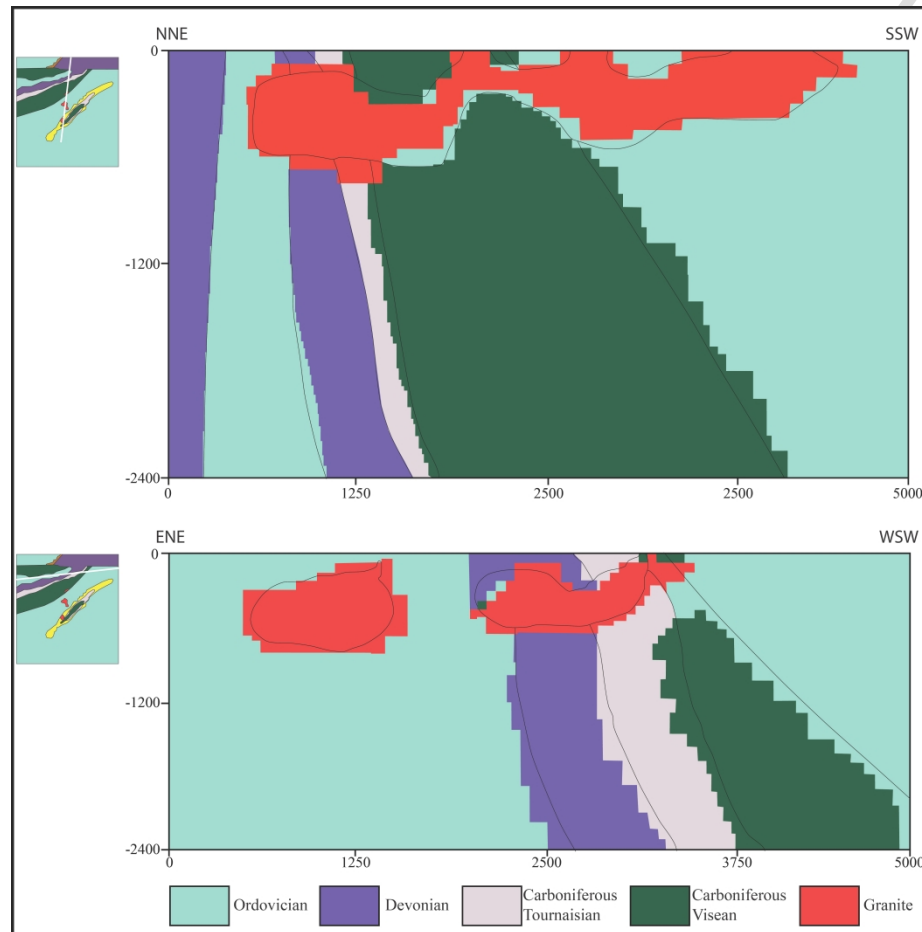
⁽³⁾ commercial software developed by BRGM and Intrepid Geophysics (Calcagno et al. 2008; Guillen et al. 2008); <http://www.geomodeller.com>

323 The initial 3D model of the proposed Tighza pluton was constructed using a maximum depth
324 around 1 km and one face of the model was fixed at the major Tighza fault. These constraints were
325 based mainly on the 2D cross-sections (e.g. Figure 4). Another constraint was the assumption that the
326 intruded body is thicker at field-identified apices (e.g. Mispikel, Mine, and Kaolin granites) and in
327 other places where the Bouguer gravity map displays strongly negative values (Figures 3, 4). The
328 initial geometry was interpolated in 3D based on eight cross-sections, as well as the surface geology
329 and topography. Intersections between cross-sections, geological boundaries, and structural features
330 were extrapolated with apparent dip. In order to make the 3D model consistent with all of the 2D
331 information, the 3D model was slightly adjusted manually where necessary, taking into account field
332 geological observations and using 2D balancing techniques. When a realistic 3D model was achieved,
333 satisfying all available data/information, it was discretized into a 3D matrix of voxels, prior to further
334 refinement in a 3D gravity inversion. This last stage of 3D discretization is necessary to ensure that all
335 complex 3D effects of the geological features are correctly taken into account and that the 3D
336 interpolation has not produced any unexpected deformation of these features.

337 In the stochastic 3D gravity inversion, implemented in 3D GeoModeller, the initial 3D geometrical
338 model is randomly modified voxel by voxel over millions of iterations (Guillen et al. 2008), with the
339 gravity effect being recomputed at each step and compared to the measured gravity map. The Monte
340 Carlo exploration of model space ensures that the “data minus model” misfit decreases during the
341 course of the iterations until an acceptable difference between the model effect and the measured data
342 is attained. For the Tighza gravity inversion, the threshold of misfit was fixed at 1 mGal, which is the
343 estimated accuracy of the Bouguer anomaly data.

344 This means that all the 3D models tested during the inversion process that had a Root Mean Square
345 (RMS) misfit lower than 1 mGal were considered as “valid 3D geometrical gravity models”. 2 million
346 iterations were run and more than 1.1 million models were thus defined as “valid”. Then, these 1.1

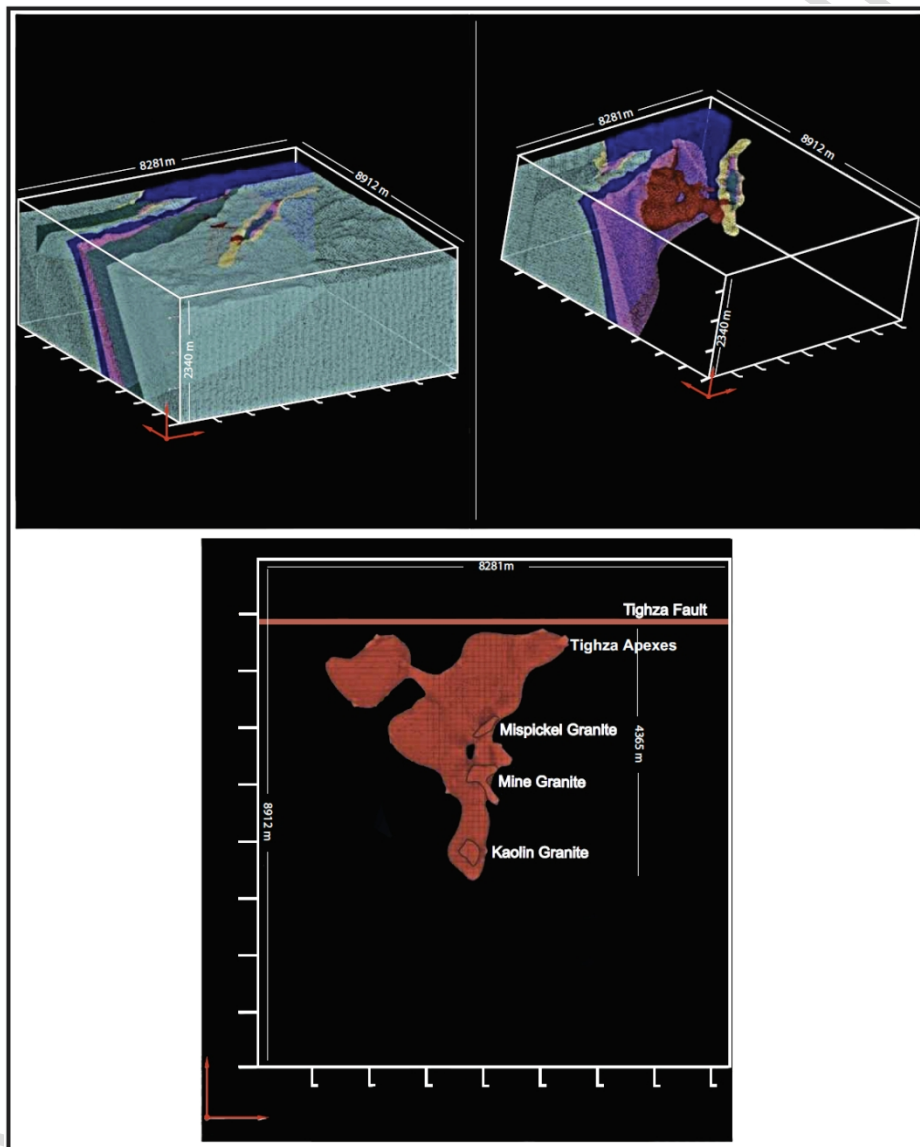
347 million models were statistically analyzed to obtain the “most probable density model” which is
 348 inferred by allocating to each voxel the density with the maximum probability computed from all the
 349 “valid density models”.



350
 351 *Figure 5 N-S (a) and NE-SW (b) cross-sections displaying the result of the 3D gravity inversion in*
 352 *terms of the most probable lithology. The black solid lines represent the starting geological model.*

353
 354 The same can be done regarding lithology. This is illustrated in Figure 5 along two cross-
 355 sections for which each pixel of the cross-sections was queried in terms of most probable lithology.
 356 These two cross-sections show that the most probable locations of the geological bodies, as derived
 357 from the gravity inversion, is globally close to the 3D geological starting model (represented by black
 358 lines in Figure 5). This confirms the overall good quality of the 2D gravity models used to constrain

359 the initial 3D geological model. Thus, by construction, the geometries of the 3D inverted model are in
 360 good accordance with the gravity data (with a RMS misfit lower than 1 mGal over the entire study
 361 area). Figure 6 displays the final 3D model of the Tighza area, including the resulting gravity
 362 inversion.
 363



364
 365 *Figure 6 At the top, two snapshots of the final 3D model of the area based on geology and gravity*
 366 *combined inversion. At the bottom, the most probable 3D architecture of the Tighza pluton in plan*
 367 *view.*

368 4. 3D numerical hydrothermal modeling

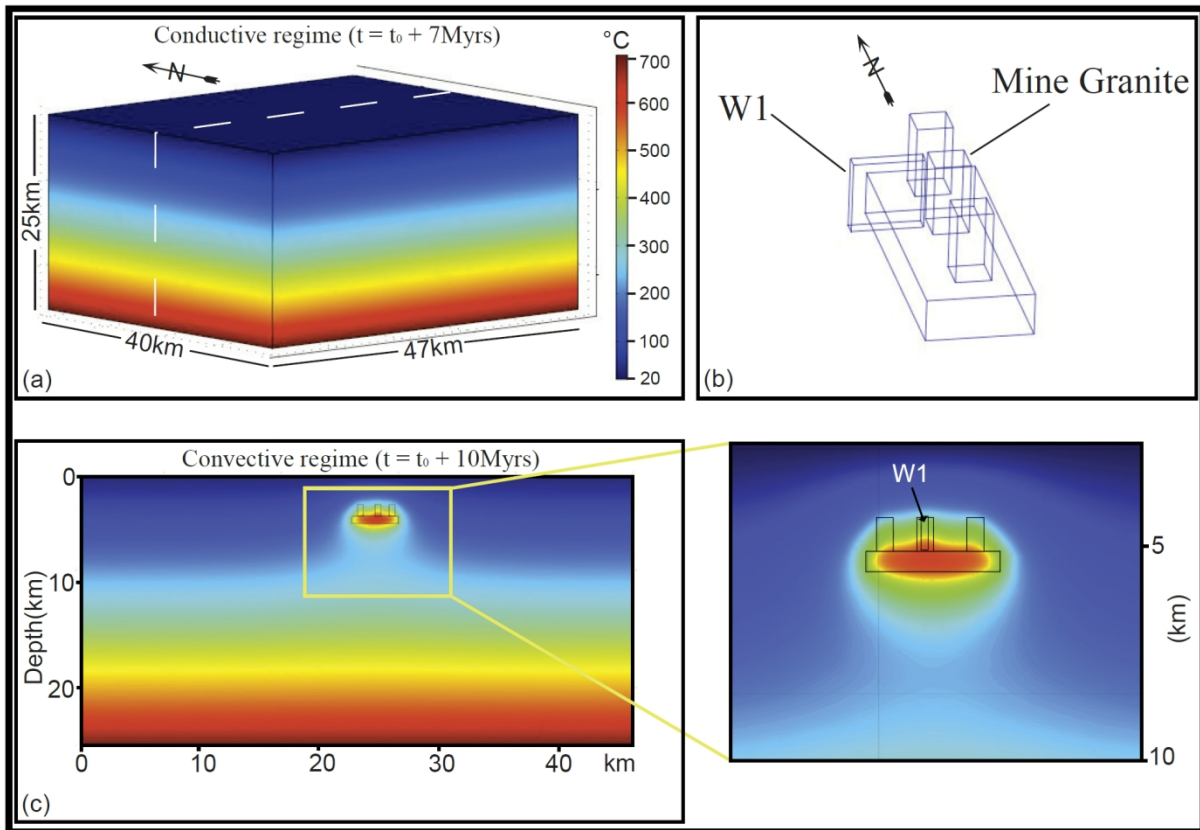
369 4.1 Hydrothermal coupling and numerical modeling construction

370 In a previous study (Eldursi et al. 2009), 2D numerical modeling of fluid circulation around granitic
371 bodies was performed and applied to various natural cases, some of them including the presence of
372 plutonic apices or proximity to a low-angle normal fault. Here, a similar numerical modeling is
373 focused on the proposed Tighza plutonic system.

374 Firstly, for this numerical investigation, we chose the best 3D geometry obtained from gravity
375 modeling: three apices connected at depth to a laccolithic intrusion as represented in the Tighza IRGD
376 model of Marcoux et al. (2015). However, the 3D intrusion geometry was simplified because using a
377 complex 3D geometry for the pluton would not be useful, as thermal diffusion tends to smooth out
378 small-scale asperities. The simplified geometry chosen was a box-like shape with dimensions of
379 $4000 \times 1200 \times 500 \text{ m}^3$, corresponding to a large-scale batholith, surmounted by three apices with unique
380 dimensions of $500 \times 500 \times 1000 \text{ m}^3$ (Figure 7). The main granitic stocks were emplaced at a depth of
381 circa 5500 m, based on fluid inclusion analyses (see model of Marcoux et al. 2015, with lithostatic
382 fluid pressure during the stage “W”).

383 Following this simplification approach, only the W1 structure was chosen to be included in the
384 model due to its economic importance (Figure 2). It was represented as a permeable fracture with a
385 vertical box-shape (200 m wide and 1000 m long) located next to the middle apex that represents the
386 Mine granite (Figure 7). Moreover, to simulate the effects of permeable fracture zones during W-Au
387 hydrothermal event, the “W” veins network (Figure 2) were considered to be important and were
388 simplified as a high-permeability zone in the west-east direction. These simplifications aim at
389 enhancing first-order controlling physical processes associated with both the presence of plutonic
390 apices and of a permeable zone (e.g. fractures/fault).

391



392

393 *Figure 7 Simplified 3D hydrothermal model of the Tighza region: a) the thermal steady-state of the*
 394 *whole model at time t_0+7 Myrs, b) the simplified 3D model geometry of the Tighza pluton and the W1*
 395 *structure, c) enlarged view of the hottest phase of emplacement (t_0+10 Myrs.). Colored scale is in ($^{\circ}$ C),*
 396 *from 20 to 750 $^{\circ}$ C.*

397

398 As for the previous 2D models of Eldursi et al. (2009), heat transfer in a porous medium was
 399 coupled with Darcy law, and mass conservation is applied. The basic numerical scheme was tested and
 400 validated by reproduction of different benchmark tests (Gerdes et al. 1998; Rabinowicz et al. 1998;
 401 McKenna and Blackwell 2004). Permeability was considered to be depth-dependent, and fluid density
 402 and viscosity were considered to be temperature-dependent. All governing equations, physical laws,
 403 boundary conditions, and initial thermal and fluid flow regimes are detailed in the Appendix.

404 The 3D hydrothermal modelling of the proposed Tighza pluton includes three phases, 1) a first

405 phase (from 0 to 7 Myrs) designed to reach thermal equilibrium (thermal steady-state before triggering
 406 the second phase); 2) a second phase during which the pluton temperature was increased regularly up
 407 to its emplacement temperature (700°C) for a duration of 3 Myrs (7 to 10 Myrs; Figure 8; details
 408 presented in Eldursi et al. 2009), and 3) a third phase from 10 to 20 Myrs, during which the granitic
 409 body was considered as a fully crystallized felsic body, cooling with time, and where petrophysical
 410 parameters (such as thermal conductivity, heat capacity, and density) varied continuously with
 411 temperature (Eldursi et al. 2009). The fault permeability followed a depth-dependent function during
 412 the first phase (see Appendix), and was set to 10^{-15} m² during the second and third phases. The host
 413 rock permeability followed a depth-dependent function during the three phases.

414 **4.2 Rock alteration index**

415 During the computation of the numerical model, temperature and velocity fields are recorded,
 416 and a Rock Alteration Index (RAI) is computed as the scalar product of fluid velocity and temperature
 417 gradient:

$$418 \quad RAI = \vec{u} \cdot \vec{\nabla} T \quad (1)$$

419 RAI is a physical index defined by Phillips (1991) and used in several studies (e.g. Zhao et al. 1998;
 420 Raffensberger and Vlassopoulos 1999; Harcouët-Menou et al. 2009). It is designed to represent a local
 421 cooling rate such that when the RAI value is negative, mineral precipitation occurs. When the RAI
 422 value is positive, i.e. when the fluid flows towards higher temperatures, mineral dissolution occurs. To
 423 account for the highest cooling rates, negative values can only be retained with fluid velocity
 424 exceeding 10^{-10} m.s⁻¹. To do this, a binary scalar, the “restricted-RAI” (referred to herein as R²AI), is
 425 set to 1. For other negative RAI values (i.e. with a velocity lower than 10^{-10} m.s⁻¹) and for positive
 426 values of the RAI, R²AI is set to 0. Consequently, patterns of R²AI illustrate at a given time the
 427 location of the most probable zones of mineralization. The use of this “restricted-Rock Alteration
 428 Index R²AI” in a previous study on 2D numerical modeling of fluid circulation around granitic bodies

429 (Eldursi et al. 2009) led to the prediction of favorable mineralized zones for different natural cases.

430 **4.3 Results of mineralization patterns**

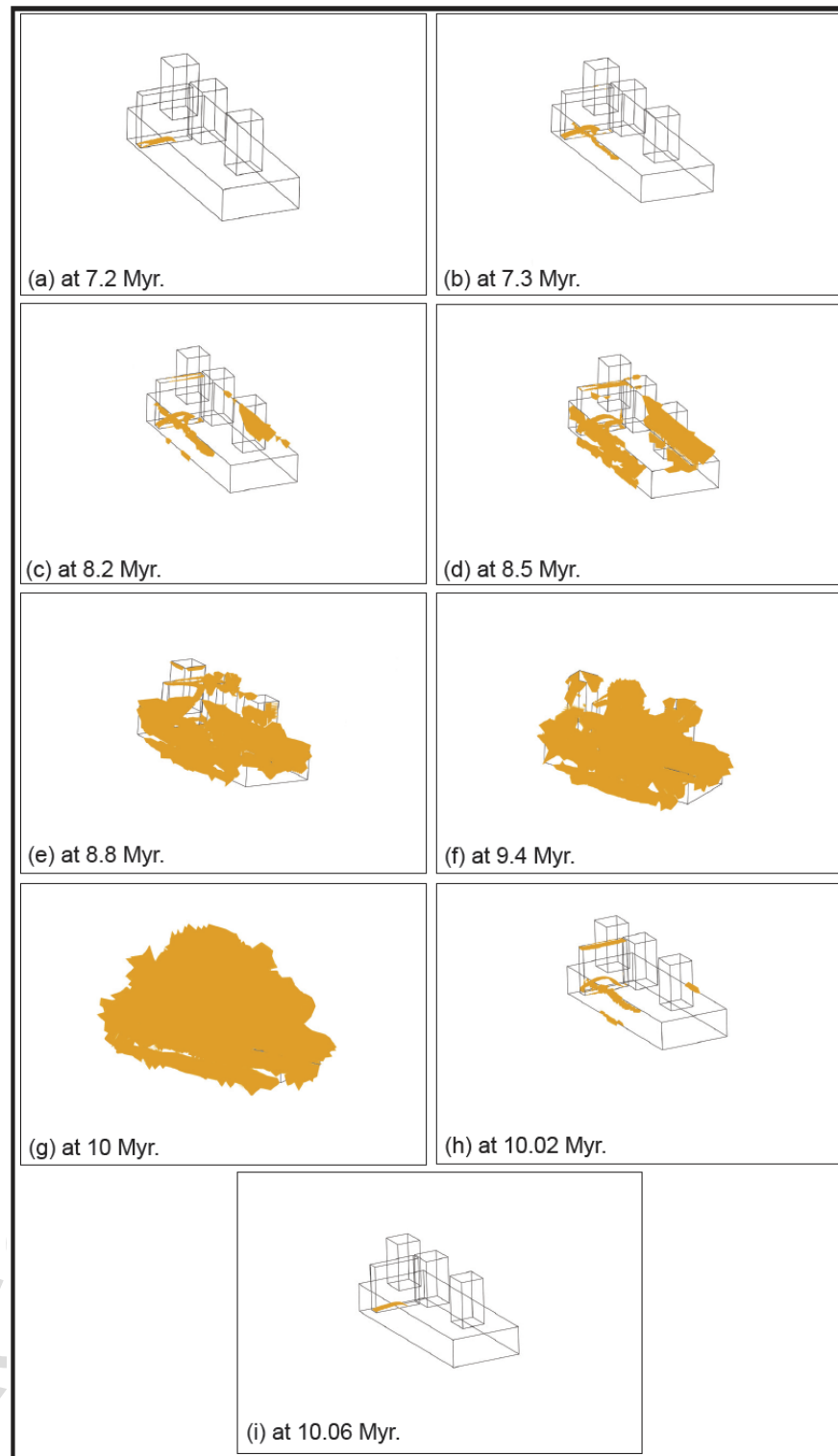
431 The history of probable zones of mineralization (PZM) obtained by applying the R^2AI factor during
432 the emplacement phase (Phase 2; 7-10 Myrs) and the cooling phase (Phase 3; 10-20 Myrs) is
433 displayed in Figure 8:

- 434 1) The first appearance of a PZM occurred, at the bottom of W1, during the first 0.2 Myr. of
435 emplacement (7.2 Myrs; Figure 8a);
- 436 2) After 0.3 Myr of emplacement, the PZM extends laterally along the margin of the magmatic
437 body and upward into the permeable fault (7.3 Myrs; Figure 8b);
- 438 3) The PZM continues to cover all of the margins of the magmatic body and rises up into the
439 permeable zone (W1 structure) after 1.8 Myrs (8.8 Myrs; Figures 8c-8e).
- 440 4) As the temperature increases due to magma emplacement, the PZM extends and covers the
441 top of the middle apex after 2.4 Myrs (9.4 Myrs; Figure 8f).
- 442 5) At the hottest phase of emplacement, the PZM covers almost the entire intruded body and
443 fills up the permeable zone (10 Myrs; Figure 8g).
- 444 6) During the cooling phase, the PZM concentrates along the western and eastern margins of
445 the intruded body, and inside the permeable zone during the first 0.02 Myr of cooling (10.02
446 Myrs; Figure 8h).
- 447 7) After 0.06 Myr of cooling, the PZM is restricted to the bottom part of the permeable zone
448 only (10.06 Myrs; Figure 8i).

449 Noteworthy, the use of R^2AI coupled with W-Au temperature deposition range determined in
450 Marcoux et al. (2015) has been tested in order to optimize the PZM: as a result, the numerical
451 PZM remain unchanged comparatively to ones obtained with $R2AI$ only. Using these results,
452 the numerically modeled time for where the PMZ best matches the observed mineralization

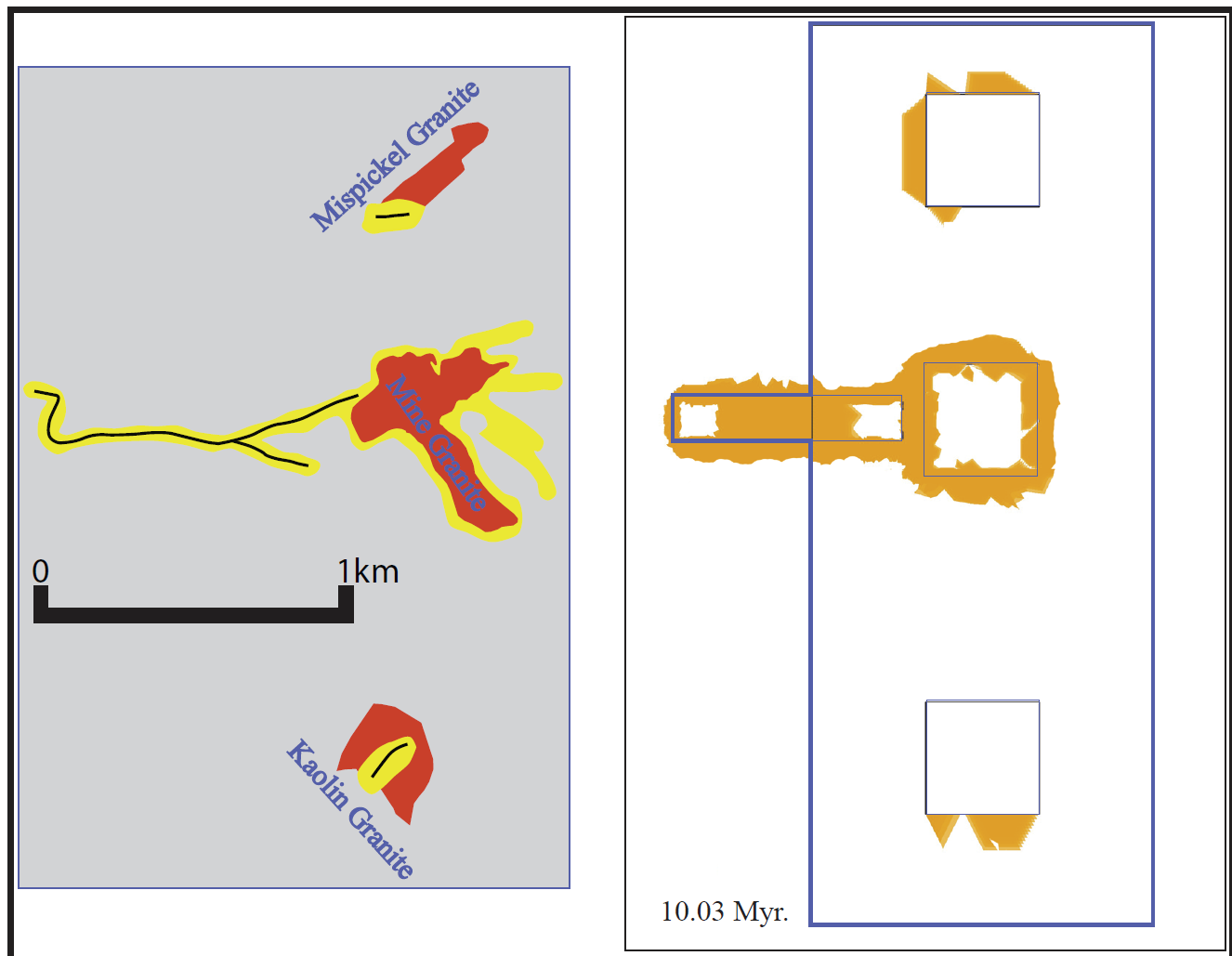
453

pattern at Tighza was determined.



454

455 *Figure 8 Evolution of R^2AI during the emplacement and cooling phases, the colored zone shows the*456 *most probable zones for mineralization (Eldursi et al. 2009).*



457

458 *Figure 9 Plan view of mineralized areas (yellow) as observed in the field (left), and (orange) inferred*
 459 *from the numerical model at time $t_0+10.03$ Myrs (right, at a depth of 4700 m).*

460 Figure 9 presents in plan-view a map of the simplified distribution pattern of the Tighza W-Au
 461 mineralization against numerically-inferred PZM determined for 0.03 Myr after the hottest phase
 462 (10.03 Myrs) at a depth of 4700 km. At this time, the, PMZ best matches the observed Tighza W-Au
 463 distribution pattern. Thus, the observed mineralization pattern, which is characteristic of the “Main”
 464 type of mineralization, occurs in the model shortly after the hottest phase of pluton emplacement (10
 465 Myrs, when the pluton produces the maximum heat energy), but it does not continue for long time, as
 466 the PZM disappears after 0.06 Myr of cooling (10.06 Myrs; Figure 8i).

467 Thus, there is a probability of “Pre-Main” mineralization that is modeled from 2.8 Myrs before

468 the hottest phase and extends more and more up to the hottest part of the emplacement at 10.0 Myrs.
469 When the pluton has completely cooled, the thermal regime returns to equilibrium as it was before the
470 emplacement and no evidence of favorable physical conditions for mineralization are observed in this
471 period.

472 **5. Discussion**

473 **5.1 Coupling 3D geological modeling and gravimetry of the Tighza pluton**

474 To build a realistic 3D model of proposed Tighza pluton, geological field data like structural
475 orientations, petrophysical measurements of different geological units, and geophysical data (gravity)
476 were integrated. The sharp contact between the positive Bouguer gravity anomalies in the north and
477 the moderate to low negative anomalies to the south reflects essentially the contact between high-
478 density host rocks and low-density intrusives, respectively, that exactly traces the Tighza fault. This
479 relationship is compatible with a model in which the proposed Tighza syntectonic pluton was rooted in
480 the Tighza fault and spread toward the south as a horizontal sheet during D2 at which time the Tighza
481 fault acted as a right-lateral oblique slip normal fault. However, to account for the D3 event, which
482 may have cut off the northern edge of the laccolith, additional arguments must be provided to confirm
483 the syntectonic nature of this pluton.

484 The results of the gravity data interpretation resulted in the definition of three different
485 categories: 1) high positive anomalies that correspond to the material composed mainly of
486 metasedimentary mica schist, 2) low negative anomalies that correspond to granitic stocks and dykes,
487 and 3) moderate negative anomalies that were modeled as a sub-horizontal magmatic sheet located
488 beneath the granitic stocks. These geophysical and geological features suggest the presence of a single
489 batholith-shaped pluton at depth associated with granitic cupolas/apices at surface. As expected, the
490 best 3D solution obtained does not allow us to depict a deeper intrusion separated from a shallower
491 intrusion by host rock panels as drawn in Rossi et al. (2017). However, this does not mean that this

492 two-intrusions scenario is wrong (e.g. telescoping etc...). Noteworthy, a single intrusion at depth does
493 not preclude a long-lived pluton construction with various “pulses” of magma inputs at different times.

494 **5.2 Timing of mineralization events and thermo-chronological constraints**

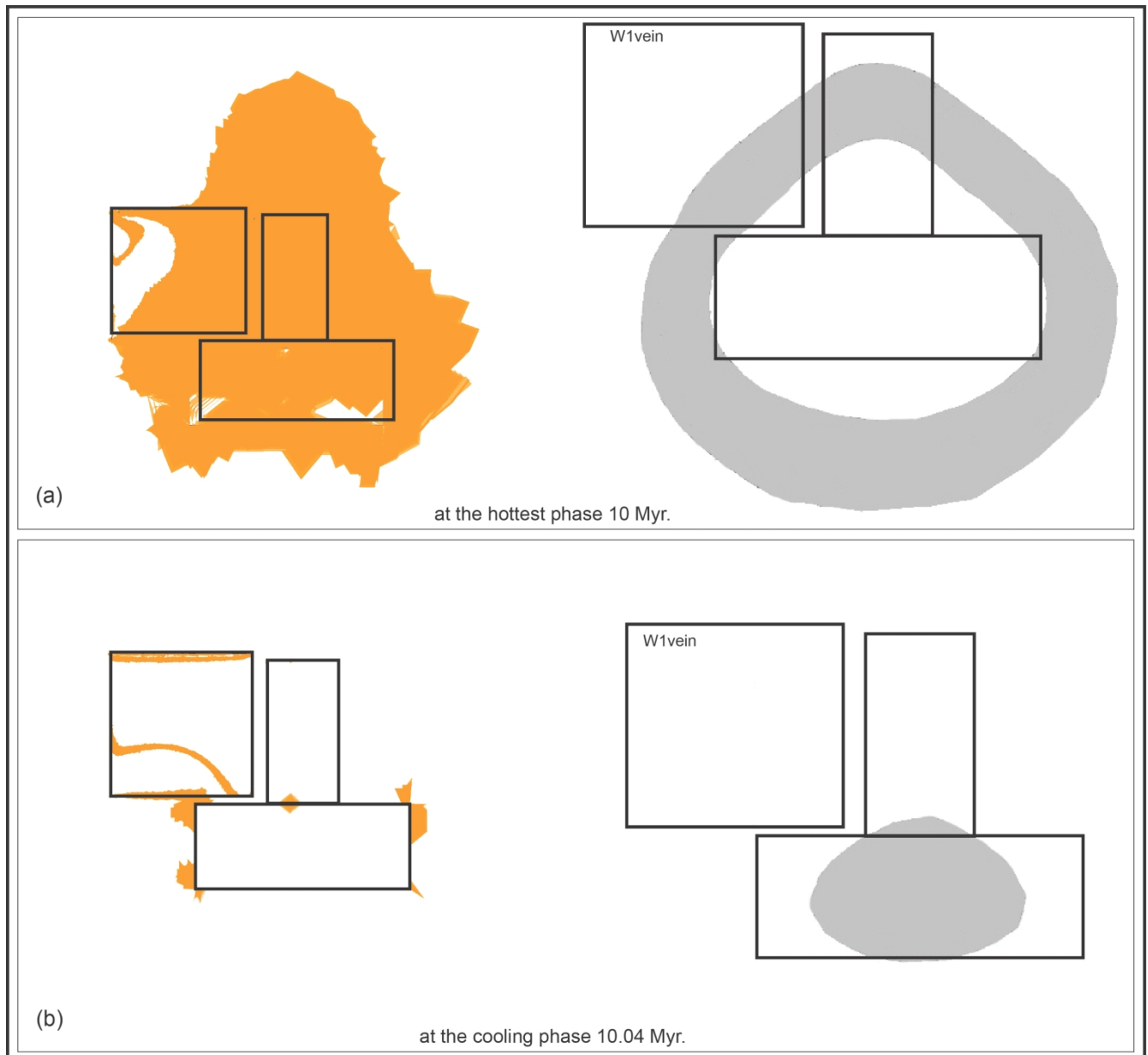
495 The results of the modeling allow apparent radiometric age discrepancies in such magmatic-
496 hydrothermal systems to be highlighted. Indeed, the presence of an early high-temperature
497 hydrothermal event (EHE, potentially economic; see above) has been demonstrated and dated at 291
498 ± 0.3 Ma on phlogopite (Marcoux et al. 2015, see their initial stage on Figure 15A). This EHE
499 developed within the W1 north vein close to the Mine granite which yields very homogenous Ar/Ar
500 ages on micas of around 286 Ma. Because the EHE and the major hydrothermal event leading to the
501 “Main” type W-Au deposition are both clearly related to magmatic activity, the presence of a
502 difference for the two mineralization types in the ages obtained using the same radiometric method is
503 consistent with the modeling results. The 3D hydrothermal models provide a possible answer for this
504 difference. Indeed, in the numerical model we assumed a magmatic emplacement period lasting 3
505 Myrs (from 7 to 10 Myrs; with the ultimate hottest phase at 10 Myrs) that is reasonable for a long-
506 living emplacement period (e.g. Annen et al. 2006). At Tighza, this hypothesis is coherent with the
507 long-lasting 295-280 Ma magmatic-hydrothermal event responsible for the W-Au mineralization and
508 related (hidden) granite intrusion proposed by Rossi et al. (2016, 2017). The application of R²AI in the
509 Tighza 3D model shows that the favorable physical conditions were detected during two periods
510 (Figure 8). The first period corresponds to a long-lived period of “Pre-Main” type mineralization
511 during which a shallow pluton (less than 8 km depth) is able to create PZM during the emplacement
512 period. This “Pre-Main” type mineralization might then correspond to the EHE observed in the W1
513 north vein system (292 Ma stage in Figure 15 of Marcoux et al., 2015). The second period corresponds
514 to a shorter-lived period of “Main” type mineralization, gathering the stages I (W), II(Au-As-Bi-Te)
515 and II(Zn-Cu) of Marcoux et al. (2015), and relates to the initiation of the cooling period of the

516 shallow pluton (Figures 8 and 9).

517 To investigate this time-lag between these two periods of hydrothermal activities; during the
518 emplacement phase of intrusion and the initial cooling phase, the isotopic closure temperatures
519 muscovite and biotite were mapped in the 3D numerical model (Figure 10). The values of 350° and
520 450°C were used for the closure temperatures of muscovite and biotite, respectively (Spear 1993; Villa
521 1998). This temperature range is shown by grey regions in Figure 10. The mapping results show
522 clearly that the W1 structure, which hosts PZM during early pluton emplacement (Figure 8) is colder
523 than the Ar^{40}/Ar^{39} closure temperature of biotite at the time of the hottest phase (i.e. 10 Myrs; Figure
524 10) while the temperature in the granite core is still higher than the closure temperature range. The
525 same temperature range crosses the granite after 0.04 Myr of cooling (Figure 10). Thus, due to
526 advective heat transfer in permeable zones, a time-lag is observed between the emplacement of the
527 intrusives and related proximal high permeability zones (i.e. W1 structure). This result is in accordance
528 with theoretical models of Eldursi et al. (2009), in which the presence of PZM before the hottest phase
529 of the pluton was demonstrated.

530 In the 3D hydrothermal modeling, this time-lag (ca. 2.8 Myrs) is shorter than the one observed
531 in the field, with the Tighza natural case showing a time-lag of around 5 Myrs, (i.e. 291 Ma to 286 Ma).
532 From the modeling work, it appears that with a longer pluton emplacement time, the time-lag (gap)
533 between “Pre-Main” and “Main” types of mineralization increases because the intruded magma
534 creates warm conditions in which the fronts of closure temperature of muscovite and biotite need a
535 longer time to pass through the W1 vein and then the parent granite. Consequently, the time-lag
536 between the mineralization formed in W1 and the parent granite becomes longer. Although the
537 numerical modeling did not take the chemical and deformation contributions into account, the results
538 show that the time-lag between mineralization and the parent intrusion can be produced by favorable
539 physical conditions for mineralization before the cooling period (crystallization phase). This time-lag

540 could be as long as the period of magma emplacement.



541
 542 *Figure 10 Biotite-Muscovite closure temperatures (the grey zone) cross the permeable zone at the*
 543 *hottest phase (a), and cross the intruded body after 0.04 Myr of cooling (b), the orange zone indicates*
 544 *the R^2AI zone.*

545 **6. Conclusion**

546 Based on the geological, petrophysical, and geophysical data, 3D geometrical modeling of the Tighza
 547 region suggests that the granitic stocks of the area are connected in the subsurface to a plutonic mass of

548 batholith-like shape that was rooted in the north at the E-W Tighza fault and that spread toward the
549 south during a D1 extensional regime. The available geophysical data and gravity inversion presented
550 in this work does not allow us to depict the scenario of two hidden plutonic bodies in the subsurface
551 proposed by Cheilltez et al. (2015) and Rossi et al. (2017).

552 The 3D numerical modeling of thermally-driven fluid flow demonstrated that: 1) suitable
553 physical conditions for mineralization can take place before the crystallization of the parent granite. 2)
554 regardless of whether the W-Au mineralization belongs to “Pre-Main” or “Main” types, the time-lag
555 between early mineralization and young parent magma intrusion could be established to be as long as
556 several hundreds of thousands to millions of years depending on the period of emplacement and/or
557 cooling rate.

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565 of the manuscript.

567 **Appendix**

568 All governing equations, physical laws, boundary conditions, and initial thermal and fluid flow
569 regimes are the same as those published in Eldursi et al. (2009).

570 **1. The physical model**

571 The model geometry for the host rock was a box-like shape with dimensions of 47x40x25 km (Figure
572 7). The geometry of the Tighza pluton was simplified and drawn based on the final geometrical model
573 produced according to the geological, petrophysical, and geophysical data. The W1 fracture/fault was
574 simulated by a box with dimensions of 0.2x1x1 km. The permeability of the W1 fault follows a depth-
575 dependent equation during the first phase, and is fixed at 10^{-15} m² during the second and third phases.
576 The permeability of the host rock follows the Manning and Ingebritsen (1999) depth-permeability
577 curve during the three phases:

$$578 \quad \log(K) = -14 - 3.2 \log(Z) \quad (1)$$

579 The permeability of the pluton for the first two phases (steady-state and emplacement of magma)
580 follows Manning and Ingebritsen (1999) depth-dependent equation, and is fixed at 10^{-24} m² during the
581 third phase (cooling phase).

582

583 **2. The governing equations and parameters**

584 The thermally-driven fluid flow model for Tighza is governed by different partial differential
585 equations such as Darcy's law, and mass and energy conservations, expressed as, respectively:

$$586 \quad \vec{u} = -\frac{k}{\mu} (\vec{\nabla} p - \rho_L \vec{g}) \quad (2)$$

$$587 \quad \frac{\partial(\Phi \rho_L)}{\partial t} = -\nabla \cdot (\rho_L \vec{u}) \quad (3)$$

$$588 \quad Q = C_{eq} \left(\frac{\partial T}{\partial t} \right) + \nabla \cdot (-\lambda_{eq} \vec{\nabla} T) + C_L \vec{u} \vec{\nabla} T \quad (4)$$

589 The fluid density depends on temperature by the linear approximation.

$$590 \quad \rho_L = \rho_0 (1 - \alpha_v (T - T_0)) \quad (5)$$

591 The fluid dynamic viscosity also varies with temperature (Rabinowicz et al.,1998)

$$592 \quad \mu(T) = 2.414 \cdot 10^{-5} \times \exp\left(\frac{247.8}{T - 140}\right) \quad (6)$$

593 where T is in K and μ is in Pa s.

594 The thermal conductivity, heat capacity, and densities of the host rock and intrusives were also varied
595 with temperature during the three phases, for more details see Eldursi (2009) and Eldursi et al. (2009).

596

597 **3. The boundary and initial conditions**

598 For the thermal part of the modeling, the bottom and upper boundaries were fixed at 700 and 20°C,
599 respectively, whereas the vertical boundaries were set as “thermally insulating”. The intruded body
600 was set as a time-dependent heat source at 500 mW.m⁻³ in order to reach 750°C during 3 Ma of
601 emplacement. Then the heat source was turned off when the magma temperature reached 700°C. For
602 the hydraulic part of the modeling, the bottom and vertical boundaries were set as impermeable
603 boundaries; whereas the upper boundary was set as a permeable boundary at a constant pressure of 10⁵
604 Pa.. The initial temperature and pressure were governed by the following equations:

$$605 \quad T = 20 + 0.024 * z \quad (7)$$

606 where 0.024 is the geothermal gradient (in °C/m),

$$607 \quad P = P_0 + \rho_L * g * z \quad (A8)$$

608 where P₀ is the atmospheric pressure.

609

610

611

612

613

614

615

616 *Table (1): units and symbols of the used parameters and variables, after Eldursi et al. (2009).*

Parameter	Unit	Symbol
Fluid velocity	m s^{-1}	u
Permeability	m^2	K
Dynamic viscosity	Pa s	μ
Fluid density	Kg m^{-3}	ρ_L
Fluid pressure	Pa	p
Gravitational acceleration	m s^{-2}	g
Depth	m	z
	km	Z
Volumetric coefficient of thermal expansion	K^{-1}	α_v
Weighted average volumetric heat capacity	$\text{J m}^{-3} \text{K}^{-1}$	C_{eq}
Specific heat capacity	$\text{J kg}^{-1} \text{K}^{-1}$	C_p
Thermal conductivity	$\text{W m}^{-1} \text{K}^{-1}$	λ
Equivalent thermal conductivity	$\text{W m}^{-1} \text{K}^{-1}$	λ_{eq}
General heat source	W m^{-3}	Q
Volumetric heat capacity of moving fluid	$\text{J m}^{-3} \text{K}^{-1}$	C_L
Temperature	K	T
Porosity, (5% in pluton, host rock and permeable zone)	-	Φ

617

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ACCEPTED MANUSCRIPT

Highlights

- Gravity data processing and inversion of Tighza area.
- 3D hydrothermal fluid flow occurred before, during, and after magma emplacement.
- The rock alteration index is useful for detection of the sites of ore deposition.
- “Pre-Main” and “Main” types of ore deposits are associated with magma emplacement.
- Time-lag issues associated with Intrusion-Related Gold Deposits are interpreted.