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13 Mantle wedge exhumation beneath the Dora-Maira (U)HP dome
14 unravelled by local earthquake tomography (Western Alps)

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24 **Abstract**

25 In continental subduction zones, the behaviour of the mantle wedge during exhumation of
26 (ultra)high-pressure [(U)HP] rocks provides a key to distinguish among competing exhumation
27 mechanisms. However, in spite of the relevant implications for understanding orogenic evolution, a
28 high-resolution image of the mantle wedge beneath the Western Alps is still lacking. In order to fill
29 this gap, we perform a detailed analysis of the velocity structure of the Alpine belt beneath the
30 Dora-Maira (U)HP dome, based on local earthquake tomography independently validated by
31 receiver function analysis. Our results point to a composite structure of the mantle wedge above the
32 subducted European lithosphere. We found that the Dora-Maira (U)HP dome lays directly above
33 partly serpentinized peridotites ($V_p \sim 7.5$ km/s; $V_p/V_s = 1.70-1.72$), documented from ~ 10 km
34 depth down to the top of the eclogitized lower crust of the European plate. These serpentinized
35 peridotites, possibly formed by fluid release from the subducting European slab to the Alpine
36 mantle wedge, are juxtaposed against dry mantle peridotites of the Adriatic upper plate along an
37 active fault rooted in the lithospheric mantle. We propose that serpentinized mantle-wedge
38 peridotites were exhumed at shallow crustal levels during late Eocene transtensional tectonics, also

39 triggering the rapid exhumation of (U)HP rocks, and were subsequently indented under the Alpine
40 metamorphic wedge in the early Oligocene. Our findings suggest that mantle-wedge exhumation
41 may represent a major feature of the deep structure of exhumed continental subduction zones. The
42 deep orogenic levels here imaged by seismic tomography may be exposed today in older (U)HP
43 belts, where mantle-wedge serpentinites are commonly associated with coesite-bearing continental
44 metamorphic rocks.

45 **Keywords:** continental subduction; ultra-high-pressure metamorphism; mantle wedge exhumation;
46 peridotite serpentinization; local earthquake tomography; Western Alps

47 **Highlights:**

- 48 - High-resolution image of the seismic velocity structure of the Alpine mantle wedge
- 49 - First geophysical evidence of mantle-wedge exhumation during continental subduction
- 50 - Mantle wedge exhumation is favoured by upper plate divergent motion

51 **1. Introduction**

52 Exhumed (ultra)high-pressure [(U)HP] rocks bear compelling evidence of the interaction
53 between subducting plates and the overlying mantle wedge (Carswell and Compagnoni, 2003;
54 Hacker et al., 2006; Ferrando et al., 2009; Scambelluri et al., 2010; Deschamps et al., 2013; Gilotti,
55 2013). However, the role played by the mantle wedge during (U)HP rock exhumation is still poorly
56 understood. Some numerical models point to a negligible mantle involvement during exhumation
57 (Yamato et al., 2008; Butler et al., 2013), whereas other models suggest that mantle rocks may be
58 strongly involved, and may follow the exhumation path of buoyant (U)HP rocks towards the Earth's
59 surface (Schwartz et al., 2001; Petersen and Buck, 2015). The behaviour of the mantle wedge
60 during (U)HP rock exhumation may thus provide a key to discriminate among competing
61 exhumation models (e.g., Agard et al. 2009; Guillot et al., 2009a; Liou et al., 2009; Warren, 2013).

62 In the Cenozoic metamorphic belt of the Western Alps, the geologic record of subduction and
63 exhumation is exceptionally well preserved (e.g., Lardeaux et al., 2006; Malusà et al., 2011), but a
64 high-resolution image of the mantle wedge is still lacking. A detailed analysis of the seismic
65 velocity structure beneath the Dora-Maira (U)HP dome, where coesite attesting deep continental
66 subduction was first described three decades ago (Chopin, 1984), may thus provide new insights on
67 the ongoing debate concerning the mechanisms triggering the exhumation of (U)HP rocks (e.g.,
68 Jolivet et al., 2003; Schwartz et al., 2001; Agard et al., 2009; Little et al., 2011; Butler et al., 2013;
69 Malusà et al., 2015; Ducea, 2016). Moreover, this kind of analysis may provide new interpretation
70 keys to understand the field relationships between mantle-wedge rocks and continental (U)HP rocks
71 in deeply unroofed pre-Cenozoic orogenic belts (e.g., van Roermund, 2009; Scambelluri et al.,
72 2010), where the geophysical record of subduction and exhumation is no longer preserved (e.g.,
73 Zhao et al. 2017).

74 In this article, we exploit a comprehensive seismic dataset, also including anomalously deep
75 earthquakes (Eva et al., 2015), to derive a local earthquake tomography model of the mantle wedge
76 beneath the Dora-Maira (U)HP dome, which is then compared with the results provided by receiver
77 function analysis along the CIFALPS transect (China-Italy-France Alps seismic survey; Zhao et al.,
78 2015). Our results indicate that part of the mantle wedge was metasomatized above the Alpine
79 subduction zone, and subsequently exhumed at shallow depth beneath continental (U)HP rocks now
80 exposed at the surface. This suggests that mantle-wedge exhumation may be a prominent feature of
81 the deep structure of many (U)HP belts, which should be integrated in future theoretical models of
82 continental subduction and (U)HP rock exhumation.

83 **2. Tectonic framework**

84 ***2.1 The orogenic wedge of the southern Western Alps***

85 The Western Alps are the result of oblique subduction of the Alpine Tethys under the Adriatic
86 microplate since the Late Cretaceous, followed by continental collision between the Adriatic and

87 European paleomargins during the Cenozoic (Coward and Dietrich, 1989; Dewey et al., 1989;
88 Lardeaux et al., 2006; Handy et al. 2010; Malusà et al., 2016a). The resulting slab structure is still
89 largely preserved (Zhao et al., 2016a), as well as the orogenic wedge formed atop the European slab
90 (Lardeaux et al., 2006; Beltrando et al., 2010; Malusà et al., 2011). In the southern Western Alps,
91 along the CIFALPS transect (X-X' in Fig. 1), the Alpine orogenic wedge mainly consists of rocks
92 derived from the Piedmont ocean-continent transition and from the adjoining European paleomargin
93 (Lemoine et al., 1986; Dumont et al., 2012). The external zone, exposed to the west of the Frontal
94 Pennine Fault (FPF in Fig. 1), includes the Pelvoux and Argentera basements and their deformed
95 Meso-Cenozoic sedimentary cover sequences (Ford et al., 2006), which record a transition from
96 thin-skinned to thick-skinned compressional tectonics during the Neogene (Schwartz et al., 2017).
97 East of the Frontal Pennine Fault, in the Alpine metamorphic wedge, the Briançonnais nappe stack
98 (Br in Fig. 1) mainly consists of Upper Paleozoic to Mesozoic metasediments and underlying pre-
99 Alpine basement rocks that underwent subduction starting from the Paleocene, and were later
100 exhumed in the Eocene - early Oligocene (Malusà et al., 2002, 2005a; Ganne et al., 2007; Lanari et
101 al., 2014). The Briançonnais nappe stack forms the core of the present-day Alpine fan-shaped
102 structure (Michard et al., 2004) that was overprinted by a dense network of extensional faults during
103 the Neogene (Sue et al., 2007; Malusà et al., 2009). The eastern part of the fan is formed by oceanic
104 metasediments of the Schistes lustrés complex (SL in Fig. 1; Lemoine et al., 1986; Lagabrielle and
105 Cannat, 1990), including boudinaged decametre-to-kilometre-sized ophiolitic bodies that were
106 deformed and metamorphosed during Alpine subduction under blueschist to transitional blueschist-
107 eclogite facies conditions (Agard et al., 2002; Tricart and Schwartz, 2006; Schwartz et al., 2009)
108 (Fig. 2A). A ductile normal fault (DF1 in Fig. 2A; Ballèvre et al., 1990) separates the Schistes
109 lustrés complex from the Viso metaophiolites (Vi in Fig. 1; Lombardo et al., 1978; Angiboust et al.,
110 2012), representing major imbricated remnants of the Tethyan oceanic lithosphere that were
111 deformed and metamorphosed under eclogite facies conditions during the Eocene (Duchêne et al.,
112 1997; Schwartz et al., 2000; Rubatto and Hermann, 2003). Another ductile normal fault (DF2 in

113 Fig. 2A; Blake and Jayko, 1990) separates the Viso eclogites from the underlying stack of deeply
114 subducted continental basement slices referred to as the Dora-Maira (U)HP dome (DM in Fig. 1;
115 Henry et al., 1993; Michard et al., 1993), which also includes the coesite-bearing Brossasco-Isasca
116 eclogitic unit (black star in Figs. 1 and 2A; Chopin et al., 1991; Compagnoni and Rolfo, 2003).
117 Along the boundary with the Po Plain, the CIFALPS transect crosses the southern tip of the Lanzo
118 massif (La in Fig. 1; Boudier, 1978; Piccardo et al., 2007), an eclogitized mantle slice separated
119 from the Dora-Maira dome by a near-vertical active fault system rooted in the upper mantle (Rivoli-
120 Marene deep fault - RMF in Fig. 1) at the southward prolongation of the Insubric Fault (Eva et al.
121 2015; Malusà et al., 2017). The Lanzo massif consists of slightly serpentinized spinel plagioclase
122 peridotites surrounded by a 3–5 km thick envelope of foliated serpentinites (Müntener et al., 2004;
123 Debret et al., 2013), and records a high-pressure metamorphic peak of early Eocene age (Rubatto et
124 al., 2008). Beneath the Po Plain, the complex transition zone between the Adriatic upper plate and
125 the Apennines, also involving rotated fragments of the Alpine orogenic wedge (Maffione et al.,
126 2008; Eva et al. 2015), is mainly covered by thick Cenozoic to Quaternary sedimentary successions.

127 ***2.2 The Dora-Maira (U)HP dome***

128 The Dora-Maira (U)HP dome is exposed all along the internal side of the southern Western
129 Alps (Chopin et al., 1991; Lardeaux et al., 2006) (Fig. 1). To the west of Torino, it is juxtaposed
130 against the Lanzo massif along the Lis-Trana deformation zone (Perrone et al., 2010), possibly
131 representing a shallow splay of the Rivoli-Marene deep fault (Eva et al., 2015). To the south, it is
132 partly buried by the sedimentary successions of the Po Plain (Fig. 1), and is exposed as a half-dome
133 including coesite-bearing eclogitic rocks (Brossasco-Isasca unit) sandwiched between quartz-
134 eclogite facies rocks, above, and blueschist facies rocks, below (Compagnoni et al. 1995; Avigad et
135 al., 2003; Compagnoni and Rolfo, 2003) (Fig. 2A). The Brossasco-Isasca unit is a coherent
136 continental crust sliver composed of granitic gneisses (Lenze and Stöckhert, 2007), whiteschists
137 (Chen et al., 2017), mafic eclogites (Groppo et al., 2007) and impure marbles (Ferrando et al.,

138 2017). It was subducted to depths greater than ~100 km by the late Eocene (e.g., Chopin et al.,
139 1991; Rubatto and Hermann, 2001; Hermann, 2003), and was exhumed close to the Earth's surface
140 by the early Oligocene, at rates faster than subduction rates (Rubatto and Hermann, 2001; Malusà et
141 al., 2015), as confirmed by low-temperature thermochronology data (Gebauer et al., 1997; Tricart et
142 al., 2007; Beucher et al., 2012). The overlying quartz-eclogite Venasca *p.p.* and Dronero units,
143 including gneisses and metasediments derived from a Permian-Triassic detrital sequence, and the
144 underlying blueschist-facies Sanfront-Pinerolo unit, consisting of orthogneisses and metasediments
145 intruded by Permian diorites (Avigad et al., 2003), were piled up together with the Brossasco-Isasca
146 and Viso units during late Eocene exhumation (Schwartz et al., 2009; Malusà et al., 2011), to
147 become part of the Eocene Eclogite belt now exposed along the upper-plate side of the Western
148 Alps (Fig. 1), at the rear of a lower-pressure Paleogene wedge (LP in Fig. 2B,C).

149 The structure and lithologic composition of the orogenic wedge beneath the Dora-Maira
150 (U)HP dome is still largely unknown. The velocity structure provided by available seismic
151 tomography models is well resolved only for the uppermost 15-20 km (Paul et al., 2001; Béthoux
152 et al., 2007). Recent tectonic reconstructions postulated the occurrence of Briançonnais crust
153 slivers down to depths greater than 30 km, and suggested that these slivers would be involved in
154 an east-vergent backfold at the scale of the whole Eclogite belt (Schmid et al., 2017). However,
155 the Dora-Maira dome shows no cartographic evidence of such large-scale backfolding, which is
156 instead observed in the Monte Rosa dome (MR in Fig. 1) of the northern Western Alps, where
157 late backfolding is possibly ascribed to progressive westward shifting of Adria indentation from
158 the Central Alps to the northern Western Alps during the Neogene (Malusà et al., 2016b). As a
159 matter of fact, alternative interpretations of the deep tectonic structure of the southern Western
160 Alps are not adequately supported by geophysical data. This information gap has so far precluded
161 a full understanding of the exhumation mechanisms that were active within the Alpine subduction
162 zone during the late Eocene.

163 *2.3 Exhumation models and implications on the deep orogenic structure*

164 In general terms, exhumation models applied to (U)HP belts can be framed within two different
165 groups, also implying alternative scenarios of mantle involvement: (i) synconvergent exhumation
166 models, either requiring fast erosion or forced circulation in a low-viscosity wedge (e.g., Beaumont
167 et al., 2001; Zeitler et al., 2001; Jamieson and Beaumont, 2013), and (ii) exhumation models that
168 consider boundary divergence within the subduction zone, with a minor role played by erosion (e.g.,
169 Dewey, 1980; Brun and Faccenna, 2008). Both categories of models have been applied to the
170 Western Alps (e.g., Malusà et al., 2011; Butler et al., 2013).

171 Classic tectonic reconstructions of the Alpine belt suggest that synconvergent exhumation
172 could be favoured by deep duplex formation via the accretion of continental material derived from
173 the lower plate (Schmid et al., 2004; Agard et al., 2009), which may be followed by indentation of
174 the upper-plate mantle beneath the accretionary wedge (Schmid and Kissling, 2000; Béthoux et al.,
175 2007). This scenario would imply that seismic velocities in the upper-plate mantle should be similar
176 beneath the orogenic wedge and in the hinterland (Fig. 2B). In case of divergent motion between the
177 upper plate and the descending slab, (U)HP rock exhumation might be instead associated to the
178 emplacement of serpentinized mantle-wedge rocks at shallow depth beneath (U)HP continental
179 rocks, provided that divergence is sufficiently high (Fig. 2C). Because of widespread mantle-wedge
180 serpentinization during subduction (Lafay et al., 2013; Plümpner et al., 2017), seismic velocities are
181 predicted to be lower in mantle-wedge rocks beneath the (U)HP dome, and higher in adjoining dry
182 mantle rocks of the upper plate (Fig. 2C).

183 These alternative scenarios would be in agreement with alternative end-member tectonic
184 reconstructions of the southern Western Alps, based on recent geophysical data from the CIFALPS
185 experiment (Zhao et al., 2016b). One possible end-member reconstruction, consistent with
186 geophysical data, invokes a thick complex of (U)HP continental slivers, in line with predictions of
187 numerical models of syn-convergent exhumation, whereas a second end-member reconstruction
188 invokes a larger volume of mantle rocks possibly exhumed at shallow depth during divergent

189 motion within the subduction zone (Zhao et al., 2015; Malusà et al., 2017). A local earthquake
190 tomography model, complementing previous studies based on receiver function analysis, would be
191 extremely useful to discriminate between these end-member tectonic reconstructions, and may
192 allow a decisive step forward in our understanding of mechanisms leading to exhumation of (U)HP
193 rocks.

194 **3. Methods**

195 ***3.1. Building the database***

196 The local earthquake tomography presented in this work is largely based on the dataset collected
197 during the CIFALPS experiment (Zhao et al., 2016b), which was integrated by data recorded in the
198 same time interval by permanent seismic networks operating in Italy and France, and complemented
199 with selected older events. The temporary network of the CIFALPS experiment (blue marks in
200 Figure 1B) includes 46 broadband seismic stations deployed along a linear WSW-ESE transect
201 from the European foreland to the western Po Plain, and 9 additional stations installed to the north
202 and to the south of the main profile. Stations operated from July 2012 to September 2013, and were
203 specifically deployed for a direct comparison between receiver function and local earthquake
204 tomography. Stations located along the main profile were conceived for receiver function analysis
205 (Zhao et al., 2015). Their spacing ranges from ~5 km in the Western Alps mountain range to ~10
206 km in the European foreland and in the western Po Plain. Off profile stations were installed to
207 improve the crossing of seismic rays for local earthquake tomography.

208 The high number of recording stations along the main CIFALPS profile may increase the
209 computational burden during local earthquake tomography (e.g. in ray tracing) without a direct
210 improvement in the final resolution. However, it ensures a number of advantages. For example, any
211 potential loss of data due to station malfunctioning is easily recovered by adjacent instruments, and
212 doubtful data can be discarded without jeopardizing the quality of the dataset. In order to improve
213 the ray coverage and ensure ray crossing from any azimuth in the study volume, we added to the

214 dataset all published phase pickings recorded by permanent seismic stations operating in France and
215 Italy during the CIFALPS experiment (red marks in Figure 1B). We additionally considered few
216 events that occurred before the experiment to fill specific spatial gaps. This was the case of the
217 intermediate depth earthquakes that were useful to sample anomalies at the bottom of the study
218 volume. Because these earthquakes are relatively rare (Eva et al., 2015), only few events were
219 recorded during the CIFALPS experiment. In summary, 270 events on a total of 1088 events
220 utilized in this work were added as supplementary entries from datasets available at French and
221 Italian seismic networks; about 80% of the remaining events were merged with existing phase
222 pickings. The final P and S ray coverage is shown in Figure 3A.

223 ***3.2. Seismic tomography setup and reconstruction test***

224 We adopted the local earthquake tomography code SIMULPS (Thurber, 1983) for tomographic
225 analysis, in its version 14 that implements the ray tracer by Virieux (1991) to cope with models of
226 regional size. We subdivided the study volume into layers containing nodes, and used an initial
227 velocity model derived from previous seismic experiments over a larger area (Scafidi et al., 2009).
228 Several tests were performed for a correct choice of the inversion parameters, and classical damping
229 trade-off curves (Eberhart-Phillips, 1986) were computed to pick up the best values for P and S
230 velocities.

231 The resolution capability of the coupling between inversion setup and data was evaluated by
232 checkerboard and reconstruction tests. These tests were useful to choose an adequate geometry of
233 the starting model and evaluate the smearing due to the contrast between high and low velocity
234 anomalies. The reconstruction test was specifically conceived to test the potential impact of the
235 high-velocity Ivrea body, a long recognized tectonic feature associated to a positive gravimetric
236 anomaly (red dotted line in Fig. 1) and interpreted as a slice of Adriatic mantle emplaced at shallow
237 depth (Closs and Labrouste, 1963; Nicolas et al., 1990). We used a “stairwell” geometry to simulate
238 a high-velocity east-dipping layer along the CIFALPS profile (Fig. 3B) and test the resolution

239 capability of the coupling between seismic dataset and inversion setup. The same geometry after
240 interpolation by the algorithm used in SIMULPS is shown in Figure 3C. A comparison with Figure
241 3B shows that the interpolation process introduces a smoothing of the anomalies and a band of fake
242 colors around them. Figure 3D shows the reconstruction of the imposed stairwell structure based on
243 our seismic dataset. The inversion of synthetic data does not consider the resolution, and Figure 3D
244 only displays the reconstructed model as if it was completely resolved except for areas that were not
245 sampled (in white). As shown in the reconstruction test, the shape of the anomaly is well
246 reproduced, but the velocity of the first and second steps is lowered from ~ 8.0 km/s (blueish) to
247 about ~ 7.5 km/s (greenish), and weak vertical and horizontal periodic stripes of yellow color appear
248 at ~ 50 km depth. These artifacts, and the undestimation of the magnitude of the high velocity
249 anomalies in the uppermost 10 km of the crust, have been considered during the subsequent phases
250 of tomography interpretation. The real data tomographic model is about 700×700 km wide, and was
251 obtained after 6 iterations on a 12 layers model of 36×36 nodes each. In the central part of the
252 model, spacing between nodes is equal to 15 km.

253 **4. Results**

254 Figure 4 shows the V_p and V_p/V_s cross-sections along the CIFALPS profile. The lighter areas
255 are those where the diagonal elements of the resolution matrix are < 0.1 . This threshold was chosen
256 as the divider between resolved and non-resolved areas based on a comprehensive comparisons
257 between different resolution indicators (Paul et al., 2001). As expected, the maximum depth of the
258 resolved area is limited by the depth of occurrence of most of the deepest events (Eva et al., 2015;
259 Malusà et al., 2017). Beneath the Dora-Maira (U)HP dome, the tomography model is well resolved
260 down to 50-60 km depth, whereas the two extremes of the CIFALPS cross section are poorly
261 resolved. Letters “a” to “k” indicate the relevant velocity features highlighted by the tomography
262 model. The main tectonic structures previously inferred from receiver function analysis (Zhao et al.

263 2015) and surface geology (Lardeaux et al., 2006; Malusà et al., 2015) are also indicated for
264 comparison (black lines in Fig. 4).

265 The most prominent feature of the tomography model is represented by the high velocity body
266 ($V_p \sim 7.5$ km/s; $V_p/V_s = 1.70$ - 1.72), labelled with “a”, which is located right below the Dora-Maira
267 (U)HP dome, at depths as shallow as ~ 10 km. Such a high-velocity body was already imaged with
268 similar velocities by previous works ($V_p \sim 7.4$ - 7.7 km/s; Paul et al., 2001; Béthoux et al., 2007), but
269 was only resolved down to depths of 15-20 km. It is still observed to the south of the CIFALPS
270 profile (Fig. 5D,E), but progressively vanishing towards the north (Fig. 5A,B). A series of N-S
271 cross sections, ranging from the Western Alps to the Po Plain (Fig. 6), shows that this high-velocity
272 anomaly is exclusively found beneath the Dora-Maira (U)HP dome (Fig. 6A), and disappears
273 farther east.

274 The mantle-wedge region labelled with “b” is located at depth of 20-45 km, in correspondence
275 with a cluster of intermediate depth earthquakes that mark the Rivoli-Marene deep fault (RMF in
276 Fig. 4A; Eva et al., 2015). This region shows higher V_p values (~ 8.0 km/s) compared to region “a”,
277 and anomalously high V_p/V_s ratios (> 1.74) that are supportive of low shear wave velocities. This
278 cluster of intermediate depth earthquakes in region “b” is not only observed along the CIFALPS
279 profile, but also in cross sections located more to the north or to the south (Fig. 5). The deepest
280 mantle wedge region resolved by the tomographic model is labelled with “c”. This region, located
281 at depth of ~ 40 - 50 km atop the European slab, shows lower V_p and V_p/V_s values compared to
282 region “b” ($V_p \sim 7.0$ - 7.5 km/s; $V_p/V_s < 1.70$), but the V_p/V_s ratio is locally higher ($V_p/V_s \sim 1.74$).

283 The well-resolved regions of the model also include some subducted European lower crust. This
284 shows a progressive increase in V_p from the region labelled with “d” ($V_p \sim 6.7$ km/s) to the region
285 labelled with “e” ($V_p \sim 7.6$ km/s), under a rather constant V_p/V_s ratio of 1.70-1.72. Such variations
286 are detected in all of the analyzed WSW-ENE transects of Figure 5. No seismic event was recorded
287 in regions “d” and “e” since 1990 (installation of permanent seismic networks) and during the
288 CIFALPS experiment (Malusà et al., 2017).

289 On the eastern side of the transect, the region labelled with “f” is located below the Adriatic
290 Moho as determined by receiver function analysis combined with gravity modelling. It shows Vp
291 values ~8.0 km/s and Vp/Vs = 1.70-1.72. This region is affected by intermediate depth earthquakes
292 that are also observed to the north and to the south of the CIFALPS transect (Fig. 5). The vertical
293 and horizontal periodic stripes of yellow color observed at 50 km depth in this region are artifacts,
294 as confirmed by the reconstruction test of Fig. 3D. Above the Adriatic Moho, measured Vp values
295 are much lower, generally <6.7 km/s, but in places they reach values as high as ~7.2 km/s. Very
296 high Vp/Vs values (>1.8) are locally observed at ~30 km depth at the base of the Adriatic crust.
297 This region, labelled with “g”, is also characterized by a cluster of seismic events that are only
298 observed in the vicinity of the main CIFALPS transect.

299 In the uppermost part of the Alpine orogenic wedge (regions “h” to “k”), Vp values are
300 invariably <6.5 km/s, but major variations in Vp/Vs ratios are locally observed. For example, the
301 region to the east of the Dora-Maira (U)HP dome (labelled with “h”) shows Vp/Vs values >1.72,
302 whereas the region corresponding to the western flank of the Dora-Maira dome (labelled with “j”)
303 shows much lower Vp/Vs ratios, even <1.66. Vp/Vs ratios <1.68 are also observed in the region
304 labelled with “k”, located beneath the Frontal Pennine Fault. The double-vergence accretionary
305 wedge located to the east of the Frontal Pennine Fault, and labelled with “i”, shows instead Vp/Vs
306 values > 1.75, and includes most of the shallow earthquakes recorded in the Western Alps area.

307 **5. Comparison with receiver function analysis**

308 Results of local earthquake tomography are compared in Figure 7 with published CIFALPS
309 results of receiver function analysis (Zhao et al., 2015). Unlike local earthquake tomography, the
310 receiver function technique is based on the analysis of teleseismic earthquakes, and enhances P-to-S
311 (Ps)-converted waves on velocity interfaces beneath an array. The polarity of the converted signal
312 depends on the sign of the velocity change, and interfaces with velocity increase can be
313 discriminated from interfaces with velocity decrease. Assumptions and arbitrary choices of the

314 receiver function approach applied to the CIFALPS transect (e.g., magnitude threshold, epicentral
315 distance, seismograms filtering, velocity model, choice of the direction of back azimuths) are
316 described in full in Zhao et al. (2015).

317 The image of Figure 7B is based on radial receiver functions from teleseismic events with
318 magnitude ≥ 5.5 , epicentral distance of 30-90°, and ENE back-azimuths (see Zhao et al., 2015). This
319 image shows two major interfaces marked by positive-polarity Ps-conversions (red-to-yellow
320 regions), which attest the downward velocity increase corresponding to the European and Adriatic
321 Mohos (thick dashed lines). The eastward-dipping European Moho is recognized from ~40 km
322 depth beneath the Frontal Pennine Fault to ~75 km depth beneath the Po Plain. The Adriatic Moho
323 is recognized from 20-30 km depth, to the east, to 10-15 km depth, to the west. The red spots
324 located at 40-55 km depth beneath the Adriatic Moho are multiples, as confirmed by synthetic tests
325 (Zhao et al., 2015). A shallow positive-polarity converted phase is also observed beneath the Dora-
326 Maira massif, between regions “a” and “h”, whereas a spot of negative-polarity Ps-conversions
327 marking a downward velocity decrease is located above region “c”, at 20-40 km depth (blue
328 region).

329 On the eastern side of the CIFALPS transect, the sharp velocity increase from $V_p < 6.5$ km/s to
330 $V_p > 8$ km/s evidenced by local earthquake tomography faithfully matches the location of the
331 downward velocity increase highlighted by receiver function analysis. Localized anomalies in
332 V_p/V_s ratios, e.g., in region “g”, match with major breaks in the alignment of positive-polarity Ps-
333 conversions. Beneath the Dora-Maira (U)HP dome, the downward increase in V_p values from
334 region “h” ($V_p < 6.5$ km/s) to region “a” ($V_p \sim 7.5$ km/s) is consistent with the observed positive-
335 polarity Ps-conversions, whereas the downward velocity decrease from regions “a” and “b” (V_p
336 ~ 7.5 km/s and > 8 km/s) to region “c” ($V_p \sim 7.0-7.5$ km/s) is consistent with the spot of negative-
337 polarity Ps-conversions located at 20-40 km depth in Figure 7B. The shape of the high-velocity
338 region labelled with “a” is also mirrored by the distribution of seismic events recorded since 1990.
339 Region “a” is virtually aseismic (Malusà et al., 2017), and earthquakes are chiefly located along its

340 external boundaries or in the surrounding regions (Fig. 7B). On the western side of the CIFALPS
341 transect, the alignment of positive-polarity Ps-conversions generated along the European Moho is
342 partly included within the resolved area of the local earthquake tomography model, and fits with a
343 downward velocity increase from ~6.7 km/s (region “d”) to ~7.6 km/s (region “e”). The velocity
344 structure unravelled by the analysis of local earthquakes is thus independently confirmed by the
345 analysis of teleseismic earthquakes (Zhao et al., 2015) and by the distribution of seismic events
346 (Eva et al., 2015; Malusà et al., 2017).

347 **6. Geologic interpretation**

348 The geologic cross section of Figure 7C shows the main features of the orogenic wedge of the
349 Western Alps, and of the mantle wedge between the European and the Adriatic plates, as inferred
350 from the velocity structure derived from local earthquake tomography along the CIFALPS profile.
351 Correlation between seismic velocity and lithology in former subduction zones is a challenging
352 task. Subducted rocks are heterogenous, and display anisotropic fabrics and velocity variations as a
353 function of direction (e.g., Rudnick and Fountain, 1995; Weiss et al., 1999). A full 3D coverage of
354 seismic rays is thus required to get a reliable characterization of the velocity structure (see Fig. 3A).

355 In the European plate, the V_p values ~6.7 km in region “d” are supportive of a relatively felsic
356 composition of the European lower crust (e.g. Rudnick and Fountain 1995; Weiss et al., 1999;
357 Goffé et al. 2003; Wang et al., 2005; Mechie et al. 2012). The homogeneous V_s values < 4 km/s
358 reported by Lyu et al. (2017) suggest that the European lower crust may be rather homogeneous at
359 the scale of seismic observations, and may consist of granulite having felsic to intermediate
360 composition. Major occurrence of granulitic metapelites can be safely excluded, because it would
361 result in much higher V_p (>6.7 km/s up to 7.2 km/s) and V_s values (~4 km/s; Rudnick and
362 Fountain, 1995).

363 The increase in V_p values evidenced at ~40 km depth by local earthquake tomography, from
364 ~6.7 km/s in region “d” to ~7.6 km/s in region “e”, may mirror a progressive eclogitization of lower

365 crust rocks with consequent density increase by metamorphic phase changes (e.g., Hacker et al.,
366 2003; De Paoli et al., 2012). Mineral equilibria at the granulite-eclogite transition depend on rock
367 composition. The eclogitization of a felsic granulite strongly increases the garnet content, and
368 consequently the density from 2.90 to 3.30 kg/dm³, and the P velocity up to a maximum of 7.6 km/s
369 (e.g., Christensen, 1989; Hacker et al., 2003, 2015; Hacker and Abers, 2004; Hetényi et al., 2007).
370 These values are consistent with the V_p values observed in region "e". The increase in P velocity
371 from region "d" to region "e" is associated with a progressive increase in S velocity up to 4.2 km/s
372 (Lyu et al., 2017), which may be either interpreted as an increase in mafic component, or as an
373 effect of metamorphic reactions under increasing pressure-temperature conditions. However, V_p
374 values in region "e" are far too low for a pure mafic eclogite (Bezacier et al., 2010; Reynard, 2013),
375 thus suggesting no major compositional changes from west to east in the European lower crust, but
376 only a progressive change in metamorphic assemblage during subduction. This interpretation also
377 explains the progressive weakening of the positive-polarity converted phases observed along the
378 European Moho, from red to yellow background colours in Fig. 7B, as previously described by
379 Zhao et al. (2015).

380 On the eastern side of the Western Alps, V_p values >8 km/s confirm the presence of Adriatic
381 mantle at shallow depth beneath the western Po Plain (10-15 km), just in correspondence with the
382 positive gravimetric anomaly classically referred to as the Ivrea body (Closs and Labrouste, 1963;
383 Nicolas et al., 1990) and in line with results of previous tomographic models (e.g., Solarino et al.,
384 1997; Paul et al., 2001; Scafidi et al., 2006; 2009; Diehl et al., 2009; Wagner et al., 2012). East of
385 the Ivrea body gravimetric anomaly, the Adriatic Moho is located at 30-35 km depth, which is a
386 much more reliable estimate of the Moho depth beneath the Po Plain compared to previous
387 estimates based on receiver function alone (Zhao et al., 2015). The locally high V_p/V_s ratios >1.8,
388 associated to V_p of 7.0-7.5 km/s (region "g"), may be supportive of gabbro (Weiss et al., 1999)
389 underplated at the base of the Adriatic lower crust. Noteworthy, Permian gabbros are indeed
390 exposed north of the Po plain, where they are intruded into lower crust rocks belonging to the

391 Adriatic (Southalpine) basement (Quick et al., 1994; Schaltegger and Brack, 2007). Above the
392 Adriatic Moho, local spots with $V_p \sim 7.2$ km/s but low V_p/V_s ratios (Fig. 5) are supportive of a
393 more heterogeneous composition of the Adriatic lower crust compared to the European lower crust,
394 and may suggest a local occurrence of granulite facies metapelites (V_p 6.7-7.2 km/s, $V_s \sim 4$ km/s;
395 Rudnick and Fountain, 1995) not only at the surface (e.g., Ewing et al., 2014), but also at depth.
396 Differences in velocity structure among crustal sections now exposed on the opposite sides of the
397 Alps probably reflect a different pre-Alpine evolution, rather than processes related to the Cenozoic
398 evolution of the Adria-Europe plate boundary zone (Guillot et al., 2009b; Carosi et al., 2012;
399 Bergomi et al., in review).

400 In the uppermost part of the Alpine wedge, the structural variability of stacked rocks is largely
401 mirrored by their variability in V_p/V_s ratios. The V_p/V_s values >1.75 observed in the double-
402 vergence accretionary wedge chiefly including Briançonnais and Schistes lustrés units (Lardeaux et
403 al., 2006), may reflect low V_s values, possibly associated to the widespread network of mesoscale
404 faults developed in these rocks since the Neogene (Tricart et al., 2004; Sue et al., 2007; Malusà et
405 al., 2009). To the east, low V_p/V_s values even <1.66 observed on the western flank of the Dora-
406 Maira dome (region “j”) may instead reflect high V_s velocities, suggesting that the poorly fractured
407 granitic gneisses exposed at the surface (Brossasco granite; Paquette et al., 1999; Lenze and
408 Stöckhert, 2007) may be also present at depth. Fracturing may be also invoked to explain the low
409 V_s values observed along the eastern boundary of the Dora-Maira dome, where (U)HP continental
410 rocks are juxtaposed against the eclogitized mantle rocks of the Lanzo massif (Kienast and
411 Pognante, 1988; Piccardo et al., 2007) along the Lis-Trana deformation zone (Perrone et al., 2010).
412 To the west of the Frontal Pennine Fault, V_p/V_s values <1.68 suggest instead that the European
413 upper crust in the External zones is poorly deformed, consistent with minor seismicity recorded in
414 that area (Fig. 7B).

415 But the most relevant results of the tomography model presented in this work is related to the
416 velocity structure beneath the Dora-Maira (U)HP dome. This information is critical to discriminate

417 between contrasting models of (U)HP rock exhumation (Malusà et al., 2011, 2015; Jamieson and
418 Beaumont, 2013), and to discern between end-member tectonic reconstructions recently proposed in
419 the light of available geophysical data (Malusà et al., 2017). The velocity structure of the mantle
420 wedge region “a”, showing Vp velocity of ~7.5 km/s from depths as shallow as ~10 km down to
421 ~30 km, is largely inconsistent with the presence of imbricated continental crust units (e.g., Schmid
422 et al., 2017) or dry mantle peridotite beneath the Dora-Maira (U)HP dome. Instead, it may suggest a
423 complex evolution of mantle-wedge rocks in terms of P-T conditions and fluid-rock interaction.
424 Such Vp values point in fact to widespread serpentinization of mantle rocks (~60% according to
425 Reynard, 2013), that may locally exceed 90% both in the uppermost part of anomaly “a” and in the
426 Lanzo massif, although velocity values in the uppermost crustal levels may be slightly
427 underestimated, as unravelled by the reconstruction tests of Fig. 3D. The degree of serpentinization
428 at 30-40 km depth is instead much lower (<30%), and consistent with the occurrence of
429 intermediate-depth earthquakes (Fig. 7B). Vp/Vs ratios are in the range of 1.70-1.72 in region “a”,
430 but sharply increase to values >1.74 in region “b”, where Vp values (~8.0 km/s) are consistent with
431 dry mantle peridotite. The high Vp/Vs ratios in region “b” point to low shear wave velocities, which
432 are in line with a potential impact of the Rivoli-Marene deep fault on the rock fabric. According to
433 previous work, the deepest part of the mantle wedge beneath the thick blue spot of negative polarity
434 conversions (region “c” in Fig. 7B) may either include serpentinites, or slivers of (U)HP rocks. On a
435 geophysical ground, serpentinites can be easily distinguished from other lithologies possibly found
436 in high-pressure mélange zones (e.g., Marschall and Schumacher, 2012) such as eclogitic
437 metasediments and mafic eclogites (Reynard, 2013). Our results indicate that the velocity values
438 observed in region “c” (Vp ~7.0-7.5 km/s; Vp/Vs <1.70) are neither consistent with eclogitic
439 metasediments (Vp ~7.0 km/s; Vp/Vs ~1.75) nor with mafic eclogite (Vp > 8.0 Vp/Vs ~1.73), but
440 are instead supportive of ultramafic rocks with a degree of serpentinization ranging between 50%
441 and 75% (Weiss et al., 1999; Reynard, 2013). However, minor slivers of eclogitic metasediments

442 could be present at ~40 km depth at the top of the European slab, in regions showing the highest
443 V_p/V_s ratios (Fig. 7A).

444 **7. Implication for (U)HP rock exhumation**

445 In the southern Western Alps, the positive gravimetric anomaly ascribed to the Ivrea body is
446 classically interpreted in terms of upper mantle indentation (e.g., Lardeaux et al., 2006; Béthoux et
447 al., 2007), in line with previous tectonic interpretations proposed for the Central Alps and for the
448 northern Western Alps (e.g., Schmid and Kissling, 2000). According to these interpretations, the
449 uppermost part of the Adriatic mantle would act as an indenter beneath the Alpine accretionary
450 wedge, and would transfer compression towards the European foreland. The main geologic
451 implications of this model include major crustal shortening in the upper plate, and fast erosion
452 focused above the indenter (Fig. 2B). These latter features are indeed observed in the Central Alps,
453 where upper mantle indentation, accommodated by back-folding of (U)HP domes (Keller et al.,
454 2005) and by backthrusting of Adriatic units (Zanchetta et al., 2015), triggered the fast erosional
455 exhumation of the amphibolite-facies rocks of the Lepontine dome (Anfinson et al., 2016; Malusà et
456 al., 2016b). However, these features are not common to the southern Western Alps, where
457 shortening in the accretionary wedge was minor during and after (U)HP rock exhumation (Malusà
458 et al., 2009; Dumont et al., 2012), and erosion was much slower compared to the Lepontine dome,
459 as attested by low-temperature thermochronometers (Malusà et al., 2005b; Vernon et al., 2008; Fox
460 et al., 2015) and by preserved Oligocene corals unconformably lying on top of Eocene eclogites
461 (Molare Fm; Quaranta et al. 2009). A tectonic scenario exclusively invoking upper-plate mantle
462 indentation beneath the accretionary wedge would also imply that seismic velocities in the upper-
463 plate mantle should be quite similar beneath the orogenic wedge and in the hinterland (Fig. 2B).
464 Major seismic velocity changes, e.g., by metamorphic phase changes triggered by fluids released by
465 the downgoing slab, would remain undetected in local earthquake tomography models, because they
466 would take place at much greater depths (Deschamps et al., 2013; Abers et al., 2017).

467 Our study points to a complex velocity structure in the upper-plate mantle of the southern
468 Western Alps. The region beneath the Dora-Maira (U)HP dome is dominated by serpentinitized
469 peridotites, documented from ~10 km depth down to the top of the European slab. To the east, these
470 rocks are juxtaposed against dry mantle peridotites of the Adriatic upper plate along a steeply
471 dipping fault rooted in the lithospheric mantle (RMF in Fig. 7C). In between, mantle rocks of the
472 Lanzo massif underwent subduction during the Alpine orogeny, and were later exhumed and
473 accreted against the Adriatic upper plate when the Dora-Maira (U)HP rocks were still buried at
474 mantle depths (Rubatto and Hermann, 2001). This scenario is supportive of (U)HP rock and mantle-
475 wedge exhumation triggered by upper plate divergent motion (Fig. 2C).

476 Serpentinitized peridotites with $V_p \sim 7.5$ km/s that are found beneath the Dora-Maira dome may
477 have favoured the exhumation of (U)HP rocks across the upper crust, in the depth range where
478 eclogitized continental crust rocks may have become neutrally buoyant (Schwartz et al., 2001).
479 According to Agard et al. (2009), exhumation of eclogitized ophiolites would be favoured by
480 accretion of continental material. Our results point instead to a decisive role played by buoyant
481 serpentinites (e.g., Hermann et al., 2000; Schwartz et al., 2001) during continental (U)HP rock
482 exhumation, within a broadly extensional tectonic framework that is common to many recent
483 tectonic reconstructions of the Central Mediterranean area (e.g., Vignaroli et al., 2008; Malusà et
484 al., 2015) (Fig. 8).

485 No exhumed mantle-wedge serpentinites are recognized so far at outcrop in the southern
486 Western Alps (Scambelluri et al., 1995; Piccardo et al., 2004; Hattori and Guillot, 2007; Deschamps
487 et al., 2013). However strong fluid-rock interactions are recognized in subducted serpentinites and
488 associated ophiolitic rocks (Scambelluri and Tonarini, 2012; Lafay et al., 2013; Plümper et al.,
489 2017), suggesting that fluid release may have occurred during oceanic and even during continental
490 subduction (e.g., Castelli et al., 2007; Ferrando et al., 2009), possibly triggering the partial
491 serpentinitization of the Adriatic mantle wedge. Part of the Adriatic mantle wedge was then exhumed
492 at shallow crustal levels during late Eocene transtension along the Western Alps subduction zone

493 (Malusà et al., 2015) and coeval rapid exhumation of the Dora-Maira (U)HP rocks (Rubatto and
494 Hermann, 2001) (step 1 in Fig. 8). The exhumed mantle wedge was finally indented beneath the
495 Alpine belt during early Oligocene tectonic shortening (Dumont et al., 2012; Jourdan et al., 2012,
496 2013) (step 2 in Fig. 8). Along the Adria-Europe plate boundary, the divergent component of
497 Eocene transtension progressively decreased towards the north to become negligible in the Central
498 Alps (Fig. 8A), where Adria was indented more deeply beneath the accretionary wedge compared to
499 the Western Alps, and rocks now exposed in the Lepontine dome were exhumed at lower rates
500 through the upper crust (Fig. 8B). We speculate that, north of the Dora-Maira dome, upper plate
501 divergence was probably insufficient to allow an effective exhumation of the mantle wedge (Fig.
502 8C). However, testing this hypothesis would require a high resolution tomographic image of the
503 northern Western Alps, which may be precluded by the lack of deep earthquakes.

504 Our results demonstrate that recent geologic cross-sections postulating a thick wedge of
505 Briançonnais eclogites beneath the Dora-Maira dome (e.g., Schmid et al., 2017) are likely incorrect.
506 The palinspastic reconstructions derived from such geologic cross-sections, and exclusively
507 considering a Cenozoic evolution within a broadly compressional framework, should be
508 reconsidered at the advantage of palinspastic reconstructions also including major episodes of
509 divergence within the plate boundary zone (e.g., Vignaroli et al., 2008; Malusà et al., 2015). Mantle
510 wedge exhumation is in fact more consistent with a late Eocene transtensional tectonic framework
511 (Fig. 8C) followed by early Oligocene convergence (Fig. 8D), accommodated by orogen-
512 perpendicular shortening in the external Alps (Dumont et al., 2012) and by transpressional tectonics
513 in the Alps-Appennines transition zone (Malusà and Balestrieri, 2012).

514 The occurrence of mantle-wedge serpentinites exhumed at shallow depth within a continental
515 subduction zone is not specific of the southern Western Alps. Mantle wedge serpentinites associated
516 with (U)HP rock are described, for example, in the Indus Suture Zone in the Himalaya, in the
517 Carribean (Guillot et al., 2001; Deschamps et al., 2012), in the Western Gneiss Region in Norway
518 (Scambelluri et al., 2010), and are inferred by geophysical evidence under the Dabie-Sulu (Liu et

519 al., 2015). Our findings suggest that orogen-scale exhumation of the mantle wedge may represent a
520 prominent, but still underestimated feature of the deep structure of many orogenic belts. As such, it
521 should be integrated in more advanced theoretical models of subduction and exhumation. Moreover,
522 widespread mantle-wedge exhumation may explain the common occurrence of boudinaged mantle-
523 wedge rocks within continental UHP rocks in the roots of old orogenic belts now unroofed by
524 erosion. In pre-Cenozoic orogenic belts such as the Dabie-Sulu or the Western Gneiss Region,
525 where the evidence of minor erosion during UHP exhumation, if any, is no longer preserved, the
526 occurrence of mantle wedge rocks at shallow depth may represent the only evidence supporting
527 (U)HP rock exhumation triggered by divergent motion between upper plate and accretionary
528 wedge.

529 **8. Conclusions**

530 The new local earthquake tomography model of the southern Western Alps, independently
531 validated by receiver function analysis, unravels a complex seismic velocity pattern consistent with
532 a composite structure of the mantle wedge above the subducted European lithosphere. Seismic
533 velocities indicate that the Dora-Maira (U)HP dome lays directly above serpentinitized peridotites,
534 documented from ~10 km depth down to the top of the eclogitized lower crust of the European
535 plate. We propose that peridotite serpentinitization was the result of fluids released to the Adriatic
536 mantle wedge during Alpine subduction. During late Eocene transtension, when the subduction
537 wedge was largely exhumed at the Earth's surface, part of the mantle wedge was also exhumed at
538 shallow crustal levels, to be finally indented under the Alpine metamorphic units in the early
539 Oligocene. Our results suggest that mantle wedge exhumation may represent an important feature of
540 the deep structure of exhumed continental subduction zones. Deep orogenic levels, as those imaged
541 by local earthquake tomography in the southern Western Alps, may be exposed today in older
542 continental subduction zones, where mantle wedge serpentinites are commonly associated to
543 continental (U)HP metamorphic rocks.

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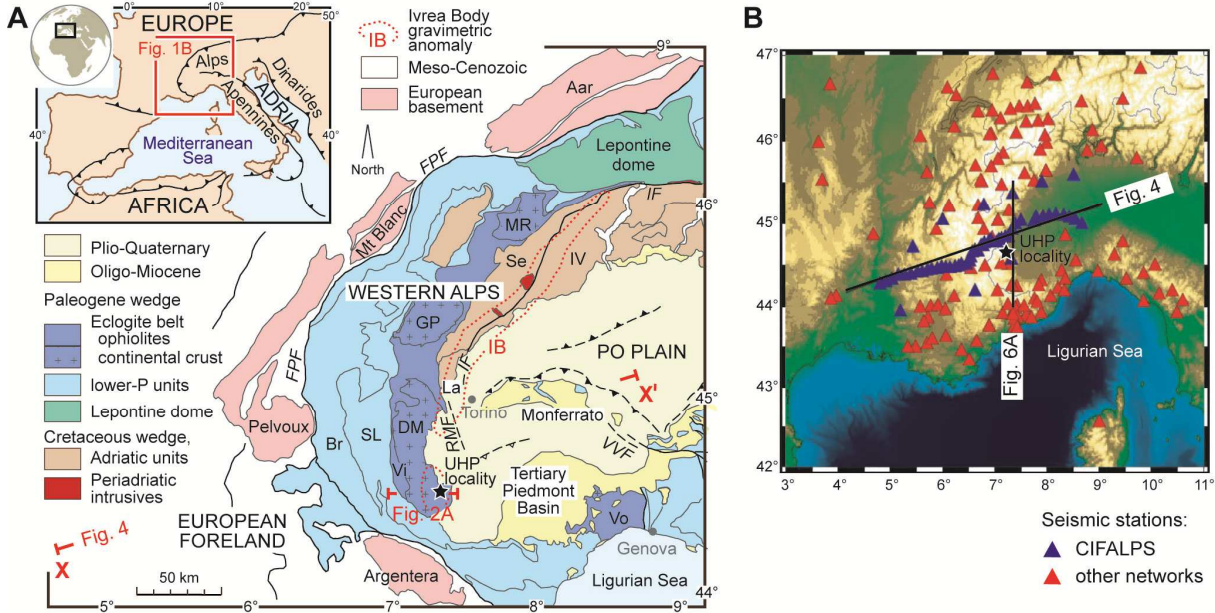
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Figure 1



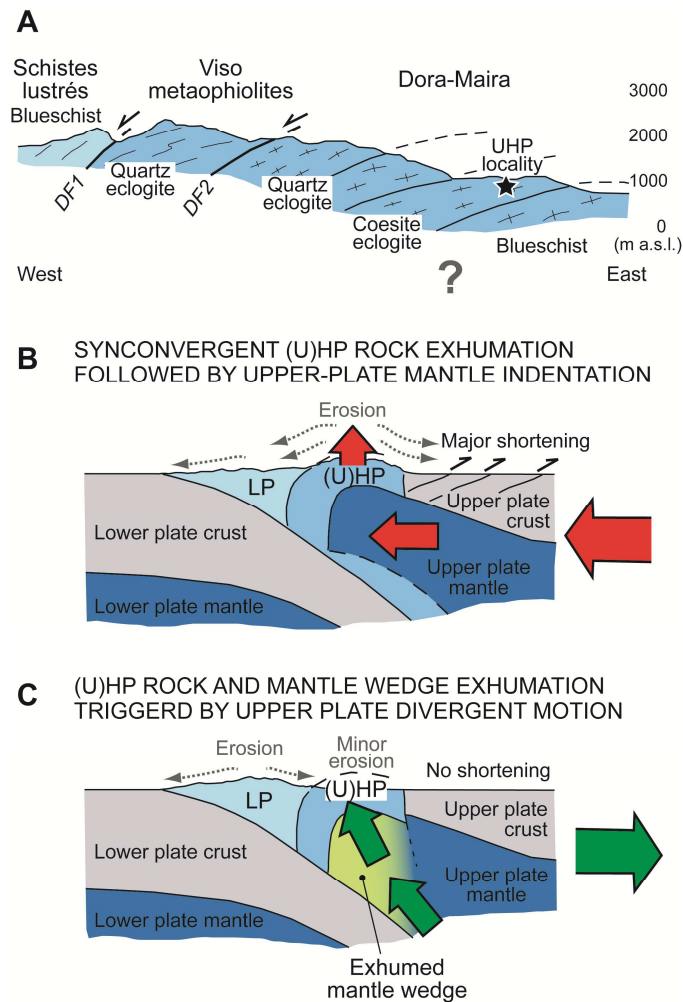
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Figure 1: A) Tectonic sketch map showing the (U)HP domes of the Western Alps (dark blue), the gravimetric anomaly of the Ivrea body (0 mGal isoline in red), and the location of the CIFALPS transect (X-X'). Acronyms: Br, Briançonnais; DM, Dora-Maira; FPF, Frontal Pennine Fault; GP, Gran Paradiso; IF, Insubric Fault; IV, Ivrea-Verbanò; La, Lanzo; MR, Monte Rosa; RMF, Rivoli-Marene deep fault; Se, Sesia-Lanzo; SL, Schistes lustrés; Vi, Viso; Vo, Voltri; VVF, Villalvernia-Varzi Fault. The black star marks the Brossasco-Isasca UHP locality. B) Seismic stations utilized in this work (blue = CIFALPS; red = other networks) and location of tomographic cross sections (black lines).

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Figure 2

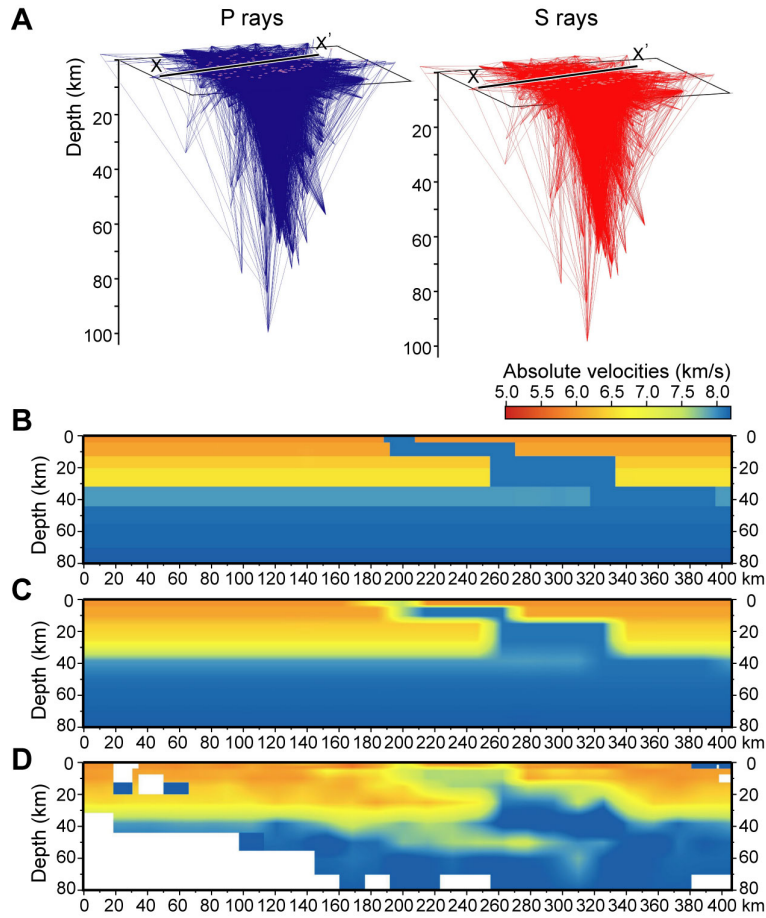


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Figure 2: **A)** Geologic cross-section of the Dora-Maira (U)HP dome (see location in Fig. 1A; based on Avigad et al., 2003; Lardeaux et al., 2006). **B, C)** Alternative scenarios of mantle involvement in (U)HP orogenic belts. In (B), synconvergent exhumation of (U)HP rocks (e.g., Butler et al., 2013), possibly associated with deep duplex formation (Schmid et al., 2017), is followed by indentation of the upper-plate mantle beneath the accretionary wedge, with consequent fast erosion of the (U)HP dome and major tectonic shortening in the upper plate (e.g., Béthoux et al., 2007). Seismic velocities in the upper-plate mantle are similar beneath the orogenic belt and in the hinterland, as indicated by the uniform dark blue colour. In (C), divergence between upper plate and accretionary wedge triggers the exhumation of (U)HP rocks (Malusà et al., 2011) and the emplacement of serpentinitized mantle-wedge rocks at shallow depth. Erosion on top of the (U)HP dome is minor at this stage, and shortening is negligible. Because of widespread serpentinitization of the mantle wedge during subduction, seismic velocities will be lower in the mantle-wedge rocks beneath the (U)HP dome (as indicated by the pale green colour), and higher in the adjoining dry mantle rocks of the upper plate (dark blue).

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Figure 3

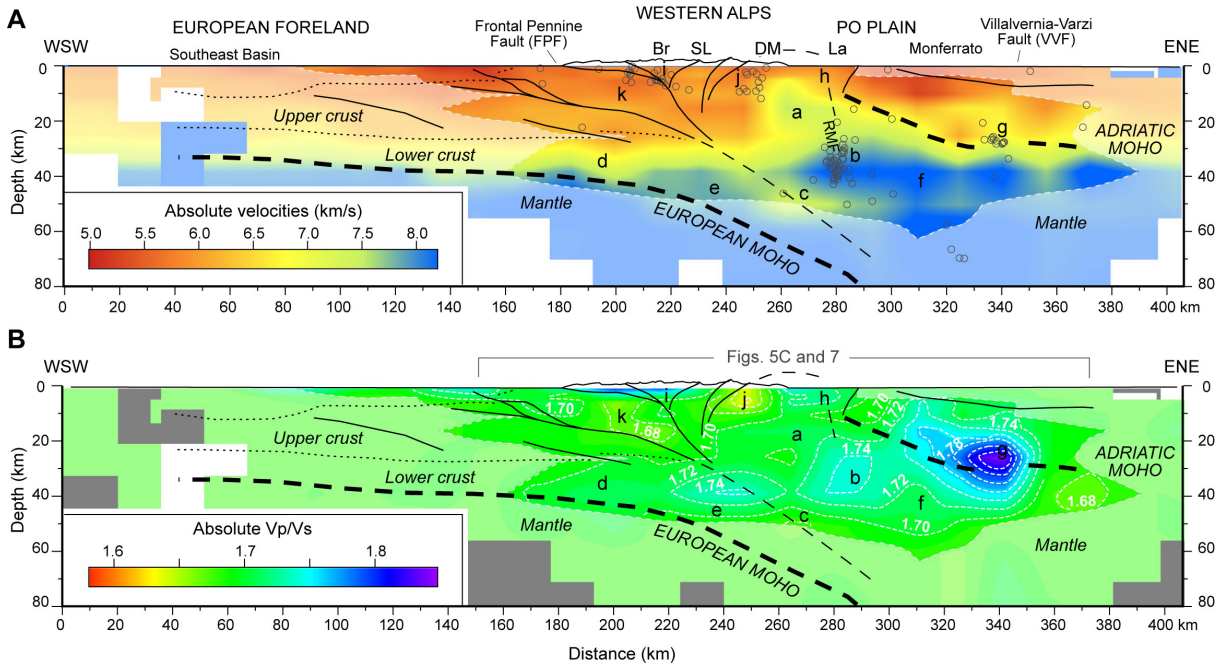


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Figure 3. A) Three-dimensional P and S ray coverage based on the seismic events considered in this study (X-X' indicates the CIFALPS transect, see Fig. 1). B) Imposed stairwell geometry along the CIFALPS transect for testing the resolution capability of the coupling between seismic dataset and inversion setup. C) Same geometry after interpolation by the algorithm used in SIMULPS, which introduces a smoothing and a thin band of fake colors around the anomalies. D) Reconstruction test showing that the shape of the imposed stairwell structure is well reproduced using our dataset, but the high velocities in the uppermost 10 km are converted to lower values (as less as 0.5 km/s); the weak vertical and horizontal periodic stripes of yellow color at 50 km depth within the blue area are artifacts; white areas are not sampled.

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Figure 4

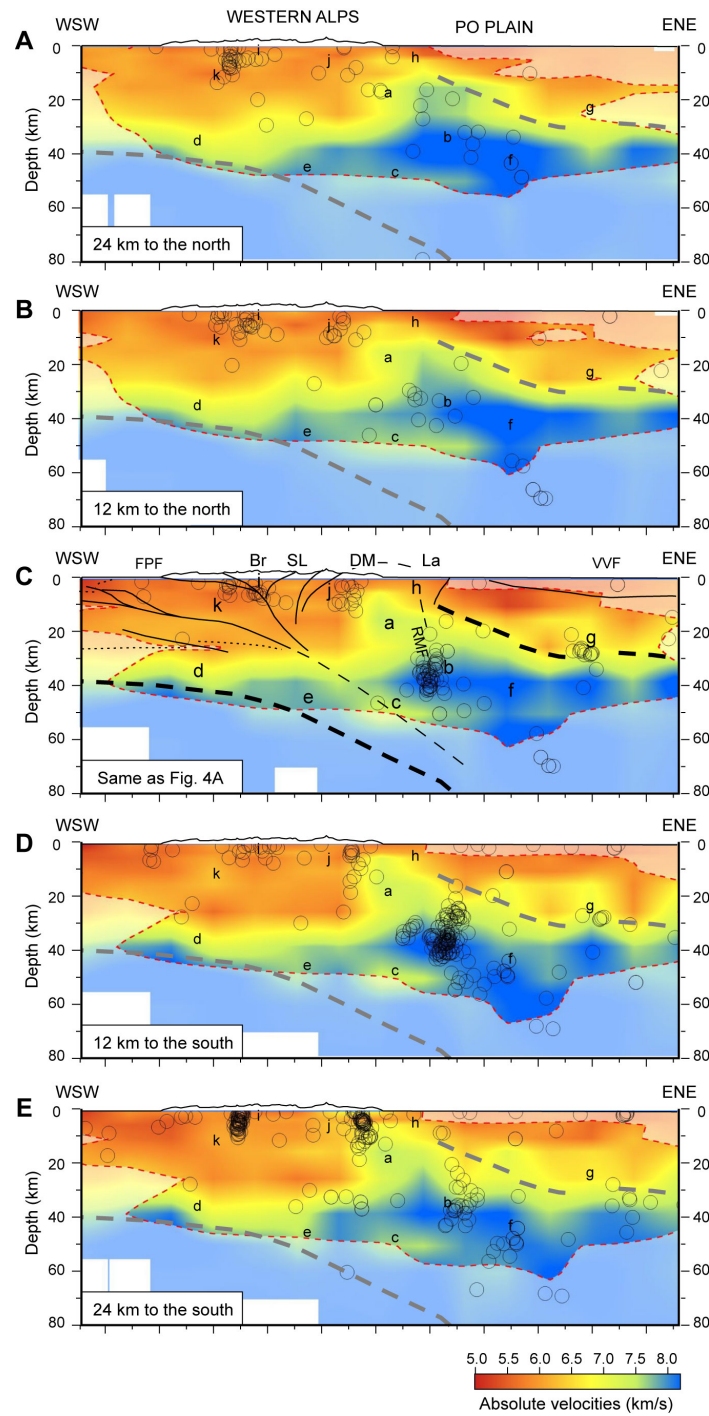


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Figure 4: Tomographic cross sections along the CIFALPS transect. A) Absolute Vp velocity. The velocity structure beneath the Dora-Maira (U)HP dome is well resolved down to 50-60 km depth (acronyms as in Fig. 1A); areas with resolution diagonal elements <0.1 are masked, white areas are not sampled; letters a to k indicate regions of the model discussed in the main text; black circles indicate earthquakes as located with the 3D model; black lines and text in italics indicate the main tectonic features previously inferred from receiver function analysis (Zhao et al., 2015; Malusà et al., 2017, see Fig. 7B). Note the prominent high velocity body (labelled with “a”) located right below the Dora-Maira (U)HP dome. The vertical and horizontal periodic stripes of yellow color at 50 km depth are artifacts, as attested by the reconstruction test of Fig. 3D. B) Vp/Vs ratios. White dashed lines are isolines of equal Vp/Vs, grey areas are not sampled (other keys as in frame A).

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Figure 5

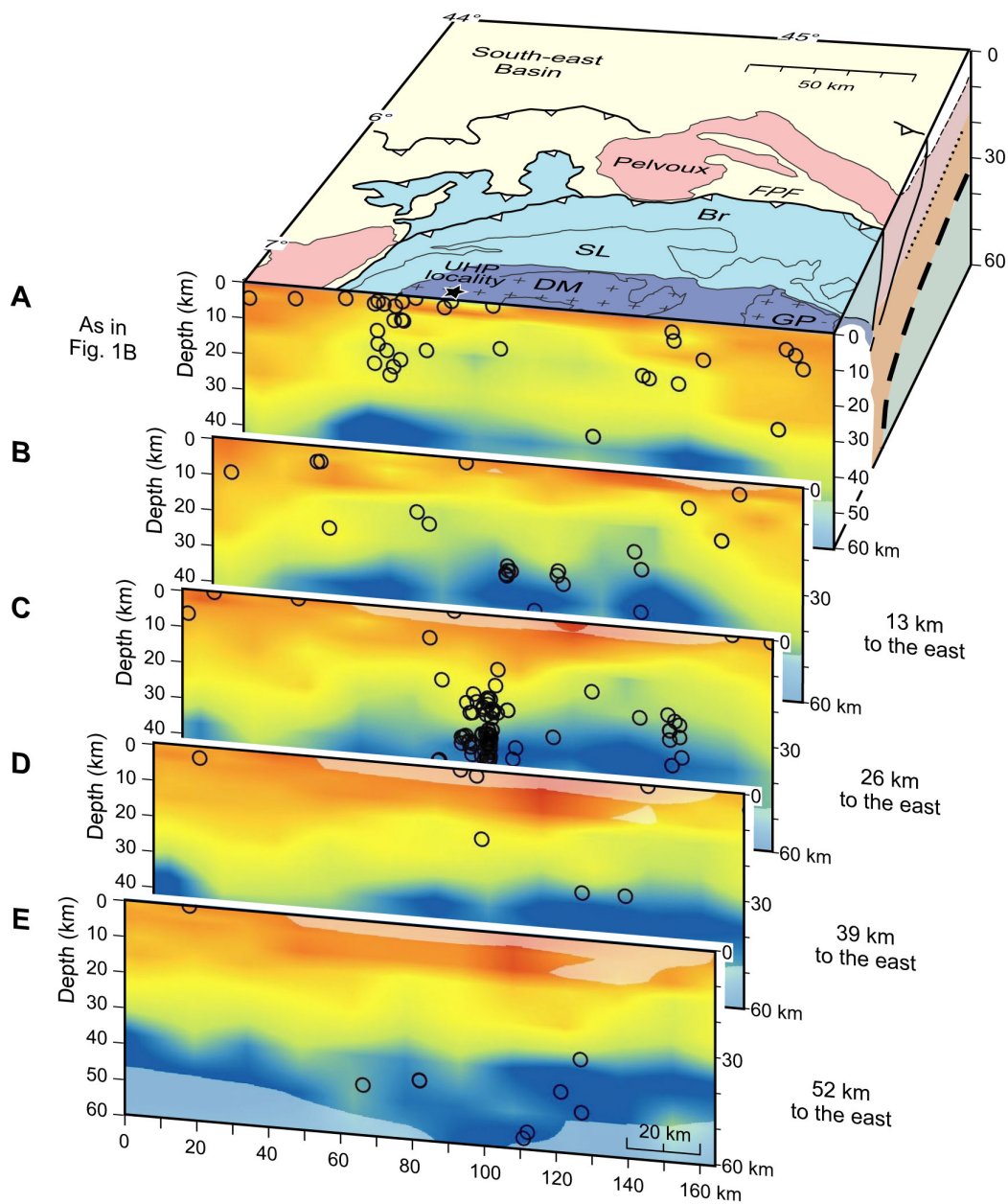


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Figure 5: Lateral variations in Vp velocity in the mantle wedge as shown in a series of WSW-ENE cross-sections lying to the north (A, B) and to the south (D, E) of the main CIFALPS transect (C). The high velocity body labelled with “a” progressively disappears moving to the north. Black circles are projected hypocentres located within ± 5 km distance off the profiles. The thick dashed lines, reported in all sections for comparison, indicate the European and Adriatic Mohos inferred from receiver function analysis (cf. Fig. 7B). Other keys as in Fig. 4.

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Figure 6

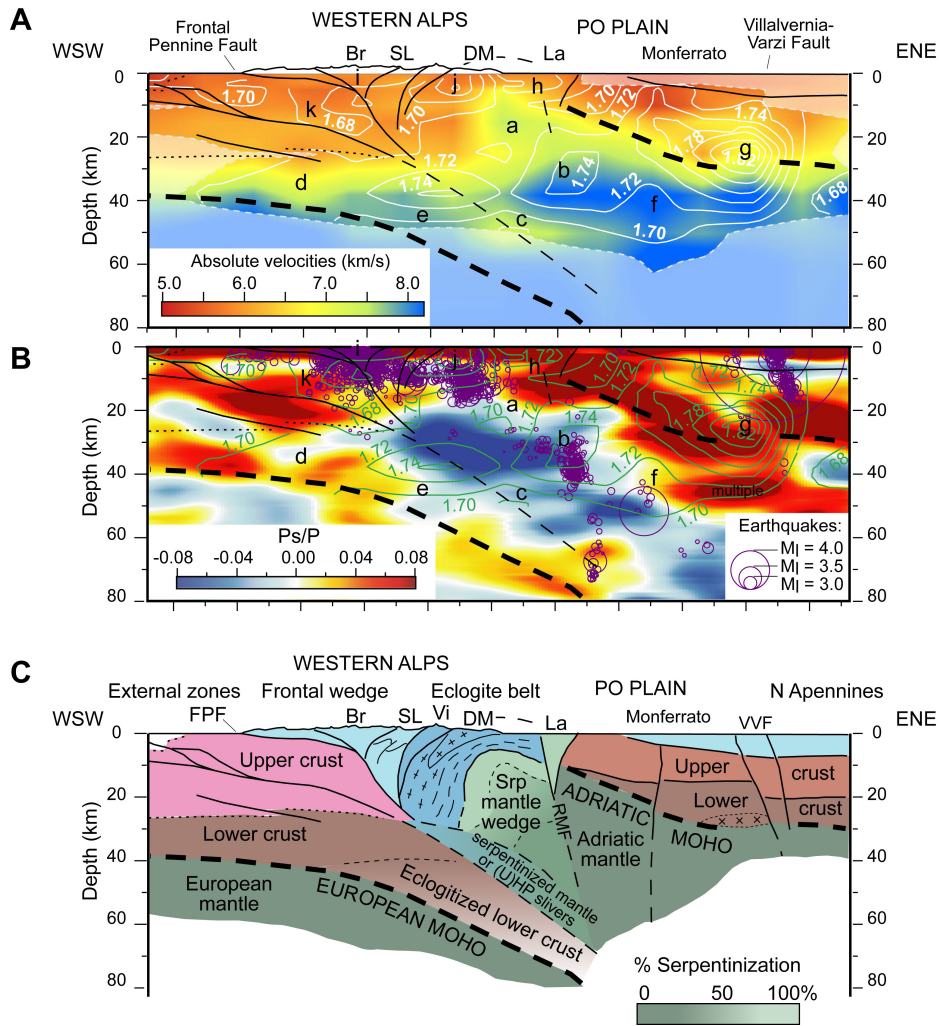


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Figure 6: Lateral variations in V_p velocity beneath the Dora-Maira (U)HP dome, as shown in a series of N-S cross-sections from the mountain range to the Po Plain. Black circles are projected hypocentres located within ± 3 km distance off the profiles. The high-velocity body labelled with “a” in Figs. 4 and 5 is exclusively found beneath the Dora-Maira dome (see cross section A) and progressively disappears towards the east. Acronyms as in Fig. 1.

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Figure 7

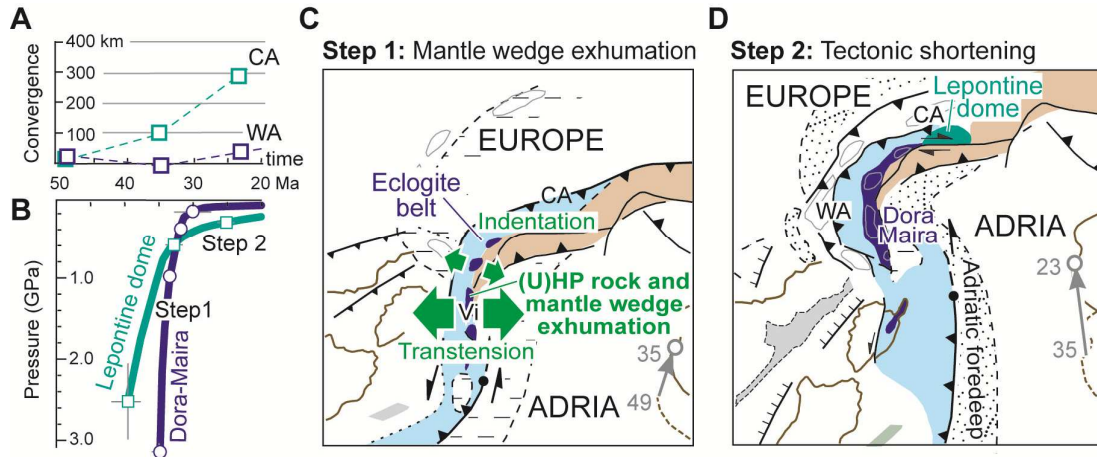


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Figure 7: Synthesis of geophysical data (A, B) and inferred mantle wedge structure (C). Black lines in A and B are tectonic features based on receiver function analysis (colors in B indicate positive- and negative-polarity Ps-converted phases, Zhao et al., 2015); contours are isolines of equal V_p/V_s ; purple circles in B are earthquakes recorded since 1990 (Malusà et al., 2017). The amount of serpentinitization in C, in the mantle wedge underlying the Dora-Maira (U)HP dome, is inferred from seismic velocities (Reynard, 2013). Note the consistency between structures unravelled by local (A) and teleseismic (B) events. Acronyms as in Fig. 1, letters a to k as in Fig. 4.

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Figure 8



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Figure 8: Geodynamic framework of mantle wedge exhumation. **A)** Trench-normal component of Adria-Europe relative motion in the Central (CA) and Western Alps (WA) segments of the Alpine subduction zone (Malusà et al., 2015). **B)** Pressure-time exhumation paths (Dora-Maira: Chopin et al., 1991; Rubatto and Hermann, 2001; Lepontine dome: Becker, 1993; Gebauer, 1996; Brouwer et al., 2004; Nagel, 2008). **C,D)** Late Eocene transtension leading to (U)HP rock and mantle wedge exhumation, and subsequent tectonic shortening in the early Oligocene; grey arrows indicate Adria motion relative to Europe (modified after Malusà et al., 2015).