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The La Bellière gold and antimony district (French Armorican Massif): a two-stage evolution model controlled by Variscan strike-slip tectonic

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Abstract

The La Bellière gold district, in the Ligerian domain of the Armorican Massif is the third one of France in term of production. It shows gold-bearing quartz veins neighboring with a complex network of Sb veins, hosted in the Neoproterozoic metasedimentary basement of the Mauges Nappe. New structural data, mineralogical, textural and geochemical assays have been carried out to clarify the gold and Sb veins relationships. Our results favor of two-step hydrothermal event with an early deep-seated gold event followed by a late shallower Sb event. The gold-bearing quartz veins are hosted by N60°E to N90°E sinistral strike-slip faults. These veins exhibit a four-stage paragenesis: 1) early barren quartz filling, 2) arsenopyrite-pyrite with minor scheelite assemblage, 3) gold and base metal sulfides crystallized after a ductile/brittle deformation stage, and 4) pyrite-carbonate coeval

with a late brittle stage. These results support a continuous evolution model from a deep emplacement level (12 to 9 km) to a shallow one (7 to < 5 km) with a brutal change in the P/T conditions that triggered the deposition of gold (stage 2) controlled by a seismic-valve mechanism. On the contrary, Sb veins correspond to open space vein types, hosted in a N20°W to N20°E conjugated strike-slip fault network arranged in a Riedel shear model, controlled by N130-140°E regional dextral strike-slip shear zone associated to the South Armorican Shear Zone system. A two-stage evolution is recognized, the first one consists of early pyrite-arsenopyrite with microcrystalline quartz, followed by a late stibnite stage formed after geodic quartz deposition. Quartz texture argues for a shallow emplacement (< 5 km) of veins controlled by a suction-pump mechanism. Chronologically, the gold vein opening complies with a NE-SW maximum shortening direction different from the late-Variscan stress field responsible for the dextral Armorican shear zone during late Carboniferous, and constrains an indirect dating. Thus, a Visean age is proposed for the formation of gold veins. The overall features of gold deposits are close to those of mesozonal orogenic gold deposits in Europe. Conversely, Sb-bearing veins are generated by a N-S shortening during the late Carboniferous. The Sb event, distinct in time from the earlier gold one, might correspond to a different expression of the so-called “epizonal” type in the orogenic gold model. It is proposed that the Late Carboniferous “Or 300” event is not the unique metallogenic period for gold deposition in the French Variscan domain. Our study highlights the role of the sinistral strike-slip shearing of possible Visean age, in the gold metallogenesis of the Ligerian domain.

Keywords. Orogenic gold; Sb mineralization; La Bellière district; Or300 hydrothermal event; Armorican Massif; Structural control; Variscan metallogeny

1. Introduction

Gold and antimony (Sb) quartz vein-type deposits are recognized and mined since Antiquity in the European Variscan orogen. They are widespread in the Bohemian Massif (Nemec and Zacharias, 2018), Rhenish Massif, and Saxothuringian zone in Germany (Dill, 1985; Wagner and Cook, 2000), Iberian Massif (Gumiel and Arribas 1987; Neiva et al., 2008), French Massif Central and Armorican Massif (Périchaud 1980; Bril and Beaufort, 1989; Chauris and Marcoux, 1994; Bellot et al., 2003; Bouchot et al., 2005), Slovakian Western Carpathians (Chovan et al., 1995, 1998). Most of them were formed during the late and post-orogenic extensional phases of the orogeny. They display a clear structural control in relation to regional strike-slip and normal faults (Wagner and Cook 2000; Bellot et al., 2003; Bouchot et al., 2005; Neiva et al., 2008). Most of gold deposits can be classified as orogenic gold type deposits (Robert et al., 1997; Groves et al., 1998, 2003; Bierlein and Crowe, 2000; Goldfarb et al., 2005). However, some of them better fit with the intrusion-related gold deposit type (Zacharias et al., 2014). According to the same authors, the Variscan Sb deposits correspond to shallow quartz veins systems postdating the orogenic gold deposits.

The French Massif Central, which displays the most important gold reserves in France, has been studied since more than thirty years (e.g. Touray et al., 1989; Boiron et al., 1989, 1990, 2003; Bonnemaïson and Marcoux, 1990; Bril et al., 1994; Marignac and Cuney, 1999; Charonnat et al., 1999; Essaraj et al., 2001; Nicaud, 2001; Bouchot et al. 2000, 2005; Chauvet et al., 2012). These authors interpreted the formation of gold and Sb-bearing veins as the result of a progressive evolution of a single hydrothermal event in which Sb was formed after the gold stage. In the Armorican Massif, the geological constraints on gold and Sb deposits are still rare, and restrained to the la Lucette Sb-Au district (Serment, 1978), Le Semnon Sb-Au district (Pochon et al., 2018), and the Quimper Sb district (Fouquet 1980; Fig.1), all located in Central- and North-Armorican domains, and to the Vendée Sb-Au district (Marcoux et al., 1984; Bailly et al., 2000) in the South-Armorican domain. Gold and Sb

deposits are considered to be formed during a large-scale hydrothermal event developed between 305-300 Ma in the French Massif Central (Marignac and Cuney, 1999; Bouchot et al., 2000, 2005 and enclosed references), and around 360 Ma (Pochon et al., 2018) in the Central-Armorican Massif. In the Ligerian domain and South-Armorican Domain, the genetic process of gold and Sb deposition, and the understanding of the spatial association of these two elements remain poorly constrained. This study focusses on the La Bellière gold district, hosted in the Variscan Mauges Nappe of the Variscan Armorican Massif (Ballèvre et al., 2009; Figs. 1-2). This district shows a peculiar context with gold-bearing quartz veins neighboring with a complex network of Sb veins. It represents the major gold district in the Armorican Massif according to its economic potential (Bouchot et al., 1997).

This paper presents the structural control of the gold and Sb quartz veins, the mineralogical and geochemical features that allow us to decipher relationships between gold and Sb mineralizations, and the veins formation mechanisms. Their paragenetic evolution, the nature of hydrothermal fluids responsible of the ore deposition, and their timing during the late stage of the Variscan orogeny are discussed.

2. Regional geological background

2.1 The Armorican Massif

With Iberian Massif, the Armorican Massif forms the westernmost part of the European Variscan belt that results from a continental collision between Gondwana and Laurussia plates during the late Devonian and Carboniferous. The Variscan orogeny ended during the late Carboniferous (Gzhelian) around 300 Ma (Matte, 1986; Faure et al., 2005; Ballèvre et al., 2009). From the geodynamic and paleogeographic points of view, the Armorican Massif (MA in the following) is subdivided from north to south into the Léon, North-, Central-, and South-Armorican domains that corresponds to four continental masses (Fig. 1, Faure et al., 2005, 2008; Ballèvre et al., 2009). The Nort-sur-Erdre fault that separates the last two domains represents the Devonian eo-Variscan suture.

Furthermore, the MA is split by crustal-scale Late Carboniferous dextral strike-slip faults namely, the North, and South branches of the South Armorican Shear Zone, and the North Armorican Shear Zone (Jégouzo, 1980). The dextral motion of these crustal-scale strike-slip shear zones was coeval with the emplacement of late syntectonic plutons, like the Questembert granite dated at ca 316 Ma (Ar-Ar on muscovite; Tartèse et al., 2011), or the Mortagne pluton (Cuney et al., 2001) formed by partial melting of metasediments (Tartèse and Boulvais, 2010).

2.2 The Ligerian domain

The Ligerian domain corresponds to the area located on both sides of the Nort-sur-Erdre eo-Variscan suture (Fig. 2). South of the suture, it consists of a stack of Variscan nappes (Fig. 2, Marchand et al., 1969a and b; Cogné, 1974; Marchand et al., 1981; Lardeux and Cavet, 1994; Wyna et al., 1998a and b; Cartier and Faure, 2004; Faure et al., 2008; Ballèvre et al., 2009). From bottom to top, the following units are recognized. The Mauves-sur-Loire Unit consists of metagreywackes and metapelites metamorphosed under greenschist facies conditions. The overlying Champtoceaux Complex is a tectonic imbrication of paragneiss, orthogneiss with lenses of fine-grained eclogites frequently retrogressed into amphibolites, and migmatite (Marchand et al., 1981; Ballèvre et al. 2009). These foliated rocks experienced a high-pressure metamorphism well preserved in the mafic protoliths. Quartz-feldspar rich protoliths of the Champtoceaux Complex are commonly melted as migmatites (Faure et al., 2008; Pitra et al., 2010). Monazite from the migmatite that yields chemical U-Th-Pb age of 387 ± 6 Ma, is interpreted as a result of the exhumation of high-pressure rocks (Cocherie et al. 2005), a second migmatitization event is also recorded between 340-330 Ma (Pitra et al., 2010). The overlying Drain Unit consists of mafic rocks (basalts, volcanoclastites, gabbros), ultramafic, and siliceous rocks. These rocks are interpreted as Variscan ophiolitic remnants. Conversely to the Champtoceaux complex, the rocks of the Drain Unit were metamorphosed under amphibolite facies conditions only. The uppermost unit, or Mauges Nappe, is formed by a 4-5 km thick Neoproterozoic greywacke and pelite series metamorphosed and foliated during the

Neoproterozoic Cadomian orogeny. The Neoproterozoic rocks are unconformably overlain by a Cambrian to Ordovician sedimentary and volcanic series (Blaise et al. 1970; Wyns and Le Metour 1983; Wyns et al., 1998a and b). The Mauges Nappe is folded as an E-W regional antiform (Cogné, 1974).

The Nort-sur-Erdre suture was firstly reworked as a sinistral strike-slip fault that controlled the Tournaisian to Visean opening of the Ancenis pull-apart basin (Diot, 1980; Rolet, 1984; Lardeux and Cavet, 1994; Cartier and Faure, 2004). Moreover, this fault has been reactivated in a dextral motion during the Gzhelian (ex-Stephanian; Diot, 1980; Lemarchand et al., 2012). To the South, the south branch of the South Armorican Shear Zone is the southern boundary of the stack of nappes along which the Late Carboniferous two mica Mortagne granitic pluton emplaced (Jegouzo, 1980). Muscovite from the La Pommeraie leucogranite yields $^{40}\text{Ar}/^{39}\text{Ar}$ Visean ages at 337-336 Ma (located in Fig. 2A; Gumiaux et al., 2004). The stack of nappes is folded in an E-W striking upright anticline, probably coeval with the dextral movement of the South Armorican Shear Zone (Bitri et al., 2003). The Late Carboniferous magmatism is also represented in this area by the Chemillé two mica granite located in the eastern part of the Mauges Nappe and inferred of late Carboniferous age (Wyns and Hottin, 1992). Its western extension to the South of the Beaupréau town (Fig. 3A) is inferred by gravimetric anomaly (Martelet et al., 2009).

2.3 Geological setting of the La Bellière ore district

2.3.1 Geology of the Mauges Nappe

The La Bellière ore field (Fig. 3A) is hosted within an alternation of metagreywackes and metapelites belonging to the Mauges Nappe. The Neoproterozoic mafic volcanic and volcanic-sedimentary rocks (called the Beaupréau formation), intercalated in the greywacke-pelite sedimentary series, form the southern boundary of the ore field. The Beaupréau formation geometrically overlies

the metagreywacke from which it is separated by a graphite-rich black shale unit showing some iron cap covering some pyrite and pyrrhotite occurrences. The northern boundary of the ore field corresponds to the metapegmatite dyke field of La Boissière-sur-Evre supposed to be of Ordovician age (Wyns and Le Métour, 1983). In the ore district, the Cadomian regional foliation of the host rocks strikes N80-120°E, and dips 10 to 50° to the south (Fig. 3B). Thus, these deposits lay on the southern limb of a N100E-N115E striking, km-sized late Variscan Mauges antiform. A crenulation lineation develops parallel to the antiform axis.

The Mauges Nappe experienced three metamorphic events. i) A greenschist facies one, well developed in the mafic volcanic Beaupréau formation, and also in the surrounding terrigenous rocks (south of St-Pierre-Montlimart). ii) It is overprinted to the W-NW by a medium pressure/medium temperature (MP/MT) event, characterized by staurolite, garnet, and locally kyanite. iii) Close to La Rouillère (Fig. 3), biotite and andalousite porphyroblasts, oblique to the regional Cadomian foliation, may represent a thermal contact metamorphism due to an underlying hidden granitic pluton, probably equivalent to the Chemillé leucogranite (Wyns, 1984). The greenschist facies metamorphism and the MP/MT assemblages (i) are related to the Cadomian event. The (ii) and (iii) events are related to the Variscan history (Wyns and Le Métour, 1983).

2.3.2 *The gold and Sb deposits of La Bellière district*

Gallic tribes, 2000 years BC, and the Romans already mined gold-bearing quartz veins in the Mauges region. From the end of the XIXth to the end of the XXth century, the La Bellière district was intensively explored, and successfully mined between 1905 and 1952 for gold. The La Bellière deposit, discovered in 1895, produced 10.32 t of gold with an average grade of 9 to 10 g/t Au (Bouchot et al., 1997 and enclosed references) from a complex quartz vein lode network extending more than 2 km in the WSW-ENE direction (Fig.3A-C). Between the middle of the seventies to the nineties the district was intermittently explored by the French Geological Survey (BRGM) and non-economic

gold-shoots were discovered. Presently, the La Bellière gold district remains the most important one in the Armorican Massif, the third in term of metropolitan gold production, and presents the strongest economic potential for gold discovery in the Armorican Massif.

During previous works, two types of ore-bearing quartz veins have been identified: i) gold with sulfides, the most important one from an economic geology point of view, and ii) Sb veins (Wyns 1984; Braux et al., 1991; Bouchot, 1992). In the La Bellière gold field, about twenty gold-bearing quartz shear veins display a roughly triangular cartographic shape open to the east, extending about 15 km in the E-W direction, with an eastward widening from 3 km to 7 km in width (Fig. 3). Two main groups of cm- to 20 m-thick gold-bearing quartz veins are recognized crosscutting the Neoproterozoic rocks of the Mauges Nappe, Le Petit Moulin and La Bellière group, to the north and La Rouillère and Bégrolle group to the south. These N60-90°E striking, steeply South dipping (60-80°) veins are discontinuously exposed (Fig. 3). The hydrothermal alteration (up to 20m) of the country-rock is characterized by bleaching, intense silicification, sericitization, and little sulfidation marked by pyrite dissemination (Wyns, 1984). Three main stages have been identified in the paragenetic sequence (Sevensma, 1942; Bouchot, 1992): (1) an early stage with massive arsenopyrite and minor pyrite, (2) an ore stage with galena, sphalerite, chalcopyrite, pyrrhotite and electrum, developed after a cataclastic episode, (3) a final barren stage. Two controversial structural models have been proposed to account for the gold-bearing quartz veins formation. Firstly, according to Wyns (1984), the gold-bearing quartz veins correspond to tension gashes opened during the activity of a N100-110°E trending sinistral strike-slip fault in the Tournaisian-Viséan. Secondly, Bouchot (1992) proposed a fault fill vein genesis under a sinistral shearing setting during the Late Carboniferous. According to the same authors, the source of the hydrothermal fluids would be related to a magmatic intrusion hidden under the ore field. However, a metamorphic fluid source has also been proposed (Boiron et al., 1990).

Sb-bearing quartz veins are rarely exposed in the field, but numerous Sb-bearing float stones are widespread in the entire district (Fig. 3; Wyns, 1984, 1987). They display a N10°W to N20°E strike with a steep dip, and less than 1m thickness. The ore consists of massive stibnite with few arsenopyrite and pyrite associated with a black chalcedonic quartz. Presently, no genetic model for these Sb veins has been proposed, however Bouchot (1992) argued that the Sb quartz veins crosscut the gold ones.

3. Sampling and analytical methods

The geometry of the quartz veins has been assessed during our new geological mapping. Due to poor outcrop conditions, the presence of quartz veins is also inferred on the basis of float stones, and remains of mining works from Gallo-Roman times to the beginning of the XXth century. The compilation of former descriptions, maps and exploration geochemical works (Mulot 1982; Wyns 1984; Bonnemaïson and Boubée de Gramont 1985; Terree 1986; Wyns 1987; Blouin and Apolinarski 1990; Braux et al. 1991; Martel-Jantin 1996; Gouin et al., 2017) allowed us to set up an exhaustive state of art (Fig. 4).

Totally 105 quartz gangue and ore samples have been collected in the field. Gold-bearing quartz veins have been sampled in the different dumps and floats (60% of the sample), and 40% in fresh outcrops around La Bellière mine, Bégrolle, and Belleville (Fig.3A) occurrences. Barren quartz veins have been sampled in the La Barbinière and Rouillère veins (Fig.3A). Sb-bearing quartz samples come from the St-Antoine vein, la Bertinière area (Fig.3A), and the active Jousselin quarry near Saint-Pierre-Montlimart. For this study, 73 thin sections have been made for microscopic studies on ores and gangue minerals using a Leica DMRX polarizing optical microscope. The accessory mineral phases were identified using a Merlin compact Zeiss Scanning Electron Microscopy (SEM), co-operated by BRGM-CNRS-Orléans University, and equipped with Energy Dispersive System (EDS) for qualitative analysis. Chemical analyses of arsenopyrite were obtained with the Cameca SX-Five Electron Probe Micro-Analyzer (EPMA) at the Institut des Sciences de la Terre d'Orléans (ISTO). The sulfides analyses were performed using an accelerating voltage of 20 kv, a beam current at 40

nA, a beam diameter of 3 μm , and a counting time of 30 s for Fe, Ni, Co, Sb, Pb, Cu, As, Zn, Ag, Bi, and 60 s for gold (detection limit around 800 ppm). The gold grain analyses were performed using an accelerating voltage of 20 kv, a beam current at 60 nA, a beam diameter of 3 μm , and a counting time of 30 s for Au, Ag, 60 s for Bi, Te, Hg, 120 s for Cu and 10 s for Sb and S.

4. Field and structural aspects

4.1 The different vein families

The orientation of 155 structures, crosscutting the regional foliation of the Neoproterozoic country rocks, recognized in the field (60% of personal observations), or interpreted from previous works (Fig. 4A-D), have been plotted on a rose-diagram (Fig.4B). Fig. 4C shows the structural data acquired on 17 outcropping quartz veins distributed all around the district. These data have been completed by a detail investigation in the Jousselin active quarry (Fig.4D-E), where well exposed quartz veins can be observed and measured. This last area offers the best outcropping conditions for kinematic determination and intersection relationships between the different vein families, and the regional fracturing (Fig.4D-E).

From the cartographic point of view, the mapped quartz veins (Fig. 4A-D) can be classified into four families according to their strikes (Fig. 4B-C-E): i) N60-90°E, ii) N20°W to N30°E, iii) N130-140°E, and iv) N90-120°E. Each family presents distinct ore contents and quartz gangue textures that will be presented in the next section. The N60-90°E and N20°W to N30°E families are associated respectively to gold deposits and Sb occurrences with a potential economic interest. On the contrary, the N90-120°E family is barren. The N130-140°E family is filled with black chalcedonic quartz with few pyrite aggregates.

4.2 The N60-90°E gold-bearing quartz veins

Economic gold ore shoots are hosted within N60-90°E-striking, 60-80°S-dipping, up to 0.5 m thick quartz veins (Fig.4) that represent opening jogs along shear zones. The infilling can be either massive

or diffuse in the fault structure, and form meter-scale quartz vein or complex stockworks. These veins are filled by greasy white-grey quartz with a minor amount of white mica. Cataclastic textures are very common (Fig.5B), whereas banded textures are rarely observed.

One of these structures has been observed in the Jousselin quarry (West of St-Pierre-Montlimart town, Fig. 4A-D-E-F and Fig.5A). The gold-bearing quartz veins are hosted in a 15m thick shear zone composed by altered and cataclased metagreywacke. This alteration consists of bleaching, silicification, and a neoformation of sericite, and sulfide (mostly pyrite). A 4 to 5m thick zone with numerous greasy white-grey quartz veins develops in the center of the cataclastic zone. Altered metagreywacke clasts are cemented by the greasy white-grey quartz and black clay. Deformed quartz veins show lenticular shapes sub-parallel to the fault structure. In detail, the quartz veins dip is variable, close to the wall rock, it is steeply southward dipping (ca 80°S), while to the top of the structure, the dip becomes subvertical. A fracture cleavage is observed near the vein walls, the metagreywacke foliation tends to become parallel to the dip of the structure, and drag folds indicate a normal component. Fault surfaces show subhorizontal slickensides with a pitch of 05° to the NE, quartz, and carbonate steps indicate a sinistral shear sense with a minor normal component that complies with previous work (Wyns 1984; Balé 1990). At macroscopic scale, the gold-bearing mineralization is characterized by deformed and anastomosed aggregates of sulfide veinlets. Arsenopyrite, pyrite, and to a lesser amount, galena, sphalerite and arsenopyrite are associated with mm-sized visible gold (Fig.5B).

4.3 The N-20°W and N10-30°E Sb-bearing quartz veins

The N10-20°W and N10-30°E striking quartz veins have been grouped together because they present the same structural features (Fig.4D-E-F). They display either a veining morphology associated with fractured, and altered host rock corridor (Fig. 5C and Fig. 6 A and B). Compared to the gold-bearing veins, the Sb ones exhibit a smaller thickness, about 0.2 to 0.5 m. Their main characteristics is a white or black chalcedonic quartz gangue (Fig.5D) with geodic or comb textures

(Fig. 5D) that argue for a formation in a brittle setting (Vearncombe, 1993; Dong et al., 1995; Simmons et al., 2005). Comb quartz growing from the vein wall toward the center, and the geodic and fibro-radiated textures argue for an open-space vein filling of these veins. In the vicinity of the quartz veins, the country rock develops a hydrothermal alteration halo of about 10 meters thick (Wyns, 1984; Fig. 5C). Close to the vein, silicification is dominant, whereas the distal part of the alteration halo is formed by sericite, sulfides and red clays. In most cases, the N10-20°W striking quartz veins present a steep dip, between 65 to 85° to the East, while the N10-30°E veins display a steep dip to the west, like the S^t-Antoine vein (Fig.5C). Whatever the vein orientation, drag folds indicate a normal component of motion. The N10-30°E slickensides with a 5° dip to the SW document a sinistral motion with a minor normal component (Fig.6A). This kinematic pattern is in agreement with the cartographic offset of the N60-90°E gold-bearing quartz veins by the N10-30°E quartz veins and faults (Fig.4A). Furthermore, detail mapping document dextral motion along the N10-20°W veins (Fig.4A-D). Sb mineralization occurs in both N10-20°W and N10-30°E striking quartz veins indicating that they formed in the same tectonic event. Thus Sb-bearing quartz veins were formed by open-space quartz vein filling mechanism controlled by strike-slip shearing. Sb develops as aggregates or acicular stibnite crystals disseminated in a black chalcedonic quartz gangue (Fig. 5D) that appears weakly mineralized. Arsenopyrite and euhedral acicular pyrite are associated to stibnite. White clay fills centimetric quartz vugs. Another type of quartz veins, close to the N10-30°E have been observed in the Jousselin quarry (Fig.4D-E-F). It shows a N0-10°E orientation with subvertical, or steeply dipping to the West, quartz comb tension gashes. They have a narrow thickness up to 5 cm without apparent offset of the wall rock. The infill is made by quartz comb with grain long direction perpendicular to the wall rock, and chalcedonic quartz with fibro-radiated quartz texture. They are rarely mineralized in acicular pyrite, however, Sb mineralization is absent. They do not postdate the Sb-bearing ones

4.4 The N130-140°E quartz veins and faults network

The N130-140°E direction is represented by two sets of structures. The first one consists of chalcedonic veins with geodic and comb quartz structures that display a 56-70° dip to the NE (Fig.4D-E-F). Several veins are located in a 0.5 to 10 m thick fault corridor. The fault planes associated to these veins exhibit subhorizontal slickensides with a pitch of 5°SE that document a dextral sense of shear with a minor normal component (Fig.6B).

The second type of N130-140°E structure is characterized by a pervasive fracture zone of 10 to 20 m thick (Fig.4D). This type of structure, only observed in the Jousselin quarry, is represented by a dense network of subvertical, NE trending joints with a 10 to 50 cm spacing. No kinematic indicator has been recognized, but on the basis of horizontal slickenside, a strike-slip displacement has been assumed by Wyns (1984).

4.5 The N85-50°S faults

N85-50°S faults, never described in previous works, have been observed in the Jousselin quarry (Fig.4D-E-F). They exhibit cataclasite and gouge clay, but no large quartz infill. Though rarely represented in the district, they have a large thickness (ca 1 m). Drag folds show a reverse motion with a northward displacement of the upper part (Fig.6C).

4.6 The N90-120°E barren quartz veins

This family is widely exposed in the field. It corresponds to the thickest quartz veins, larger than 1 m, with field evidence of shearing. For instance, the La Barbinière vein (Fig.3A) is a N100E striking, 80°S dipping structure filled by milky massive white quartz in its core (Fig.6D). The vein walls display slickenlines dipping 10° to the SE. Kinematic indicators show a sinistral strike-slip motion with a minor normal component. The entire shear structure, wider than 20 m, exhibits an alteration halo with sericite and clay whereas the quartz vein margin is silicified. In contrast with the gold-bearing veins, only a few sulfides have been observed in this structure. The mineralization hosted in these shear veins is limited to pyrite and marcasite clusters disseminated in brecciated quartz.

5. Ore and gangue mineralogy

5.1 Gold-bearing quartz veins

The La Bellière gold deposit formed throughout multistage deformations. On the basis of gangue and ore mineral textural observation, 4 main depositional stages are recognized, namely: stage 1 with early quartz gangue (Fig.7A), and syntectonic deposition of scheelite, massive arsenopyrite (Fig.7B-E), and rare white micas; stage 2 with base metals and gold (Fig.8 A to G); stage 3 with illite and dolomite (Fig.8 H); and stage 4 with pyrite, marcasite, and carbonates. The first stage is subdivided in stage 1A and stage 1B corresponding to the pre- and early-ore deposition.

Stage 1: early vein filling followed by shearing and scheelite-arsenopyrite paragenesis. Stage 1A corresponds to the initial quartz filling recorded only in some preserved parts of the vein. It is represented by weakly deformed, smoky, macrocrystalline quartz (Qz1) with abundant fluid inclusions, and weakly deformed plagioclase (Fig.7A).

During stage 1B, an heterogeneous deformation episode affected the previous assemblage. The Qz1 grains experienced a dynamic recrystallization (Stipp et al., 2002) shown by undulose extinction, subgrain rotation recrystallization with golfed or jagged shape of the Qz1 boundary, bulging and quartz neograin (Qz2) formation at the expense of the early Qz1 (Figs.7A-B; 8C-E). The Qz2 displays undulose extinction and grain size reduction. Qz2 is located along 10 to more than 100 μm shear bands that surround the edges of Qz1, or crosscut it. When Qz2 is well developed, it overprints the previous gangue leaving only relics of deformed Qz1 (Fig.7A-B). This deformation episode is coeval with deposition of massive arsenopyrite \pm pyrrhotite (Fig.7C). Arsenopyrite is the dominant sulfide of stage 1B. It shows cm-scale subeuhedral aggregates frequently associated with pyrite and muscovite. EPMA analysis of arsenopyrite does not indicate the presence of invisible gold but visible gold inclusions of 1-10 μm size are observed in arsenopyrite and pyrite crystals. This first gold occurrence is present as electrum with a mean of 27 wt % Ag and 73 wt % gold (n = 4). The silver content based on Morrison et al. (1991) shows a “medium gold fineness” of 728 (all data are reported

in Appendix A.1). For the first time, large grains of subeuhedral scheelite and apatite (up to 2 cm) have been found only close to Qz1, and plagioclase. Plagioclase alteration mainly observed by its transformation in iron dolomite occurred after the ductile/brittle deformation episode. This transformation is highlighted by relics of early plagioclase with polysynthetic twin included in iron-bearing dolomite (Fig.7F).

Stage 2: main ore stage with gold and base metals deposition. At microscopic scale, this stage is characterized by intense microfracturing and brecciation of the Qz2 and ore formed during stage 1. Arsenopyrite and scheelite with sharp grain boundaries, are highly brecciated. Fractures are filled by a new quartz generation noted Qz3 (Fig.7C-H, G) that displays a weak deformation with limited undulose extinction of hyaline elongated or fibrous grains habitus, perpendicular to sulfides or scheelite clasts, and few fluids inclusions. Fibrous grains habitus display solid or liquid inclusion trails perpendicular to the walls of the veins (Fig.7H). Comb and hyaline quartz have also been observed rarely in the same fracture with sphalerite, galena and gold. Sphalerite, displaying pyrrhotite inclusions, and galena are the dominant sulfides coeval with gold deposition (Fig.8A-B-C). Sphalerite is also present as inclusions in other minerals crystallized during stage 2. Chalcopyrite is observed either alone in the fractured arsenopyrite (Fig.8C), or in ovoidal micro-inclusions scattered in sphalerite (Fig. 8D). Such a “chalcopyrite disease” texture formed by partial exsolution of sphalerite by chalcopyrite (Barton and Bethke, 1987). Traces of ullmanite (Fig.8F) and bismuth-tellurides (Fig.8G) have been observed as inclusions in galena and chalcopyrite. Two different types of Bi-Te have been recorded: i) a Bi-Te mineral close to hedleyite (Bi_7Te_3) in chemical composition (Fig. 8G1), and ii) a Bi, Pb and Ag sulfide (Fig. 8G2) that seems either close to the vikingite-gustavite series or belonging to native bismuth with contamination during EDS analysis by traces of Pb and Ag. In La Bellière vein, only a few traces of native bismuth have been observed (Picot et al. 1981), whereas traces of Bi-Te minerals are recognized in La Rouillère vein (Fig. 2). The presence of Bi-Te phases in the two main veins (La Rouillère and Bégrolle) supports the hypothesis of a structural link

between these two veins (See Fig.3 for location). Chalcopyrite, pyrite, marcasite, and pyrrhotite are unfractured (Fig.7D, E and Fig8.A-B-C).

Gold is present as electrum (from 10 μ m to more than 1 mm in size) with a mean of 14 wt % Ag and 86 wt % gold (n = 17). The silver content is depleted compared to the first gold deposition during Stage 1B. Based on Morrison et al. (1991) electrum from the Stage 2 shows a “high gold fineness” of 860 (all data are reported in Appendix A.1). Electrum exhibits various associations : (i) trapped in the neoformed Qz2 quartz network (Fig.8E), (ii) between the deformed Qz1 quartz grains, (iii) inside neoformed quartz bands, (iv) frequent inclusions in galena, sphalerite (Fig.8A-B), and iron-dolomite, and (v) rarely present in pyrite and chalcopyrite (Fig.8C).

Iron-dolomite is commonly associated with stage 2 minerals (Fig.7D-E). It is characterized by euhedral crystals up to 100 μ m in size with numerous galena or electrum inclusions (Fig.8C). It constitutes up to 40 vol% of the gangue and has an average composition of 28.4 wt % MgO, 11 wt % CaO, 15.6 wt % FeO and 1 wt % MnO (n = 29). Analytical data on carbonates geochemistry are provided in Appendix A.2.

Stage 3: late opening stage. Brittle fracturing events postdating stage 2, were responsible for the brecciation of previous Qz2, sphalerite and galena. This fracturing represents the final brittle event coeval with filling of microcrystalline quartz (Qz4) associated with illite and low iron dolomite (18.3 wt % MgO, 30.45 wt % CaO, 5.7 wt % FeO and 1.45 wt % MnO). It is also characterized by massive deposition of pyrite and marcasite from early pyrrhotite replacement.

Stage 4: late reopening stage. At least one late brittle reopening event is recorded after stage 3. It displays the same paragenesis as stage 3, with another generation of carbonate veinlets with subeuhedral pyrite crosscutting stage 3 assemblage.

The four stages described above are synthesized on figure 9, this evolution represents increments of a single progressive crystallization-deformation evolution.

5.2 Sb-bearing-quartz veins

The mineralogic and textural analysis of the La Bellière Sb veins allowed us to subdivide the paragenetic evolution into 4 stages, namely, stage 1: early silicification; stage 2: arsenopyrite-pyrite deposition; stage 3: main stibnite deposition; and stage 4: reopening of veins.

Stage 1: early silicification. This stage starts with brecciation of the country rocks, and deposition of microcrystalline quartz (Qz1) that cemented host-rock clasts (Fig.10A). Stage 1 is not well represented in our samples. Though available observations are limited, it seems that Sb was absent from this stage.

Stage 2: arsenopyrite-pyrite deposition. This stage corresponds to a large opening event of the vein, represented by the deposition of cm sized subhedral, undeformed comb quartz (Qz2) with fibro-radiated texture. The initial microcrystalline Qz1 quartz acted as nucleus for the growth of the fibro-radiated Qz2 quartz (Fig.10A-B), which is similar to the “flamboyant quartz texture” in the classification of Dong et al., (1995). Pyrite and marcasite are intimately associated in mixing and intergrowth with Qz2 deposition (Fig.10A-C). Their sharp shape and the porous surface suggest that they formed as pseudomorphs of previous pyrrhotite crystals (Picot and Johan, 1977; Bonnemaïson and Marcoux, 1990; Epp et al., 2019). They are surrounded by euhedral rhombic arsenopyrite (Fig.10C). Isolated arsenopyrite crystals are present in the quartz gangue.

Stage 3: main stibnite deposition. This stage begins with quartz (Qz3) overgrowths on Qz2 comb quartz (Fig.10B-D). It is cloudier than Qz2, and displays golfed and corroded edge shapes (Fig.10D). Massive stibnite is the main Sb ore mineral presents as: i) infilling voids formed during stage 2, ii) corroding the Qz2 grain boundaries, and iii) penetrating the Qz2 quartz comb texture (Fig.10B-D). Stibnite is present as aggregates or acicular shape that postdates arsenopyrite, and pyrite deposition (Fig.10E). Berthierite is frequently replaced by stibnite along fractures and grain boundaries (Fig.10F) with the neoformation of xenomorphic pyrite (5-10 μ m). Rare chalcopyrite inclusions are observed in berthierite (Fig.10F). Fuloppite and micro inclusions of native Sb (2-5 μ m) appear along stibnite cleavages (Fig.10G). Visible gold has never been observed in the Sb ore.

Stage 4: reopening of veins. A late fracturing stage developed at the end of the mineralization history with the deposition of huge amount of marcasite and pyrite-marcasite upon stibnite pseudomorphs (Fig.10H). Sphalerite (Fig.10H) might belong to this stage.

The four stages described above are summarized on figure 11, this evolution represents increments of a single progressive crystallization-deformation evolution.

6. Arsenopyrite chemistry and thermometry

Chemical composition of La Bellière gold deposit arsenopyrite (stage 1) ranges from 43.69 to 44.82 wt % As (31.01-31.93 at % As), with Sb below the detection limit. Arsenopyrites from La Bellière district Sb occurrences (stage 2) display 36.03 to 41.36 wt % As (26.48-29.89 At % As), with Sb average content of 4.75 wt % (standard deviation = 1.71 wt % Sb). Arsenopyrite from Sb-bearing quartz veins in la Bellière district displays a distinct geochemical signature compared to gold-bearing quartz veins, characterized by an enrichment in Sb. Sb-rich arsenopyrite is also described in other Sb vein-type deposit in the Central-Armorican domain (Fouquet, 1980; Pochon et al., 2018) and in the Massif Central (Sandras, 1988; Bellot et al., 2003).

Arsenopyrite crystallization temperatures can be calculated using the geothermometer of Kretschmar and Scott (1976). Considering that arsenopyrite is in equilibrium with pyrite, the calculated temperature is around 430 °C for the La Bellière stage 1 gold deposit, while the average temperature for the La Bellière Sb occurrences are below 300 °C (Table.1). From the Sb occurrences, more precise estimation may be considered with caution, because of the presence of Sb in the arsenopyrite. All analytical data are reported in Appendix A.3.

7. Discussion and interpretation

7.1 Structural controls of the gold and Sb deposition

The N60-90°E sinistral strike slip fault-hosted gold-bearing quartz veins are controlled by ductile/brittle shearing, according to quartz gangue texture (Fig.7a, c). We note that the relative

chronology between the gold-bearing-quartz vein and the N90-N120°E barren ones have never been observed. However, on the basis of similar kinematics, ductile/brittle deformation record, and quartz textural characteristics of both veins, we assume that their emplacement occurred simultaneously.

Our detail mapping highlights the presence of the strike-slip N130-140°E veins and faults (Fig. 10). Its fault orientations, kinematics, vein filling, mineralogy, deformation features are consistent with the regional faults that control the location of the Evre and La Bellière rivers (Fig. 12A). The N130-140°E veins and faults display similar features with the N20°W-N30°E Sb-bearing quartz, we assume that they can be related to the same tectonic event. Moreover, the cartographic offset of the N60-90°E gold-bearing quartz veins (Fig. 12A) due to the N20°W to N30°E quartz veins and faults implies two independent tectonic events: i) an early gold event, and ii) a late Sb one.

The single-stage model proposed for the formation of the gold and Sb quartz veins cannot explain all the structural features of the entire ore district (Wyns, 1984). Our observations of highly sheared and recrystallized quartz aggregates (Qz1-Qz2), spatially associated with gold mineralization and barren N90-120°E quartz veins are incompatible with a simple tensional opening of the veins, as proposed by Wyns, (1984). The marked difference in Sb-content of arsenopyrite supports also two-stage emplacement (Fig. 12B). The subhorizontal slickensides argue for a strike-slip regime, which is accommodated by a NE-SW maximum shortening direction (Z1-a). Thus, the Z1-a direction (Fig. 12B) accounts for the development of the N60-90°E gold-bearing, and N90-120°E barren veins in sinistral shear zones.

The second deformation phase was characterized by an NNE-SSW maximum shortening direction (Z1-b in Fig. 12B) compatible with formation of the N20°W and N30°E conjugate strike-slip faults associated with Sb-quartz vein, and chalcedonic quartz gangue. Furthermore, the N130-140°E brittle dextral Evre fault complies with a Riedel model in which the Sb-bearing veins and fault correspond to the R and R' faults (Fig. 12B). In this model, the N20-10°W quartz veins with Sb ore represent the synthetic R dextral faults. The N10-35°E with Sb ore, like the St-Antoine vein correspond to the R'

sinistral faults. The Riedel shear model would also account for the N10°E chaledonic, and geodic quartz with pyrite veins, observed in the Jousselin quarry that would correspond to tension gashes (T type fracture; see Fig. 12B). The E-W striking reverse faults observed in the La Bellière mine (Fig. 3C), and in the Jousselin quarry (Fig. 6C) would correspond to the FI type structure (Fig. 12B). The FI structures cross cutting the gold-bearing quartz veins (Fig. 3C) confirm their formation after the gold-bearing event. In the current state of knowledge, this two-stage evolution model allows us to account for most of the veins, especially those with economic potential.

7.2 Mineralogical evolution and emplacement model for the gold mineralization

The gold-bearing quartz veins of the la Bellière district display a four-stage paragenesis (Fig.13). The first stage (1A) represents the early vein filling with minor amount of plagioclase with low-to null deformation. The crystallization temperature is assumed higher than 500 °C (Lang and Baker, 2001). Then, a ductile/brittle deformation stage 1B developed under an intermediate temperature, ca 400 to 500 °C, resulted in the dynamic recrystallization of the Qz1 quartz (Stipp et al., 2002). The B1 stage was coeval with scheelite and massive arsenopyrite deposition with a discrete generation of electrum with medium silver content. This range of temperature emplacement is also supported by arsenopyrite geothermometry. A fluid inclusion study performed on Qz1 quartz confirms its high temperature formation, (~400 °C) under 200 to 150 MPa (Boiron et al., 1990). Ductile/brittle deformation and high temperature of deposition argue for a deep-seated formation between 12 to 9 km (Scholz, 1988).

The transition from stages 1 to 2 (Fig.13) corresponds to another deformation characterized by grain-size micro fracturing of arsenopyrite or scheelite. The fracture-sealing by Qz3 that lacks of recrystallization texture argues for a shallower emplacement depth under a more brittle conditions compared to the stage 1 (Ramsay 1980; Cox, 1987; Cox et al., 1991). Quartz fibers and inclusions trails perpendicular to the vein walls are in favor of a crack-and-seal mechanism (Ramsay, 1980).

The pressure and temperature drop from lithostatic (stage 1) to hydrostatic regime (stage 2) at the brittle-ductile transition can be interpreted as a consequence of a fault-valve mechanism (Sibson et al., 1988; Cox et al., 1991; Cox, 1995; Robert et al., 1995; Gaboury and Daigneault, 2000; Sibson, 2001). This mechanism might explain the deposition of gold with low silver content, during stage 2 under emplacement conditions at 200-300 °C, and 7 to <5 km depth, under hydrostatic regime (Scholz, 1988). This hypothesis is consistent with our structural analysis, which has demonstrated the synchronism between the fault activity and the formation of gold-bearing quartz vein. Generally, gold deposit models controlled by fault-valve behavior, have been set up for a compressional stress field (Sibson et al., 1988; Sibson, 2001; Gaboury and Daigneault, 2000; Robert et al., 1995; Groves et al., 1998). However, this kind of tectonic regime is not observed in the la Bellière district. A strike slip stress field, coeval with the genesis of orogenic gold deposits, may also be favorable for a fault-valve mechanism, as documented by Faleiros et al. (2014) on the orogenic gold deposit of Morro do Ouro (Brazil).

7.3 Mineralogical evolution and emplacement model for the Sb mineralization

The N20°W-N30°E quartz veins, suggest an open-space quartz vein filling mechanism for their emplacement controlled by strike-slip faults motion. Deposition of microcrystalline quartz following country rocks fracturation, and deposition of comb and fibro-radiated quartz with arsenopyrite. The formation of the main Sb stage began with minor amount of berthierite and was followed by massive stibnite with minor native Sb (Fig.14). A late stage characterized by replacement of stibnite by pyrite-marcasite and sphalerite corresponds to a low temperature one recorded in the Sb-bearing quartz veins in the Vendée district (Marcoux et al., 1984). It may be related to the Pb-Zn-Ba event, of Triassic age (ca 210-230 Ma), widely recognized in the Armorican Massif (Fouquet, 1980; Marcoux et al., 1984; Chauris and Marcoux, 1994).

Comb, geodic, and fibro-radiated quartz textures, without deformation, are very similar to those observed in epithermal vein deposits formed between 50 to 1500 m depth at temperatures ranging from 100 to 300 °C (e.g. Dowling and Morison, 1988; White and Hedenquist, 1990; Dong et al., 1995; Kontak et al., 1996; Simmons et al. 2005; Chauvet et al. 2006; Shimizu 2014). In the La Bellière district, these textures suggest that Sb deposition took place at shallow crustal level, probably less than 2 km in brittle deformation conditions. The early microcrystalline quartz is interpreted as the result of crystallization from a siliceous gel during rapid drop in the fluid pressure (Bouchot et al., 1994). Fibro-radiated quartz texture corresponds to recrystallisation from early supersaturation of silica in the solution (Fournier, 1985; Dong et al., 1995). All these features, and the absence of ductilely deformed quartz grains, argue for an open-space vein infilling, as found in epithermal environments. To explain the genesis of Sb vein in such structural level and the vein formation mechanisms, we propose a model similar to the suction pump mechanism (Sibson, 1987).

7.4 Timing of the gold and Sb ore deposition

The NNE-SSW shortening direction (Z1-a) controlling the emplacement of la Bellière gold-bearing quartz veins is incompatible with the dextral motion documented along South Branch of the South Armorican Shear Zone (SASZ, Jegouzo, 1980; Rollin and Colchen, 2001; Lemarchand et al. 2011). Moreover, in the southern part of the Armorican Massif, several lines of evidence support a sinistral strike-slip faulting developed earlier than the Late Carboniferous SASZ. Firstly, the neighboring Tournaisian-Viséan Ancenis basin opened as a pull-apart controlled by the left-lateral reactivation of the Nort-sur-Erdre fault (Fig.2; Diot, 1980; Rolet, 1984; Cartier and Faure, 2004; Ducassou et al., 2011). Secondly, farther North, in the Lanvaux unit, the NW-SE striking sinistral shear zones were developed before the Late Carboniferous dextral shearing (Cogné et al. 1983; Faure and Cartier 1998). According to our structural control model, and the regional geological context, the La Bellière hydrothermal gold event would occur during the Viséan, i.e. around 340-325 Ma.

Presently, radiometric dating on U, Th rich phases coeval with ore deposition are unavailable. Unfortunately, the analyzed scheelite, apatite, and carbonates minerals yield too much common lead, to be used to obtain valuable radiometric ages.

In the Massif Central, the late Variscan gold deposits were emplaced later under a relatively short period between 310-300 Ma, defining the so-called “Or-300 event” associated to a regional NW-SE extensional direction (Marignac and Cuney, 1999; Bouchot et al., 2005).

Moreover, at the Variscan orogeny scale, two separate gold deposition events were already recognized in the Bohemian Massif. The first gold event is well established at about 350-340 Ma by numerous Re-Os molybdenite ages (Zacharias and Stein, 2001; Zacharias et al., 2001, 2014 and Zacharias, 2016). $^{40}\text{Ar}/^{39}\text{Ar}$ dating on muscovite from the Jilové deposit (Zacharias et al., 2013) yield an age of 339 ± 1.5 Ma. The Sb-Au deposits (Krásná Hora and Příčovy deposits; Němec and Zachariáš, 2018) occurring in the same area, although not dated yet, clearly postdate lamprophyre dykes (minette) that intruded at 338 ± 0.5 Ma ($^{40}\text{Ar}/^{39}\text{Ar}$ on biotite, Žák et al., 1997).

The second gold event is estimated to about 300-290 Ma, based on indirect geological evidence only, being represented mainly by the Ag-Au Roudný deposit (Zachariáš et al., 2004, 2009). This deposit differs markedly from the former group (350-340 Ma) with a highly Sb admixture in arsenopyrite (up to 4.2 wt.% Sb) and in absence of Bi-Te phases.

Based on the example of the gold mineralizations in the Bohemian massif, and our results in the La Bellière district, we can hypothesize that at least two gold deposition events might be present in the French Variscan massifs, namely an early event around 340-325 Ma, closely after the Variscan collision, and a second one, the “Or-300” event during the late-orogenic event.

On the contrary, the N-S shortening direction (Z1-b) controlling the development of Sb-bearing quartz veins complies with the dextral brittle motion along the South Armorican Shear Zone around

310-300 Ma (Jegouzo, 1980; Rollin and Colchen, 2001; Lemarchand et al. 2011). This argues for an anti-clockwise rotation of the shortening direction during the late Carboniferous period at the scale of the Mauges Nappe. This two-stage model seems more efficient, at local scale, to explain the emplacement of the late Carboniferous mineralized quartz veins compared to reconstruction tectonic model of the Armorican Massif proposed by Gumiaux et al. (2004) predicting clockwise model with NNE-SSW shortening direction for the 310-300 Ma period.

A similar North-South shortening direction is also inferred in other Sb districts of the Armorican Massif. In the Vendée district (Marcoux et al., 1984), in La Lucette district, where Z1 strikes N05-10°E (Bouchot et al., 1997), and in Quimper district, the shortening direction is oriented N30-20°W (Fouquet, 1980). The Sb veins crosscut the leucrogranite of Odet dated at 324 ± 5 Ma (U/Pb on zircon; Cocherie, 2001) and place relative time-constraints on this mineralizing event. Therefore, it might be assumed that most of the Sb deposits of the Armorican Massif were generated during the same tectonic event around 310-300 Ma. Numerous Sb deposits, widespread in the French Massif Central (Bouchot et al., 2005) or in the European Variscan belt, were also formed during this period (Bellot et al., 2003). In the Central-Armorican Domain, Le Semnon Sb-Au district seems to have been formed earlier, around 360 Ma in local compressional regime (Pochon et al., 2018). The possibility of two Sb hydrothermal events successively developed around 360 and 300 Ma cannot be ruled out.

7.5 Genetic model and comparison with other Variscan deposits

7.5.1 Gold deposits

The gold-bearing quartz veins of La Bellière district show many similarities with deep-seated gold deposits of the Massif Central like the Saint-Yrieix or Salsigne districts (Touray et al., 1989; Bonnemaïson and Marcoux, 1990; Boiron et al., 1989, 1990, 2003; Bril et al., 1994; Marignac and Cuney, 1999; Charonnat et al., 1999; Essaraj et al., 2001; Nicaud, 2001; Bouchot et al. 2000, 2005; Chauvet et al., 2012). It includes the dominant structural control with strike-slip, ductile/brittle shear

zone developed in green-schist facies sedimentary rocks, and a late emplacement of economic ores with high fineness visible gold. They are also typical of mesozonal type in the orogenic gold deposit classification (Robert et al., 1997; Groves et al., 1998, 2003; Marignac and Cuney, 1999; Bierlein and Crowe, 2000; Goldfarb et al., 2001-2005; Bouchot et al. 2005).

The high CO₂ and CH₄ concentrations recorded in stage 1 fluids inclusions (Boiron et al., 1990), and high temperature (400-500 °C) coeval ductile/brittle deformation, suggest a deeply-sourced hydrothermal fluid. The limited presence of magmatic-source-elements related ore phases such as bismuth tellurides, wolframite, lollingite or molybdenite also suggests that the ore-forming fluids are not related to a pure magmatic process (Lang and Baker, 2001), but could possibly have a strong metamorphic origin. Another heat source must be found, and could be researched in the Champtoceaux metamorphic units. Indeed, the southern part of the Champtoceaux metamorphic units experienced a partial melting event coeval with the gold-deposits emplacement time. Champtoceaux migmatites were dated around 345-335 Ma (⁴⁰Ar/³⁹Ar on muscovite; Pitra et al., 2010), and its associated magmatic products, as the La Pommeraie leucogranite at 337-336 Ma (⁴⁰Ar/³⁹Ar on muscovite; Gumiaux et al., 2004). This fertile geological context seems very similar to that established for the Saint-Yrieix gold-district in the Massif Central (Touray et al., 1989; Boiron et al., 1989, 1990, 2003; Marignac and Cuney, 1999; Essaraj et al., 2001; Nicaud, 2001; Bouchot et al. 2000, 2005), and could explain their metallogenic similarities, although their deposition events seem separated in time.

7.5.2 *Sb deposits*

The studied Sb-bearing quartz veins system would be 30-40 Ma younger than the orogenic gold deposit of La Bellière and belonging to a distinct hydrothermal event from gold one. In this model, it would be coeval with the regional high geothermal gradient, responsible for the South Armorican migmatite and leucogranite emplacement at the end of the post-collisional exhumation (e.g. Gapais et al., 1993; Brown and Dallmeyer, 1996; Turrillot et al., 2011; Ballouard et al., 2015). The shallow

emplacement of the Sb quartz veins (<5 km) with a strong structural control and “epithermal” gangue textures could correspond to a different expression of “epizonal Au-Sb mineralization” (<5 km, 130-300 °C) in the classification of the orogenic gold deposits proposed by Groves et al. (1998).

Sb-bearing quartz veins of the La Bellière district are very similar to those of other districts in the Armorican Massif such as the Vendée or Quimper ones (Fig.1). This is also true with the late Variscan deposits in the Massif Central (Périchaud 1980; Bril and Beaufort, 1989; Munoz et al., 1992; Bril et al., 1994; Marignac and Cuney, 1999; Bellot et al., 2003), and in Western Europe (Chovan et al., 1998; Wagner and Cook, 2000; Neiva et al., 2008; Nemeč and Zacharias, 2018). The Sb deposits might be ascribed to a common genetic model, but the source of the mineralizing fluid remains unknown. Due to the lack of volcanic activity during this period, lack of spatial and temporal links in the La Bellière district with the late Variscan leucogranite dykes, and the shallow crustal emplacement setting, pure magmatic fluids can be excluded.

The existence of two independent hydrothermal events associated to distinct geological contexts seems to be present in the La Bellière district with: i) an early gold-bearing one at ~ 335 Ma contemporaneous of the emplacement of the metamorphic and magmatic complex of Champtoceaux followed by ii) an Sb one at 310-300 Ma, in a more superficial emplacement condition, in relation with uplift of the basement and the hydrothermal system of the SBSASZ at this time (Lemarchand et al. 2011). Such a large time gap has not been reported in the main gold and Sb districts of the Massif Central (Saint-Yrieix, le Chatelet-Villeranges) where gold and Sb deposition are considered as parts of the same hydrothermal event (Touray et al., 1989; Boiron et al., 1989, 1990, 2003; Marignac and Cuney, 1999; Essaraj et al., 2001; Nicaud, 2001; Bouchot et al. 1997, 2000, 2005). Sb is commonly reported in several gold deposits with different accessory phases (Sb-bearing arsenopyrite, sulfosalts, stibnite...) as Laurieras deposit. In the Chatelet-Villeranges gold district, Sb may become predominant, Sb-ore corresponds to stibnite quartz-veins up to one-meter thickness. Introduction of

the Sb-ore belong to the same hydrothermal event than gold-one around 300 Ma and is considered later than the gold stage (Boiron et al., 1989).

8. Conclusion

The La Bellière district is the unique area in the Armorican Massif where the relationships between gold-bearing and Sb-bearing quartz veins can be documented.

The gold-bearing quartz veins are located along N60-90°E striking sinistral strike-slip shear zones, controlled by a NE-SW maximum shortening at 340-325 Ma. This deformation episode that occurred earlier than the dextral-strike slip motion along the SASZ, was very important for gold metallogenesis in the Ligerian domain. Fluid overpressure and associated decrease in temperature, from with early ductile/brittle transition towards a shallower depth (7 to <5 km depth) and a more brittle context triggered gold and base metals precipitation. The La Bellière gold veins exhibit characteristic features of a mesozonal orogenic gold deposit, similar to other Late Variscan gold deposits in Europe. The structural control of the La Bellière gold deposit argues for an indirect Viséan age of deposition, which must be assessed by radiogenic dating. It is proposed that at least two major gold emplacement periods are recorded in the French Variscan Massifs, i) an early one between 340-325 Ma after the Variscan metamorphic peak, and ii) a later, “Or-300”, event during the late-orogenic evolution.

Sb mineralization emplacement is separated in time from gold-bearing-quartz veins and formed at 310-300 Ma according to our interpretation of the regional strain patterns. Sb-bearing-quartz veins at the la Bellière district were developed in N10-20°W and N0-30°E conjugate strike-slip faults developed in a dextral N130-140°E dextral Riedel shear model, controlled at the Ligerian domain by the SBSASZ which acted as crustal drains for the Sb-bearing fluids.

Sb Quartz texture argues for a deposition in open-space vein filling and shallow environment (< 5 km). This event might belong to the different expression of “epizonal Sb-Au mineralization”. It displays the same features than those recognized in other late Variscan Sb deposits in Europe, formed at ca 310-300 Ma. It could belong to the “Or-300” event in the Armorican Massif.

Our study provides new elements to better target the regional exploration of orogenic gold and Sb veins in the Armorican Massif. The study of Sb occurrences is a powerful tool to constrain the cartographic shape and continuity of gold ore shoots.

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Figure Captions

Fig. 1. Simplified geological map of the Armorican Massif (modified from Chantraine et al., 1996) with the location of Sb and gold occurrences and deposits (modified from Chauris and Marcoux, 1994). NASZ: North-Armorican Shear Zone; SASZ (N) and SASZ (S): northern and southern branches of the South-Armorican Shear Zone, respectively; NSEF: Nort-sur-Erdre Fault; A: Lopérec district; B: Quimper district; C: La Lucette district; D: Le Semnon district; E: La Bellière district; F: Vendée district. Sb: stibnite; Au: gold; Cin: cinnabar; Gn: galena; Sp: sphalerite.

Fig. 2. (A) Geological and structural map of the Ligerian domain with location of the study area (Fig. 3). P: La Pommeraié 2 micas granite, C: Chemillé 2 micas granite. Modified after Cartier and Faure, (2004). (B) NW-SE cross section of the Ligerian domain showing the structural relationships between the North-and-Central Armorican Domain and the South-Armorican Domain. The Mauges Nappe is the uppermost tectonic unit in the South-Armorican nappe stack. Modified after Cartier and Faure, (2004).

Fig.3. (A) Local geological map of the la Bellière district in the Mauges Nappe with the main gold and Sb veins and occurrences (modified from Wyns et al., 1998a,1998b). (B) Stereogram (lower hemisphere projection) of the poles of the regional foliation (Fo) in the Neoproterozoic alternation of metagraywacke and metapelite with the density contours per 1% area, n: number of measurements. (C) North-South cross section across the la Bellière deposit showing the geometry of the gold-bearing quartz vein, and the late faults crosscutting the gold mineralization (modified from Sevensma, 1942).

Fig.4. (A) Structural map with the main veins and faults of the la Bellière district classified on the basis of their metal content and quartz type infilling, (personal field observations correspond to 60 % of mapped structures and compilation of historical data: Blouin and Apolinarski, 1990; Braux et al., 1991; Mulot, 1982; Wyns, 1984; Bonnemaïson and Boubée de Gramont, 1985; Wyns, 1987; Terree, 1986; Martel-Jantin, 1996). The fault kinematics are inferred from field observations and cartographic offsets of mineralized structures faults or quartz veins. See text for discussion. (B) Rose diagram of the mineralized quartz veins orientations observed on the field and inferred from previous works. Four orientation families of mineralized and barren veins and faults have been recognized. (C) Stereogram (lower hemisphere projection) of poles of mineralized and barren quartz veins measured in the field. (D) Detailed structural map of the Jousselin quarry in the la Bellière district with orientation dip, quartz and ores infilling. (E) Stereogram (lower hemisphere projection) of poles of faults, and quartz veins with ore infilling, 40 structures have been measured on the field and 15 come from Wyns, (1984). (F) Stereogram (lower hemisphere projection) showing great circle and related kinematics of

the different quartz veins and faults. Number of faults and veins measured with clear kinematics indicators = 12. See text for discussion about the kinematics.

Fig.5. Field and hand sample pictures of gold and Sb-bearing structures. (A) Gold-bearing lenticular quartz veins in a N65°E striking shear zone with hydrothermalized Neoproterozoic country rocks in the Jouselin quarry. Quartz veins contain traces of pyrite, galena, chalcopyrite and sphalerite. Kinematic indicators show a sinistral strike-slip fault with a normal component. (B) Representative ore sample of the la Bellière gold deposit, with sphalerite (Sp), galena (Gn), visible gold, and carbonates in a grey-white sheared quartz. Sample from dumps of Bonair shaft. (C) Sb-bearing chalcedonic and comb quartz vein oriented N25°E with stibnite and pyrite crosscutting the hydrothermalized Neoproterozoic country rock (S^t-Antoine vein). (D) Representative ore sample of an Sb-bearing quartz vein showing chalcedonic black quartz with dissemination of stibnite and pyrite. Sample from S^t-Antoine vein. See Figure 4A for outcrops location.

Fig.6. Field and hand sample pictures of Sb and barren veins. (A) N22°E striking subhorizontal slickenslide with chalcedonic and comb quartz with pyrite and arsenopyrite. A sinistral strike-slip movement can be inferred from kinematic criteria. (B) N140°E striking subhorizontal slickenslide with chalcedonic and comb quartz with pyrite and arsenopyrite. The slickenslide indicates a sinistral strike-slip movement. (C) Reverse fault zone striking N85E, dipping 50°S, with cataclasite and drag folds showing a top to the North shearing. (D) La Barbinière massive barren quartz vein. Subhorizontal slickenslide indicate a sinistral strike-slip movement with a normal component. See Figure 4 for outcrops location.

Fig.7. Microphotographs showing mineralogy of the initial scheelite infilling stage 1 to the gold stage (stage 2) of the la Bellière gold deposit paragenesis. Samples come from dumps of Bonair shaft (A to F) and from dumps of Verger shaft (photos G-H). (A) Weakly deformed plagioclase (Pl) and initial quartz (Qz1). The plagioclase is crosscut by a second quartz generation (Qz2) corresponding to neograins due to dynamic recrystallisation of the Qz1 during the shearing. Relic of Qz1 clasts are present in the upper right part of the photograph. (B) Initial quartz Qz1 highly sheared crosscut or in relic in the neoformed quartz (Qz2). (C) Large scheelite (Sch) grain crosscutted by galena (Gn), sphalerite (Sp), chalcopyrite (Ccp) veinlets. (D) SEM back-scattered electron image of apatite (Ap) crosscutting by galena (Gn) surrounded by iron dolomite (Dlo-Fe). (E) Relic of polysynthetic twin from an early plagioclase (Pl) in iron-bearing dolomite (Dol-Fe), and scheelite (Sch). This early paragenesis is crosscut by sphalerite (Sp), galena (Gn) and comb quartz (Qz3) veinlet belonging to the stage 2 paragenesis. (F) Scheelite (Sch) crosscut by fibrous quartz habitus (Qz3). (G) Detail of brecciated arsenopyrite crystal (Asp) cemented by fibrous quartz habitus with inclusion trails (Qz3). (H) Same microphotography as the previous one with a focus on the inclusion trails in the quartz Qz3.

Fig.8. Microphotographs of ore minerals and their relationships with gangue quartz of the gold stage (stage 2). Samples come from dumps of St-Antoine shaft (photos A-B-C), from dumps of Verger shaft (photo D), Bonair shaft (photos E-F-H). Sample from G microphotograph come from Bégrolle occurrence dump (see location in Fig. 3). (A) Clusters of fractured arsenopyrite (Asp) filled by sphalerite (Sp), electrum (El), galena (Gn) and chalcopyrite (Ccp). (B) Representative paragenesis of the main gold ore (stage 2) with clusters of arsenopyrite (Asp) fractured and filled by galena (Gn), electrum (El), sphalerite (Sp). (C) Cluster of fractured arsenopyrite (Asp) filled with sub-automorphous pyrite (Py) and chalcopyrite (Ccp). Pyrite (Py) includes electrum (El), sphalerite (Sp) and chalcopyrite (Ccp). Chalcopyrite (Ccp) also contains electrum (El) inclusions. (D) Sphalerite (Sp) with numerous micro-inclusions of chalcopyrite (Ccp). (E) High strain quartz gangue (Qz2) with numerous recrystallized grains and electrum (El) located inside the highly deformed recrystallized quartz grains (Qz2). (F) SEM back-scattered electron image of galena (Gn) with ullmannite (Ull) inclusions. (G) SEM back-scattered electron image of galena (Gn) with bismuth-tellurides inclusions, hedleyite (Hed) and Bi-Pb-Ag sulphosalt. (H) Late stage 3 with microcrystalline quartz (Qz4) with illite (Ill) and pyrite-marcasite mixing (Py/Mar2), brecciated Qz1 and Qz2 quartz.

Fig.9. Paragenetic succession of the la Bellière gold deposit. See text for details.

Fig.10. Microphotographs of ore minerals and their relationships with gangue quartz deformation observed in Sb-bearing quartz veins. Samples have been collected on the S^t-Antoine vein striking N30°E (photos, A-E) a N20°W vein (photos, B-D), close to the S^t-Antoine vein and la Bertinière occurrences striking N20°W (photos, C-F-G) located in the eastern part of the district (Fig. 3). (A) Comb quartz veinlet with pyrite and marcasite (Py/Mar) and mosaic texture showing initial microcrystalline quartz (Qz1) acting as nucleus for the growth of fibro-radiated quartz (Qz2). (B) Initial filling of a microcrystalline quartz (Qz1) with no deformation surrounding by fibro-radiated quartz (Qz2) fractured and infilling with stibnite noted Sb, (S^t-Antoine quartz vein). (C) SEM back-scattered electron image of arsenopyrite (Asp) encrusting pyrite-marcasite mixing note Py/Mar, (pseudomorph after early pyrrhotite note Po) and crystals with rhombic habitus. (D) Detail of a fibro-radiated quartz (Qz2) fractured and partially corroded by stibnite (Sb). (E) Stibnite (Sb) clusters with residual arsenopyrite (Asp). (F) Berthierite (Ber) and inclusions of chalcopyrite (Ccp) pseudomorphs by late stibnite (Sb). (G) Stibnite (Sb) with residual inclusions of berthierite (Ber) and native Sb (Sb-nat). (H) Stibnite (Sb) being replaced by late marcasite (Mar) and sphalerite (Sp).

Fig.11. Paragenetic succession of Sb-bearing chalcedonic and comb quartz veins. See text for details.

Fig.12. (A) Synthetic structural map of the la Bellière district with main mineralized quartz veins, mineralization occurrences, faults, and relative chronology. (B) Structural model showing two main deformation stages with (1) gold

and (2) Sb deposition stages during the regional left-lateral strike-slip deformation. Z1-a corresponds to the interpreted maximum shortening direction, and probably the maximum principal stress direction during gold emplacement. Z1-b corresponds to the maximum shortening direction, interpreted as the maximum principal stress direction, during Sb deposition. The theoretical Riedel-shear model is from Wilcox et al. (1973).

Fig.13. Textural sketch showing the paragenetical evolution associated to deformation processes responsible for the formation of the La Bellière gold-bearing quartz vein. The early stage corresponds to the initial quartz vein filling with deposition of weakly deformed quartz gangue with plagioclase in high temperature conditions. During Stage 1, an important ductile/brittle deformation, coeval with arsenopyrite-pyrite and minor scheelite deposition, led to the early gold paragenetic stage under lithostatic pressure, and high temperature (around 430°C according to the arsenopyrite geothermometer). After this stage 1, a sharp P-T conditions decrease occurred in the hydrothermal system with a fast evolution to a more superficial and brittle deformation setting. This change was recorded by hydrostatic pressure as indicated by the crack-and-seal texture coeval with the deposition of economic gold stage 2 with base metal under at 300-250°C temperature conditions according to mineral paragenesis. Finally, the stage 3 was a late carbonate-microcrystalline quartz-pyrite-marcasite paragenesis (Py/Mar) that indicates the end of the hydrothermal system controlled by a temperature decrease from stage 2.

Fig.14. Sketch illustrating the paragenesis evolution associated to deformation processes responsible for the formation of Sb-bearing quartz veins. Stage 1 is an early event of silification represented by brecciation of the country rock and development of microcrystalline vein filling with pyrrhotite. Stage 2 was a fracturing event that allowed the creation of voids into which deposition of fibro-radiated and comb quartz. This texture was coeval with pyrrhotite pseudomorph after pyrite and marcasite. Arsenopyrite encrusting developed under a relatively low temperature (ca; 280-300°C). The final stage 3 was a reopening event that allowed the deposition of massive stibnite in open-space filling.

Table.1. Summary of EPMA analysis on arsenopyrite and comparison of the chemical composition. Estimated crystallization temperatures of arsenopyrite from gold-bearing quartz vein (la Bellière gold deposit) and Sb-bearing quartz vein (S^t-Antoine vein and la Bertinière occurrences). The geothermometer is based on the projection of As activity (Kretschmar and Scott, 1976).

Declaration of interests

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

Highlights:

- Study describes relationships between gold-bearing quartz veins and antimony-bearing quartz veins in the French Armorican Massif.
- New structural, mineralogical and geochemical data are given.
- Two stage of emplacement are recorded with different structural control, gold bearing-quartz veins correspond to orogenic gold deposit emplaced under mesothermal environment, antimony quartz veins are shallower.

- Chronologically gold event is emplaced during Viséan age, while antimony event corresponds to Late Carboniferous metallogenic peak.

Tables and appendices

Table 1

Samples	Number of analysis	S (W%)	Std	Fe	Std	As	Std	Sb	Std	Ni	Std	Total	As (at%)	Std	T (°C) Kretschmar and Scott, 1976
Asp from la Bellière gold deposit	13	20.71	0.22	34.97	0.60	44.24	0.36	bdl	0	0.17	0.46	100.09	31.66	0.31	430°C
Asp from Sb occurrences (la Bellière district)	30	20.86	0.32	34.32	0.39	38.97	1.61	4.75	1.71	bdl	0	98.9	28.51	1.08	< 300°C

1 Electron microprobe analyses of electrum from la Bellière gold deposit.

Sample	Comment	Ag (w%)	Au (w%)	Hg (w%)	Sb (w%)	S (w%)	Total	Det.Lim Ag (ppm)	Au (ppm)	Gold finesse (Morrisson et al 1990)
DGSTP_751_A	Au in Qz2-3 (Stage 2)	11.39	86.04	bdl	bdl	bdl	97.42	716	2388	883
DGSTP_751_B	Au in Qz2-3 (Stage 2)	11.75	87.24	bdl	bdl	bdl	98.99	709	2409	881
DGSTP_751_C	Au in Qz2-3 (Stage 2)	12.97	86.98	bdl	bdl	bdl	99.95	720	2429	870
MPSTP1311_A	Au in Qz2-3 (Stage 2)	8.76	93.13	bdl	bdl	bdl	101.90	713	2375	914
DGSTP_751_D	Au in inclusion in Gn and Sp (Stage 2)	12.68	86.31	bdl	bdl	bdl	98.99	724	2514	872
DGSTP_751_E	Au in inclusion in Gn and Sp (Stage 2)	12.74	85.68	bdl	bdl	bdl	98.42	710	2444	871
DGSTP_751_F	Au in inclusion in Gn and Sp (Stage 2)	12.72	86.13	bdl	bdl	bdl	98.85	716	2468	871
DGSTP_751_G	Au in inclusion in Gn and Sp (Stage 2)	13.05	86.81	bdl	bdl	bdl	99.86	727	2462	869
DGSTP_751_H	Au in inclusion in Gn and Sp (Stage 2)	12.82	86.41	bdl	bdl	bdl	99.23	724	2459	871
MPSTP1311_B	Au in inclusion in Gn and Sp (Stage 2)	15.75	86.04	bdl	bdl	bdl	101.78	718	2394	845
MPSTP1311_C	Au in inclusion in Gn and Sp (Stage 2)	16.17	85.51	bdl	bdl	bdl	101.68	737	2364	841
MPSTP1311_D	Au in inclusion in Gn and Sp (Stage 2)	16.44	85.23	bdl	bdl	bdl	101.67	732	2378	838
MPSTP1311_E	Au in inclusion in Gn and Sp (Stage 2)	16.82	84.31	bdl	bdl	bdl	101.13	715	2434	834
MPSTP1311_F	Au in inclusion in Gn and Sp (Stage 2)	16.68	84.46	bdl	bdl	bdl	101.14	737	2381	835
MPSTP1311_G	Au in inclusion in Gn and Sp (Stage 2)	17.41	83.84	bdl	bdl	bdl	101.25	746	2421	828
MPSTP1311_H	Au in inclusion in Gn and Sp (Stage 2)	10.62	89.84	bdl	bdl	bdl	100.46	710	2437	894
MPSTP1311_I	Au in inclusion in Gn and Sp (Stage 2)	16.76	84.70	bdl	bdl	bdl	101.46	729	2411	835
DGSTP_754_A	Au in inclusion in Py and Asp (Stage 1B)	28.21	73.17	bdl	bdl	bdl	101.38	773	2388	722
DGSTP_754_B	Au in inclusion in Py and Asp (Stage 1B)	27.94	72.50	bdl	bdl	bdl	100.44	770	2344	722
DGSTP_754_C	Au in inclusion in Asp (Stage 1B)	26.71	74.49	bdl	bdl	bdl	101.20	588	1855	736
DGSTP_754_D	Au in inclusion in Asp (Stage 1B)	26.50	72.05	bdl	bdl	bdl	98.54	610	1930	731

IO	MnO	Total	Weight% Mg	Si	Ca	Fe	Mn	Total	Det.Lim ppm Mg	Si	Ca	Fe	Min
77	1.09	56.15	6.23	0.01	19.97	13.03	0.84	56.33	500	604	949	2353	2151
22	0.90	55.65	6.52	0.00	19.82	12.61	0.69	55.65	448	0	1000	2892	2391
93	0.84	56.19	6.47	0.00	19.79	13.16	0.65	56.26	518	0	898	3002	1848
1.04	56.03	6.81	0.00	20.16	12.04	0.80	0.80	56.03	637	637	974	3007	2231
0.00	1.09	55.24	6.62	0.02	20.11	11.66	0.84	55.23	463	547	923	2266	2076
35	0.99	56.70	6.07	0.00	20.21	13.49	0.77	56.69	502	595	993	1914	2220
08	1.30	57.19	7.18	0.00	20.65	11.72	1.01	57.19	517	565	910	2652	1859
09	0.69	55.76	6.24	0.01	19.73	13.28	0.54	55.76	506	582	958	2938	2535
29	1.08	56.53	6.06	0.02	20.05	13.44	0.84	56.53	458	541	944	2379	2219
94	0.87	55.63	6.21	0.00	20.39	12.39	0.67	55.63	417	0	976	3038	2419
04	0.95	55.16	6.61	0.00	20.16	11.69	0.73	55.16	456	0	959	2253	1938
44	1.18	55.32	6.68	0.00	20.46	11.23	0.91	55.58	534	0	1021	2902	1865
78	1.24	54.83	6.92	0.01	20.24	10.71	0.96	54.99	384	552	959	3414	2555
05	0.93	55.88	6.45	0.00	20.16	12.47	0.72	55.88	403	0	905	2380	1749
22	1.13	57.05	6.87	0.00	20.95	11.83	0.88	57.30	487	0	924	2378	1938
54	1.08	57.34	6.70	0.00	20.45	12.86	0.83	57.34	441	0	969	2937	2257
73	0.79	56.40	6.89	0.00	21.06	11.45	0.61	56.64	550	0	925	2252	2682
39	1.06	55.70	6.29	0.02	19.86	12.74	0.82	55.70	460	532	927	3134	1540
87	1.04	57.28	7.02	0.01	20.53	12.33	0.80	57.37	450	553	875	2617	1936
08	1.06	55.38	6.38	0.00	20.48	11.72	0.82	55.38	491	592	1026	2388	1310
57	0.98	56.25	7.30	0.03	20.40	11.33	0.76	56.42	410	522	956	1971	1756
39	1.12	56.22	6.77	0.00	20.36	11.97	0.86	56.29	495	0	936	2837	2106
42	1.29	55.42	7.12	0.01	20.64	10.43	1.00	55.49	483	532	944	3046	1947
54	1.08	57.34	6.70	0.00	20.45	12.86	0.83	57.34	441	0	969	2937	2257
73	0.79	56.40	6.89	0.00	21.06	11.45	0.61	56.64	550	0	925	2252	2682
39	1.06	55.70	6.29	0.02	19.86	12.74	0.82	55.70	460	532	927	3134	1540
87	1.04	57.28	7.02	0.01	20.53	12.33	0.80	57.37	450	553	875	2617	1936
99	0.85	55.92	6.98	0.01	20.36	11.65	0.66	55.92	426	524	928	2124	1552
19	1.19	57.20	7.19	0.00	20.73	11.72	0.92	57.20	454	0	947	2731	1755
49	1.00	55.85	7.27	0.01	20.21	11.26	0.78	55.85	431	586	912	2956	1554
63	0.97	58.24	6.81	0.02	20.93	12.93	0.75	58.23	418	594	919	2193	2345
84	2.55	57.37	10.84	0.01	22.15	4.54	1.98	57.37	390	482	961	2136	2475
195	3.95	55.38	10.88	0.00	22.24	3.33	1.51	55.48	373	0	1017	3194	1771
18	1.65	56.54	11.61	0.01	22.47	3.25	1.28	56.54	377	517	962	1997	1043
71	1.32	57.12	12.14	0.00	22.85	2.89	1.02	57.12	412	0	998	2883	1351
01	1.27	56.25	10.01	0.24	21.34	6.22	0.98	56.32	389	537	953	2405	1580
93	1.38	56.28	10.93	0.01	22.03	4.61	1.07	56.28	481	594	993	2424	1427
16	1.27	57.06	11.38	0.03	22.66	4.01	0.98	57.06	379	516	912	2283	1997
47	1.38	57.36	11.60	0.00	22.36	4.25	1.07	57.36	408	593	1066	2418	2629
19	1.53	56.72	11.44	0.01	22.16	4.04	1.18	56.72	459	506	996	2282	2170
57	1.32	56.46	11.31	0.02	21.99	4.33	1.03	56.46	381	528	1035	2536	2330
47	1.52	57.03	11.33	0.00	22.34	4.25	1.18	57.03	376	0	955	1841	2338
19	1.75	55.66	10.90	0.01	21.89	4.03	1.35	55.66	503	470	1031	2978	1804
62	1.50	56.56	11.10	0.00	22.18	4.37	1.16	56.55	393	0	945	2653	2756
97	1.23	54.79	9.98	0.00	20.75	6.20	0.95	54.79	430	0	1043	2647	1579
30	1.28	55.98	11.24	0.03	21.94	4.12	0.99	55.98	403	559	1013	1464	2331
46	1.49	57.35	11.96	0.03	22.42	3.58	1.15	57.35	399	560	952	1661	1809
46	1.19	56.98	11.19	0.01	21.26	5.80	0.92	56.98	398	582	1064	2976	2744
52	1.25	54.63	10.22	0.07	20.56	5.85	0.97	54.71	394	528	1000	2276	2155
11	1.38	54.31	10.39	0.04	21.08	4.75	1.07	54.39	455	482	1020	2651	1586
94	1.23	54.62	10.99	0.03	21.55	3.84	0.95	54.62	434	446	952	1997	2333
06	1.27	55.68	10.82	0.00	21.73	4.71	0.99	55.68	478	0	944	1834	2477
72	1.34	57.11	11.09	0.00	21.30	4.44	1.04	55.26	370	0	973	2143	2621
92	1.73	57.11	11.42	0.00	22.53	3.83	1.34	57.11	369	0	1013	2544	2179
85	1.29	55.83	11.01	0.01	21.73	4.55	1.00	55.82	416	599	995	2411	2620
63	1.43	54.50	10.82	0.00	21.07	4.38	1.11	54.50	526	0	924	1994	1993

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Appendix A.3 Electron microprobe analyses of arsenopyrite from la Bellière district.

Location	Sample	Weight% S	Fe	As	Sb	Ni	Total	Atomic% S	Fe	As	Sb	Ni	Total	Det. Lim ppm S	Fe	As	Sb	Ni
La Bellière gold deposit	DG754_Asp2	21.03	35.47	44.10	bdl	bdl	100.60	34.90	33.79	31.31	0	0	100	327	322	815	0	0
La Bellière gold deposit	DG754_Asp3	20.80	35.43	44.82	bdl	bdl	100.70	34.57	33.80	31.63	0	0	100	299	314	810	0	0
La Bellière gold deposit	DG754_Asp4	20.54	35.47	44.82	bdl	bdl	100.83	34.19	33.89	31.93	0	0	100	317	305	819	0	0
La Bellière gold deposit	DG754_Asp5	20.70	35.23	44.64	bdl	bdl	100.57	34.48	33.69	31.83	0	0	100	319	317	828	0	0
La Bellière gold deposit	DG754_Asp6	21.12	35.66	43.69	bdl	bdl	100.47	35.04	33.96	31.01	0	0	100	349	321	810	0	0
La Bellière gold deposit	FCSTP209A_Asp1	20.40	35.05	44.55	bdl	bdl	100.00	34.23	33.77	32.00	0	0	100	341	317	809	0	0
La Bellière gold deposit	FCSTP209A_Asp1	20.51	34.78	44.54	bdl	bdl	99.83	34.45	33.54	32.02	0	0.16	100	349	313	817	0	0
La Bellière gold deposit	FCSTP209A_Asp5	20.75	34.81	43.97	bdl	0.18	99.70	34.78	33.50	31.55	0	0.77	100	323	321	819	0	367
La Bellière gold deposit	FCSTP209A_Asp6	20.82	33.78	43.89	bdl	0.84	99.33	35.02	32.61	31.59	0	0.06	100	332	314	815	0	365
La Bellière gold deposit	FCSTP209A_Asp7	20.37	33.77	44.44	bdl	1.01	99.58	34.34	32.68	32.06	0	0.93	100	336	319	824	0	358
La Bellière gold deposit	FCSTP209A_Asp8	20.62	35.19	44.17	bdl	bdl	99.98	34.53	33.82	31.65	0	0	100	322	318	828	0	0
La Bellière gold deposit	FCSTP209A_Asp9	20.77	35.18	43.86	bdl	0.07	99.88	34.75	33.79	31.40	0	0.06	100	331	319	816	0	367
La Bellière gold deposit	FCSTP209A_Asp10	20.77	34.85	44.00	bdl	0.06	99.68	34.83	33.55	31.57	0	0.05	100	334	315	812	0	356
Sb occurrence of la Bertinière	FCSTP239_Asp1	20.47	34.46	39.32	4.57	bdl	98.81	35.12	33.95	28.87	2.06	0	100	339	317	794	558	0
Sb occurrence of la Bertinière	FCSTP239_Asp2	21.30	34.79	39.04	4.08	bdl	99.20	36.07	33.82	28.29	1.82	0	100	318	323	792	551	0
Sb occurrence of la Bertinière	FCSTP239_Asp3	20.58	34.80	41.16	2.24	bdl	98.78	35.03	34.00	29.97	1.00	0	100	321	320	810	525	0
Sb occurrence of la Bertinière	FCSTP239_Asp4	21.42	34.52	37.53	5.28	bdl	98.74	36.50	33.77	27.36	2.37	0	100	344	311	800	546	0
Sb occurrence of la Bertinière	FCSTP239_Asp5	20.69	34.21	36.63	7.30	bdl	98.82	35.72	33.90	27.06	3.32	0	100	343	319	789	562	0
Sb occurrence of la Bertinière	FCSTP239_Asp6	21.03	34.43	36.03	7.38	bdl	98.87	36.17	33.98	26.51	3.34	0	100	319	329	783	551	0
Sb occurrence of la Bertinière	FCSTP239_Asp7	20.68	34.53	39.38	4.70	bdl	99.29	35.30	33.83	28.76	2.11	0	100	324	313	787	547	0
Sb occurrence of la Bertinière	FCSTP239_Asp8	20.73	34.69	39.94	3.87	bdl	99.22	35.29	33.89	29.09	1.73	0	100	328	325	798	525	0
Sb occurrence of la Bertinière	FCSTP239_Asp9	21.12	35.00	38.90	3.96	bdl	98.98	35.85	34.11	28.26	1.77	0	100	332	325	788	547	0
Sb occurrence of la Bertinière	FCSTP239_Asp10	21.66	33.87	37.88	6.46	bdl	99.87	36.71	32.94	27.47	2.88	0	100	326	318	806	554	0
Sb occurrence of la Bertinière	FCSTP239_Asp11	21.11	34.44	36.99	5.85	bdl	98.40	36.24	33.94	27.18	2.64	0	100	328	327	787	551	0
Sb occurrence of la Bertinière	FCSTP239_Asp12	20.71	34.43	41.29	2.68	bdl	99.12	35.20	33.58	30.02	1.20	0	100	320	318	803	535	0
Sb occurrence of la Bertinière	FCSTP239_Asp13	20.72	34.74	38.03	5.64	bdl	99.13	35.47	34.13	27.85	2.54	0	100	319	331	791	548	0
Sb occurrence of la Bertinière	FCSTP239_Asp14	20.67	34.58	38.10	5.60	bdl	98.94	35.46	34.05	27.96	2.53	0	100	343	314	794	544	0
Sb occurrence of la Bertinière	FCSTP239_Asp15	20.20	34.33	39.57	4.72	bdl	98.82	34.78	33.93	29.15	2.14	0	100	312	322	807	551	0
Sb occurrence of la Bertinière	FCSTP239_Asp16	20.92	34.76	40.02	3.50	bdl	99.21	35.50	33.87	29.07	1.57	0	100	321	313	794	537	0
Sb occurrence of la Bertinière	FCSTP239_Asp17	20.78	34.35	40.97	2.93	bdl	99.02	35.34	33.54	29.82	1.31	0	100	330	322	803	550	0
Sb occurrence of la Bertinière	FCSTP239_Asp18	20.77	34.16	41.01	2.60	bdl	98.54	35.44	33.45	29.94	1.17	0	100	312	320	802	519	0
Sb occurrence of la Bertinière	FCSTP239_Asp19	20.44	34.17	40.79	3.75	bdl	99.15	34.94	33.53	29.84	1.69	0	100	336	324	798	534	0
Sb occurrence of la Bertinière	FCSTP239_Asp20	21.03	34.78	40.25	3.18	bdl	99.23	35.61	33.81	29.16	1.42	0	100	320	317	808	559	0
Sb occurrence of la Bertinière	FCSTP239_Asp21	20.86	34.59	41.36	2.76	bdl	99.57	35.27	33.58	29.92	1.23	0	100	340	316	813	540	0
Sb occurrence of la Bertinière	FCSTP239_Asp22	21.11	33.74	37.20	7.01	bdl	99.06	36.25	33.25	27.33	3.17	0	100	336	321	792	568	0
Sb occurrence of la Bertinière	FCSTP239_Asp23	20.86	33.72	36.54	7.96	bdl	99.08	35.99	33.40	26.98	3.62	0	100	315	325	800	566	0
Sb occurrence of la Bertinière	FCSTP239_Asp24	21.05	33.85	37.00	7.43	bdl	99.32	36.13	33.34	27.17	3.36	0	100	326	320	786	558	0
Sb occurrence of la Bertinière	FCSTP239_Asp25	20.55	34.19	39.40	5.21	bdl	99.35	35.18	33.60	28.87	2.35	0	100	339	325	801	553	0
Sb occurrence of la Bertinière	FCSTP239_Asp26	20.99	34.12	36.83	7.57	bdl	99.51	35.98	33.58	27.02	3.42	0	100	343	330	789	568	0
Sb occurrence of la Bertinière	FCSTP239_Asp27	20.54	33.65	39.21	4.39	bdl	97.78	35.54	33.42	29.03	2.00	0	100	329	318	798	540	0
Sb occurrence of la Bertinière	FCSTP239_Asp28	20.97	33.41	38.24	4.23	bdl	96.85	36.39	33.28	28.39	1.93	0	100	314	314	788	541	0
Sb occurrence of la Bertinière	FCSTP239_Asp29	21.29	34.19	40.18	2.92	bdl	98.58	36.16	33.33	29.20	1.30	0	100	321	312	805	542	0
Sb occurrence of la Bertinière	FCSTP239_Asp30	20.51	34.30	38.63	4.75	bdl	98.19	35.38	33.95	28.51	2.16	0	100	337	322	797	535	0
Sb occurrence of la Bertinière	FCSTP239_Asp31	20.81	34.17	40.80	2.65	bdl	98.43	35.53	33.48	29.80	1.19	0	100	330	314	814	533	0

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