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**Plio-Pleistocene actinopterygian fishes from Santa Maria Island (Azores:
North-eastern Atlantic Ocean): palaeoecological and
palaeobiogeographical implications**

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Abstract

Fossil fish are among the rarest in volcanic oceanic islands, their presence providing invaluable data for the understanding of more general (palaeo)biogeographical patterns and processes. Santa Maria Island (Azores Archipelago) is renowned for its palaeontological heritage, with representatives of several phyla, including the Chordata. Herein, we report on the fossil fishes, resulting in an increase of the number of Pliocene fish from the Azores to 11 taxa: 7 Chondrichthyes and 4 Actinopterygii. The genus *Sparisoma* is reported for the first time in the fossil record. The presence of fossil remains of the parrotfish *Sparisoma cretense* in Last Interglacial outcrops is significant, because it posits a setback for the theory that most of the present-day Azorean marine species colonized the area after the last glacial episode. Our multidisciplinary approach combines palaeontological data with ecological and published genetic data, offering an alternative interpretation. We suggest that most of the Azorean shallow-water subtropical and temperate marine species living in the archipelago during the Last Interglacial were not affected by the decrease in sea surface temperatures during the last glacial episode. We also predict low genetic diversity for fish species presently living in the Azores and ecologically associated with fine sediments. These are viewed as post-glacial colonizers or as ‘bottleneck’ survivors from the Last Glaciation.

Keywords

Actinopterygii; Azores archipelago; glacial-interglacial cycles; sea-surface temperature; *Sparisoma cretense*; *Labrodon pavementatum*; Sparidae

48 **Introduction**

49 Oceanic islands have long been considered as ideal places to study biogeographical
50 patterns and processes, and to test evolutionary theories (Darwin, 1859; Wallace,
51 1880). A detailed knowledge of an island's geological history and evolution is of
52 paramount importance, influenced by the ontogenic state of the island as determined
53 by a series of factors [e.g. geological age of the oldest subaerial lavas, terrestrial area,
54 maximum altitude, latitude; see Whittaker et al. (2010); Fernández-Palacios et al.
55 (2016)]. Such factors ultimately regulate and explain both the number of species that
56 live on the island, as well as the functional composition of the biological assemblage
57 itself. The marine fossil record for volcanic oceanic islands may be considered as rare
58 or poor, especially when compared to that for vastly larger continents (Ávila et al.,
59 2018). However, the dating of an island's fossil fauna and flora when coupled with
60 checklists validated by experts provides valuable insights for a number of disciplines
61 such as palaeobiogeography, palaeoceanography, phylogenetic and climate-change
62 studies (Ávila et al., 2019).

63 The remarkable marine fossil record of Santa Maria Island in the Azores
64 Archipelago has been intensively studied over the last decades (for a review, see Ávila
65 et al. (2018) and references therein). The result has been 204 Pliocene and 146
66 Pleistocene (Last Interglacial) species and 15–20 further ichnospecies reported for the
67 island (Santos et al., 2015; Uchman et al., 2016, 2017; Raposo et al., 2018). The vast
68 majority of specimens collected are invertebrates: molluscs (Zbyszewski & Ferreira,
69 1962; Ávila et al., 2002, 2015a, 2015b, 2016; Janssen et al., 2008), brachiopods (Kroh
70 et al., 2008), crustacean balanids (Winkelmann et al., 2010), echinoderms (Madeira et
71 al., 2011) and ostracods (Meireles et al., 2012). A small number of fossil vertebrate

remains also have been studied, most notably cetaceans (Estevens & Ávila, 2007; Ávila et al., 2015c) and sharks (Ávila et al., 2012). Fossil bony fishes are known from the island, but the last account for Santa Maria fossil Actinopterygii fishes' dates from the 1950's and 60's (Zbyszewski & Ferreira, 1962; Ferreira, 1955). With few exceptions (Ávila et al., 2012; Betancort et al., 2016), the majority of the published palaeoecological and palaeobiogeographical studies on the marine fauna and flora of the Macaronesian archipelagos targeted sessile or low-motility species, such as the marine molluscs, echinoderms, bryozoans, barnacles, ostracods and algae (Meco, 1977; Meco et al., 1997, 2015, 2016; Ávila et al., 2008a, 2008b, 2009; 2015a; Winkelmann et al., 2010; Madeira et al., 2011; Santos et al., 2011, 2012a, 2012b; Meireles et al., 2012; Ávila et al., 2013; Baarli et al., 2013, 2017; Betancort et al., 2014; Johnson et al., 2014; Meireles et al., 2014; Rebelo et al., 2014; Tuya et al., 2017). Still largely unresolved, an interesting question is how the evolutionary patterns and processes inferred from these studies do relate with species' dispersal capabilities (Ávila et al., 2019). Therefore, this study aims to: 1) review and update the status of all existing Actinopterygii fossil material from Santa Maria Island; 2) use this new data to test previous statements by several authors (e.g., Briggs, 1974; Santos et al., 1995; Domingues et al., 2006, 2008) who concluded that the shallow littoral fauna of the Azores was extirpated during the Last Glacial episode; and 3) compare the palaeobiogeographical and palaeoecological patterns known from sessile and low-dispersive species (e.g. non-planktotrophic marine molluscs), with that of highly motile species such as fishes.

Geographical and Geological Setting

96 *Azores Archipelago*

97 The nine volcanic oceanic islands that form the Azores Archipelago are located in the
98 north-eastern Atlantic Ocean (Fig. 1) and rise from a bathymetric anomaly (the Azores
99 Plateau) with depths of about 2,000 m. The most distant islands (Flores, in the western
100 group, and Santa Maria in the eastern group) are 615 km apart. Flores and Corvo
101 belong to the North American tectonic plate, whereas the other seven islands are
102 located around the most-studied area called the Azores Triple Junction (Laughton &
103 Whitmarsh, 1974), a place where three major tectonic plates, North American, Nubian
104 and Eurasian, interact. The Azores Triple Junction is limited by the East Azores Fracture
105 Zone to the south, the Mid-Atlantic Ridge (MAR) to the west and the Terceira's Rift and
106 Gloria Fault to the northeast and east (Searle, 1980) (Fig. 1A). This area (i.e. the Azores
107 Plateau, the Azores Triple Junction and the islands that compose the Azores
108 Archipelago), was named the 'Azores Geosyndrome' (Vogt & Jung, 2018), which is
109 derived from a series of anomalies related to 'crustal thickness, rock composition,
110 basement depth, plate boundary morphology, seismicity, gravity and geoid, and upper
111 mantle seismic velocity structure'.

112

113 *Santa Maria Island*

114 This island has a complex geological history (see Ramalho et al. (2017) for a review). It
115 is the oldest island of the Azores Archipelago, first rising above sea level at about 6 Ma,
116 initially by Surtseyan activity (the Cabrestantes Formation), and then by subaerial,
117 monogenetic volcanism associated to the Porto Formation. High rates of magma
118 production formed a shield volcano (the Anjos volcanic complex, dating from 5.8 to 5.3
119 Ma), which substantially increased the area of the Santa Maria protoisland, as attested

by the present-day insular shelf (Ricchi et al., 2018). Although the entire volcanic edifice likely has undergone subsidence since extrusion of the first lavas of the submarine edifice, at around 5.3 Ma, a prolonged subsidence rate of about 100 m/myr lasted until 3.5 Ma (Ramalho et al., 2017). Volcanic activity associated with the Anjos shield volcano ceased at about 5.3 Ma and was followed by a period of about 1.2 myr with almost no volcanism, during which the protoisland nearly vanished due to erosion (Ávila et al., 2012; Ramalho et al., 2017), thus originating a large, flat, shallow seamount. The sediments of this unit (the Touril volcano-sedimentary complex, 5.30 to 4.13 Ma) are highly fossiliferous and correspond to early Pliocene outcrops formerly described in the literature as of Miocene in age (see Madeira et al. (2007) for a review). Volcanic activity initiated again at about 4.1 Ma, during the first rejuvenated stage, with eruptive foci associated with the Pico Alto volcanic successions, first submarine and later subaerial in nature (Ramalho et al., 2017), thus originating the revival of Santa Maria as an island. This stage lasted to 3.5 Ma, when the subsidence trend was reversed, and the island's volcanic edifice initiated an uplift trend continuing to the present day. Initially, a rate of 59 m/myr occurred from 3.5 to 2.15 Ma, thereafter at a slower rate of 42 m/myr between 2.15 Ma and the present (Ricchi et al., 2018). This uplift trend, coupled with Pleistocene glacio-eustatic sea-level fluctuations, produced a staircase of 10 subaerial and of 5 submerged marine terraces, the former particularly evident across the island's western sector at elevations ranging from 7–11 to 210–230 m (Ramalho et al., 2017), and the latter on the shelf all around the island at depths ranging between -40/-50 and -120/-140 m (Ricchi et al., 2018). A second rejuvenated stage occurred from 3.2 to 2.8 Ma, associated with monogenetic volcanism of the Feteiras Formation. Uplift has continued until today, but erosion

became the dominant agent impacting the island's landscape during the last 2.8 myr (Ramalho et al., 2017).

Institutional Abbreviations

DBUA-F: fossil collection of the Department of Biology of the University of the Azores (Portugal).

LAQ-F: fossil collection of the Liceu Antero de Quental, Ponta Delgada, São Miguel Island (Azores, Portugal).

MGM: fossil collection of the former Museu Geológico e Mineiro, now Museu Geológico, LNEG, Lisbon (Portugal).

Methods

Most of the 20 known fossiliferous outcrops on Santa Maria Island are early Pliocene in age, but scattered Last Interglacial deposits from the Pleistocene also are known from the north (Lagoinhas; Ávila et al., 2009a, 2009b) and southern shores (Prainha, Praia do Calhau and Vinha Velha outcrops (Ávila et al., 2002, 2015a; Fig. 1B). The fossil fish specimens reported, herein, were found in sedimentary layers from a total of eight outcrops (cf. Table 1). Six of these contain fauna that deposited during the early Pliocene: Cré (Kroh et al., 2008; Fig. 2A), Figueiral (Rebelo et al., 2016a; Ávila et al., 2018a; Fig. 2B), Malbusca (Rebelo et al., 2016b; Johnson et al., 2017; Fig. 3A), West Fault of Malbusca (Uchman et al., 2017 Fig. 3B), Pedra-que-pica (Ávila et al., 2015b, 2018a; Uchman et al., 2016; Fig. 3C), and the yet undescribed Ponta dos Frades outcrop. The two remaining outcrops [Prainha (Fig. 2C) and Vinha Velha (Fig. 2D)] include faunas from the Last Interglacial period (c. 116–130 kyr), of which the warmest

interval is known as MIS 5e (Marine Isotopic Stage 5e) (Ávila et al., 2015a). Figure 1B shows the geographical location of each outcrop and Figure 4 displays representative composite sections of the studied outcrops.

All specimens described were collected during fourteen fieldwork campaigns of the *Palaeontology in Atlantic Islands* workshops, held on Santa Maria Island from 2002 to 2017. In compliance with the legislation that rules the “PalaeoPark Santa Maria”, all fossil specimens collected were deposited in the fossil collection at the Department of Biology of the University of the Azores (DBUA-F). Over 1,300 lots containing fossil specimens collected at Santa Maria Island were screened for fossil Actinopterygii fishes, yielding a total of 42 lots. Systematics for Actinopterygii conform to the World Register of Marine Species (WoRMS, 2018). As the oral and pharyngeal tooth plates are frequently well preserved in the fossil record and as they display diagnostic features (Bellwood & Schultz, 1991), direct comparisons of the fossil remains with fossil and recent remains (whenever available) were used to classify our material. The teeth terminology for Scaridae is in accordance with Day (2002).

All data generated or analysed during this study are included in this published article.

Results

Actinopterygian fish taxa were found in the fossiliferous sediments of Santa Maria Island: at least two extinct Sparidae species with a probable Mio-Pliocene temporal range (Fig. 5A-Q); *Sparisoma cretense* (Linnaeus, 1758) (Fig. 6), an extant Scaridae herein reported for the Pliocene and Pleistocene (Last Interglacial, MIS 5e) of the

Azores; and *Labrodon pavimentatum* Gervais, 1857 (Fig. 7), an extinct Labridae with a Mio-Pliocene temporal range. In total, 54 teeth and several pharyngeal plates were found.

SYSTEMATIC PALAEONTOLOGY

Class ACTINOPTERYGII Klein, 1885

Order PERCIFORMES Bleeker 1859

Family SPARIDAE Rafinesque, 1818

Genus and species indet.

Figure 5A–Q

Material Examined—54 teeth in total: Pedra-que-pica (26 teeth), Malbusca (17), Cré (4), West Fault of Malbusca (2), Figueiral (1): DBUA-F 165-1, 165-2, 204-2, 205-2, 250-1, 260, 286, 313, 344-2, 350, 382, 448, 460-1, 477-2, 508-1, 508-2, 595, 596, 597, 598, 599, 601, 647, 667, 668, 669, 671, 672, 673, 674, 966-2, 991, 1014, 1017, 1073, 1083. Cré, formerly also known as Pedreira dos Frades (2): MGM 11311, 11312 (one spheroidal and one oval molariform teeth, respectively); Santa Maria Island (4 teeth), no locality: MGM 11313.

Description—The 50 teeth examined from the DBUA collection represent three types of dentition: molariform, conical and incisiform teeth. We have examined a total of 39 molariform teeth, corresponding to the lingual series (36 spheroidal in shape (Fig. 5A-F, 5L), ranging from 2.87 to 17.34 mm in diameter; and 3 oval in shape, 5.89–11.51 mm), 9 conical teeth (Fig. 5G-K), corresponding to the labial series (2.41–7.93 mm in diameter) and 2 ‘spatulate’ incisiform teeth (1.45–1.53 mm in diameter; DBUA-F 668, 671). Some

of the largest spheroidal molariform teeth exhibit a central depression; these are considered to belong to the upper lingual series (Fig. 5A-C, 5F). Some of the largest teeth show radial grooves at the base of the crown. The general shape and size resemble some of the teeth described by (Betancort et al., 2016).

The four teeth housed at MGM (Lisbon) that we examined are all lateral teeth, with a conical shape, and rounded, worn crowns (Fig. 5M-Q).

Remarks—this material represents at least two taxa, including a small one and a larger one (possibly *Diplodus* and *Pagrus*, respectively). Some of these teeth might also belong to *Archosargus* (Cutwa & Turingan, 2000). However, any generic/specific assignment would be highly tentative without any additional osteological evidence.

Stratigraphic and geographic range—Extinct *Archosargus* species are reported from the Miocene and Pliocene of the Mediterranean Sea (Obrador & Mercadal, 1973; Vicens & Rodríguez-Perea, 2003; Mas & Antunes, 2008) and from the Miocene of France (Gagnaison, 2017). In the Atlantic, *Archosargus* species are reported from the Miocene of Portugal (Zbyszewski & Moitinho d’Almeida, 1950; Jonet et al., 1975) and Morocco (Lecointre, 1952), the Pliocene of Canary Islands (Gran Canaria Island) (Betancort et al., 2016) and from the Pliocene of the Azores (Santa Maria Island) (Zbyszewski & Moitinho d’Almeida, 1950; Ferreira, 1955; Zbyszewski & Ferreira, 1962; Ávila et al., 2015b). The extinct species formerly known as *D. jomnitanus* is reported from the Miocene of France (Gagnaison, 2017), Portugal (Zbyszewski & Moitinho d’Almeida, 1950; Jonet et al., 1975) and the Mediterranean Sea (Vicens & Rodríguez-Perea, 2003), and from the Pliocene of the Azores (Santa Maria Island) (Zbyszewski & Moitinho d’Almeida, 1950; Ferreira, 1955; Zbyszewski & Ferreira, 1962). Extinct *Pagrus* species are reported from Mio-Pliocene

sediments from the Mediterranean Sea and from Europe (Arambourg, 1927; Bauzá, 1948; Menesini, 1969; Bauzá & Plans, 1973; Jonet, 1975; Antunes et al., 1981).

Family LABRIDAE Cuvier, 1816

Tribe SCARINI Rafinesque, 1810

Genus *SPARISOMA* Swainson, 1839

Sparisoma cretense (Linnaeus, 1758)

Figure 6

Material Examined—Seven pharyngeal plates (4 lower and 3 upper) plus one fragmentary tooth plate: Ponta dos Frades (3 lower and 3 upper pharyngeal plates), Vinha Velha (1 lower pharyngeal plate) and Pedra-que-pica (one fragmentary tooth plate): DBUA-F 825, 1030, 1069, 1214-1 to 1214-5.

Description—This species has a jaw formed from modified hindmost gill-arch, with tooth-bearing bones above, denominated upper pharyngeal plates, and a large, single bone below, the lower pharyngeal plate. Our fossil specimens' pharyngeal plates are very similar to those of Recent specimens of *Sparisoma cretense*. The tooth rows of the lower pharyngeal plate range from 11 to 12 in the fossil material, and from 10 to 14 in the recent material. As all fossil lower pharyngeal plates examined are smaller than the recent ones (see below), we attribute the intraspecific variation of the pharyngeal jaw elements (i.e., the lower number of tooth rows in the fossil material, as well as the slightly different angles of the lateral bony articulation in smaller fossil specimens) to ontogeny. A fresh specimen of *S. cretense* was weighted and measured (1.585 kg; 45.50 cm in total length; 40.50 cm standard length). Its lower pharyngeal plate was 4.11 cm wide, which compares with those from fossil material, which measured 2.62 cm wide

(DBUA-F 1030), 2.82 cm (DBUA-F 1214-1), 2.84 cm (DBUA-F 1214-2) and an inferred 3.82 cm (DBUA-F 825 broken specimen measuring 2.84 cm). The fragmentary tooth plate collected at Pedra-que-pica (DBUA-F 1069) measures 1.22 cm.

Today, there are 15 species of *Sparisoma* in the Atlantic Ocean and the Mediterranean Sea (Table 2). Of these, only 3 species are reported for the Eastern Atlantic: two tropical species (*Sparisoma choati* Rocha, Brito & Robertson, 2012, reported for Cabo Verde, São Tomé and Príncipe and to the tropical west African shores; *S. frondosum* (Agassiz, 1831), endemic to Cabo Verde); and the eurythermic *S. cretense*, with a wider geographical distribution, occurring in all the Macaronesian archipelagos, the Mediterranean Sea, Iberian shores and Senegal (Abecasis et al., 2008; Randall, 1990).

Ecology—The parrotfish *Sparisoma cretense* is an omnivorous daytime feeder species with specialized pharyngeal bones and muscles (Bullock & Monod, 1997). It is a sexually dimorphic species that lives on rocky bottoms covered by algae and sea grass beds, where it feeds on algae, sea grass and small invertebrates with its fused, beak-like jaw (Quignard & Pras, 1986; Abecasis et al., 2008). This shallow-water species is common to about 50 m depth (Guidetti & Boero, 2002).

Stratigraphic and geographic range—This extant species is presently reported from the Mediterranean Sea and from the eastern Atlantic shores of Portugal, the Azores, Madeira, Canary Islands, Cabo Verde and Senegal (Abecasis et al., 2008; Randall, 1990). A search for fossil *Sparisoma* species on the PaleoBiology Database and on Bellwood et al. (2019) yielded no occurrences. With the exception of otoliths from the Late Eocene of France referred to *Sparisoma* [Nolf (1988) but see remarks by Bellwood et al. (2019), in whose opinion these otoliths are not secure evidence of *Sparisoma*], this genus is reported herein for the first time from the fossil record based on dental remains. *S.*

cretense is herein reported from the Pliocene (Pedra-que-pica and Ponta dos Frades outcrops) and Pleistocene (Vinha Velha outcrop; Last Interglacial, MIS 5e) of Santa Maria Island, Azores.

Tribe HYP SIGENYINI Günther, 1861

Genus *LABRODON* Gervais, 1857

† *Labrodon pavimentatum* Gervais, 1857

Figure 7

Material Examined—Two lower pharyngeal plates, both from Pedra-que-pica outcrop (Fig. 7A–E): DBUA-F 386, 678. One lower pharyngeal plate from Santa Maria Island (no locality; Fig. 7F–G): MGM 11318.

Description—The dentition of Labridae is characterized by subtriangular pharyngeal plates composed by small, contiguous, flattened teeth, which may be rounded or sub-oval in shape (Jonet, 1968). The two lower pharyngeal plates collected from the Pedra-que-pica coquina in Santa Maria Island (Azores), as well as that deposited at MGM collection (Lisbon), conform with both the figures and the description provided by Betancort et al. (2016). Cocchi (1864) was the first to describe pharyngeal plates similar to our samples, which were attributed to *Labrodon multidentis* (Sacco, 1916), and to *L. pavimentatum* (Simonelli, 1889).

Stratigraphic and geographic range—This extinct species is reported from the Miocene of France (Sauvage, 1875; Gagnaison, 2017), Portugal (Jonet et al., 1975), Poland (Bellwood et al., 2019), and the Mediterranean (Cocchi, 1864; Mas & Antunes, 2008). It is also reported from the western Atlantic, from the Miocene of Costa Rica (Laurito et al., 2014) and also from the Pliocene of Morocco (Lecointre, 1952), France (de Lumley,

1988) and the Azores (Santa Maria Island: Zbyszewski & Moitinho d'Almeida, 1950; Ferreira, 1955; Zbyszewski & Ferreira, 1962; this work).

Discussion

During the late Miocene, the once continuous and large tropical Miocene European-West African palaeobiogeographical molluscan province (Ávila et al., 2016) stretched between about 50° N to the equator and further south to Angola (Baarli et al., 2017), at about 15° S (Fig. 8A). The waters around the Azores thus exhibited tropical temperatures during the late Miocene/early Pliocene, with mean annual sea surface temperatures (SSTs) estimated to be about 3.7–6.3°C higher than the present-day 20.6°C (Ávila et al., 2016). The extinct Pliocene fish species reported from the fossiliferous sediments of Santa Maria Island in the Azores (two Sparidae indet. and *Labrodon pavimentatum*) add at least three more taxa to an increasing number of known thermophilic taxa that went extinct or locally disappeared from this volcanic oceanic island (as well as from other Atlantic islands/archipelagos) during the pronounced climatic cooling events that characterized the mid-Piacenzian stage, towards the end of the Pliocene at about 3.3–3.6 Ma (or sometime prior), consequence of a global glaciation (the MIS M2; see Lisiecki & Raymo, 2005; Fig 8B). Several warm-water species are also known to have disappeared from the Canary Islands at about 4.2–4.1Ma (Meco et al., 2007). As a direct consequence of the MIS M2 glaciation, a southward geographical range contraction of the biogeographical provinces took place (Ávila et al., 2016; Fig. 8B). Table 3 lists examples of several extinctions or extirpations of Pliocene species from the marine fauna of Santa Maria

Island across a wide spectrum of phyla: molluscs (5 bivalves, 1 benthic
macrogastropod, 1 neustonic gastropod and 4 holoplanktonic gastropods; Janssen et
al., 2008; Beu, 2017; Ávila et al., 2015, 2016, 2018a), 4 echinoderms (Madeira et al.,
2011), 1 brachiopod (Kroh et al., 2008), 1 barnacle (Winkelmann et al., 2010) and 5
sharks (Ávila et al., 2012). This pattern of extinctions/extirpations from volcanic
oceanic islands during the late Pliocene also is reported for other Atlantic archipelagos,
such as Madeira (Santos et al., 2012), Canary Islands (Meco et al., 2007, 2015, 2016;
Baarli et al., 2017) and Cabo Verde (Baarli et al., 2017).

Between 5.3 and 4.13 Ma, the protoisland of Santa Maria probably was entirely
destroyed, originating a large seamount where marine sediments rich in organic debris
were deposited. In a broader palaeoceanographical context, the pattern of sea surface
and bottom circulation in the area should have been different from the present day, as
the connection between the Atlantic and the Pacific oceans was still open at the
Central American Seaway (Garcia-Castellanos et al., 2009), and the Gibraltar Seaway
had just reopened at 5.33 Ma, in an event known as the Zanclean or post-Messinian
flood (Blanc, 2002). As attested by the fossil record of their geographical distribution,
most of the Actinopterygii (this work) and Chondrichthyes fishes (Ávila et al., 2012)
reported from the Pliocene of the Azores are also reported from Pliocene deposits in
the northeast Atlantic and the Mediterranean Sea, with some species being reported
as well from Pliocene sediments along the north-western shores of Africa, and no
Western Atlantic species reported. This predominantly eastern Atlantic pattern of
biogeographical relationships for the Azorean Pliocene marine fauna derives from the
sea surface currents set in place by the easterly trade winds that facilitated dispersal of
marine species from European shores towards the western Atlantic, a few reaching

and colonizing American Atlantic shores (Harzhauser et al., 2002; Baarli et al., 2017) and even the Pacific (Meco et al., 2016).

The final closure of the Central American Seaway, caused by re-emergence of the Panamanian land bridge (Briggs, 1995) at about 2.8 Ma (O'Dea et al., 2016), and the associated strengthening of the Gulf Stream, altered the north Atlantic system of sea surface circulation. The effect of a stronger Gulf Stream is testified by the shallow-water (<50 m depth) marine molluscs of the Azores, which show a gradual increase in the number of Western-Atlantic related species from 10.0% to 12.3%, when Last Interglacial (MIS 5e) checklists of fossil marine molluscs from the Azores (Ávila et al., 2002, 2009a, 2009b, 2015a) are compared to recent checklists available from this archipelago (Ávila, 2000; Cordeiro et al., 2015; Fig. 8C).

Intensification of the Pleistocene glacial episodes clearly impacted the Azorean fish pool in a similar manner to that described for the marine molluscs (Ávila et al., 2008a, 2008b, 2009a). We here expand on those conclusions, which were based on the MIS 5e mollusc fossil record from Santa Maria Island, and we postulate that range expansions of tropical and subtropical fishes with a Cabo Verdean origin probably occurred during 'windows of opportunity' towards the end of glacial Terminations (i.e. the short periods of time, c. 6 kyr, between the end of a glacial episode and the inception of a full interglacial episode). During those intervals, some fish species might have been able to reach, colonize and establish viable populations in the Canaries, Madeira and, in less numbers, the Azores Archipelago (Fig. 8D). We further hypothesize that, again in a similar manner to the marine molluscs (Ávila et al., 2008a, 2008b, 2009a), during the subsequent glacial episode, (1) all tropical species with a Cabo Verdean provenience that became established in the northern archipelagos

(Azores, Madeira and Canaries) during the previous interglacial were extirpated (Ávila et al., 2019; Fig. 8E), and (2) shallow-water fish species associated with fine sediments were functionally extirpated (i.e. locally disappeared from most volcanic oceanic islands) or their insular populations passed through a ‘bottleneck’ effect, resulting in a substantial net decrease in population size. The latter event occurs because the fine-sand habitat is mostly lost whenever sea level falls below the erosive edge of the insular shelf during glacial episodes (Ávila et al., 2008b, 2015a; Fig. 8G-I). The trigger for this outcome resides with the slope of the insular volcanic edifice being so steep below the erosive shelf edge, that it no longer becomes possible to physically retain fine sediments around the island, ultimately being lost to the abyssal depths (Fig. 8H). Either way, be it post-glacial colonizers or ‘bottleneck’ glacial survivors, we predict lower genetic diversity of the populations of shallow-water insular fish species associated with sandy bottoms [e.g. the following extant species in the Azores: the European finless eel *Apterichtus caecus* (Linnaeus, 1758); the wide-eyed flounder *Bothus podas* (Delaroche, 1809); the striped red mullet *Mullus surmuletus* Linnaeus, 1758; the Atlantic lizard fish *Synodus saurus* (Linnaeus, 1758); the greater weever *Trachinus draco* Linnaeus, 1758; and the pearly razorfish *Xyrichtys novacula* (Linnaeus, 1758)], when compared with continental populations. This prediction can be tested using molecular tools.

In contrast, fish species associated with algae-covered rocky shores, such as the omnivorous parrotfish *Sparisoma cretense*, were not affected by the drop in SSTs associated with the inception of glacial episodes, which was estimated at about 2–3°C (Crowley, 1981) in the region of the Azores. The lineage that led to the two eastern Atlantic species, *S. cretense* (presently reported from the Mediterranean Sea and from

the eastern Atlantic shores of mainland Portugal, the Azores, Madeira, Canaries, Cabo Verde and Senegal; cf. Table 2), and *S. strigatum* (Günther, 1862) (endemic to Saint Helena and Ascension islands; Table 2), originated at about 10 Ma from a West Atlantic ancestor (Bernardi et al., 2000). The oldest known fossil specimens of *S. cretense* were collected in two Pliocene outcrops from Santa Maria Island (Pedra-que-pica and Ponta dos Frades) and dated as 4.78 ± 0.13 to 4.13 ± 0.19 myr in age (Table 1), thus our data are in agreement with the estimates of Bernardi and colleagues (2000).

Last interglacial (MIS 5e) fossils of *S. cretense* were also collected from a Pleistocene outcrop (Vinha Velha; cf. Table 1). Today, *S. cretense* lives as far north as the eastern Atlantic shores of mainland Portugal, where it breeds during the summer, from July to September (Abecasis et al., 2008; Froese & Pauly, 2018), with mean SSTs varying from 17 to 19°C. The present average SSTs range in the Azores from 15° C in the winter, to about 23–24°C in the summer (Ávila et al., 2008c). The gonadosomatic index of *S. cretense* (which is closely associated with mean temperature) increases in the Azores from June to August and then drops to a minimum in November (Afonso et al., 2008). Hence, although it mainly breeds from mid-July to mid-September, it is possible that this species extends its reproductive activity through October (Afonso et al., 2008). Therefore, in our opinion, neither temperature nor habitat shift were an impediment for the survival of this species in the Azores throughout the last glacial episode. Our view differs from that of previous authors, who suggested that ‘the drop of sea temperatures (...) that occurred during the Pleistocene probably resulted in mass extinctions’ (of fishes) and so ‘most of the organisms now present would have reached the Azores in the last 17,000 years’ (Santos et al., 1995). Our reasoning also contradicts the conclusions of Domingues et al. (2008), who stated that ‘persistence of

(...) *S. cretense* in the Azores archipelago during the Pleistocene glaciations is difficult to admit'. Moreover, the very abundant and well-studied MIS 5e (i.e. Last Interglacial) fossil record from Santa Maria Island disproves the conclusions of these former authors.

The finding of fossil pharyngeal plates belonging to *S. cretense* in MIS 5e sediments does not necessarily exclude the hypothesis that this species locally disappeared during the Last Glaciation, only to recolonize the archipelago post-glacially. Nevertheless, we consider it more plausible to argue that such extirpation did not happen, taking in consideration the following: (1) detailed studies on the Azorean MIS 5e and recent marine molluscs showed that the drop of the SSTs during the Last Glaciation only affected two groups of species: the tropical species that reached the Azores probably during a 'window of opportunity' associated with the final phase of Termination 2, and bivalve species associated to fine-sand sediments; both groups were extirpated (Ávila et al., 2008a, 2008b, 2009b, 2015a); (2) as temperature is known to play a key role on marine fish species biology (Pepin, 1991; van der Kraak & Pankhurst, 1997; Pankhurst & Porter, 2003; Dolomatova et al., 2013; Poloczanska et al., 2016) and in light of the premises pointed above, the drop of SSTs was not sufficient enough to severely affect the survival, growth and reproduction of the temperate parrotfish *S. cretense* in the Azores during the Last Glacial episode; (3) finally, if *S. cretense* was a post-glacial colonizer of the Azores, the genetic diversity of the Azorean populations should be low, in comparison with southern Madeiran populations (postulated to have not been affected by the drop of SSTs), but that is not the case, as pointed out by Domingues et al. (2008). Thus, combining our palaeontological data with ecological and genetic data from other authors (Abecasis et

al., 2008; Afonso et al., 2008; Domingues et al., 2008), we believe the most plausible scenario is the survival of most (if not all) temperate and subtropical fish species in the Azores during the Last Glacial episode, so long as their ecological traits are not tied on sandy habitats.

This contribution expands the number of Pliocene fishes reported from the Azores to at least 11 taxa: seven taxa of Chondrichthyes (Ávila et al., 2012): *Carcharias acutissima* (Agassiz, 1833), *Carcharhinus* cf. *leucas* (Valenciennes, 1839 in Müller and Henle, 1839–1841), *Cosmopolitodus hastalis* (Agassiz, 1833), *Isurus oxyrinchus* Rafinesque, 1810, *Otodus* (*Megaselachus*) *megalodon* (Agassiz in Charlesworth, 1837), *Notorynchus primigenius* (Agassiz, 1833) and *Paratodus benedenii* (Le Hon, 1871); and four of Actinopterygii (this work): at least two Sparidae indet. taxa, *Labrodon pavimentatum* and *Sparisoma cretense*. Most of these are also reported for the Pliocene of Canaries (Betancort et al., 2016), the exceptions being *Carcharhinus* cf. *priscus* (Agassiz, 1843), *Diodon scillae* Agassiz, 1843 and *Galeocerdo* cf. *aduncus* Agassiz, 1843, only reported for Canary Islands; and Sparidae indet., *N. primigenius* and *S. cretense*, only reported for the Azores. The high number of shared species on archipelagos over 1,100 km apart is understandable if it is taken into consideration that during the Pliocene both archipelagos were part of the same Atlantic biostratigraphic molluscan province [the Pliocene Mediterranean-West African Province of Ávila et al. (2016); see Fig.8B] sharing similar tropical palaeoclimatic conditions.

Conclusions

The genus *Sparisoma* Swainson, 1839 is reported, herein, for the first time in the fossil record. The discovery of several pharyngeal plates of *S. cretense* on Pliocene and on MIS 5e sediments (Last Interglacial) of Santa Maria Island challenges former authors (Briggs, 1974; Santos et al., 1995; Domingues et al., 2006, 2008) who argued that mass extinctions affected the shallow-water marine fauna of the Azores during the course of the Last Glaciation and that, as a consequence, most of the present Azorean marine species colonized the area during the last 17–18 kyr. A multidisciplinary approach, merging our palaeontological data with ecological and genetic data, offers an alternative, more plausible explanation, suggesting that most (if not all) of the Azorean shallow-water subtropical and temperate marine species living in the archipelago during the Last Interglacial (and probably in previous interglacials as well) and associated to non-mobile sediments (e.g., rocky shores, covered or not by algae), were not affected by the drop of the SSTs during the Last Glacial episode.

Finally, we also predict that the only marine fish species that were extirpated from the Azores during the course of the Last Glaciation were those associated with fine sand (Fig. 8G-I), and the thermophilic species that colonized the archipelago during the Last Interglacial/Termination 2 ‘window of opportunity’ (Fig. 8D-F). Whether fish species ecologically associated with fine sediments and presently living in the Azores are post-glacial colonizers or ‘bottleneck’ survivors from the Last Glaciation should be assessed by genetic tools, as we anticipate low genetic diversity of the Azorean populations, when compared with continental counterparts.

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527 **Author Contributions**

528 S.P.A and R.V. designed the research plan. S.P.A., P.M., R.C. and R.V. collected the
529 fossil material. C.S.M. constructed all figures. S.P.A., M.E.J and R.V. wrote the
530 manuscript, with input from all authors, who discussed and commented on the
531 manuscript.

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534 **Additional Information**

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Figure Legends

Figure 1. A: Location of the Azores Archipelago in the NE Atlantic, and location and geotectonic setting of Santa Maria Island within the Azores Archipelago and within the Azores triple junction, respectively. The light blue area represents the Azores plateau (see text). MAR: Mid-Atlantic Ridge; EAFZ: East Azores fault zone; GF: Gloria fault; NA: North American plate; Eu: Eurasian plate; Nu: Nubian (African) plate. **B:** location of the Pliocene (red circles) and Pleistocene fossiliferous outcrops (yellow triangles) of Santa Maria Island. The Pleistocene outcrops are restricted to the warmest interval of the Last Interglacial Period (c. 116–130 kyr), which is known as MIS 5e (Marine Isotopic Substage 5e). [Intended for page width].

Figure 2. Aerial (drone) views of the fossiliferous outcrops from Santa Maria Island. **A–B:** Pliocene outcrops; **C–D:** Pleistocene (MIS 5e) outcrops. The white line delimits the outcrops. **A:** Cré. **B:** Figueiral. **C:** Prainha (the white arrows point to the areas where fossiliferous sediments are better preserved). **D:** Vinha Velha. [Intended for page width].

Figure 3. Aerial (drone) views of the Pliocene fossiliferous outcrops from Santa Maria Island. The white line delimits the outcrops. **A:** Malbusca. **B:** West fault of Malbusca. **C:** Pedra-que-pica. At West fault of Malbusca (**B**), the fossil strata are displaced by a fault slip of about 20 m. Pedra-que-pica (**C**) is probably the largest worldwide multispecific coquina ever described from a volcanic oceanic island, with a total estimated area of >23,400 m² and a total thickness for the sediments of 10–11 m (Ávila et al., 2015b). [Intended for page width].

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932 **Figure 4.** Stratigraphic columns of the studied outcrops: Malbusca (Rebelo et al.,
933 2016a; Uchman et al., 2017), Figueiral (Ávila et al., 2018), Pedra-que-pica (Ávila et al.,
934 2015b), West fault of Malbusca (Uchman et al., 2017), Cré (Janssen et al., 2008),
935 Prainha and Vinha Velha (Ávila et al., 2010, 2015a). [Intended for page width].

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937 **Figure 5.** Actinopterygian fishes from Early Pliocene deposits of Santa Maria Island
938 (Azores). **A–L:** Teeth of Sparidae indet. **A–F, L:** Upper lingual hemispheroidal
939 molariform. Note the central depression in teeth photographed in **A–C** and **F**. **A–C**
940 (DBUA-F 165-1); **D–E** (LAQ, not numbered); **F** (DBUA-F 674). **G–K:** anterior teeth. **G–H**
941 (DBUA-F 667); **I** (DBUA-F 599); **J–K** (DBUA-F 477-2); **L** (MGM 11321). **M–Q:** Teeth of
942 Sparidae indet.; these specimens were collected by Georges Zbyszewski at Santa Maria
943 Island (no information for locality) and originally referred to *Diplodus jomnitanus*.
944 Original label by the former Serviços Geológicos de Portugal (presently housed at the
945 Museu Geológico e Mineiro, Lisbon: MGM 11313). **M, P:** Apical view of the teeth; **N–O,**
946 **Q:** Lateral view. [Intended for page width].

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948 **Figure 6.** Actinopterygian fishes from Early Pliocene (DBUA-F 825, 1069, 1214) and Late
949 Pleistocene (Last Interglacial, MIS 5e; DBUA-F 1030) deposits of Santa Maria Island
950 (Azores). **A:** Lower grinding face view of Early Pliocene fragmentary tooth plate (DBUA-
951 F 1069) of *Sparisoma cretense*. **B–E:** Upper grinding face view of pharyngeal plates of
952 *Sparisoma cretense*. **B–C, E:** Lower pharyngeal plates. **B** (DBUA-F 1030), **C** (DBUA-F
953 1214-1), **E** (DBUA-F 825). **D:** Upper pharyngeal plate (DBUA-F 1214-2). **F–G:** Upper

grinding face view of recent pharyngeal plates from a dissected *S. cretense*. **F**: Upper pharyngeal plate. **G**: Lower pharyngeal plate. [Intended for page width].

Figure 7. Actinopterygian fishes from Early Pliocene deposits of Santa Maria Island (Azores). **A–G**: *Labrodon pavimentatum* Gervais, 1857. **A–E**: Fragmentary pharyngeal plates collected at Pedra-que-pica outcrop (**B, D**: DBUA-F 678; **C, E**: DBUA-F 386); **F–G**: Fragmentary pharyngeal plate collected at Santa Maria (no locality; MGM 11318). **C, D, F**: Upper grinding face view of lower pharyngeal plate. **E, G**: Lower grinding face view of lower pharyngeal plate. [Intended for page width].

Figure 8. **A–C**: NE Atlantic Biogeographic Molluscan Provinces from late Miocene to Pliocene (**A**, 6.0–5.33 Ma), early Pliocene to the end of the mid-Piacenzian Warm Period (**B**, 5.33–2.95 Ma) and the present (**C**) (Ávila et al., 2016). **D–F**: Comparison of large-scale evolutionary and biogeographical patterns, as a result of long-distance dispersal of marine species between oceanic islands located at different climate settings (subtropical versus temperate latitudes) for the last 150 kyr, encompassing the Last Interglacial (MIS 5e), the Last Glacial episode, and the present interglacial⁶. **D**: During the final phase of Termination 2 and/or the beginning of the Last Interglacial, a subset of marine subtropical species expands their geographical ranges towards higher latitudes (red arrow), reaching temperate archipelagos and establishing viable populations in those islands (e.g. small red star in **D**). In a similar way, a subset of marine temperate species expands their geographical ranges towards higher latitudes, reaching boreal/arctic archipelagos and establishing viable populations as well (blue arrow). **E**: During the course of the Last Glacial episode, the thermophilic species that

established on temperate islands are extirpated and it is expected that a subset of marine temperate species (small blue circles), adapted to cool temperatures may expand their geographical ranges towards lower latitudes, reaching subtropical archipelagos and establishing viable populations. **F**: the subsequent episode of global warming that led to the present interglacial, extirpates the cold-adapted species that reached subtropical islands during the previous glacial episode, and range expansion of species towards higher latitudes is documented. **G–I**: Mobile sediment response to glacio-eustatic sea level fluctuations (Ávila et al., 2008b, 2010). **G**: During the Last interglacial (MIS 5e), as a result of marine and fluvial erosion, sediments are transported to the island shelf where they accumulate between the shoreline angle (i.e., the angle between the cliff of the island and the shore platform) and the erosive shelf edge. **H**: The inception of sea-level lowstands during the Last Glacial episode, promotes the remobilization and transport downslope of sediments from the island shelf, when relative mean sea level falls below the erosive shelf edge. Whenever this happens, sediments are lost to the abyssal depths that surround the insular edifice. **I**: Sediments accumulate again on the island shelf, as a result of marine and fluvial erosion, during the present sea-level highstand. [Intended for page width].

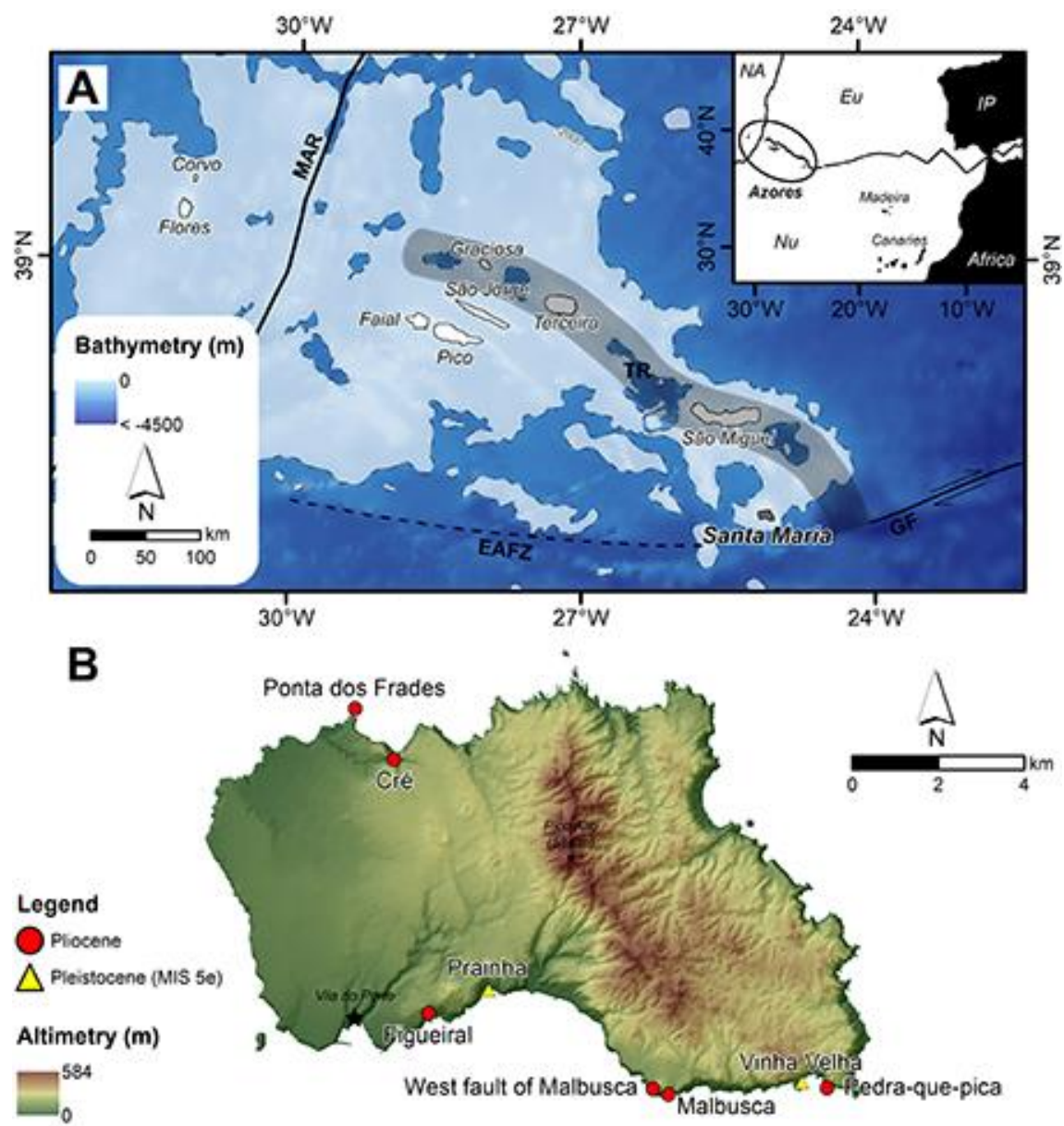


Figure 1. A: Location of the Azores Archipelago in the NE Atlantic, and location and geotectonic setting of Santa Maria Island within the Azores Archipelago and within the Azores triple junction, respectively. The light blue area represents the Azores plateau (see text). MAR: Mid-Atlantic Ridge; EAFZ: East Azores fault zone; GF: Gloria fault; NA: North American plate; Eu: Eurasian plate; Nu: Nubian (African) plate. **B:** location of the Pliocene (red circles) and Pleistocene fossiliferous outcrops (yellow triangles) of Santa Maria Island. The Pleistocene outcrops are restricted to the warmest interval of the

1007 Last Interglacial Period (c. 116–130 kyr), which is known as MIS 5e (Marine Isotopic
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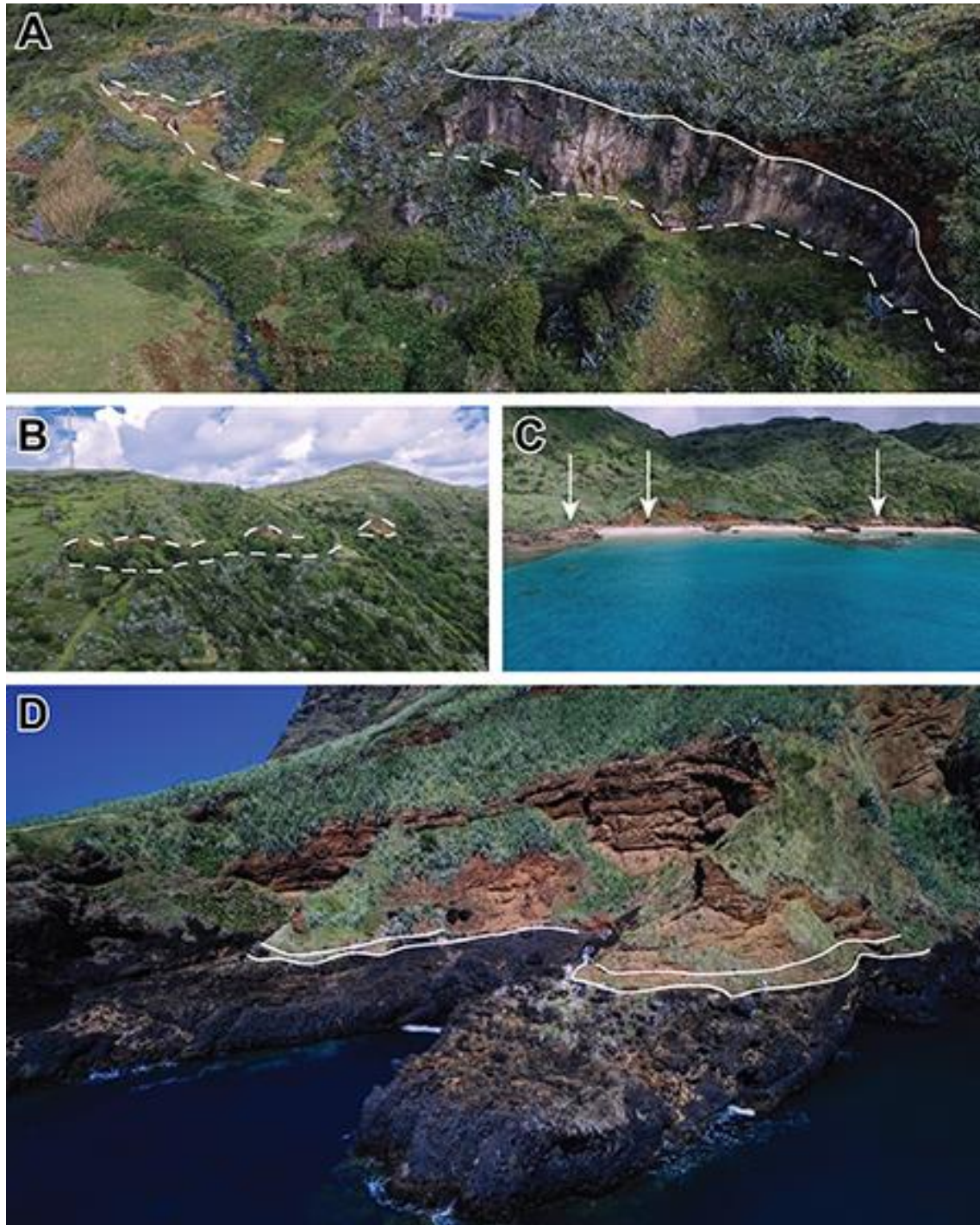


Figure 2. Aerial (drone) views of the fossiliferous outcrops from Santa Maria Island. **A–B:** Pliocene outcrops; **C–D:** Pleistocene (MIS 5e) outcrops. The white line delimits the outcrops. **A:** Cré. **B:** Figueiral. **C:** Prainha (the white arrows point to the areas where fossiliferous sediments are better preserved). **D:** Vinha Velha. [Intended for page width].

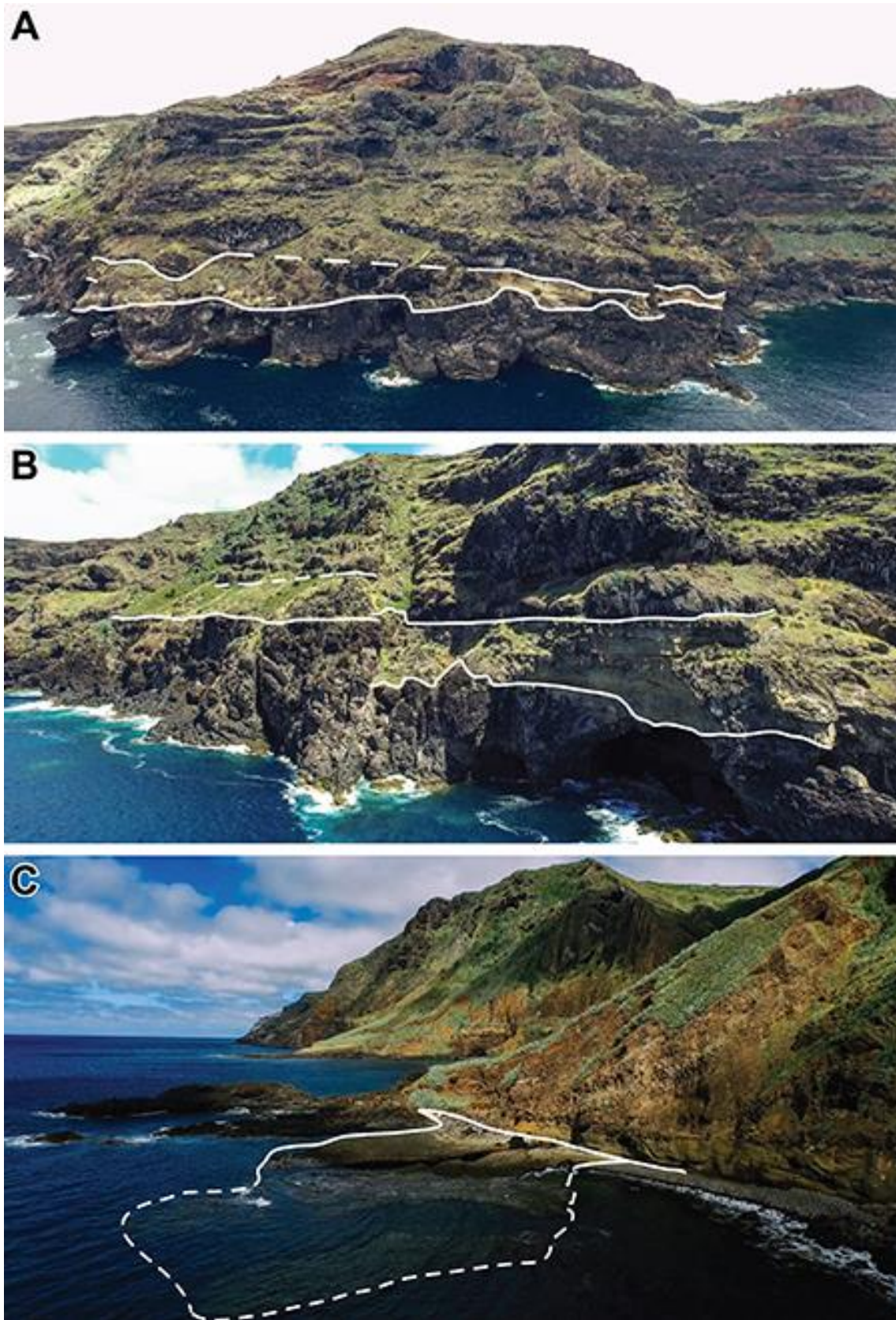


Figure 3. Aerial (drone) views of the Pliocene fossiliferous outcrops from Santa Maria Island. The white line delimits the outcrops. **A:** Malbusca. **B:** West fault of Malbusca. **C:**

1020 Pedra-que-pica. At West fault of Malbusca (**B**), the fossil strata are displaced by a fault
1021 slip of about 20 m. Pedra-que-pica (**C**) is probably the largest worldwide multispecific
1022 coquina ever described from a volcanic oceanic island, with a total estimated area of
1023 >23,400 m² and a total thickness for the sediments of 10–11 m (Ávila et al., 2015b).
1024 [Intended for page width].

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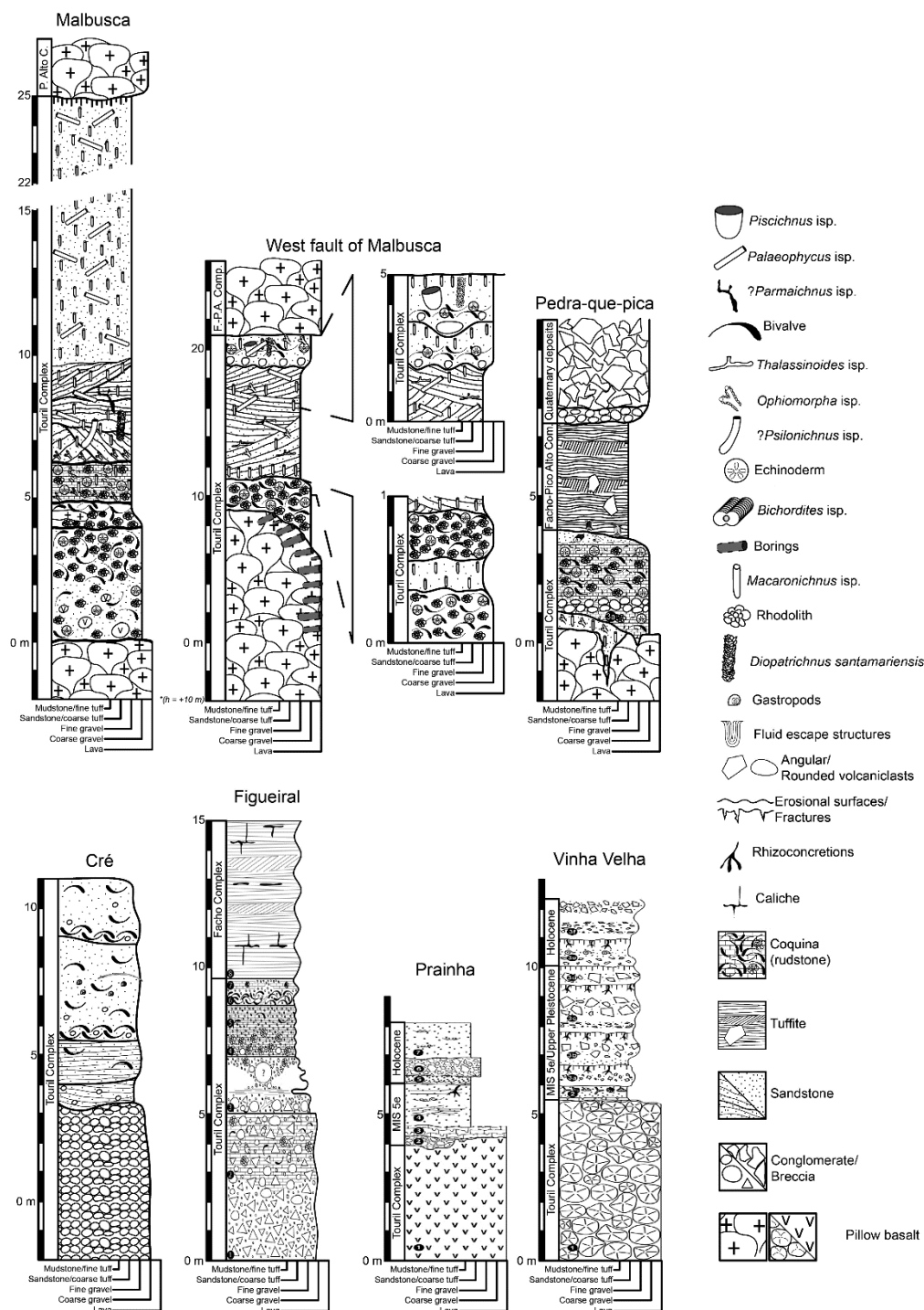


Figure 4. Stratigraphic columns of the studied outcrops: Malbusca (Rebello et al., 2016a; Uchman et al., 2017), Figueiral (Ávila et al., 2018), Pedra-que-pica (Ávila et al., 2015b), West fault of Malbusca (Uchman et al., 2017), Cré (Janssen et al., 2008), Prainha and Vinha Velha (Ávila et al., 2010, 2015a). [Intended for page width].

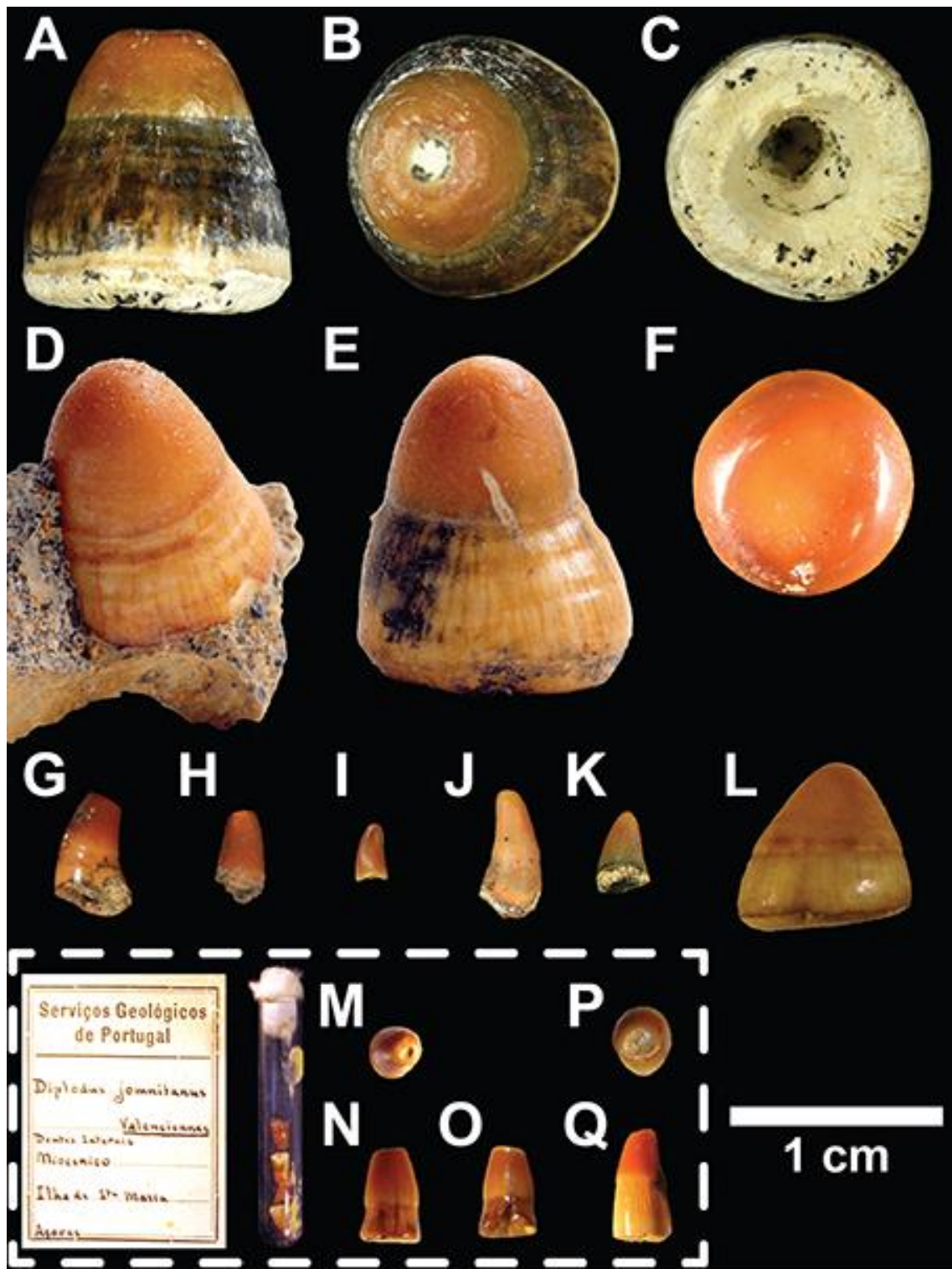


Figure 5. Actinopterygian fishes from Early Pliocene deposits of Santa Maria Island (Azores). **A–L:** Teeth of Sparidae indet. **A–F, L:** Upper lingual hemispheroidal molariform. Note the central depression in teeth photographed in **A–C** and **F**. **A–C** (DBUA-F 165-1); **D–E** (LAQ, not numbered); **F** (DBUA-F 674). **G–K:** anterior teeth. **G–H**

1038 (DBUA-F 667); **I** (DBUA-F 599); **J–K** (DBUA-F 477-2); **L** (MGM 11321). **M–Q**: Teeth of
1039 Sparidae indet.; these specimens were collected by Georges Zbyszewski at Santa Maria
1040 Island (no information for locality) and originally referred to *Diplodus jomnitanus*.
1041 Original label by the former Serviços Geológicos de Portugal (presently housed at the
1042 Museu Geológico e Mineiro, Lisbon: MGM 11313). **M, P**: Apical view of the teeth; **N–O**,
1043 **Q**: Lateral view. [Intended for page width].

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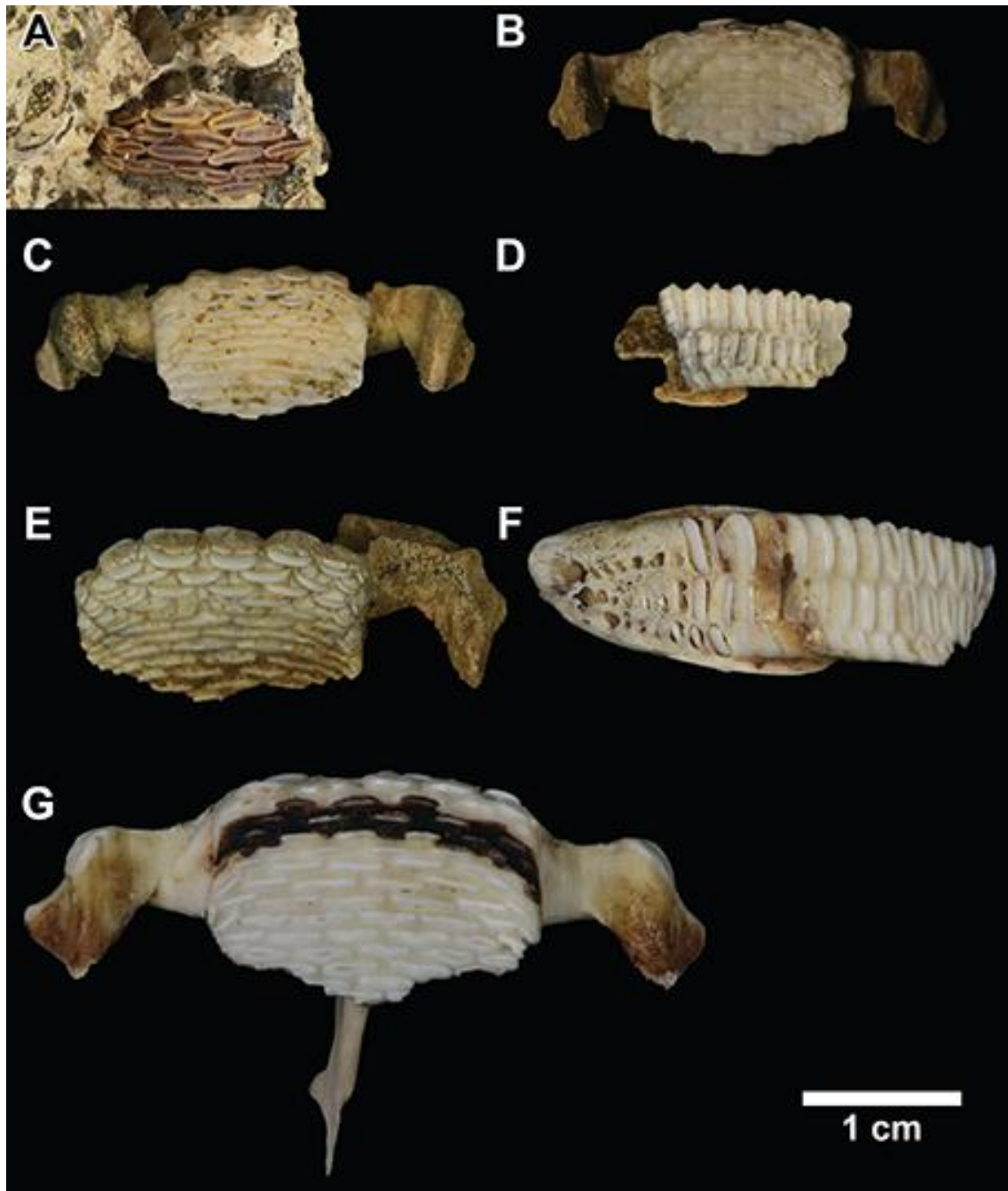


Figure 6. Actinopterygian fishes from Early Pliocene (DBUA-F 825, 1069, 1214) and Late Pleistocene (Last Interglacial, MIS 5e; DBUA-F 1030) deposits of Santa Maria Island (Azores). **A:** Lower grinding face view of Early Pliocene fragmentary tooth plate (DBUA-F 1069) of *Sparisoma cretense*. **B–E:** Upper grinding face view of pharyngeal plates of *Sparisoma cretense*. **B–C, E:** Lower pharyngeal plates. **B** (DBUA-F 1030), **C** (DBUA-F 1214-1), **E** (DBUA-F 825). **D:** Upper pharyngeal plate (DBUA-F 1214-2). **F–G:** Upper

1053 grinding face view of recent pharyngeal plates from a dissected *S. cretense*. **F**: Upper
1054 pharyngeal plate. **G**: Lower pharyngeal plate. [Intended for page width].
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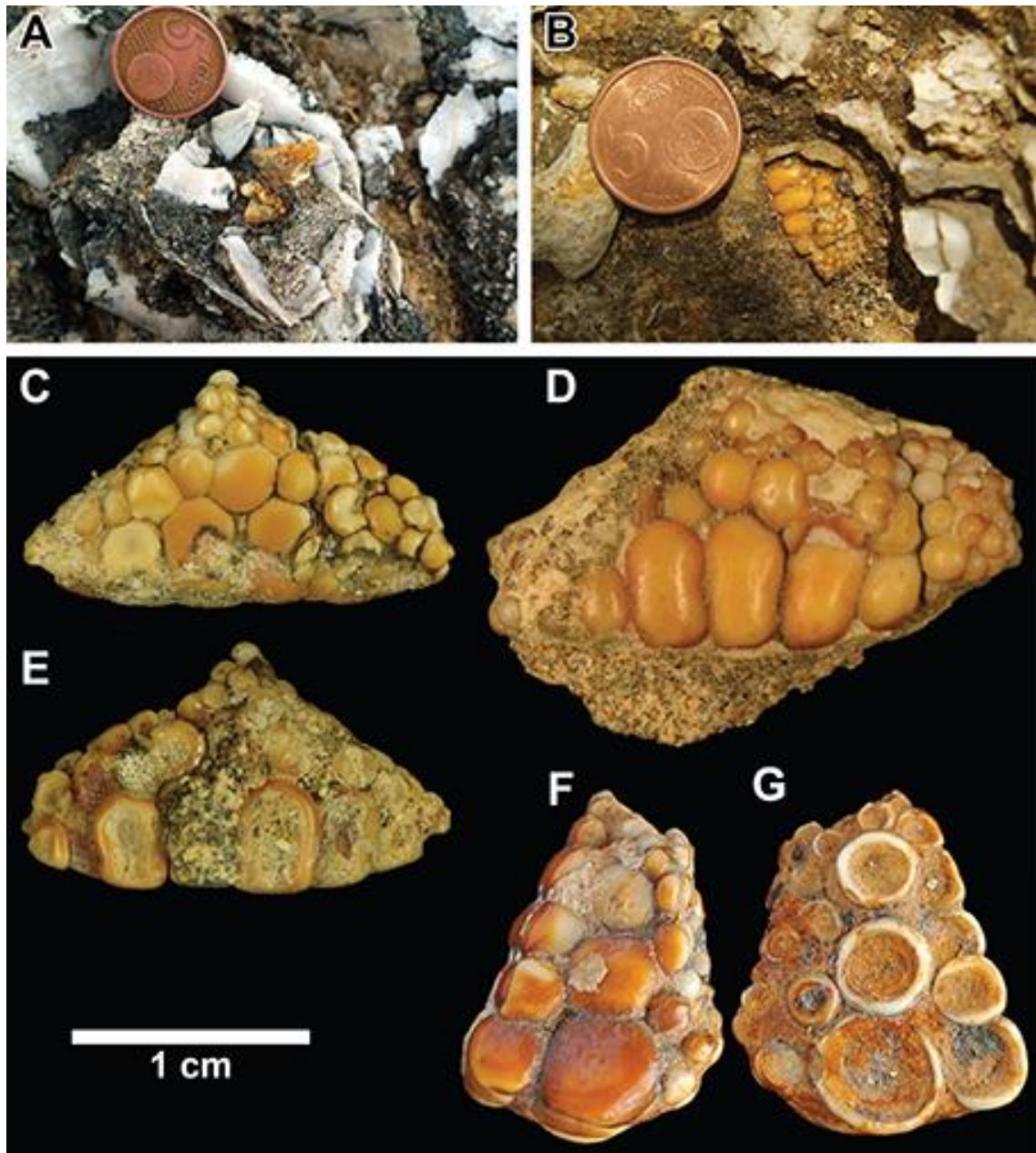


Figure 7. Actinopterygian fishes from Early Pliocene deposits of Santa Maria Island (Azores). **A–G:** *Labrodon pavimentatum* Gervais, 1857. **A–E:** Fragmentary pharyngeal plates collected at Pedra-que-pica outcrop (**B, D:** DBUA-F 678; **C, E:** DBUA-F 386); **F–G:** Fragmentary pharyngeal plate collected at Santa Maria (no locality; MGM 11318). **C, D, F:** Upper grinding face view of lower pharyngeal plate. **E, G:** Lower grinding face view of lower pharyngeal plate. [Intended for page width].

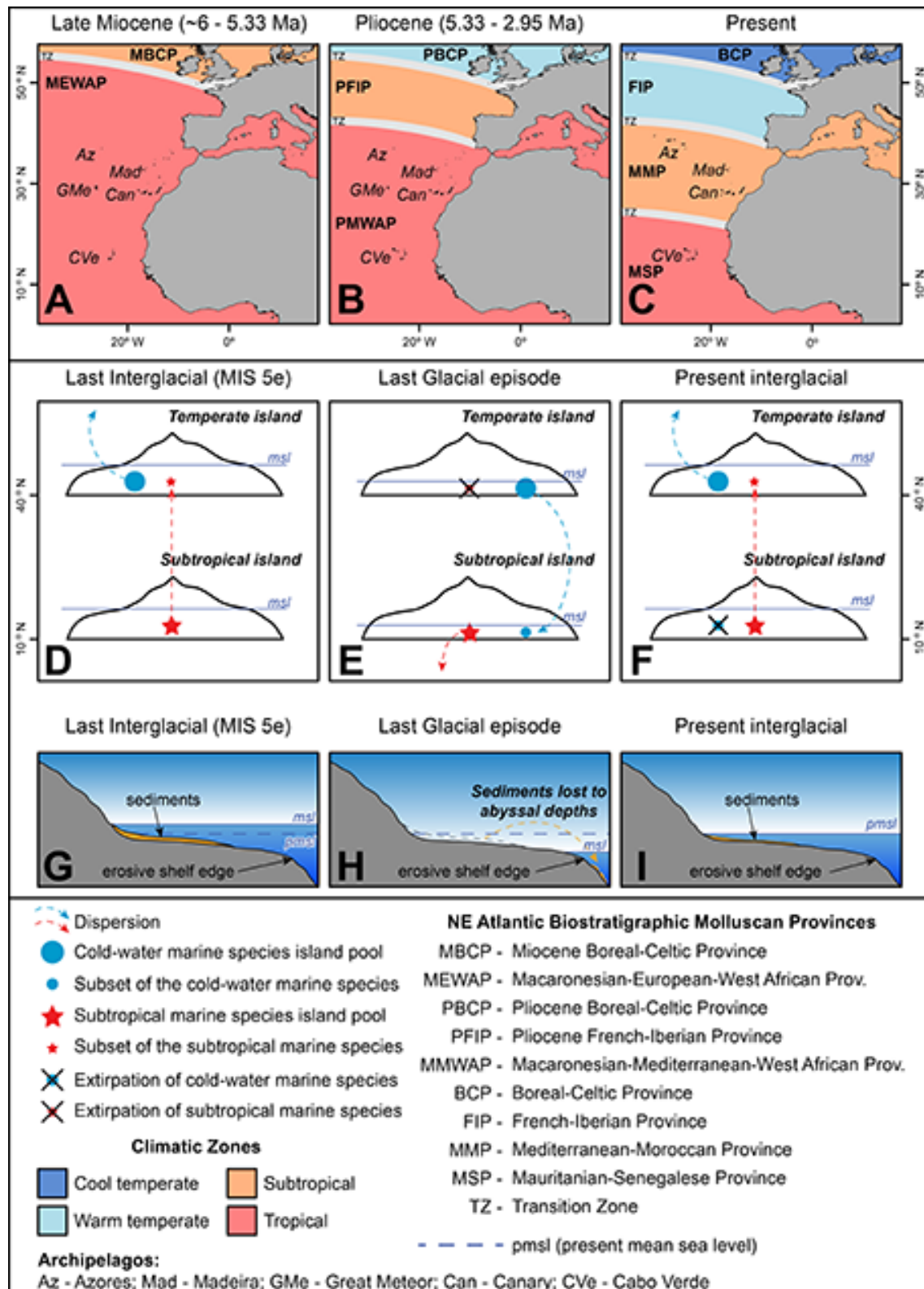


Figure 8. A–C: NE Atlantic Biogeographic Molluscan Provinces from late Miocene to Pliocene (A, 6.0–5.33 Ma), early Pliocene to the end of the mid-Piacenzian Warm Period (B, 5.33–2.95 Ma) and the present (C) (Ávila et al., 2016). **D–F:** Comparison of

1069 large-scale evolutionary and biogeographical patterns, as a result of long-distance
1070 dispersal of marine species between oceanic islands located at different climate
1071 settings (subtropical versus temperate latitudes) for the last 150 kyr, encompassing the
1072 Last Interglacial (MIS 5e), the Last Glacial episode, and the present interglacial (Ávila et
1073 al., 2019). **D**: During the final phase of Termination 2 and/or the beginning of the Last
1074 Interglacial, a subset of marine subtropical species expands their geographical ranges
1075 towards higher latitudes (red arrow), reaching temperate archipelagos and
1076 establishing viable populations in those islands (e.g. small red star in **D**). In a similar
1077 way, a subset of marine temperate species expands their geographical ranges towards
1078 higher latitudes, reaching boreal/arctic archipelagos and establishing viable
1079 populations as well (blue arrow). **E**: During the course of the Last Glacial episode, the
1080 thermophilic species that established on temperate islands are extirpated and it is
1081 expected that a subset of marine temperate species (small blue circles), adapted to
1082 cool temperatures may expand their geographical ranges towards lower latitudes,
1083 reaching subtropical archipelagos and establishing viable populations. **F**: the
1084 subsequent episode of global warming that led to the present interglacial, extirpates
1085 the cold-adapted species that reached subtropical islands during the previous glacial
1086 episode, and range expansion of species towards higher latitudes is documented. **G–I**:
1087 Mobile sediment response to glacio-eustatic sea level fluctuations (Ávila et al., 2008b,
1088 2010). **G**: During the Last interglacial (MIS 5e), as a result of marine and fluvial erosion,
1089 sediments are transported to the island shelf where they accumulate between the
1090 shoreline angle (i.e., the angle between the cliff of the island and the shore platform)
1091 and the erosive shelf edge. **H**: The inception of sea-level lowstands during the Last
1092 Glacial episode, promotes the remobilization and transport downslope of sediments

1093 from the island shelf, when relative mean sea level falls below the erosive shelf edge.
1094 Whenever this happens, sediments are lost to the abyssal depths that surround the
1095 insular edifice. I: Sediments accumulate again on the island shelf, as a result of marine
1096 and fluvial erosion, during the present sea-level highstand. [Intended for page width].
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1098 **Table 1.** Relative age of the outcrops (in millions of years) from Santa Maria Island (Azores) containing fossil remains of Actinopterygii fishes.

Outcrop	Age (Ma)	Bed	Inferred depth of deposition (m)	Number of teeth/pharyngeal plates per species		
				Sparidae	<i>Sparisoma cretense</i>	<i>Labrodon pavimentatum</i>
Cré	4.78 ± 0.13 to 4.13 ± 0.19	Sandstone	?	4		
Figueiral	4.78 ± 0.13 to 4.13 ± 0.19	Sandstone	?	1		
Ponta dos Frades	4.78 ± 0.13 to 4.13 ± 0.19	Sandstone	?		6	
Pedra-que-pica	4.78 ± 0.13 to 4.13 ± 0.19	Coquina	-40	26	1	2
Malbusca	4.32 ± 0.06 to 4.02 ± 0.06	Sandstone	-60	17		
West fault of Malbusca	4.32 ± 0.06 to 4.02 ± 0.06	Sandstone	-60	2		
Vinha Velha	Last Interglacial (MIS 5e)	Sand	+1 to +2		1	

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1101 **Table 2.** Present geographical distribution of the genus *Sparisoma* in the Atlantic Ocean and the Mediterranean Sea. AZO – Azores Archipelago;
 1102 MAD – Madeira Archipelago; SEL – Selvagens Archipelago; CAN – Canaries Archipelago; CAB – Cabo Verde Archipelago; STP – São Tomé and
 1103 Príncipe Archipelago; IBE – Atlantic Iberian shores from Finisterra south to the Straits of Gibraltar; MED – Mediterranean Sea; TWAF – Tropical
 1104 west African shores, from Cape Blanc (Senegal) south to Angola; NWA – Atlantic coast of Northwest Africa, from the Straits of Gibraltar south
 1105 to Senegal; ASC – Ascension Island; STH – Saint Helena Island; SWA – Southwest Atlantic, including Brazil and its oceanic islands; WAT –
 1106 Western Atlantic, including the Caribbean Sea.

Species	AZO	MAD	SEL	CAN	CAB	STP	IBE	MED	TWAF	NWA	ASC	STH	SWA	WAT
<i>Sparisoma choati</i>	0	0	0	0	1	1	0	0	1	0	0	0	0	0
<i>Sparisoma cretense</i>	1	1	1	1	1	0	1	1	1	1	0	0	0	0
<i>Sparisoma frondosum</i>	0	0	0	0	1	0	0	0	0	0	0	0	0	0
<i>Sparisoma amplum</i>	0	0	0	0	0	0	0	0	0	0	0	0	1	0
<i>Sparisoma atomarium</i>	0	0	0	0	0	0	0	0	0	0	0	0	0	1
<i>Sparisoma aurofrenatum</i>	0	0	0	0	0	0	0	0	0	0	0	0	0	1
<i>Sparisoma axillare</i>	0	0	0	0	0	0	0	0	0	0	0	0	1	0
<i>Sparisoma chrysopterum</i>	0	0	0	0	0	0	0	0	0	0	0	0	0	1
<i>Sparisoma griseorubrum</i>	0	0	0	0	0	0	0	0	0	0	0	0	0	1
<i>Sparisoma radians</i>	0	0	0	0	0	0	0	0	0	0	0	0	1	1
<i>Sparisoma rocha</i>	0	0	0	0	0	0	0	0	0	0	0	0	1	0
<i>Sparisoma rubripinne</i>	0	0	0	0	0	0	0	0	0	0	0	0	0	1
<i>Sparisoma strigatum</i>	0	0	0	0	0	0	0	0	0	0	1	1	0	0
<i>Sparisoma tuiupiranga</i>	0	0	0	0	0	0	0	0	0	0	0	0	1	0
<i>Sparisoma viride</i>	0	0	0	0	0	0	0	0	0	0	0	0	1	1

1108 **Table 3.** Taxa/species reported from the Pliocene sediments of Santa Maria Island and
 1109 that are presently extinct or were extirpated (local disappearances) from the marine
 1110 fauna of the island during the global climate deterioration that characterizes the late
 1111 Pliocene.

Phylum	Taxa	Status
Mollusca	<i>Aequipecten macrotis</i> (Sowerby, 1847)	extinct
	<i>Gigantopecten latissimus</i> (Brocchi, 1814)	extinct
	<i>Chlamys hartungi</i> (Mayer, 1864)	extinct
	<i>Lopha plicatuloides</i> (Mayer, 1864)	extinct
	<i>Pecten dunkeri</i> Mayer, 1864	extinct
	<i>Persististrombus coronatus</i> (Defrance, 1827)	extinct
	<i>Janthina typica</i> (Brönn, 1861)	extinct
	<i>Cavolinia grandis</i> (Bellardi, 1873)	extinct
	<i>Cavolinia marginata</i> (Brönn, 1862)	extinct
	<i>Cuvierina intermedia</i> (Bellardi, 1873)	extinct
	<i>Bowdenathea jamaicensis</i> Collins, 1934	extinct
Echinodermata	<i>Clypeaster altus</i> (Leske, 1778)	extinct
		local
	<i>Eucidaris tribuloides</i> (Lamarck, 1816)	disappearance
		local
	<i>Echinoneus</i> cf. <i>cyclostomus</i> Leske, 1778	disappearance
	<i>Schizobrissus</i> sp.	extinct
Brachiopoda	<i>Novocrania turbinata</i> (Poli, 1795)	extinct

Table 3. (Continued)

Arthropoda	<i>Zullobalanus santamariaensis</i> Buckeridge &	
(Crustacea)	Winkelmann, 2010	extinct
Chordata		
(Elasmobranchii)	<i>Notorynchus primigenius</i> (Agassiz, 1833)	extinct
	<i>Carcharias acutissima</i> (Agassiz, 1833),	extinct
	<i>Cosmopolitodus hastalis</i> (Agassiz, 1833),	extinct
	<i>Paratodus benedenii</i> (Le Hon, 1871),	extinct
	<i>Megaselachus megalodon</i> (Agassiz in Charlesworth,	
	1837)	extinct
Chordata		
(Actinopterygii)	Sparidae indet.	extinct
	<i>Labrodon pavimentatum</i> Gervais, 1857	extinct

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