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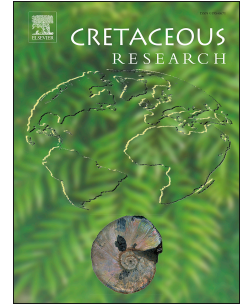
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- D. BERT: let the work, wrote the article, made field-trip, made analyses;
- S. Reboulet: wrote the article, made field-trip;
- B. Vernet: wrote the article, made field-trip;
- S. Bersac: wrote the article, made analyses
- L. Canut: made field-trip

All the co-authors participated to the study, and have approved the final version of the manuscript.

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1 **Early *Crioceratites* (heteromorphic ammonites) from the lower Hauterivian**  
2 **of south-eastern France: systematics, intraspecific variation and**  
3 **biostratigraphic implications**

4

5 Didier Bert<sup>a,b,c\*</sup>, Stéphane Reboulet<sup>d,c</sup>, Benjamin Vernet<sup>b</sup>, Stéphane Bersac<sup>c</sup>, Léon Canut<sup>c</sup>

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7 \* Corresponding author: [didier.paleo@gmail.com](mailto:didier.paleo@gmail.com)8 <sup>a</sup>: Réserve naturelle nationale géologique de Haute-Provence, Conseil départemental des Alpes de  
9 Haute-Provence, 13 rue du Docteur Romieu, CS 70216, F-04995 Digne-Les-Bains Cedex 9, France.10 <sup>b</sup>: Laboratoire Géosciences, UMR-CNRS 6118, Université de Rennes-1, campus Beaulieu, bâtiment  
11 15, F-35042 Rennes cedex, France.12 <sup>c</sup>: Laboratoire du Groupe de recherche en Paléobiologie et biostratigraphie des Ammonites (GPA), 65  
13 Grand rue, F-04170 La Mure-Argens, France.14 <sup>d</sup>: Université de Lyon, UCBL, ENSL, CNRS, LGL TPE, Bâtiment Géode, 2 rue Dubois, 69622  
15 Villeurbanne, France.

16

17 **Highlights**

- 18 • First Hauterivian acme of the
- Crioceratites*
- belongs to
- C. claveli*
- sp. nov. only
- 
- 19 •
- C. claveli*
- variation is under dimorphism, multipolar variation of order two and
- 
- 20 heterochronies
- 
- 21 • Multipolar variation is not under the strict Buckman's 1st Law in
- C. claveli*
- 
- 22 • New
- C. claveli*
- (Sub-)Zones are introduced in replacement of
- C. loryi*
- (Sub-)Zones
- 
- 23 •
- Crioceratites jurensis*
- is revised; its synonymy with
- Davouxiceras nolani*
- is rejected

24

25 **Abstract**

26 Two acmes of the genus *Crioceratites* were recorded in the current *Crioceratites loryi* Zone  
27 (lower Hauterivian) in the Lower Cretaceous series of south-eastern France (Vocontian  
28 Basin), both in the classic section of the Cheiron ravine (Castellane area) and in the  
29 Hauterivian GSSP area at La Charce. The revision of an abundant material allows us to define  
30 *Crioceratites claveli* sp. nov. The first acme and the first part of the second acme correspond  
31 to the range of *C. claveli* sp. nov. and not to *C. loryi* as previously admitted. The remaining of  
32 the second acme is related to this latter species solely. The great abundance of *C. claveli* sp.  
33 nov. in the lower interval Horizon allowed a detailed analysis of its variation around 3 factors:  
34 (1) dimorphism; (2) multipolar variation of order two here characterized by an absence of  
35 correlation between the ornamental and morphological shell parameters (non-covariant  
36 factors); and (3) ontogenic heterochronies, which depend on the robustness of the  
37 ornamentation only. On the other hand, the northern edge of the Provençal platform delivered  
38 *C. jurensis* (revised here) from the top part of the *Acanthodiscus radiatus* Zone, which  
39 reduces the gap with the older *Crioceratites* known in the uppermost Valanginian. Finally, the  
40 study of the stratigraphic distribution of *C. claveli* sp. nov. and *C. loryi* leads to propose the  
41 new *C. claveli* (Sub)Zone instead of the *C. loryi* (Sub)Zone. This proposal for the south-  
42 eastern France may have extension potential for the Mediterranean Province of the  
43 Mediterranean Caucasian Subrealm (Tethyan Realm).

44

45 **Keywords:** Heteromorphic ammonites; *Crioceratites*; Hauterivian; Intraspecific variability;  
46 Biozonation; France.

47

## 48 1. Introduction

49

50 The heteromorphic ammonite group related to *Davouxiceras nolani* (Kilian, 1910) s.l. was  
51 excluded from the *Crioceratites* Léveillé, 1837 s. str. and is now considered an  
52 Emericiceratidae Vermeulen, 2004 (see Klein et al., 2007, p. 67). Consequently, the  
53 appearance of the genus *Crioceratites* was recognized to coincide with that of *Crioceratites*  
54 *loryi* Sarkar, 1955 (index species of the current *Crioceratites loryi* Zone, Reboulet et al.,  
55 2018), that is to say widely beyond the base of the Hauterivian stage. This situation changed  
56 with the discovery of ‘primitive *Crioceratites*’ at the top part of the Valanginian by Reboulet  
57 (1996). However, the link between the Hauterivian and Valanginian *Crioceratites* was not  
58 then evident due to the hiatus that remained throughout the thick *Acanthodiscus radiatus* Zone  
59 (Reboulet, 1996, p. 277).

60 Research carried out in the south-east of France has shown that the stratigraphic distribution  
61 of the *Crioceratites* in the current *C. loryi* Zone is organized around two acmes (Bulot, 1995,  
62 tab. 16; Reboulet, 1996, fig. 23; and this work). However, when the stratigraphy is known  
63 with sufficient precision, only the elements of the second occurrence have been cited and  
64 figured in the literature (i.e. Thieuloy, 1972; Ropolo, 1995; Bulot, 1995, p. 93-94; Reboulet,  
65 1996, p. 178, pl. 24, fig. 6): all these specimens correspond to the original representation of *C.*  
66 *loryi* from Sarkar (1955) and to the very accurate description of this species by Thieuloy  
67 (1972). It was therefore important to verify if the *Crioceratites* of the first acme indeed  
68 correspond to the same species, and if it does not, to verify their relations with *C. loryi* and to  
69 draw all the possible important biostratigraphic implications, *C. loryi* being a zonal index.

70 This research was carried out in the classic section of the Cheiron ravine (south-eastern  
71 France, Alpes de Haute-Provence department, south of the Vocontian Basin), which, although  
72 cited for a long time in the literature, has been the subject of only few studies and no precise  
73 survey (Cotillon, 1971, fig. 3). The detailed bed-by-bed survey of the Hauterivian in the  
74 Cheiron ravine was carried out for the first time by Vernet and Bert (Vernet, 2017,

75 unpublished Master thesis) as part of an internship at the Réserve naturelle nationale  
76 géologique de Haute-Provence (RNNGHP). For the present work, this section was resampled  
77 and yielded numerous ammonites including many *Crioceratites*. In the lower part of the  
78 current *C. loryi* Zone (the first acme evoked previously; = new *Crioceratites claveli* Zone, this  
79 work) all the *Crioceratites* belong to the same species (here *C. claveli* sp. nov.). The study  
80 material of Reboulet (1996) from La Charce (the Hauterivian GSSP) was also reviewed, as  
81 well as new elements from the complementary Pommerol section (Reboulet, 2015; Aguirre-  
82 Urreta et al., 2019 – Drôme department, central part of the Vocontian Basin). The abundant  
83 material from the first acme made it possible to perform a statistical analysis and identify the  
84 variation of the sample studied on the basis of the intraspecific laws of variability recently  
85 explored by Bert (2014a,b; 2019).

86 Additional research in the neritic domain of the Arc de Castellane (Provençal platform  
87 boarder, southeastern France) did not allow finding the two occurrences of *Crioceratites* of  
88 the current *C. loryi* Zone. However, several specimens of a related species [*Crioceratites*  
89 *jurensis* (Nolan, 1894) – rehabilitated in this work] have been collected in the upper part of  
90 the *A. radiatus* Zone, making it currently the oldest recognized *Crioceratites* s. str. species in  
91 the Hauterivian.

92

## 93 **2. Geological setting**

94

95 The three main sections studied belong to the Vocontian Basin series (Paquier, 1900 – Fig. 1).  
96 During the Early Cretaceous, the Vocontian Basin was located at a palaeolatitude of ~30°N in  
97 a marginal marine position of the western Ligurian Tethys Ocean. This west-east oriented  
98 basin corresponded to a deep (a few hundred meters) depositional environment (pelagic  
99 series), and included the Diois-Baronnies region and the tectonic arc of Digne-Castellane in

100 the west and east parts, respectively (Ferry, 2017). It was surrounded by slopes (transitional  
101 domain, hemiplegic series) and platforms (neritic series; Vercors, Vivarais and Provence  
102 regions). The dilated/expanded Hauterivian sections of the Vocontian Basin s.l. provide good  
103 conditions to follow the succession and evolution of ammonite faunas with great precision.

104  
105 The Cheiron ravine (CHE – Castellane; Cotillon, 1971) provided the main part of the fossils  
106 studied. Although located in the pelagic domain of the southern part of the Vocontian Basin  
107 (the Eastern Subalpine Basin in the sense of Ferry, 2017), this section is very close to the  
108 hemipelagic domain from which it receives some influences (Fig. 1). From a structural point  
109 of view, this section belongs to the southern subalpine chains of the Arc de Castellane, which  
110 are the outermost units of the Alpine Arc. The La Charce (LCH – Hauterivian GSSP;  
111 Reboulet et al., 1992; Reboulet, 1996; Reboulet and Atrops, 1999; Mutterlose et al., 2020)  
112 and Pommerol sections (POM – Drôme department; Reboulet, 2015; Aguirre-Urreta et al.,  
113 2019) are located in the western part of the Vocontian Basin (the Vocontian through in the  
114 sense of Ferry, 2017), in the pelagic domain (Fig. 1).

115 The successions of the three sections studied (CHE, LCH, POM) are characterized by  
116 marlstone-limestone alternations, which have been interpreted as the result of cycles in the  
117 production of calcareous nannoplankton caused by climatic fluctuations in the Milankovitch  
118 frequency band (Cotillon et al., 1980; Giraud et al., 1995). Alternatively, Reboulet et al.  
119 (2003) have proposed a model of cyclic export of carbonate mud from shallow platform  
120 environments towards the basin for the Valanginian successions of the Vocontian Basin.  
121 Occasionally, synsedimentary slumps occurred in the basin.

122 Some sections of the Provençal platform boarder (Arc de Castellane) were sampled for the  
123 present work (Rougou and La Palud-sur-Verdon area in the Alpes de haute-Provence  
124 department: CAR, PV2 – Cotillon, 1971; Bulot et al., 1995; Reboulet, 1996 – and Bargème



125 area in the Var department: COU1; Fig. 1). These sections are characterized by neritic  
126 deposits with often reduced and sometimes discontinuous sedimentation, in particular in the  
127 lower Hauterivian.

128

### 129 **3. Material and methods**

130

131 **Stratigraphy.** The Crioceratitidae are relatively rare in the upper Valanginian and more  
132 abundant in the lower Hauterivian substage (Reboulet, 1996, figs. 22-23). For this  
133 stratigraphic interval, numerous works dealing with paleontology and biostratigraphy were  
134 made during the last four decades; they allowed to improve the zonal scheme of the upper  
135 Valanginian–lower Hauterivian of the south-east France basin (Thieuloy, 1977; Reboulet et  
136 al., 1992; Bulot et al., 1993, 1995; Thieuloy and Bulot, 1993; Bulot, 1995; Bulot and  
137 Thieuloy, 1995; Reboulet, 1996, 2015; Reboulet and Atrops, 1999). These studies provide a  
138 sound base for the standard Mediterranean ammonite scale accepted by the Kilian Group  
139 (Hoedemaeker et al., 2003; Reboulet et al., 2006; 2009; 2011; 2014; 2018); this Lower  
140 Cretaceous standard zonation is applied here.

141 The terminologies used in this paper (biostratigraphy and chronostratigraphy versus  
142 geochronology) respect the standards of the International Commission on Stratigraphy  
143 (Salvador, 1994).

144 **Material.** This study focuses on the discovery of new bed-by-bed sampled material of the  
145 genus *Crioceratites* in the lower Hauterivian. The main part of this material (122 specimens)  
146 comes from the sections CHE (88 specimens), LCH (26 specimens) and POM (8 specimens)  
147 and belongs to the species *C. claveli* sp. nov. Five other specimens of the little-known species  
148 *C. jurensis* were collected in the sections CAR, PV2 and COU1.

149 All the sections studied in the Alpes de Haute-Provence and Var departments belong to the  
150 protected perimeter of the ‘Réserve naturelle nationale géologique de Haute-Provence’  
151 (RNNGHP), managed by the Departmental Council of the Alpes de Haute-Provence on behalf  
152 of the French State. The section LCH is the GSSP of the Hauterivian stage; this site is  
153 protected as an ‘Espace naturel sensible’ (Mutterlose et al., 2020).

154 **Repositories.** The material from LCH and POM sections is stored in the palaeontological  
155 collections of the Claude Bernard University of Lyon (France, collections of Reboulet, and  
156 Reboulet and Noclin). The material collected in the CHE section (Bert’s collection) is curated  
157 by the RNNGHP. Some interesting specimens were added to the study albeit they are in a  
158 private collection (L. Canut): they are a minority, and of course they are never used as type  
159 specimens; the curating of these specimens is monitored by the RNNGHP.

160 Collections acronyms: MHNG=Museum of Natural history of Geneva (Switzerland);  
161 FSL=Faculté des Sciences de Lyon; RNNGHP=Réserve naturelle nationale géologique de  
162 Haute-Provence; DBT=D. Bert; LCT=L. Canut.

163 **Measurements.** The ammonoid fauna is relatively well preserved, but the fragmentation of  
164 specimens can be frequent in some beds. Dissolution of the ammonoid shells is the norm and  
165 specimens are preserved as internal calcareous moulds. The compaction of ammonites  
166 sampled in the pelagic part of the Vocontian Basin is relatively important, particularly for the  
167 phragmocone part of the shell. Consequently, some characteristics such as the thickness of  
168 whorl section cannot be observed and compared most of the time. The suture lines are not  
169 always well preserved (especially in pelagic deposits), however their systematic presence  
170 indicates the beginning of the body chamber. The abbreviations used for the measurements  
171 (mm) are: D=diameter; H=whorl height; W=whorl thickness; U=umbilical diameter; h=spiral  
172 hiatus; Nt/2=number of tuberculate (main) ribs per half whorl; Ni/2=number of smooth ribs  
173 (intercalatories) per half whorl. When measurements were taken at multiple diameters on the

174 same specimen, the angle alpha between each successive measurement has been shown in the  
175 tables.

176 **Analysis.** In the present work, the sample of *Crioceratites claveli* sp. nov. from the basis of its  
177 stratigraphic range (the first acme previously evoked, here the *C. claveli* lower interval  
178 Horizon) is the only sufficiently large and homogeneously preserved to be statistically  
179 analyzed; this would also allows avoiding time-averaging as much as possible. Statistical  
180 analyses and tests were carried out using PAST, version 3 (Hammer et al., 2001). The method  
181 used has already been described by Bersac and Bert (2012) and Bert (2014a), and will not be  
182 given in detail here again. A digital matrix was made from different measures of the  
183 quantifiable characters (continuous and discrete variables) and ratios of some of them.  
184 Continuous parameters tested were D, H, U and h. The specimens often have a side embedded  
185 in the rock and most are compressed post mortem, so the variables associated with W were  
186 not tested due to lack of effective (significance level not reached). This is also the case for the  
187 diameters of the ontogenic stages, which could not be integrated into the statistical analysis  
188 for the same reasons, but they were taken into account in the descriptive analysis.  
189 Conventional ratio H/D (relative height), U/D (relative umbilicus) and U/H were calculated.  
190 The discrete variables tested are Nt/2 and Ni/2. Significance level of probability tests was set  
191 at  $p=0.05$ . Univariate and bivariate analyses were employed to visualize and test the  
192 homogeneity of the sample, the possible variation through ontogeny, and correlations between  
193 parameters two by two. Multivariate analysis allows (1) investigation of the contribution of  
194 each variable relative to the total morphospace occupation of the sample, (2) investigation of  
195 the relationships between variables, and (3) to test groups recognized in the descriptive  
196 analysis (i.e. dimorphs) considering all the variables simultaneously.

197

198 **4. Description of the sections**

199

200 **4.1. The Cheiron ravine**

201

202 The palaeontological site of the Cheiron ravine has been known very probably since the  
203 beginning of the 19th century and has been visited for a very long time by local  
204 paleontologists (Astier, Jaubert, Duval-Jouve, Emeric, Audoul, Puzos, etc.). This long  
205 frequentation allowed the introduction of at least 30 species of ammonites from the Lower  
206 Cretaceous series, including zonal indexes [for example *Plesiospidiscus ligatus* (d'Orbigny,  
207 1841)], which make this site one of the most important for the Hauterivian stage in the south-  
208 east of France. Despite this historical past, the only study of this locality was from Cotillon  
209 (1971) who gave an imprecise synthetic geological section.

210

211 The Hauterivian lithology of the Cheiron ravine is composed of a relatively regular alternation  
212 of marlstone and limestone beds over 205 m thick, more compact at the top of the series. Only  
213 a limited slump area disturbs the Hauterivian succession in the lower part of the *Subsaynella*  
214 *sayni* Zone and causes an about 5 m thick invisibility. The beds and interbeds are thick in the  
215 *Acanthodiscus radiatus* Zone, and then they tend to thin out in the current *Crioceratites loryi*  
216 Zone (Fig. 2; = new *C. claveli* Zone, this work). The top part of the current *C. loryi* Subzone  
217 (= new *C. claveli* Subzone, this work) is formed by a small marly ledge between the beds  
218 194-198. A second marly ledge is located at the limit between the *Olcostephanus jeannoti*  
219 Subzone and the *Lyticoceras nodosoplicatum* Zone (beds 206-214). In the current *C. loryi*  
220 Zone (= new *C. claveli* Zone, this work), three beds are exceptionally rich in fossils: bed 179  
221 at the base of the zone (first acme of the genus *Crioceratites*), the very marly bed 195 (10 cm  
222 thick), which constitutes a Konzentrat-Lagerstätte type deposit (bioclastic accumulation), and  
223 bed 197 just below the *O. jeannoti* Subzone. Beds 195-197 form the second acme of the

224 genus. In bed 195, the fossils are very compressed and fragile structures such as spines are  
225 often kept in place; we also note the presence of many aptychi and very scattered grains of  
226 glaucony. The fauna is mainly made up of *C. claveli* sp. nov. in beds 179 and 195, which  
227 contain the First Occurrence (FO) and Last Occurrence (LO) of the species respectively, and  
228 by *C. loryi* in beds 196-197. In the other beds of the current *C. loryi* Subzone (= new *C.*  
229 *claveli* Subzone, this work), *Neolissoceras grasianum* (d'Orbigny, 1841) is often the more  
230 abundant. The first *Crioceratites curnieri* Ropolo, 1992 have been observed to date in bed  
231 201.

232

#### 233 **4.2. The La Charce-Pommerol sections**

234

235 A detailed lithology of the upper Valanginian-lower Hauterivian at La Charce was made  
236 previously by Reboulet et al. (1992), Bulot et al. (1993; 1995), Bulot (1995), Reboulet (1996;  
237 2015) and Reboulet and Atrops (1999). In the upper Valanginian, the *Saynoceras verrucosum*  
238 Zone is a marly-dominated interval; the relative abundance of carbonate in the marlstone-  
239 limestone succession increases gradually during the *Neocomites peregrinus* and  
240 *Criosarasinella furcillata* zones. Carbonate-rich marlstone-limestone alternations  
241 predominates the lower Hauterivian beds (*Acanthodiscus radiatus*, *Crioceratites loryi* and  
242 *Lyticoceras nodosoplicatum* zones).

243 In the La Charce section, the continuity of the marlstones-limestone alternation deposits is  
244 interrupted by two slumps in the lower Hauterivian observed at the top part of the *A. radiatus*  
245 Zone and across the boundary between the current *C. loryi* and *L. nodosoplicatum* zones.  
246 However, the stratigraphic intervals equivalent to these two slumps were correlated to the  
247 nearby Pommerol section characterized by undisturbed marlstone-limestone alternations  
248 (Reboulet, 2015; Aguirre-Urreta et al., 2019). These authors provided additional data allowing

249 to complete more precisely the lower Hauterivian ammonoid distribution previously  
250 established for the La Charce section.

251

## 252 **5. Descriptive palaeontology**

253

254 Class: Cephalopoda Cuvier, 1798

255 Order: Ammonoidea Agassiz, 1846

256 Suborder: Ammonitina Hyatt, 1889

257 Superfamily: Perisphinctoidea Steinmann, 1890

258 Family: Crioceratitidae Gill, 1871

259 Genus *Crioceratites* Léveillé, 1837

260

261 **Type species.** *Crioceratites duvalii* Léveillé, 1837.

262

263 ***Crioceratites jurensis* (Nolan, 1894)**

264 Figs. 3-5

265 **Synonymy.**

266 *vm* 1861 *Ancyloceras duvalii*, 'type 1'; Pictet and Campiche: p. 37-39, p. 49; pl. 47bis, fig. 1a-  
267 b.

268 1894 *Crioceras picteti* var. *jurensis* nov.; Nolan: p. 192.

269 1910 *Crioceras jurense*; Kilian: p. 270.

270 ?1936 *Crioceratites* sp. juv. gr. *jurensis*; Breistroffer: p. 145.

271 *non* 1967 *Crioceratites jurensis*; Dimitrova: p. 44; pl. XVIII, fig. 1.

272 **Holotype by monotypy.** Specimen MHNG GEPI 16787 (Fig. 3) of the Pictet's collection,  
273 curated in the Museum of Natural history of Geneva (Switzerland).

274 **Type locality.** The Ravin Saint-Martin area, near Escragnoles (Alpes-Maritimes,  
275 southeastern France).

276 **Type horizon.** Unspecified in the original work of Pictet and Campiche (1861), but the type  
277 specimen is said by these authors (p. 39) to come from a bed that yielded also *Acanthodiscus*  
278 *radiatus* (Bruguière, 1789), *Oosterella cultrata* (d'Orbigny, 1841), *Neolissoceras grasianum*  
279 (d'Orbigny, 1841) and *Lyticoceras cryptoceras* (d'Orbigny, 1840). In the type locality, the  
280 Hauterivian basis is condensed in a fossiliferous level with numerous *Acanthodiscus* Uhlig,  
281 1905 and *Leopoldia* Mayer-Eymar, 1887 (Kilian and Zürcher, 1896), and it is highly probable  
282 that this level has also delivered the type specimen of *Crioceratites jurensis*.

283 **Geographic distribution.** Southeastern France (neritic domain of the 'Arc de Castellane'),  
284 Switzerland ('Hauterive marlstones' in Pictet and Campiche, 1861, p. 49).

285 **Stratigraphic distribution.** New specimens were collected (LC and DB) at La Palud-sur-  
286 Verdon and Carajuan (Rougou, south-east of France) together with numerous *Leopoldia*  
287 above the beds rich in *Acanthodiscus radiatus* (Carajuan, bed 88 in Bulot et al., 1995, fig. 7),  
288 just below the gap of most of the current *C. loryi* Zone (= new *C. claveli* Zone, this work).  
289 This species is therefore restricted to the upper part of the *A. radiatus* Zone, without further  
290 details in the current state of knowledge (= *Leopoldia buxtorfi* Horizon of Bulot, 1995).

291 **Material studied (N=6).** Six specimens, all from the platform boarder of the 'Arc de  
292 Castellane' (South-East of France), including the cast of the holotype provided by Lionel  
293 Cavin (curator of the Museum of Natural history of Geneva). Two specimens (DBT.BC47 and  
294 DBT.BC48) are from the bed 88 of the Carajuan section (Rougou, Alpes de Haute-Provence  
295 department – see Bulot et al., 1995, fig. 7). Two specimens are from the same equivalent level  
296 in an unpublished section near La Palud-sur-Verdon (Alpes de Haute-Provence, DBT.BG31  
297 and LCT.7Z). The last specimen is from the locality of Bargème (unpublished section – Var

298 department, DBT.BC49), in a condensed bed assigned to the *A. radiatus* Zone with  
299 ferruginous ooliths.

300 **Description.** The shell is cryptoconic, rather weakly uncoiled, with a spiral hiatus which  
301 increases regularly with growth. The whorl section is relatively depressed, nearly sub-circular,  
302 more compressed in the inner whorls, but this is increased by the greater post mortem  
303 compression at this level. Taking into account the length of the adult body chamber (more  
304 than 1/3 of a whorl), the reconstructed size reached 120 mm (DBT.BC47 – Fig. 4) to 165 mm  
305 in diameter (LCT.7Z – Fig. 5; DBT.BG31). The ontogenetic sequence is as follows (note that  
306 the ornamentation of the ventral area is known on the studied specimens only from the  
307 ontogenetic stage 4):

308 - Stage 1 ‘initial trituberculate’ up to about 15-20 mm in diameter: the ornamentation  
309 is differentiated with main trituberculate ribs bearing large nodular tubercles; the lateral  
310 tubercle is located towards the mid-flank. The intercalary ribs are smooth and few in number.

311 - Stage 2 ‘with nodular periventral tubercles’ up to D=50-55 mm: the ornamentation is  
312 weakly differentiated, with main bituberculate ribs bearing tubercles near the umbilicus and  
313 on the periventral area. The latter tubercle is distinctly nodular. Fasciculations (several sickle-  
314 ribs that start from the same periumbilical tubercle) are exceptional and limited to a very few  
315 ribs only. The periventral tubercle is thorny compared to the periumbilical one; it is well  
316 marked and lengthened in the longitudinal direction. The ribs are slightly tipped forward in  
317 the middle of the flanks, which gives them a flexuous appearance.

318 - Stage 3 ‘with spiny periventral tubercles’, up to D=110-130 mm: the ribs are better  
319 differentiated with the total loss of the fasciculation and the acquisition of true trituberculate  
320 main ribs generally separated by 5-6 smooth interribs. The tubercles, however, remain  
321 inconspicuous on the material studied; the periventral tubercle loses its nodular appearance  
322 characteristic of the previous stage.



323 - Stage 4 'adult', up to D=150-165 mm. The ornamentation becomes more irregular  
324 than in the previous stage; the intercalary ribs are reinforced, spaced out, and their distinction  
325 from the main ribs becomes sometimes less obvious. The main ribs (and sometimes the  
326 intercalaries when they are strong) remain trituberculate with a lateral tubercle. They can be  
327 grouped into doublets separated by a weak constriction with a flat bottom. The bifurcations on  
328 the top of the flanks are rare. All the ribs cross the ventral area, and the main ones substantially  
329 strengthen. The body chamber (probably adult) starts between 90 and 110 mm in diameter.

330 The suture lines are very indented, of the ELUI type according to the terminology used by  
331 Wiedmann (1966), with a well marked and deep lateral lobe L. It is trifold and each of the three  
332 branches is narrow and well developed. The umbilical lobe U is also trifold, shallower than L  
333 and more stocky. The siphonal lobe E is shallower than L; it is bifid with elongated and  
334 narrow terminal branches. The following formula is respected:  $L > E > U$ .

335 **Discussion.** Dimorphism is not yet known in *C. jurensis*, probably due to a lack of specimens.  
336 By comparison with the other related species for which the dimorphic pairs have been  
337 recognised (*C. loryi* and *C. claveli* sp. nov.), it seems that the specimens described herein all  
338 belong to the macroconch dimorph.

339 The specimen from the Pictet's collection, shown by Pictet and Campiche (1861, pl. 47 bis,  
340 fig. 1), and which was later used as type for *C. jurensis* by Nolan (1894) was never refigured.  
341 It could be found by Lionel Cavin in the collections of the Museum of Geneva and the  
342 synonymy with *Davouxiceras nolani* proposed by Sarkar (1955), and followed by Klein et al.  
343 (2007), is rejected here. Actually, Pictet and Campiche (1861, p. 38-39, 49) themselves  
344 recognized the differences between this specimen (their fig. 1) and the one in their figure 2  
345 (Pictet and Campiche, 1861, pl. 47bis), which the latter would later become the type of *D.*  
346 *nolani*. The two species are clearly not of the same group, *C. jurensis* belonging to the genus  
347 *Crioceratites* s.str.

348

349 Thus, *Crioceratites jurensis* is also very different from *Davouxiceras coniferus* (Busnardo et  
350 al., 2003), although they can be found in the same levels. This latter species has a much more  
351 open coiling. New specimens (LC, DB and Poupon collections) show that it could reach a  
352 large to very large adult size, between 360 mm and up to probably around 500 mm. We do not  
353 recognize in *D. coniferus* the ontogenic stages described in *C. jurensis*: in the former, the  
354 trituberculation is constant with ribs always differentiated into very strong main ribs and  
355 smooth intercalaries, which seem to become scarce or even disappear on the adult body  
356 chamber. This is the morphology usually accepted by authors as characteristic of the group of  
357 '*Davouxiceras nolani*' sensu lato (Klein et al., 2007). The type specimen of *D. coniferus* and  
358 the specimen studied by Reboulet (1996; = *Crioceratites* n. sp. 1, pl. 24, fig. 7), reported to  
359 this species by Busnardo et al. (2003), are assigned to the lowermost part of the Hauterivian  
360 (*A. radiatus* Zone). Since then, Vašíček (2005) has reported *D. coniferus* in the upper  
361 Valanginian (*Criosarasinella furcillata* Zone and Subzone).

362

363 *C. jurensis* is closer to *C. loryi* (Fig. 6), with which it shares the general morphology of the  
364 shell, and with which it has many affinities. The adult size of *C. jurensis* appears to be slightly  
365 smaller than that of *C. loryi* (150-165 mm vs. 180 mm). The initial trituberculate stage is  
366 common to the both species, but it is distinctly shorter in *C. loryi* (5-7 mm vs. 15-20 mm).  
367 Stage 2 is slightly longer in *C. jurensis* (50-55 mm vs. 40 mm), but the characteristic  
368 fasciculations of *C. loryi* are almost absent or only limited to a few ribs. *C. jurensis* stage 3 is  
369 also longer (110-130 mm vs. 90 mm), with the presence of systematic trituberculate main ribs,  
370 which is not the case in *C. loryi* where the ribs are bituberculate. Finally, the adult stages 4 of  
371 the both species are quite similar, but the ribs are still trituberculate in *C. jurensis*. Note that  
372 this aspect is less evident in the type of *C. jurensis* due to wear of the specimen.

373

374 *C. jurensis* has a more regular coiling than *C. curnieri*, which varies between crioconic to  
375 subaspinoceratic poles. In stages 3 and 4, *C. jurensis* shows fairly constant trituberculate ribs  
376 pattern that is never the case with *C. curnieri*, which is bi- or even monotuberculate. The  
377 ornamentation of *C. jurensis* is markedly less regular than that of *C. curnieri*. Finally, *C.*  
378 *curnieri* is more recent, as it appears in the *O. jeannoti* Subzone and becomes more abundant  
379 in the *L. nodosoplicatum* Zone (Ropolo, 1992; Bulot, 1995; Reboulet, 2015; Aguirre-Urreta et  
380 al., 2019).

381

382 Some specimens of *Crioceratites heterocostatus* Mandov, 1976 could have an aspinoceratic  
383 coiling of the adult body chamber (see Reboulet, 1996, pl. 24, fig. 1), which is never the case  
384 in *C. jurensis*. *C. heterocostatus* is also smaller with a diameter up to 100 mm (vs. at least  
385 150-165 mm). We recognize almost the same succession of ontogenic stages (in particular  
386 stages 1 and 2, which are quite similar), with the exception of stage 3 ‘with spiny periventral  
387 tubercles’, absent in *C. heterocostatus*, which de facto prolongs stage 2 up to 60-70 mm in  
388 diameter (vs. 50-55 mm in *C. jurensis*). There are still other differences between these two  
389 species: the strongly trituberculate juvenile stage 1 is shorter in *C. jurensis* (15-20 mm vs. 20-  
390 30 mm); in adult stage 4, *C. heterocostatus* is bituberculate while *C. jurensis* is trituberculate;  
391 rib bifurcations on the upper flanks are rarer in *C. jurensis* (stage 4). On the other hand, the  
392 duplication of ribs on the adult body chamber is present in the both species, as in *C. loryi*.  
393 Following Vašíček (2005, p. 252, see also Klein et al., 2007, p. 44, note 31) and Company (in  
394 Klein et al., 2007, p. 40, note 30), we consider that *Crioceratites primitivus* Reboulet, 1996 is  
395 a junior synonym of *C. heterocostatus* Mandov 1976. This latter is reported from the  
396 Menhegini-Cryptoceras Zone of the Western Balkanids, which places it in the lower  
397 Hauterivian for Mandov (1976), while it would be more likely uppermost Valanginian for

398 Thieuloy (1977, p. 127). This older age, confirmed by Vašíček (1995, p. 171), reinforces the  
399 hypothesis of synonymy between the two species, with *C. heterocostatus* taking precedence.

400

401 ***Crioceratites claveli* sp. nov.**

402 Figs. 7-12

403 **Synonymy.**

404 1997 *Crioceratites* sp.; Faraoni et al.: p. 69, pl. 7, fig. 17.

405 **Derivation of name.** This species is dedicated to the memory of our French paleontologist  
406 colleague Bernard Clavel (1938-2018†).

407 **Zoobank Record.** LSID urn:lsid:zoobank.org:act:790B5346-0463-44D4-961B-  
408 B111EA4D17AA.

409 **Holotype.** Specimen DBT.BB48 (Fig. 7a-b) of the Bert's collection, curated by the  
410 RNNGHP.

411 **Type locality.** The Cheiron ravine, near Castellane, section CHE (Alpes de Haute-Provence,  
412 southeastern France).

413 **Type horizon.** The bed CHE/179 of the type locality, at the very basis of the *Crioceratites*  
414 *loryi* Zone in its current acceptance (= new *C. claveli* Zone, this work).

415 **Geographic distribution.** The species is currently known in south-east of France (Alpes de  
416 Haute-Provence and Drôme departments) and in Italy (Marche Apennines).

417 **Stratigraphic distribution.** The lower stratigraphic limit of *C. claveli* sp. nov. is the bed with  
418 the first acme of the genus *Crioceratites* at the lowermost part of the current *C. loryi* Zone (=  
419 new *C. claveli* Zone, this work) and its upper limit is in the second acme of the genus, just  
420 below the appearance of *C. loryi* at the top part of the current *C. loryi* Subzone (= new *C.*  
421 *claveli* Zone, this work and see also Reboulet, 1996, fig. 23). In the Cheiron ravine, this

422 distribution includes the beds CHE/179-195 and the beds LCH/219-229 in the La Charce  
423 section.

424 **Diagnosis.** Dimorphic crioconic shell with weak spiral hiatus in the inner whorls. Four stages  
425 during ontogeny: stage 1 'initial trituberculate' up to D=9 mm; stage 2 with 'nodular  
426 periventral tubercles' up to approximately D=20-35 mm for the microconchs and D=45-50  
427 mm for the macroconchs (differentiated ribs; main ribs bituberculate, exceptionally  
428 trituberculate; sometimes fasciculation with 2-3 ribs); stage 3 'with spiny periventral  
429 tubercles', up to D=56-70 mm in microconchs and D=95-115 mm in macroconchs, with less  
430 regular and flexuous ornamentation than in the previous stage (irregular appearance of  
431 trituberculation – lateral tubercle); stage 4 'adult', up to D=78-90 mm for the microconchs  
432 and D=135-145 mm, even 165 mm, for the macroconchs, with well individualized  
433 trituberculate main ribs and frequent duplicating of the main ribs. Stages 2-4 show long and  
434 highly developed spines on the periumbilical, lateral and periventral tubercles.

435 **Material studied (N=124).** Ninety specimens are from the Cheiron ravine (CHE) near  
436 Castellane (Alpes de Haute-Provence, France): 32 are from bed CHE/179, which are  
437 numbered DBT.BB48, DBT.BB49, DBT.BB51 to DBT.BB76 and DBT.BB92 to DBT.BB94  
438 (DBT.BB55 bears two specimens), 3 are from bed CHE/180, which are numbered  
439 DBT.BB71, DBT.BB74 and DBT.BB95, 1 is from bed 194 (DBT.BG49), and 54 are from  
440 bed 195 (DBT.BB79, DBT.BG49 to DBT.BG64 and DBT.BG67 to DBT.BG88, some  
441 numbers having several specimens). In addition some unnumbered fragmentary specimens  
442 where collected in beds 189, 190, 193 and 195 (Fig. 2). Height specimens are from Pommerol  
443 (POM), all from bed POM/219. They are temporarily numbered POM 219-2, 3, 5, 7, 8, 12, 20  
444 and 23. Twenty-six specimens are from La Charce (LCH). They are numbered FSL 487738  
445 (LCH/226), FSL 487736 a-j (LCH/220), FSL 487733 a-l (LCH/219), FSL 487739  
446 (LCH/225), FSL 487738 (LCH/226) and FSL 487743 a-b (LCH/229).

447 **Description.** The shell is heteromorphic, of the crioconic type. The spiral hiatus is weak in  
448 the inner whorls, but this parameter can be parasitized by the significant post mortem crushing  
449 of the specimens in the phragmocone, and it tends to grow significantly in the adult body  
450 chamber (Fig. 8). Dimorphism was recognized by the adult size and the diameter reached by  
451 the various ontogenic stages. There are 4 stages during ontogenesis:

452 - Stage 1 'initial trituberculate' up to  $D=9$  mm on average, with differentiated ribs, the  
453 main ones having large lateral tubercles at mid-flanks (Fig. 7f). The ribs are simple, straight  
454 or slightly retroverted. The section of the whorl is more rounded than on the rest of the shell.

455 - Stage 2 'with nodular periventral tubercles' up to approximately  $D=20-35$  mm for  
456 microconchs and  $D=45-50$  mm for macroconchs: the ornamentation is quite irregular from  
457 one specimen to another. The ribs are differentiated; the main ribs are bituberculate,  
458 exceptionally trituberculate (lateral tubercle very discreet when present – FSL 487736f).  
459 There is some variation in the strength and density of the ornamental pattern, and the more  
460 slender specimens have markedly less marked rib differentiation with very inconspicuous  
461 tubercles. Most often, fasciculation (several ribs starting from the same periumbilical  
462 tubercle) concerns only a small number of ribs (2-3); it is usually rare but more frequent in  
463 slender specimens. The periumbilical tubercle is thorny, slightly elongated in the direction of  
464 the rib which supports it. The periventral tubercle is quite nodular and sometimes groups  
465 together several ribs by fibulation (two lateral ribs which meet at the top of the flank – Fig.  
466 9b-c). When the preservation conditions allow it (as for example in the CHE/195 bed), it is  
467 possible to observe that all the tubercles (periumbilical, lateral when present and periventral)  
468 are extended by a triangular elongated spine, more or less thin, even for poorly developed  
469 tubercles (Fig 7b and 10 shows example of the lateral and periventral spines respectively; note  
470 that the periumbilical and lateral spines are unfortunately most often broken during the  
471 mechanical preparation process). The ribs are generally quite flexible: projected forward near

472 the umbilical wall, they straighten between the first third and half of the flank. They do not  
473 pass through the siphonal area or are much erased and broaden there because of the  
474 thickening-out associated with the nodular appearance of the adjacent periventral tubercle.

475 - Stage 3 'with spiny periventral tubercles', up to  $D=56-70$  mm in microconchs and  
476  $D=95-115$  mm in macroconchs, is close to the previous stage with an even less regular and  
477 flexuous ornamentation. The ribs are hardly differentiated in slender specimens. The tubercles  
478 are thorny and lose the nodular appearance of the previous stage. This stage is mainly marked  
479 by the irregular appearance of the trituberculation (lateral tubercle) from  $D=40$  and  $50$  mm  
480 respectively in the robust and slender microconch specimens, and  $D=55$  mm in the  
481 macroconchs. Note that the lateral tubercle is absent in very slender specimens. As in the  
482 previous stage, all tubercles are extended by long, well-developed spines.

483 - Stage 4 'adult', up to  $D=78-90$  mm for microconchs and  $D=135-159$  mm for  
484 macroconchs (reconstructed up to  $D=165$  mm on specimen DBT.BB52, see below – Fig. 8).  
485 The ornamentation becomes again a little more flexible than in the previous stage, with very  
486 well individualized trituberculate main ribs, especially in macroconchs. They are much  
487 reinforced with strong tubercles in robust specimens, especially the macroconchs, and a little  
488 less in slender specimens and microconchs. All the ribs strengthen on the ventral area.  
489 Duplication of the main ribs is frequent. In some microconchs, the ribs may become stronger  
490 and become slightly spaced. The adult body chamber starts at  $D=55-65$  mm in microconchs  
491 and  $D=94-100$  mm in macroconchs. At this stage again the tubercles are extended by spines.

492 **Dimorphism.** Among the 42 macroconchs identified, four appear almost complete and can be  
493 considered as adults (preservation of stage 4 with approximation of the reinforced main ribs  
494 and more marked adoral projection): (1) the specimen DBT.BB52 is the largest with  $140$  mm  
495 in diameter ( $D_{ph}=100$  mm – Fig. 8) and a body chamber preserved on  $1/4$  of a whorl  
496 (reconstructed at more than  $1/3$ , it would measure approximately  $D=165$  mm in diameter); (2)

497 DBT.BB48 (Fig. 7a-b) and (3) DBT.BG58 of which only the end of the body chamber is  
498 missing (respectively  $D_{ph}=94$  and  $99.82$  mm). (4) Specimen DBT.BG49 (Fig. 11) is an  
499 almost complete adult body chamber of  $159$  mm in diameter. Specimen DBT.BB55a (Fig. 9a)  
500 is also a macroconch, but it is almost wholly septate with a  $72$  mm diameter phragmocone,  
501 suggesting that it is probably an immature specimen. The other macroconchs are more  
502 fragmentary, but the ontogenic stage reached according to the diameter (often the ornamental  
503 'adult' stage present on the body chamber) makes it possible to identify them.

504 Forty-three sufficiently complete microconchs have been identified. The most complete are  
505 specimens POM 219.2 ( $D=86$  mm – Fig. 12a-b) and BDT.BB49 ( $D=78$  mm – Fig. 7c-d), with  
506 a complete (adult) body chamber more than  $1/3$  of a whorl and a phragmocone around  $55$  mm  
507 in diameter, and DBT.BG50 ( $D=78.4$  mm – Fig. 10a), which shows approximated septa. Five  
508 other specimens present their phragmocone and at least part of the body chamber  
509 (DBT.BB54, DBT.BB57 to DBT.BB61 and FSL 487736g – Figs. 7f-g, 9b-c). The last  
510 specimens are fragments of body chambers alone (DBT.BB74) or showing the transition with  
511 the phragmocone (DBT.BB75).

512 The last 38 specimens are either juveniles that have not reached a sufficient development to  
513 be able to be attributed to one or the other dimorph, or too incomplete specimens that do not  
514 show the diameter of transition between the ontogenic stages. In all cases, below  $45$  mm in  
515 diameter the both antidimorphs are identical. Visually, the dimorphs can only be separated on  
516 the basis of their adult size and the ontogenetic stages achieved according to the diameter, but  
517 neither by the dimensional parameters of the shell nor by the aspect of the ornamentation (at  
518 equivalent stage).

519 The suture line is very indented and follows the ELUI formula (terminology of Wiedmann,  
520 1966), with  $L>E>U$ . The lateral lobe  $L$  is trifid, well developed, relatively symmetrical and  
521 deep. The umbilical lobe  $U$  is also trifid, shallower than  $L$  and more stocky. The siphonal lobe



522 E is barely visible due to the compaction of the specimens. The lateral and umbilical saddles  
523 are well incised by an accessory lobe.

524 **Discussion.** The specimen of *Crioceratites* sp. described by Faraoni et al. (1997, p. 69, pl. 7,  
525 fig. 17) from the Marche Apennines (Mt Catria, Maiolica Formation, central Italy) could  
526 belong to the species *C. claveli* sp. nov. (microconch) because of its morphology, the presence  
527 of main trituberculate ribs at least from D=35 mm (the very inner whorls are poorly  
528 preserved), the presence of nodular lateroventral tubercles, which encompass several ribs  
529 before the onset of trituberculation, and the presence of long well-defined spines. This  
530 specimen is from bed 407 of the Chiaserna section assigned to the *A. radiatus* Zone (non-  
531 basal) by Faraoni et al. due to the presence of *Tescheniceras flucticulus* (Thieuloy, 1977) and  
532 *Spitidiscus meneghini* (Rodighiero, 1919) respectively in beds 392 and 405. The last terms of  
533 the geological section could actually be more recent, since at La Charce these species are  
534 reported up quite high in the *A. radiatus* Zone, and even up to the base of the current *C. loryi*  
535 Zone (= new *C. claveli* Zone, this work) for *S. meneghini* (Reboulet, 1996, fig. 23). Under  
536 these conditions, it is highly probable that bed 407 of the Chiaserna section is to be assigned  
537 to the base of the current *C. loryi* Zone (= new *C. claveli* Zone, this work).

538  
539 *C. claveli* sp. nov. is closely related to *C. loryi*. The double acme in the distribution of *C. loryi*  
540 (LCH/219-220 and LCH/230-235) recorded by Reboulet (1996) at La Charce and at the  
541 Cheiron ravine (CHE/179-180 and CHE/194-195 – this work) is a result which led to a more  
542 careful attention in order to verify whether these two acmes could correspond to two different  
543 species or if the total vertical distribution corresponds to the solely species *C. loryi*. The  
544 present work clearly goes in favor of the first hypothesis. The differences between *C. claveli*  
545 sp. nov. and *C. loryi* are mainly focused on the ontogenic stages and the systematic presence  
546 of highly developed spines in *C. claveli* sp. nov. (Figs. 7a-b, c, g, 10a-c). In *C. loryi*, the

547 'initial trituberculate' stage 1 is shorter (5-7 mm), stage 2 is very characteristic with the  
548 systematic presence of numerous fasciculated dense and homogeneous ribs, almost all similar,  
549 starting from the periumbilical tubercle. This fasciculate pattern is much rarer in *C. claveli* sp.  
550 nov. and always limited to a few ribs only. The stage 3 in *C. loryi* shows very discrete  
551 trituberculation only on the extremely rare hyper-robust specimens; the ornamentation is more  
552 regular with a periumbilical tubercle always more discreet than in *C. claveli* sp. nov. Finally,  
553 in stage 4 (adult) of most macroconchs (and much more discreetly in microconchs and hyper-  
554 slender macroconchs) the main ribs are clearly trituberculate in *C. claveli* sp. nov., with a  
555 sometimes very developed lateral tubercle extended by a long spine (as for example in the  
556 specimen BB48 – Fig. 7a-b), while the ribs are most often bituberculate in *C. loryi*. In some  
557 rare macroconchs of this latter species, trituberculation can be developed irregularly and  
558 fleetingly at the end of stage 4 with an always inconspicuous and unsystematic lateral tubercle  
559 that never shows the presence of a spine. In terms of adult size, *C. loryi* appears to be larger  
560 (105 mm versus 90 mm for microconchs and 180 mm versus 165 mm for macroconchs, see  
561 Fig. 6a), although this maximum size is not expressed in all adult macroconchs of *C. loryi*.

562  
563 *C. curnieri* has a more variable coiling between a crioconic to a subaspinoceratic pole.  
564 However, the coiling of *C. claveli* sp. nov. is narrower and more regular since *C. curnieri*  
565 shows an increase in the uncoiling of the adult body chamber at the start of a shaft and a short  
566 hook (elliptical coiling), even in the case of the crioconic morphology. In all cases, *C. curnieri*  
567 is bituberculate, or even monotuberculate, without ever being trituberculate as in *C. claveli* sp.  
568 nov. Tuberculation is always very inconspicuous, unlike *C. claveli* sp. nov. in which every  
569 tubercles is prolonged by a well-developed spine. The ornamentation of *C. curnieri* is more  
570 regular in the alternation between the main and intercalary ribs. The periventral tubercles  
571 appear later in ontogeny, after the periumbilical tubercles (Ropolo, 1992, p. 67), which is not

572 the case in *C. claveli* sp. nov. where the innermost whorls are trituberculate (stage 1). Then,  
573 the lateral tubercle disappears, making concomitant both the two periventral and periumbilical  
574 tubercles (stage 2), before its subsequent reappearance at stage 3. Adult stage 4 is also  
575 different, despite the strengthening of the main and secondary ribs in *C. curnieri*, which are  
576 more and more similar and spaced (sometimes bifurcated on the top of the flanks as can also  
577 be observed in some specimens of *C. claveli* sp. nov.), and despite the presence of  
578 constrictions. The ribs, on the other hand, remain there clearly differentiated and more spaced  
579 in *C. claveli* sp. nov.; in macroconchs, the tubercles are clearly reinforced. Finally, *C. curnieri*  
580 is more recent, from the *Olcostephanus jeannoti* Subzone, and mostly in the *Lyticoceras*  
581 *nodosoplicatum* Zone (Ropolo, 1992; Bulot, 1995).

582  
583 *C. claveli* sp. nov. is very close to *C. jurensis*, with an adult size that could be comparable on  
584 the basis of our current data. The coiling is identical between the two species. Ontogenically,  
585 the stages are on average shorter in *C. claveli* sp. nov. (stage 1: 9 mm vs. 15-20 mm; stage 2:  
586 45-50 mm vs. 50-55 mm; stage 3: 95-115 mm vs. 110-130 mm). Both species are  
587 trituberculate at stage 3, from roughly the same diameter. In *C. claveli* sp. nov., the  
588 ornamentation is generally less regular, with slightly more flexuous ribs. Rib bifurcations at  
589 the umbilical tubercle (fasciculation) are only exceptional in *C. jurensis*. Unlike *C. claveli* sp.  
590 nov., we do not observe in *C. jurensis* the peculiar appearance of the periventral tubercles  
591 which encompass several ribs so as to form a fibula, although they are nodular in both  
592 species. Rib bifurcations on the upper flanks of the adult body chamber (stage 4) are rare in  
593 the both species. Finally, *C. jurensis* is older than *C. claveli* sp. nov. (late *Acanthodiscus*  
594 *radiatus* Zone).

595

596 *C. claveli* sp. nov. has certain affinities with *C. heterocostatus* (= *C. primitivus*). However,  
597 some specimens of *C. heterocostatus* present a marked uncoiling of the body chamber, which  
598 gives them a general elliptical shape (aspinoceratic, see Reboulet, 1996, pl. 24, fig. 1); this is  
599 never the case in *C. claveli* sp. nov. whether in micro- or macroconch specimens. The  
600 ontogeny of *C. heterocostatus* is different, with a single trituberculate stage (stage 1) longer in  
601 the innermost whorl (20-30 mm vs. 9 mm), whereas in *C. claveli* sp. nov. the lateral tubercle  
602 comes back at stage 3 to be well expressed in the adult body chamber. Ornamentally  
603 speaking, in *C. heterocostatus* the fasciculate pattern of stage 2 is made with bundles of  
604 usually two ribs starting from the umbilical tubercle. But unlike *C. claveli* sp. nov. in which  
605 we also recognize this tendency, in *C. heterocostatus* these tubercles show a characteristic  
606 lengthening in the direction of the rib that carries it (character known in the ‘*Sarasinella*’  
607 Uhlig, 1905 of the late Valanginian – see Reboulet, 1996, p. 170). *C. heterocostatus* shows  
608 more radial ribs than in *C. claveli* sp. nov., but also more bifurcations of the ribs on the top of  
609 the flanks at stage 4 ‘adult’ (without being numerous). Ornamentation is interrupted at the  
610 siphonal edge, even at stage 4 (except in putative macroconchs), unlike *C. claveli* sp. nov.  
611 where all the ribs systematically pass over the ventral area on the body chamber. Both species,  
612 as well as *C. loryi*, show the duplications of the main ribs on the adult body chamber (stage  
613 4). *C. claveli* sp. nov. consistently has long and developed spines from all tubercles, which  
614 does not appear to be the case in *C. heterocostatus*. Finally, *C. heterocostatus* is older than *C.*  
615 *claveli* sp. nov., since this species is recorded in the late Valanginian *Criosarasinella*  
616 *furcillata* Zone.

617

618 The relatively open coiling with a fairly low growth in whorl height known in *D. coniferus*,  
619 associated with the total absence of fasciculation and the constant trituberculation of the main  
620 ribs throughout ontogeny, which are regularly interspersed with numerous smooth ribs,

621 discard this species from the genus *Crioceratites* stricto sensu to which *C. claveli* sp. nov.  
622 belongs.

623

## 624 **6. Biometric and statistical analyzes of *Crioceratites claveli* sp. nov.**

625

### 626 **6.1. Univariate and bivariate analyses**

627

628 The three histograms H/D, U/D and U/H (Fig. 13) of the univariate analysis are clearly  
629 unimodal, relatively symmetrical and agree with a normal distribution (Shapiro-Wilk tests  
630 respectively of  $W=0.9823$ ,  $p=0.8215$ ;  $W=0.9705$ ,  $p=0.4585$  and  $W=0.9793$ ,  $p=0.7359$ ): the  
631 assumption of the homogeneity of the sample is not challenged on the basis of the univariate  
632 analysis.

633

634 In the framework of the bivariate analysis (Fig. 14), the regression curves of the relations  
635 between H and U as a function of D show an harmonic growth according to the function  
636  $Y=bD^a$ , with  $a$  significantly different from 1 ( $p<<0.05$ ), indicating an allometric growth of the  
637 shell. The dispersion is low around the mean (coefficients of determination  $R^2>0.98$ ): the  
638 results are therefore very significant and the hypothesis of the homogeneity of the sample is  
639 reinforced (one single growth curve). However, we observe a slight disparity in the plot  
640 distribution between microconch and macroconch specimens on the diagrams for the values  
641 of U (mainly U/H), with a cluster associated with the high values of U/D and U/H in the  
642 microconchs. This latter ratio was recommended by Marchand (1986) to detect a possible  
643 bimodality associated with sexual dimorphism. In ammonites, dimorphism is generally  
644 characterized by a difference in coiling between the antidimorphs: at the acquisition of sexual  
645 maturity, and at equivalent diameter, the microconch generally has a more open umbilicus due

646 to the decreasing whorl height from its dorsal edge. Here, this feature is very weak and was  
647 not highlighted by the descriptive analysis only; this shows the importance to perform a  
648 biometric analysis whenever possible.

649 The  $h=f(D)$  diagram shows a very strong linear increase in the spiral hiatus as a function of  
650 the diameter, but the dispersion is important ( $R^2=0.68$ ), which shows the great variability of  
651 this parameter.

652 The  $Nt/2$  and  $Ni/2$  diagrams show a relative stability in the number of ribs per half-whorl  
653 during growth, although this parameter is extremely variable (large dispersion around the  
654 mean curve).

655

## 656 **6.2. Multivariate analysis**

657

658 A Principal Component Analysis (PCA – Fig. 15) was performed. The parameters  $h$  and  $Nt/2$   
659 have a distribution that significantly departs from normality (grayed out in Table 3). The  
660 Dornik and Hansen omnibus test shows departure from normality for the data taken as a  
661 whole ( $p \ll 0.05$ ), which requires the log transformation of the data to practice the PCA in  
662 order to normalize their distribution, reduce heteroscedasticity and resolve the problem of  
663 non-additivity of the independent variables, without modification of the variance. The  
664 correlation matrix was therefore used in order to standardize the data.

665 The first 3 principal components of the PCA have an eigenvalue close or greater than 1 (4.05,  
666 1 and 0.95), and they retain 85.72% of the total inertia of the variance (Table 4), of which  
667 72.17% for PC1 and PC2 alone. The analysis is therefore considered relevant. The results are  
668 shown in Figure 15 and Table 5. The ornamentation vectors are orthogonal to the shell  
669 morphological vectors, indicating an apparent lack of correlation between these vector groups,  
670 as previously recognized for some other *Crioceratites* (Ropolo, 1995; Bert, 2014a). The

671 diagram also shows an inverse correlation between H/D and the other morphological  
672 parameters: the relative height decreases as a function of (1) the diameter of the shell, (2) the  
673 opening of the umbilicus and (3) the spiral hiatus. These correlations are consistent with the  
674 descriptive analysis and to what is expected under the intraspecific laws in ammonites (Bert,  
675 2014a), even considering the lack of W data. In contrast, the PCA diagram shows an inverse  
676 relationship between the number of primary and secondary ribs, which was not demonstrated  
677 by descriptive analysis; this apparent relationship needs to be tested (see below the correlation  
678 by Spearman rank).

679 The convex hulls formed by the antidimorphs recognized during descriptive analysis overlap  
680 widely. Even though the distribution of plots shows that macroconchs are more influenced by  
681 PC1 than microconchs, the distinction observed in bivariate analysis (Fig. 14) with respect to  
682 the high values of U/H and U/D in microconchs is not found in the plot distribution on the  
683 PCA (Fig. 15). This lack of graphic discrimination is essentially due to the fact that the U/H  
684 and U/D vectors have the same orientation as the D and h vectors, which favors the  
685 superposition of the plots associated with the microconchs with those of the large  
686 macroconchs. Thus, the largest positive values of the parameters related to the size (D and h)  
687 override the parameters related to the opening of the umbilicus. On PC2, on the other hand,  
688 the microconchs are even more influenced than the macroconchs by the ornamental  
689 parameters. Indeterminate antidimorphs form a group opposed to the D vector, which is  
690 logical since only very small specimens are affected by this group (see reasons in descriptive  
691 section, Chapter 5).

692

693 In addition to the PCA, a correlation test by Spearman rank ( $r_s$ ) was performed in order to  
694 verify and quantify the relationships between the variables. The results (Table 6 and Fig. 15)  
695 confirm the descriptive analysis and partially the PCA; they clearly show for this group of

696 *Crioceratites* that the ornamental parameters are independent from the shell morphological  
697 parameters (non-covariant factors – see also Bert, 2014b, p. 234). The analysis of the PCA  
698 alone could suggest that the number of main tuberculate ribs would be inversely correlated  
699 with the number of smooth ribs, but this point is not confirmed by the correlation by  
700 Spearman rank ( $r_s=0.0099851$ ;  $p=0.96764$ ). Ornamental parameters also do not depend on the  
701 diameter.

702 On the other hand, the strong correlation between the morphological parameters of the shell  
703 (including the spiral hiatus  $h$ ) and the diameter is confirmed; it is positive in all cases except  
704 for the relative height, which decreases proportionally with the size.

705 The frequency histogram on PC1 is unimodal and normal ( $W=0.9734$ ,  $p=0.5242$  – Fig. 15). In  
706 contrast, the histogram attached to PC2 is clearly not normal ( $W=0.9095$ ,  $p=0.006$ ). The  
707 explanation lies to the fact that the two dimorphs are well discriminated by the analysis, even  
708 if they do not stand out very clearly graphically on the PCA, whose distribution of points is  
709 rather homogeneous.

710  
711 The morphological differences between the antidimorphs recognised by the descriptive  
712 analysis were statistically tested by an analysis of similarity (ANOSIM, Hammer and Harper,  
713 2006 – Table 7) with 100,000 permutations. This non-parametric test is more suitable for  
714 small samples than a multivariate non-parametric analysis of variance (NPMANOVA –  
715 Anderson, 2001). To ensure the homoscedasticity between the variables, they were  
716 standardized to mean zero and unit standard deviation (Quinn and Keough, 2002, p. 415). Our  
717 data are standardized measures, thus Euclidean distance was used to perform the ANOSIM  
718 (Quinn and Keough, 2002, p. 409). Pairwise comparison was performed only as a post hoc  
719 test (in case of  $p<0.05$ ) with sequential Bonferroni significance of pairwise comparison  $p$   
720 values (Table 8) in order to reduce type I errors (Hammer and Harper, 2006, p. 36).



721

722 Consistently with the rest of the study, statistically significant differences were found between  
723 the both recognized antidimorphs based on descriptive analysis, but logically not with the  
724 indeterminate group which acts somewhat as a control group.

725

## 726 **7. Intraspecific variation in *Crioceratites claveli* sp. nov.**

727

### 728 **7.1. Background on the laws of intraspecific variability in ammonites**

729

730 Ammonites shells are known to display a sometimes significant morphological intraspecific  
731 variability (see synthesis in De Baets et al., 2015). The laws of intraspecific variability in  
732 ammonites can be generalized (Bert, 2019) and relate to at least 8 potential factors of  
733 variation, separated into two categories:

734 - (1) Discontinuous variability:

735 - changes between sexual dimorphs can affect shell size, coiling, peristome  
736 appearance, ornamentation, or combinations;

737 - non-sexual polymorphism can concern a particular character (for example:  
738 supernumerary tubercle);

739 - fluctuating asymmetry results from the inability of species to absorb ecological stress  
740 during their development;

741 - dextrality/senestria is linked to chirality.

742 - (2) Continuous variability:

743 - variations in the rate of development modulate the expression of ontogeny  
744 (heterochronies) and can restructure the shell by simple shift if the stages involved are of  
745 sufficiently different morphologies (see synthesis in De Baets et al., 2015);

746 - the secondary erasing of the ornamentation does not seem to depend on any other  
747 intrinsic factor (probable role of the environment – see Bersac and Bert, 2012);

748 - variation in the coiling (heteromorphy) is unstable depending on the groups, and  
749 could also reflect the influence of the environment (Reboulet et al., 2005; Bert and Bersac,  
750 2013);

751 - multipolar variation is one of the most important in ammonites. It results from the  
752 generalization in two (<sup>2</sup>) or three (<sup>3</sup>) poles of the historical law known as the ‘1st Buckman  
753 law’ (Westermann, 1966) now outdated, which shows a correlation (to varying degrees and  
754 not always in the same direction) between the robustness of the ornamentation, the thickness  
755 of the section, the height of the whorl and the opening of the umbilicus; the extreme poles are  
756 interconnected by all the intermediaries. Other examples of multipolarity (order<sup>8</sup>) have been  
757 described recently in the literature in the Crioceratitidae *Balearites* Sarkar, 1954 /  
758 *Pseudothurmannia* Spath, 1923 (Matamales-Andreu and Company, 2019), but they arise from  
759 problems of interpretation by nesting discontinuous (non-sexual polymorphism concerning a  
760 supernumerary tubercle) and continuous factors (heteromorphy and ‘classic’ multipolar  
761 variation of order<sup>2</sup>).

762 These different factors of variation may or may not be expressed simultaneously and be  
763 covariant with each other (example of morphological-dependent heterochronies – Bert,  
764 2014a). A few exceptional cases seem to deviate from these laws: the significant anarchic  
765 variation of *Chelonicerias cornuelianum* (d'Orbigny, 1841) (Douvilleiceratidae Parona and  
766 Bonarelli, 1897) from the lower Aptian of the Argiles à Plicatules Formation (Bersac and  
767 Bert, 2018), can be explained by the context of exceptional environmental stress in connection  
768 with the Oceanic Anoxic Event 1a (OAE1a – Schlanger and Jenkyns, 1976; Arthur et al.,  
769 1985; and see discussion in Bersac and Bert, 2018). Subsequently, the *Douvilleicerias* de  
770 Grossouvre, 1894 took up a classic multipolar variation (Courville and Lebrun, 2010).

771 In the case of the *Crioceratites*, the observed variability relates to dimorphism, multipolar  
772 variation and heterochronic variation. Significant variations in the type of coiling, such as for  
773 example in *C. curnieri* (see Ropolo, 1992) have not been observed in *C. claveli* sp. nov.

774

## 775 **7.2. Laws of intraspecific variability applied to *C. claveli* sp. nov.**

776

### 777 **7.2.1. Dimorphism**

778

779 The study of abundant material of *C. claveli* sp. nov. (this work) made it possible to highlight  
780 the presence of two distinct morphologies, interpreted here as sexual dimorphs (for example,  
781 compare Figs. 7a-b, 8 versus Figs. 7c-d, 12a-b). Both the two antidimorphs are present on the  
782 same sites studied, in the same samples, have the same stratigraphic distribution, present the  
783 same ornamental characteristics, substantially the same characteristics of shell morphology  
784 and above all have the same ontogenetic succession (except for heterochronic variations).  
785 This dimorphism was clearly recognized in the descriptive analysis based on the diameter  
786 reached by the different ontogenetic stages by the two antidimorphs. Biometrics (bivariate  
787 analysis – Fig. 14) and statistical analyzes (ANOSIM – Table 7) subsequently made it  
788 possible to reinforce this result and to highlight small differences in construction concerning  
789 the relative umbilicus (U/H and U/D), not detected by the descriptive analysis alone: at  
790 equivalent diameter, the microconchs have values of these parameters very slightly higher  
791 than those of the macroconchs. Note that before the end of ontogenetic stage 2, at an average of  
792 about 45 mm in diameter, it is not possible to discriminate the antidimorphs. Without taking  
793 into account individuals which are too small (1/3 of the sample) the observed sex ratio seems  
794 close to 50%. Based on our study sample, the size ratio shows that macroconchs are on  
795 average around twice the size of microconchs. This order of magnitude is relatively consistent

796 with that observed in *C. loryi* (compare Fig. 6a and c). These results confirm those obtained  
797 by Ropolo (1995) who had already recognized the existence of a dimorphism in *Crioceratites*,  
798 including *C. loryi* and *C. curnieri* among others, and the work of Reboulet (1996) which  
799 seemed to show its existence in *C. heterocostatus* (= *C. primitivus* in Reboulet, 1996). The  
800 microconch dimorph has so far not been clearly demonstrated in *C. jurensis* due to a lack of  
801 specimens. By comparison with related species (*C. claveli* sp. nov., *C. loryi*, *C. curnieri* and  
802 probably also *C. heterocostatus*), where it is known, the existence of dimorphism in this  
803 species is highly probable on the same way.

804

### 805 **7.2.2. Multipolar variation**

806

807 The inverse relationship between relative height (H/D) and relative umbilicus (U/D) has been  
808 well demonstrated by multivariate analysis (Fig. 15; Table 6). This relationship is usually  
809 known as the Buckman's 1st law of variation; it is common to most ammonites and represents  
810 one aspect of the classical continuous multipolar variation (Bert, 2014a). There is also some  
811 variation in the strength and density of the ornamental pattern: the more slender specimens  
812 have markedly less rib differentiation with very inconspicuous tubercles. On the other hand,  
813 the multivariate analysis shows an absence of correlation between the ornamental parameters  
814 and the morphology of the shell.

815 Despite of this, continuous variation in *C. claveli* sp. nov. looks to be of the dipolar type  
816 (order<sup>2</sup>), even if caution is required because of the lack of parameters based on the values of  
817 W in the analysis (this work). However, it is clear that the importance of W in the  
818 intraspecific variation in *Crioceratites* has never been highlighted in the literature.

819

### 820 **7.2.3. Heterochrony**

821

822 While the appearance diameter of the ornamental stages does not seem to depend on the  
823 morphological parameters of the shell in *C. claveli* sp. nov., this is not the same for the  
824 heterochronic variation. It was observed in the descriptive analysis, at least in the  
825 microconchs, that the appearance of trituberculation of stage 3 during ontogeny is slightly  
826 fluctuating and depends on the robustness of the ornamentation. Ontogeny is slightly delayed  
827 for robust specimens and slightly accelerated for slender specimens.

828

## 829 **8. Biostratigraphic implications**

830

831 The discovery of *C. claveli* sp. nov. and the new understanding of the stratigraphic repartition  
832 of *C. loryi*, lead us to propose some changes in the biostratigraphic chart currently admitted  
833 for the early Hauterivian. This proposal for the south-eastern France may have extension  
834 potential for the Mediterranean Province of the Mediterranean Caucasian Subrealm (Tethyan  
835 Realm).

836

### 837 ***Crioceratites claveli* Zone (new)**

838 **Index species.** *Crioceratites claveli* sp. nov.

839 **Reference section.** The Cheiron ravine (CHE, see Chapter 4.1) in the Castellane area  
840 (southeastern France, Vocontian Basin) with the beds interval 179-212. This interval is  
841 correlable to the bundle comprising the beds 219-241e of the La Charce-Pommerol sections  
842 (see Chapter 4.2 and Aguirre-Urreta et al., 2019, fig. 8). Note that at La Charce, the bed  
843 succession is disturbed by a slump at the boundary between the *C. claveli* (new) and *L.*  
844 *nodosoplicatum* zones.

845 **Status.** The finding of *C. claveli* sp. nov. in a stratigraphic position that was once considered  
846 to be the first acme of *C. loryi*, gives rise to some reflections since it appears that *C. claveli*  
847 sp. nov. is de facto the base marker of the current *C. loryi* Zone (for example see Bulot, 1995;  
848 Reboulet, 1996). Note that *C. claveli* sp. nov. shows a vertical distribution widely beyond its  
849 FO, since it extends to the second acme of the genus (at the top part of the current *C. loryi*  
850 Subzone; see Reboulet, 1996 for the quantitative approach of the 2 acmes). *C. loryi* does not  
851 exist at the base of its current zone and it only appears during the second acme of the genus,  
852 to disappear a short time later.

853 These findings allow us to propose the new *C. claveli* Zone instead of the *C. loryi* Zone for  
854 the Vocontian Basin (Table 9). Its lower limit is defined by the First Appearance Datum  
855 (FAD) of *C. claveli* sp. nov. and its upper limit by the FAD of *L. nodosoplicatum* (in the  
856 Cheiron reference section the FO of *C. claveli* sp. nov. is recorded in bed 179, and in the bed  
857 219 of the La Charce and Pommerol sections). This proposal has the advantage of avoiding  
858 any possible confusion between the former *C. loryi* Zone and a new definition that the  
859 revision of the stratigraphic distribution of its index species would have led.

860 The stratigraphic distribution of *C. claveli* sp. nov. is precisely known, and because of its high  
861 frequency (mainly in pelagic deposits; there apparent absence from the peri-Vocontian  
862 platform borders are due to lacks in the deposits – see Bulot, 1995, p. 93-94) it is a serious  
863 and legitimate candidate to the status of index species. Its potential for long-distance  
864 correlations has now to be tested outside the Vocontian Basin, but its presence is also attested  
865 in Italy (Faraoni et al., 1997 and see descriptive section in Chapter 5).

866 **Faunal assemblage.** See below for the new *C. claveli* Subzone, and Reboulet (1996, fig. 23)  
867 and Aguirre-Urreta et al. (2019, fig. 8) for the *O. jeannoti* Subzone.

868

869 ***Crioceratites claveli* Subzone (new)**

870 **Index species.** *Crioceratites claveli* sp. nov.

871 **Reference section.** The Cheiron ravine (CHE, see Chapter 4.1) in the Castellane area  
872 (southeastern France, Vocontian Basin), with the beds interval 179-197. This interval is  
873 correlable to the bundle of beds 219-231 of the la Charce-Pommerol sections (see Chapter 4.2  
874 and Aguirre-Urreta et al., 2019, fig. 8).

875 **Status.** For the same reasons as for the new *C. claveli* Zone, it is necessary to replace the *C.*  
876 *loryi* Subzone with the new *C. claveli* Subzone. Its lower limit is defined by the FAD of *C.*  
877 *claveli* and its upper limit by the FAD of *O. jeannoti*. Within the new *C. claveli* Subzone of  
878 the Cheiron ravine and the La Charce-Pommerol sections, certain beds or bundles of beds are  
879 noticeably rich in representatives of the genus *Crioceratites*. We propose to recognize them as  
880 interval horizons; they are as follows:

881 - (1) the *C. claveli* lower interval Horizon (CC1 in Fig. 2): at the Cheiron ravine, it is  
882 represented by the bed 179 (= bed 219 in the La Charce-Pommerol sections), of which the  
883 lower limit is the First Occurrence (FO) of *C. claveli*. This horizon corresponds to the first  
884 acme of the genus *Crioceratites* evoked previously;

885 - (2) the *C. claveli* upper interval Horizon (CC2 in Fig. 2): at the Cheiron ravine, it is  
886 represented by the beds 194 and 195 (their exact correspondence in the La Charce-Pommerol  
887 sections is not yet clear, probably around the bed 229). The latter is the bed with the Last  
888 Occurrence (LO) of *C. claveli* sp. nov.;

889 - (3) the *C. loryi* interval Horizon (CL in Fig. 2): at the Cheiron ravine, it is  
890 represented by the beds 196 (containing the FO of *C. loryi*) and 197 (= beds 230-231 in the La  
891 Charce-Pommerol sections).

892 Both the upper *C. claveli* and *C. loryi* interval horizons represent the second acme of the  
893 genus *Crioceratites* evoked previously.

894 **Faunal assemblage.** *C. claveli* sp. nov. dominates largely the ammonite faunas of both the  
895 two *C. claveli* lower and upper interval horizons. The index species is present but rarer in the  
896 rest of its zone, especially compared to *Neolissoceras grasianum* for example. The other  
897 ammonites found in the subzone are (Fig. 2): *Phyllopachyceras winkleri* (Uhlig, 1882), *Ph.*  
898 *infundibulum* (d'Orbigny, 1841), *Hypophylloceras tethys* (d'Orbigny, 1841), *Protetragonites*  
899 *cf. quadrisulcatus* (d'Orbigny, 1841), *Olcostephanus sayni* (Kilian, 1896), *Spitidiscus*  
900 *meneghinii*, *Bochianites kiliani* Turner, 1962, *Himantoceras* sp., *Oosterella* sp. and *Saynella*  
901 *aff. clypeiformis* (d'Orbigny, 1841). In the *C. loryi* interval Horizon, the latter species replaces  
902 *C. claveli* sp. nov. and dominates largely the ammonite fauna. We can also notice the presence  
903 of *Lyticoceras* sp. nov., *Valdedorsella* sp. and *Davouxiceras aff. sablieri* (Astier, 1851).

904

## 905 **9. Conclusions**

906

907 (1) The research carried out in the Lower Cretaceous series of south-eastern France  
908 (Vocontian Basin) made it possible to highlight the two acmes of the genus *Crioceratites*  
909 during the former *C. loryi* Zone (early Hauterivian), so far reported to the single species *C.*  
910 *loryi*. The first acme marks the beginning of the zone and was considered to correspond more  
911 or less to the appearance of the genus *Crioceratites* in the early Hauterivian despite the  
912 existence of older representatives of this genus in the late Valanginian (Reboulet, 1996).  
913 However, only *Crioceratites* of the 2nd acme were depicted in the literature. The new  
914 material from the classic Cheiron ravine site (Alpes de Haute-Provence), as well as the  
915 revision of the material cited by Reboulet (1996) from the former *C. loryi* Zone of the La  
916 Charce (Hauterivian GSSP) and Pommerol (Drôme) reference sections, revealed that only  
917 part of the second acme actually belongs to *C. loryi*. The *Crioceratites* from the interval  
918 between the first acme and the first part of the second acme belongs to *C. claveli* sp. nov. only



919 (including the specimen of *C. sp.* depicted by Faraoni et al., 1997); *C. claveli* sp. nov. and *C.*  
920 *loryi* relay each other in the stratigraphy. An interesting point is that *C. claveli* sp. nov. shows  
921 a great morphological and ornamental stability between its appearance and disappearance,  
922 while the shift to *C. loryi* seems particularly sudden.

923

924 (2) *C. claveli* sp. nov. and *C. loryi* are differentiated mainly by the duration of the ontogenic  
925 stages reached during growth and by the presence of a marked spiny trituberculation in *C.*  
926 *claveli* sp. nov. The abundant material from the *C. claveli* lower interval Horizon (the first  
927 acme) allowed a detailed analysis of the variation of this species for a very short period of  
928 time (avoiding time-averaging). Under the laws of intraspecific variability in ammonites, 3  
929 factors are currently known in *C. claveli* sp. nov.:

930 - the dimorphism has been recognized by palaeontological analysis (description of the  
931 both macro- and microconch morphs), biometry (discrimination of the both dimorphs on the  
932 U/D and H/D crossplots) and statistics (ANOSIM and pairwise comparison). The observed  
933 sex ratio is around 50%;

934 - the multipolar variation of order<sup>2</sup> with an inverse correlation between the relative  
935 height and the relative umbilicus between two poles (morphological parameters) corresponds  
936 only partially to the classic Buckmann's 1st law of covariation since the robustness of the  
937 ornamentation (ornamental parameter) is here independent from the shell morphological  
938 parameters (non-covariant factors);

939 - the heterochronies concern the onset diameters of the ontogenic stages depending on  
940 the robustness of the ornamentation but not on the morphological parameters of the shell  
941 (ontogeny is slightly delayed for robust specimens and slightly accelerated for slender  
942 specimens).

943

944 (3) The analysis of the stratigraphic distribution of *C. claveli* sp. nov. and *C. loryi* leads us to  
945 propose the new *C. claveli* Zone and Subzone to replace the *C. loryi* Zone and Subzone for  
946 the south-east of France. This proposal seems preferable to avoid a revision of the *C. loryi*  
947 Zone, which would lead to confusion with its current definition. Three interval horizons are  
948 recognized: the *C. claveli* lower interval Horizon (= the first acme of the *Crioceratites*); the *C.*  
949 *claveli* upper interval Horizon (= the first part of the second acme of the *Crioceratites*, which  
950 corresponds with the LO of *C. claveli* sp. nov.); and the *C. loryi* interval Horizon (= the  
951 second part of the second acme of the *Crioceratites*, which corresponds with the FO of *C.*  
952 *loryi*). The presence of *C. claveli* sp. nov. at least in France (this work) and Italy (*Crioceratites*  
953 sp. in Faraoni et al., 1997) is a first step towards a generalization of this zonal proposal for the  
954 Mediterranean Province of the Mediterranean Caucasian Subrealm (Tethyan Realm).

955

956 (4) The northern edge of the Provençal platform (neritic zone of the Arc de Castellane)  
957 delivered some *Crioceratites* s. str. from the top part of the *A. radiatus* Zone, thus reducing  
958 the gap between the older *Crioceratites* of the uppermost Valanginian and those of the new *C.*  
959 *claveli* Zone. The specimens collected all belong to the species *C. jurensis*, which is  
960 rehabilitated and revised in this work, making it possible to refuse its synonymy with  
961 *Davouxiceras nolani* as proposed by Sarkar (1955 – see also Klein et al., 2007, p. 70): the  
962 type of *C. jurensis* is figured here for the first time since the original drawing by Pictet and  
963 Campiche (1861).

964

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975

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979

## 980 **References**

981 Agassiz, L., 1846. Nomina systematica generum molluscorum tam viventium quam fossilium. In: Jent,  
982 Gassmann (Eds.), Nomenclator zoologicus, vol. 9, 98 pp.

983 Aguirre-Urreta, B., Martinez, M., Schmitz, M., Lescano, M., Omarini, L., Tunike, M., Kuhnert, H. Concheyro,  
984 A., Rawson, P. F., Ramos, V. A., Reboulet, S., Noclin, N., Frederichs, Th., Nickl, A.-L., Pälke, H., 2019.  
985 Interhemispheric radio-astrochronological calibration of the time scales from the Andean and the Tethyan areas  
986 in the Valanginian–Hauterivian (Early Cretaceous). *Gondwana Research* 70, 104–132,  
987 <https://doi.org/10.1016/j.gr.2019.01.006>

988 Anderson, M.J., 2001. A new method for non-parametric multivariate analysis of variance. *Austral Ecology* 26,  
989 32–46.

990 Arnaud, H., 2005. The South-East France Basin (SFB) and its Mesozoic evolution. *Géologie Alpine, Série*  
991 *spéciale "Colloques et Excursions"*, 7, p. 5–28.

992 Arthur, M.A., Dean, W.E., Schlanger, S.O., 1985. Variations in the global carbon cycle during the cretaceous  
993 related to climate, volcanism, and changes in atmospheric CO<sub>2</sub>. In: Sundquist, E.T., Broecker, W.E. (Eds.), *The*  
994 *carbon cycle and atmospheric CO<sub>2</sub>: natural variations archean to present. Geophysical monograph series*, 32, pp.  
995 504–529.

- 996 Astier, J.-E., 1851. Catalogue descriptif des *Ancyloceras* appartenant à l'étage Néocomien d'Escragnolles et des  
997 Basses-Alpes. Annales des sciences physiques et naturelles d'agriculture et d'industrie 3 (2), 435–456.
- 998 Bersac, S., Bert, D., 2012. Ontogenesis, variability and evolution of the Lower Greensand Deshayesitidae  
999 (Ammonoidea, Lower Cretaceous, Southern England): reinterpretation of literature data; taxonomic and  
1000 biostratigraphic implications. Annales du Muséum d'Histoire naturelle de Nice 27, 197–270.
- 1001 Bersac, S., Bert, D., 2018. Revision of the lower Aptian (Lower Cretaceous) ammonite species *Chelonicerus*  
1002 *cornuelianum* (d'Orbigny, 1841). Annales de Paléontologie 104 (1), 45-70.  
1003 <https://doi.org/10.1016/j.annpal.2018.01.001>.
- 1004 Bert, D., 2014a. Factors of intraspecific variability in ammonites, the example of *Gassendicerus alpinum*  
1005 (d'Orbigny, 1850) (Hemihoplitidae, Upper Barremian). Annales de Paléontologie 100 (3), 217–236.
- 1006 Bert, D., 2014b. L'influence de la variabilité intraspécifique sur la taxinomie, la biostratigraphie et l'évolution des  
1007 ammonites : une approche paléobiologique - Exemples pris dans le Jurassique supérieur et le Crétacé inférieur  
1008 (Unpubl. PhD thesis). Université de Rennes 1, 736 pp.
- 1009 Bert, D., 2019. Les lois de la variabilité intraspécifique chez les ammonites. Les cahiers de la Réserve Naturelle  
1010 de Sainte-Victoire, numéro spécial congrès APF 2019, 12.
- 1011 Bert, D., Berasc, S., 2013. Evolutionary patterns – tested with cladistics – and processes in relation to  
1012 palaeoenvironments of the upper Barremian genus *Gassendicerus* (Ammonitina, Lower Cretaceous).  
1013 Palaeontology 56 (3), 631–646.
- 1014 Breistroffer, M., 1936. Révision de la faune Hauterivienne du Néron en Chartreuse (Isère). Travaux du  
1015 Laboratoire de Géologie de la Faculté des Sciences de l'Université de Grenoble 18, 131–155.
- 1016 Bruguière, J.-G., 1789. Encyclopédie méthodique : histoire naturelle des vers. Panckoucke ed., 555 pp.
- 1017 Bulot, L., 1995. Les formations à ammonites du Crétacé inférieur dans le Sud-Est de la France (Berriasien à  
1018 Hauterivien) : biostratigraphie, paléontologie et cycles sédimentaires (Unpubl. PhD thesis). Muséum National  
1019 d'Histoire Naturelle de Paris, 375 pp.
- 1020 Bulot, L., Thieuloy, J.-P., 1995. Les biohorizons du Valanginien su Sud-Est de la France: un outil fondamental  
1021 pour les corrélations au sein de la Téthys occidentale. Géologie Alpine Mémoire Hors Série 20(1994), 15–41.
- 1022 Bulot, L., Thieuloy, J.-P., Blanc, E., Klein, J., 1993. Le cadre stratigraphique du Valanginien supérieur et de  
1023 l'Hauterivien du Sud-Est de la France: définition des biochronozones et caractérisation de nouveaux biohorizons.  
1024 Géologie Alpine 68(1992), 13–56.

- 1025 Bulot, L., Thieuloy, J.-P., Arnaud, H., Delanoy, G., 1995. The Lower Cretaceous Cephalopod Team First Field  
1026 Metting (Digne, 1990). *Géologie Alpine Mémoire Hors Série* 20(1994), 383–399.
- 1027 Busnardo, R., Charollais, J.-J., Weidmann, M., Clavel, B., 2003. Le Crétacé inférieur de la Veveyse de Châtel  
1028 (Ultrasuisse des Préalpes externes; canton de Fribourg, Suisse). *Revue de Paléobiologie* 22(1), 1–174.
- 1029 Cotillon, P., 1971. Le Crétacé inférieur de l'arc subalpin de Castellane entre l'Asse et le Var – stratigraphie et  
1030 sédimentologie. *Bureau de Recherches Géologiques et Minières* 68, 313.
- 1031 Cotillon, P., Ferry, S., Gaillard, Ch., Jautée, E., Latreille, G., Rio, M., 1980. Fluctuation des paramètres du  
1032 milieu marin dans le domaine Vocontien (France Sud-Est) au Crétacé inférieur : mise en évidence par l'étude des  
1033 formations marno-calcaires alternantes. *Bulletin de la Société Géologique de France* 22(5), 735–744.
- 1034 Courville, P., Lebrun, P., 2010. L'Albien (Crétacé) de la région de Troyes (Aube) et ses ammonites: Hoplitidae  
1035 et Douvilleiceratidae. *Fossiles* 4, 6–30.
- 1036 Cuvier, G., 1798. *Tableau élémentaire de l'histoire naturelle des animaux*. Baudouin imprimeur 662, 754.
- 1037 De Baets, K., Bert, D., Hoffmann, R., Monnet, C., Yacobucci, M., Klug, Ch., 2015. Ammonoid intraspecific  
1038 variability. In Klug, Ch. et al. (eds), *Ammonoid Paleobiology : from anatomy to ecology*. *Topics in Geobiology*  
1039 43, pp. 359–426, DOI 10.1007/978-94-017-9630-9\_9.
- 1040 Dimitrova, N., 1967. Les fossiles de Bulgarie. 4. Crétacé inférieur. Cephalopoda (Nautiloidea et Ammonoidea).  
1041 Académie Bulgare des Sciences 424.
- 1042 Faraoni, P., Flore, D., Marini, A., Pallini, G., Pezzoni, N., 1997. Valanginian and early Hauterivian ammonite  
1043 successions in the Mt Catria group (Central Apennines) and in the Lessini Mts (Southern Alps), Italy.  
1044 *Palaeopelagos* 7, 59–100.
- 1045 Ferry, S., 2017. Summary on Mesozoic carbonate deposits of the Vocontian Trough (Subalpine Chains, SE  
1046 France). In Granier, B. (ed.), *Some key Lower Cretaceous sites in Drôme (SE France)*. *Carnets de Géologie*,  
1047 CG2017\_B01, pp. 9–42.
- 1048 Gill, Th., 1871. Arrangement of the families of mollusks. *Smithsonian Miscellaneous Collections* 227, 49.
- 1049 Giraud, F., Beaufort, L. and Cotillon, P., 1995. Periodicities of carbonate cycles in the Valanginian of the  
1050 Vocontian Trough: a strong obliquity control. *Geological Society, London, Special Publications* 85(1), 143–164.
- 1051 Grossouvre, A. de, 1894. *Recherches sur la Craie supérieure*. Deuxième partie : Paléontologie. Les ammonites de  
1052 la Craie supérieure. *Mémoires du service de la carte géologique détaillée de la France* (1893), 264 pp.
- 1053 Hammer, O., Harper, D. A. T., 2006. *Palaeontological data analysis*. Blackwell publishing, Oxford, 351 pp.

- 1054 Hammer, O., Harper, D. A. T., Ryan, P. D., 2001. PAST: Palaeontological Statistics Software Package for  
1055 Education and Data Analysis. *Palaeontologia Electronica* 4(1), 9 pp.
- 1056 Hoedemaeker, P.J., Reboulet, S., (reporters), Aguirre-Urreta, M.B., Alsen, P., Aoutem, M., Atrops, F., Barragán,  
1057 R., Company, M., González Arreola, C., Klein, J., Lukeneder, A., Ploch, I., Raisossadat, N., Rawson, P.F.,  
1058 Ropolo, P., Vašíček, Z., Vermeulen, J., Wippich, M.G.E., 2003. Report on the 1st International Workshop of the  
1059 IUGS Lower Cretaceous Ammonite Working Group, the “Kilian Group” (Lyon, 11 July 2002). *Cretaceous*  
1060 *Research* 24, 89–94, and erratum p. 805.
- 1061 Hyatt, A., 1899. Cephalopoda. In: Zittel, K. A. von (Ed.), (1896–1900), *Textbook of Palaeontology*. Transl.  
1062 Eastman, C.R. Macmillan, pp. 502–604.
- 1063 Kilian, W., 1896. Notice stratigraphique sur les environs de la région de Sisteron et contributions à la  
1064 connaissance des terrains secondaires du Sud-Est de la France. *Bulletin de la Société Géologique de France* 23,  
1065 211–227.
- 1066 Kilian, W., 1910. Un nouvel exemple de phénomène de convergence chez des ammonitidés. Sur les origines du  
1067 groupe de l' « *Ammonite bicurvatus* » Mich. (sous-genre *Saynella* Kil.). Extrait des Comptes rendus des séances  
1068 de l'Académie des Sciences 150, 150–153.
- 1069 Kilian, W., Zürcher, Ph., avec la collaboration d'Adrien Guebhard, 1896. Notice sur la région d'Escragnolles  
1070 (Alpes Maritimes). *Bulletin de la Société Géologique de France* 3(23), 952–969.
- 1071 Klein, J., Busnardo, R., Company, M., Delanoy, G., Kakabadze, M.V., Reboulet, S., Ropolo, P., Vašíček, Z.,  
1072 Vermeulen, J., 2007. Lower Cretaceous Ammonites III. Bochianitoidea, Protancyloceratoidea, Ancyloceratoidea,  
1073 Ptychoceratoidea. *Fossilium Catalogus I : Animalia*, vol. 144. Backhuys Publishers, p. 381.
- 1074 Lévillé, Ch, 1837. Description de quelques nouvelles coquilles fossiles du Département des Basses-Alpes.  
1075 *Mémoires de la Société Géologique de France* 2, 313–315.
- 1076 Mandov, G. K., 1976. L'étage Hauterivien dans les Balkanides occidentales (Bulgarie de l'ouest) et sa faune  
1077 d'ammonites. *Annuaire de l'Université de Sofia, Livre 1, Géologie* 67, 11–99.
- 1078 Marchand, D., 1986. L'évolution des Cardioceratinae d'Europe occidentale dans leur contexte  
1079 paléobiogéographique (Callovien supérieur-Oxfordien moyen) (Unpubl. PhD thesis). Université de Dijon,  
1080 601 pp.
- 1081 Matamales-Andreu, R., Company, M., 2019. Morphological variability patterns in the *Balearites-*  
1082 *Pseudothurmannia* genera boundary (Ammonoidea, late Hauterivian): taxonomic and biostratigraphic  
1083 implications. *Journal of Systematic Palaeontology* 17 (13), 869–895.

- 1084 Mayer-Eymar, K., 1887. Systematisches Verzeichniss der Kreide- und Tertiär-Versteinerungen der Umgegend  
1085 von Thun, nebst Beschreibung der neuen Arten. Beiträge zur geologischen Karte der Schweiz, Beilage 24(2),  
1086 i–xxviii + 1–128.
- 1087 Mutterlose, J., Rawson, P., Reboulet, S., Baudin, F., Bulot, L., Emmanuel, L., Gardin, S., Martinez, M., Renard,  
1088 M., 2020. The Global Boundary Stratotype Section and Point (GSSP) for the base of the Hauterivian Stage  
1089 (Lower Cretaceous), La Charce, southeast France. Communication of IUGS Geological Standards, 22pp.  
1090 <https://doi.org/10.18814/epiiugs/2020/020072>.
- 1091 Nolan, H., 1894. Note sur les *Crioceras* du groupe de *Crioceras Duvali*. Bulletin de la Société Géologique de  
1092 France 3(22), 183–196.
- 1093 Orbigny, A. d', 1840-1842. Paléontologie française - Terrains crétacés - Tome 1, Céphalopodes (liv. 16).  
1094 Masson, pp. 121–430.
- 1095 Paquier, V. L., 1900. Recherches géologiques dans le Diois et les Barronies orientales. Bulletin de la Société de  
1096 Statistique des Sciences Naturelles et des Arts industriels du Département de l'Isère 5, 77–476.
- 1097 Parona, C.F., Bonarelli, G., 1897. Fossili albiani d'Escagnolles, del Nizzardo e della Liguria occidentale.  
1098 Paleontographia Italica 2, 53–112.
- 1099 Pictet, F., Campiche, G., 1861. Description des fossiles du terrain Crétacé des environs de Sainte-Croix, part 2.  
1100 Matériaux pour la Paléontologie Suisse 3, 1–144.
- 1101 Quinn, G.P., Keough M.J., 2002. Experimental design and data analysis for biologists. Cambridge University  
1102 Press, 537 pp.
- 1103 Reboulet, S., 1996. L'évolution des ammonites du Valanginien-Hauterivien inférieur du bassin Vocontian et de  
1104 la plate-forme provençale (Sud-Est de la France): relations avec la stratigraphie séquentielle et implications  
1105 biostratigraphique. Documents des Laboratoires de Géologie de Lyon, 137, 371 pp.
- 1106 Reboulet, S., 2015. Le Valanginien-Hauterivien inférieur du SE de la France. 2015, Livret-Guide des Excursions  
1107 du Groupe Français du Crétacé, 74 pp.
- 1108 Reboulet, S., Atrops, F., 1999. Comments and proposals about the Valanginian-Lower Hauterivian ammonite  
1109 zonation of south-eastern France. Eclogae geologicae Helvetiae 92, 183–197.
- 1110 Reboulet, S., Atrops, F., Ferry, S., Schaaf, A., 1992. Renouvellement des ammonites en fosse Vocontienne à la  
1111 limite Valanginien-Hauterivien. Geobios 25(4), 469–476.

- 1112 Reboulet, S., Mattioli, E., Pittet, B., Baudin, F., Olivero, D., Proux, O., 2003. Ammonoid and nannoplankton  
1113 abundance in Valanginian (early Cretaceous) limestone-marl successions from the southeast France Basin:  
1114 carbonate dilution or productivity? *Palaeogeography, Palaeoclimatology, Palaeoecology* 201, 113–139.
- 1115 Reboulet, S., Giraud, F., Proux, O., 2005. Ammonoid abundance variations related to changes in trophic  
1116 conditions across the Oceanic Anoxic Event 1d (Latest Albian, SE France). *Palaios* 20, 121–141.
- 1117 Reboulet, S., Hoedemaeker, Ph. J., Aguirre-Urreta, M., B., Alsen, P., Atrops, F., Baraboshkin, E., Y., Company,  
1118 M., Delanoy, G., Dutour, Y., Klein, J., Latil, J.-L., Lukeneder, A., Mitta, V., Mourgues, F. A., Ploch, I.,  
1119 Raisossadat, N., Ropolo, P., Sandoval, J., Tavera, J. M., Vašíček, Z., Vermeulen, J., Arnaud, H., Granier, B.,  
1120 Premoli-Silva, I., 2006. Report on the 2nd international meeting of the IUGS lower Cretaceous ammonite  
1121 working group, the “Kilian Group” (Neuchâtel, Switzerland, 8 September 2005). *Cretaceous Research* 27,  
1122 712–715.
- 1123 Reboulet, S., Klein, J., (reporters), Barragán, R., Company, M., González-Arreola, C., Lukeneder, A.,  
1124 Raisossadat, S.N., Sandoval, J., Szives, O., Tavera, J.M., Vašíček, Z., Vermeulen, J., 2009. Report on the 3rd  
1125 International Meeting of the IUGS Lower Cretaceous Ammonite Working Group, the “Kilian Group” (Vienna,  
1126 Austria, 15th April 2008). *Cretaceous Research* 30, 496–502.
- 1127 Reboulet, S., Rawson, P.F., Moreno-Bedmar, J.A., (reporters), Aguirre-Urreta, M.B., Barragán, R., Bogomolov,  
1128 Y., Company, M., González-Arreola, C., Idakieva Stoyanova, V., Lukeneder, A., Matrimon, B., Mitta, V.,  
1129 Randrianaly, H., Vašíček, Z., Baraboshkin, E.J., Bert, D., Bersac, S., Bogdanova, T.N., Bulot, L.G., Latil, J.-L.,  
1130 Mikhailova, I.A., Ropolo, P., Szives, O., 2011. Report on the 4th International Meeting of the IUGS Lower  
1131 Cretaceous Ammonite Working Group, the “Kilian Group” (Dijon, France, 30th August 2010). *Cretaceous*  
1132 *Research* 32, 786–793.
- 1133 Reboulet, S., Szives, O., (reporters), Aguirre-Urreta, B., Barragán, R., Company, M., Idakieva, V., Ivanov, M.,  
1134 Kakabadze, M.V., Moreno-Bedmar, J.A., Sandoval, J., Baraboshkin, E.J., Çağlar, M.K., Fözy, I., González-  
1135 Arreola, C., Kenjo, S., Lukeneder, A., Raisossadat, S.N., Rawson, P.F., Tavera, J.M., 2014. Report of the 5th  
1136 International Meeting of the IUGS Lower Cretaceous Ammonite Working Group, the Kilian Group (Ankara,  
1137 Turkey, 31st August 2013). *Cretaceous Research* 50, 126–137.
- 1138 Reboulet, S., Szives, O., Aguirre-Urreta, B., Barragán, R., Company, M., Frau, C., Kakabadze, M.V., Klein, J.,  
1139 Moreno-Bedmar, J.A., Lukeneder, A., Pictet, A., Ploch, I., Raisossadat, S.N., Vašíček, Z., Baraboshkin, E.J.,  
1140 Mitta, V.V., 2018. Report on the 6th International Meeting of the IUGS Lower Cretaceous Ammonite Working  
1141 Group, the Kilian Group (Vienna, Austria, 20th August 2017). *Cretaceous Research* 91 (4), 100–110.



- 1142 Rodighiero, A., 1919. Il sistema Cretaceo del Veneto occidentale compreso fra l'Adige e il Piave con speciale  
1143 riguardo al Neocomiano dei Setti Comuni. *Palaeontographica Italica* 25, 70-125.
- 1144 Ropolo, P., 1992. *Crioceratites curnieri* nov. sp. une nouvelle espèce d'ammonite hétéromorphe préfigurant  
1145 l'acquisition des coquilles triparties de l'Hauterivien inférieur (Ammonoidea, Ancyloceratina). *Mésogée* 51,  
1146 65–73.
- 1147 Ropolo, P., 1995. Implications of variation in coiling in some Hauterivian (Lower Cretaceous) heteromorph  
1148 ammonites from the Vocontian basin, France. *Memorie Descrittive della Carta Geologica d'Italia* 51, 137–165.
- 1149 Salvador, A., 1994. International stratigraphic guide: a guide to stratigraphic classification, terminology, and  
1150 procedure, second ed. The Geological Society of America, Boulder, Colorado, p. 214.
- 1151 Sarkar, S., 1954. Sur un genre nouveau d'ammonite déroulée. *Compte Rendu Sommaire des Séances de la*  
1152 *Société géologique de France*, 97–98.
- 1153 Sarkar, S.S., 1955. Révision des ammonites déroulées du Crétacé inférieur du Sud-Est de la France. *Memoires*  
1154 *de la Société Géologique de France* 34 (72), 1–176.
- 1155 Schlanger, S. O., Jenkyns, H. C., 1976. Cretaceous oceanic anoxic events: causes and consequences. *Geologie en*  
1156 *Mijnbouw* 55(3-4), 179–184.
- 1157 Spath, L.F., 1923. In: *Monograph of the Ammonoidea of the Gault*, Vol. 1. Palaeontographical Society, pp. 1–72.
- 1158 Steinmann, G., 1890. In: Steinmann, G., Doederlein, L. (Eds.), *Elemente der Paläontologie*, p. 848. Leipzig.
- 1159 Thieuloy, J.-P., 1972. Biostratigraphie des lentilles à Pérégrinelles (Brachiopodes) de l'Hauterivien de Rotier  
1160 (Drôme, France). *Geobios*, 5(1), 5–53.
- 1161 Thieuloy, J.-P., 1977. La zone à *callidiscus* du Valanginien supérieur Vocontian (Sud-Est de la France).  
1162 Lithostratigraphie, ammonitologie, limite Valanginien-Hauterivien, corrélations. *Géologie Alpine* 53, 83–143.
- 1163 Thieuloy, J.-P., Bulot, L.G., 1993. Ammonites du Crétacé inférieur du Sud-Est de la France: 1. Nouvelles  
1164 espèces à valeur stratigraphique pour le Valanginien et l'Hauterivien. *Géologie Alpine* 68(1992), 85–103.
- 1165 Turner, J., 1962. Quelques remarques sur les genres *Bochianites* d'Orbigny et *Baculina* d'Orbigny. *Travaux du*  
1166 *Laboratoire de Géologie de la Faculté des Sciences de Grenoble* 38 : 241–248.
- 1167 Uhlig, V., 1882. Zur Kenntnis der Cephalopoden der Rossfeldschichten. *Jahrbuch der Kaiserlich-Königlichen*  
1168 *Geologischen Reichsanstalt* 32, 373–396.
- 1169 Uhlig, V., 1905. Einige Bemerkungen über die Ammonitengattung *Hoplites* Neumayr. *Sitzungsberichte der*  
1170 *Keiserlichen Akademie der Wissenschaften in Wien, mathematisch-naturwissenschaftliche Klasse* 114, 591–636.

- 1171 Vašíček, Z., 1995. Lower Cretaceous ammonite biostratigraphy in the Western Carpathians (the Czech and  
 1172 Slovak Republics). *Géologie Alpine Mémoire Hors Série* 20(1994), 169–189.
- 1173 Vašíček, Z., 2005. The oldest (Late Valanginian) Crioceratitinae (heteromorphic ammonoids) from the Central  
 1174 Western Carpathians (Slovakia). *Geologica Carpathica* 56, 245–254.
- 1175 Vermeulen, J., 2004. Vers une nouvelle classification à fondement phylogénétique des ammonites  
 1176 hétéromorphes du Crétacé inférieur méditerranéen. Le cas des Crioceratitidae Gill, 1871, nom. correct. Wright,  
 1177 1952, des Emericiceratidae fam. nov. et des Acrioceratidae fam. nov. (Ancylocerataceae Gill, 1871). *Riviéra*  
 1178 *Scientifique* 88, 69–92.
- 1179 Vernet, B., 2017. Etude de la coupe du Ravin du Cheiron (Unpubl. M2 thesis). Université de Rennes 1, 37 pp.
- 1180 Westermann, G.E.G., 1966. Covariation and taxonomy of the Jurassic ammonite *Sonninia adicra* (Waagen).  
 1181 *Neues Jahrbuch für Geologie und Paläontologie Abteilung* 124(3), 289–312.
- 1182 Wiedmann, J., 1966. Stammesgeschichte und System der posttriadischen Ammonoideen. 2. Teil. *Neues Jahrbuch*  
 1183 *für Geologie und Paläontologie, Abhandlungen* 127, 13–81.

1184

#### 1185 **Captions of the figures:**

1186 **Fig. 1.** Palaeogeographic map of the Vocontian Basin for the Hauterivian (southeastern France, modified after  
 1187 Arnaud, 2005). The stars point out the sections mentioned in the text: CAR = Rougon; CHE = the Cheiron  
 1188 ravine; COU1 = Bargème; ESC = Escragnolles; LCH = La Charce GSSP; POM = Pommerol; PV2 = La Palud-  
 1189 sur-Verdon.

1190 **Fig. 2.** Lithology of the Cheiron ravine (CHE) with ammonite fauna and biostratigraphy (this work). The interval  
 1191 horizons highlighted in this section are: CC1 = *Crioceratites claveli* lower interval Horizon; CC2 = *Crioceratites*  
 1192 *claveli* upper interval Horizon; CL = *Crioceratites loryi* interval Horizon.

1193 **Fig. 3.** *Crioceratites jurensis* (Nolan, 1894), cast of the type specimen MHNG GEPI 16787, section ESC,  
 1194 *Acanthodiscus radiatus* Zone. The arrow points out the end of the phragmocone.

1195 **Fig. 4.** *Crioceratites jurensis* (Nolan, 1894), specimen DBT.BC47, section CAR, upper part of the  
 1196 *Acanthodiscus radiatus* Zone; c is face B of the body chamber of the same specimen. The arrow points out the  
 1197 end of the phragmocone.

1198 **Fig. 5.** *Crioceratites jurensis* (Nolan, 1894), specimen LCT.7Z, section PV2, upper part of the *Acanthodiscus*  
 1199 *radiatus* Zone. The arrow points out the end of the phragmocone.

1200 **Fig. 6.** *Crioceratites loryi* (Sarkar, 1955); a-b: macroconch specimen DBT.AV98 from the Moriez area; c-d:  
1201 microconch specimen DBT.BB50 from the Moriez area; e: microconch specimen DBT.BB84 from CHE, bed  
1202 197. All the specimens are from the *Crioceratites loryi* interval Horizon in the new *Crioceratites claveli* Zone  
1203 and Subzone. The arrows point out the end of the phragmocone.

1204 **Fig. 7.** *Crioceratites claveli* sp. nov.; a-b: macroconch specimen DBT.BB48, holotype; c-d: microconch  
1205 specimen DBT.BB49; e: microconch specimen DBT.BB65; f: microconch specimen DBT.BB60; g: microconch  
1206 specimen DBT.BB57. All the specimens are from CHE, bed 179, in the *C. claveli* lower interval Horizon, new  
1207 *C. claveli* Zone and Subzone. The arrows point out the end of the phragmocone.

1208 **Fig. 8.** *Crioceratites claveli* sp. nov., macroconch specimen DBT.BB52, from CHE, bed 179, in the *C. claveli*  
1209 lower interval Horizon, new *C. claveli* Zone and Subzone. The arrow points out the end of the phragmocone.

1210 **Fig. 9.** *Crioceratites claveli* sp. nov.; a: macroconch specimen DBT.BB55; b-c: microconch specimen  
1211 DBT.BB54, note the nodular periventral tubercles of the ontogenic stage 2. All the specimens are from CHE,  
1212 bed 179, in the *C. claveli* lower interval Horizon, new *C. claveli* Zone and Subzone. The arrows point out the end  
1213 of the phragmocone.

1214 **Fig. 10.** *Crioceratites claveli* sp. nov., a: microconch specimen DBT.BG50; b: microconch specimen  
1215 DBT.BG51; c: macroconch specimen DBT.BG75. All the specimens are from CHE, bed 195, in the *C. claveli*  
1216 upper interval Horizon, new *C. claveli* Zone and Subzone. The arrows point out the end of the phragmocone.

1217 **Fig. 11.** *Crioceratites claveli* sp. nov., macroconch specimen DBT.BG49, from CHE, bed 194, in the *C. claveli*  
1218 upper interval Horizon, new *C. claveli* Zone and Subzone. The arrow points out the end of the phragmocone.

1219 **Fig. 12.** *Crioceratites claveli* sp. nov.; a-b: microconch specimen POM 219.2 from POM, bed 219, b is the  
1220 mould of the imprint; c: microconch specimen FSL 487733a; d: microconch specimen FSL 487743a; e:  
1221 microconch specimen FSL 487733e; f: microconch specimen FSL 487733f; g: undetermined antidimorph  
1222 specimen FSL 487733j. Specimens c-f are from LCH, in bed 219; specimen g is from bed 220 of the same  
1223 section. All the specimens are from the *C. claveli* lower interval Horizon, new *C. claveli* Zone and Subzone. The  
1224 arrows point out the end of the phragmocone.

1225 **Fig. 13.** Diagrams of the univariate analysis. Normality curve in black, Kernel density in red (refer to the  
1226 electronic version of the paper).

1227 **Fig. 14.** Bivariate analysis crossplots. Red dots: microconchs; blue dots: macroconchs; black dots: undetermined  
1228 antidimorphs (refer to the electronic version of the paper).

1229 **Fig. 15.** Diagrams of the multivariate analysis: principal component analysis crossplot, 'loading' of the variables  
1230 on the first two Principal Components, frequency histograms of the projections of the plots on the first two PC  
1231 and matrix of the Spearman's rank correlation. Red dots: microconchs; blue dots: macroconchs; black dots:  
1232 undetermined antidimorphs; normality curve in black, kernel density in red (refer to the electronic version of the  
1233 paper).

1234

### 1235 **Caption of the tables:**

1236 **Table 1.** Measurements of the main specimens of *Crioceratites jurensis*. In bold are the diameters expected  
1237 adult.

1238 **Table 2.** Measurements of the main specimens of *Crioceratites claveli* sp. nov. In bold are the expected adult  
1239 diameters.

1240 **Table 3.** Normality tests of the multivariate analysis.

1241 **Table 4.** Variance and eigenvalue reached by the principal components.

1242 **Table 5.** Loading of the variables on the 7 PC.

1243 **Table 6.** Correlation by Spearman rank. In grey: dependent variables ( $p < 0.05$ ).

1244 **Table 7.** The ANOSIM.

1245 **Table 8.** p value of the sequential Bonferoni significance: in grey are the significant differences.

1246 **Table 9.** Biostratigraphic chart.

No. specimen	D	H	W	U	h	Nt/2	Ni/2	H/D	W/D	U/D	W/H	U/H	Alpha
MHNG GEPI 16787 – Fig. 3	139.78	33.20	35.50	79.62	c16.00	-	-	0.24	0.25	0.57	1.07	2.40	170°
	88.00	24.70	19.34	44.58	5.22	8	30	0.28	0.22	0.51	0.78	1.80	
DBT.04171-CAR.BC47 – Fig. 4	116.50	32.52	26.48	67.18	9.56	6	26	0.28	0.23	0.58	0.81	2.07	120°
	ph89.48	26.38	23.68	47.54	4.24	-	-	0.30	0.26	0.53	0.90	1.80	
LCT.04144-PV2.7Z – Fig. 5	<b>155.34</b>	<b>42.10</b>	<b>39.61</b>	<b>84.64</b>	<b>c13.00</b>	<b>7</b>	<b>38</b>	<b>0.27</b>	<b>0.26</b>	<b>0.55</b>	<b>0.94</b>	<b>2.01</b>	140° 260°
	ph109.00	33.02	32.14	59.40	6.52	-	-	0.30	0.30	0.55	0.97	1.80	
	54.30	18.00	-	23.14	2.42	10	-	-	0.33	-	0.43	-	
DBT.04144-PV2.BG31	<b>164.00</b>	<b>43.00</b>	<b>35.00</b>	<b>88.76</b>	<b>9.26</b>	-	-	<b>0.26</b>	<b>0.21</b>	<b>0.54</b>	<b>0.81</b>	<b>2.04</b>	

No. specimen	Dimorph	D	H	W	U	h	Nt/2	Ni/2	H/D	W/D	U/D	W/H	U/H	Alpha
DBT.04039-CHE/179.BB48 – Fig. 7a-b	[M]	<b>115.00</b>	<b>30.84</b>	<b>19.90</b>	<b>58.20</b>	<b>5.30</b>	<b>6</b>	<b>30</b>	<b>0.27</b>	<b>0.17</b>	<b>0.51</b>	<b>0.65</b>	<b>1.89</b>	
		98.00	27.54		50.60	6.40			0.28		0.52		1.84	90°
		77.82	25.34		35.60	1.60			0.33		0.46		1.40	90°
		64.20	19.38		28.68	1.30	6	32	0.30		0.45		1.48	90°
		50.52	17.08		22.00	0.90	7	35	0.34		0.44		1.29	90°
DBT.04039-CHE/179.BB49 – Fig. 7c-d	[m]	<b>78.30</b>	<b>24.00</b>	<b>16.18</b>	<b>38.50</b>	<b>3.20</b>	<b>6</b>	<b>32</b>	<b>0.31</b>	<b>0.21</b>	<b>0.49</b>	<b>0.67</b>	<b>1.60</b>	
		60.00	20.84	12.76	28.20	0.00			0.35	0.21	0.47	0.61	1.35	90°
DBT.04039-CHE/179.BB51	[m]	43.88	14.22		19.42	1.10	0	37	0.32		0.44		1.37	
DBT.04039-CHE/179.BB52 – Fig. 8	[M]	<b>140.20</b>	<b>41.30</b>		<b>68.00</b>	<b>6.14</b>	<b>6</b>	<b>38</b>	<b>0.29</b>		<b>0.49</b>		<b>1.65</b>	
DBT.04039-CHE/179.BB53	[m]	42.98	15.56		18.68	1.80	9	28	0.36		0.43		1.20	
		34.72	13.88		13.44	0.60	9	30	0.40		0.39		0.97	90°
DBT.04039-CHE/179.BB54 – Fig. 09b-c	[m]	56.84	19.40		26.76	1.98	10	28	0.34		0.47		1.38	
DBT.04039-CHE/179.BB55a – Fig. 09a	[M]	75.40	26.60		31.36	2.22	9	40	0.35		0.42		1.18	
DBT.04039-CHE/179.BB56	[M]	111.74	34.30	20.70	57.64	6.14			0.31	0.19	0.52	0.60	1.68	
DBT.04039-CHE/179.BB57 – Fig. 7g	[m]	47.24	16.66		18.98	1.18	9	51	0.35		0.40		1.14	
DBT.04039-CHE/179.BB58	[M]	47.18	15.38		19.42	2.42			0.33		0.41		1.26	
DBT.04039-CHE/179.BB59	[m]	48.94	15.60		22.24	1.66	9	39	0.32		0.45		1.43	
DBT.04039-CHE/179.BB60 – Fig. 7f	[m]	66.00	23.30		29.30	2.35			0.35		0.44		1.26	
DBT.04039-CHE/179.BB64	?	24.30	8.94		9.22		7	44	0.37	0.00	0.38		1.03	
DBT.04039-CHE/179.BB65 – Fig. 7e	[m]	38.34	13.58	8.70	17.00	1.40	7	47	0.35	0.23	0.44	0.64	1.25	
DBT.04039-CHE/194.BG49 – Fig. 11	[M]	<b>159.00</b>	<b>44.32</b>		<b>87.76</b>				<b>0.29</b>		<b>0.55</b>		<b>1.98</b>	
DBT.04039-CHE/195.BG50 – Fig. 10a	[m]	<b>78.38</b>	<b>23.30</b>		<b>37.00</b>	<b>3.50</b>	<b>9</b>	<b>24</b>	<b>0.30</b>		<b>0.47</b>		<b>1.59</b>	
DBT.04039-CHE/195.BG51 – Fig. 10b	[m]	<b>85.54</b>	<b>25.62</b>		<b>40.06</b>	<b>3.84</b>			<b>0.30</b>		<b>0.49</b>		<b>1.56</b>	
DBT.04039-CHE/195.BG58	[M]	133.16	39.50		66.44	5.44	8	34	0.30		0.50		1.68	
DBT.04039-CHE/195.BG59	[m]	61.12	19.32		27.60	3.06	9	29	0.32		0.45		1.43	
DBT.04039-CHE/195.BG74	[m]	62.38	19.62		28.82	2.5	12	21	0.32		0.46		1.47	
POM 219.8	[M]	52.90	17.52		22.60	1.60	10	33	0.33		0.43		1.29	
POM 219.5	[m]	41.22	13.90		17.12	1.62			0.34		0.42		1.23	
POM 219.2 – Fig. 12a-b	[m]	<b>86.00</b>	<b>26.00</b>		<b>44.00</b>	<b>2.60</b>			<b>0.30</b>		<b>0.51</b>		<b>1.69</b>	
POM 219.12	[M]	58.70	18.00		28.54	1.70	7	32	0.31		0.49		1.59	
FSL 487736 c	?	27.36	10.92	5.30	10.10		6	36	0.40	0.19	0.37	0.49	0.92	
FSL 487736 f	?	29.40	11.30		12.40		11		0.38		0.42		1.10	
FSL 487736 g	[m]	63.48	20.54	12.20	27.14	2.70	8	29	0.32	0.19	0.43	0.59	1.32	
FSL 487736 h	[m]	40.10	12.30		17.86	3.00			0.31		0.45		1.45	
FSL 487736 i	[M]	57.90	18.78			1.54			0.32		0.00		0.00	
FSL 487733 a – Fig. 12c	[m]	54.08	16.88		25.06	2.42			0.31		0.46		1.48	
FSL 487733 b	[m]	46.30	15.20		22.34	2.20			0.33		0.48		1.47	
FSL 487733 e – Fig. 12e	[m]	40.16	13.30				9	46	0.33		0.00		0.00	
FSL 487733 f – Fig. 12f	[m]	32.20	11.40				11	30	0.35		0.00		0.00	
FSL 487733 j – Fig. 12g	?	20.18	7.48		8.62		6	30	0.37		0.43		1.15	
FSL 487733 l	?	37.10	13.50		14.30				0.36		0.39		1.06	
FSL 487738	[m]	48.80	17.00		20.80				0.35		0.43		1.22	
FSL 487743 a – Fig. 12d	[m]	55.66	19.04	8.10	26.86	2.80	11	36	0.34	0.15	0.48	0.43	1.41	

Normality tests	D	h	Nt/2	Ni/2	H/D	U/D	U/H
N	36	29	20	20	36	35	35
Shapiro-Wilk W	0.9861	0.8557	0.8731	0.938	0.983	0.9715	0.9898
p(normal)	0.9228	0.0009995	0.01335	0.2194	0.8415	0.4855	0.9828

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PC	Eigenvalue	% variance
1	4.05054	57.865
2	1.00108	14.301
3	0.948792	13.554
4	0.571191	8.1599
5	0.266392	3.8056
6	0.159531	2.279
7	0.00247875	0.035411

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	PC 1	PC 2	PC 3	PC 4	PC 5	PC 6	PC 7
D	0.42444	0.0075554	0.15133	-0.35641	0.80451	0.14899	0.020707
h	0.35504	-0.072193	0.11709	0.89564	0.20663	-0.10065	0.0051982
Nt/2	-0.13147	0.62224	0.75266	-0.0030951	-0.049202	-0.16308	-0.0012766
Ni/2	-0.10706	-0.76078	0.62253	-0.067397	-0.099499	0.08799	-0.0049681
H/D	-0.4626	0.1036	0.0042742	0.22423	0.19762	0.72089	0.40773
U/D	0.46358	0.13295	0.084777	-0.021826	-0.37895	0.64311	-0.45008
U/H	0.16806	0.22109	-0.1251	0.71473	0.017449	0.017449	0.56262

	h	Nt/2	Ni/2	H/D	U/D	U/H
D	rs=0.61082 p=0.00043261	rs=-0.18903 p=0.42476	rs=-0.14793 p=0.53369	rs=-0.73707 p=2.9521E-07	rs=0.75062 p=2.0688E-07	rs=0.77383 p=4.9882E-08
h		rs=-0.20082 p=0.4558	rs=-0.18942 p=0.46652	rs=-0.60044 p=0.0005739	rs=-0.6745 p=8.2778E-05	rs=0.68515 p=5.7536E-05
Nt/2			rs=0.0099851 p=0.96764	rs=0.31352 p=0.17829	rs=-0.24228 p=0.30339	rs=-0.27324 p=0.24377
Ni/2				rs=0.11195 p=0.63841	rs=-0.21551 p=0.3615	rs=-0.19781 p=0.40316
H/D					rs=-0.79611 p=1.0787E-08	rs=-0.94336 p=2.2428E-17
U/D						rs=0.94638 p=9.2953E-18

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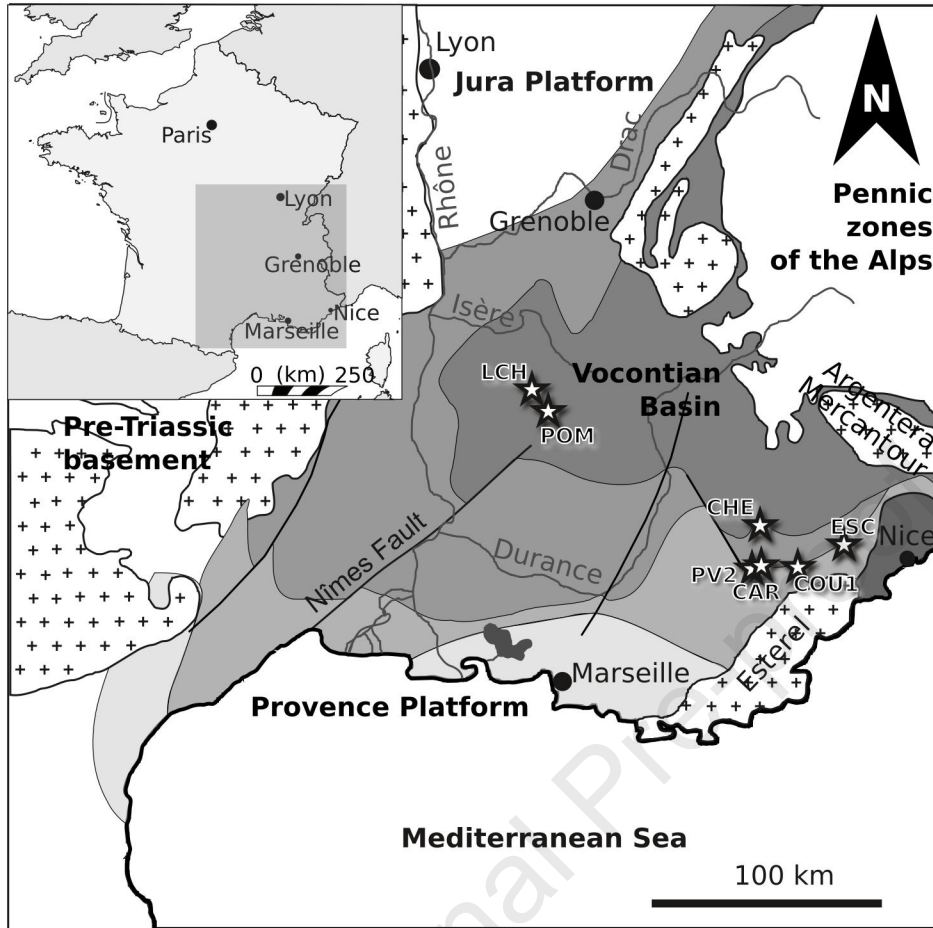
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Mean rank between:	329.9
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$p$ (same):	0.07049






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Dimorphs	Macro	micro	indet
Macroconchs		0.00529	1
microconchs	0.00529		0.4511
indeterminate antidimorphs	1	0.4511	

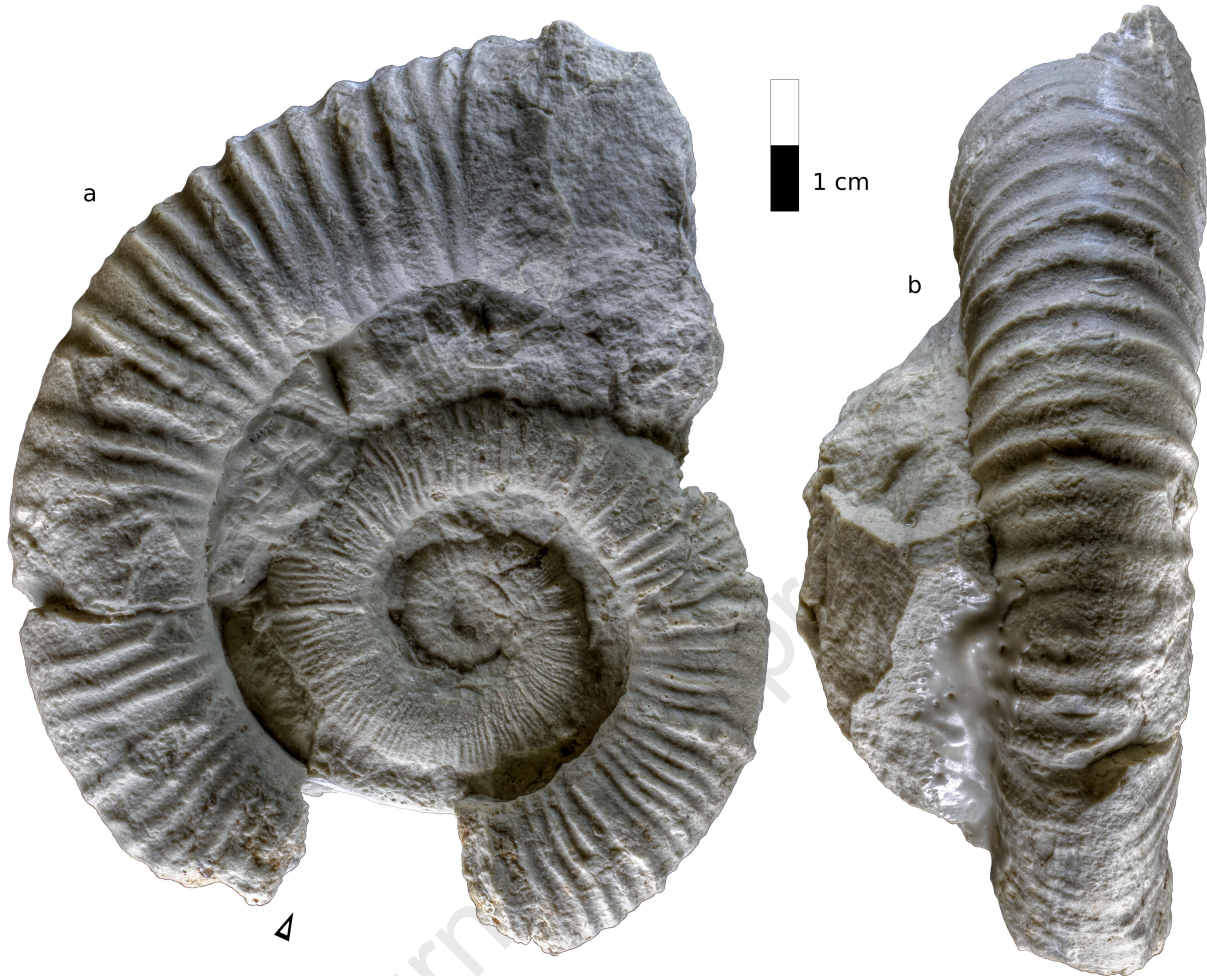
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Standard zonation (from <a href="#">Reboulet et al., 2018</a> )				Zonation proposed in the present work (in grey: not studied)		
Stage	Zone	Subzone	Interval Horizon	Zone	Subzone	Interval Horizon
lower Hauterivian	<i>Lyticoceras nodosoplicatum</i>			<i>Lyticoceras nodosoplicatum</i>		
			<i>O. variegatus</i>			<i>O. variegatus</i>
	<i>Crioceratites loryi</i>	<i>Olcostephanus jeannoti</i>		<i>Crioceratites claveli</i> (new)	<i>Olcostephanus jeannoti</i>	
		<i>Crioceratites loryi</i>			<i>C. claveli</i> (new)	<i>C. loryi</i>
						<i>C. claveli</i> upper
				<i>C. claveli</i> lower		
	<i>Acanthodiscus radiatus</i>			<i>Acanthodiscus radiatus</i>		
		<i>Breistrofferella castellanensis</i>			<i>Breistrofferella castellanensis</i>	



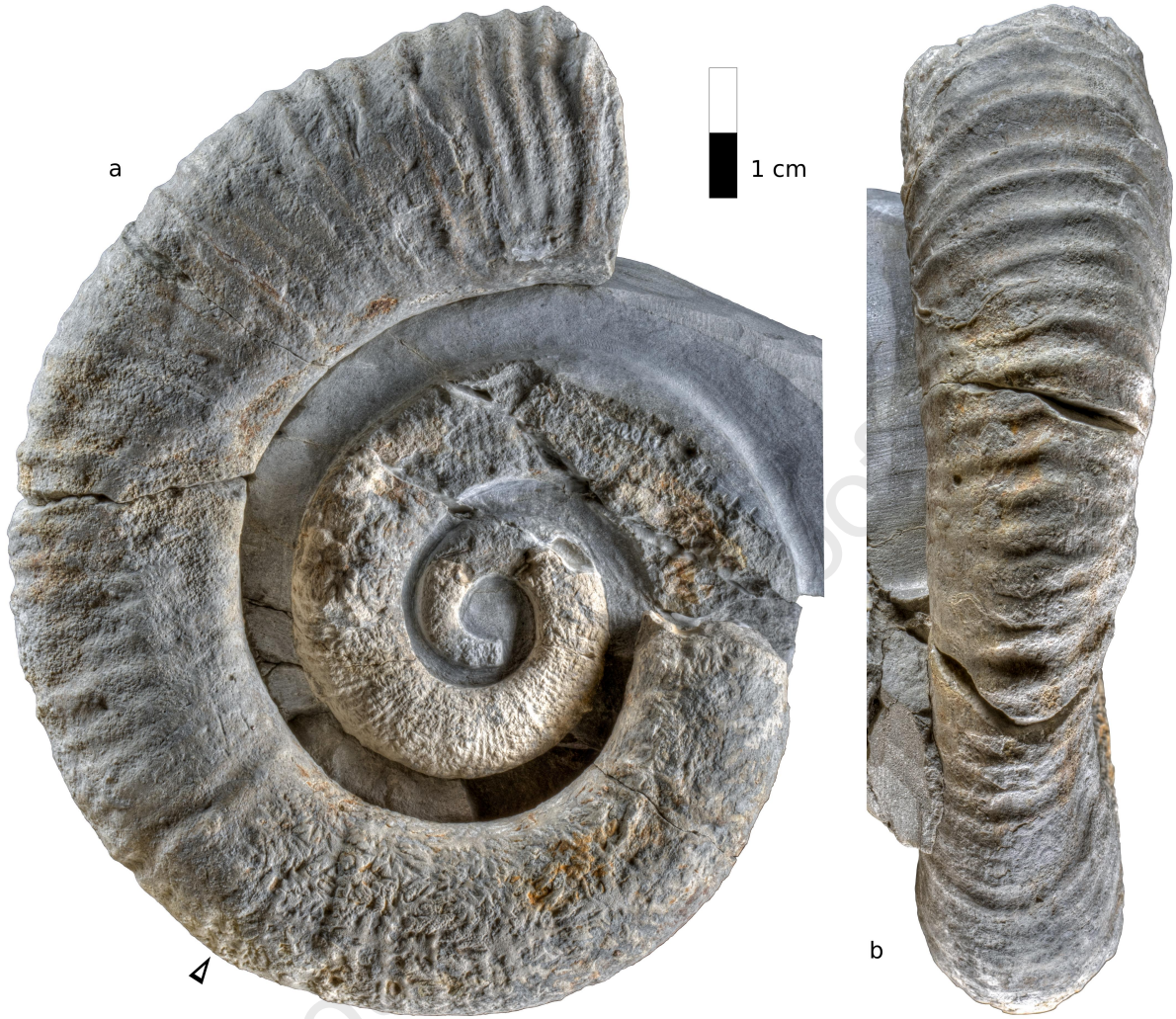
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|---|--|
|  Deep open marine environment (pelagic)        |  Drowned platform (bioclastic facies) |
|  Shallow open marine environment (hemipelagic) |  Drowned platform (condensed series)  |
|  Drowned platform (marly limestones)           |  |









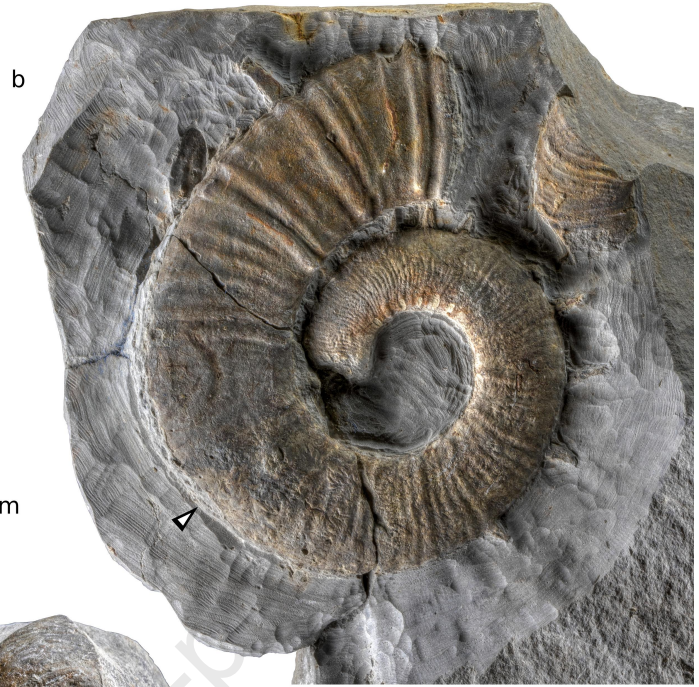
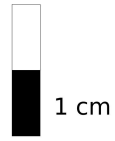
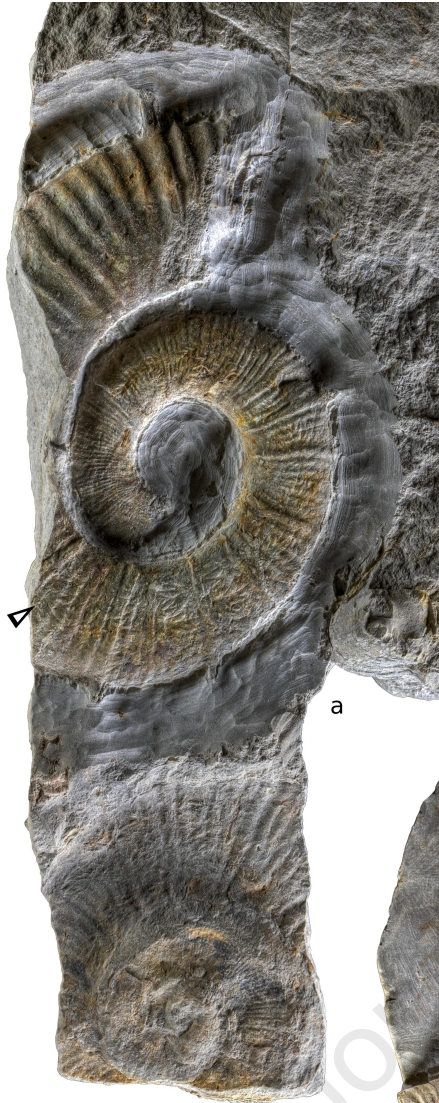






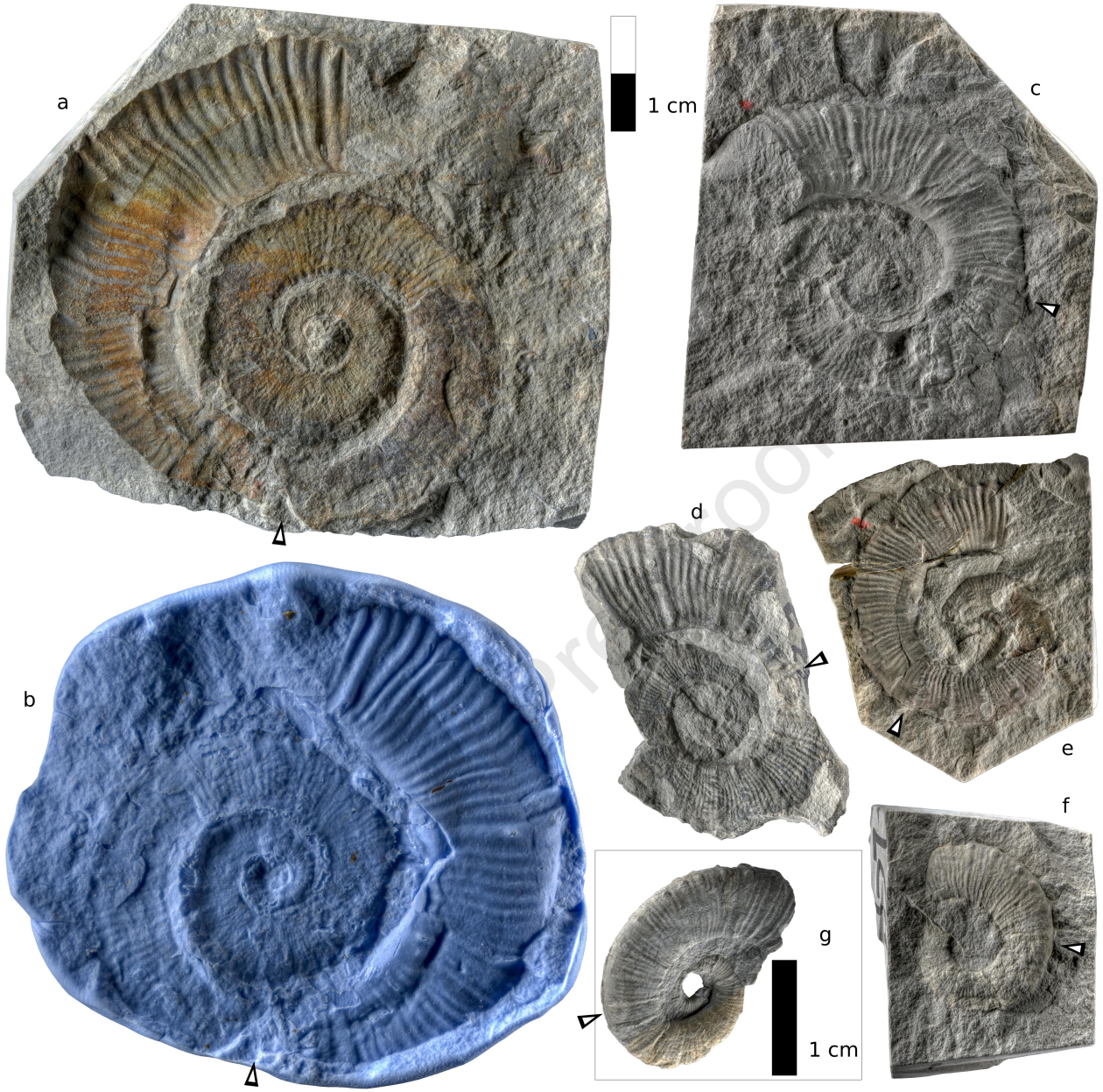


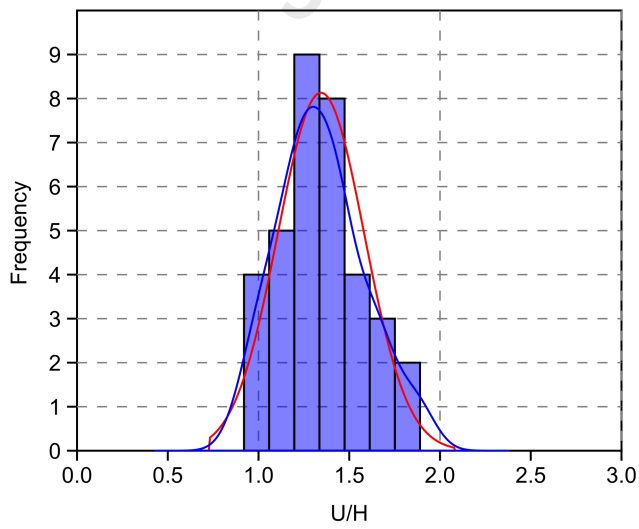
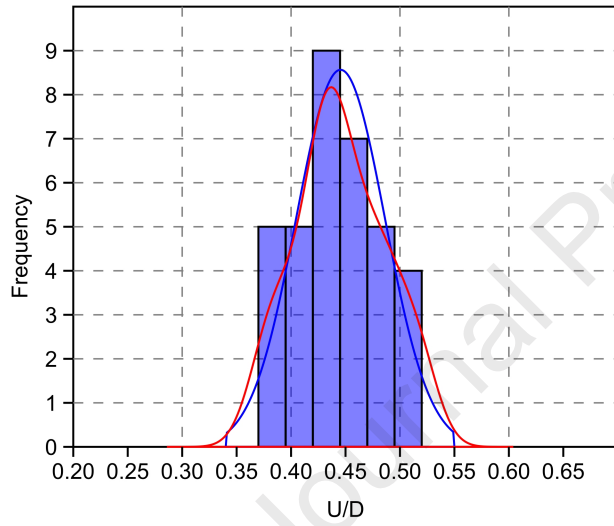
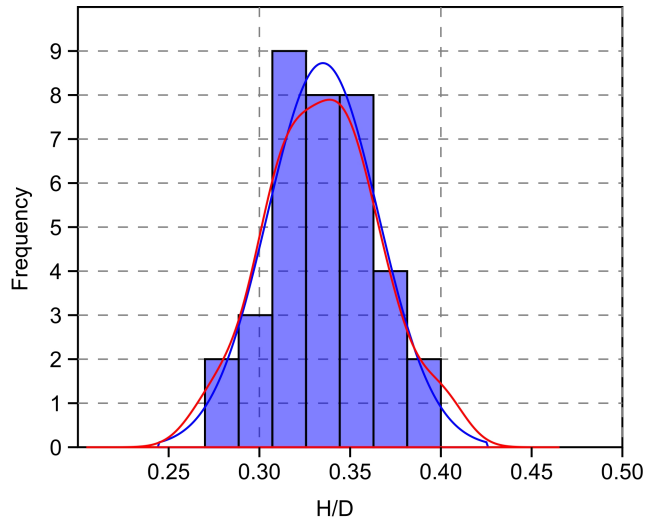


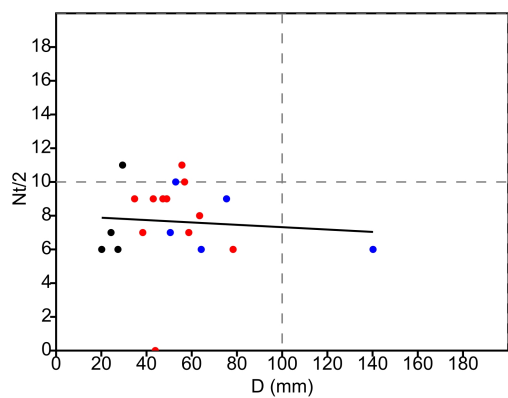
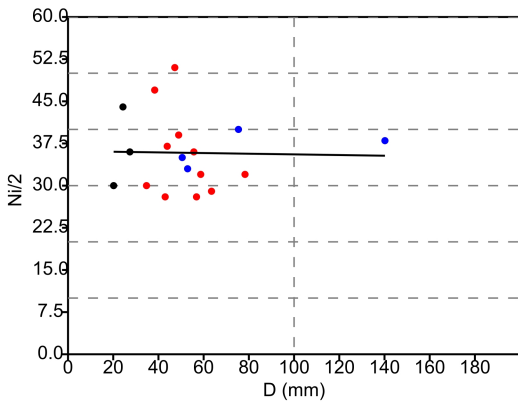
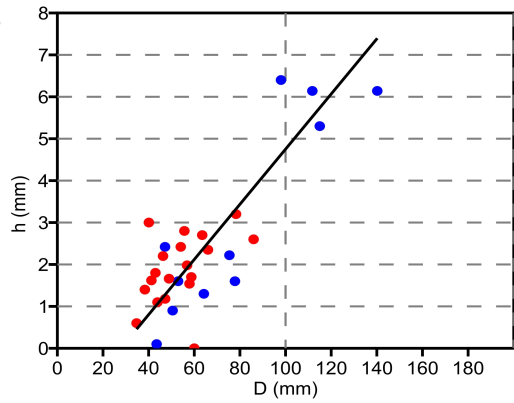
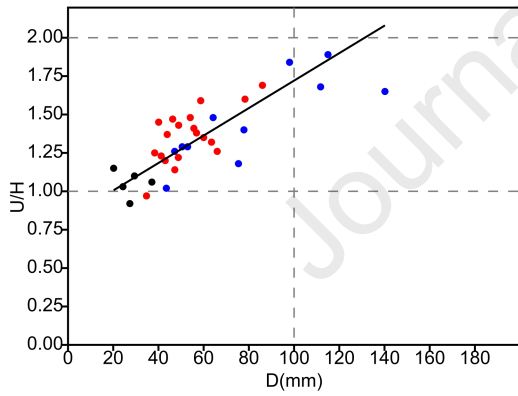
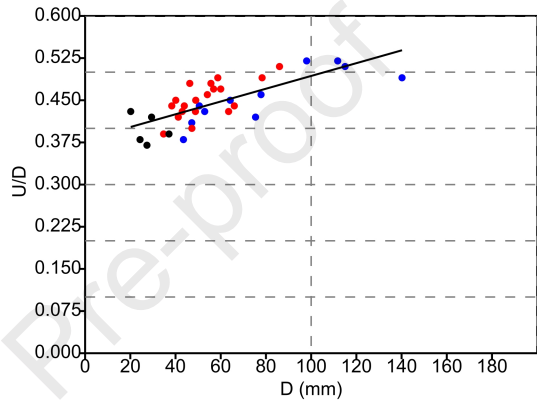
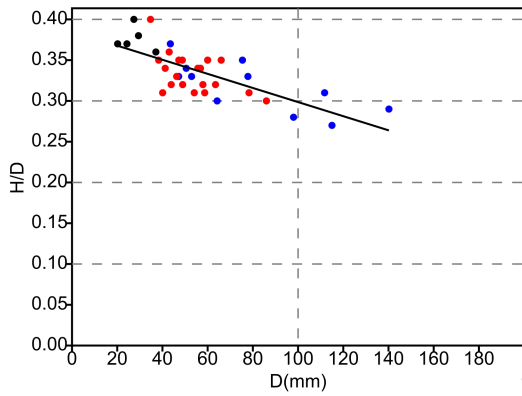
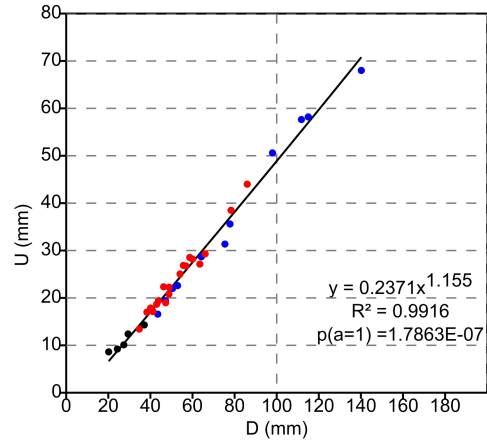
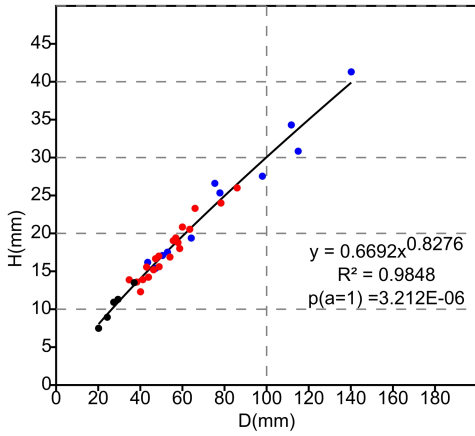


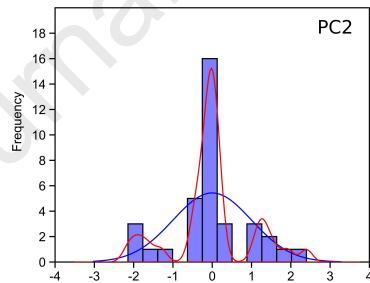
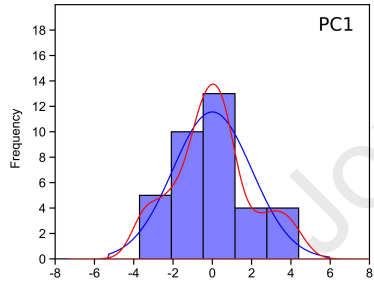
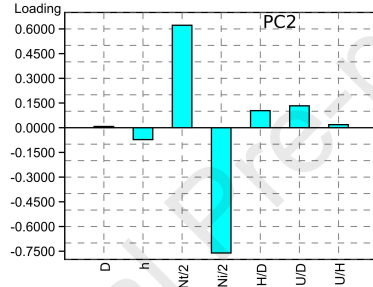
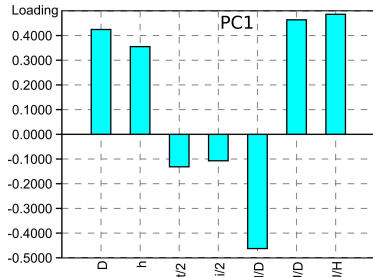
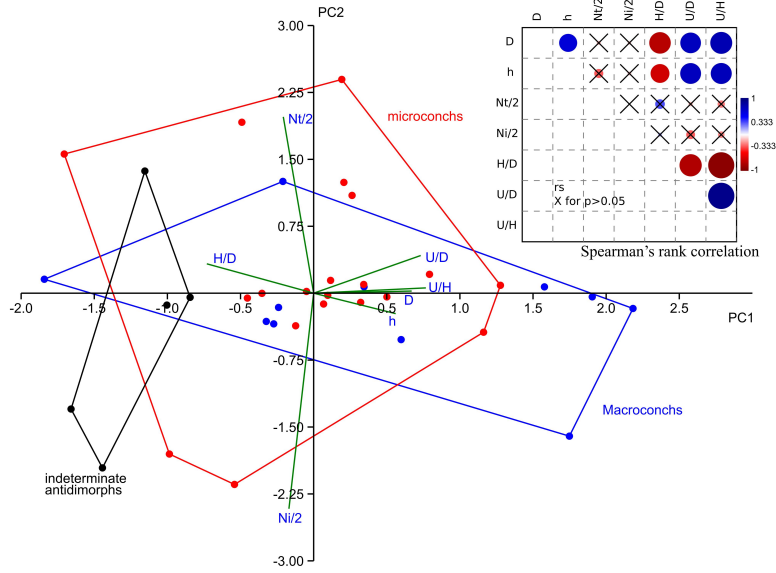












**Declaration of interests**

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

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