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1 **Interactions of plutons and detachments,**
2 **comparison of Aegean and Tyrrhenian granitoids**
3
4

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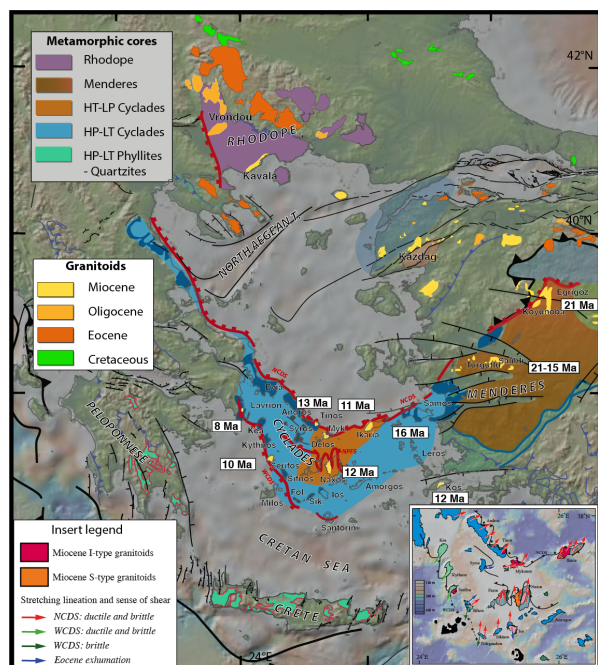
14 **Abstract:** Back-arc extension superimposed on mountain belts leads to distributed normal
15 faults and shear zones, interacting with magma emplacement in the crust. The composition of
16 granitic magmas emplaced at this stage often involves a component of crustal melting. The
17 Miocene Aegean granitoids were emplaced in metamorphic core complexes (MCC) below
18 crustal-scale low-angle extensional shear zones and normal faults. Intrusion in such contexts
19 interacts with extension and shear along detachments, from the hot magmatic flow within the
20 pluton root zone to the colder ductile and brittle deformation along the detachment. A
21 comparison of the Aegean plutons with the Elba Island MCC in the back-arc region of the
22 Apennines subduction shows that these processes are characteristic of pluton-detachment
23 interactions in general and we discuss a conceptual emplacement scenario, tested by numerical
24 models. Mafic injections within the partially molten lower crust above the hot asthenosphere
25 trigger the ascent within the core of the MCC of felsic magmas, controlled by the strain
26 localization on persistent crustal scale shear zones at the top that guide the ascent until the brittle
27 ductile transition is reached during exhumation. Once the system definitely enters the brittle
28 regime, the detachment and the upper crust are intruded while new detachments migrate upward
29 and in the direction of shearing. Numerical models reproduce the geometry and the kinematic
30 evolution deduced from field observations.
31
32
33



34 1. Introduction

35

36 In the deep parts of orogens, the flow of melts is coupled with ductile deformation and
37 controlled by buoyancy and tectonic forces (Brown, 1994; Brown and Solar, 1998; Brown,
38 2007). Migmatites, which are weak crustal material as long as they are kept at high temperature,
39 are the source of magmas that concentrate within plutons of various sizes. On the other hand,
40 interactions between magmatism and lithospheric deformation, and more specifically
41 interactions of plutons with crustal-scale tectonics, depend first of all upon the rate of magma
42 production and, to a second order, strain rates. The rate of magma transfer to the crust is indeed
43 so large compared to tectonic strain rates that the construction of plutons is thought in a first
44 approach to be little influenced by the tectonic setting, especially when small plutons are
45 concerned (de Saint Blanquat et al., 2011).



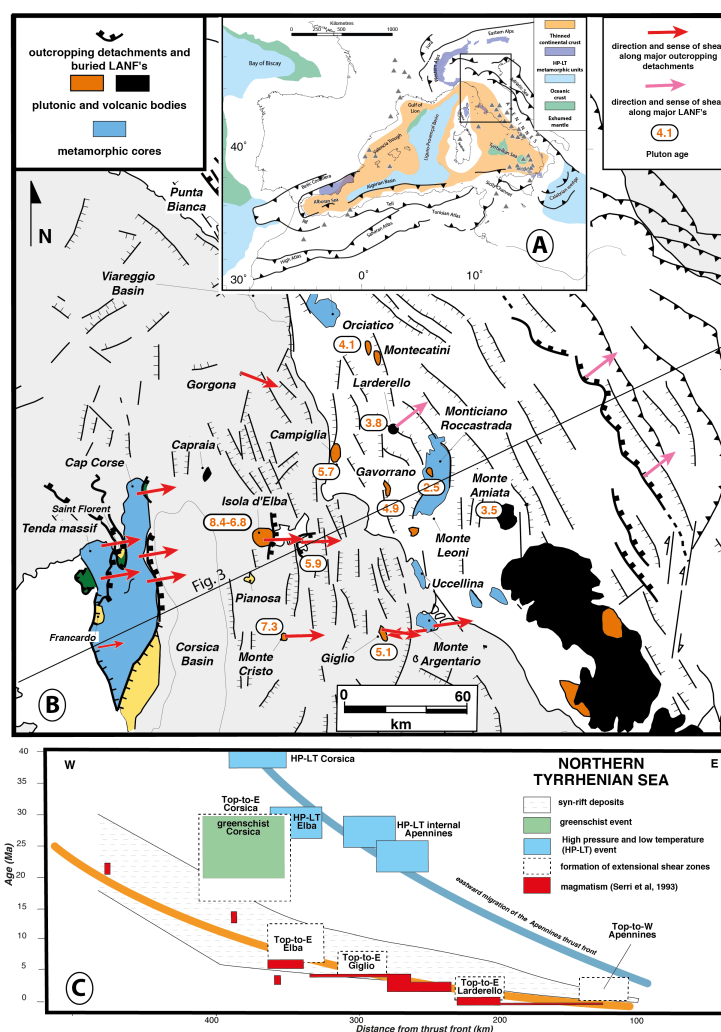
46

47 *Figure 1: Tectonic map of the Aegean region showing the successive generations of granitoids*
48 *since the late Cretaceous. Modified from Jolivet et al. (2015).*

49 The Miocene Aegean plutons (figure 1), emplaced in an extensional context within
50 metamorphic core complexes (MCCs), may however depart from this behaviour. Despite a
51 moderate volume, they have indeed recorded the complete evolution from syn-tectonic
52 magmatic flow to localized mylonitic deformation along the main detachment (Faure and
53 Bonneau, 1988; Urai et al., 1990; Faure et al., 1991; Lee and Lister, 1992; Gautier et al., 1993;



54 Laurent et al., 2015; Rabillard et al., 2015; Bessi re et al., 2017; Rabillard et al., 2018). All of
 55 them moreover show a systematic magmatic and tectonic evolution of the host MCCs with
 56 several magmatic pulses and a series of detachments forming sequentially during exhumation
 57 (Rabillard et al., 2018). Several of them also show an association of mixed or mingled felsic
 58 and mafic magmas, with an evolution from a significant component of crustal melting toward
 59 more mafic composition, a trend that is common in post-orogenic magmas (Bonin, 2004).



60

61 *Figure 2: Tectonic map of the Central and Western Mediterranean with the distribution of*
 62 *recent magmatism (Savelli, 1988; Serri et al., 1993; Savelli, 2002b, a; Duggen et al., 2005;*
 63 *Avanzinelli et al., 2009; Savelli, 2015). Tectonic map (B) of the northern Tyrrhenian region*
 64 *and Northern Apennines and (C) a diagram showing the evolution of the ages of syn-rift basins,*
 65 *metamorphic events and magmatism along a cross-section from Corsica to the Apennines,*
 66 *modified from Jolivet et al. (1998).*

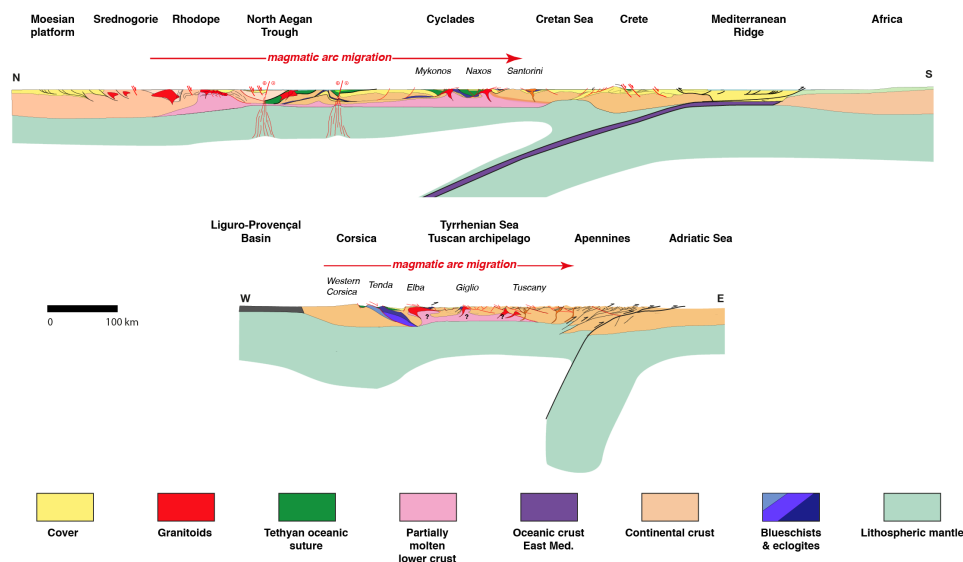


67

68 Whether these features are characteristic of syn-extension plutons in post-orogenic back-arc
69 environments is the question we address in this paper, through a comparison of the Aegean
70 plutons with those of the northern Tyrrhenian Sea and Tuscany, with a focus on Elba Island in
71 the Tuscan archipelago (figure 2). Striking similarities can indeed be observed between the two
72 contexts in terms of tectonic and magmatic evolution. We propose and test a scenario of
73 formation and emplacement of plutons in a back-arc post-orogenic context below crustal-scale
74 detachments.

75 2. Geodynamic context

76 The Aegean and North Tyrrhenian granitoids were emplaced during the Miocene and part
77 of the early Pliocene in the back-arc regions of the Hellenic and Apennines subduction,
78 respectively (Serri et al., 1993; Jolivet et al., 1998; Pe-Piper and Piper, 2002, 2007; Avanzinelli
79 et al., 2009; Jolivet et al., 2015; Rabillard et al., 2018) (figure 3).



80

81 *Figure 3: Two lithospheric-scale cross-sections of the Aegean domain (Jolivet and Brun,*
82 *2010) and the Northern Tyrrhenian Sea and the Apennines (Jolivet et al., 1998).*

83

84 These two subduction zones started to retreat approximately at the same time, 30-35 Ma
85 ago (Jolivet and Faccenna, 2000). A first-order change in the geodynamics of this region indeed
86 occurred at this period, also coeval with the hard collision between Africa and Eurasia in the
87 eastern and westernmost Mediterranean. The subducting African lithosphere, locked between
88 two collision zones, continued to subduct northward but with a significant component of retreat.



89 Since that time subduction has been continuous, with however several episodes of slab
90 detachment and tearing (Wortel and Spakman, 2000; Spakman and Wortel, 2004; Faccenna and
91 Becker, 2010; Faccenna et al., 2014). The Aegean plutons studied in this paper were emplaced
92 during the formation of a large tear in the subducting lithosphere between 16 and 8 Ma (Jolivet
93 et al., 2015). The oldest North Tyrrhenian pluton is dated around 7 Ma in Elba (Westerman et
94 al., 2004) and the youngest ones, Pliocene in age (Serri et al., 1993), are currently exploited for
95 geothermal energy in Tuscany (Rossetti et al., 2008; Rochira et al., 2018). All these plutons
96 contain a significant component of crustal melts and some of them are linked with migmatite
97 domes such as on Naxos, Mykonos and Ikaria (Jansen, 1977; Urai et al., 1990; Denèle et al.,
98 2011; Beaudoin et al., 2015; Vanderhaeghe et al., 2018). Mixing and mingling with mafic
99 magmas are also observed in some of these plutons and the general evolution shows an increase
100 of the mantle component with time.

101 Most of these plutons are associated with low-angle normal faults (LANF) and shear
102 zones and they were emplaced in the core of MCCs (Faure et al., 1991; Lee and Lister, 1992;
103 Lister and Baldwin, 1993; Daniel and Jolivet, 1995; Jolivet et al., 1998; Berger et al., 2013;
104 Bessièrè et al., 2017; Rabillard et al., 2018). These LANF and associated ductile shear zones
105 (we use the term "detachment" for the whole structure, brittle and ductile) started to form before
106 the emplacement of the plutons, in both regions. The main difference between the two regions
107 are the kinematics of these detachments (figure 3) (Jolivet et al., 2008) and the role of tectonic
108 inheritance. In the Aegean, most of the MCCs are capped by north-dipping detachments except
109 in the southwest where south-dipping detachments are observed. The north-dipping
110 detachments probably partly reactivate former thrusts related to the building of the Hellenides
111 orogenic wedge. In the Northern Tyrrhenian Sea and in Tuscany, all detachments dip eastward,
112 i.e. toward the subduction zone. In that case, the detachments cannot reactivate the former
113 thrusts of the internal Apennines that dip westward. Only in the case of the oldest detachments,
114 found in Alpine Corsica, can they correspond to reactivated thrusts. The case of Elba Island
115 shows the detachments cutting down-section eastward within the stack of former nappes (Keller
116 and Pialli, 1990; Collettini and Holdsworth, 2004). Whatever the nature (i.e. reactivated
117 structure or not) and the sense of shear of these detachments, the interaction with the plutons
118 follows a similar pattern that we recall below, first briefly for the Aegean region and then for
119 Elba Island (see Rabillard et al., 2018, for details on the Aegean).

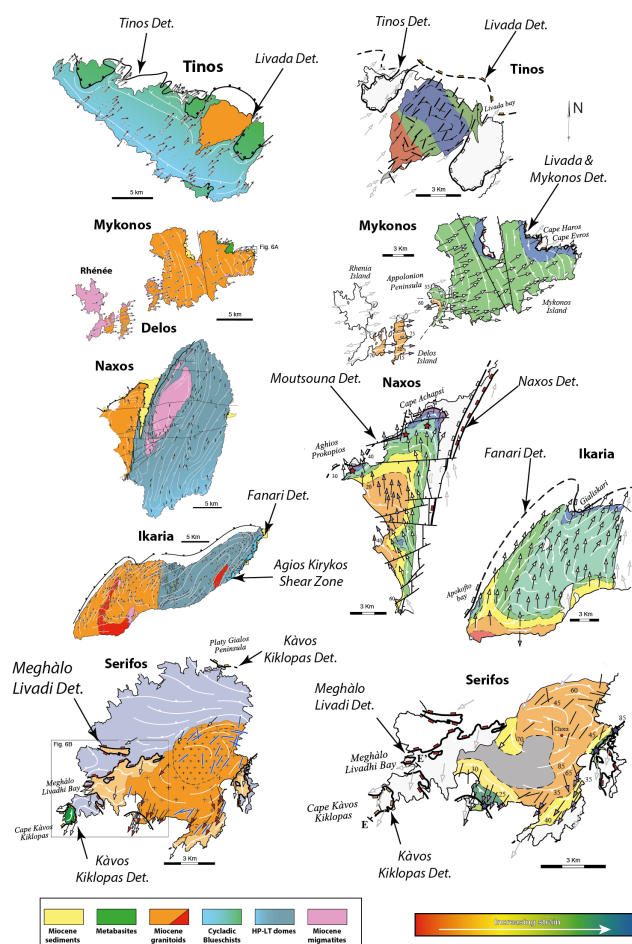
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121 3. Aegean plutons

122

123 The Miocene Aegean and Menderes plutons were emplaced during a short time period
 124 between ~20 Ma and 8 Ma, the oldest cropping out in the Menderes massif and the youngest in
 125 the western part of the Aegean region (figures 1, 3) (Jolivet et al., 2015). Those occupying the
 126 Cycladic domain are all associated with detachments, either north or south-dipping (figure 4)
 127 (Grasemann and Petrakakis, 2007; Rabillard et al., 2018).



128

129 *Figure 4: Five examples of the Aegean granitoids showing the interactions between*
 130 *deformation and intrusion, after Rabillard et al. (2018) and references therein. The left*
 131 *column show maps of the entire islands and the right column shows the internal fabrics of the*
 132 *plutons. These maps were obtained based on deformation grades observed in the field, a scale*
 133 *of grades was designed for each pluton to describe the gradients. Arrows show stretching*
 134 *lineations and sense of shear and black bars on the Tinos and Serifos plutons shows the*
 135 *direction of the magnetic lineation.*



136

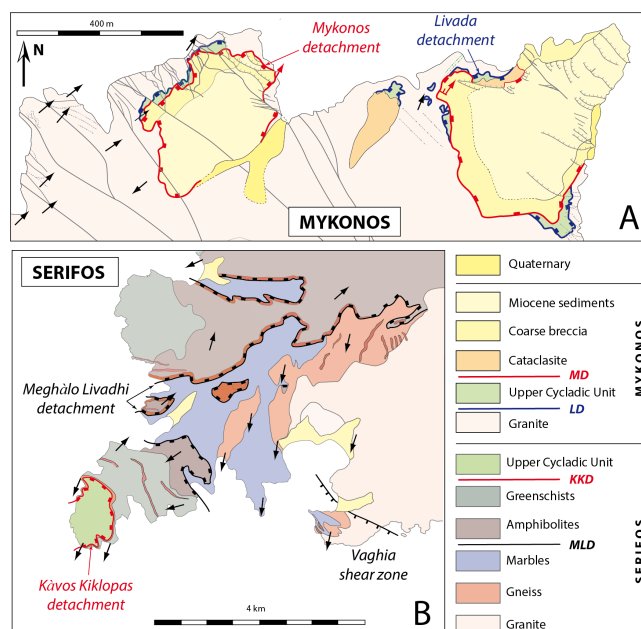
137 Except for Serifos and Lavrion plutons, associated with the West Cycladic Detachment
138 System (WCDS) (Grasemann and Petrakakis, 2007; Berger et al., 2013; Scheffer et al., 2016),
139 the plutons crop out in the core of MCCs exhumed by north-dipping detachments, such as the
140 North Cycladic Detachment System (NCDS) (Gautier and Brun, 1994b, a; Jolivet et al., 2010)
141 or the Naxos-Paros Fault System (NPFS) (Urai et al., 1990; Gautier et al., 1993; Vanderhaeghe,
142 2004; Bargnesi et al., 2013; Cao et al., 2017). The detachments upper plate is made of the Upper
143 Cycladic Nappe, a remnant of the Pelagonian domain, made of greenschists-facies metabasites
144 or serpentinite with, in a few cases, early to late Miocene sediments deposited during extension
145 (Angelier et al., 1978; Sanchez-Gomez et al., 2002; Kuhleemann et al., 2004; Menant et al.,
146 2013). The MCCs are made of various units of the Cycladic Blueschists, more or less
147 retrograded in the greenschist-facies, or the Cycladic basement, showing HT-LP metamorphic
148 facies and even anatexis conditions on several islands, such as Naxos, Paros, Mykonos or Ikaria
149 (Buick and Holland, 1989; Urai et al., 1990; Buick, 1991; Keay et al., 2001; Duchêne et al.,
150 2006; Seward et al., 2009; Kruckenberg et al., 2011; Beaudoin et al., 2015; Laurent et al., 2015;
151 Rabillard et al., 2015; 2018). The plutons intruded these MCCs and were sheared at the top by
152 the detachments during their emplacement (Rabillard et al., 2018).

153 The granitoids show a variety of facies and composition, but most of them have a crustal
154 melting component and some are closely associated with migmatites, as on Ikaria or Mykonos
155 (Vanderhaeghe, 2009; Kruckenberg et al., 2011; Denèle et al., 2011; Beaudoin et al., 2015;
156 Jolivet et al., 2021). Compositions show a common trend for these plutons indicating that they
157 crystallized primarily from I-type magmas with some contamination by the continental crust
158 and little fractionation (figure A1, Appendix A). Field evidence show a close association of
159 these I-type intrusions with two-micas granites (in Ikaria for instance), migmatites, or both
160 (Ikaria, Naxos, Paros, Rheneia-Delos) (Pe-Piper et al., 1997; Pe-Piper, 2000; Pe-Piper et al.,
161 2002; Vanderhaeghe, 2004; Bolhar et al., 2010; Bolhar et al., 2012; Bargnesi et al., 2013;
162 Beaudoin et al., 2015; Laurent et al., 2015; Jolivet et al., 2021). Tinos, Ikaria and Serifos
163 granitoids were emplaced in several magma batches with an evolution through time,
164 characterized by more and more mafic compositions and a decrease of the grain size
165 (Grasemann and Petrakakis, 2007; Ring, 2007; Bolhar et al., 2010; Petrakakis et al., 2010;
166 de Saint Blanquat et al., 2011; Bolhar et al., 2012; Beaudoin et al., 2015; Laurent et al., 2015;
167 Rabillard et al., 2015; Ducoux et al., 2016). On Serifos and Naxos, the farthest parts of the
168 pluton from the detachment show an enrichment in mafic enclaves and evidence for magma



169 mixing and mingling in the roots of the rising plutons (Rabillard et al., 2015; Bessière et al.,
170 2017; Rabillard et al., 2018).

171 A common evolution is observed in several of these plutons during their interaction with
172 the system of detachments exhuming their host MCC (Rabillard et al., 2018). A series of two
173 or three detachments is observed (figure 4, figure 5).



174

175 *Figure 5: Details of the two detachments on Mykonos and Serifos. A: northeastern Mykonos*
176 *(see location on figure 5), B: Southwest Serifos. The Mykonos (MD) and Livada (LD)*
177 *detachments on Mykonos and the mineralized veins and normal faults (baryte and iron-*
178 *hydroxides) – grey- are after Menant et al. (2013). The Kavos Kiklopas (KKD) and Meghàlo*
179 *Livadi (MLD) detachments on Serifos are after Grasemann and Petrakakis (2007) and*
180 *Ducoux et al. (2016).*
181

182 The deepest one is mostly ductile and has started to act long before the granitic intrusion
183 that ultimately intrudes it. The upper detachments are mostly brittle and are locally intruded by
184 dykes and sills emanating from the main pluton. When a sedimentary basin is present, it is
185 deposited on top of the uppermost detachment during extension and can be partly affected by
186 mineralized veins (Menant et al., 2013). All plutons show a gradient of shearing deformation
187 toward the detachment with an evolution from ductile to brittle (Figure 4). The maps shown in
188 figure 4 were drawn after detailed field observations and the construction of a scale of up to 7
189 grades of progressive deformation, from non-deformed granitic texture to ultra-mylonites, with
190 the progressive appearance of foliation, stretching lineation, localization of C and C' shear

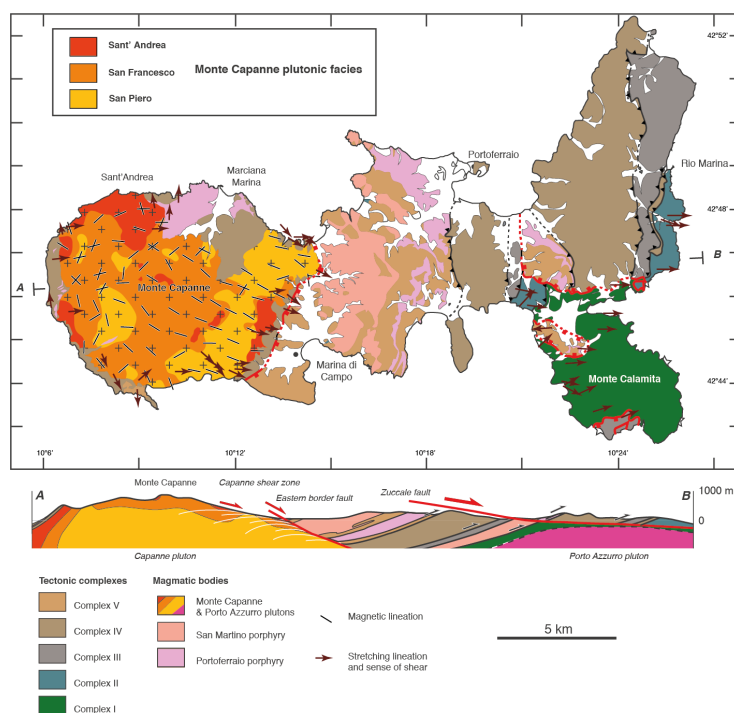


191 bands (Berthé et al., 1979; Lister and Snoke, 1984), for details see Rabillard et al. (2018). The
192 inner parts of the plutons show mixing of acidic and mafic magmas and a co-magmatic
193 deformation, co-axial with the post-solidus deformation along the detachment (Rabillard et al.,
194 2015; 2018). The flow of magma is thus oriented by the regional strain field. Serifos shows (i)
195 a decrease of grain size through time with an inner facies with smaller grain size and finally
196 fine-grained dykes and (ii) evidence for hydrothermalism in the root zone of the pluton,
197 suggesting that the magmatic system was open upward with a possible volcano-plutonic system
198 (Rabillard et al., 2015; 2018).
199

200 4. North Tyrrhenian plutons

201

202 The Monte Capanne pluton on Elba island (figure 6) is the oldest of a series of plutons
203 cropping out in the Tuscan archipelago and onshore Tuscany (Serri et al., 1993; Westerman et
204 al., 2004; Avanzinelli et al., 2009).



205

206 *Figure 6: Tectonic map and cross-section of Elba Island showing the main tectonic units and*
207 *the main extensional shear zones and detachments, modified after Bianco et al. (2015).*
208 *Details of the internal structure of the Monte Capanne intrusions are reported based on*
209 *Farina et al. (2010).*



210

211 These plutons belong to magmas migrating from west to east between the end of the
212 Oligocene and the Quaternary, mimicking the migration of the Apennines thrust system and the
213 HP-LT metamorphism of the internal Apennines and Tuscan Archipelago (Serri et al., 1993;
214 Jolivet et al., 1998) (figure 2). This situation is thus very similar to the Aegean Sea. The
215 decrease of the time lag between the recording of HP-LT metamorphism or the activation age
216 of the thrust front and the magmatism has been interpreted as a consequence of slab steepening
217 during retreat (Jolivet et al., 1998; Brunet et al., 2000). Pluton ages decrease eastward from ~8
218 Ma to 2-3 Ma (figure 2). Among the youngest plutons are those powering the active geothermal
219 fields of Larderello and Monte Amiata (Camelli et al., 1993; Brogi et al., 2003; Rossetti et al.,
220 2008). The oldest plutons are observed offshore on Elba (Monte Capanne and Porto Azzuro
221 plutons), Monte Cristo and Giglio islands (Westerman et al., 1993) (figure 2). These four
222 plutons are granodiorites/monzogranites and they all display a contamination with crustal
223 magmas with a main source thought to be lower crustal anatexis (Serri et al., 1993; Innocenti
224 et al., 1997). They were emplaced within an overall extensional context during the rifting of the
225 Northern Tyrrhenian Sea in the back-arc region of the Apennines (Jolivet et al., 1998). First
226 evidenced in Alpine Corsica and on Elba island, a series of east-dipping low-angle detachments
227 controlled the kinematics of extension along the Corsica-Apennines transect from the
228 Oligocene onward (Jolivet et al., 1998). Extension is shown to migrate from west to east with
229 time and it is active at present in the highest altitude regions of the Apennines just west of Corno
230 Grande peak with however west-dipping normal faults (D'Agostino et al., 1998). The youngest
231 east-dipping low-angle normal faults are seismically active in the Alto Tiberina region
232 (Collettini and Barchi, 2002, 2004; Pauselli and Ranalli, 2017). Evidence for top-to-the east
233 shearing deformation is found within the plutons of the Tuscan archipelago, but the
234 detachments crop out nicely mostly on Elba island (Keller and Piali, 1990; Daniel and Jolivet,
235 1995; Collettini and Holdsworth, 2004; Liotta et al., 2015).

236

237 **4.1. Monte Capanne pluton, Elba Island**

238 Elba, the largest island of the Tuscan archipelago, shows the relations between
239 peraluminous magmatic bodies and two east-dipping low-angle shear zones cutting down-
240 section within the Tuscan nappe stack emplaced before extension started (figure 6) (Keller and
241 Piali, 1990; Bouillin et al., 1993; Pertusati et al., 1993; Daniel and Jolivet, 1995; Westerman
242 et al., 2004; Bianco et al., 2015). Five thrust packages (complexes I to V) are separated by west-
243 dipping low-angle reverse faults (Trevisan, 1950; Barberi et al., 1967; Perrin, 1975; Pertusati



244 et al., 1993; Bianco et al., 2015; Bianco et al., 2019). Long thought free of any *HP-LT* imprint,
245 at variance with the nearby Gorgona and Giglio islands where the presence of Fe-Mg-carpholite
246 attest for blueschist-facies metamorphism (Rossetti et al., 1999a; 1999b), the nappe stack has
247 only recently revealed *HP-LT* parageneses along the east coast of the island (Bianco et al.,
248 2015). Through a correlation with the *HP-LT* units of Gorgona (Rossetti et al., 1999), where
249 $^{40}\text{Ar}/^{39}\text{Ar}$ dating on micas yielded ages around 25 Ma (Brunet et al., 2000), the Elba blueschists
250 are attributed to the Oligocene-Early Miocene.

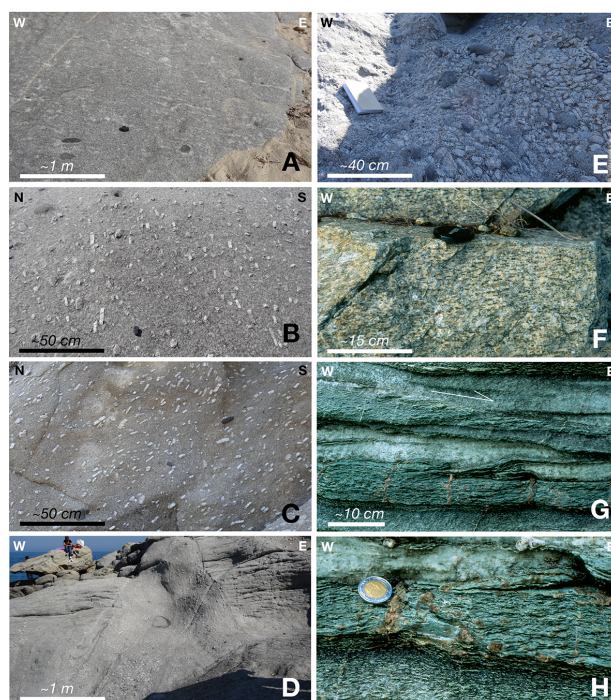
251 The Nappe stack is intruded by the shallow-level San Martino and Portoferraio porphyries
252 coeval with the Monte Capanne pluton (figure 6), spanning a short period between 8 and 6.8
253 Ma, showing that the magma has intruded the detachment in a late stage (Saupé et al., 1982;
254 Juteau et al., 1984; Ferrara and Tonarini, 1985; Bouillin et al., 1994; Westerman et al., 2004).
255 The Monte Capanne intrusion makes the major part of the western half of the island and the
256 highest peak. It is surrounded by a contact metamorphic aureole developed at the expense of
257 the nappe stack (Duranti et al., 1992; Dini et al., 2002; Rossetti et al., 2007; Rossetti and Tecce,
258 2008). The metamorphic parageneses within the aureole suggest an emplacement at a depth of
259 4-5 km (Dini et al., 2002; Rocchi et al., 2002; Farina et al., 2010; Pandeli et al., 2018). The
260 pluton shows an internal deformation with a gradient of shearing toward the east attested by the
261 magnetic fabric, stretching lineation and sense of shear (Bouillin et al., 1993; Daniel and Jolivet,
262 1995). The pluton and the metamorphic aureole are separated from the nappe stack by an east-
263 dipping low-angle shear zone (Capanne shear zone) evolving into a brittle east-dipping fault
264 (eastern border fault) (Daniel and Jolivet, 1995). Syn-kinematic contact metamorphism
265 minerals coeval with top-to-the east kinematic indicators attest for the syn-kinematic nature of
266 the intrusion (Daniel and Jolivet, 1995; Pandeli et al., 2018).

267 The eastern part of the island shows granitic dykes emanating from the buried younger
268 Porto Azzuro pluton intruding the Calamiti schists complex (Complex I, figure 6) (Daniel and
269 Jolivet, 1995; Maineri et al., 2003; Musumeci and Vaselli, 2012). Here too, evidence for top-
270 to-the east shearing at the time of intrusion have been described (Daniel and Jolivet, 1995)
271 (figure 7). The pluton and the Calamiti schists are topped by the Zuccale low-angle normal fault
272 that cuts down-section across the entire nappe stack with clear evidence of top-to-the east
273 shearing (figure 8) (Keller and Pialli, 1990; Keller et al., 1994; Collettini and Holdsworth,
274 2004).

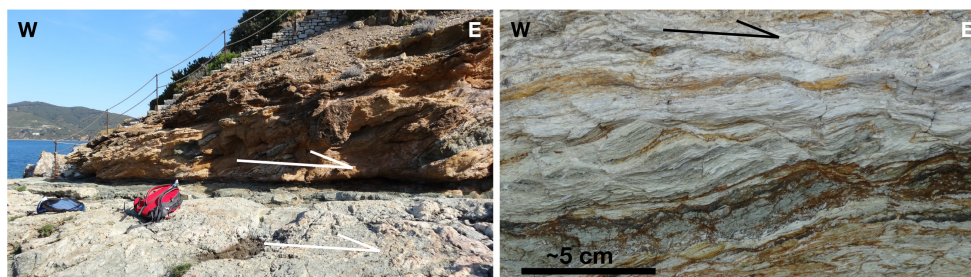
275 The main facies of the Monte Capanne pluton exhibits a constant, peraluminous,
276 monzogranitic composition (Poli et al., 1989; Dini et al., 2002; Gagnevin et al., 2004) while the
277 mafic microgranular enclaves (MME) varies from tonalitic-granodioritic to monzogranitic. The



278 leucogranitic dykes are syenogranitic in composition (Gagnevin et al., 2004). Gagnevin et al.
279 (2004) proposed a multiphase magmatic emplacement from peraluminous magmas issued from
280 melting of a metasedimentary basement and hybridized with mantle-derived mafic magmas
281 whose heat supply possibly enhanced wall-rock assimilation. In addition, injection of mantle-
282 derived magma in the Sant' Andrea facies would have triggered extensive fractionation and
283 mixing of the basic magma with the resident monzogranitic mush (Poli and Tommasini, 1991).



284
285 *Figure 7: Photographs of the Sant' Andrea facies in the Monte Capanne pluton and of the*
286 *deformation along the eastern margin of the pluton, in the pluton itself and in the contact*
287 *metamorphic aureole. A: general view of the orientation of K-feldspar megacrysts and some*
288 *mafic enclaves. B: Detailed view of the oriented K-feldspar megacrysts (horizontal plane). C:*
289 *zoom on the orientation of K-feldspar megacrysts (vertical plane). D: Cotoncello dyke. E:*
290 *cluster of K-feldspar megacrysts in the vicinity of the Cotoncello dyke. F: Mylonitic foliation*
291 *within the Monte Capanne shear zone. G: Sigmoidal foliation and top-to-the-east sense of*
292 *shear within the metamorphic aureole of the Monte Capanne pluton. H: detailed view of syn-*
293 *kinematic contact metamorphism garnets in veins perpendicular to the regional stretching*
294 *direction.*
295



296

297 *Figure 8: Photographs of the Zuccale detachment and its internal structure (see Collettini*
298 *and Holdsworth, 2004). Upper: overview of the detachment fault. Lower left: Detail of the*
299 *contact zone and the truncated foliation in the hanging wall. Lower right: detail of the shear*
300 *bands indicating top-to-the east kinematics.*
301

302 The internal magmatic structure of Monte Capanne pluton has been described based on the
303 abundance of large alkali-feldspar phenocrysts (Farina et al., 2010). Three main facies
304 corresponding to different magma batches emplaced within a too short period to be
305 discriminated based on geochronology are reported with downward fining of grain size (figure
306 6). The largest grain size characterizes the upper Sant' Andrea facies that mainly crops out in
307 the northwest of the pluton, while the finest grain size is observed in the lower San Piero facies
308 cropping out mainly in its eastern part, within the zone affected by the most intense shearing.
309 These three facies delineate an asymmetric dome-shaped bulk structure compatible with the
310 general top-to-the east sense of shear. In the westernmost part of the Monte Capanne pluton
311 near Sant' Andrea, mafic products are observed as large enclaves, with evidence of magma
312 mixing and mingling. These mafic enclaves are mostly found in the Sant' Andrea facies that
313 was emplaced first. Their occurrence in the westernmost part of the plutonic body, the farthest
314 from the detachment, with a geometry similar to that observed on Serifos island in the Cyclades
315 (Rabillard et al., 2015), suggests that they are associated with the root of the pluton.

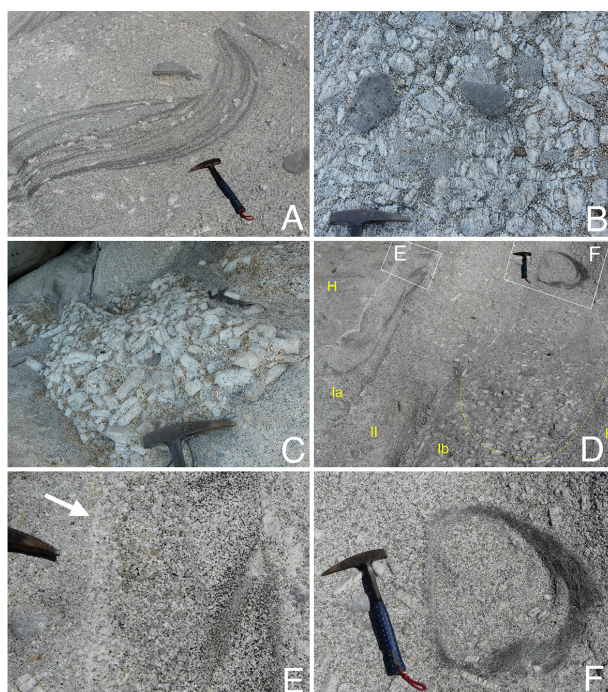
316 Assuming that the three main felsic facies correspond to three successive intrusion
317 batches, one observes an evolution toward finer grain size through time, an evolution that is
318 compatible with progressive exhumation and also with a shorter residence time in the magma
319 chamber, suggesting opening of the magmatic plumbing toward the surface leading to volcanic
320 activity, as recorded above the detachment. The last episodes of intrusive activity are seen as a
321 series of felsic dykes striking N-S or NE-SW, due to eastward extensional brittle deformation
322 while the pluton was at near solidus conditions.

323

324 *4.2. Orientation of K-feldspar megacrysts*



325 A detailed study of the orientation of feldspar megacrysts was conducted along the shore
326 near Sant'Andrea (figures 7, 9, 10, 11). The pluton is there characterized by a high
327 concentration of megacrysts and of mafic enclaves reaching several meters in size. Megacrysts
328 show euhedral shapes in general.



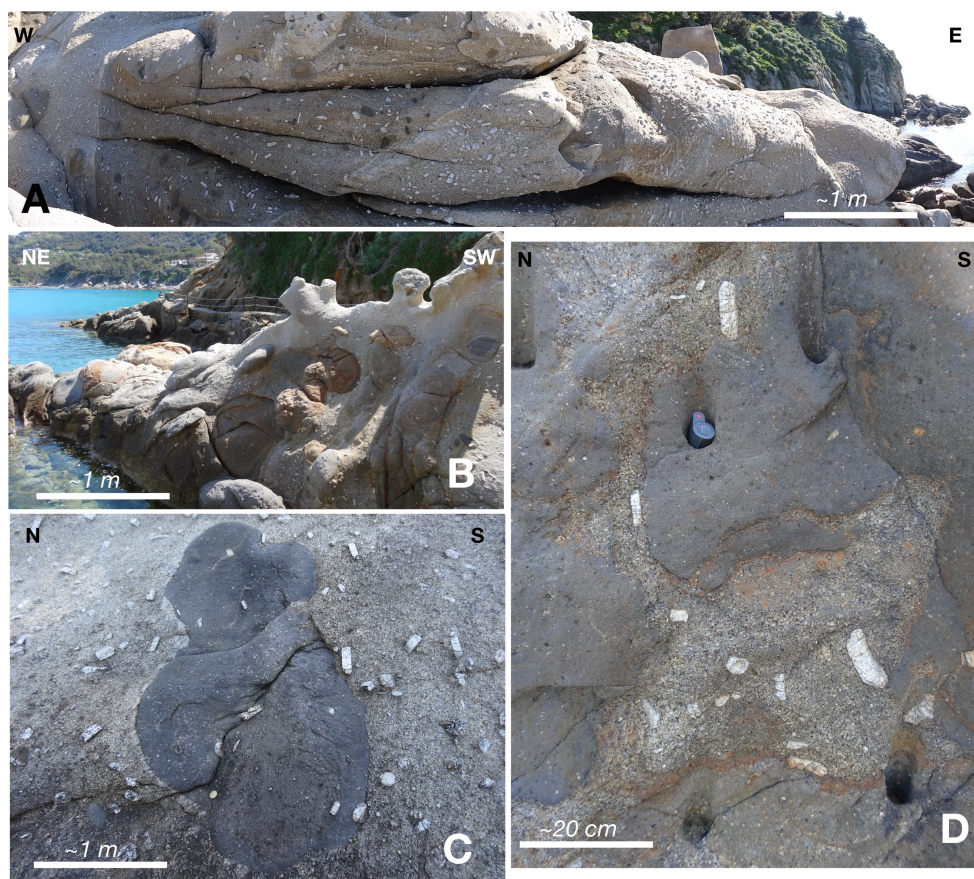
329

330 *Figure 9: Photographs of the root zone of the Sant'Andrea facies in the melange zone in and*
331 *around the Cotoncello dyke. A: Schlieren with cross-bedding. B: Mush zone with cluster of*
332 *K-feldspar megacrysts and mafic enclaves. C: Isolated blob of mush zone with large K-*
333 *feldspar megacrysts. D: Folded alternation of leucocratic and melanocratic layers with*
334 *schlierens. E: Detail of D, schlieren. F: detail of D: schlieren tube.*

335

336 Although the orientation of megacrysts is quite stable at the scale of a few hundred meters,
337 in the westernmost region the presence of the large enclaves is associated with a disorientation
338 of the megacrysts (figure 10A), indicating increasing tortuosity of the flow wrapping around
339 them associated with local turbulence in pressure shadows. Smaller enclaves are in general
340 aligned with the megacrysts. Some of the mafic enclaves with lobate shapes show a sharp
341 boundary with the felsic matrix, suggesting quenching of a hot mafic magma within the cooler
342 felsic magma (Fernandez and Barbarin, 1991; van der Laan and Wyllie, 1993; Fernandez and
343 Gasquet, 1994).

344



345

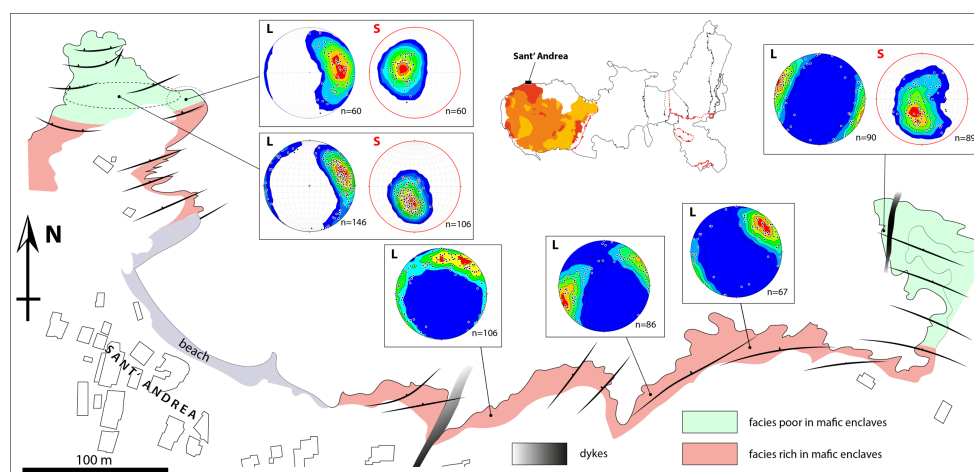
346 *Figure 10: Photographs of large mafic enclaves and melange zones in the westernmost part*
347 *of the Sant'Andrea facies showing the disorientation of K-feldspar megacrysts and melange*
348 *facies (mingling).*

349

350 Other enclaves are less mafic and show evidence of magma mingling-mixing. The
351 disaggregation process of the mafic magma responsible of these enclaves could happened
352 either in a deep-seated magmatic chamber (Christofides et al., 2007), or more likely in the
353 ascent conduit as a result of remelting of chilled mafic margins (Fernández and Castro, 2018)
354 and subsequent viscous fingering dynamics (Perugini et al., 2005). Megacrysts contain
355 inclusions of biotite, plagioclase and quartz and show euhedral shapes in general, although
356 resorption surface has been noticed (Gagnevin et al., 2008). Other enclaves are less mafic and
357 show evidence of magma mingling-mixing. These enclaves are associated with an aureole
358 where feldspar crystals are concentrated, showing that the assimilation of the enclave occurred
359 at the magmatic stage. Megacrysts are sometimes included within the mafic enclaves, showing
360 that they were already present before the solidification of enclaves and thus providing evidence



361 of low viscosity contrast between the enclaves and the host magma at the magmatic stage. All
362 these observations suggest that this western zone is a mixing between a mafic magma of mantle
363 origin and a felsic magma partly issued from crustal anatexy and that this part of the pluton is
364 close to the main feeder. This conclusion is confirmed by AMS (anisotropy of magnetic
365 susceptibility) showing that the magnetic foliation and lineation are steeper there than anywhere
366 else in the pluton (Bouillin et al., 1993).



367

368 *Figure 11: Detailed study of the orientation of K-feldspar megacrysts in the Sant-Andrea*
369 *facies with foliation trajectories.*

370

371 Further to the east, still within the Sant' Andrea facies, the main granite is intruded by a
372 N-S syeno-granitic dyke-like structure near Cotoncello headland, made of a finer-grained facies
373 and a lower concentration of megacrysts and enclaves (figure 7D, figure 9). In its vicinity, the
374 host granite contains folded schlierens with cross-bedding (Figure 9A). Within these large
375 schlierens, the megacrysts are aligned parallel with the folded foliation of biotite-rich layers.
376 From place to place, decametric megacryst-rich mush zones enriched in decametric and
377 rounded mafic enclaves occur in this host facies (Figure 9B). In addition, isolated blobs of
378 mush, characterized by an irregular shape, are observed in the coarse-grained, megacryst-poor
379 domains (Figure 9C). These blobs originate from the disruption of preexisting mush zones
380 within the root zone by subsequent magma injection as illustrated by the dyke-like structure
381 (Rodríguez and Castro, 2019). This structure is composed of three successive injections
382 characterized by undulating and fuzzy boundaries (Figure 9D). The westernmost injection
383 (injection Ia, figure 9D) shows folded alternating leucocratic flow-sorted layers made of quartz,
384 K-feldspar and plagioclase with more melanocratic layers rich in biotite that can be described
385 as schlierens (Figure 9E). These schlierens are folded and cross-cut by a subsequent and final



386 injection (Figure 9D, injection II). In the easternmost injection (Figure 9D, injection Ib), K-
387 feldspar megacrysts are accumulated and their orientation defines a concave upwards foliation.
388 Such mineral fabric is similar to those described by Rocher et al. (2018) in finger and drip
389 structures developed at the margins of the Asha pluton (NW Argentina) and interpreted as
390 mechanical accumulation in a downward localized multiphase magmatic flow. In addition, this
391 megacrysts accumulation is associated at its top with ring schlieren that could represent a cross-
392 section of a schlieren tube (e.g. Žák and Klomínský, 2007) (figure 9F). Ring schlierens are also
393 associated with drip structures in the Asha pluton among others (Paterson, 2009; Rocher et al.,
394 2018). The most external rim between the host body H and injection Ia (Figure 9D) is associated
395 with a reaction zone with recrystallization of quartz and K-feldspar (Figure 9E, white arrow).
396 Outside the injections, the mineral fabric shown by the K-feldspar megacrysts tends to
397 reorientate parallel to the rims.

398 All these observations point out to an injection of a low viscosity, crystal poor, magma
399 with a viscosity contrast of about one order of magnitude lower with respect to its host magma
400 (Wiebe et al., 2017). Mineral fabrics and accumulation, folded and ring schlieren indicate that
401 the structures were formed by localized multiphase magmatic flow when the crystallizing host
402 magma remained partially molten, probably containing around 50% of crystals (Weinberg et
403 al., 2001). The Cotoncello dyke-like structure is thus co-magmatic with the Sant'Andrea facies,
404 but the pluton was already enough crystallized to allow the formation of N-S cracks in the
405 crystal mush capable of transmitting tectonic stress where the magma was injected.

406 Between the Cotoncello dyke and the root zone to the west, the proportion of enclaves
407 and megacrysts is everywhere high. Systematic measurements of feldspar megacrysts were made
408 (figure 11). Mineral foliation and lineation represents the main orientation distributions of the
409 orientation of (010) faces and [001] major axis of the measured crystals, respectively. At the
410 scale of a few hundred meters the fabric shows a consistent pattern with a low-angle north-
411 dipping foliation more prominent in regions poorer in mafic enclaves. The lineation is in
412 average E-W trending. Late mafic and acidic dykes strike perpendicular to the lineation. Within
413 the mélange zone, the mineral fabric is often perturbed approaching enclave swarms. Then,
414 the fabric becomes more uniform with variations around an average ENE-WSW trend from
415 N30 to N100°E for the long axes of megacrysts.

416 As the megacrysts were formed in early magmatic conditions (Vernon, 1986; Vernon and
417 Paterson, 2008), they were in suspension within the melt. Such a preferential orientation is due
418 to a rigid rotation of isolated crystals within a viscous matrix submitted to magmatic flow
419 (Fernandez and Laporte, 1984). In the present case, the various observations attesting for a co-



420 magmatic fabric show that the preferential orientation of the megacrysts foliation results from
421 a fossilization of the magmatic flow. The large-scale variations of the foliation attitude suggest
422 in addition that the E-W to ENE-WSW flow was laminar in general, except in the immediate
423 vicinity of the large enclaves where the flow wrapping around these stronger bodies was more
424 turbulent.

425 These detailed observations show that the internal magmatic fabric of the pluton is similar
426 in orientation with its overall tectonic fabric, including the sub-solidus deformation along the
427 eastern margin due to the detachment with a main stretching direction oriented WNW-ESE, as
428 shown by magnetic susceptibility studies (Bouillin et al., 1993) and deformation features near
429 the main eastern contact within the eastern extensional shear zone (Daniel and Jolivet, 1995).
430 This focussing of the pluton fabric, from the magmatic stage to the brittle stage around an E-W
431 stretching direction compatible with the extensional shear along the main detachment, suggests
432 that the magmatic flow was oriented parallel to the main direction of extension active at crustal
433 scale since the magmatic stage. A continuum is thus observed from the magmatic stage to the
434 sub-solidus deformation and the localization of the detachment, and this continues during the
435 emplacement of the younger Porto-Azzuro pluton and the formation of the Zuccale low-angle
436 normal fault. This evolution recalls that of the Aegean plutons summarized above.

437

438 **5. Discussion and modelling**

439

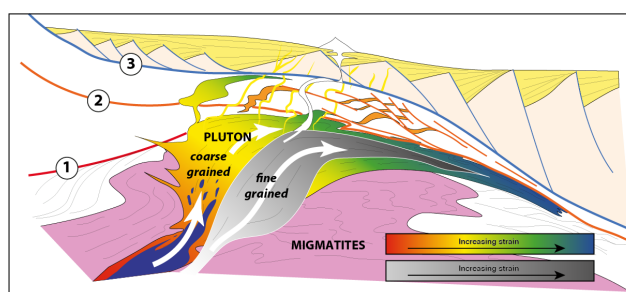
440 *5.1. Synthesis of observations*

441

442 The coaxiality of the structures measured in the Monte Capanne pluton from its magmatic
443 stage to the tectonic overprint is similar to observations made on the Cycladic plutons,
444 especially Ikaria and Serifos where a similar gradual transition is observed from the magmatic
445 stage to the localisation of strain along the main detachment. The similarity goes further as the
446 root of the pluton shows a mixture of mafic and felsic facies. On Serifos (figure 4), field
447 observations show that the root of the pluton is characterized by vertical or steep dykes and
448 some of them are dilacerated by the top-south flow while the magma is still viscous (Rabillard
449 et al., 2015). Moving toward the detachment, the sub-solidus deformation takes over with a N-
450 S trending stretching lineation and top-south kinematic indicators. A similar evolution can be
451 observed in the Raches pluton of Ikaria island in the Cyclades (Laurent et al., 2015). Emplaced
452 below to top-to-the north detachment, the magma shows a steep foliation in the south far from



453 the detachment and it flattens toward the north to become parallel to the detachment plane.
454 Evidence of co-magmatic stretching and shearing parallel to the regional stretching direction is
455 observed in the southern side of the pluton and sub-solidus mylonitization and ultra-mylonites
456 on the northern side. A similar situation can be described in the case of the Naxos granodiorite
457 (Bessière et al., 2017). All cases show the syn-kinematic character of the pluton, the best
458 evidence being the syn-kinematic contact metamorphism.



459

460 *Figure 12: Schematic section showing a conceptual model of the relations between syn-*
461 *kinematic plutons and the detachments, based on the examples of the Aegean and North*
462 *Tyrrhenian plutons. Modified from Rabillard et al. (2018).*
463

464 The Monte Capanne pluton thus shows clear similarities with the Aegean plutons. Figure
465 12 shows a simplified scheme of the geometrical and kinematic relations between detachments
466 and plutons based on the examples of the Aegean and the Northern Tyrrhenian, modified from
467 Rabillard et al. (2018). The root zone of the pluton, characterized with an association of mafic
468 and acidic magmas, shows a steeper upward magmatic flow and evidence of co-magmatic
469 stretching and shearing parallel to the regional direction of extension with a kinematics similar
470 to that of the main detachments. During the emplacement of the pluton, the magma chamber
471 progressively opens toward the surface and the granitoids evolve toward finer-grained facies.
472 Progressive extension and exhumation is accompanied by the inflation of the pluton and
473 injection of dykes across the ductile detachment. New detachments are formed above
474 sequentially.

475

476 **5.2. A conceptual model based on published numerical experiments**

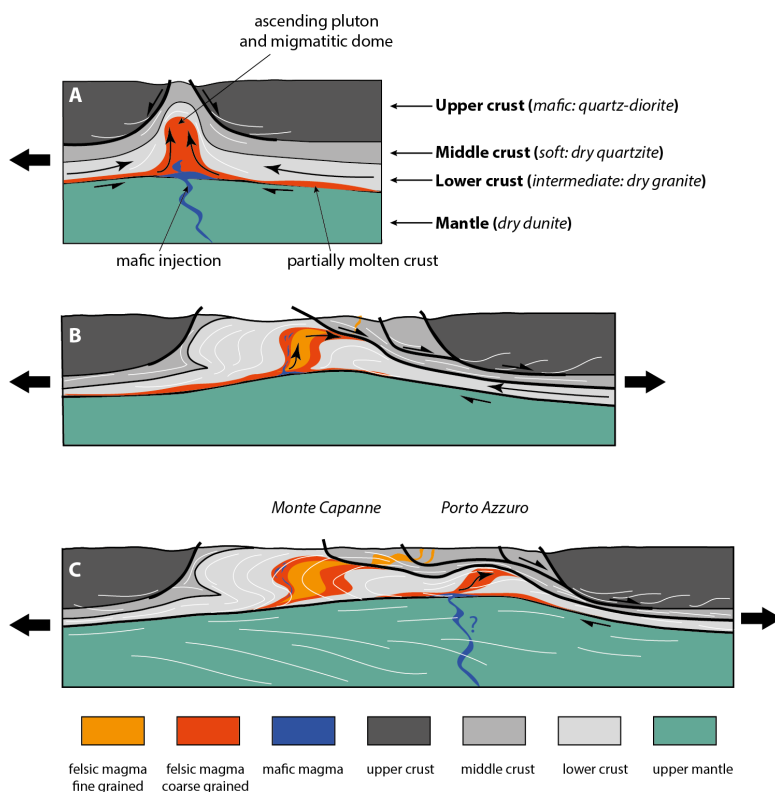
477

478 This evolution can be compared with numerical models. Thompson and Connolly (1995)
479 summarize the three ways of melting lower continental crust as (1) supplying water to the crust
480 to lower the solidus, (2) decreasing pressure and (3) providing additional heat to the lower crust.
481 They also state that extension alone of a thickened crust is unlikely to reach the conditions of



482 lower crustal melting unless some additional heat is given by the mantle. Back-arc regions
 483 above retreating slabs, where the lithosphere is thinned and the asthenosphere advected upward
 484 underneath the crust seem to the first order to fit these conditions (Roche et al., 2018).

485 Schubert et al. (2013) have explored numerically the effect of the injection of molten
 486 mafic material in an extending crust. They show that the injection of this hot material in the
 487 lower crust will induce melting and trigger the formation of felsic magmas that will then ascent
 488 along steep normal faults all the way to the upper crust, forming the observed plutons. This is
 489 a situation that can easily be compared with the Aegean or the Tuscan archipelago where the
 490 granitoids are associated in their root zones with coeval mafic magmas and the felsic plutons
 491 ascend along low-angle detachments. In figure 13, we propose a further conceptual model based
 492 on numerical experiments of post-orogenic extensional deformation with low-angle shear zones
 493 (Huet et al., 2011).



494
 495
 496
 497
 498
 499

Figure 13: Conceptual model of the succession of events leading to the emplacement of a plutonic system below an active series of detachments, based on Huet et al. (2011) and Schubert et al. (2013).



500 In this series of numerical experiments, the thermal gradient and Moho temperature were
501 varied as well the rheological stratification with either a classical rheological stratification or
502 an inverted crustal structure resulting from the formation of the pre-extension nappe stack, the
503 latter setup being used in figure 13, see also Labrousse et al. (2016) for more details on the
504 dynamics of this system with inverted rheological profiles. This latter choice is designed to
505 mimic the Aegean orogenic wedge where the Cycladic Blueschists Unit is sandwiched between
506 the Cycladic Basement and the Upper Cycladic Unit (UCU) (Huet et al., 2009; Jolivet and Brun,
507 2010; Ring et al., 2010). The UCU belongs to the Pelagonian paleogeographic domain and is
508 largely composed of an ophiolite, denser and stronger than the CBU (Labrousse et al., 2016) as
509 well as other basement lithologies (Reinecke et al., 1982; Katzir et al., 1996; Soukis and
510 Papanikolaou, 2004; Martha et al., 2016; Lamont et al., 2020). Asymmetric lateral boundary
511 conditions are applied with 1 cm/yr on the left side and no displacement of the right side as in
512 Tirel et al. (2004). The upper surface is free and the base is driven by hydrostatic forces. No
513 prescribed discontinuity is introduced in the model, strain localization is only due to the use of
514 random noise in the cohesion value of the upper crust (for more details, see Huet et al., 2011).
515 The results shown here represent a case where the rheological stratification is inverted and the
516 thermal gradient is high, a likely situation in the Aegean or Tyrrhenian post-orogenic and back-
517 arc contexts.

518 The conceptual model of the interactions between the numerical model dynamics and the
519 intrusions is that we assume that a batch of mafic magmas, issued from partial melting of the
520 mantle, is injected at the base of the lower crust where it triggers the melting of felsic materials.
521 This leads to the formation of migmatites and collection of the felsic melts in a rising pluton
522 progressively caught in the detachment dynamics as it reaches the upper parts of the crust. The
523 felsic magma is thus deformed while it is still partly liquid and then mylonitized once it has
524 cooled down below the solidus. While extension proceeds, the overburden is removed by the
525 activity of the detachment and the molten material that comes next is injected in lower pressure
526 conditions and finds a faster access to the surface because of extension, thus leading to smaller
527 grainsize plutonic facies and probable volcanism at the surface. While the system of
528 detachments migrates toward the right and a new dome forms, the same situation can be
529 reproduced and a new pluton is emplaced below a detachment further to the right, closer to the
530 active detachment. This evolution is reminiscent of the evolution of Elba Island with the
531 formation of the Monte Capanne pluton in a first stage and the Porto Azzuro pluton in a later
532 stage.

533

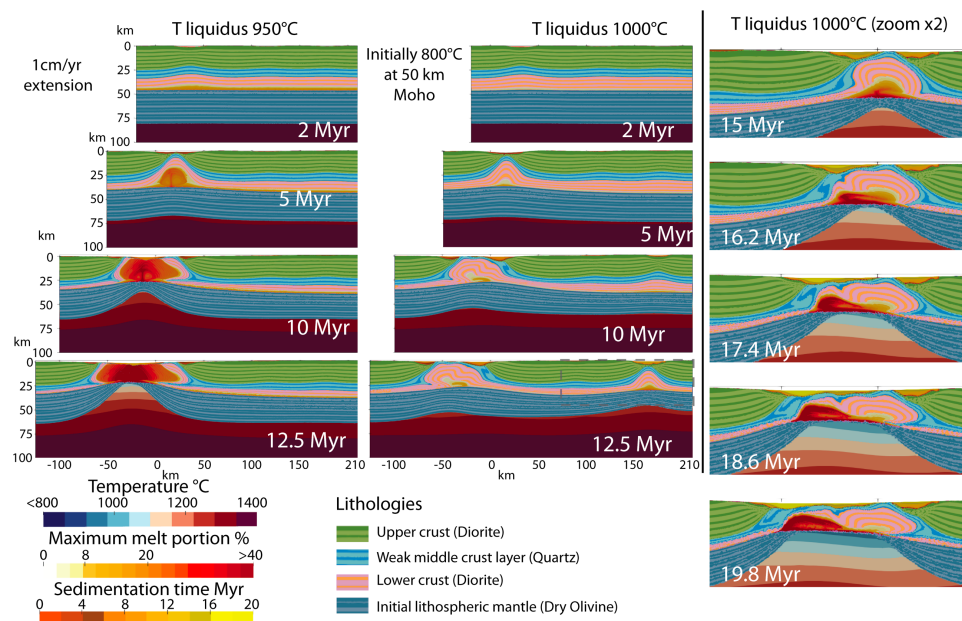


534 **5.3. Testing the concept with numerical experiment**

535

536 This conceptual model is now tested with new numerical experiments involving the
537 emplacement of magmas like in Schubert et al. (2013), but in a different situation where low-
538 angle detachments form, to see whether the introduction of a low-viscosity material in the
539 model developed by Huet et al. (2011) would drastically change the system dynamics or not.

540 This has been done for figures 14, 15 and 16.



541

542 *Figure 14: Snapshots of the numerical experiments with two different liquidus temperatures*
543 *(950°C and 1000°C) from 2 to 12.5 Myr. White line limits the molten lower crust with melt*
544 *ratio above 40% (pluton).*

545

546 The kinematics of exhumation produced by the nappe stacking experiments of Huet et al.
547 (2011) produces extension along long-lived detachment better resembling Mediterranean
548 example than diapiric spreading models that are produced by models with no intermediate weak
549 layers as in Tirel et al. (2004, 2008) or Rey et al. (2009). Hence, in order to test how molten
550 rocks interacts with detachments, we decided to build on our experience and start from this set
551 up which is a 210 km wide model domain submitted to 1cm/yr of extension on its left side for
552 10 Myr or more with an initial lithospheric column constituted from 25 km upper crust, 10 km
553 weak middle crust, 15 km thick lower crust overlying 40 km of lithospheric mantle. The Moho
554 located at 50 km depth is initially at a temperature of 830°C. We have taken the same



555 rheological parameters which are reported in Table A1 (Appendix A). The four major
556 differences with Huet et al. (2011) are :

557

- 558 i) Erosion and sedimentation applied on the top boundary,
559 ii) the deforming Wrinkler foundation at the LAB has been replaced by inflow of
560 asthenospheric material with higher thermal diffusivity to simulate small scale
561 convection and keep the base of the lithosphere at 1300°C during the experiments
562 as it was the case in Huet et al. (2011) study,
563 iii) the numerical code used in this study is pTatin2d (May et al., 2014, 2015) that solves
564 the same momentum equation

$$565 \quad \nabla \cdot \sigma = \rho g$$

566 For velocity v , as well as heat conservation

567

$$568 \quad -\nabla \cdot (-\kappa \nabla T + vT) + H = \frac{\partial T}{\partial t}$$

569

570 for Temperature T as Huet et al. (2011). However, it uses an incompressible visco-plastic
571 rheology minimizing the stress between a dislocation creep regime and Drucker Prager failure

572

$$573 \quad \begin{aligned} \nabla \cdot v &= 0 \\ \sigma &= \min \left(\sin \phi + 2 C \cos \phi, \dot{\epsilon}^{\frac{1}{n}} A \frac{-1}{n} e^{\frac{Q+VP}{nRT}} \right) \end{aligned}$$

574

575 to evaluate an effective viscosity :

$$576 \quad \eta_r = \frac{\sigma}{2\dot{\epsilon}}$$

577

578 instead of visco-elasto-plastic rheology based on dislocation creep and Mohr Coulomb failure
579 criteria.

580

- 581 iv) We added a simplified parametrization in order to account for the mechanical effect
582 of a melt in the simulations. The melt fraction M_f has been introduced as a linear
583 function of solidus (T_s) and liquidus (T_l) temperature

584

$$585 \quad M_f = \min \left(\max \left(\frac{T - T_s}{T_l - T_s}, 0 \right), 1 \right)$$



586

587 following Gerya and Yuen (2003). Based on melt fraction, the density and viscosity of passive
588 markers are modified following algebraic averaging for density

589

$$590 \quad \rho = M_f \rho_m + (1 - M_f) \rho_r,$$

591

592 and harmonic averaging for viscosity

593

$$594 \quad \eta = \left(\frac{M_f}{\eta_m} + \frac{1-M_f}{\eta_r} \right)^{-1}.$$

595

596 The solidus dependence on pressure P (in GPa) is implemented following wet granite
597 solidus of Miller et al. (2003), but we also added a variable temperature offset ΔT to account
598 for more mafic granitic composition as follows:

599

$$600 \quad T_s^c = 590 + \frac{250}{10(P + 0.1)} + \Delta T$$

601

602 and the dependence of liquidus to pressure is modeled following:

603

$$604 \quad T_l^c = T_s^c|_{P=0} + 10 + 200P.$$

605

606 The mantle is also allowed to melt following Hirshmann et al. (2000) solidus law

607

$$608 \quad T_s^m = -5.904P^2 + 139.44P + 1108.08$$

609

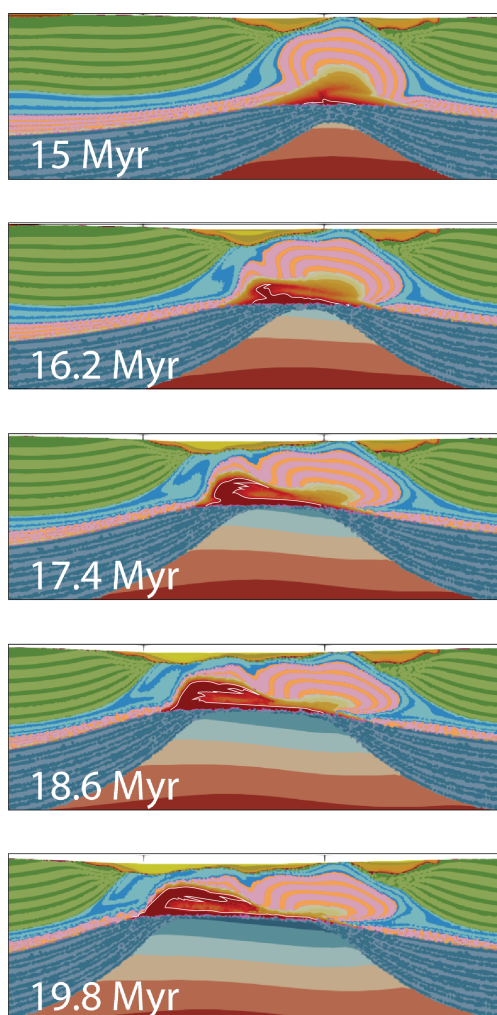
$$T_l^m = T_s^m + 600$$

610

611 All solidus and liquidus are represented on figure A2 (Appendix A) as a function of
612 pressure and temperature. For low melting temperature of the lower crust (wet granite solidus
613 with ΔT up to 100°C), the crust is largely molten in the initial conditions and buoyancy effects
614 dominate forming “spreading domes” in the classification of Huet et al. (2011), and this despite
615 the presence of a weak middle crustal layer. For higher melting temperature (ΔT from 150°C)
616 melt proportion remains below the 8% melt connectivity threshold described by Rosenberg
617 and Handy (2005) during most of the simulation. We interpret elements with markers that never



618 crossed that critical threshold as migmatites. In that case, melting does not disrupt the typical
619 asymmetric detachment kinematics observed in Elba and in the Cyclades that was well
620 reproduced by Huet et al. (2011) study. In the late stage of deformation, when the lithospheric
621 mantle is sufficiently attenuated by boudinage, the temperature of the lower crust reaches a
622 sufficient temperature to melt more generously, generating plutons which we define as markers
623 with a melt proportion greater than 40%.

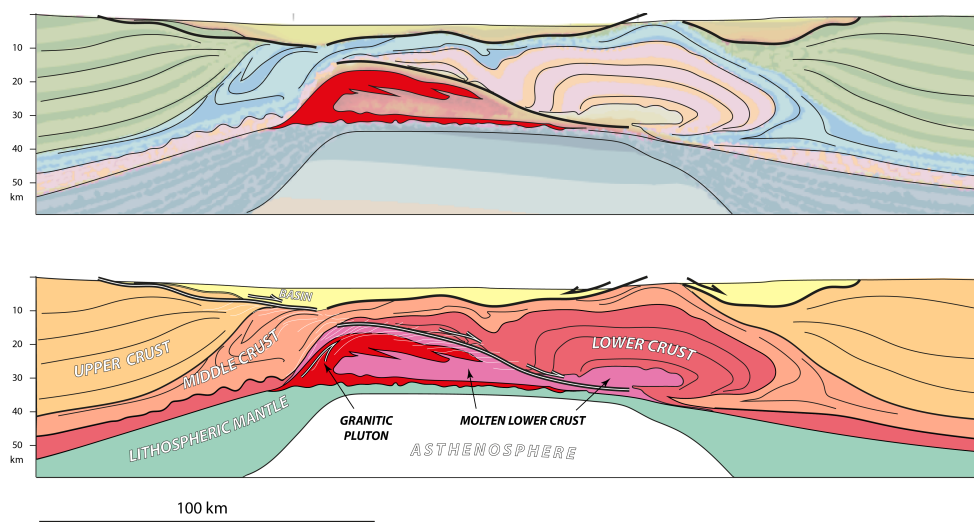


624
625 *Figure 15: Zoom of the last 5 Ma of evolution of the model of figure 15 with a liquidus*
626 *temperature of 1000°C from 15Myr to 19.8 Myr (black dotted rectangle on figure 15 for*
627 *location).*

628
629



630 With a temperature of 950°C at the surface for the liquidus, the molten layer is initially
631 thicker (figure 14) and strain develops into a spreading dome geometry like in the models of
632 Tirel et al. (2004, 2008) or Rey et al. (2009), with symmetric strain pattern and limited strain
633 localization. Shear is indeed progressively relocated on newly formed shear zones at the top
634 of newly exhumed hot material at the structure axis. When the temperature of the liquidus is
635 higher, reaching 1000°C in surface conditions (figure 14), the molten layer is initially thinner,
636 the deformation is persistently more localized on a detachment on one edge of the dome
637 structure and the model evolves with a detachment on the side of a dome with a limited rate of
638 partial melting in the lower crust (<8%). The deformation ultimately migrates to form a second
639 dome where a syn-kinematic low viscosity body develops with a melt ratio >40%, which we
640 interpret as analogous to a granite intrusion. The late evolution of this second dome shows a
641 strongly asymmetric geometry with the shape of the syn-kinematic intrusion controlled by the
642 asymmetry in strain pattern and a low-angle shear zone (figure 15). Figure 16 shows a structural
643 interpretation of the final step of the model for the second dome highlighting this asymmetry:
644 the dome is bounded by two antithetic crustal scale persistent shear zones, with steep and
645 shallow dip, and an eccentricity of the intrusion feeding pipe within the dome. This overall
646 asymmetrical strain and intrusion localization in the case of limited partial melting rate hence
647 reproduces the main features of the strain and intrusion pattern described in the field in the
648 Cyclades and Elba.



649
650 *Figure 16: Structural cross-section based on the most evolved stage of the numerical model at*
651 *19.8 Myr.*
652



653

654 Although the model does not show the details of the interactions between the dome and
655 the pluton, which involve percolation of melts, drainage through migmatites and dyke-swarms,
656 and progressive intrusion of the detachment by the rising granite, the overall geometry and
657 kinematics is similar with the natural case. The observed geometry is better reproduced in runs
658 with higher crustal melting temperatures and limited melt production. Low temperature melting
659 reactions for the continental crust are the wet solidus and the muscovite dehydration solidus,
660 while biotite and hornblende dehydration melting reactions could represent higher temperature
661 melting reactions (Weinberg and Hasalová, 2015). The model does not show either the role of
662 mafic injections at the base of the model. In the case of the Aegean and Monte Capanne, we
663 have postulated that mafic melts, generated by partial melting of the mantle in the arc and back-
664 arc region, intrude the lower crust and trigger the generation of felsic melts that then rise within
665 the dome. This is an additional input of heat in the model that would also localize the weakest
666 layers and thus the deformation and likely favor the evolution of the model in the same direction
667 as in the model presented here with a lower melting temperature. Similar evolution can arise
668 with an additional input of water in the lower crust. This water may originate from amphibole-
669 rich gabbros (sanukitoides) that could act as water donors enhancing lower crustal partial
670 melting to further produce secondary I-type granites (Castro, 2020). In the Aegean arc, whereas
671 amphibole bearing, I-type granites (Naxos and Serifos granodiorites among others) likely
672 reassemble secondary I-type granites as described by Castro et al. (2020 and references therein),
673 this water origin remains elusive as amphibole-bearing mantle magmas are not yet evidenced
674 interacting with the migmatites, except in the Mykonos-Delos-Rheneia MCC (Jolivet et al.,
675 2021).

676 The metamorphic parageneses associated with contact metamorphism in the case of Elba
677 suggests a depth of emplacement of the pluton of 4-5 km. In the numerical model, the genesis
678 of the pluton starts in the lower crust at a larger depth. The contact metamorphism observed in
679 the field characterizes the upper part of the pluton after it had risen within the crust, which may
680 explain this apparent contradiction.

681 One additional factor has not been considered in this study, the heterogeneity of the crust
682 inherited from earlier tectonic events. It has been shown by numerical experiments that dipping
683 heterogeneities in the crust mimicking structures inherited from nappe stacking help localizing
684 deformation, favoring the development of asymmetrical extensional structures and the
685 development of MCCs (Le Pourhiet et al., 2004; Huet et al., 2011; Lecomte et al., 2011; 2012).
686 Using this sort of initial conditions with melting would also favor the localization of



687 deformation on a single detachment. It may alternatively favor the development of several
688 asymmetric domes with low-angle detachments. Future studies should focus on testing such
689 initial conditions and also test these processes in 3-D.

690 The comparison of the Aegean and North Tyrrhenian plutons shows that the model of
691 interactions between plutons and detachments proposed by Rabillard et al. (2018) is
692 reproducible in similar contexts in different regions and can thus be probably generalized. The
693 comparison of this field-based conceptual model with numerical models moreover suggests that
694 this conceptual model is physically feasible. It requires the concomitance of post-orogenic
695 extension, thus extension set on an orogenic wedge, and a back-arc context to provide the
696 necessary heat and water for the generation of magmas in the mantle and the crust. The
697 possibility of a slab tear would be even more favorable as it increases the possibilities of
698 advecting hot asthenosphere directly below the extending crust (Roche et al., 2018).

699 As shown in the 3-D numerical experiments of Roche et al. (2018), slab retreat and back-
700 arc extension lead to the boudinage of the lithosphere with a spacing of ~100 km which will
701 then localize the formation of crustal domes and detachments. This type of evolution may
702 explain the formation of the lines of domes observed in the Aegean and the Menderes Massif.
703 Whether the boudinage also focuses the collection of mantle-derived magmas below the domes
704 is a question that should be addressed by further modelling. This question is important also
705 because the interactions between plutons and detachments described here and in Rabillard et
706 al. (2018) may provide guides for geothermal exploration. The case of the Menderes Massif is
707 exemplary of the intimate relations between active geothermal fields and crustal-scale
708 detachments (Roche et al., 2018). The case of the Tuscan Archipelago is partly similar with the
709 geothermal fields of Larderello and Monte Amiata developed above recent shallow plutons in
710 a context of asymmetric extension with top-to-the east low-angle shear zones (Jolivet et al.,
711 1998; Brogi et al., 2003; Rochira et al., 2018). The Monte Capanne and Porto Azzuro plutons
712 on Elba are associated with hydrothermal activities and mineralization (Maineri et al., 2003;
713 Rossetti and Tecce, 2008; Liotta et al., 2015) that make them good exhumed analogues of active
714 geothermal fields, a situation that is found also in the Cyclades with the mineralizations
715 observed on Mykonos or Serifos in the Cyclades (Salemink, 1985; St. Seymour et al., 2009;
716 Menant et al., 2013; Tombros et al., 2015; Ducoux et al., 2016).

717

718 **6. Conclusions**

719



720 The comparison between the Aegean and Tyrrhenian Miocene plutons shows striking
721 similarities in their interactions with coeval detachments. These plutons were all emplaced
722 underneath low-angle ductile shear zones and brittle detachments in a post-orogenic back-arc
723 environment where extra heat is provided by advected asthenospheric mantle. The roots of
724 those felsic plutons show a mixing with mafic magmas. The magmatic fabric is steep in the
725 vicinity of the roots and shallows toward the detachments. The plutons record substantial
726 stretching and shearing coaxial with the regional deformation while they still contain a
727 significant amount of melt. In sub-solidus conditions, the granitoids are then mylonitized when
728 approaching the detachment, until the formation of ultra-mylonites and pseudotachylytes. In
729 both cases too, the felsic magma intrudes the detachment and invades the upper plate. A
730 migration of detachments is then observed from the deep and ductile detachments to more
731 brittle and surficial ones. Late magmatic batches show a smaller grain size compatible with an
732 opening of the magma chamber toward the surface suggestive of the volcano-plutonic context
733 of these plutons. To account for these similarities suggesting that this model can be generalized
734 we proposed a conceptual model where mafic magmas batches are injected in the lower crust
735 of an extending orogenic wedge in a back-arc region with low-angle detachments. These mafic
736 injections trigger the melting of the lower crustal felsic material and the ascent of felsic plutons
737 in the crust, controlled by the low-angle detachments. The migration of the detachments through
738 time explains the migration of plutons and detachments observed in the Tuscan Archipelago
739 and in Tuscany from the Late Miocene to the Late Pliocene, as well as in the Aegean. This
740 conceptual model is tested with a numerical approach showing the impact of melt supply in the
741 development of the dome strain pattern. The observed asymmetry of strain localization and
742 intrusion are reproduced for a limited melting rate, while a higher melting rate would lead to
743 the development of a completely different dome structure. The geometry and kinematics
744 observed in the field are well reproduced by the model. These intimate interactions between
745 plutons and detachments can be foreseen as useful guides for the prospection and understanding
746 of geothermal and associated mineralization.

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748 Roche and Aurélien Rabillard did the field work in the Aegean, Loïc Labrousse and Laetitia Le
749 Pourhiet designed the modeling procedure and ran the numerical experiments. All authors
750 contributed to the writing of the manuscript.

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756

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1225
1226



1227 **Appendix A Table A1**

1228

| | | Sediments (2) | Upper crust (1) | Middle crust(2) | Lower crust(1) | Lithospheric mantle(3) | Asthenospheric mantle(3) |
|---|----------|------------------------|------------------------|----------------------|----------------------|---------------------------|-----------------------------|
| Density(kg.m ⁻³) | ρ_r | 2400 | 2700 | 2700 | 2700 | 3300 | 3300 |
| Pre exp. Factor (MPa ⁻ⁿ s ⁻¹) | A | 2.4 | 3.3 | 2.4 | 3.3 | 3.5 | 3.5 |
| Activation energy(kJ) | Q | 156 | 186 | 156 | 219 | 532 | 532 |
| Stress exponent | n | 6.7 10 ⁻⁶ | 2 10 ⁻⁶ | 6.7 10 ⁻⁶ | 1.3 10 ⁻³ | 2.5 10 ⁴ | 2.5 10 ⁴ |
| Melt density(kg.m ⁻³) | ρ_m | - | 2400 | 2400 | 2400 | 2800 | 2800 |
| Melt viscosity | η_m | - | 10 ¹⁷ | 10 ¹⁷ | 10 ¹⁷ | 10 ¹² | 10 ¹² |
| Heat production (Wm ⁻³) | H | 1.67 10 ⁻¹⁰ | 1.67 10 ⁻¹⁰ | 0 | 0 | 0 | 0 |
| Thermal diffusivity | κ | 10 ⁻⁶ | 10 ⁻⁶ | 10 ⁻⁶ | 10 ⁻⁶ | 10 ⁻⁶ | 5 10 ⁻⁶ |

1229

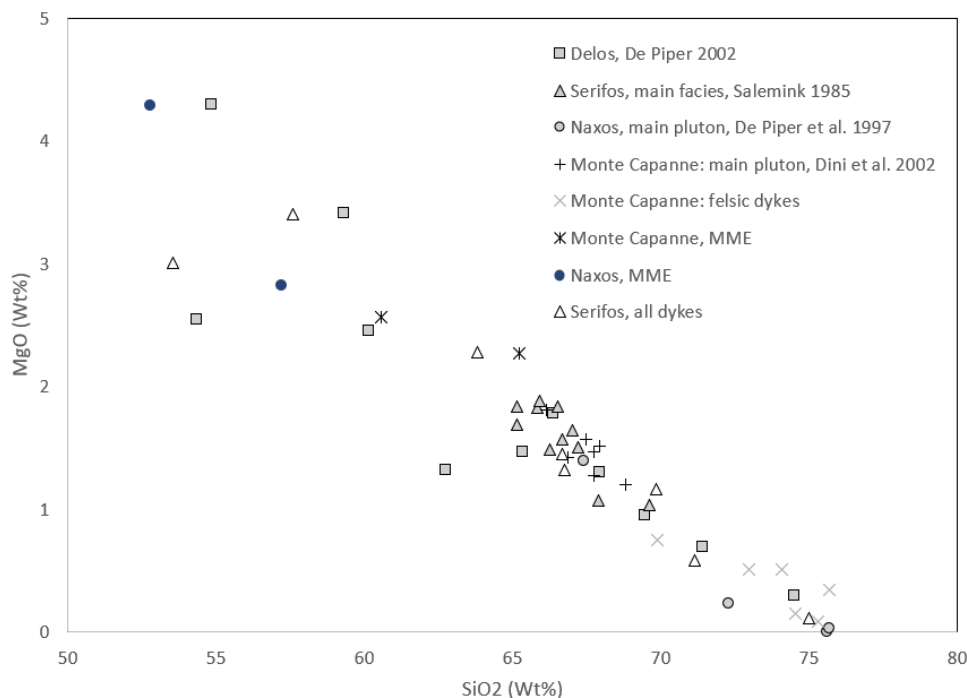
1230 *Table A1 : Values and notation for variable physical parameters. Parameters for dislocation*
 1231 *creep come from 1 (Ranalli and Murphy, 1987), 2 (Hansen and Carter, 1982) and 3 (Chopra*
 1232 *and Paterson, 1984). Other constant parameters include: friction ϕ and cohesion C which*
 1233 *varies linearly with plastic strain in range [0,1] respectively from 30° to 20° and from 20*
 1234 *MPas to 2 MPa; Coefficient of thermal expansion and compressibility that are set to 3.10⁻⁵K⁻¹*
 1235 *and 10⁻¹¹ Pa⁻¹.*

1236



1237 **Appendix A, figure A1 and A2:**

1238

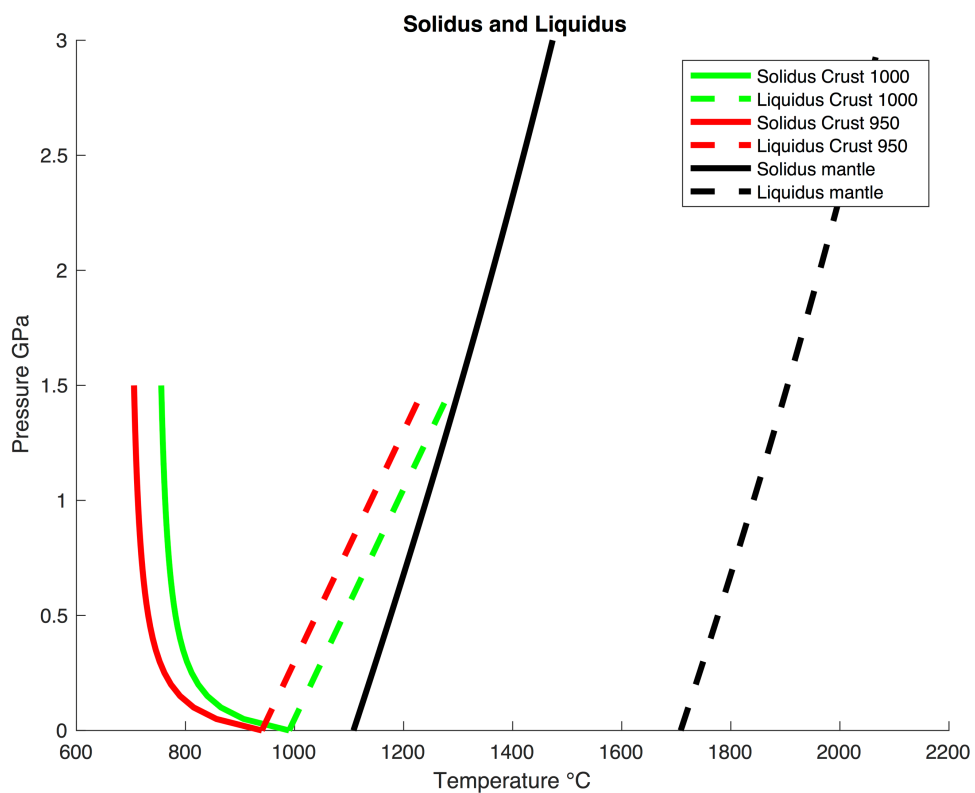


1239

1240

1241 *Figure A1: MgO v. SiO₂ Harker diagram showing the negative correlation between whole rock*
1242 *MgO and SiO₂ content for three, I-type hornblende-biotite bearing, representative*
1243 *Aegean granites. Data from Delos intrusion (Pe-Piper et al., 2002), Serifos (Salemink,*
1244 *1985) and Naxos (Pe-Piper et al., 1997). Mixing/mingling processes between mafic*
1245 *mantle-derived melts and acid magmas produce composite batholiths (Poli and*
1246 *Tommasini, 1990) as illustrated by the case of the Elba Island magmatic complex showed*
1247 *for comparison (See Dini et al. 2002 for explanation). MME = Mafic Microgranular*
1248 *Enclaves.*

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1251 *Figure A2: Solidus and liquidus as a function of pressure and temperature for mantle and*
1252 *crust (950 and 1000°C).*

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