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1 **Competition within low-density bacterial populations as an**  
2 **unexpected factor regulating carbon decomposition in bulk soil**

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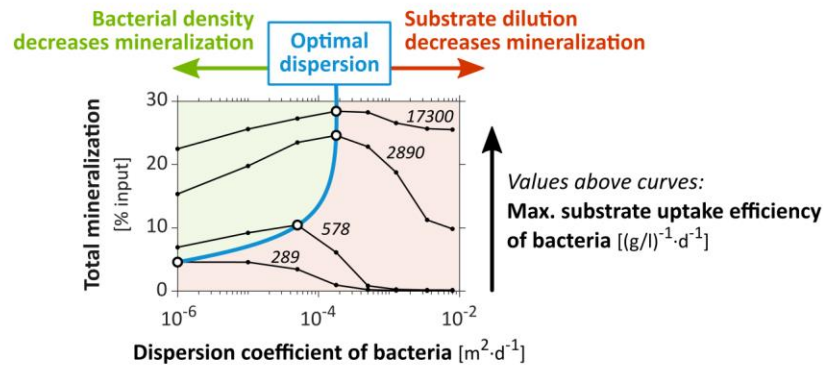
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15 **Abstract**

16 Bacterial decomposition of organic matter in soils is generally believed to be mainly  
17 controlled by the access bacteria have to organic substrate. The influence of bacterial traits  
18 on this control has, however, received little attention. Using the substrate-dependent  
19 Monod growth model, we develop a bioreactive transport model to screen the interactive  
20 impacts of spatial dispersion and bacterial traits on mineralization. Bacterial traits primarily

21 involved in the bacterial response to the substrate concentration, such as the maximum  
22 specific uptake rate and efficiency, the adaptation time of the uptake rate and the initial  
23 population density, are considered. We compare the model results with two sets of  
24 previously performed cm-scale soil-core experiments in which the mineralization of the  
25 pesticide 2,4-D was measured under well-controlled initial distributions and transport  
26 conditions. Bacterial dispersion away from the initial substrate location induced a significant  
27 increase in 2,4-D mineralization. It reveals an increase of specific uptake rates at lower  
28 bacterial densities, more than compensating the decrease of specific uptake rates caused by  
29 substrate dilution. This regulation of bacterial activities by density, caused by the local  
30 depletion of substrate by competing bacteria, becomes dominant for bacteria with an  
31 efficient uptake of substrate at low substrate concentrations (a common feature of  
32 oligotrophs). Such oligotrophs, commonly found in soils, compete with each other for  
33 substrate even at remarkably low population densities. The ratio-dependent Contois growth  
34 model, which includes a density regulation in the expression of the uptake efficiency, is more  
35 accurate and convenient to calibrate than the substrate-dependent Monod model, at least  
36 under these conditions. In view of their strong interactions, bioreactive and transport  
37 processes cannot be handled independently but should be integrated, in particular when  
38 reactive processes of interest are carried out by oligotrophs.

39 *Keywords:* biodegradation of organic matter; heterogeneous spatial distributions;  
40 bioreactive transport model; competition for substrate; bacterial traits; ratio-dependent  
41 growth



42

### 43 Highlights

- 44 - The impact of spatial distributions on decomposition depends on bacterial traits
- 45 - Decomposition can be reduced by competition between bacteria even at low densities
- 46 - Bacterial density regulation counterbalances substrate accessibility regulation
- 47 - Regulation of decomposition by bacterial density is more acute for oligotrophs

### 48 1. Introduction

49 Organic carbon is involved in most ecological functions provided by soils (Bünemann et al.,  
 50 2018). Its cycling in soil depends upon the activity of microorganisms. Soluble organic  
 51 molecules are taken up as substrates by specific populations of soil bacteria, and degraded  
 52 inside the cells by endoenzymes to provide carbon and energy. This is precisely the case for  
 53 the 2,4-Dichlorophenoxyacetic acid (2,4-D) used in this study as a generic model compound  
 54 (Don and Weightman, 1985; Pieper et al., 1988; Boivin et al., 2005). Bacterial degradation of  
 55 soil carbon has generally been modeled with the Monod equation, where the specific  
 56 substrate uptake rate is controlled by substrate concentration and bacterial traits such as  
 57 the maximum specific growth rate, the yield (or carbon use efficiency) and the “maximum

58 specific uptake efficiency” (e.g. Monod, 1949; Sinton et al., 1986; Cheyns et al., 2010). With  
59 the Monod equation, at the lowest substrate concentration, the specific uptake rate is  
60 linearly proportional to the substrate concentration. The proportionality factor is referred to  
61 here as the “maximum uptake efficiency” and it reflects the maximal ability of the cell to  
62 capture substrate molecules that collide with its membrane (Button, 1978, 1983). The  
63 maximum uptake efficiency can also be understood as the volume from which a cell can  
64 harvest substrate per unit of time, as used in some studies (Desmond-Le Quéméner and  
65 Bouchez, 2014; Nunan et al., 2020; Ugalde-Salas et al., 2020). Each bacterium is assumed to  
66 be exposed to the whole substrate concentration of its surroundings, without any limitation  
67 by the population density (Lobry and Harmand, 2006).

68 The direct contact (exposure) between bacteria and substrate depends on their spatial  
69 distributions (Holden and Firestone, 1997; Nunan et al., 2007). Bacteria and substrate are  
70 both heterogeneously distributed as a result of numerous biotic and abiotic processes  
71 (Dechesne et al., 2014; Kuzyakov and Blagodatskaya, 2015). There are complex feedback  
72 loops between these distributions, dispersive transport processes such as diffusion and  
73 hydrodynamic dispersion (Madsen and Alexander, 1982; Breitenbeck et al., 1988), and the  
74 bacterial activity itself such as consumption and growth (Poll et al., 2006).

75 Aggregated bacterial distributions, as observed at the mm-scale for 2,4-D degraders (Vieublé  
76 Gonod et al., 2003), have been shown to decrease degradation rates when the distribution  
77 of substrate is homogeneous, because of local substrate depletion (Pallud et al., 2004;  
78 Dechesne et al., 2010). Yet, the role of bacterial metabolic traits on the impact of bacterial  
79 and substrate distributions on degradation remains mostly unknown, especially when  
80 substrate and bacteria are heterogeneously and dynamically redistributed in soils over

81  $\mu\text{m}$ -to-cm scales by numerous spatial disturbances (Madsen and Alexander, 1982;  
82 Breitenbeck et al., 1988; König et al., 2020). We investigated the extent to which bacterial  
83 activity and transport processes can be treated independently or should be integrated to  
84 characterize, understand and predict degradation under various advective, diffusive and  
85 dispersive conditions. The simultaneous characterization of the impacts of bacterial traits  
86 and transport parameters through their mutual interactions is methodologically challenging.  
87 It requires several well-controlled experiments in comparable degradation conditions, with  
88 specific spatial distributions of substrate and degraders in specific transport conditions, and  
89 a spatiotemporal monitoring of the different carbon pools.

90 Among the scarce relevant datasets (e.g. Dechesne et al., 2010), we used the two sets of  
91 cm-scale soil-core experiments performed by Pinheiro et al. (2015, 2018), in which the  
92 degradation of 2,4-D under different initial spatial distributions and transport conditions was  
93 measured in similar repacked soil columns. Mostly reported independently, they have shown  
94 first that the proximity between bacteria and the initial location of a heterogeneously  
95 distributed substrate exerts a strong control on mineralization. Mineralization was greater  
96 when bacteria were close to the initial location of substrate, even though most of the initial  
97 dissolved substrate diffused away from its initial location. This was attributed to the fact that  
98 bacteria located far from the initial substrate location were only exposed to highly diluted  
99 substrate concentrations (Babey et al., 2017). However, the hydrodynamic dispersion of  
100 both bacteria and substrate away from their initial location caused a greater than four-fold  
101 increase in the mineralization of substrate that was not leached out, to the point that it  
102 almost reached the same performance as in homogeneous conditions in which there was no  
103 dilution (Pinheiro et al., 2018). The surprising increase in mineralization suggests a regulation

104 of mineralization by population density compensating the effect of substrate dilution, the  
105 activity of bacteria being enhanced when their density is diluted by the dispersive  
106 percolation events. While such regulations by bacterial density have not yet been considered  
107 in soils, presumably because of the extremely low apparent bacterial densities found in soils  
108 (Young et al., 2008), they are well known in bioreactors, where they are usually modeled by  
109 the ratio-dependent Contois growth law (Contois, 1959; Harmand and Godon, 2007).

110 In order to determine the relevance of the putative bacterial decomposer density effect on  
111 decomposition, we developed a quantitative approach to model the two sets of experiments  
112 within the same unified framework (section 2). We assessed the relevance of previously  
113 developed models, improved the calibration of a Monod-based model and investigated an  
114 alternative Contois-based model (section 3). We discuss the implication of the results on the  
115 controlling factors of soil organic carbon cycling, on the relevant bacterial growth models  
116 and on the possible bacterial strategies (section 4).

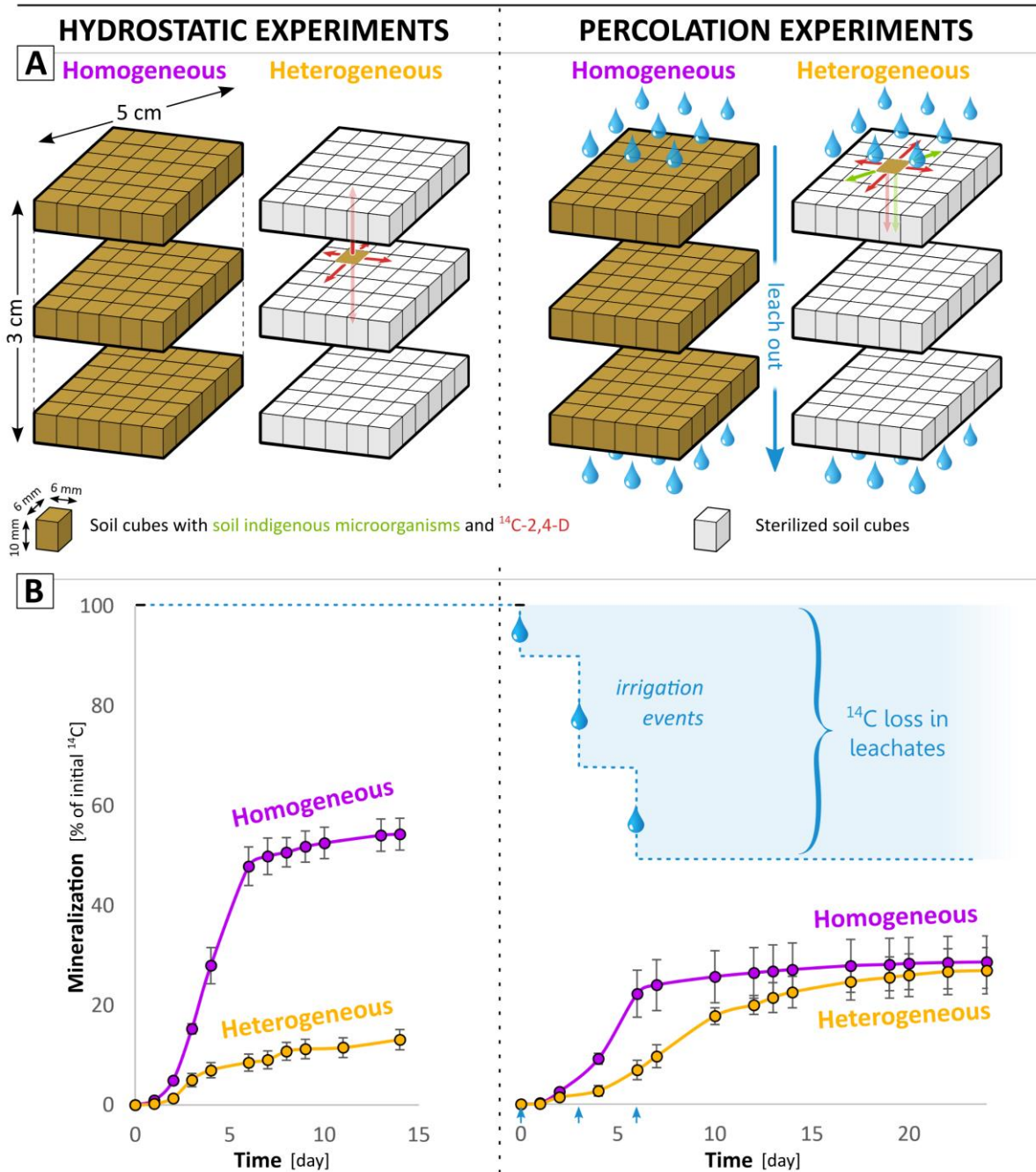
## 117 **2. Models and methods**

### 118 ***2.1. Experiment scheme, geometry and initial distributions***

119 We briefly introduce the experiments performed previously and highlight aspects of the  
120 experiments that are important for the modeling (Fig. 1). The full experimental setting is  
121 presented in the supplementary materials (Fig. S1 and Table S1) for the sake of  
122 completeness. Soil columns were packed with two homogeneous or heterogeneous  
123 arrangements of soil cubes, either sterilized, or hosting the indigenous microbial  
124 communities (referred to as “degraders”) and amended with <sup>14</sup>C-labelled 2,4-D (referred to

125 as “substrate”). Two sets of experiments, referred to as “hydrostatic” and “percolation”  
126 conditions, were performed respectively with only substrate diffusion (Pineiro et al., 2015),  
127 or with additional substrate and bacterial advection and dispersion caused by water  
128 percolation (Pineiro et al., 2018). The initial locations of the bacteria and substrate were  
129 set in the model according to the experimental conditions (Fig 1A). Initial concentrations  
130 used in the model are detailed in Table 1. In the experiments, the mass of mineralized  $^{14}\text{C}$   
131 derived from the degradation of the labelled 2,4-D was monitored at the core scale during  
132 at least two weeks (**Fig. 1B**). These data were used to confront the model processes with a  
133 physical system, as detailed in section 2.5.





134

135 **Fig. 1.** Model experimental design, geometry and initial distributions (A) based on previously  
 136 performed experiments in hydrostatic (Pineiro et al., 2015) and percolation (Pineiro et al.,  
 137 2018) conditions. The red and green arrows refer respectively to the 2,4-D and degrader  
 138 modeled displacements. (B) Experimental cumulated production of CO<sub>2</sub> (adapted from  
 139 Pineiro et al. (2015, 2018), permission for reproduction granted by Elsevier).

140 **2.2. Bioreactive model**

141 The bioreactive model extends the model published by Babey et al. (2017) (Fig. 2) to account  
 142 for Contois growth law as an alternative to Monod's. The sorption processes, the bacterial  
 143 lag phase and the biomass recycling described below were previously discussed and their  
 144 use justified in Babey et al. (2017) to consistently represent the experimental data. The  $r(\cdot)$   
 145 notation expresses the reaction rates of the biochemical dynamics that are expressed as  
 146 follows:

$$r(S) = k_{AS} A - k_{SA} S - k_R S - \frac{\mu}{y} B + m_t \chi B \quad (1)$$

$$r(A) = k_{SA} S - k_{AS} A \quad (2)$$

$$r(R_S) = k_R S \quad (3)$$

$$r(CO_2) = \frac{(1-y)}{y} \mu B \quad (4)$$

$$r(B) = \mu B - m_t B \quad (5)$$

$$r(R_B) = m_t (1 - \chi) B \quad (6)$$

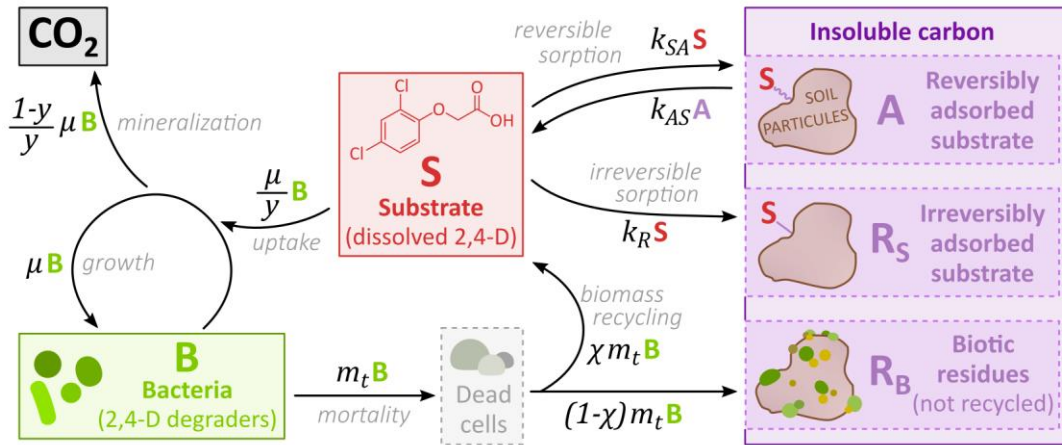
147 All variable and parameter definitions are listed in Table 1. The dynamics of the specific  
 148 growth rate  $\mu$  are given, for the Monod-based model, by:

$$\frac{\partial \mu}{\partial t} = \alpha \left( \mu_{max} \frac{S}{\kappa_M + S} - \mu \right) \quad (7)$$

149 and, for the Contois-based model, by:

$$\frac{\partial \mu}{\partial t} = \alpha \left( \mu_{max} \frac{S/B}{\kappa_C + S/B} - \mu \right) = \alpha \left( \mu_{max} \frac{S}{\kappa_C B + S} - \mu \right) \quad (8)$$

150 where  $\mu = 0$  at  $t = 0$ .



151  
 152 **Fig. 2.** Graphical representation of the biochemical model and carbon fluxes identified by the  
 153 arrows. Under low substrate concentrations  $S$ , the specific uptake rate  $(1/y) \cdot \mu$  becomes  
 154 equal to  $S \cdot (1/y) \cdot \mu_{max}/K_M$ , where  $(1/y) \cdot \mu_{max}/K_M$  is referred to as the “maximum specific  
 155 uptake efficiency”.

156 **Table 1.**

157 Values and range of values of the reactive transport model. The effective dispersion  
 158 coefficient  $d_{disp}$  applies only to heterogeneous percolation experiments.  $B(t = 0)$  is the initial  
 159 density of bacteria in the natural cubes. It is considered 1.6 times smaller in the percolation  
 160 experiments than in the hydrostatic experiments according to the initial experimental  
 161 measurements.

Parameter description	Symbol	Unit	Fixed values and admissible ranges for screening
Initial substrate concentration	$S(t = 0)$	$\mu\text{g} \cdot \text{g}^{-1}$ (mass of substrate carbon per mass of dry soil)	0.825 <sup>b</sup>
		$\mu\text{g} \cdot \text{g}^{-1}$	6.52 <sup>b</sup>
Reversible adsorption coefficient	$k_{SA}$	$\text{d}^{-1}$	0.09207
Reversible desorption coefficient	$k_{AS}$	$\text{d}^{-1}$	4.361
Irreversible adsorption coefficient	$k_C$	$\text{d}^{-1}$	0.01296

Uptake yield		$y$	-	0.5206
Maximum specific uptake rate		$(1/y) \cdot \mu_{max}$	$d^{-1}$	[0.0190 – 19.5]
Specific uptake efficiency at the lowest substrate concentration		$(1/y) \cdot \mu_{max} / \kappa^a$ <i>where <math>\kappa</math> is <math>\kappa_M</math> or <math>B(t=0) \cdot \kappa_C</math></i>	$g \cdot \mu g^{-1} \cdot d^{-1}$ (mass of dry soil per mass of bacterial carbon per unit of time)	[0.0152 – 159] <sup>c</sup>
Accommodation rate		$\alpha$	$d^{-1}$	[0.00934 – 934]
Initial degrader population density	Hydrostatic experiments	$B(t=0)$	$\mu g \cdot g^{-1}$ (mass of bacterial carbon per mass of dry soil)	[0.0161 – 1.61] <sup>d</sup>
	Percolation experiments	$B(t=0)$	$\mu g \cdot g^{-1}$	[0.0101 – 1.01] <sup>d</sup>
Mortality rate		$m_t$	$d^{-1}$	0.0602
Biomass recycling yield		$\chi$	-	0.6010
Effective diffusion coefficient		$d_{diff}$	$m^2 \cdot d^{-1}$	$1 \cdot 10^{-5}$ <sup>e</sup>
Effective dispersion coefficient		$d_{disp}$	$m^2 \cdot d^{-1}$	[0 – $\infty$ ]
Leaching rates (days 0; 3; 6)	Homogeneous experiments	$v$	-	0.108; 0.226; 0.180
	Heterogeneous experiments			0.107; 0.223; 0.178

162 <sup>a</sup> The half-saturation constant  $\kappa$  corresponds to  $\kappa_M$  for the Monod-based model and  
163  $B(t=0) \cdot \kappa_C$  for the Contois-based model (where  $B(t=0)$  is the value from the hydrostatic  
164 experiments).

165 <sup>b</sup> The initial substrate concentration  $S(t=0)$  is set equal to the <sup>14</sup>C-2,4-D concentration  
166 amended in the experiments.

167 <sup>c</sup> The values of  $(1/y) \cdot \mu_{max} / \kappa$  correspond to ranges of  $[1.65 - 1.73 \cdot 10^4]$   $l \cdot g^{-1} \cdot d^{-1}$  (volume of  
168 water per mass of bacteria per unit of time)

169 <sup>d</sup> The values of  $B(t=0)$  correspond respectively to ranges of  $[1.48 \cdot 10^{-4} - 1.48 \cdot 10^{-2}]$   $g \cdot l^{-1}$  (mass  
170 of bacteria per volume of water) for the hydrostatic experiments and  $[9.24 \cdot 10^{-5} - 9.24$   
171  $\cdot 10^{-3}]$   $g \cdot l^{-1}$  for the percolation experiments.

172 <sup>e</sup> The value of  $d_{diff}$  has been calibrated on a  $3 \times 6 \times 6$  grid in similar conditions (Babey et al.,  
173 2017).

174 The dissolved substrate  $S$  is either reversibly adsorbed to soil particles (pool  $A$ ) or irreversibly  
175 adsorbed (pool  $R_S$ ) (Eqs. (1), (2), (3)), or taken up by bacteria  $B$  (Eq. (1)) and metabolized into  
176  $CO_2$  (Eq. (4)) and new biomass  $B$  (Eq. (5)).  $k_{SA}$  and  $k_{AS}$  are the reversible sorption coefficients.  
177  $k_R$  is the irreversible one. Bacteria death occurs at a constant rate  $m_t$  (Eq. (5)) and a fraction  
178 of the bacterial necromass is considered to return to the dissolved substrate pool  $S$  to  
179 account for biomass recycling (Eq. (1)), while the rest is transformed to biotic residues  $R_B$   
180 (Eq. (6)). The remobilization of carbon previously absorbed by bacteria is necessary to  
181 adequately predict the slower dynamics of mineralization that takes place once most of the  
182 dissolved substrate has been consumed, observed after 5 days in homogeneous experiments  
183 and after respectively 3 or 10 days in the heterogeneous hydrostatic or percolation  
184 experiments. This remobilization is modeled in the form of a biomass recycling in order to  
185 be consistent with the model previously published in Babey et al. (2017), but similar effects  
186 could be achieved by other mechanisms, like a slower mineralization of biogenic residues  
187 (Fig. S9). Its impact on the final mineralization does not account for more than 10% of the  
188 substrate that is not leached out. The adsorbed substrate and biotic residues form the pool  
189 of insoluble carbon  $A + R_S + R_B$ . The substrate  $S$  is consumed by bacteria  $B$  according to their  
190 specific uptake rate  $(1/y) \cdot \mu$  expressed either by the substrate-dependent Monod growth law  
191 (Eq. (7)) (Monod, 1949) or by the ratio-dependent Contois growth law (Eq. (8)) (Contois,  
192 1959). The yield coefficient  $y$  relates the specific uptake rate  $(1/y) \cdot \mu$  to the specific growth  
193 rate  $\mu$ .  $\mu_{max}$  is the maximum specific growth rate.  $K_M$  and  $K_C$  are Monod and Contois constants  
194 respectively. The effective uptake is delayed by the accommodation rate  $\alpha$ , which explicitly  
195 takes into account the “memory” effects of the bacteria when adapting to new conditions  
196 (Patarinska et al., 2000). This delay is necessary to capture the mineralization lag time at the  
197 beginning of the experiments (see Fig. S10). Over long time periods ( $t \gg 1/\alpha$ ),  $\mu$  follows the

198 exact expression of the Monod or Contois equations. All modeled pools ( $S$ ,  $B$ ,  $CO_2$ ,  $A$ ,  $R_S$  and  
199  $R_B$ ) were expressed as carbon concentrations in  $\mu\text{g}\cdot\text{g}^{-1}$  (mass of carbon per mass of dry soil)  
200 considering a soil water content of  $0.205\text{ g}\cdot\text{g}^{-1}$  (mass of water per mass of dry soil), a bulk  
201 density of the soil column of  $1.3\ 10^3\text{ g}\cdot\text{l}^{-1}$  (mass of dry soil per apparent soil volume) and an  
202 average bacterial dry weight of  $2.8\ 10^{-13}\text{ g}$  corresponding to  $1.49\ 10^{-13}\text{ g}$  of carbon per cell.  
203 These values of water content and bulk density were those set up in the experiments, the  
204 latter corresponding to a water potential adjusted at  $-31.6\text{ kPa}$  (pF 2.5). The average bacterial  
205 weight was assumed based on Dechesne et al. (2010) and Pinheiro et al. (2015). The water-  
206 filled pore space (54%, volume of water per volume of pores) was such that oxygen was not  
207 considered a limiting factor for 2,4-D degradation.

### 208 ***2.3. Reactive transport model***

209 The transport model is based on the substrate diffusion model of Babey et al. (2017) to which  
210 advective-dispersive processes explored in the experiments of Pinheiro et al. (2018) are  
211 added. Hydrodynamic leaching out and dispersion of bacteria and substrate were observed  
212 in the percolation experiments while the substrate was reported to diffuse in all  
213 experiments. Hydrodynamic leaching and dispersion were modeled independently, as they  
214 result from, respectively, bypass flow through large pores and complex hydrodynamic  
215 dispersion processes coming not only from usual flow mechanisms but also from large  
216 saturation variations and local redistribution of moisture in the pore network. Due to the  
217 lack of adequate experimental data to characterize the details of the dispersion process, we  
218 applied a simple isotropic dispersion coefficient. Complementary numerical simulations  
219 show that other anisotropic dispersion parameterizations are only weakly sensitive (Fig. S3

220 and S4). Bacterial and substrate transports by percolation were described with the same  
 221 advective and dispersive parameters. This assumption did not significantly alter the results  
 222 (Fig. S5 and S6). Coupled to the equations of the bioreactive model ((1)-(8)), the full reactive  
 223 transport model is given by:

$$\frac{\partial S}{\partial t} = r(S) + \nabla(d_{diff} \nabla S) + G(\nabla(d_{disp} \nabla S) - \nu S) \quad (9)$$

$$\frac{\partial B}{\partial t} = r(B) + G(\nabla(d_{disp} \nabla B) - \nu B) \quad (10)$$

$$\frac{\partial U}{\partial t} = r(U) \quad \text{for } U = A, R_B, R_S \text{ and } CO_2 \quad (11)$$

224 where  $d_{diff}$  is the effective molecular diffusion coefficient of  $S$ ,  $d_{disp}$  is the effective  
 225 hydrodynamic dispersion coefficient of  $S$  and  $B$  and  $\nu$  is their leaching rate. Note that the  
 226 dispersion coefficient  $d_{disp}$  mostly affected the spreading of bacteria, given that substrate  
 227 was mainly spread by diffusion, as confirmed by consistent results from equivalent models  
 228 without hydrodynamic dispersion of  $S$  (Fig. S7 and S8). Effective diffusion and dispersion  
 229 processes were assumed to be isotropic and uniform at the column-scale. Dispersion and  
 230 leaching were active only during the observed 1-hour percolation events at days 0, 3 and 6  
 231 as controlled by the function  $G$  defined as:

$$G(t) = 1 \quad \text{for } t = [0d - 0d1h]; [3d - 3d1h]; [6d - 6d1h] \quad (12)$$

$$G(t) = 0 \quad \text{otherwise.}$$

232 No-flow boundary conditions were imposed at the edges of the soil core ( $\nabla S = 0$  and  $\nabla B = 0$ )  
 233 during periods outside of the percolation events. The transient evolutions of the water  
 234 content and their effects on concentrations were not considered because of the short

235 duration of the percolation events (1 h) and the absence of detectable effects on the  
236 experimental mineralization curve around the percolation events (Fig. 1D). Hydration  
237 conditions were considered constant, constrained by the water potential adjusted  
238 to -31.6 kPa. No bacterial mobility was observed in the hydrostatic experiments, suggesting  
239 that the bacterial mobility observed in the percolation experiments resulted primarily from  
240 hydrodynamic dispersion.

241 Carbon pools concentration dynamics were simulated on a  $3 \times 6 \times 6$  regular mesh grid.  
242 Although the shape of the grid was slightly different from that of the cylindrical soil-core, it  
243 did not have any observable impact (Babey et al., 2017). We recall that substrate and  
244 bacteria were initially co-located in the same cube(s). Each cube was considered to be  
245 physically, chemically and biologically homogeneous. Diffusion and dispersion were  
246 simulated using a finite-difference scheme (Iserles, 2009) and coupled with the bioreactive  
247 model, itself solved by the 4<sup>th</sup> order Runge-Kutta integration method function of MATLAB  
248 (Shampine and Reichelt, 1997). The coupling of transport and bioreactive models was  
249 achieved with a sequential non-iterative operator-splitting method, in which the equations  
250 are resolved within each time step in a sequence of one transport step followed by one  
251 bioreactive step (Carrayrou et al., 2004; Lagneau and van der Lee, 2010). The time steps were  
252 smaller than the characteristic diffusion and reaction times to avoid any coupling issues.

#### 253 ***2.4. Exploratory screening***

254 Parameters and their values are listed in Table 1. Sorption parameters and the diffusion  
255 coefficient were set at values that were calibrated and validated by Babey et al. (2017) in  
256 independent experiments without degradation. The mortality rate and the biomass recycling



257 yield were also kept at the values calibrated in Babey et al. (2017) as they were considered  
258 to be well constrained by the residual mineralization dynamics of the homogeneous  
259 hydrostatic experiment (Fig. 1D). The four biological parameters primarily involved in the  
260 biological response of bacteria to the concentration of substrate were determined to be  
261  $(1/y) \cdot \mu_{max}$ ,  $\alpha$ ,  $B(t=0)$  and either  $(1/y) \cdot \mu_{max}/\kappa_M$  for the Monod-based model or  
262  $(1/y) \cdot \mu_{max}/(B(t=0) \cdot \kappa_C)$  for the Contois-based model. Each of these four parameters were  
263 sampled over 7 logarithmically-distributed values within the theoretically and physically  
264 relevant ranges given by Babey et al. (2017), and all possible combinations of values were  
265 screened (Table S2). We recall that the maximum specific uptake efficiency  $(1/y) \cdot \mu_{max}/\kappa_M$   
266 characterizes the bacterial uptake of substrate at the lowest substrate concentration  
267 (Button, 1991), while the maximum specific uptake rate  $(1/y) \cdot \mu_{max}$  characterizes the  
268 bacterial uptake at the highest substrate concentration. Note that the uptake yield  $y$  was  
269 fixed at the value calibrated by Babey et al. (2017) with a high degree of certainty. The initial  
270 maximum uptake efficiency  $(1/y) \cdot \mu_{max}/(B(t=0) \cdot \kappa_C)$  in the Contois-based model was  
271 screened in the same range as  $(1/y) \cdot \mu_{max}/\kappa_M$ . The accommodation rate  $\alpha$  of the degrader  
272 response ranged from a negligible delay of few minutes ( $\alpha = 934 \text{ d}^{-1}$ ) to a prolonged delay of  
273 around 10 days ( $\alpha = 9.34 \cdot 10^{-2} \text{ d}^{-1}$ ).  $B(t=0)$  values were screened around the initial  
274 experimental measurements of the *tfdA* gene copy number, assuming that one *tfdA*  
275 sequence corresponded to one bacterium. They ranged over two orders of magnitude to  
276 account for the uncertainty of the conversion of *tfdA* copy number into alive 2,4-D degraders  
277 (Bælum et al., 2006, 2008). Bacterial density in the uptake efficiency expression will also be  
278 expressed in  $\text{g} \cdot \text{l}^{-1}$  (mass of bacteria per volume of water) for a more direct comparison with  
279 the relevant literature.

280 The spatial distribution of bacteria observed at the end of the experiments could not be used  
 281 to determine the effective dispersion coefficient  $d_{disp}$  (Fig. S2). While they qualitatively  
 282 ascertained that bacteria spread orthogonally to the percolation direction, experimental  
 283 data were not sufficiently resolved to be used quantitatively. The dispersion coefficient was  
 284 thus screened over 10 values ranging from no dispersion ( $d_{disp} = 0$ ) to complete instant  
 285 homogenization of the soil core ( $d_{disp} = \text{inf}$ ) (Table S2). In order to analyze the result of  
 286 bacterial dispersion in terms of distance from the initial location of the substrate, we  
 287 compute the root-mean-square displacement of bacteria, defined as the root-mean-square  
 288 of their spreading during the duration  $\Delta t$  of one percolation event and expressed as  
 289  $\sqrt{6 d_{disp} \Delta t}$  (Stana, 2020). The effective diffusion coefficient  $d_{diff}$  had been calibrated  
 290 independently from percolation conditions (Pineiro et al., 2015; Babey et al., 2017). The  
 291 leaching rates  $v$  were determined based on the experimental masses of leached  $^{14}\text{C}$  (Pineiro  
 292 et al., 2018) (Table 1). Detailed values for the screened parameters are listed in Table S2.

### 293 **2.5. Model to data comparison**

294 The comparison between the results of the model and the experimental data was based on  
 295 the core-scale data of mineralization deduced from the carbon mass  $m_{CO_2}$  of  $^{14}\text{CO}_2$  emissions:

$$m_{CO_2}(t) = \int_V CO_2(x, t) dx \quad (13)$$

296 with  $V$  the volume of the soil cores. Mineralization at a given time  $t$  was expressed as the  
 297 carbon mass of cumulated  $^{14}\text{CO}_2$  emissions ( $m_{CO_2,q}(t)$ ) per initial carbon mass of  
 298  $^{14}\text{C}$ -substrate  $S$  ( $m_{S,q}(t = 0)$ ) where the index  $q$  identifies the experiment at hand. Indices

299 1, 2, 3 and 4 are respectively given to the homogeneous hydrostatic, heterogeneous  
 300 hydrostatic, homogeneous percolation and heterogeneous percolation experiments.  
 301 Data-to-model adequacy was assessed for each of the experiments by a classical root-mean-  
 302 square evaluation function  $J_q$  comparing the modeled mineralization of Eq. (4) to the  
 303 measured mineralization at the  $n_q$  available sampling times  $t_i$ :

$$J_q = \left( \frac{1}{n_q} \sum_{i=1}^{n_q} \left( \frac{m_{CO_2,q}^{mod}(t_i) - m_{CO_2,q}^{data}(t_i)}{m_{S,q}(t=0)} \right)^2 \right)^{\frac{1}{2}} \quad (14)$$

304 Discrepancies over the full set of experiments  $J_{1234}$  were thus expressed as:

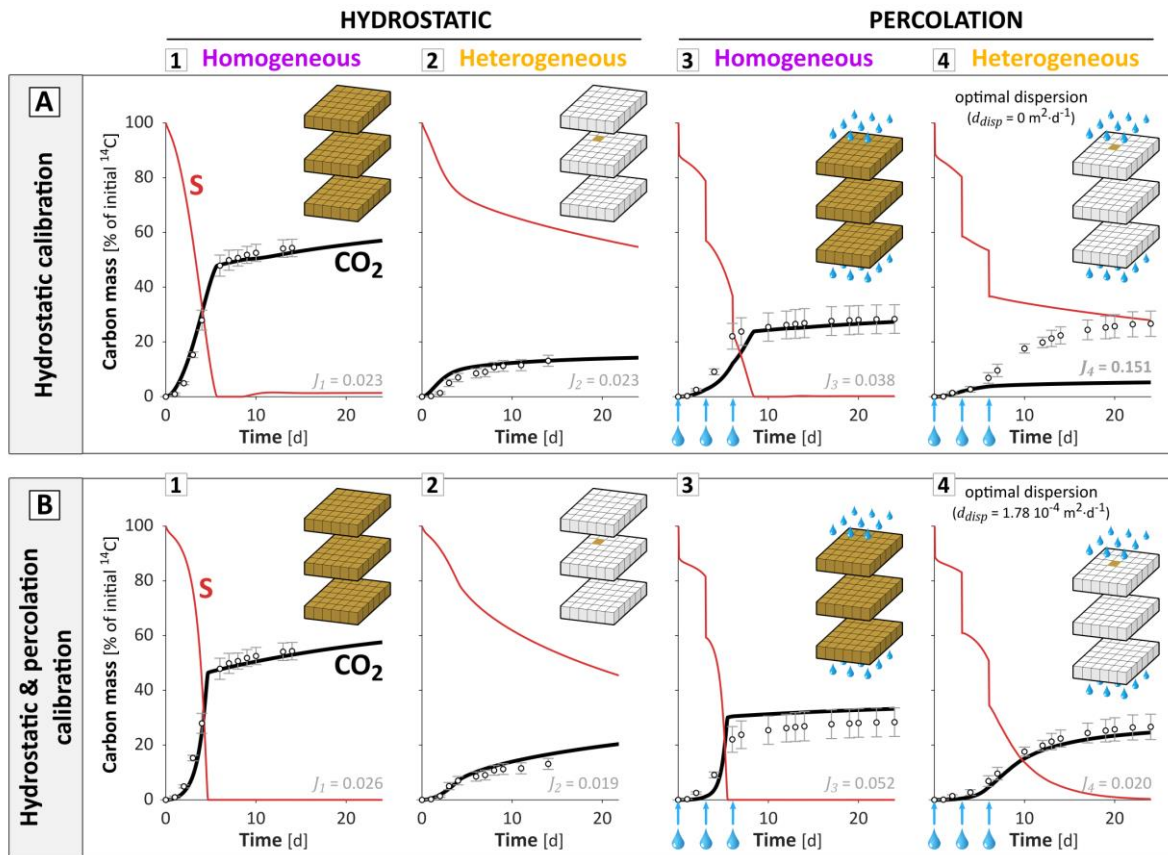
$$J_{1234} = \left( \frac{1}{4} \sum_{k=1}^4 J_k^2 \right)^{\frac{1}{2}} \quad (15)$$

305 Following the systematic parameter screening described in section 2.4, the parameter set  
 306 minimizing  $J_{1234}$  was determined and referred to as the set calibrated on both hydrostatic  
 307 and percolation experiments. The measurement errors were in average 1.7 times higher in  
 308 the percolation experiments than in the hydrostatic experiments. This was assumed to be  
 309 due to differences in experimental setup between the two sets of experiments of Pinheiro  
 310 et al. (2015, 2018). This error difference contributed to limit the weight of the percolation  
 311 experiments when determining the best-fitting parameter set over the whole set of  
 312 experiments ( $J_{1234}$ ). We made the choice to give an equal weight to all experiments by only  
 313 taking into account the average CO<sub>2</sub> values.

## 314 **3. Results**

### 315 **3.1. Model calibration**

316 The calibration of the bioreactive transport model carried out using only the hydrostatic  
317 experimental data (Babey et al., 2017) led to a minimal discrepancy between data and model  
318 of  $J_{12} = 0.023$  (Fig. 3-A1 and A2). This pre-existing parameterization was used to provide blind  
319 predictions of the percolation experiments, with the effective dispersion coefficient  $d_{disp}$  as  
320 an additional fitting parameter. It gave a reasonable prediction of mineralization in the  
321 homogeneous percolation experiment ( $J_3 = 0.038$ , Fig. 3-A3) but failed in the heterogeneous  
322 percolation experiment ( $J_4 = 0.151$ , Fig. 3-A4), regardless of the dispersion coefficient values.  
323 The smallest discrepancy  $J_4$  was surprisingly obtained without any bacterial dispersion  
324 ( $d_{disp} = 0$ ) in contradiction with the bacterial spread observed in the experimental data  
325 (Fig. S2). The final predicted mineralization was highest when bacteria remained aggregated  
326 close to the initial location of the substrate. The highest predicted mineralization was  
327 however four times lower than the experimental data. The large gap between the  
328 experimental data and the modeled scenario suggests that bacterial proximity to the initial  
329 substrate location is not the underlying explanatory mechanism for the high mineralization  
330 rates. On the contrary, it suggests that mineralization might rather be increased by the  
331 dispersion of bacteria towards more diluted substrate concentrations, and that the identified  
332 bacterial traits do not match this increase of mineralization with dispersion.



333

334

**Fig. 3.** Mineralization dynamics predicted with the Monod-based model calibrated on the hydrostatic experiment only (A) and on both hydrostatic and percolation experiments (B).

335

The related experimental setups are indicated in the top right corner of each graph. The

336

agreement between experiments and model is indicated by the value of discrepancy  $J$

337

displayed at the bottom and can be visually assessed by the proximity between the black line

338

and the dots representing respectively the model results and experimental data. The red line

339

refers to the carbon mass of substrate remaining in the soil core. In the percolation

340

experiments (A3,4 and B3,4), around 51% of the initial mass of  $^{14}\text{C}$  was lost through leaching

341

after the three percolation events ( $t = 0, 3$  and  $6$  days, blue arrows). The carbon balance

342

among the different pools is detailed in Fig. S11. Note that the reversible sorption eventually

343

accounted for less than 2% of the initial carbon mass and therefore did not significantly alter

344

the results.

345

346

In order to investigate the capacity of the reactive transport model to fit both hydrostatic

347

and percolation experimental data, the biological parameters ( $(1/y) \cdot \mu_{max} / \kappa_M$ ,  $(1/y) \cdot \mu_{max}$ ,  $\alpha$ ,

348

$B(t=0)$ ) and the dispersion coefficient ( $d_{disp}$ ) were calibrated on both hydrostatic and

349 percolation experiments following the screening approach given in section 2.4 to minimize  
350  $J_{1234}$ . The mineralization dynamics were adequately predicted in all four experiments with  
351 the biological parameter set giving the lowest overall discrepancy ( $J_{1234} = 0.032$ ) and a  
352 non-zero dispersion coefficient ( $d_{disp} = 1.78 \cdot 10^{-4} \text{ m}^2 \cdot \text{d}^{-1}$ ) (Fig. 3 and Table 2). The non-zero  
353 dispersion coefficient indicates that the calibrated model accounts for a positive impact of  
354 bacterial dispersion on degradation. The model results suggest that this effect is necessary  
355 to successfully predict the high degree of degradation in the experimental data. Compared  
356 to the parameters calibrated only using the hydrostatic experiments, the parameter set  
357 calibrated on both hydrostatic and percolation experiments also displayed a much higher  
358 maximum uptake efficiency ( $(1/y) \cdot \mu_{max} / \kappa_M = 26.5 \text{ g} \cdot \mu\text{g}^{-1} \cdot \text{d}^{-1}$  (mass of dry soil per mass of  
359 bacterial carbon per unit of time) (Table 2). The systematic exploration of the parameter  
360 space showed that high maximum uptake efficiency was a common feature of the 1%  
361 best-fitting parameterizations over both hydrostatic and percolation experiments (smallest  
362  $J_{1234}$ ), with values of 159 and  $26.5 \text{ g} \cdot \mu\text{g}^{-1} \cdot \text{d}^{-1}$ , corresponding respectively to  $1.73 \cdot 10^4$  and  
363  $2.89 \cdot 10^3 \text{ l} \cdot \text{g}^{-1} \cdot \text{d}^{-1}$  (volume of water per mass of bacteria per unit of time). It underlines the  
364 essential role of the maximum uptake efficiency for modulating the impact of dispersion on  
365 degradation, further detailed and explained in section 3.2.3.

366 **Table 2.**

367 Parameters for the Monod-based model calibrated by the screening approach (section 2.2)  
368 on the hydrostatic experiments only (Babey et al., 2017) and on both hydrostatic and  
369 percolation experiments, and for the Contois-based model calibrated on both hydrostatic  
370 and percolation experiments, as described in section 2.4

Parameter symbol	Unit	Monod model calibration		Contois model calibration	
		on the sole hydrostatic experiments	on both hydrostatic & percolation experiments	on both hydrostatic & percolation experiments	
$(1/y) \cdot \mu_{max}$	d <sup>-1</sup>	1.22	9.73	4.86	
$(1/y) \cdot \mu_{max}/\kappa^a$	g·μg <sup>-1</sup> ·d <sup>-1</sup> (mass of dry soil per mass of bacterial carbon per unit of time)	2.65 <sup>b</sup>	26.5 <sup>b</sup>	2.65 <sup>b</sup>	
$\alpha$	d <sup>-1</sup>	9.341 10 <sup>-1</sup>	9.34 10 <sup>-2</sup>	9.34 10 <sup>-2</sup>	
$B(t = 0)$	Hydrostatic experiments	μg·g <sup>-1</sup> (mass of bacterial carbon per mass of dry soil)	1.61 10 <sup>-1</sup>	3.23 10 <sup>-2</sup>	3.76 10 <sup>-1</sup>
	Percolation experiments	μg·g <sup>-1</sup>	1.01 10 <sup>-1</sup>	2.01 10 <sup>-2</sup>	2.34 10 <sup>-1</sup>
$d_{disp}$	m <sup>2</sup> ·d <sup>-1</sup>	0 <sup>c</sup>	1.78 10 <sup>-4 c</sup>	10 <sup>-5 c</sup>	
$J_{1234}$	-	0.079	0.032	0.022	

371 <sup>a</sup> The half-saturation constant  $\kappa$  corresponds to  $\kappa_M$  for the Monod-based model and  
372  $B(t = 0) \cdot \kappa_C$  for the Contois-based model (where  $B(t = 0)$  is the value from the hydrostatic  
373 experiments).

374 <sup>b</sup> Values of  $(1/y) \cdot \mu_{max}/\kappa$  correspond respectively to 2.89 10<sup>2</sup>, 2.89 10<sup>3</sup> and 2.89 10<sup>2</sup> l·g<sup>-1</sup>·d<sup>-1</sup>  
375 (volume of water per mass of bacteria per unit of time).

376 <sup>c</sup> The corresponding spreading values induced by the hydrodynamic dispersion  
377 (root-mean-square displacements) for each percolation events are respectively 0, 6.7 and  
378 1.6 mm, to be compared to the 25 mm radius of the soil column.

### 379 **3.2. Analysis of the controls exerted on degradation by substrate dilution and** 380 **bacterial density**

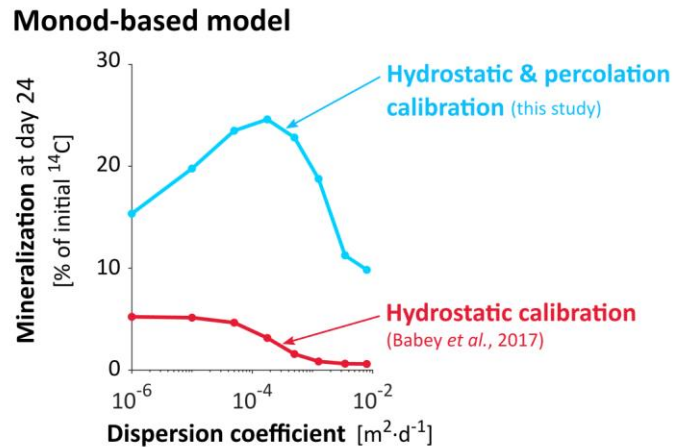
381 The effect of dispersion on degradation differed greatly between the two calibrated sets of  
382 biological parameters described in section 3.1. We therefore conducted a more systematic

383 investigation of the coupled impact of bacterial dispersion and bacterial traits on  
384 degradation, revealing its control by substrate dilution and bacterial density.

### 385 3.2.1 Impact of dispersion on degradation

386 We used the mineralization at the end of the experimental time (day 24) as a proxy for  
387 degradation and determined its sensitivity to dispersion, as a function of the  
388 parameterization of bacterial traits. Fig. 4 shows the impact of the dispersion coefficient  $d_{disp}$   
389 on the final predicted mineralization for the two calibrated biological parameter sets, all  
390 other parameters being kept constant (thick red and blue lines). For the biological parameter  
391 set calibrated on hydrostatic experiments, the final mineralization decreased monotonically  
392 with dispersion (Fig. 4, red line). For the parameter set calibrated on both hydrostatic and  
393 percolation experiments, the final mineralization first increased, reached a maximum around  
394  $d_{disp} \approx 10^{-4} \text{ m}^2 \cdot \text{d}^{-1}$  and then decreased (Fig. 4, blue line). These two kinds of behaviors were  
395 observed regardless of the parameters  $\alpha$ ,  $(1/y) \cdot \mu_{max}$  and  $B(t=0)$  as long as  $(1/y) \cdot \mu_{max} / \kappa M$   
396 remained the same (Fig. S12). The non-monotonic impact of dispersion on degradation  
397 highlights the existence of an optimal bacterial dispersion for which mineralization is the  
398 highest. The comparison between the red and blue lines on Fig. 4 suggests that the optimal  
399 dispersion value depends on the bacterial uptake efficiency. Note that, although the optimal  
400 dispersion value varied with time due to the spatial dynamics of both bacteria and substrate  
401 (Fig. S14), it tended towards a limit that was mostly reached within 4 to 7 days and is thus  
402 represented at day 24 on Fig. 4.





403  
 404 **Fig. 4.** Influence of the dispersion coefficient  $d_{disp}$  on mineralization predicted at day 24  
 405  $m_{CO_2}(t = 24)$  for the biological parameter set calibrated on the sole hydrostatic experiments  
 406 (A, thick red line) and on both hydrostatic and percolation experiments (B, thick blue line).  
 407 Note that for the model calibrated on both hydrostatic and percolation experiments, the  
 408 value of  $d_{disp}$  leading to the highest final mineralization ( $d_{disp} = 1.78 \cdot 10^{-4} m^2 \cdot d^{-1}$ , thick blue  
 409 line) is also equal to its calibrated value leading to the best adequacy with mineralization  
 410 kinetics (Table 2). Note that the optimal dispersion value remains the same when  
 411 representing the remaining dissolved substrate instead of the mineralization (Fig. S13).

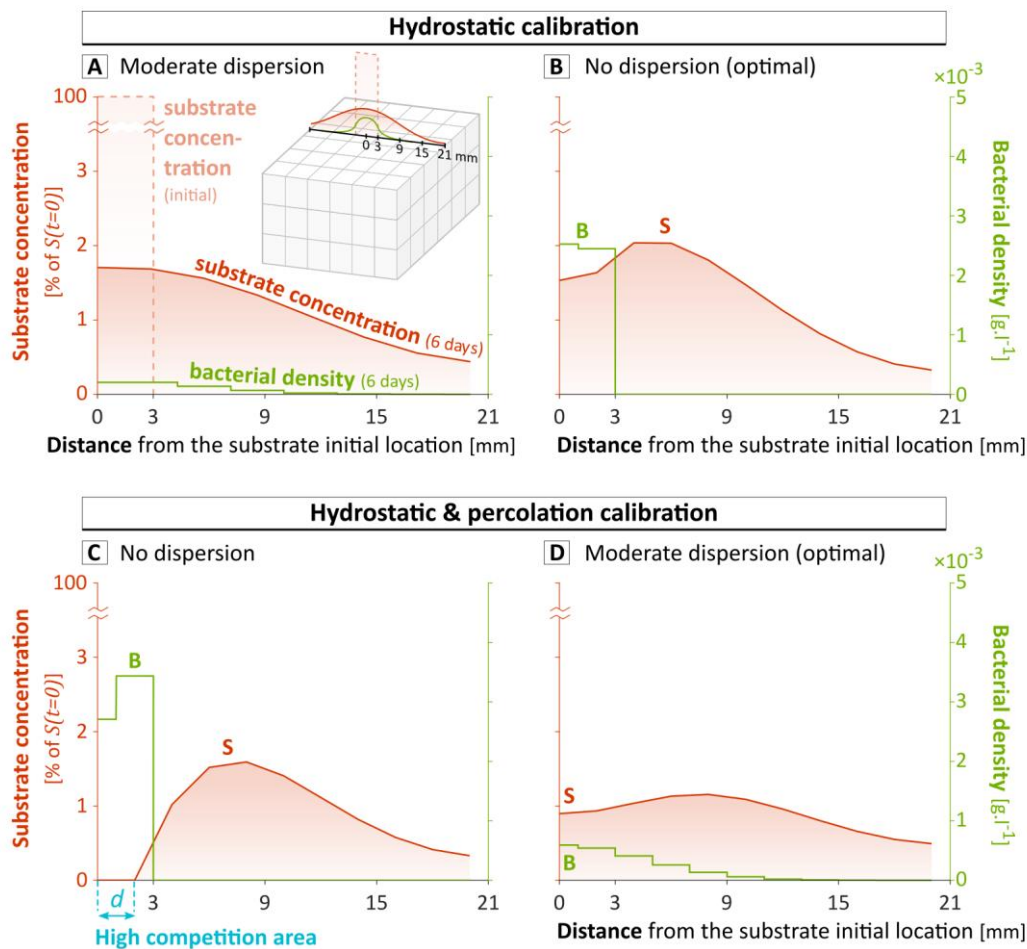
### 412 3.2.2 Double control of degradation by substrate dilution and bacterial density

413 The non-monotonic effect of bacterial dispersion on degradation is an unusual and key  
 414 feature of the model calibrated on both hydrostatic and percolation experiments. In the  
 415 following we will present an explanation for how such relationships between dispersion and  
 416 degradation could arise, resulting from a non-monotonic spatial substrate profile, itself  
 417 derived from the respective effects of substrate dilution and bacterial density.

418 In the model, the instant exposure of bacteria to their substrate is maximal if all the bacteria  
 419 are located inside the voxel(s) with the highest substrate concentration. In the hydrostatic  
 420 calibrated parameter set, the profile of substrate concentration primarily resulted from its  
 421 initial heterogeneity (bell-shape red curve on Fig. 5A and pseudo bell-shape red curve on

422 Fig. 5B). The flux of substrate reaching each bacterium was therefore mostly determined by  
423 the distance between the bacterium and the initial location of substrate. The exposure of a  
424 single bacterium to the substrate decreased with its distance from the substrate initial  
425 location. This effect is referred to as “substrate dilution”. In these cases (Fig. 5A and B),  
426 mineralization was mainly regulated by substrate dilution, and therefore reduced by  
427 bacterial dispersion (Fig. 4, blue line). However, for the parameter set calibrated on both  
428 hydrostatic and percolation experiments, local degradation by aggregated bacteria reshaped  
429 the substrate spatial profile, thus critically changing the voxel(s) with the highest substrate  
430 concentration. The bacteria aggregated at their initial location consumed the substrate much  
431 faster than it was replenished by backward diffusion and dispersion, creating a critical  
432 inversion of the substrate gradient, which led to an intra-population competition for  
433 substrate (Fig. 5C). The competition was critical for bacterial densities as small as  $3.5 \cdot 10^{-3}$   
434  $\text{g}\cdot\text{l}^{-1}$  (Fig. 5C). In contrast, the dispersion of bacteria reduced competition by diluting the  
435 highest bacterial densities, thus flattening the substrate gradient inversion induced by  
436 bacterial local degradation, resulting in a better overall exposure of bacteria to the substrate  
437 concentrations, and thus an enhanced mineralization (Fig. 5D). In these cases (Fig. 5C and D),  
438 mineralization was mainly regulated by bacterial density, or in other words by the distances  
439 among bacteria. The relationship between bacterial density and the limitation of their  
440 exposure to the substrate is not instantaneous and is mediated by the local depletion of the  
441 substrate concentration. This is expressed in the model equations through the dependence  
442 of bacterial activity  $\mu(t)$  on substrate concentration  $S(t)$  (Eq. (7)) and the dependence of the  
443 substrate concentration  $S(t)$  on degradation  $\mu(t)\cdot B(t)$  (Eq. (1)), within each voxel. However,  
444 when bacterial dispersion was too great, substrate dilution became the dominant control  
445 again. This suggests that an optimal bacterial spatial spread exists for which the dilution of

446 substrate is compensated by the dilution of high local bacterial densities. The modeled  
 447 scenario illustrated by the two calibrated parameter sets were also observed for most of the  
 448 other parameter sets. The optimal dispersion coefficient for the 300 best-fitting  
 449 parameterizations to both hydrostatic and percolation experiments (smallest  $J_{1234}$  values)  
 450 was on average  $d_{disp} \approx 2 \cdot 10^{-5} \text{ m}^2 \cdot \text{d}^{-1}$  (Fig. S15), corresponding to a root-mean-square  
 451 displacement of bacteria of 2.2 mm during each percolation event.



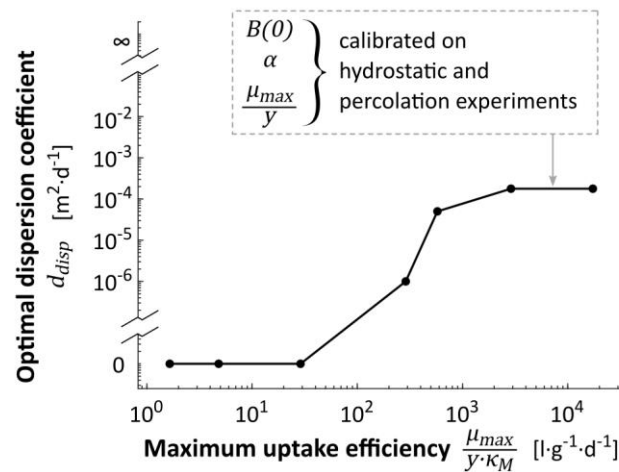
452  
 453 **Fig. 5.** Predicted substrate and bacterial spatial concentration profiles after 6 days of  
 454 diffusion and dispersion in the conditions of heterogeneous percolation experiment, in  
 455 which bacteria and substrate are initially located exclusively in the central cube (from 0 to  
 456 3 mm). Results are simulated on a  $9 \times 18 \times 18$  grid obtained by subdividing the  $3 \times 6 \times 6$  grid  
 457 used for the screenings. The results are represented for the parameter set calibrated using  
 458 only the sole hydrostatic experiment, either with a moderate dispersion ( $d_{disp} = 1.78 \cdot 10^{-4}$   
 459  $\text{m}^2 \cdot \text{d}^{-1}$ ) (A) or with the calibrated dispersion (no dispersion) (B), and for the biological

460 parameter set calibrated on both hydrostatic and percolation experiments, either without  
461 dispersion (C) or with the calibrated dispersion ( $d_{disp} = 1.78 \cdot 10^{-4} \text{ m}^2 \cdot \text{d}^{-1}$ ) (D). On one hand,  
462 bacteria are exposed to smaller substrate concentrations if they are far from the source  
463 (right part of the substrate concentration profiles). On the other hand, bacteria undergo  
464 competition if they are too close from each other (left part of the substrate concentration  
465 profiles). In (C), the bacteria aggregated below  $d$  consume the substrate faster than it is  
466 replenished by backward diffusion and dispersion. The total number of bacteria within the  
467 whole soil column at day 6 is similar in (A), (B), (C) and (D), respectively equal to  $6.0 \cdot 10^5$ ,  
468  $9.5 \cdot 10^5$ ,  $11.5 \cdot 10^5$  and  $11.3 \cdot 10^5$ . The final mineralization at day 24 is however strongly different  
469 between scenario, reaching respectively 3.2%, 5.3%, 9.1% and 24.7% of the initial mass of  
470  $^{14}\text{C}$ .

### 471 3.2.3 *Effect of bacterial uptake efficiency on the impact of dispersion on degradation*

472 A non-monotonic substrate concentration profile only occurs when bacterial degradation  
473 locally depletes the substrate faster than it is replenished by diffusion. This area of high local  
474 competition for substrate results from either high local densities of bacteria or high  
475 competitiveness or both. Bacterial competitiveness is related to their maximum uptake  
476 efficiency  $(1/y) \cdot \mu_{max}/\kappa_M$ , which also describes their capacity to maintain their activity and  
477 growth under dilute substrate concentrations (Healey, 1980; Button, 1991; Lobry et al.,  
478 1992). Bacteria with high maximum uptake efficiency are thus expected to benefit more  
479 from dispersion. Fig. 6 shows the optimal dispersion coefficient as a function of the  
480 maximum uptake efficiency, with all other parameters equal to those of the model calibrated  
481 on both hydrostatic and percolation experiments. The optimal dispersion coefficient,  
482 defined as the dispersion coefficient maximizing the final mineralization, increased with the  
483 maximum uptake efficiency. For small maximum uptake efficiencies of  $30 \text{ l} \cdot \text{g}^{-1} \cdot \text{d}^{-1}$  and below,  
484 mineralization was highest in the absence of dispersion, suggesting a regulation dominated  
485 by substrate dilution. For larger maximum uptake efficiencies, dispersion impacted positively

486 mineralization, suggesting that degradation shifted from being regulated by substrate  
 487 dilution to being regulated by bacterial densities, as bacteria were both more prone to  
 488 competition between themselves and more efficient under diluted substrate conditions. In  
 489 other words, the proximity to other bacteria constrained activity more than the proximity to  
 490 the substrate initial location enhanced it. This combined effect of the maximum uptake  
 491 efficiency and the bacterial dispersion on degradation was a general relationship common  
 492 to all parameterizations (Fig. S16).



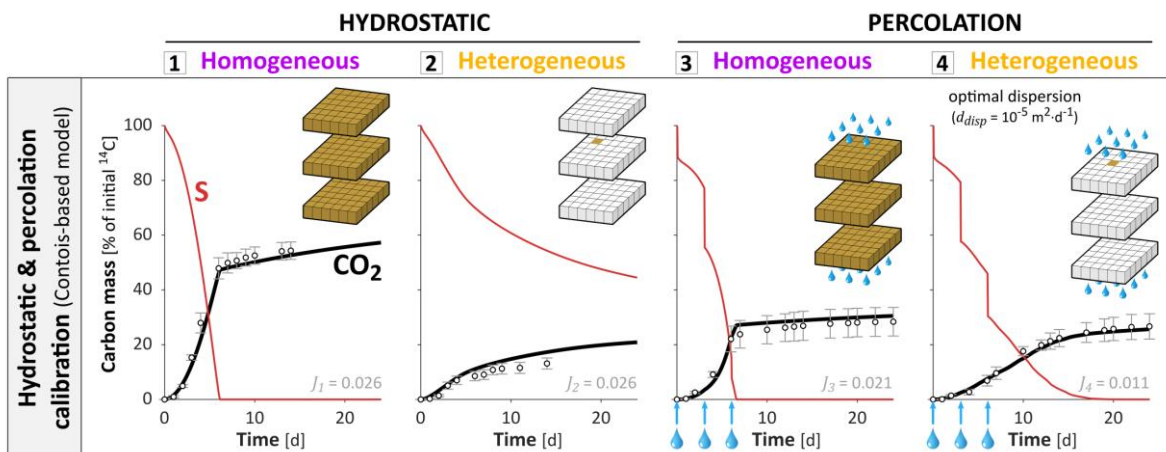
493  
 494 **Fig. 6.** Dispersion coefficient giving the highest predicted mineralization at day 24 as a  
 495 function of maximum uptake efficiency, all other parameters equal to those of the model  
 496 calibrated on both hydrostatic and percolation experiments.

### 497 **3.3. The Contois-based model as an alternative to Monod**

498 Given that degradation is regulated by both substrate dilution and bacterial density, and that  
 499 their relative importance is modulated by bacterial uptake efficiency at the lowest substrate  
 500 concentration,  $(1/y) \cdot \mu_{max}/K_M$ , we investigated the relevance of the Contois model by  
 501 applying the calibration methodology of section 2.5, as used in section 3.1. The interest in  
 502 the Contois growth law (Eq. (8)) stems from the inclusion of a regulation by density in the

503 expression of the uptake efficiency at the lowest substrate concentration, becoming  
 504  $(1/y) \cdot \mu_{max} / (B(t) \cdot \kappa_C)$ .

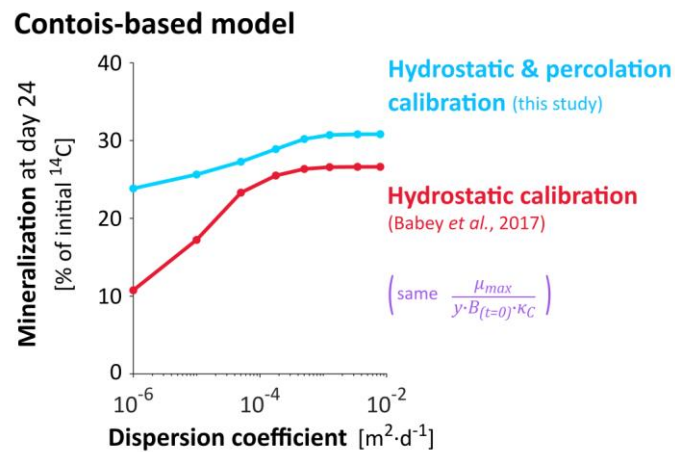
505 In comparison with the Monod-based model, the predictions of the experimental  
 506 observations of Pinheiro et al. (2015, 2018) were facilitated with the Contois-based model,  
 507 on three levels. First, the Contois-based model captured the degradation dynamics better  
 508 than the Monod-based model, especially for the 1% best-fitting parameterizations (smallest  
 509  $J_{1234}$  values) (Fig. S17). The calibrated Contois-based model had an overall discrepancy of  
 510  $J_{1234} = 0.022$  (Fig. 7), which was smaller than the lowest value of  $J_{1234} = 0.032$  obtained for  
 511 the calibrated Monod-based model (Fig. 3).



512  
 513 **Fig. 7.** Mineralization dynamics predicted with the Contois-based model calibrated on both  
 514 hydrostatic and percolation experiments. For representation and legend, see Fig. 3. The  
 515 carbon balance among the different pools is detailed in Fig. S11.

516 Second, the parameter sets that fitted homogeneous experiments also performed well  
 517 under heterogeneous conditions, as long as the dispersion coefficient  $d_{disp}$  was calibrated as  
 518 well (Fig. S18). It is an important advantage as it confers a better capacity to predict  
 519 degradation kinetics for heterogeneous and varying distributions, once the model is  
 520 calibrated in homogeneous conditions, which are more appropriate for the experimental

521 measurement of bacterial parameters. Besides, using a dispersion coefficient value different  
 522 from the calibrated one weakened the predictions of the mineralization dynamics but not  
 523 the predictions of the mineralization after 24 days, which remained satisfying regardless of  
 524 the dispersion coefficient. More precisely, the prediction of the final mineralization became  
 525 mostly independent of the dispersion coefficient, as shown for the calibrated model (Fig. 8).  
 526 This is because, in the Contois model at low substrate concentrations, the number of active  
 527 bacteria in a soil volume is exactly counterbalanced by the regulation of their uptake  
 528 efficiency by population density (Eq. (8)), resulting in limited effects of bacterial spreading  
 529 on overall mineralization (Fig. 8, constant part of the curves).



530  
 531 **Fig. 8.** Influence of the dispersion coefficient on mineralization at day 24 for the  
 532 Contois-based models calibrated on the sole hydrostatic experiments (thick red line) and on  
 533 both hydrostatic and percolation experiments (thick blue line). For representation and  
 534 legend, see Fig. 4.

## 535 4. Discussion

### 536 **4.1. Relevance of density control for 2,4-D degradation and soil carbon cycling**

#### 537 4.1.1 *Density control of soil oligotroph bacteria*

538 Bulk soil and highly-diluted environments are usually found to be dominated by bacteria with  
539 high maximum uptake efficiency, also called oligotrophs (Fierer et al., 2007; Nunan et al.,  
540 2020). Their high maximum uptake efficiency differentiates their life-history strategies and  
541 conditions their ability to thrive in resource poor environments (Button, 1993), also  
542 assimilated to K-strategy (Tecon and Or, 2017), by opposition to copiotrophic bacteria  
543 adapted to rich environments (r-strategy). The maximum uptake efficiency values of the 1%  
544 best-fitting parameter sets were of the order of  $10^3$ - $10^4$  l·g<sup>-1</sup>·d<sup>-1</sup> (volume of water per mass  
545 of bacteria per unit of time), within the range proposed by Button (1991) to define  
546 oligotrophs. Similar or higher maximum uptake efficiency values of the order of  
547  $10^4$ - $10^5$  l·g<sup>-1</sup>·d<sup>-1</sup> have been reported for soil oligotrophs (Ohta and Taniguchi, 1988; Zelenev  
548 et al., 2005). Values up to  $1.64 \cdot 10^5$  have been reported by Tuxen et al. (2002) for 2,4-D  
549 degraders in an aerobic aquifer and even greater values might also be possible (see section  
550 S5). The high maximum uptake efficiencies predicted in section 3.1 for the best-fitting  
551 parameterizations are therefore a plausible bacterial trait among 2,4-D degraders as well as  
552 bulk soil bacteria in general. It suggests that density control might be relevant for a  
553 component of soil bacteria, which would benefit from dispersion as suggested by Fig. 6. The  
554 calibrated model has shown in section 3.2.2 that the values of densities from which  
555 competition became critical were around  $3.5 \cdot 10^{-3}$  g·l<sup>-1</sup>, corresponding to  $7.5 \cdot 10^{-7}$  g·g (mass  
556 of bacteria per mass of dry soil), ranging in the low end of usual total soil bacterial densities



557 (Raynaud and Nunan, 2014; Kuzyakov and Blagodatskaya, 2015). This suggests that  
558 competition might play a significant role even under the low bacterial densities observed in  
559 bulk soils, at least in similar substrate conditions. Reciprocally, the model suggests that  
560 competition for substrate between copiotrophic bacteria only appears at much larger  
561 population densities, such as those found in soil biofilms (Holden et al., 1997, Or et al., 2007).  
562 Interestingly, copiotrophic bacteria have been reported to cohabit with oligotrophic bacteria  
563 even in diluted environments (Gözdereliler et al., 2013). Results from the screening suggest  
564 that, for densities of copiotrophs as low as for oligotrophs, their impact on overall  
565 decomposition in dilution-dominated environments would be much lower due to their  
566 poorly adapted uptake efficiency (Fig. 4A). Conversely, this striking density regulation might  
567 be one of the main limitations of the overall population densities in soils. Note that this  
568 density regulation occurs within a single population with homogeneous biological constants.  
569 Spatial heterogeneities and low substrate concentrations, common in bulk soil, may indeed  
570 shift competition from the inter-population level to the intra-population level (Pfeiffer et al.,  
571 2001; Roller and Schmidt, 2015).

#### 572 *4.1.2 A new perspective on Regulatory Gate hypothesis*

573 Density regulation might partially contribute to explain the common paradox of the apparent  
574 uncoupling between the overall mineralization of a soil volume and the size of its microbial  
575 population (Kemmitt et al., 2008). The rate of soil carbon mineralization remains the same  
576 even if 90% of the microbial decomposers are killed. This observation is commonly explained  
577 by the Regulatory Gate hypothesis, where mineralization is assumed to be controlled by an  
578 abiotic process, such as desorption or diffusion, that limits the availability of the substrate,  
579 resulting in mineralization rates that are independent of the degrader abundance. We

580 propose that the density regulation of decomposition in oligotrophic environments may  
581 contribute to this phenomenon, through competition for substrate or other biological  
582 interactions. In the case of competition-related density regulation, it reduces the  
583 dependence of the overall carbon mineralization on degrader abundance, as any increase of  
584 population density counterbalances the effect of the increased population size. Note that  
585 the involved abiotic process, namely the substrate diffusion backward to bacteria (see  
586 section 3.2), is well limiting but only in situations of high bacterial competition.

#### 587 ***4.2. Relevance of the ratio-dependent Contois model in soils***

588 As argued in section 3.3, ratio-dependence might facilitate decomposition modeling in the  
589 soil conditions typical of the experiments analyzed here. The Contois model's  
590  $(1/y) \cdot \mu_{max} / (\kappa_C \cdot B)$  calibrated in homogeneous conditions might be used in heterogeneous  
591 conditions more reliably than the Monod model's  $(1/y) \cdot \mu_{max} / \kappa_M$ , at least for soil systems in  
592 which the competition for the substrate plays a substantial role within the degrader  
593 population. The similarity between  $\kappa_M$  and  $\kappa_C B$  suggests the need to consider population  
594 density when measuring the apparent maximum uptake efficiency of soil bacteria to avoid  
595 underestimating it by unintentionally including density regulation. Moreover, the better  
596 predictions obtained with the Contois model in the soil conditions represented by the  
597 experiments suggest that the Contois ratio-dependence includes not only the effect of  
598 competition for substrate at the scale of measurement, but it can also reasonably reflect  
599 other density processes such as the spatial variability of bacterial distributions at finer scales  
600 related to their high degree of local aggregation in microcolonies (Raynaud and Nunan,  
601 2014). Moreover, ratio-dependence may also include the cumulative effects of ecological

602 interactions other than competition (Sibly and Hone, 2002). Note that the methodological  
603 approach used in this study for both Monod and Contois models is based on an effective  
604 representation of concentrations and parameters at the mm- to cm-scale of measurements.  
605 These effective concentrations and parameters conceptually integrate the smaller-scale  
606 processes highlighted by other studies (Ebrahimi and Or, 2014; Portell et al., 2018; Tecon et  
607 al., 2018). Such microscale processes should be addressed for further generalization beyond  
608 the conditions of the soil experiments analyzed here. Despite its advantages, Contois models  
609 have also a drawback with the fact that the modeled uptake efficiency of bacteria  
610 approaches infinity for low densities, which does not correspond to any physical nor  
611 biochemical process (Gleeson, 1994; Abrams, 2015). However, this side effect mostly affects  
612 a negligible fraction of the bacteria and the substrate, as it was the case in the soil conditions  
613 represented by the experiments.

614 Further work is required to confront the relevance of the Contois model to other soil  
615 systems. To the best of our knowledge, ratio-dependent growth models such as the Contois  
616 model have not yet been considered for the modeling of microbial degradation in soils.  
617 However, the Contois growth equation is generally accepted to be more appropriate than  
618 the Monod equation for modeling immobilized, heterogeneously distributed or mixed  
619 microbial cultures (Arditi and Saiah, 1992; Harmand and Godon, 2007), all of which are  
620 characteristics of soils. The regulation of individual activity by population density has  
621 frequently been justified as a “crowding effect” associated with high population densities  
622 leading to competition for substrate (Lobry and Harmand, 2006; Harmand and Godon, 2007;  
623 Krichen et al., 2018). However, little is known about possible density regulation when  
624 apparent microbial densities are low, as is observed in bulk soil (Raynaud and Nunan, 2014;

625 Kuzyakov and Blagodatskaya, 2015), although some studies have mentioned  
626 ratio-dependence in highly-diluted environments such as aquifers (Hansen et al., 2017). As  
627 discussed in section 4.1.1, the high maximum uptake efficiencies commonly observed for soil  
628 bacteria adapted to oligotrophic environments are relevant to draw attention on the  
629 potential significance of density control at low densities in oligotrophic soils, and thus  
630 ratio-dependent models, among which the Contois model is a consistent choice.

### 631 ***4.3. Hypothetical relationship between bacterial traits and their spatial*** 632 ***strategies***

633 Density regulation might be at the origin of a relationship between bacterial oligotrophy,  
634 their location in soil and their mobility strategy. Soil copiotroph bacteria have a maximum  
635 uptake efficiency mostly between  $100 \text{ l}\cdot\text{g}^{-1}\cdot\text{d}^{-1}$  (Button, 1991) and  $800 \text{ l}\cdot\text{g}^{-1}\cdot\text{d}^{-1}$  (Daugherty  
636 and Karel, 1994; Zelenev et al., 2005). For copiotrophs with maximum uptake efficiency  
637 values below  $288 \text{ l}\cdot\text{g}^{-1}\cdot\text{d}^{-1}$ , bacterial dispersion was largely detrimental to their activity (Fig. 4  
638 blue line, Fig. 6), in agreement with the results of Pagel et al. (2020), suggesting that  
639 copiotrophs have more aggregated distributions than oligotrophs. The negligible  
640 mineralization even without dispersion (Fig. 3-A4) also highlights the fact that copiotrophs  
641 are particularly inefficient at degrading substrates that diffuse in the environment, as also  
642 evidenced by Babey et al. (2017). To maintain significant activity, soil copiotrophs are likely  
643 to remain immobile in the close surroundings of the substrate source or any immobile  
644 substrate, likely attached to surfaces or embedded in EPS matrices. If not, they would be  
645 dispersed towards more diluted area where their low maximum uptake efficiency would  
646 result in negligible uptake. On the contrary, to survive and develop, soil oligotrophs should

647 be able to easily disperse and escape high competition areas. Given that soil is a poor and  
648 heterogeneous environment, this dispersion would be essentially passive (Nunan et al.,  
649 2020), through advective processes for example. We therefore suggest the existence of a  
650 theoretical relationship between proximity to substrate sources (respectively remoteness),  
651 copiotrophy (respectively oligotrophy) and attachment (respectively mobility).

## 652 **5. Conclusions**

653 Heterogeneous distributions of degraders and substrate in soils strongly control soil organic  
654 matter degradation through their interactions with the bacterial activity. Taking 2,4-D as a  
655 model organic solute substrate for soil bacteria, we investigated the coupled effects of  
656 bacteria and substrate distributions on one side and bacterial traits on the other side on  
657 substrate degradation. The analysis of published experiments with contrasted spreading  
658 conditions of both bacteria and substrate reveals that, in addition to the distance of bacteria  
659 from high substrate concentrations, mineralization is also surprisingly limited by the  
660 bacterial density even under the low bacterial densities commonly observed in bulk soils.  
661 Moreover, the impact of bacterial dispersion on solute substrate degradation can shift from  
662 negative to positive depending on the bacterial maximum uptake efficiency. The activity of  
663 soil oligotrophs may be mostly regulated by bacterial density rather than by substrate  
664 dilution, echoing the population size paradox regularly observed. It follows that the  
665 ratio-dependent Contois model might be more relevant to model bulk soil mineralization in  
666 the heterogeneous conditions investigated than the substrate-dependent Monod model. To  
667 predict the impact of spatial distributions on degradation in oligotrophic soil, and more  
668 particularly the impact of bacterial dispersion, we suggest that bacterial densities might be

669 a more useful measurement than the volumes of soil devoid or occupied with bacteria. With  
670 respect to the current lack of direct microscale data on microbial processes and distributions,  
671 we propose some key perspectives on the bacterial kinetics and distributions.

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## 681 **Appendix A. Supplementary Data**

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