



HAL
open science

Multiphase rheology as a cause for stick-slip like melt extraction

Jean-Louis Vignerese, Bénédicte Cenki

► **To cite this version:**

Jean-Louis Vignerese, Bénédicte Cenki. Multiphase rheology as a cause for stick-slip like melt extraction. *Earth and Planetary Science Letters*, 2022, 577, 10.1016/j.epsl.2021.117269 . insu-03661255

HAL Id: insu-03661255

<https://insu.hal.science/insu-03661255>

Submitted on 5 Jan 2024

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial 4.0 International License

1 Multiphase rheology as a cause for stick-slip like melt extraction

2 Jean-Louis Vigneresse^{1*} and Bénédicte Cenko²

3 ¹GéoRessources, UMR 7539, Université de Lorraine, 54506 Vandœuvre /Nancy Cédex,
4 France, jean-louis.vigneresse@univ-lorraine.fr * Corresponding author

5 ²Géosciences Montpellier UMR5243, Université de Montpellier, CNRS, 34095 Montpellier
6 cedex 5, France, benedicte.cenko@umontpellier.fr

7

8 Highlights

- 9 ○ Numerical model of melt extraction
- 10 ○ Sporadic motion of melt and an analogue with dry friction (stick-slip)
- 11 ○ Hysteresis and memory effects during changes in input parameters
- 12 ○ Global expression of the free energy variation, with a term akin to velocity and a
13 second term to acceleration
- 14 ○ Interactions between strain, chemistry and entropy (order)

15

16 **ABSTRACT**

17 When partial melting is parametrized in numerical thermo-mechanical models in a
18 straightforward manner, it directly relates to variations in temperature. However, melt
19 extraction is, to our knowledge rarely taken into account in these models, even if melt
20 extraction and escape have a key influence on the mechanical behavior of the entire system.

21 Here, we propose a first step towards a numerical simulation for melt extraction in partially
22 molten portions of the continental crust. We use a Lagrangian description, considering a
23 discrete mineral distribution at sample scale. A cellular automaton mimics the mobile melt
24 phase moving from cell to cell. An infinite source and sink for the melt and lateral periodic
25 boundaries avoid melt accumulation. Time is represented by a proxy through successive
26 iterations. The major controlling factors are the initial bulk melt percentage and a threshold
27 controlling melt escape. We test pure and simple shear conditions to simulate strain
28 partitioning between the melt and its matrix. The parameters, provided they do not vary,
29 determine a regular or steady state amount of melt extraction after a few iterations. In
30 contrast, for varying parameters during a single run, the quantity of melt extracted at each step
31 presents alternate variations in amplitude (σ_μ) around an average value. Variations are neither
32 periodic in the sense of a sinusoidal wave, nor of equal amplitude. Melt extraction behaves
33 similar to a stick-slip motion in dry friction. Since we cannot have access to the prediction of
34 melt amount, we test the origin of the sporadic motion. We examine variations of the input
35 parameters (initial melt ϕ , permeability ζ , applied stress σ or τ). The transient effects mark by
36 spikes ($\Delta\mu$) with two contrasted situations. Return to a steady state is exponential, similar to
37 the discharge of a capacitor. Nevertheless, a first mode occurs when the changes are made on
38 separate runs, reinitializing the parameters between them. These experiments show a linear
39 dependence of the variation in amplitude of σ_μ with the parameters. Alternatively, when the
40 changes are successive during the same run, then σ_μ decreases accordingly. The exponential
41 variation is interpreted as a memory effect or hysteretic behavior of melt extraction. Similarly,
42 the amplitude of the stiction spikes ($\Delta\mu$) does not show the same behavior when permeability
43 increases or decreases. By analogy with friction, the change from sticking to slipping is
44 interpreted as a manifestation of pinning and depinning, leading to an avalanche-like
45 phenomena. Mechanically they correspond to a change in friction velocity. Physically, it

46 points to the change in the free energy of the system, with a first derivative akin to velocity
47 and a second derivative akin to acceleration. During melting, the mechanical barriers
48 correspond to the energy variations due to chemical reactions between melt and crystals.
49 Depinning corresponds to enhanced melting whereas pinning links with back reactions
50 leading to garnet growth. The sporadic melt extraction links dry friction to poroelastic flow,
51 or more generally to a consequence of two-phases rheology. Hence the sporadic motion of the
52 weak phase originates from interactions with its matrix. Hence the flow (velocity) suffers
53 modulations due to interactions with chemical potential (back-reaction, melting), entropy
54 (ordering through dislocation creep), and strain (aperture of porous flow). Such interactions
55 should be considered when dealing with quantitative mush (melt+ matrix) modeling
56 (rejuvenation, extrusion).

57

58 **Graphical Abstract**

59

60 **Keywords:** stick-slip; melt extraction; friction; hysteresis; pinning/depinning

61 Running title Melt extraction: stick-slip motion

62 Submitted to Earth and Planetary Science Letters

63 Version *2/11/2022*

64 Statistics words (text) 5000 characters 30000 Figures 9 Ref 46

65

66 Corresponding author : J.L. Vignerese, GéoRessources, UMR 7539, Université de
67 Lorraine, 54506 Vandoeuvre/Nancy Cédex, France

68 Tel 00 333 8368 4737 Fax 00 333 8368 4701

69 e-mail jean-louis.vignerese@univ-lorraine.fr

70

71

72 **1. Introduction**

73 Temperature rules melt generation in the crust, most often through heat diffusion, or
74 advection, which are both continuous processes. The extraction of melt out of its matrix is by
75 contrast discontinuous, initiating from melt pockets that aggregate to form veins to escaping
76 and feeding isolated plutons (Kisters et al., 1998). According to such field observations, the
77 paradigm of melt extraction and transport turned from diapirism or dyking to discrete solitary
78 waves (Connolly and Podladchikov, 2000; Rabinowicz and Vigneresse, 2004) that propagate
79 from themselves, driven by alternate changes in permeability. Such models of porous flow
80 also apply to propagating fluids, melt or water within a stressed crust, with applications to
81 magmatism, geothermic and water circulation into the basement. They form the basis of
82 poroelasticity (Cheng, 2016). The paradigm switch from a viscous material more mobile into
83 a viscous matrix to a problem of permeability suggests that hydraulic parameters should rule
84 crustal fluid transport.

85 Modelling melt extraction thus points to the rheology of multi-phase material, i.e. the
86 differential response of a mobile fluid circulating into a deforming matrix. The viscosity
87 contrast between the melt and its matrix i.e. a weak and a strong phase, may be apprehended
88 from several points of view including estimates of an effective viscosity, intermediate
89 between the two end-members (Pistone et al., 2016; Agard et al., 2018). Three factors then
90 add complexity to such a direct approach. The first is the dependence of the viscosity of the
91 strong phase on the applied strain rate whereas the melt is simply Arrhenian. This makes the
92 experimental studies hard to be transposed to natural conditions (Pistone et al., 2016) because
93 the fast strain rate imposed by the duration of experiments. The second factor is strain
94 partitioning (Vigneresse and Tikoff, 1999) between the weak and the strong phase (Goodwin
95 and Tikoff, 2002). Their respective abundance determines which phase globally controls the
96 effective rheology. When the weak phase is abundant, the rheology is governed by the

97 distribution and geometry of rigid blocks (Agard et al., 2018; Ioannidi et al., 2021) as
98 observed at the interface of subduction zones. In contrast, when the weak phase is in minor
99 amount, the strong rigid matrix controls the bulk rheology (Vachon and Hyeronimus, 2017).
100 The two approaches can be considered as a variant of a direct problem, overdetermined, and
101 always offering a solution. However, they do not tackle the third factor, which is the exchange
102 of matter between the melt and its matrix, forming feedback loops. Hence melting occurs
103 when adequate minerals are close together, thus lowering the energy of melting. Conversely
104 back reactions, especially involving ferro-magnesian minerals, precipitate elements,
105 consuming energy to form incongruent minerals (Cenki et al., 2002; Nyström and Kriegsman,
106 2003). The two reactions act in modifying the free Gibbs energy budget of the reactions.
107 However, their extent, i.e. the amount of material involved that rules the energy budget cannot
108 be controlled exactly in time and space. It thus lowers the accuracy of a direct modelling.

109 To partly avoid and solve these ambiguities, a cellular automaton has been constructed
110 and modified (Vigneresse and Burg, 2005; Vigneresse et al., 2020), giving insights to the
111 mechanisms of two-phase rheology in partially molten rocks (see below).

112 The process of melt fusion and extraction can be directly compared with frictional motion,
113 that rules displacement of two materials in contact under an applied force (Persson, 1995;
114 Marone, 1998). Friction characterizes the relative motion, usually between two solid phases
115 (Byerlee, 1978). It links the tangential (τ) to normal (σ) forces (Fig. 1). The former controls
116 motion imposed by an applied force, whereas the second akin to the gravitational force resists
117 to lateral motion. The static coefficient of friction (μ_0) is a linear function of the applied
118 forces, and it is constant for most materials (Byerlee, 1978) (Fig. 1a). The irregularities of the
119 contact make lateral motion sporadic in time. A kinetic coefficient of friction (Dieterich,
120 1979; Ruina, 1983) quantifies the derivative of the displacement, akin to a velocity.

121 Irregularities, or asperities in the planar surface separating the two bodies resist to motion.
122 The frictional motion becomes transient owing to the change in the velocity of the motion
123 (Fig. 1b). Stick-slip is a special case of motion in which a prolonged stagnation (the “stick”
124 phase) is interrupted by short bursts of movement (the “slip” phase). The motion is not
125 periodic in the sense that the time interval between the short bursts is not always identical,
126 depending on the resisting forces (Fig. 1c). The amplitude of the bursts that vary around an
127 average value are not identical neither. When the initial conditions locally vary, a transient
128 motion develops with sudden acceleration in motion. Such transient effects are not well
129 understood, though they represent the responses of the system to the initial parameters
130 variations

131 While the discontinuous nature of melt extraction could be easily been transposed from
132 discrete permeability waves, the transient effects are not intuitively transposable to
133 displacements between viscous bodies or to poroelasticity. Hence, the discontinuous nature of
134 seismicity for example, yield to the concepts of asperity, small irregularity in the contact
135 between materials. They represent pinning points restraining motion and the sudden change in
136 contact area, or depinning leads to sliding. Following this model, frictional motion manifests
137 by alternate sticking and sliding phases, known as stick-slip motion (Baumberger and Caroli,
138 2006; Brace and Byerle, 2008).

139 Illustrating and interpreting frictional motion between a fluid, melt or water, and its matrix
140 represents an example of alternate concepts between rheology and hydraulics applied to
141 magmatism. The starting point of this study is a numerical simulation of melt extraction using
142 a cellular automaton (Vigneresse and Burg, 2005). The origin of a sporadic motion is
143 examined in relation with the initial parameters and internal changes in free energy through
144 the interactions between its three components (entropy, strain and chemical potential).

145

146 2. Numerical modelling set up

147 A numerical modelling using a Lagrangian formulation has been built from a cellular
148 automaton modified from a preceding model (e.g. available code in Vigneresse and Burg,
149 2005). The initial system consists in discrete cells, organized in a hexagonal pattern with
150 dimension 64 by 38 cells (Fig. 1). Cells contain indifferently melt and solid in proportion
151 varying from 0 to 100 (Vigneresse et al., 2021). The solid phase involves quartz and
152 plagioclase, two micas as muscovite and biotite, and some incongruent or residual
153 ferromagnesian minerals. The initial mineral and melt proportions are fixed, with a textural
154 disposition, either random, in layers or small pockets (Vigneresse et al., 2021). A
155 stoichiometric equation rules dehydration melting (Gardien et al., 2000). Similarly, back
156 reactions are introduced for the ferromagnesian (Cenki et al., 2002). The melt sink is infinite,
157 as well the heat source. Those two conditions assume a stationary condition for melt
158 evolution. The remaining solid phase is replaced by the moving melt from one cell to the next,
159 according to simple and/or pure shear geometry. Shear strain localization is possible by
160 moving melt to more than one cell. Periodic lateral boundaries avoid melt accumulation. The
161 initial melt content can be modulated, or by partial refilling in melt mimicking temperature
162 burst. Superimposed to the initial conditions and geometric parameters, two rules control melt
163 motion and extraction (Vigneresse et al., 1996). In addition to a liquid percolation threshold
164 (LPT) a melt escape threshold (MET) controls extraction. The effect of stress and strain rate is
165 also examined in both situations of pure (σ) and simple (τ) shear. However, the adopted 2D
166 geometry biases the results since melt may escape the system in the third direction. The time
167 is represented by an iterative process. As in most cellular automata, that are qualitative by
168 essence, the time aspects are only apprehended through iterations, i.e. without exact
169 calibration in time.

171 3. Results concerning melt extraction

172 Results allow to identify the effects on melt extraction of the main parameters: i.e. the
173 initial amount of melt (ϕ_0) and the threshold for extraction (MET). The former is supposed to
174 overcome melt percolation through the system. The latter is equivalent to fluid draining, thus
175 renamed as ζ (Fig. 2d, e). All parameters vary on a scale from 0 to 100. The extreme values of
176 met (> 50), i.e. with more melt than matrix, are not representative of nature, but allow a
177 control on the evolution of the numerical model.

178 Bulk results show a stationary state when there is no significant change in the input
179 parameters, but a spike in melt extraction occurs when the ζ value changes, and more
180 commonly when sporadic extraction occurs. The three modes (stationary, spike and sporadic)
181 show a good analogy with frictional motion, pointing to interpretation with the changes in
182 friction, i.e. in terms analogue to a frictional law (e.g. Marone, 1998).

183 3.1. Preliminary assessment of the results

184 Melt extraction is scaled in relative quantity of melt extracted versus iterations, i.e. time,
185 without possibility of more precise scaling (e.g. hours or years) yielding a relative time scale.
186 Similarly, the energy budget of the process is not quantitatively available.

187 The amount of melt is another parameter used during the computation. The possible
188 amount of melt in each cell ranges from 0 to 100. Consequently, the extracted melt fraction
189 (σ) at each step also ranges from 0 to 100 (Fig. 3). The fraction of extracted melt per step
190 shows, as expected, a spike at the start of the model, that rapidly converges toward a
191 stationary value. The spike is positive when the initial melt fraction is high, negative when
192 melt fraction is low. Curves with different initial melt fractions converge after about 70-80
193 steps. In a second stage, the percentage of melt extracted (σ) at each step is normalized to the
194 initial melt (ϕ_0) and the dimension of the model ($N_x * N_y$), providing the coefficient of melt

195 extraction (μ) (Fig. 3). After normalization, the value converges after about 150 steps toward
196 a value ranging 0 to about 1.5 (Fig. 3) increasing with the initial melt (ϕ_0). In this example, the
197 linear relation between μ and ϕ_0 defines a bulk coefficient of melt extraction (μ_0) that equals
198 0.17 ± 0.002 whatever the initial melt percentage, and using a value of 0.25 for ζ assuming a
199 one order less variation of (μ_0) as a function of ζ (Vigneresse and Burg, 2005).

200 A note of caution seems necessary, here, prior to interpreting the results. The coefficient
201 of melt extraction reaches a value of 1.0 for ϕ_0 larger than 50 and values higher correspond to
202 more melt extracted than produced. It means that the model diverges for high values of initial
203 melt, i.e. when there is more liquid than solid. Such a situation, mathematically possible, does
204 not reflect reality. However, it allows examining the results in a larger range of parameters.
205 We have indicated on the figures 6 to 7 the range of values for melt extraction, as estimated in
206 nature.

207 The coefficient of melt extraction (μ) shows two behaviours, regular or stationary, and
208 transient (Fig. 4), depending on whether the other parameters are constant or not during a
209 single run. Both cases are examined separately. The first case leads to a steady state regime.
210 Conversely, the transient mode due to varying ζ , or strain rate (ϵ° and γ°) induces a spike of
211 amplitude $\Delta\mu$.

212 The return to the steady state values follows a classical damping effect reflecting an
213 exponential decrease of the sharp increase with a characteristic time (D_c) about 70-80 steps
214 (Fig. 4). The lack of definition in time (i.e. iterations) precludes a more detailed investigation
215 on the variability of D_c . The spike is positive or negative, depending on the lack or enhanced
216 quantity of melt. It thus, like a charge and discharge of a capacitor, requires using a
217 logarithmic function, with the sharp increase normalized to the initial value. The function

218 takes the form of a $\ln(\phi/\phi_0)$. Such functions are applied to changes in ζ and strain rate (ϵ° and
219 γ°)

220 3.2. Towards a law of melt extraction

221 The preceding diagrams allow quantifying the relative variations as a function of the input
222 parameters. The results can be summarized as follow and interpreted by a simplified law of
223 melt extraction that takes into account the linear and non-linear effects, associated to the
224 potential barriers of melt extraction mimicking a frictional law (Vigneresse and Burg, 2005):

$$225 \mu = \mu_0 + a (\phi/\phi_0) + b (\zeta/\zeta_0) + c (\sigma/\sigma_0) + d \ln (\delta\phi/\phi_0) + e \ln (\delta\sigma^\circ/\sigma^\circ_0)$$

226 with the numerical values $\mu_0 = .022$, $a = .0018$, $b = -.0015$, and $c = .016$ (simple shear)
227 or $= .066$ (pure shear) (Vigneresse and Burg, 2005). The two functions in $\delta\phi$ and stress rate
228 ($\delta\sigma^\circ$) variations are non-linear and could not be determined quantitatively. Globally, the
229 effective production of melt ranges from 15 % to a maximum of 40 % depending on protolith
230 composition (Vigneresse et al., 2021). The potential barriers represent variations one order of
231 magnitude less than the coefficient of melt extraction μ_0 .

232 It should be compared to the suggested frictional laws (Dieterich 1972), with the modification
233 about displacement (Ruina 1983).

$$234 \mu = \mu_0 + a \ln(V/V_0) + b \ln(V_0 \theta / D_c) \text{ and } d\theta/dt = 1 - V\theta/D_c$$

235 Nevertheless, because of the poor access to time, the problems of strength recovery (healing)
236 interpreted from the state variable θ , without more specification (Ruina, 1983), have not been
237 considered in the actual study. The analogy points to the effect of velocity, either as the
238 velocity of traction during dry friction, or melt flux under strain partitioning on the viscous
239 melt, eventually enhanced by acting stress. In addition, a second term related to imposed
240 strain is present in the equations, that may have an influence on the sporadic motion during

241 friction or melt extraction. In a first approach, this term is a derivative of the motion, akin to
242 acceleration. It suggests that melt motion is represented by the relative value of melt volume
243 (ϕ/ϕ_0), whereas the remaining other parameters (ζ , strain) act as enhancing/damping the flow.

244

245 **4. Sporadic melt extraction**

246 One difficulty in interpreting the irregular sporadic melt extraction patterns mostly links
247 with the poor determination related to temporal calibration. It includes difficulties in a true
248 scaling in time units, but also difficulty in interpreting a discontinuous diagram of irregular
249 amount of extracted quantity of melt. It manifests during friction by a stick-slip motion (Fig.
250 2c), as a second order manifestation, interpreted by the hysteretic nature of friction. Hence,
251 friction is a dissipative function of energy, that needs additional energy to recover the initial
252 state. We provisionally adopt the term stick-slip behaviour for describing melt extraction.

253 4.1. Variations in stick-slip amplitude (σ_μ)

254 Stick-slip manifests during friction by alternate periods of no or restricted motion, or
255 stick phase, followed by slip at constant velocity (Fig. 2c). Similarly, the amplitude of melt
256 extraction, varies with iterations, but remains constrained between two values defining a
257 dynamic and static extraction modes (Fig. 2f). The difference is referred to as σ_μ . This value
258 is examined as a function of the other parameters. The amplitude of σ_μ varies linearly
259 between 0.0015 and 0.013 for an initial melt content varying from 5 to 85 %, though the last
260 value is unrealistic in nature, but brackets the range of variables. It slightly varies by about
261 0.003 for a variation of 50 in the initial melt content (ϕ_0), i.e. from dry to saturated melt. The
262 values increase with ζ by about 0.035 for a change of 20 in ζ (Fig. 5). Thus, the effect is more

263 pronounced depending on ζ rather than on ϕ_0 . The linear effect also excludes artifacts such as
264 errors or noise. It correlates with a non-strictly periodic effect in amplitude.

265 The very moderate decrease, about one order less than the coefficient of melt
266 extraction with the initial melt content points to an enhanced melt extraction with the total
267 melt content (Fig. 5), suggesting a possible lubrication of the interface between melt and its
268 matrix. Nevertheless, this effect remains weak compared to the effect of ζ . It also points to
269 enhanced/damping effects occurring during transient changes compared to the variations
270 induced by varying the initial parameters.

271

272 4.2. Variation of amplitude depending on the system history

273 The amplitude increase of the sporadic melt extraction has been tested against the other
274 parameters evidenced in the melt extraction law (ϕ , ζ , ε°) (Vigneresse and Burg, 2005). This
275 has been realized under two modes. In a first set of runs, these parameters remain fixed along
276 the whole run. The separate runs (previous sections) show moderate variations in σ_μ , with a
277 bulk increasing linear trend whatever the change in all parameters (Fig. 6).

278 Alternatively, a second set of runs, that are successive, allows abrupt changes in the
279 parameters during iterations, indicated by enhanced melt extraction spikes before a
280 progressive return to the previous extraction mode. Those spikes, with amplitude $\Delta\mu$, are
281 positive or negative depending on charge or discharge of melt. By contrast, the amplitude of
282 melt extraction (σ_μ) shows a pronounced decrease, for all parameters (ϕ , ζ , ε°), when one of
283 them is kept constant (Fig. 6). The decrease in amplitude is significant (0.12 to 0.0), since σ_μ
284 tends to zero when the changes in input parameters modify the coefficient of melt extraction
285 by a factor 2 or 3. In addition the variation in amplitude for separate runs largely overcomes

286 the variation for separate runs (Fig. 6). To further explore the situation, their amplitudes have
287 been plotted during successive variations of the inputs (Fig. 7a). The amplitudes decrease with
288 iteration (i.e. time) according to a law proportional to $\zeta^{-1} \ln(\zeta/\zeta_0)$ with a coefficient α and ζ_0
289 the initial permeability (Fig. 7b). It points to a memory effect acting on the stick-slip
290 amplitude, and decreasing with the system memory. Nevertheless, as stated before, the poor
291 variability in the amplitude of Dc, i.e. some quantification of the time constant of the
292 relaxation, precludes insights to the energetical aspects of melt extraction. An interpretation
293 would be an enhanced permeability due to healing. Once porosity is enlarged, it should keep
294 its high value, leading to enhanced melt extraction.

295

296 4.3. Return on the spike amplitudes

297 The fact that σ_μ decrease with the melt extraction has an impact on the amplitude of the
298 enhanced melt extraction spikes ($\Delta\mu$), but no real trend could be estimated from the spike
299 amplitudes. Taking into account a possible memory effect allows to reconsider differently the
300 spikes. From the various runs, the spikes amplitudes have been plotted as a function of ζ
301 variations, distinguishing the positive from the negative variations. These plot shows a weak
302 coherence, adopting a double curve, not symmetric to each other, but mimicking some kind of
303 hysteresis cycle as a function of the quantity of melt refilling or escaping. It corroborates the
304 asymmetric plot of $\Delta\mu$ versus ϕ variations (Fig. 4).

305

306 5. Implications and discussion

307 The large variability observed in the modelled melt extraction has some direct
308 consequences.

309

310 5.1. Hysteresis

311 Hysteresis is the dynamic time lag between input and output of the system, reflecting its
312 dependence on its previous history. It is associated with the need to consume or deliver
313 additional energy for the return to initial conditions. Examples of hysteresis in geological
314 systems include remanent magnetism, plastic deformation, phase change during melting and
315 crystallization, or surface tension (Persson, 1995; Hill, 1998). A simplistic representative
316 example is a drop of water on a glass. In case of a porous medium, with a two-phase flow (air
317 and water) the curve of capillary pressure versus the wetting phase saturation also describes a
318 hysteretic cycle, with a wetting stage, and return along drying curve. Due to the solubility of
319 the gaseous phase, the hysteretic curve shrinks with successive cycles of drying and wetting.

320

321 5.2. Pinning and depinning

322 The sporadic melt extraction, with alternate bursts, links to variations in the melt
323 velocity, i.e. de-/acceleration. As stated in friction, it relates to a state variable θ (Ruina,
324 1983), essentially defined by its derivative with time ($d\theta/dt$). It is characterized by alternate
325 episodes of lock (stick) and motion (slip). The frictional interpretation commonly relates to
326 those stages to surface conditions. Indeed, they present asperities, or small stronger areas of
327 relief, arresting the motion. If the interpretation is correct, the alternate stages of
328 enhanced/reduced melt extraction should also be controlled by such points of resistance, i.e.
329 pinning and depinning, and their equivalent in terms of chemical thermodynamics is
330 consuming/releasing Gibbs free energy.

331 Experimental studies on stick-slip, using small transducers and fast camera recording
332 allow a better understanding of the phenomena traduces in a force vs time diagram during dry

333 friction (Fig. 8a). The equivalent diagram in displacement vs time indicates periods of no
334 motion corresponding to the stick stage followed by a sudden jump in velocity (Fig. 8a).
335 However, more sensible experiments from the group of Jay Fineberg (Hebrew University,
336 Jerusalem) tackled the fine motion during the sticking stage in a bi-phase material interface.
337 They record high-speed measurements of the real contact area, whereas the strain field is
338 measured in the region surrounding propagating rupture tips (Svetlizky and Fineberg, 2014).
339 They observe highly localized slip pulses accompanied by a large local reduction of the
340 normal stress near the rupture tip. The pulses propagate along the direction of motion of the
341 softer material at a maximal velocity, evolving while propagating. When the rupture front
342 velocity overcomes the seismic S-waves velocity, it determines a supershear front (Rosakis et
343 al., 1999) that propagates in the opposite direction, suggesting crack-like rupture. Such results
344 corroborate the ideas of dislocations steadily propagating along a material interface. In
345 parallel, the transition from static to dynamic friction is quantitatively described by classical
346 singular solutions for the motion of a rapid shear crack (Rubinstein et al., 2004). They
347 observed propagations at sub-Rayleigh velocity, the sound speed along a free surface, but also
348 a slower front that could be the trace of healing at contact area, before slip arrest. This is
349 similar to ageing that immediately commences as contact area increases. Its characteristic
350 timescale is determined by the local memory of the system, according to its effective contact
351 time before slip arrest (Ben-David et al., 2010).

352 Such evolution at the microscale of dislocations has been surveyed in detail on a single
353 LiF crystal since it allows simultaneous observation of evolution of the dislocation structures
354 through the Bauschinger effect, a piling-up of dislocations at grain boundaries. Indeed, the
355 microscopic distribution of stress alters the stress/strain characteristics of the material (Harea
356 et al., 2013). This property is used to trace the dislocation pattern during deformation
357 (Perfilyev et al., 2014). Observations show that the indenter velocity increases nonlinearly

358 during the transition from stick to slip. in the mid-part of the slip phase, the indenter reaches
359 its maximum velocity, while the friction force decreases abruptly (Fig. 8b). The average
360 velocity of the indenter in the stick phase drops by about 300 times than the maximum
361 velocity in the slip phase. It drastically affects the behaviour of the dislocation regime. Three
362 stages of creep are evidenced between two slip events (Perfilyev et al., 2014). First, a primary
363 creep links with the formation of bonds (pinning effect). Their number increases during the
364 second long-time steady state. In contrast, the tertiary creep is characterized by the rupture of
365 bonds (depinning effect).

366 The link between pinning/depinning and dislocations (Anderson et al., 2017) brings
367 some highlights on the asperity model since it marks by acoustic emission that may be
368 recorded during the stick-slip motion. Known as the Barkhausen noise in ferro-magnetism
369 (Santa-Aho et al., 2019) it results from the rapid size variations of magnetic domains, with
370 their magnetic spins equally oriented. The produced noise is recorded and is currently in use
371 to detect defects in the alloys.

372

373 5.3. Implications on melt extraction processes

374 The preceding section provides mechanical examples involving the role of
375 dislocations in pinning/depinning effects during frictional stick-slip motion (Fig. 9a).
376 Dislocations creep correspond to a change in the ordering of a crystal, i.e. a change in
377 entropy. It may apply to solid-solid contacts but fails when a viscous material is present
378 (Marone, 1998). Another model of stick-slip behaviour derived from poroelasticity (Cheng,
379 2016) takes into account the surface tension and the solubility between the components, solid,
380 liquid and air (Fig. 9c). It corresponds to an interaction between strain of the flow and strain
381 imposed by the matrix. In both cases a locking effect, or barrier, retards motion while an

382 acceleration occurs when additional energy overcomes the resisting forces. In both cases, the
383 processes are the sum of a static equivalent to the free energy of the system plus a term
384 function of the its first derivative, equivalent to velocity, and a second derivative, equivalent
385 to acceleration. The energy barrier linked to dislocations has a fixed amplitude, also linked to
386 a progressive increase in pinning, thus favouring a limit to a dynamic friction coefficient.

387 In between those two end-members, melt extraction results from a viscous-viscous
388 interface. There is no dislocation, by lack of a true crystalline network, and there is no non-
389 percolating third phase as air bubbles. Hence the matrix is a crystalline rock undergoing
390 melting, but with a heterogeneous amount of solid phases either products or unreacted
391 inherited minerals. The melt is a mixture of silicate polymers and ions in a supercritical state.
392 Metamorphic reactions between the melt and its matrix may be considered as a cause of for
393 pinning/depinning effects (Fig. 9b). Retrograde, melt-consuming reactions, producing solid
394 phase are the counterpart of prograde melting and may produce new minerals phases acting as
395 asperities causing stick-slip or sporadic melt extraction. Both type of reactions are commonly
396 observed in migmatites, rocks that have undergone partial melting (Stüwe, 1997; Kriegsman,
397 2001).

398 Dehydration of hydrous minerals occurs during prograde melting, mainly muscovite
399 above 730 °C and the biotite around 800°C (Clemens, 1990). Hence the leucocratic portions
400 of migmatites, or leucosomes, are mostly composed of quartz, feldspars and rare micas. These
401 are interpreted as crystallized melt usually of granitic composition. The darker
402 ferromagnesian minerals form the melanosome and may segregate aside the leucosomes by
403 flow differentiation or reactive transport (Kriegsman, 2001). The minerals from the
404 melanosome can either be relics of reactants, products of melt consuming reactions or newly
405 formed incongruent mineral phases.

406 During muscovite breakdown with temperature increase, mineral phases disappear
407 whereas melt-consuming reactions may form new solid phase upon cooling (Kriegsman,
408 2001). Those may be quartz, plagioclase and muscovite (Nyman et al., 1995; Konn et al.,
409 1997) followed by biotite, sillimanite, quartz and plagioclase (Cenki et al., 2002). In contrast
410 garnet and cordierite may grow as incongruent porphyroblasts (Nyström and Kriegsman,
411 2003). During growth, they may incorporate large quartz rounded inclusions indicating
412 surface tension and immiscibility. Upon cooling, retrograde phases may form coronas around
413 these porphyroblasts, surrounding by the leucosome, illustrating back reactions between the
414 restitic phases and the melt. Ti-rich biotites also grow at the expense of ferro-magnesian
415 incongruent phases (Cesare et al., 2008). Melt consuming reactions include mica growth, fine-
416 grained intergrowths in micas, breakdown of incongruent mineral phases and coronas
417 formation (Cenki et al., 2002; Cesare et al., 2008). The amount of melt production, or
418 consumption, depends on the ability of elements to enter such reactions. It thus has no
419 bounding values, making the amplitude of stick slip non periodic.

420 Interestingly, the formation of garnet coronas through a process of diffusion is slow.
421 Thus, depending on the heat flow variations, a so-called tectonic “hysteresis” in the melt
422 extraction may result (Loomis, 1977). The retrograde reactions, added to changes in the pore
423 size, determine zones of pinning and depinning (Fig. 9b).

424 The sporadic mode of melt extraction corroborates with the observation of segregated
425 melt from its matrix in discrete pockets (Kisters et al., 1998). Such discrete pockets
426 commonly form shrinking patterns with subvertical pipe-like bodies. They further connect
427 together to evacuate a felsic melt into veins following the regional foliation planes. Discrete
428 melt batches also corroborate numerical models of porosity waves (Connolly and
429 Podladchikov, 2000; Rabinowicz and Vigneresse, 2004).

430 Another conclusion from such competition between melting and solidification (back-
431 reactions) during melt extraction may have implications on the interpretation of field samples.
432 Hence, magma crystallization is commonly coeval with fluids release, due to space shrinking,
433 rapidly reaching fluid saturation (e.g. second boiling, Candela, 1997). The additional presence
434 of a compressible fluid phase within the competing enhanced melting and crystallization (i.e.
435 external pressures variations on the system) is coherent with an variation in the extraction rate
436 during solidification with sudden pressure changes (Pistone et al., 2020). The stick-slip like
437 melt motion thus adds a supplementary positive feedback loop to the process. Conversely,
438 during initial melting, the freed fluid from hydrous minerals does not develop in a separate
439 phase since it immediately enters the melt because of its solubility. In addition, the poor
440 volume change due to biotite dehydration (Rushmer, 2001) does not alter the bulk volume of
441 the system, thus avoiding pressure related loops.

442 Hysteresis and differences in melt motion increase the symmetry breaking between
443 melting and solidification. In particular, the order of minerals crystallization, commonly
444 accepted from ferro-magnesian, plagioclase and quartz, i.e. following a decrease in chemical
445 potential, is perturbed when fluids (water-dominated) enter the process. It has implications
446 on the segregation or not of incompatible elements that do not follow a similar trend
447 compared to congruent minerals. The second important difference between melting and
448 crystallization relates to the fate of water, inducing a large scale more or less independent
449 hydrothermal activity during magma crystallization, whereas the processes of dry or hydrous
450 melting imply changes in the melt final composition. However, developing such ideas is out
451 of the scope of the present paper.

452 5.4. Analogy between dry friction and melt extraction

453 The dry friction coefficient (μ_d) characterises the relationship between a force (F) applied to a
454 mass (m) pulled by a spring. The usual relation between the force applied on a mass and its
455 acceleration applies. The force is proportional to the first derivative of the motion (dx/dt)
456 through the coefficient of friction (μ_d). Thus

$$457 \quad m \delta^2 x / \delta t^2 = - \mu_d \delta x / \delta t$$

458 The solution to this differential equation is given by

$$459 \quad \delta x(t) / \delta t = \delta x(0) / \delta t \exp [-(\mu_d / m) t]$$

460 It relates the mass-force relationship to the stress-strain relationship. It also points the double
461 effect of the velocity and its variations (i.e. acceleration), thus providing an explanation to the
462 “state variable” (θ) introduced in dry friction (Ruina, 1983). In terms of melt extraction, it
463 corresponds to variation in the melt flow, either due to enhanced motion under applied stress
464 or enhanced/retarded melting.

465

466 5.5. Notes of caution

467 The ideas presented here are derived from the numerical model that examines the motion
468 of melt in migmatites. As such, it should be considered as a model, i.e. either representing
469 the effects of the identified parameters, or in contrast, as a way to bracket the values of the
470 parameters to better fit field observations. This requires a note of caution on the way the
471 results are interpreted. In its actual shape, all parameters have been varied in a full range.
472 For instance, the initial amount of melt ranges 0 to 100, though such high values are
473 obviously not observed in partially molten rocks. However, the extreme values are
474 considered as providing the numerical trends that rule the underlying process. Similarly, the
475 coefficient of melt extraction presents values greater than 1. It corresponds to the effect of an

476 infinite source and sink for the melt. They are not realistic. Such large values correspond to
477 an initial melt larger than 50. Such notes of caution being stated, what are the main results of
478 the model.

479 First, the model illustrates the role of sharp variations in the ability of melt to be
480 extracted and move (ζ), formerly assimilated to a melt escape threshold (Vigneresse and
481 Burg, 2005). These variations show a transient behaviour. It implies that the definition of the
482 melt escape threshold as a purely geometric threshold corresponding to percolation is
483 incorrect since it is essentially static. Hence, the ζ parameter is dynamic, determining
484 stiction spikes ($\Delta\mu$) and the amplitude of the stick-slip motion (σ_μ). It corresponds to a
485 variation of the flux, i.e. an acceleration, positive or negative, of the flux.

486 Secondly, chemical reactions develop between the matrix and the melt as prograde
487 reactions, but also between the melt and its matrix, as retrograde reactions. They consume or
488 produce melt, thus modifying the melt flow velocity. Such reactions of chemical exchange
489 are important when considering the free energy balance of the system. the corresponding
490 interactions between strain (flow velocity) and chemical potential (melting and back-
491 reactions) modify the bulk energy of the system, inducing sporadic variations. Because they
492 were not considered in other forward models (Ioannidi et al., 2021), they resulted in a
493 stationary motion, that certainly does not fully consider the exact nature of two-phase
494 rheology. Similarly, variations due to the energy required to overcome dislocations induces
495 instabilities due to strain (flow) and entropy due to ordering modification (dislocation creep)
496 as in dry friction (Fig. 9a). In poroelastic flow, the interactions develop between strain
497 (flow) and changes in the pore aperture (external strain) determine sporadic motion (Fig.
498 9c). The specific case of melt extraction implies chemical exchanges (melting and back
499 reactions) between the melt and its matrix (Fig. 9b).

500 Thirdly, the transient character of the flow presents some memory effects, due to the
501 irreversibility of the reactions and the thermodynamic metastability of some minerals
502 phases. However, such memory effects should not be confused with a time-related effect. In
503 particular, it does not implicitly point any re-use of veins.

504 At last, because the model is constructed from a cellular automata, the model cannot
505 be scaled in time, but uses a proxy through iterations. It is therefore not aimed to prediction,
506 e.g. for quantitative mush (crystal-melt) interaction, as described for magma chamber
507 rejuvenation or in magma eruptability. Similarly, a purely chemical approach is no more
508 predictive, and a dry friction modelling does not inform on the next earthquake to occur. In
509 contrast, and we focus on the message here delivered, any modeling of a multi-phase
510 material should incorporate the interactions between the three components of energy, i.e. the
511 external variables such as entropy, strain and chemical potential.

512

513 **6. Conclusions**

514 The present model examines the transient flow of melt extraction depending on variations
515 of the controlling parameters, mostly the changes in flux velocity (amount of melt). The
516 occurrence of spikes reflecting sporadic melt extraction with a progressive return to a
517 stationary state is remarkable. We interpret these results with an analogy to the stick-slip
518 behaviour in mechanics that might develop resulting in an irregular, non-periodic amount of
519 melt extracted. This effect shows hysteresis, essentially due to the presence of phases
520 consumed during prograde (melting) or retrograde (garnet and sillimatite growth)
521 metamorphic reactions. The interactions between strain, chemical potential and entropy
522 modify the free energy of the system. They act as pinning and depinning during dry friction
523 and other flows of a weaker phase within a stronger matrix. The model highlights a dynamic
524 concept of melting and places melt extraction between dry friction and poroelasticity.

525

526

527 **Acknowledgements** The actual paper started from a collaboration with Jean-Pierre Burg
528 (ETH Zurich). At that time, we were mostly interested in the rheology of a two-phase
529 material, with intrinsic feedback loops and non-linearity. More recent discussions focused on
530 the role of factors that alter continuous melting due to heat diffusion toward a discontinuous
531 motion of the melt. In particular, we focused on the transient behaviours, yielding an analogue
532 to stick-slip motion, pointing to pinning/depinning processes that illustrate discontinuous
533 motion. The analogies with dry friction and poroelasticity should bring insights on irregular
534 motions observed in those disciplines. BC acknowledges the funding from the European
535 Union's Horizon 2020 research and innovation program under grant agreement No 793978.
536 We appreciated the constructive comments of one reviewer, focusing on the link back to field
537 observations. The general comments also obliged us to focus more on the final message of the
538 paper, i.e. interactions between the energy components.

539

540

541 **References**

- 542 Agard, P., Plunder, A., Angiboust, S., Bonnet, G., Ruh, J., 2018. The subduction plate
543 interface: rock record and mechanical coupling (from long to short timescales). *Lithos* 320-
544 321, 537-566. <https://doi.org/10.1016/j.lithos.2018.09.029>
- 545 Anderson, P.M., Hirt, J.P., Lothe, J., 2017. *Theory of Dislocations*. Cambridge University
546 Press, 718 pp.
- 547 Baumberger, T., Caroli, C., 2006. Solid friction from stick–slip down to pinning and aging.
548 *Advances in Physics* 55, 279-348. <https://doi.org/10.1080/00018730600732186>
- 549 Ben-David, O., Rubinstein, S., Fineberg, J., 2010. Slip-stick and the evolution of frictional
550 strength. *Nature* 463, 76-79 <https://doi.org/10.1038/nature08676>
- 551 Brace, W.F., Byerlee, J.D., 2008. Stick-slip as a mechanism for earthquakes. *Science* 153,
552 990-992. <https://doi.org/10.1126/science.153.3739.990>
- 553 Byerlee, J.D., 1978. Friction of rocks. *Pure and Applied Geophysics* 116, 615-627.
554 https://doi.org/10.1007/978-3-0348-7182-2_4
- 555 Candela, A., 1997. A review of shallow, ore-related granites: textures, volatiles, and ore
556 metals. *Journal of Petrology* 38, 1619–1633. <https://doi.org/10.1093/petroj/38.12.1619>
- 557 Cenki, B., Kriegsman, L.M., Braun, I., 2002. Melt-producing and melt-consuming reactions
558 in anatectic granulites: P–T evolution of the Achankovil cordierite gneisses, South India.
559 *Journal of Metamorphic Geology* 20, 543-561. <https://doi.org/10.1046/j.1525->
560 1314.2002.00388.x
- 561 Cesare, B., Satish-Kumar, M., Cruciani, G., Pocker, S., Nodari, L., 2008. Mineral chemistry
562 of Ti-rich biotite from pegmatite and metapelitic granulites of the Kerala Khondalite Belt
563 (southeast India): Petrology and further insight into titanium substitutions. *American*
564 *Mineralogist* 93, 327-338. <https://doi.org/10.2138/am.2008.2579>
- 565 Cheng, A.H.D., 2016. *Poroelasticity*. Springer Heidelberg, 903 pp.
- 566 Clemens, J.D., 1990. The granulite-granite connexion. In: Vielzeuf, D., Vidal, P. (Eds)
567 *Granulites and Crustal Evolution*. NATO ASI Series C, 311. Springer, Dordrecht. 25-
568 36. https://doi.org/10.1007/978-94-009-2055-2_3
- 569 Connolly, J.A.D., Podladchikov, Y.Y., 2000. Temperature-dependent viscoelastic compaction
570 and compartmentalization in sedimentary basins. *Tectonophysics* 324, 137-168.
571 [https://doi.org/10.1016/S0040-1951\(00\)00084-6](https://doi.org/10.1016/S0040-1951(00)00084-6).
- 572 Dieterich, J.H., 1972. Time-dependent friction in rocks. *Journal of Geophysical Research*
573 B77, 3690-3697 <https://doi.org/10.1029/JB077i020p03690>

574 Dieterich, J.H., 1979. Modeling of rock friction: 1. Experimental results and constitutive
575 equations. *Journal of Geophysical Research* B84, 2161-2168
576 <https://doi.org/10.1029/JB084iB05p02161>.

577 Gardien, V., Thompson, A. B., Ulmer, P., 2000. Melting of biotite + plagioclase + quartz
578 gneisses: the role of H₂O in the stability of amphibole. *Journal of Petrology* 41, 651-666.
579 <https://doi.org/10.1093/petrology/41.5.651>

580 Goodwin, L. B., Tikoff, B., 2002. Competency contrast, kinematics, and the development of
581 foliations and lineations in the crust. *Journal of Structural Geology* 24, 1065-1085.
582 [https://doi.org/10.1016/S0191-8141\(01\)00092-X](https://doi.org/10.1016/S0191-8141(01)00092-X)

583 Harea, E., Lapsker, I., Laikhtman, A., Rapoport, L., 2013. Bauschinger's effect and
584 dislocation structure under friction of LiF single crystals. *Tribology Letters* 52, 205–212.
585 <https://doi.org/10.1007/s11249-013-0206-y>

586 Hill, R., 1998. *The Mathematical Theory of Plasticity*. Oxford University Press, 372 pp.

587 Ioannidi, P. I., Le Pourhiet, L., Agard, P., Angiboust, S., Oncken, O., 2021. Effective
588 rheology of a two-phase subduction shear zone: Insights from numerical simple shear
589 experiments and implications for subduction zone interfaces. *Earth and Planetary Science*
590 *Letters* 256, 116913 <https://doi.org/10.1016/j.epsl.2021.116913>

591 Kisters, A. F. M., Ward, R. A., Anthonissen, C. J., Vietze, M.E., 1998. Melt segregation and
592 far-field melt transfer in the mid-crust. *Journal of the Geological Society* 166, 905-918.
593 <https://doi.org/10.1144/0016-76492009-012>

594 Konn, M.J., Spear, F.S., Valley, J.W., 1997. Dehydration melting and fluid recycling during
595 metamorphism: Rangeley Formation, New Hampshire, USA. *Journal of Petrology* 38,
596 1255–1277 <https://doi.org/10.1093/petroj/38.9.1255>

597 Kriegsman, L.M., 2001. Partial melting, partial melt extraction, and partial back reaction in
598 anatectic migmatites. *Lithos* 56, 75-96. [https://doi.org/10.1016/S1464-1895\(01\)00052-7](https://doi.org/10.1016/S1464-1895(01)00052-7)

599 Loomis, T.P., 1977. Kinetics of garnet granulite reaction. *Contributions to Mineralogy and*
600 *Petrology* 62, 1–22. <https://doi.org/10.1007/BF00371024>

601 Marone, C., 1998. Laboratory-derived friction laws and their application to seismic faulting.
602 *Annual Reviews in Earth and Planetary Science* 26, 643-696.
603 <https://doi.org/10.1146/annurev.earth.26.1.643>

604 Nyman, M.W., Pattison, D.R., Ghent, E.D., 1995. Melt extraction during formation of K-
605 feldspar + sillimanite migmatites, west of Revelstoke, British Columbia. *Journal of*
606 *Petrology* 36, 351–372. <https://doi.org/10.1093/petrology/36.2.351>

607 Nyström, A.I., Kriegsman, L.M., 2003. Prograde and retrograde reactions, garnet zoning
608 patterns, and accessory phase behaviour in SW Finland migmatites, with implications for
609 geochronology. In: Vance, D., Müller, W., Villa, I.M. (Eds). *Geochronology: Linking the*
610 *Isotopic Record with Petrology and Textures*. Geological Society of London, Special
611 Publications 220, 213-230. <https://doi.org/10.1144/GSL.SP.2003.220.01.13>

612 Perfilyev, V., Moshkovich, A., Lapsker, I., Laikhtman, A., Rapoport, L., 2014. Dislocation
613 structure and stick–slip phenomenon. *Tribology Letters* 55, 295–301.
614 <https://doi.org/10.1007/s11249-014-0358-4>

615 Persson, B.N.J., 1995. Theory of friction: Stress domains, relaxation, and creep. *Physical*
616 *Review B* 51, 13568. <https://doi.org/10.1103/PhysRevB.51.13568>

617 Pistone, M., Cordonnier, B., Ulmer, P., Caricchi, L., 2016. Rheological flow laws for
618 multiphase magmas: An empirical approach. *Journal of Volcanology and Geothermal*
619 *Research* 321, 158-170. <https://doi.org/10.1016/j.jvolgeores.2016.04.029>

620 Pistone, M., Baumgartner, L.P., Bégué, F., Jarvis, P.A., Bloch, E., Robyr, M., Müntener, O.,
621 Sisson, T.W., Blundy, J.D., 2020. Felsic melt and gas mobilization during magma
622 solidification: An experimental study at 1.1 kbar. *Frontiers in Earth Sciences* 8:175.
623 <https://doi.org/10.3389/feart.2020.00175>

624 Rabinowicz, M., Vigneresse, J.L., 2004. Melt segregation under compaction and shear
625 channeling: Application to granitic magma segregation in a continental crust. *Journal of*
626 *Geophysical Research* B109, B04407. <https://doi.org/10.1029/2002JB002372>

627 Reichhardt, C., Olson-Reichhardt, C.J., 2017. Depinning and nonequilibrium dynamic phases
628 of particle assemblies driven over random and ordered substrates: a review. *Report on*
629 *Progress in Physics* 80, 026501. <https://doi.org/10.1088/1361-6633/80/2/026501>

630 Rosakis, A.J., Samudrala, O., Coker, D., 1999. Cracks faster than the shear wave speed.
631 *Science* 284, 1337–1340. <https://doi.org/10.1126/science.284.5418.133>

632 Rubinstein, S.M., Cohen, G., Fineberg, J., 2004. Detachment fronts and the onset of dynamic
633 friction. *Nature* 430, 1005–1009. <https://doi.org/10.1038/nature08676>

634 Ruina, A. 1983. Slip instability and state variable friction laws. *Journal of Geophysical*
635 *Research* B88, 10359-10370. <https://doi.org/10.1029/JB088iB12p10359>

636 Rushmer, T., 2001. Volume change during partial melting reactions: implications for melt
637 extraction, melt geochemistry and crustal rheology. *Tectonophysics* 34, 342, 389-405.
638 [https://doi.org/10.1016/S0040-1951\(01\)00172-X](https://doi.org/10.1016/S0040-1951(01)00172-X)

- 639 Santa-Aho, S., Laitinen, A., Sorsa, A., Vippola, M., 2019. Barkhausen noise probes and
640 modelling: A review. *Journal of Nondestructive Evaluation* 38, 94.
641 <https://doi.org/10.1007/s10921-019-0636-z>
- 642 Stüwe, K., 1997. Effective bulk composition changes due to cooling: a model predicting
643 complexities in retrograde reaction textures. *Contributions to Mineralogy and Petrology*
644 129, 43–52. <https://doi.org/10.1007/s004100050322>
- 645 Svetlizky, I., Fineberg, J., 2014. Classical shear cracks drive the onset of dry frictional
646 motion. *Nature* 509, 205–208. <https://doi.org/10.1038/nature13202>
- 647 Vachon, R., Hieronymus, C. F., 2017. Effect of host-rock rheology on dyke shape, thickness
648 and magma overpressure. *Geophysical Journal International* 208, 1414–1429.
649 <https://doi.org/10.1093/gji/ggw448>
- 650 Vigneresse, J. L., Tikoff, B., 1999. Strain partitioning during partial melting and crystallizing
651 felsic magmas. *Tectonophysics* 312, 117-132. [https://doi.org/10.1016/S0040-
652 1951\(99\)00167-5](https://doi.org/10.1016/S0040-1951(99)00167-5)
- 653 Vigneresse, J.L, Barbey, P., Cuney, M., 1996. Rheological transitions during partial melting
654 and crystallization with application to felsic magma segregation and transfer. *Journal of
655 Petrology* 37, 1579-1600. <https://doi.org/10.1093/petrology/37.6.1579>
- 656 Vigneresse, J.L., Burg, J.P., 2005. Simulation of crustal melt segregation through cellular
657 automata: Insight on steady and non-steady state effects under deformation. *Pure and
658 Applied Geophysics* 162, 62, 987–1011. <https://doi.org/10.1007/s00024-004-2652-9>
- 659 Vigneresse, J.L., Cenki, B., Kriegsman, L. M., 2021. Melting by numbers: assessing the
660 effective melt fertility of crustal rocks. *Lithos* 386-387, 106006.
661 <https://doi.org/10.1016/j.lithos.2021.106006>

662

663

664 **Figure captions**

665 Figure 1. Model of melt extraction using a cellular automata (c) with minerals
666 distributed in cells. A melting equation (e) implies muscovite, biotite, plagioclase and quartz,
667 resulting in melt, Kfeldspar and garnet (undifferentiated ferro-magnesian). All minerals are
668 plotted according to a given colour. Melt production depends on the textural distribution, random
669 (r) or layered (l) and according to the chemical equation (e) of dehydration. Melting (m) occurs when
670 adequate minerals are adequately and proportionally distributed. Two thresholds (t) control liquid
671 percolation (LPT), i.e. melt motion, and melt extraction (MET). Particles motion may move under
672 pure shear (PS) or simple shear (SS) modifying the disposition of cells. Values with melt larger than
673 50, i.e. more melt than the matrix, are not considered.

674

675 Figure 2. Schematic representation of frictional law and its analogue for melt
676 extraction. a) Stationary friction of a block submitted to a force F. A coefficient of friction (μ)
677 is defined as the ratio of normal (σ) to tangential (τ) stress components. b) Dynamic friction.
678 The applied force varies from a velocity V_0 to V . The coefficient of friction presents a spike
679 ($\Delta\mu$) function of the relative change in velocity, with a return to stationary conditions after a
680 delay D_c . The new coefficient of friction may be higher or lower than the initial μ_0 . c) The
681 basal contact of the block is no more homogenous, with either two phases or asperities. An
682 irregular motion, with successive sticking and sliding states is observed, with amplitude (σ_μ).
683 A stiction spike is also observed, as well as return to stick-slip when the velocity varies. d)
684 Melt extraction coefficient (μ) defined as the ratio of initial melt to permeability (z) taken as a
685 melt escape threshold. e) Similar to friction, the initial melt (ϕ) varies, as well as the
686 permeability (ζ) and added stress (τ). A spike is observed before return to a stationary state. f)

687 irregular amount of melt is extracted mimicking stick-slip behaviour when the melt flow
688 presents reactions damping its amount and velocity.

689

690 Figure 3. Results of the numerical simulation for different values in the initial amount
691 of melt. a) The amount of melt extracted by step (σ) presents a stick-slip behaviour that
692 attenuates with the quantity of initial melt (ϕ_0). Stiction spike appears when this quantity is
693 suddenly changed. b) normalizing the amount of melt produced at each step by the initial
694 permeability (ζ) and dimension of the system results in curves converging toward a
695 coefficient of melt extraction (μ).

696

697 Figure 4. Two regimes, regular or stationary and transient, are observed depending on
698 whether the initial conditions vary during runs. Changing the melt proportion (ϕ_0) induces
699 spikes ($\Delta\mu$) and stick-slip motion with amplitude σ_μ . Spikes ($\Delta\mu$) in melt extraction with
700 iterations taken as a proxy for time (Dc), when the acceleration (ζ) is suddenly varied.
701 Changes are made at different iteration intervals, and shown with three initial melt percentage
702 (ϕ_0).

703

704 Figure 5. Variation of the amplitude of the stick-slip effects (σ_μ) as a function of the
705 initial melt (ϕ_0) when the permeability (ζ) varies. A weak linear variation shows the poor
706 effect of the initial melt content. Also indicated the range of values for melt (m) escape (in
707 pale grey) as estimated in nature.

708

709 Figure 6. a) Variation of the amplitude of the stick-slip effects (σ_μ) in function of the
710 melt extraction coefficient in two different situations, when the changes operate in separate
711 runs, or in succession during a single run. The changes concern the initial melt (ϕ_0), the
712 permeability (ζ) or the strain rate (ε°). The linear relationship indicates the dependence of σ_μ
713 on the coefficient of melt extraction (μ). It corroborates all variations with the equivalent
714 velocity parameters (ϕ_0 , ζ , ε°). The range of melt amount in nature (m) is indicated in light
715 grey. Conversely, the transient effects illustrate a neat decrease of σ_μ during the successive
716 variations of the equivalent velocity parameters, pointing to some memory effect. b)
717 Amplitudes of σ_μ plotted for individual values of ζ , show an exponential decrease with the
718 iteration numbers, taken as a proxy for time. It reinforces the memory effects, damping with
719 time.

720

721 Figure 7. Amplitude of the spikes ($\Delta\mu$) when the permeability value varies along a
722 single run. The successive changes are positive or negative, yielding similar changes in $\Delta\mu$.
723 The curves have been separated for increasing and decreasing permeability changes.

724

725 Figure 8. Experiments on stick slip on a single crystal of LiF, redrawn from Perfilyev et
726 al. (2014). a) The displacement (mm) as a function of time (s) shows the stick-slip behaviour,
727 following increases and sudden drop in frictional force (mN). Each episode occurs in about
728 180 s with a displacement of about 7 mm, implying a strain rate of about $5.5 \cdot 10^{-3} \text{ s}^{-1}$. b) The
729 corresponding velocity of the indenter varies by a factor about 300 in less than 1 s (note the
730 change in time scale between a and b). It is interpreted as depinning and pinning effect
731 corresponding to dislocation creep.

732

733 Figure 9. Cartoon illustrating the analogies between (a) frictional law, (b) melt
734 extraction and (c) poroelasticity. Friction implies solid-solid contact, i.e. crystalline networks
735 in a polycrystal material, with inherent dislocations (inverted tau) that act as pinning points
736 for motion. At the other end, poroelasticity describes a liquid flow, percolating through the
737 system, but with discrete air bubbles, non-percolating. The different surface energy values
738 between air, liquid and solid make the system hysteretic. They also rule pinning and
739 depinning when the local porosity changes due to the pore size. In between, melt extraction
740 consists in a polymerized silicate melt (tetrahedra) and undifferentiated ions (circles) flowing
741 in between crystallized matrix ongoing melting reactions. Minerals consist in quartz,
742 plagioclase, biotite and muscovite (white to black, with pale and dark grey), with incongruent
743 mineral (coloured hexagons), among which garnet (red) and sillimanite (stripes). Dehydration
744 melting occurs when the four quoted minerals are in contact (plain circle), feeding the melt
745 layer. They act as depinning points. In contrast, retrograde reactions manifest by garnet
746 coronas (dashed circles), consuming elements of the melt. They act as pinning points.

747

748

749

