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1 Multiphase rheology as a cause for stick-slip like melt extraction

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- 7
- 8 Highlights

9	0	Numerical model of melt extraction
10	0	Sporadic motion of melt and an analogue with dry friction (stick-slip)
11	0	Hysteresis and memory effects during changes in input parameters
12	0	Global expression of the free energy variation, with a term akin to velocity and a
13		second term to acceleration
14	0	Interactions between strain, chemistry and entropy (order)
15		

16 ABSTRACT

When partial melting is parametrized in numerical thermo-mechanical models in a
straightforward manner, it directly relates to variations in temperature. However, melt
extraction is, to our knowledge rarely taken into account in these models, even if melt
extraction and escape have a key influence on the mechanical behavior of the entire system.

Here, we propose a first step towards a numerical simulation for melt extraction in partially 21 22 molten portions of the continental crust. We use a Lagrangian description, considering a discrete mineral distribution at sample scale. A cellular automaton mimics the mobile melt 23 phase moving from cell to cell. An infinite source and sink for the melt and lateral periodic 24 boundaries avoid melt accumulation. Time is represented by a proxy through successive 25 26 iterations. The major controlling factors are the initial bulk melt percentage and a threshold 27 controlling melt escape. We test pure and simple shear conditions to simulate strain partitioning between the melt and its matrix. The parameters, provided they do not vary, 28 determine a regular or steady state amount of melt extraction after a few iterations. In 29 30 contrast, for varying parameters during a single run, the quantity of melt extracted at each step 31 presents alternate variations in amplitude (σ_{μ}) around an average value. Variations are neither periodic in the sense of a sinusoidal wave, nor of equal amplitude. Melt extraction behaves 32 similar to a stick-slip motion in dry friction. Since we cannot have access to the prediction of 33 34 melt amount, w test the origin of the sporadic motion. We examine variations of the input parameters (initial melt ϕ , permeability ζ , applied stress σ or τ). The transient effects mark by 35 spikes ($\Delta\mu$) with two contrasted situations. Return to a steady state is exponential, similar to 36 the discharge of a capacitor. Nevertheless, a first mode occurs when the changes are made on 37 separate runs, reinitializing the parameters between them. These experiments show a linear 38 39 dependence of the variation in amplitude of σ_{μ} with the parameters. Alternatively, when the changes are successive during the same run, then σ_{μ} decreases accordingly. The exponential 40 41 variation is interpreted as a memory effect or hysteretic behavior of melt extraction. Similarly, 42 the amplitude of the stiction spikes ($\Delta\mu$) does not show the same behavior when permeability increases or decreases. By analogy with friction, the change from sticking to slipping is 43 44 interpreted as a manifestation of pinning and depinning, leading to an avalanche-like phenomena. Mechanically they correspond to a change in friction velocity. Physically, it 45

46	points to the cha	inge in the free energ	gy of the system	n, with a f	irst derivati	ve akin to velocity
47	and a second derivative akin to acceleration. During melting, the mechanical barriers					
48	correspond to the energy variations due to chemical reactions between melt and crystals.					
49	Depinning corre	sponds to enhanced	melting wherea	as pinning	links with	oack reactions
50	leading to garne	t growth. The sporad	lic melt extract	ion links c	lry friction	to poroelastic flow,
51	or more general	ly to a consequence	of two-phases r	heology.	Hence the s	poradic motion of the
52	weak phase orig	inates from interacti	ons with its ma	trix. Henc	the flow (velocity) suffers
53	modulations due to interactions with chemical potential (back-reaction, melting), entropy			nelting), entropy		
54	(ordering throug	h dislocation creep)	, and strain (apo	erture of p	orous flow)	. Such interactions
55	should be consid	lered when dealing v	with quantitativ	e mush (n	nelt+ matrix) modeling
56	(rejuvenation, ex	xtrusion).				
57						
58	Graphical Abst	ract				
59						
60	Keywords: stick	<-slip; melt extractio	n; friction; hyst	teresis; pii	nning/depin	ning
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72 **1.** Introduction

73 Temperature rules melt generation in the crust, most often through heat diffusion, or advection, which are both continuous processes. The extraction of melt out of its matrix is by 74 75 contrast discontinuous, initiating from melt pockets that aggregate to form veins to escaping 76 and feeding isolated plutons (Kisters et al., 1998). According to such field observations, the paradigm of melt extraction and transport turned from diapirism or dyking to discrete solitary 77 78 waves (Connolly and Podladchikov, 2000; Rabinowicz and Vigneresse, 2004) that propagate from themselves, driven by alternate changes in permeability. Such models of porous flow 79 also apply to propagating fluids, melt or water within a stressed crust, with applications to 80 81 magmatism, geothermic and water circulation into the basement. They form the basis of poroelasticity (Cheng, 2016). The paradigm switch from a viscous material more mobile into 82 a viscous matrix to a problem of permeability suggests that hydraulic parameters should rule 83 crustal fluid transport. 84

85 Modelling melt extraction thus points to the rheology of multi-phase material, i.e. the 86 differential response of a mobile fluid circulating into a deforming matrix. The viscosity contrast between the melt and its matrix i.e. a weak and a strong phase, may be apprehended 87 from several points of view including estimates of an effective viscosity, intermediate 88 89 between the two end-members (Pistone et al., 2016; Agard et al., 2018). Three factors then 90 add complexity to such a direct approach. The first is the dependence of the viscosity of the 91 strong phase on the applied strain rate whereas the melt is simply Arrhenian. This makes the 92 experimental studies hard to be transposed to natural conditions (Pistone et al., 2016) because the fast strain rate imposed by the duration of experiments. The second factor is strain 93 94 partitioning (Vigneresse and Tikoff, 1999) between the weak and the strong phase (Goodwin and Tikoff, 2002). Their respective abundance determines which phase globally controls the 95 effective rheology. When the weak phase is abundant, the rheology is governed by the 96

distribution and geometry of rigid blocks (Agard et al., 2018; Ioannidi et al., 2021) as 97 98 observed at the interface of subduction zones. In contrast, when the weak phase is in minor amount, the strong rigid matrix controls the bulk rheology (Vachon and Hyeronimus, 2017). 99 100 The two approaches can be considered as a variant of a direct problem, overdetermined, and always offering a solution. However, they do not tackle the third factor, which is the exchange 101 of matter between the melt and its matrix, forming feedback loops. Hence melting occurs 102 103 when adequate minerals are close together, thus lowering the energy of melting. Conversely 104 back reactions, especially involving ferro-magnesian minerals, precipitate elements, consuming energy to form incongruent minerals (Cenki et al., 2002; Nyström and Kriegsman, 105 106 2003). The two reactions act in modifying the free Gibbs energy budget of the reactions. However, their extent, i.e. the amount of material involved that rules the energy budget cannot 107 be controlled exactly in time and space. It thus lowers the accuracy of a direct modelling. 108 To partly avoid and solve these ambiguities, a cellular automaton has been constructed 109 and modified (Vigneresse and Burg, 2005; Vigneresse et al., 2020), giving insights to the 110 111 mechanisms of two-phase rheology in partially molten rocks (see below). The process of melt fusion and extraction can be directly compared with frictional motion, 112 that rules displacement of two materials in contact under an applied force (Persson, 1995; 113 Marone, 1998). Friction characterizes the relative motion, usually between two solid phases 114 (Byerlee, 1978). It links the tangential (τ) to normal (σ) forces (Fig. 1). The former controls 115 motion imposed by an applied force, whereas the second akin to the gravitational force resists 116 117 to lateral motion. The static coefficient of friction (μ_0) is a linear function of the applied forces, and it is constant for most materials (Byerlee, 1978) (Fig. 1a). The irregularities of the 118 119 contact make lateral motion sporadic in time. A kinetic coefficient of friction (Dieterich, 1979; Ruina, 1983) quantifies the derivative of the displacement, akin to a velocity. 120

Irregularities, or asperities in the planar surface separating the two bodies resist to motion. 121 122 The frictional motion becomes transient owing to the change in the velocity of the motion (Fig. 1b). Stick-slip is a special case of motion in which a prolonged stagnation (the "stick" 123 phase) is interrupted by short bursts of movement (the "slip" phase). The motion is not 124 periodic in the sense that the time interval between the short bursts is not always identical, 125 depending on the resisting forces (Fig. 1c). The amplitude of the bursts that vary around an 126 127 average value are not identical neither. When the initial conditions locally vary, a transient motion develops with sudden acceleration in motion. Such transient effects are not well 128 understood, though they represent the responses of the system to the initial parameters 129 130 variations

While the discontinuous nature of melt extraction could be easily been transposed from 131 discrete permeability waves, the transient effects are not intuitively transposable to 132 displacements between viscous bodies or to poroelasticity. Hence, the discontinuous nature of 133 seismicity for example, yield to the concepts of asperity, small irregularity in the contact 134 135 between materials. They represent pinning points restraining motion and the sudden change in contact area, or depinning leads to sliding. Following this model, frictional motion manifests 136 by alternate sticking and sliding phases, known as stick-slip motion (Baumberger and Caroli, 137 138 2006; Brace and Byerle, 2008).

Illustrating and interpreting frictional motion between a fluid, melt or water, and its matrix represents an example of alternate concepts between rheology and hydraulics applied to magmatism. The starting point of this study is a numerical simulation of melt extraction using a cellular automaton (Vigneresse and Burg, 2005). The origin of a sporadic motion is examined in relation with the initial parameters and internal changes in free energy through the interactions between its three components (entropy, strain and chemical potential).

2.

Numerical modelling set up

A numerical modelling using a Lagrangian formulation has been built from a cellular 147 automaton modified from a preceding model (e.g. available code in Vigneresse and Burg, 148 149 2005). The initial system consists in discrete cells, organized in a hexagonal pattern with dimension 64 by 38 cells (Fig. 1). Cells contain indifferently melt and solid in proportion 150 varying from 0 to 100 (Vigneresse et al., 2021). The solid phase involves quartz and 151 152 plagioclase, two micas as muscovite and biotite, and some incongruent or residual ferromagnesian minerals. The initial mineral and melt proportions are fixed, with a textural 153 disposition, either random, in layers or small pockets (Vigneresse et al., 2021). A 154 155 stoechiometric equation rules dehydration melting (Gardien et al., 2000). Similarly, back reactions are introduced for the ferromagnesian (Cenki et al., 2002). The melt sink is infinite, 156 as well the heat source. Those two conditions assume a stationary condition for melt 157 evolution. The remaining solid phase is replaced by the moving melt from one cell to the next, 158 according to simple and/or pure shear geometry. Shear strain localization is possible by 159 160 moving melt to more than one cell. Periodic lateral boundaries avoid melt accumulation. The initial melt content can be modulated, or by partial refilling in melt mimicking temperature 161 burst. Superimposed to the initial conditions and geometric parameters, two rules control melt 162 163 motion and extraction (Vigneresse et al., 1996). In addition to a liquid percolation threshold (LPT) a melt escape threshold (MET) controls extraction. The effect of stress and strain rate is 164 also examined in both situations of pure (σ) and simple (τ) shear. However, the adopted 2D 165 geometry biases the results since melt may escape the system in the third direction. The time 166 167 is represented by an iterative process. As in most cellular automata, that are qualitative by essence, the time aspects are only apprehended through iterations, i.e. without exact 168 calibration in time. 169

171 **3.**

Results concerning melt extraction

172 Results allow to identify the effects on melt extraction of the main parameters: i.e. the 173 initial amount of melt (ϕ_0) and the threshold for extraction (MET). The former is supposed to 174 overcome melt percolation through the system. The latter is equivalent to fluid draining, thus 175 renamed as ζ (Fig. 2d, e). All parameters vary on a scale from 0 to 100. The extreme values of 176 met (> 50), i.e. with more melt than matrix, are not representative of nature, but allow a 177 control on the evolution of the numerical model.

Bulk results show a stationary state when there is no significant change in the input parameters, but a spike in melt extraction occurs when the ζ value changes, and more commonly when sporadic extraction occurs. The three modes (stationary, spike and sporadic) show a good analogy with frictional motion, pointing to interpretation with the changes in friction, i.e. in terms analogue to a frictional law (e.g. Marone, 1998).

183 3.1. Preliminary assessment of the results

Melt extraction is scaled in relative quantity of melt extracted versus iterations, i.e. time, without possibility of more precise scaling (e.g. hours or years) yielding a relative time scale. Similarly, the energy budget of the process is not quantitively available.

187 The amount of melt is another parameter used during the computation. The possible amount of melt in each cell ranges from 0 to 100. Consequently, the extracted melt fraction 188 (σ) at each step also ranges from 0 to 100 (Fig. 3). The fraction of extracted melt per step 189 shows, as expected, a spike at the start of the model, that rapidly converges toward a 190 stationary value. The spike is positive when the initial melt fraction is high, negative when 191 192 melt fraction is low. Curves with different initial melt fractions converge after about 70-80 steps. In a second stage, the percentage of melt extracted (σ) at each step is normalized to the 193 initial melt (ϕ_0) and the dimension of the model (Nx * Ny), providing the coefficient of melt 194

195 extraction (μ) (Fig. 3). After normalization, the value converges after about 150 steps toward 196 a value ranging 0 to about 1.5 (Fig. 3) increasing with the initial melt (ϕ_0). In this example, the 197 linear relation between μ and ϕ_0 defines a bulk coefficient of melt extraction (μ_0) that equals 198 0.17 ± 0.002 whatever the initial melt percentage, and using a value of 0.25 for ζ assuming a 199 one order less variation of (μ_0) as a function of ζ (Vigneresse and Burg, 2005).

A note of caution seems necessary, here, prior to interpreting the results. The coefficient of melt extraction reaches a value of 1.0 for ϕ_0 larger than 50 and values higher correspond to more melt extracted than produced. It means that the model diverges for high values of initial melt, i.e. when there is more liquid than solid. Such a situation, mathematically possible, does not reflect reality. However, it allows examining the results in a larger range of parameters. We have indicated on the figures 6 to 7 the range of values for melt extraction, as estimated in nature.

207 The coefficient of melt extraction (μ) shows two behaviours, regular or stationary, and 208 transient (Fig. 4), depending on whether the other parameters are constant or not during a 209 single run. Both cases are examined separately. The first case leads to a steady state regime. 210 Conversely, the transient mode due to varying ζ , or strain rate (ε° and γ°) induces a spike of 211 amplitude $\Delta\mu$.

The return to the steady state values follows a classical damping effect reflecting an exponential decrease of the sharp increase with a characteristic time (Dc) about 70-80 steps (Fig. **4**). The lack of definition in time (i.e. iterations) precludes a more detailed investigation on the variability of Dc. The spike is positive or negative, depending on the lack or enhanced quantity of melt. It thus, like a charge and discharge of a capacitor, requires using a logarithmic function, with the sharp increase normalized to the initial value. The function 218 takes the form of a $\ln(\phi/\phi_0)$. Such functions are applied to changes in ζ and strain rate (ε° and 219 γ°)

220 3.2. Towards a law of melt extraction

The preceding diagrams allow quantifying the relative variations as a function of the input parameters. The results can be summarized as follow and interpreted by a simplified law of melt extraction that takes into account the linear and non-linear effects, associated to the potential barriers of melt extraction mimicking a frictional law (Vigneresse and Burg, 2005):

225
$$\mu = \mu_0 + a (\phi/\phi_0) + b (\zeta/\zeta_0) + c (\sigma/\sigma_0) + d \ln (\delta\phi/\phi_0) + e \ln (\delta\sigma^{\circ}/\sigma^{\circ}_0)$$

with the numerical values $\mu_0 = .022$, a = .0018, b = -.0015, and c = .016 (simple shear) or = .066 (pure shear) (Vigneresse and Burg, 2005). The two functions in $\delta \phi$ and stress rate ($\delta \sigma^{\circ}$) variations are non-linear and could not be determined quantitatively. Globally, the effective production of melt ranges from 15 % to a maximum of 40 % depending on protolith composition (Vigneresse et al., 2021). The potential barriers represent variations one order of magnitude less than the coefficient of melt extraction μ_0 .

It should be compared to the suggested frictional laws (Dieterich 1972), with the modificationabout displacement (Ruina 1983).

234
$$\mu = \mu_0 + a \ln(V/V_0) + b \ln(V_0 \theta / D_c) \text{ and } d\theta/dt = 1 - V\theta/D_c$$

Nevertheless, because of the poor access to time, the problems of strength recovery (healing) interpreted from the state variable θ , without more specification (Ruina, 1983), have not been considered in the actual study. The analogy points to the effect of velocity, either as the velocity of traction during dry friction, or melt flux under strain partitioning on the viscous melt, eventually enhanced by acting stress. In addition, a second term related to imposed strain is present in the equations, that may have an influence on the sporadic motion during friction or melt extraction. In a first approach, this term is a derivative of the motion, akin to acceleration. It suggests that melt motion is represented by the relative value of melt volume (ϕ/ϕ_0) , whereas the remaining other parameters (ζ , strain) act as enhancing/damping the flow.

244

245

4.

Sporadic melt extraction

One difficulty in interpreting the irregular sporadic melt extraction patterns mostly links with the poor determination related to temporal calibration. It includes difficulties in a true scaling in time units, but also difficulty in interpreting a discontinuous diagram of irregular amount of extracted quantity of melt. It manifests during friction by a stick-slip motion (Fig. **2c**), as a second order manifestation, interpreted by the hysteretic nature of friction. Hence, friction is a dissipative function of energy, that needs additional energy to recover the initial state. We provisionally adopt the term stick-slip behaviour for describing melt extraction.

253 4.1. Variations in stick-slip amplitude (σ_{μ})

Stick-slip manifests during friction by alternate periods of no or restricted motion, or 254 stick phase, followed by slip at constant velocity (Fig. 2c). Similarly, the amplitude of melt 255 extraction, varies with iterations, but remains constrained between two values defining a 256 dynamic and static extraction modes (Fig. 2f). The difference is referred to as σ_{μ} . This value 257 is examined as a function of the other parameters. The amplitude of σ_{μ} varies linearly 258 between 0.0015 and 0.013 for an initial melt content varying from 5 to 85 %, though the last 259 value is unrealistic in nature, but brackets the range of variables. It slightly varies by about 260 0.003 for a variation of 50 in the initial melt content (ϕ_0), i.e. from dry to saturated melt. The 261 values increase with ζ by about 0.035 for a change of 20 in ζ (Fig. 5). Thus, the effect is more 262

263 pronounced depending on ζ rather than on ϕ_0 . The linear effect also excludes artifacts such as 264 errors or noise. It correlates with a non-strictly periodic effect in amplitude.

The very moderate decrease, about one order less than the coefficient of melt extraction with the initial melt content points to an enhanced melt extraction with the total melt content (Fig. 5), suggesting a possible lubrication of the interface between melt and its matrix. Nevertheless, this effect remains weak compared to the effect of ζ . It also points to enhanced/damping effects occurring during transient changes compared to the variations induced by varying the initial parameters.

271

4.2. Variation of amplitude depending on the system history

The amplitude increase of the sporadic melt extraction has been tested against the other parameters evidenced in the melt extraction law (ϕ , ζ , ε°) (Vigneresse and Burg, 2005). This has been realized under two modes. In a first set of runs, these parameters remain fixed along the whole run. The separate runs (previous sections) show moderate variations in σ_{μ} , with a bulk increasing linear trend whatever the change in all parameteers (Fig. **6**).

Alternatively, a second set of runs, that are successive, allows abrupt changes in the 278 parameters during iterations, indicated by enhanced melt extraction spikes before a 279 280 progressive return to the previous extraction mode. Those spikes, with amplitude $\Delta \mu$, are positive or negative depending on charge or discharge of melt. By contrast, the amplitude of 281 melt extraction (σ_{μ}) shows a pronounced decrease, for all parameters ($\phi, \zeta, \varepsilon^{\circ}$), when one of 282 them is kept constant (Fig. 6). The decrease in amplitude is significant (0.12 to 0.0), since σ_{μ} 283 tends to zero when the changes in input parameters modify the coefficient of melt extraction 284 by a factor 2 or 3. In addition the variation in amplitude for separate runs largely overcomes 285

3065.Implications and discussion

The large variability observed in the modelled melt extraction has some directconsequences.

310 5.1. Hysteresis

Hysteresis is the dynamic time lag between input and output of the system, reflecting its 311 312 dependence on its previous history. It is associated with the need to consume or deliver additional energy for the return to initial conditions. Examples of hysteresis in geological 313 314 systems include remanent magnetism, plastic deformation, phase change during melting and 315 crystallization, or surface tension (Persson, 1995; Hill, 1998). A simplistic representative 316 example is a drop of water on a glass. In case of a porous medium, with a two-phase flow (air and water) the curve of capillary pressure versus the wetting phase saturation also describes a 317 318 hysteretic cycle, with a wetting stage, and return along drying curve. Due to the solubility of the gaseous phase, the hysteretic curve shrinks with successive cycles of drying and wetting. 319

320

321 5.2. Pinning and depinning

The sporadic melt extraction, with alternate bursts, links to variations in the melt 322 velocity, i.e. de-/acceleration. As stated in friction, it relates to a state variable θ (Ruina, 323 1983), essentially defined by its derivative with time ($d\theta/dt$). It is characterized by alternate 324 episodes of lock (stick) and motion (slip). The frictional interpretation commonly relates to 325 those stages to surface conditions. Indeed, they present asperities, or small stronger areas of 326 relief, arresting the motion. If the interpretation is correct, the alternate stages of 327 enhanced/reduced melt extraction should also be controlled by such points of resistance, i.e. 328 329 pinning and depinning, and their equivalent in terms of chemical thermodynamics is consuming/releasing Gibbs free energy. 330

331 Experimental studies on stick-slip, using small transducers and fast camera recording332 allow a better understanding of the phenomena traduces in a force vs time diagram during dry

friction (Fig. 8a). The equivalent diagram in displacement vs time indicates periods of no 333 334 motion corresponding to the stick stage followed by a sudden jump in velocity (Fig. 8a). However, more sensible experiments from the group of Jay Fineberg (Hebrew University, 335 336 Jerusalem) tackled the fine motion during the sticking stage in a bi-phase material interface. They record high-speed measurements of the real contact area, whereas the strain field is 337 338 measured in the region surrounding propagating rupture tips (Svetlizky and Fineberg, 2014). 339 They observe highly localized slip pulses accompanied by a large local reduction of the normal stress near the rupture tip. The pulses propagate along the direction of motion of the 340 softer material at a maximal velocity, evolving while propagating. When the rupture front 341 342 velocity overcomes the seismic S-waves velocity, it determines a supershear front (Rosakis et al., 1999) that propagates in the opposite direction, suggesting crack-like rupture. Such results 343 corroborate the ideas of dislocations steadily propagating along a material interface. In 344 345 parallel, the transition from static to dynamic friction is quantitatively described by classical singular solutions for the motion of a rapid shear crack (Rubinstein et al., 2004). They 346 347 observed propagations at sub-Rayleigh velocity, the sound speed along a free surface, but also a slower front that could be the trace of healing at contact area, before slip arrest. This is 348 similar to ageing that immediately commences as contact area increases. Its characteristic 349 350 timescale is determined by the local memory of the system, according to its effective contact time before slip arrest (Ben-David et al., 2010). 351

Such evolution at the microscale of dislocations has been surveyed in detail on a single LiF crystal since it allows simultaneous observation of evolution of the dislocation structures through the Bauschinger effect, a piling-up of dislocations at grain boundaries. Indeed, the microscopic distribution of stress alters the stress/strain characteristics of the material (Harea et al., 2013). This property is used to trace the dislocation pattern during deformation (Perfilyev et al., 2014). Observations show that the indenter velocity increases nonlinearly

during the transition from stick to slip. in the mid-part of the slip phase, the indenter reaches 358 its maximum velocity, while the friction force decreases abruptly (Fig. 8b). The average 359 velocity of the indenter in the stick phase drops by about 300 times than the maximum 360 361 velocity in the slip phase. It drastically affects the behaviour of the dislocation regime. Three stages of creep are evidenced between two slip events (Perfilyev et al., 2014). First, a primary 362 creep links with the formation of bonds (pinning effect). Their number increases during the 363 364 second long-time steady state. In contrast, the tertiary creep is characterized by the rupture of bonds (depinning effect). 365

The link between pinning/depinning and dislocations (Anderson et al., 2017) brings some highlights on the asperity model since it marks by acoustic emission that may be recorded during the stick-slip motion. Known as the Barkhausen noise in ferro-magnetism (Santa-Aho et al., 2019) it results from the rapid size variations of magnetic domains, with their magnetic spins equally oriented. The produced noise is recorded and is currently in use to detect defects in the alloys.

372

373 5.3. Implications on melt extraction processes

374 The preceding section provides mechanical examples involving the role of dislocations in pinning/depinning effects during frictional stick-slip motion (Fig. 9a). 375 376 Dislocations creep correspond to a change in the ordering of a crystal, i.e. a change in entropy. It may apply to solid-solid contacts but fails when a viscous material is present 377 (Marone, 1998). Another model of stick-slip behaviour derived from poroelasticity (Cheng, 378 379 2016) takes into account the surface tension and the solubility between the components, solid, liquid and air (Fig. 9c). I corresponds to an interaction between strain of the flow and strain 380 imposed by the matrix. In both cases a locking effect, or barrier, retards motion while an 381

acceleration occurs when additional energy overcomes the resisting forces. In both cases, the processes are the sum of a static equivalent to the free energy of the system plus a term function of the its first derivative, equivalent to velocity, and a second derivative, equivalent to acceleration. The energy barrier linked to dislocations has a fixed amplitude, also linked to a progressive increase in pinning, thus favouring a limit to a dynamic friction coefficient.

In between those two end-members, melt extraction results from a viscous-viscous 387 388 interface. There is no dislocation, by lack of a true crystalline network, and there is no nonpercolating third phase as air bubbles. Hence the matrix is a crystalline rock undergoing 389 melting, but with a heterogeneous amount of solid phases either products or unreacted 390 391 inherited minerals. The melt is a mixture of silicate polymers and ions in a supercritical state. Metamorphic reactions between the melt and its matrix may be considered as a cause of for 392 pinning/depinning effects (Fig. 9b). Retrograde, melt-consuming reactions, producing solid 393 phase are the counterpart of prograde melting and may produce new minerals phases acting as 394 asperities causing stick-slip or sporadic melt extraction. Both type of reactions are commonly 395 396 observed in migmatites, rocks that have undergone partial melting (Stüwe, 1997; Kriegsman, 2001). 397

Dehydration of hydrous minerals occurs during prograde melting, mainly muscovite 398 399 above 730 °C and the biotite around 800°C (Clemens, 1990). Hence the leucocratic portions of migmatites, or leucosomes, are mostly composed of quartz, feldspars and rare micas. These 400 401 are interpreted as crystallized melt usually of granitic composition. The darker ferromagnesian minerals form the melanosome and may segregate aside the leucosomes by 402 flow differentiation or reactive transport (Kriegsman, 2001). The minerals from the 403 404 melanosome can either be relics of reactants, products of melt consuming reactions or newly formed incongruent mineral phases. 405

During muscovite breakdown with temperature increase, mineral phases disappear 406 407 whereas melt-consuming reactions may form new solid phase upon cooling (Kriegsman, 2001). Those may be quartz, plagioclase and muscovite (Nyman et al., 1995; Konn et al., 408 409 1997) followed by biotite, sillimanite, quartz and plagioclase (Cenki et al., 2002). In contrast garnet and cordierite may grow as incongruent porphyroblasts (Nyström and Kriegsman, 410 2003). During growth, they may incorporate large quartz rounded inclusions indicating 411 412 surface tension and immiscibility. Upon cooling, retrograde phases may form coronas around these porphyroblasts, surrounding by the leucosome, illustrating back reactions between the 413 restitic phases and the melt. Ti-rich biotites also grow at the expense of ferro-magnesian 414 415 incongruent phases (Cesare et al., 2008). Melt consuming reactions include mica growth, finegrained intergrowths in micas, breakdown of incongruent mineral phases and coronas 416 formation (Cenki et al., 2002; Cesare et al., 2008). The amount of melt production, or 417 418 consumption, depends on the ability of elements to enter such reactions. It thus has no bounding values, making the amplitude of stick slip non periodic. 419

Interestingly, the formation of garnet coronas through a process of diffusion is slow.
Thus, depending on the heat flow variations, a so-called tectonic "hysteresis" in the melt
extraction may result (Loomis, 1977). The retrograde reactions, added to changes in the pore
size, determine zones of pinning and depinning (Fig. 9b).

The sporadic mode of melt extraction corroborates with the observation of segregated melt from its matrix in discrete pockets (Kisters et al., 1998). Such discrete pockets commonly form shrinking patterns with subvertical pipe-like bodies. They further connect together to evacuate a felsic melt into veins following the regional foliation planes. Discrete melt batches also corroborate numerical models of porosity waves (Connolly and Podladchikov, 2000; Rabinowicz and Vigneresse, 2004).

Another conclusion from such competition between melting and solidification (back-430 431 reactions) during melt extraction may have implications on the interpretation of field samples. Hence, magma crystallization is commonly coeval with fluids release, due to space shrinking, 432 rapidly reaching fluid saturation (e.g. second boiling, Candela, 1997). The additional presence 433 of a compressible fluid phase within the competing enhanced melting and crystallization (i.e. 434 external pressures variations on the system) is coherent with an variation in the extraction rate 435 436 during solidification with sudden pressure changes (Pistone et al., 2020). The stick-slip like melt motion thus adds a supplementary positive feedback loop to the process. Conversely, 437 during initial melting, the freed fluid from hydrous minerals does not develop in a separate 438 439 phase since it immediately enters the melt because of its solubility. In addition, the poor volume change due to biotite dehydration (Rushmer, 2001) does not alter the bulk volume of 440 the system, thus avoiding pressure related loops. 441

Hysteresis and differences in melt motion increase the symmetry breaking between 442 melting and solidification. In particular, the order of minerals crystallization, commonly 443 accepted from ferro-magnesian, plagioclase and quartz, i.e. following a decrease in chemical 444 potential, is perturbated when fluids (water-dominated) enter the process. It has implications 445 on the segregation or not of incompatible elements that do not follow a similar trend 446 compared to congruent minerals. The second important difference between melting and 447 crystallization relates to the fate of water, inducing a large scale more or less independent 448 hydrothermal activity during magma crystallization, whereas the processes of dry or hydrous 449 450 melting imply changes in the melt final composition. However, developing such ideas is out of the scope of the present paper. 451

452 5.4. Analogy between dry friction and melt extraction

453	The dry friction coefficient (μ_d) characterises the relationship between a force (F) applied to a
454	mass (m) pulled by a spring. The usual relation between the force applied on a mass and its
455	acceleration applies. The force is proportional to the first derivative of the motion (dx/dt)
456	through the coefficient of friction (μ_d). Thus
457	$m \ \delta^2 x / \delta t^2 = - \ \mu_d \ \delta x / \delta t$
458	The solution to this differential equation is given by
459	$\delta x(t)/\delta t = \delta x(0)/\delta t \exp \left[-(\mu_d / m) t\right]$
460	It relates the mass-force relationship to the stress-strain relationship. It also points the double
461	effect of the velocity and its variations (i.e. acceleration), thus providing an explanation to the
462	"state variable" (θ) introduced in dry friction (Ruina, 1983). In terms of melt extraction, it
463	corresponds to variation in the melt flow, either due to enhanced motion under applied stress
464	or enhanced/retarded melting.
465	
466	5.5. Notes of caution
467	The ideas presented here are derived from the numerical model that examines the motion
468	of melt in migmatites. As such, it should be considered as a model, i.e. either representing
469	the effects of the identified parameters, or in contrast, as a way to bracket the values of the
470	parameters to better fit field observations. This requires a note of caution on the way the
471	results are interpreted. In its actual shape, all parameters have been varied in a full range.
472	For instance, the initial amount of melt ranges 0 to 100, though such high values are
473	obviously not observed in partially molten rocks. However, the extreme values are
474	considered as providing the numerical trends that rule the underlying process. Similarly, the

475 coefficient of melt extraction presents values greater than 1. It corresponds to the effect of an

476 infinite source and sink for the melt. They are not realistic. Such large values correspond to
477 an initial melt larger than 50. Such notes of caution being stated, what are the main results of
478 the model.

First, the model illustrates the role of sharp variations in the ability of melt to be extracted and move (ζ), formerly assimilated to a melt escape threshold (Vigneresse and Burg, 2005). These variations show a transient behaviour. It implies that the definition of the melt escape threshold as a purely geometric threshold corresponding to percolation is incorrect since it is essentially static. Hence, the ζ parameter is dynamic, determining stiction spikes ($\Delta\mu$) and the amplitude of the stick-slip motion (σ_μ). It corresponds to a variation of the flux, i.e. an acceleration, positive or negative, of the flux.

Secondly, chemical reactions develop between the matrix and the melt as prograde 486 reactions, but also between the melt and its matrix, as retrograde reactions. They consume or 487 produce melt, thus modifying the melt flow velocity. Such reactions of chemical exchange 488 are important when considering the free energy balance of the system. the corresponding 489 490 interactions between strain (flow velocity) and chemical potential (melting and backreactions) modify the bulk energy of the system, inducing sporadic variations. Because they 491 were not considered in other forward models (Ioannidi et al., 2021), they resulted is a 492 stationary motion, that certainly does not fully consider the exact nature of two-phase 493 rheology. Similarly, variations due to the energy required to overcome dislocations induces 494 495 instabilities due to strain (flow) and entropy due to ordering modification (dislocation creep) as in dry friction (Fig. 9a). In poroelastic flow, the interactions develop between strain 496 497 (flow) and changes in the pore aperture (external strain) determine sporadic motion (Fig. 9c). The specific case of melt extraction implies chemical exchanges (melting and back 498 499 reactions) between the melt and its matrix (Fig. 9b).

500 Thirdly, the transient character of the flow presents some memory effects, due to the 501 irreversibility of the reactions and the thermodynamic metastability of some minerals 502 phases. However, such memory effects should not be confused with a time-related effect. In 503 particular, it does not implicitly to point any re-use of veins.

504 At last, because of the model is constructed from a cellular automata, the model cannot be scaled in time, but uses a proxy through iterations. It is therefore not aimed to prediction, 505 506 e.g. for quantitative mush (crystal-melt) interaction, as described for magma chamber rejuvenation or in magma eruptability. Similarly, a purely chemical approach is no more 507 508 predictive, and a dry friction modelling does not inform on the next earthquake to occur. In 509 contrast, and we focus on the message here delivered, any modeling of a multi-phase material should incorporate the interactions between the three components of energy, i.e. the 510 external variables such as entropy, strain and chemical potential. 511

512

513 **6.** Conclusions

The present model examines the transient flow of melt extraction depending on variations 514 of the controlling parameters, mostly the changes in flux velocity (amount of melt). The 515 occurrence of spikes reflecting sporadic melt extraction with a progressive return to a 516 stationary state is remarkable. We interpret these results with an analogy to the stick-slip 517 518 behaviour in mechanics that might develop resulting in an irregular, non-periodic amount of 519 melt extracted. This effect shows hysteresis, essentially due to the presence of phases 520 consumed during prograde (melting) or retrograde (garnet and sillimatite growth) metamorphic reactions. The interactions between strain, chemical potential and entropy 521 522 modify the free energy of the system. They act as pinning and depinning during dry friction and other flows of a weaker phase within a stronger matrix. The model highlights a dynamic 523 concept of melting and places melt extraction between dry friction and poroelasticity. 524

525

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541 **References**

- 542 Agard, P., Plunder, A., Angiboust, S., Bonnet, G., Ruh, J., 2018. The subduction plate
- 543 interface: rock record and mechanical coupling (from long to short timescales). *Lithos* 320-
- 544 321, 537-566. https://doi.org/10.1016/j.lithos.2018.09.029
- Anderson, P.M., Hirt, J.P., Lothe, J., 2017. *Theory of Dislocations*. Cambridge University
 Press, 718 pp.
- Baumberger, T., Caroli, C., 2006. Solid friction from stick–slip down to pinning and aging.
 Advances in Physics 55, 279-348. https://doi.org/10.1080/00018730600732186
- 549 Ben-David, O., Rubinstein, S., Fineberg, J., 2010. Slip-stick and the evolution of frictional
 550 strength. *Nature* 463, 76-79 https://doi.org/10.1038/nature08676
- Brace, W.F., Byerlee, J.D., 2008. Stick-slip as a mechanism for earthquakes. *Science* 153,
- 552 990-992. https://doi.org/10.1126/science.153.3739.990
- 553 Byerlee, J.D., 1978. Friction of rocks. *Pure and Applied Geophysics* 116, 615-627.
- 554 https://doi.org/10.1007/978-3-0348-7182-2_4
- Candela, A., 1997. A review of shallow, ore-related granites: textures, volatiles, and ore
 metals. *Journal of Petrology* 38, 1619–1633. https://doi.org/10.1093/petroj/38.12.1619
- 557 Cenki, B., Kriegsman, L.M., Braun, I., 2002. Melt-producing and melt-consuming reactions
- in anatectic granulites: P–T evolution of the Achankovil cordierite gneisses, South India.
- Journal of Metamorphic Geology 20, 543-561. https://doi.org/10.1046/j.1525-
- 560 1314.2002.00388.x
- 561 Cesare, B., Satish-Kumar, M., Cruciani, G., Pocker, S., Nodari, L., 2008. Mineral chemistry
- of Ti-rich biotite from pegmatite and metapelitic granulites of the Kerala Khondalite Belt
- 563 (southeast India): Petrology and further insight into titanium substitutions. *American*
- 564 *Mineralogist* 93, 327-338. https://doi.org/10.2138/am.2008.2579
- 565 Cheng, A.H.D., 2016. *Poroelasticity*. Springer Heidelberg, 903 pp.
- 566 Clemens, J.D., 1990. The granulite-granite connexion. In: Vielzeuf, D., Vidal, P. (Eds)
- 567 *Granulites and Crustal Evolution*. NATO ASI Series C, 311. Springer, Dordrecht. 25-
- 568
 36. https://doi.org/10.1007/978-94-009-2055-2_3
- 569 Connolly, J.A.D., Podladchikov, Y.Y., 2000. Temperature-dependent viscoelastic compaction
- and compartmentalization in sedimentary basins. *Tectonophysics* 324, 137-168.
- 571 https://doi.org/10.1016/S0040-1951(00)00084-6.
- 572 Dieterich, J.H., 1972. Time-dependent friction in rocks. Journal of Geophysical Research
- 573 B77, 3690-3697 https://doi.org/10.1029/JB077i020p03690

- 574 Dieterich, J.H., 1979. Modeling of rock friction: 1. Experimental results and constitutive
- equations. *Journal of Geophysical Research* B84, 2161-2168
- 576 https://doi.org/10.1029/JB084iB05p02161.
- 577 Gardien, V., Thompson, A. B., Ulmer, P., 2000. Melting of biotite + plagioclase + quartz
- gneisses: the role of H₂O in the stability of amphibole. *Journal of Petrology* 41, 651-666.
 https://doi.org/10.1093/petrology/41.5.651
- 580 Goodwin, L. B., Tikoff, B., 2002. Competency contrast, kinematics, and the development of
- foliations and lineations in the crust. *Journal of Structural Geology* 24, 1065-1085.

582 https://doi.org/10.1016/S0191-8141(01)00092-X

- Harea, E., Lapsker, I., Laikhtman, A., Rapoport, L., 2013. Bauschinger's effect and
- dislocation structure under friction of LiF single crystals. *Tribology Letters* 52, 205–212.
 https://doi.org/10.1007/s11249-013-0206-y
- 586 Hill, R., 1998. *The Mathematical Theory of Plasticity*. Oxford University Press, 372 pp.
- 587 Ioannidi, P. I., Le Pourhiet, L., Agard, P., Angiboust, S., Oncken, O., 2021. Effective
- 588 rheology of a two-phase subduction shear zone: Insights from numerical simple shear
- experiments and implications for subduction zone interfaces. *Earth and Planetary Science Letters* 256, 116913 https://doi.org/10.1016/j.epsl.2021.116913
- 591 Kisters, A. F. M., Ward, R. A., Anthonissen, C. J., Vietze, M.E., 1998. Melt segregation and
- far-field melt transfer in the mid-crust. *Journal of the Geological Society* 166, 905-918.
- 593 https://doi.org/10.1144/0016-76492009-012
- Konn, M.J., Spear, F.S., Valley, J.W., 1997. Dehydration melting and fluid recycling during
 metamorphism: Rangeley Formation, New Hampshire, USA. *Journal of Petrology* 38,
 1255–1277 https://doi.org/10.1093/petroj/38.9.1255
- 597 Kriegsman, L.M., 2001. Partial melting, partial melt extraction, and partial back reaction in
- anatectic migmatites. *Lithos* 56, 75-96. https://doi.org/10.1016/S1464-1895(01)00052-7
- Loomis, T.P., 1977. Kinetics of garnet granulite reaction. *Contributions to Mineralogy and Petrology* 62, 1–22. https://doi.og/10.1007/BF00371024
- Marone, C., 1998. Laboratory-derived friction laws and their application to seismic faulting.
 Annual Reviews in Earth and Planetary Science 26, 643-696.
- 603 https://doi.org/10.1146/annurev.earth.26.1.643
- 604 Nyman, M.W., Pattison, D.R., Ghent, E.D., 1995. Melt extraction during formation of K-
- 605 feldspar + sillimanite migmatites, west of Revelstoke, British Columbia. *Journal of*
- 606 *Petrology* 36, 351–372. https://doi.org/10.1093/petrology/36.2.351

- Nyström, A.I., Kriegsman, L.M., 2003. Prograde and retrograde reactions, garnet zoning
 patterns, and accessory phase behaviour in SW Finland migmatites, with implications for
- 609 geochronology. In: Vance, D., Müller, W., Villa, I.M. (Eds). *Geochronology: Linking the*
- 610 *Isotopic Record with Petrology and Textures*. Geological Society of London, Special
- 611 Publications 220, 213-230. https://doi.org/10.1144/GSL.SP.2003.220.01.13
- 612 Perfilyev, V., Moshkovich, A., Lapsker, I., Laikhtman, A., Rapoport, L., 2014. Dislocation
- 613 structure and stick–slip phenomenon. *Tribology Letters* 55, 295–301.
- 614 https://doi.org/10.1007/s11249-014-0358-4
- Persson, B.N.J., 1995. Theory of friction: Stress domains, relaxation, and creep. *Physical Review* B51, 13568. https://doi.org/10.1103/PhysRevB.51.13568
- 617 Pistone, M., Cordonnier, B., Ulmer, P., Caricchi, L., 2016. Rheological flow laws for
- multiphase magmas: An empirical approach. *Journal of Volcanology and Geothermal Research* 321, 158-170. https://doi.org/10.1016/j.jvolgeores.2016.04.029
- 620 Pistone, M., Baumgartner, L.P., Bégué, F., Jarvis, P.A., Bloch, E, Robyr, M., Müntener, O.,
- 621 Sisson, T.W., Blundy, J.D., 2020. Felsic melt and gas mobilization during magma
- 622 solidification: An experimental study at 1.1 kbar. *Frontiers in Earth Sciences* 8:175.
- 623 https://doi.org/10.3389/feart.2020.00175
- 624 Rabinowicz, M., Vigneresse, J.L., 2004. Melt segregation under compaction and shear
- 625 channeling: Application to granitic magma segregation in a continental crust. *Journal of*
- 626 *Geophysical Research* B109, B04407. https://doi.orgi/10.1029/2002JB002372
- 627 Reichhardt, C., Olson-Reichhardt, C.J., 2017. Depinning and nonequilibrium dynamic phases
- 628 of particle assemblies driven over random and ordered substrates: a review. *Report on*
- 629 Progress in Physics 80, 026501. https://doi.org/10.1088/1361-6633/80/2/026501
- 630 Rosakis, A.J., Samudrala, O., Coker, D., 1999. Cracks faster than the shear wave speed.
- 631 *Science* 284, 1337–1340. https://doi.org/10.1126/science.284.5418.133
- Rubinstein, S.M., Cohen, G., Fineberg, J., 2004. Detachment fronts and the onset of dynamic
- 633 friction. *Nature* 430, 1005–1009. https://doi.org/10.1038/nature08676
- Ruina, A. 1983. Slip instability and state variable friction laws. *Journal of Geophysical*
- 635 *Research* B88, 10359-10370. https://doi.org/10.1029/JB088iB12p10359
- Rushmer, T., 2001. Volume change during partial melting reactions: implications for melt
- extraction, melt geochemistry and crustal rheology. *Tectonophysics* 34, 342, 389-405.
 https://doi.org/10.1016/S0040-1951(01)00172-X

- 639 Santa-Aho, S., Laitinen, A., Sorsa, A., Vippola, M., 2019. Barkhausen noise probes and
- 640 modelling: A review. *Journal of Nondestructive Evaluation* 38, 94.
- 641 https://doi.org/10.1007/s10921-019-0636-z
- 642 Stüwe, K., 1997. Effective bulk composition changes due to cooling: a model predicting
- 643 complexities in retrograde reaction textures. *Contributions to Mineralogy and Petrology*

```
644 129, 43–52. https://doi.org/10.1007/s004100050322
```

- 645 Svetlizky, I., Fineberg, J., 2014. Classical shear cracks drive the onset of dry frictional
 646 motion. *Nature* 509, 205–208. https://doi.org/10.1038/nature13202
- Vachon, R., Hieronymus, C. F., 2017. Effect of host-rock rheology on dyke shape, thickness
 and magma overpressure. *Geophysical Journal International* 208, 1414–1429.
- 649 https://doi.org/10.1093/gji/ggw448
- 650 Vigneresse, J. L., Tikoff, B., 1999. Strain partitioning during partial melting and crystallizing
- 651 felsic magmas. *Tectonophysics* 312, 117-132. https://doi.org/10.1016/S0040-
- **652 1951(99)00167-5**
- Vigneresse, J.L, Barbey, P., Cuney, M., 1996. Rheological transitions during partial melting
 and crystallization with application to felsic magma segregation and transfer. *Journal of Petrology* 37, 1579-1600. https://doi.org/10.1093/petrology/37.6.1579
- Vigneresse, J.L., Burg, J.P., 2005. Simulation of crustal melt segregation through cellular
- automata: Insight on steady and non-steady state effects under deformation. *Pure and*
- 658 *Applied Geophysics* 162, 62, 987–1011. https://doi.org/10.1007/s00024-004-2652-9
- 659 Vigneresse, J.L., Cenki, B., Kriegsman, L. M., 2021. Melting by numbers: assessing the
- effective melt fertility of crustal rocks. *Lithos* 386-387, 106006.
- 661 https://doi.org/10.1016/j.lithos.2021.106006
- 662

664 **Figure captions**

Model of melt extraction using a cellular automata (c) with minerals 665 Figure 1. distributed in cells. A melting equation (e) implies muscovite, biotite, plagioclase and quartz, 666 667 resulting in melt, Kfeldspar and garnet (undifferentiated ferro-magnesian). All minerals are plotted according to a given colour. Melt production depends on the textural distribution, random 668 669 (r) or layered (l) and according to the chemical equation (e) of dehydration. Melting (m) occurs when adequate minerals are adequately and proportionally distributed. Two thresholds (t) control liquid 670 671 percolation (LPT), i.e. melt motion, and melt extraction (MET). Particles motion may move under pure shear (PS) or simple shear (SS) modifying the disposition of cells. Values with melt larger than 672 50, i.e. more melt that the matrix, are not considered. 673

674

675 Figure 2. Schematic representation of frictional law and its analogue for melt extraction. a) Stationary friction of a block submitted to a force F. A coefficient of friction (μ) 676 is defined as the ratio of normal (σ) to tangential (τ) stress components. b) Dynamic friction. 677 The applied force varies from a velocity V_0 to V. The coefficient of friction presents a spike 678 $(\Delta \mu)$ function of the relative change in velocity, with a return to stationary conditions after a 679 delay Dc. The new coefficient of friction may be higher or lower than the initial μ_0 . c) The 680 basal contact of the block is no more homogenous, with either two phases or asperities. An 681 682 irregular motion, with successive sticking and sliding states is observed, with amplitude (σ_{μ}). A stiction spike is also observed, as well as return to stick-slip when the velocity varies. d) 683 Melt extraction coefficient (μ) defined as the ratio of initial melt to permeability (z) taken as a 684 melt escape threshold. e) Similar to friction, the initial melt (ϕ) varies, as well as the 685 permeability (ζ) and added stress (τ). A spike is observed before return to a stationary state. f) 686

687 irregular amount of melt is extracted mimicking stick-slip behaviour when the melt flow688 presents reactions damping its amount and velocity.

689

690	Figure 3.	Results of the numerical simulation for different values in the initial amount
691	of melt. a) The a	mount of melt extracted by step (σ) presents a stick-slip behaviour that
692	attenuates with	the quantity of initial melt (ϕ_0). Stiction spike appears when this quantity is
693	suddenly change	ed. b) normalizing the amount of melt produced at each step by the initial
694	permeability (ζ)	and dimension of the system results in curves converging toward a
695	coefficient of m	elt extraction (μ).

696

Figure 4. Two regimes, regular or stationary and transient, are observed depending on whether the initial conditions vary during runs. Changing the melt proportion (ϕ_0) induces spikes ($\Delta\mu$) and stick-slip motion with amplitude σ_{μ}). Spikes ($\Delta\mu$) in melt extraction with iterations taken as a proxy for time (Dc), when the acceleration (ζ) is suddenly varied. Changes are made at different iteration intervals, and shown with three initial melt percentage (ϕ_0).

703

Figure 5. Variation of the amplitude of the stick-slip effects (σ_{μ}) as a function of the initial melt (ϕ_0) when the permeability (ζ) varies. A weak linear variation shows the poor effect of the initial melt content. Also indicated the range of values for melt (m) escape (in pale grey) as estimated in nature.

709	Figure 6. a) Variation of the amplitude of the stick-slip effects (σ_{μ}) in function of the				
710	melt extraction coefficient in two different situations, when the changes operate in separate				
711	runs, or in succession during a single run. The changes concern the initial melt (ϕ_0), the				
712	permeability (ζ) or the strain rate (ϵ°). The linear relationship indicates the dependence of σ_{μ}				
713	on the coefficient of melt extraction (μ). It corroborates all variations with the equivalent				
714	velocity parameters (ϕ_0 , ζ , ϵ°). The range of melt amount in nature (m) is indicated in light				
715	grey. Conversely, the transient effects illustrate a neat decrease of σ_{μ} during the successive				
716	variations of the equivalent velocity parameters, pointing to some memory effect. b)				
717	Amplitudes of σ_{μ} plotted for individual values of ζ , show an exponential decrease with the				
718	iteration numbers, taken as a proxy for time. It reinforces the memory effects, damping with				
719	time.				
720					
721	Figure 7. Amplitude of the spikes $(\Delta \mu)$ when the permeability value varies along a				
722	single run. The successive changes are positive or negative, yielding similar changes in $\Delta\mu$.				
723	The curves have been separated for increasing and decreasing permeability changes.				

Figure 8. Experiments on stick slip on a single crystal of LiF, redrawn from Perfilyev et al. (2014). a) The displacement (mm) as a function of time (s) shows the stick-slip behaviour, following increases and sudden drop in frictional force (mN). Each episode occurs in about 180 s with a displacement of about 7 mm, implying a strain rate of about 5.5 10^{-3} s⁻¹. b) The corresponding velocity of the indenter varies by a factor about 300 in less than 1 s (note the change in time scale between a and b). It is interpreted as depinning and pinning effect corresponding to dislocation creep.

733	Figure 9.	Cartoon illustrating the analogies between (a) frictional law, (b) melt
734	extraction and (c) p	poroelasticity. Friction implies solid-solid contact, i.e. crystalline networks
735	in a polycrystal ma	terial, with inherent dislocations (inverted tau) that act as pinning points
736	for motion. At the	other end, poroelasticity describes a liquid flow, percolating through the
737	system, but with di	screte air bubbles, non-percolating. The different surface energy values
738	between air, liquid	and solid make the system hysteretic. They also rule pinning and
739	depinning when the	e local porosity changes due to the pore size. In between, melt extraction
740	consists in a polym	erized silicate melt (tetrahedra) and undifferentiated ions (circles) flowing
741	in between crystall	ized matrix ongoing melting reactions. Minerals consist in quartz,
742	plagioclase, biotite	and muscovite (white to black, with pale and dark grey), with incongruent
743	mineral (coloured l	nexagons), among which garnet (red) and sillimanite (stripes). Dehydration
744	melting occurs whe	en the four quoted minerals are in contact (plain circle), feeding the melt
745	layer. They act as c	lepinning points. In contrast, retrograde reactions manifest by garnet
746	coronas (dashed cir	ccles), consuming elements of the melt. They act as pinning points.

