

Use of convolutional neural networks with encoder-decoder structure for predicting the inverse operator in hydraulic tomography

Abderrahim Jardani, T. M. Vu, Pierre Fischer

▶ To cite this version:

Abderrahim Jardani, T. M. Vu, Pierre Fischer. Use of convolutional neural networks with encoder-decoder structure for predicting the inverse operator in hydraulic tomography. Journal of Hydrology, 2022, 604, 10.1016/j.jhydrol.2021.127233. insu-03661813

HAL Id: insu-03661813 https://insu.hal.science/insu-03661813v1

Submitted on 8 Jan 2024

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution - NonCommercial 4.0 International License

1 Use of Convolutional Neural Networks with encoder-decoder structure for

2 predicting the inverse operator in hydraulic tomography.

3	A. Jardani $^{(1*)}$, T.M.Vu $^{(1)}$, and P.Fischer $^{(2)}$					
4						
5	(1) Normandy University, UNIROUEN, UNICEAN, UMR CNRS 6143 M2C					
6	5 Morphodynamique Continentale et Côtière, Université de Rouen.					
7	(2) HydroSciences Montpellier, Univ. Montpellier, CNRS, IRD, Montpellier, France					
8						
9						
10	• Corresponding author : abderrahim.jardani@univ-rouen.fr					
11						
12	Abstract: In this manuscript, we discuss the capabilities of a deep learning algorithm					

implemented with the Conventional Neural Network concept to characterize the hydraulic 13 properties of aquifers. The algorithm called CNN-HT is designed to predict the inverse operator 14 15 of hydraulic tomography using a synthetic training dataset in which the hydraulic head data associated with pumping tests are linked to hydraulic transmissivity field. This approach relies 16 on an adaptation of the SegNet network that was initially developed to process image 17 18 segmentation. The SegNet is composed of encoders and decoders networks. In the encoder, 19 sequential operations with multiple filters, as convolution, batch normalization, max-pooling are performed to identify feature maps of the input data. In the decoder, the up-sampling, 20 21 convolution, batch normalization and regression operations are used to prepare the output by recovering the loss of spatial resolution that occurred in the encoder process. In this adaptation, 22 we used the least-square iterative formulation at the initial iteration with Jacobian matrix to 23 resize the hydraulic head data to match the size of the output (transmissivity field). This protocol 24 was applied to the hydraulic head data computed numerically by solving the groundwater flow 25 26 equation for a given transmissivity field, generated geostatistically with Gaussian and spherical variograms. A part of this data was used for training the network and the other part to test its 27 performance. The test step confirmed the effectiveness of this tool in reconstructing the main 28 29 heterogeneities of the hydraulic properties, and its effectiveness is related to the nature and

quantity of the training data. Moreover, the CNN-HT method provided inversion results of the 30 31 same quality than those obtained with the Gauss-Newton algorithm using the finite difference or adjoint state method in the computation of the Jacobian matrix. However, the computational 32 time is longer in CNN-HT but this time can be less or of the same order as that of Gauss-Newton 33 using finite difference method. 34

Introduction 35 1.

36

Since the 1980s, the hydraulic tomography has been adopted in hydrology as an effective 37 technique for mapping the heterogeneity of the hydraulic properties of aquifers (Neuman, 1987; 38 Gottlieb and Dietrich, 1995; Bohling et al., 2002; Yeh and Liu, 2000). Indeed, this approach 39 can effectively provide insights into the spatial variability of the hydraulic transmissivity and 40 storage coefficient of a porous or fractured aquifer through a combined interpretation of 41 hydraulic data obtained from several pumping tests (Bohling et al., 2002; Berg and Illman, 42 2011; Cardiff et al., 2013; Fischer et al., 2018). The approach consists of using an inversion 43 algorithm, either deterministic or stochastic, to retrieve the best models that could match the 44 observed piezometric data (Fu & Gómez-Hernández, 2009a; Kitanidis, 1997). In general, the 45 46 inverse solution is not unique for that reason the inverse formulation usually incorporates a regularization term in the optimization of the objective function $\Psi(m)$ that is defined as the 47 sum of data misfit and regularization terms, as below (Tarantola and Valette, 1982): 48

-т

where the data misfit term assesses the pertinence of the model in term of matching the data 50 and is expressed as a sum of square differences between the observed hydraulic data $\mathbf{h}_{\mathrm{obs}}$ and 51 the numerical hydraulic data derived by solving a forward operator F(m). In hydraulic 52 tomography, the forward problem involves a numerical discretization for solving the 53

groundwater flow equation subject to Darcy's law in transient or steady modes. $m = -\log[T]$ 54 is the spatial distribution of the negative logarithm of the hydraulic transmissivity T, which is 55 regarded as an unknown field and is expressed logarithmically to ensure its positivity during 56 the inverse process. \mathbf{C}_{d} is a diagonal matrix for incorporating data uncertainty in the 57 optimization. On the other side, the regularization term is expressed with a prior model \boldsymbol{m}_0 and 58 its covariance matrix \mathbf{C}_{m} . This term can be obtained from geological, geophysical and tracer 59 investigations to constrain the inversion to offer only plausible solutions (Lochbühler et al., 60 2013; Zhao et al., 2016; Soueid et al., 2016). In general, the use of geostatistical constraints is 61 still widely applied in hydraulic tomography by using statistical properties such as mean and 62 covariance to obtain an imagery of hydraulic properties with smooth features (Kitanidis, 1997; 63 Yeh and Liu, 2000). 64

Once the parameters of the objective function are formulated, its optimization is processed 65 iteratively using one of the three categories of optimization algorithms: deterministic (Li et al., 66 2005), stochastic (Jimenez et al., 2016) or global (Castagna and Bellin, 2009). These algorithms 67 involve a repetitive numerical solving of the forward problem by applying finite elements or 68 69 finite difference methods to assess the ability of the proposed model to reach the convergence state. Therefore, the computation time is dependent on the numerical tools, the dimensionality 70 of the unknown parameters and the nature of the optimization algorithm employed. In general, 71 72 deterministic algorithms are the most used in high-dimensional inverse problems due to their advantage of reaching convergence in a short time compared to stochastic or global algorithms 73 74 by exploiting the geometric properties of the objective function (gradient and Hessian) (Tarantola and Valette, 1982). Nevertheless, their efficiency depends strongly on the choice of 75 the initial model as they only provide local minima. On the other hand, stochastic methods (such 76 77 as Markov chain Monte Carlo McMC) are based on the concept of sampling where the solution

is obtained by exploring the optimization performance of many randomly generated models and 78 79 selecting the best of them in terms of minimization of the objective function. Most of the time, algorithms belonging to this category are easy to implement and insensitive to the initial model. 80 However, their applicability to inverse problems with a large number of unknown parameters 81 remains limited due to the high repetitive and heavy computation of the forward problem during 82 the sampling process (Oliver et al., 1997; Fu and Gómez-Hernández, 2009b; Wang et al. 2017; 83 84 Jardani et al., 2012; Elsheikh et al., 2012). Other global optimization algorithms such as particle swarm optimization, genetic, simulated annealing can also be used to identify the best solution 85 by iteratively evaluating the objective function until the state of convergence. However, 86 87 convergence can be slow, especially in the case of high-dimensional inverse problems (Scales et al., 1990; Fernández-Martínez et al., 2011). 88

In this paper, we test a new generation of deep learning algorithm that can be considered as a 89 global optimizer in which prediction does not depend on an initial model and the computation 90 of the sensitivity matrix as it is the case for gradient-based methods (Zio, 1997). The concept 91 92 of this approach is different from previous optimization methods as the process focuses on approximating the inverse function by finding a universal relationship linking the input and 93 output data (in the case of hydraulic tomography between hydraulic pressure and hydraulic 94 95 transmissivity). Indeed, the deep learning is a powerful generalizer of non-linear and complex functions by identifying a set of parameters such as weights and biases allocated to the hidden 96 layer neurons that process the input data and link them to their corresponding output. The first 97 generation of deep learning algorithms uses a set of hidden layers with a large number of 98 neurons fully connected via high-dimensional weights, which complicates computation and 99 100 requires a lot of time, especially for handling the imaging tasks (Shen, 2018). However, the emergence of new deep learning architectures with Convolutional Neural Networks (CNN) 101 102 concept that have been proven in high-resolution image processing have opened the way for

broadening their applicability in the hydraulic characterization (LeCun et al., 1998; Sun, 2018; 103 104 Zhu and Zabaras, 2018). Indeed, the CNN concept has succeeded in reducing drastically the image processing time thanks to the convolution calculation, which allows to retrieve the local 105 106 features of the image through small filters (Indolia et al., 2018). These filters cover the whole input image but their activation is done zone by zone with local connections between pixels. 107 This local convolution reduces the number of parameters to optimize during the learning 108 process (LeCun et al., 1998). Among the first attempts at adopting the CNN concept in the 109 realm of tomography by inversion in geosciences, we cite: Sun (2018) trained the Generative 110 Adversarial networks for linking the hydraulic head map and the spatial distribution of 111 112 hydraulic conductivity. Zhu and Zabaras (2018) proposed a deep convolutional encoderdecoder network to reconstruct the image of hydraulic conductivity from hydraulic head map. 113 Laloy et al. (2018) used the Generative Adversarial networks to generate randomly in 2D and 114 115 3D binary hydraulic conductivity fields and MCMC to perform the inversion process. In geophysics realm, Wu and Lin et al. (2019) applied CNN with encoder-decoder networks for 116 mapping the subsurface velocity from seismogram data. Puzyrev and Swidinsky. (2019) used 117 118 the CNN network to determine the vertical subsurface heterogeneity of electrical conductivity by training electromagnetic data. 119

120 In this article, we discuss the use of the convolution neural networks formed according to the SegNet architecture to process hydraulic tomography. This technique was initially designed for 121 the semantic segmentation of objects learned on the images (Badrinarayanan, et al. 2017). To 122 illustrate the inversion principle with the CNN-HT structure, we organize this manuscript with 123 the following outline: the first section will be devoted to the generation of data used in training, 124 125 validation and test, and the second to the introduction of concept of the CNN-SegNet structure. In the application sections, we discuss the relevance of the approach on theoretical cases under 126 different conditions on training data with high and low resolution and contaminated or not by 127

the noise. We also devote a section to the comparison of the results obtained with CNN-SegNetcode with those determined with the Gauss-Newton algorithm.

130 2. Theoretical concept:

131 **2.1.Preparation of Training data:**

In deep learning algorithm, the first step concerns the construction of training data that will condition the accuracy of the result obtained from the predicted inversion operator (see figure 1). Given the impossibility of having a real dataset connecting the spatial distribution of hydraulic properties of aquifers and the piezometric responses, we use synthetic models in which the hydraulic parameter fields are generated geostatistically and their corresponding piezometric responses are obtained by solving the numerical groundwater equation in steady state (forward problem).

139
$$\mathbf{F}(\boldsymbol{m}):\begin{cases} \nabla .(T\nabla h) = q_p \delta(\boldsymbol{x}_p - \boldsymbol{x}), \\ h = h_0 , \Gamma_d \end{cases}, \qquad (2)$$

140 where *m* is the negative of logarithm of transmissivity $T=10^{-m}$ randomly generated and 141 distributed on the studied domain. *h* is the hydraulic head responses due to water extraction 142 represented by a point source term $q_p \delta(x_p - x)$ defined at pumping well locations x_p (δ is 143 Dirac delta function). The pumping tests will be done sequentially on several wells. The 144 forward problem is solved by finite element technique using the Comsol software with a 145 constant hydraulic level h_0 imposed as Dirichlet boundary at all limit of the domain Γ_d that are 146 positioned far from the studied area for reducing their impact.

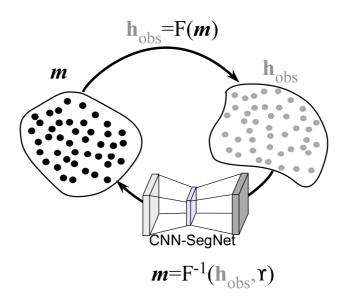
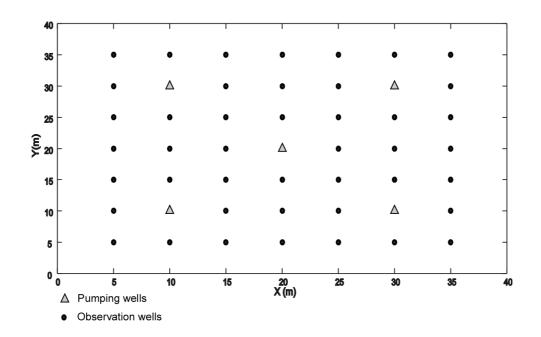




Figure 1: A sketch describing the concept of hydraulic tomography with the CNN-SegNet. This concept is based on the prediction of inverse operator $\hat{F}^{-1}(\mathbf{h}_{obs}, \Upsilon)$ by the CNN-SegNet algorithm using a training dataset in which the hydraulic transmissivity fields are generated geostatistically and the hydraulic head responses due to pumping tests are obtained numerically by solving the groundwater flow equation (Forward problem $F(\mathbf{m})$).

As the nature of the training models conditions the result of the CNN-HT inversion, we have to 153 use the prior information on the studied field to establish a cluster of models for training and 154 targeting a type of solution with certain characteristics. In this paper, prior information is 155 expressed in statistical terms, with a Gaussian distribution wherein the mean and covariance, 156 are held constants during model generation. Once the parameters of the distribution are defined, 157 we launch the generation of thousands of hydraulic transmissivity fields using the SGeMS code 158 implemented in Matlab (Remy et al., 2009). The generated transmissivity fields are then 159 assigned to a domain of [40 m × 40 m] on which we set up a dense distribution of wells to better 160 cover the heterogeneity of the field. On this configuration, we conduct 5 pumping tests in such 161 162 a way to hydraulically disturb the whole area. Each pumping test allows us to retrieve 44 measurements from the remaining observation wells; in the end, we obtain 220 measurements 163 164 for each field (see figure 2).



165

Figure 2: Position of the wells used in the generation of the hydraulic head data. 5 pumping
tests were carried out successively in the wells marked with a triangle and the responses of the
hydraulic head were collected in 44 wells with the dot symbol. This configuration provides 220
hydraulic data that are then used in the reconstruction of the hydraulic transmissivity.

To predict an inversion operator using convolutional neural networks with the encoder-decoder architecture, we need to reshape the piezometric data into a field with the same dimension than the hydraulic transmissivity field in order to link tow images with the same size. To do this, we can use one of two methods:

(i) The first consists in interpolating the piezometric data recorded on the observation wells
for each pumping test in order to form images with the same resolution in the target
transmissivity field. In this case, the input data are composed of a number of maps equal to the
number of pumping tests. In our example, we have 5 pumping tests and we will then obtain 5
maps that will be treated jointly as 5 channels in convolutional neural networks.

(ii) This second is more compact than the first in which all piezometric data set recorded
for different pumping tests are introduced by a single map using a projection operation derived
from the Gauss-Newton formulation. This formulation results from the minimization of the

183 objective function presented in the introduction section $\left(\frac{\partial \Psi(m)}{\partial m}=0\right)$ (Tarantola & Valette,

185
$$\boldsymbol{m}_{i+1} = \boldsymbol{m}_{i} + \left[\mathbf{J}_{i}^{T} \mathbf{C}_{d}^{-1} \mathbf{J}_{i} + \mathbf{C}_{m}^{-1} \right]^{-1} \left[\mathbf{J}_{i}^{T} \mathbf{C}_{d}^{-1} \left(\mathbf{h}_{obs} - \mathbf{F}(\boldsymbol{m}_{i}) \right) + \mathbf{C}_{m}^{-1} \left(\boldsymbol{m}_{i} - \boldsymbol{m}_{0} \right) \right], \quad (3)$$

186 where \mathbf{J}_{i} is Jacobian matrix with a size of $N \times M$, (*N* and *M* are number of the data and the 187 unknown parameters respectively). At the first iteration, usually we start with the prior model 188 then:

189
$$\boldsymbol{m}_{1} = \boldsymbol{m}_{0} + \left[\mathbf{J}_{0}^{\mathrm{T}} \mathbf{C}_{d}^{-1} \mathbf{J}_{0} + \mathbf{C}_{m}^{-1} \right]^{-1} \mathbf{J}_{0}^{\mathrm{T}} \mathbf{C}_{d}^{-1} \left(\mathbf{h}_{obs} - \mathbf{F}(\boldsymbol{m}_{0}) \right), \qquad (4)$$

and this equation can be reformulated under this following form:

191
$$\boldsymbol{m}_{1} = \boldsymbol{\beta} + \boldsymbol{\Lambda} \mathbf{h}_{\text{obs}}, \qquad (5)$$

192 where
$$\boldsymbol{\beta} = \boldsymbol{m}_0 - \left[\mathbf{J}_0^{\mathrm{T}} \mathbf{C}_d^{-1} \mathbf{J}_0 + \mathbf{C}_m^{-1} \right]^{-1} \mathbf{J}_0^{\mathrm{T}} \mathbf{C}_d^{-1} \mathbf{F}(\boldsymbol{m}_0) , \qquad (6)$$

193 and
$$\mathbf{\Lambda} = \left[\mathbf{J}_0^{\mathrm{T}} \mathbf{C}_d^{-1} \mathbf{J}_0 + \mathbf{C}_m^{-1} \right]^{-1} \mathbf{J}_0^{\mathrm{T}} \mathbf{C}_d^{-1},$$
 (7)

194 m_1 is the hydraulic transmissivity model resulting from the first iteration; obviously, it does not 195 fit the observed data, but it can be applied as a projection term for all piezometric data in order 196 to convert them into matrices with same size as the transmissivity fields. J_0 is the sensitivity 197 matrix of the prior homogenous model m_0 representing the mean of the fields. In this case, the 198 Jacobian can be derived analytically using the Thiem's equation for confined aquifers to express 199 the radial flow field and its adjoint state operator (for details, see Appendix). To further simplify this projection term so that it is less dependent on prior information, we remove the bias term β , replace the covariance matrices \mathbf{C}_{m} and \mathbf{C}_{d} by simple identity matrices, now the new projection term takes this form:

203
$$\mathbf{h}_{obs}^{p} = \mathbf{\Lambda} \mathbf{h}_{obs} \text{ With } \mathbf{\Lambda} = \left[\mathbf{J}_{0}^{T} \mathbf{J}_{0} + \mathbf{I}_{m} \right]^{-1} \mathbf{J}_{0}^{T}.$$
(8)

Thanks to this simple multiplication, the hydraulic data \mathbf{h}_{obs} will be transformed into a vector \mathbf{h}_{obs}^{p} of the same size as the output hydraulic transmissivity field and used in matrix form as the input field in the network. This formulation allows to link linearly \mathbf{h}_{obs} and \mathbf{h}_{obs}^{p} and preserves all the hydraulic information to establish a relationship between \mathbf{h}_{obs}^{p} and transmissivity field.

The comparison of the results of the two methods (linear interpolation and projection) in terms of inverse operator prediction is presented in the last section. However, the first sections all applications are done with the projection method.

211 **2.2.** Inversion by Convolutional Neural Networks with encoder-decoder structure

In this article, we will explore for the first time the efficiency of SegNet architecture to process the inversion in hydraulic tomography. The approach uses the Conventional Neural Networks concept built with an encoder-decoder architecture (see figure 3). This approach was originally conceived to process semantic segmentation of images by delineating the shapes of learned objects (Badrinarayanan, et al., 2017).

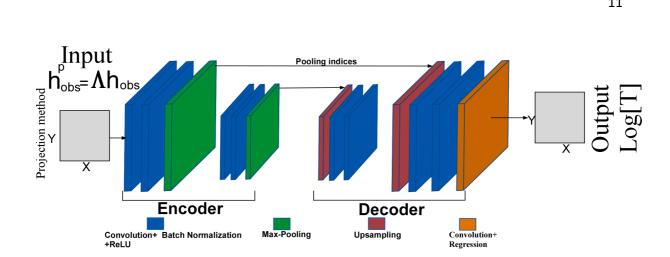


Figure 3: SegNet Convolutional Neural Networks with encoder-decoder structure. *Modified* 218 from (Badrinarayanan, et al., 2017). Here, the piezometric data can be resized to the same 219 dimension of the output field using the projection method 220

217

222 In the encoder network, the input image is processed sequentially by a set of encoders. In our case, we use only 2 encoders chosen after an analysis of the inversion results obtained with 223 networks having 1, 2 and 3 encoders. The network with 2 encoders gave the best result in terms 224 225 of training and generalization with a low sensitivity to overfitting. Each encoder starts with a convolutional process in which multiple filters with the small size $(3\times3\times64)$ are convoluted on 226 227 the input layer zone by zone to identify their features (Badrinarayanan, et al., 2017). This calculus permits to establish a local connectivity between pixels with these shared small filters, 228 which remain simple to predict in the training process. The convolutional operations are 229 230 followed by Batch normalization which is a linear operation that plays the role of a regularizer in scaling the convolution outputs in order facilitate and accelerate the training (Ioffe and 231 Szegedy, 2015). This step is followed by the application of the ReLU function, which 232 introduces the nonlinearity in the process. Then, the dimensionality of the result of the previous 233 sequence will be considerably reduced in the max-pooling layer with stride of 2 and 2×2 234 windows to keep only the main features. Therefore, the construction of feature maps in the 235 encoder network leads to a decrease in the spatial resolution that will be restored in the decoder 236 network. Indeed, the decoder network is configured in a symmetrical form with 2 decoders to 237

prepare the output by recovering the resolution lost in the decoder with up-sampling operations 238 239 from the max-pooling indices obtained in the previous max-pooling layers carried out in each encoder (Badrinarayanan, et al., 2017). These results of up-sampling will be convoluted with 240 241 other filters to build feature maps with more details. Then the Batch normalization and ReLU are applied successively on the convolutional outputs as in the encoder network. At the end of 242 243 the encoder network, we add a convolution layer having the same size of the output and a 244 regression layer to assess the performance of the training. This last layer replaces the Softmax layer used in the original version of SegNet network for the segmentation task. The set of 245 weights and biases Υ used in the various encoder-decoder operations are chosen in way to 246 match the training data $\left\{\mathbf{m}_{i},\mathbf{h}_{obs,i}^{p}\right\}$ by minimizing this objective function: 247

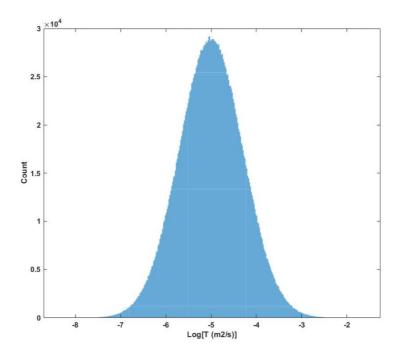
248
$$\Upsilon = \arg\min\left\{\sum_{i=1}^{N_t} \left\|\boldsymbol{m}_i - \hat{F}^{-1}(\mathbf{h}_{obs,i}^p, \boldsymbol{\Upsilon})\right\|\right\},\tag{9}$$

where i is the sample index of the training data set with Nt as size of the training data, 249 $\hat{F}^{-1}(\mathbf{h}_{obs,i}^{p}, \boldsymbol{\Upsilon})$ denotes the inverse function to identify by using the training data in the prediction 250 of the networks parameters Υ with ADAM optimization algorithm (Kingma and Ba, 2014). 251 This optimizer was run with a batch size of 70 and a learning rate starting at 0.01 and decreasing 252 253 by 0.1 every 50 epochs. The computation was performed on a workstation with (Intel(R))Xeon(R) Silver 4110 CPU @ 2.10GHz) and 128 G of RAM with a single GPU. After training 254 is complete, we use a sample of generated data not used during training phase to check the 255 quality of the inversion results by the CNN-HT code. 256

257 3. Applications to synthetic cases

In this section, we apply the inversion concept with CNN encoder-decoder networks on theoretical cases where the log of hydraulic transmissivity fields are generated randomly with

a Gaussian variogram and their hydraulic head responses by the groundwater equation as 260 explained in the data preparation section. This distribution has $\mu = 10^{-5}$ (m²/s) as mean and an 261 isotropic Gaussian variogram with 0.5 and 10 m as variance and range respectively. With this 262 variogram model, we performed 11000 realizations in which the transmissivity is ranging over 263 $[10^{-8.3}, 10^{-1.6}]$ (m²/s); however 95% of these models have transmissivities between $10^{-6.4}$ and $10^{-1.6}$ 264 $^{3.5}$ (m²/s) (see figure 4). Then, these models are used in a forward operator that takes 2s per 265 simulation to get the hydraulic head responses. These realizations aim to evaluate the 266 effectiveness of the CNN-SegNet approach in the reconstruction of hydraulic transmissivity 267 field and to analyze the sensitivity of this reconstruction to the size of the training dataset and 268 the amount and uncertainties of the hydraulic head data collected during the pumping tests. 269



270

Figure 4: Gaussian distribution of the log of transmissivity fields used in the training ofCNN-SegNet network.

273

3.1. Effect of training data size

In the first analysis, we examine the impact of training data size on the reliability of the 275 inverse operator prediction and determine the amount of data required to obtain a satisfactory 276 prediction. For this purpose, we performed three assessments with different training data sizes: 277 T1=1000, T2=3000, T3=10000, the validation was done with 1000 models and the test with 278 500 models different from those used in the learning phase (training and validation). During the 279 optimization process, we followed the quality of the minimization of loss functions with the 280 training and validation data until they reached optimal values to stop the process manually or 281 wait until the maximum number of epochs was met (see figure 5). In this way, we avoid the 282 problem of overfitting which appears especially when we tried to build the network with a 283 limited number of training data (T1=1000). 284

The results obtained with these different sizes of training data are presented in the Table 1a, 285 in which we can easily identify through the analysis of the correlation coefficient between the 286 real and predicted models of test sample that the accuracy of the prediction improved with 287 increasing training data size. Indeed, the reliability of the predictive model derived from a deep 288 289 learning algorithm is highly dependent on the amount of data used in the training, in which a large amount of data guarantees a good prediction. This implies a long computational time in 290 the generation of training data, in any case the degree of accuracy desired remains linked to the 291 292 nature of the problem studied.

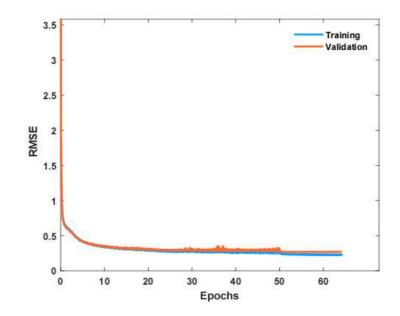
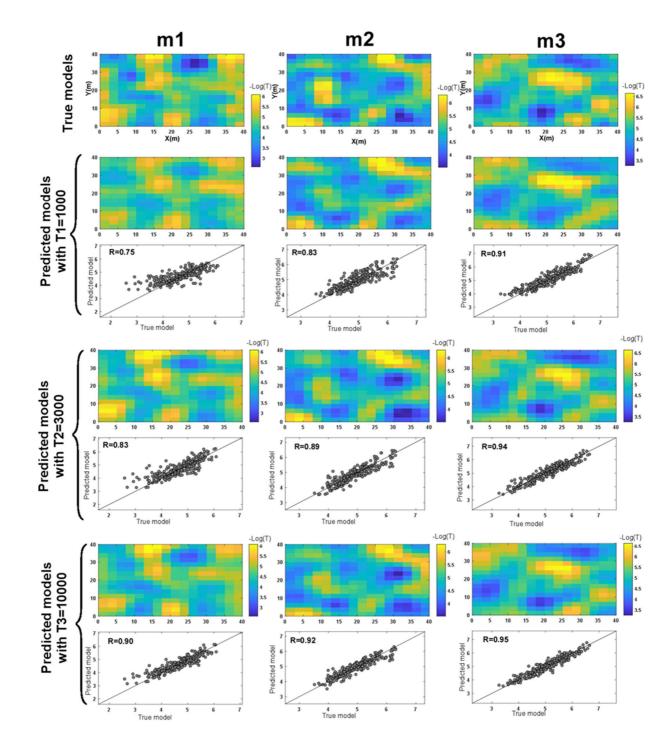


Figure 5: Both curves represent the quality of predictions of the network on validation and
 training data (10000 models for training, 1000 models for validation) during the optimization
 process.(RMSE: Root Mean Square Error).

In our case, training with $T_1=1000$ data allowed us to predict 500 test models with a 297 correlation coefficient R between real and predicted models ranging from 0.74 to 0.93 and a 298 299 mean of 0.86 (see Table 1a). From these models, we choose 3 models: the first one (m1) with a relative poor prediction (R=0.75), a second model (m2) (R=0.83) close to the mean 300 representing quality of the majority of predicted models, and the last one (m3) with a good 301 reconstruction (R=0.91) (see figure 6). These models will be used to analyze the impact of 302 training data size on the quality of their estimations (see figure 6). The first model **m1** is 303 characterized by a high degree of variability with the small heterogeneities at the boundary of 304 the domain and a logarithm of the hydraulic transmissivity that varies between [-6, -2.5]. Over 305 this range, there are outliers in the Gaussian distribution of training data in particular which are 306 around -2.5. Thus, in this case only a few models have been seen in the training with this feature, 307 which explains the poor reconstruction (see figure 6). The same problem occurs in the 308 prediction of models with low outliers < -6.5. However, the models having variabilities well 309 310 covered by training models such as **m2** and **m3** with transmissivities values between [-6.5, -

311 3.5], the reconstructions are more accurate. However, if the number of training models 312 increases, the models with outliers will be more frequent and will be reconsidered in the 313 generalization as shown in the table, where the reconstructions of these models have been well 314 improved.



315

Figure 6: Three models selected from test models to illustrate the prediction quality of the CNN-HT tool by comparing their true and predicted transmissivities. In this figure, the

- 318 predictions of these models were made with three different amounts of training data, in which
- a significant improvement in the reconstructions was obtained when the amount of training data
- 320 becomes important.

3.2 Effect of amount of observation wells

In the second analysis, we study the influence of the number of hydraulic data collected 323 during the pumping tests on the efficiency of the network to approximate the inverse operator. 324 325 In this analysis, we first reduce the number of observation wells to half (22 observation wells) 326 and in the second step to one quarter (11 observations wells). However, we maintain the number of pumping tests and the size of the spatial discretization of the transmissivity fields (16×16), 327 328 as used in previous test. These new well configurations are illustrated in the figure 7. In practice, only the hydraulic responses of abandoned wells are removed from previous training data. Thus, 329 these data acquired with the two new observation well configurations are used as training input 330 331 to determine the corresponding inverse operators, and the results obtained were compared to the previous configuration with 44 wells. The training operations were performed with 10,000 332 models, and the tests were also done on the previous unseen samples. The reconstruction results 333 of the test models with different series of piezometric data expressed in terms of correlation 334 coefficients reveal a degradation of the accuracy of prediction with the decrease in the number 335 336 of hydraulic data (see Table 1b). Because the number of piezometers becomes insufficient to cover all the heterogeneities and the complexities of the medium, and this also occurs with 337 classical inversion methods. This is further confirmed by analyzing the spatial aspect of the 338 339 reconstructions of models (m1, m2 and m3), in which we observe a loss of resolution and an increase in smoothness degrees with the reduction of hydraulic data (figure 8). However, the 340 main heterogeneities can still be identified with less data. 341

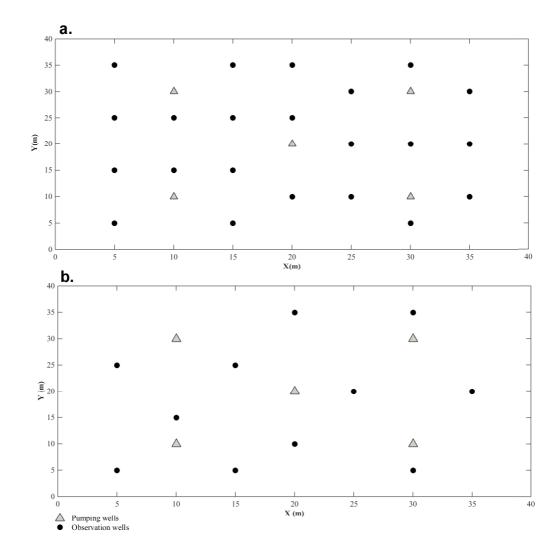


Figure 7: Two new observation well configurations were retained to study the effect of the
reduction in hydraulic data on the quality of the prediction. In the first configuration, the number
of observation wells was reduced by half (22 observation wells) and in the second, by a quarter
(11 observation wells) from the initial configuration shown in Figure 2.

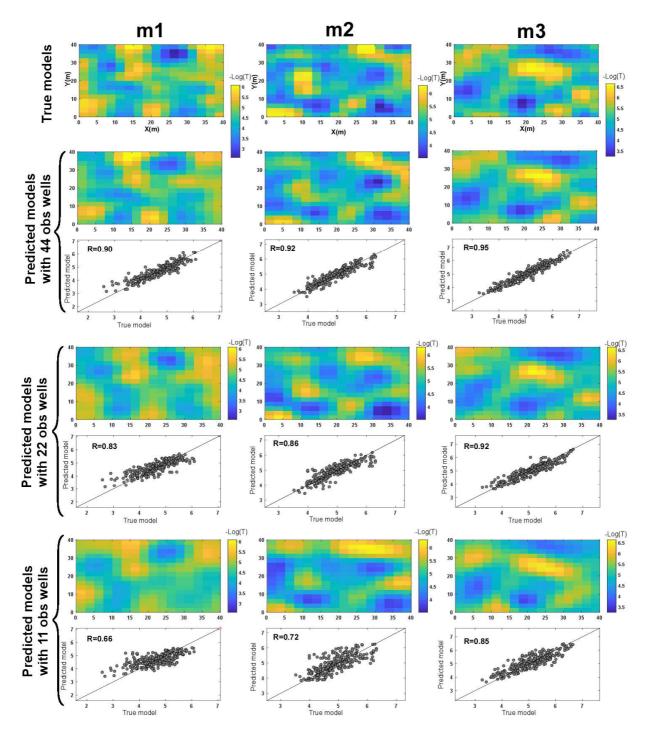


Figure 8: Predictions of the three models are provided with a set of training data (T3=10000) in which piezometric responses are collected with 3 different well configurations (44, 22 and 11 wells). These predictions showed a degradation in the reconstructions of hydraulic transmissivity with a decrease in the number of wells used in the survey.

352

353

354

Table 1 Summarizing metric evaluations of the predictions quality (in term of correlation coefficient) of 500 test models with correlation coefficient. **a.** Evaluation of the impact of the size of training data (T) on predictions. **b.** Evaluation of the impact of the number of observation wells on predictions. **c.** Evaluation of the effect of uncertainties in the piezometric data on predictions with incorporation or not of these uncertainties in the learning process.

360

a Size of training data	Comp	lation coeff	isiont
a. Size of training data	Correlation coefficient R		
	min	mean	max
T1=1000	0.74	0.86	0.93
T2=3000	0.77	0.90	0.95
T3=10000	0.84	0.92	0.96
b. Number of observation data	Corre	lation coeff	ïcient
		R	
	min	mean	max
T3=10000, N° data=220	0.84	0.92	0.96
T3=10000 N° data=110	0.73	0.85	0.93
T3=10000, N° data=55	0.57	0.77	0.91
c. Degree of noise on the data	Correlation coefficient		ïcient
	R		
	min	mean	max
T3=10000, N° data=220,	0.84	0.92	0.96
std =0 free noise			
T3=10000, N° data=220,			
std=0.5 on Training and Test	0.55	0.84	0.93
data			
T3=10000, N° data=220,	0.49	0.79	0.92
std=0.5 only on Test data			

361

362 3.3. Effect of observation uncertainties

In this section, we analyze the effect of uncertainties that may be associated with hydraulic pressure measurements collected during pumping tests. This involves contaminating the numerical piezometric data with a random Gaussian noise of 0.5 as the standard deviation, and verifying its impact on the quality of hydraulic field reconstructions of test models.

As a first analysis, we use an inverse operator established on the basis of uncontaminated 367 training data to interpret hydraulic head data from a test set that are contaminated by noise. 368 Here, the training operation did not take into account the noise. The result of this test is 369 presented by the benchmark models in figure 9 and table 1c of correlation coefficients, which 370 reveal a sharp degradation in the quality of the predictions compared to the noise-free test 371 models, with a decrease in the correlation coefficient (from 0.92 to 0.79 mean for all test 372 models). In the second strategy, noise is taken into account in the estimation of the inversion 373 374 function by adding noise to the hydraulic data used in the learning process. This method has led to a significant improvement in the prediction quality of the test models, as shown by the three 375 models (m1,m2 & m3) and the correlation coefficient noted in table 1c with the average R of 376 the set of models increasing from 0.79 to 0.84. However, this prediction is less accurate when 377 compared to the noise-free training and test data, as shown by the R-values. We can, therefore, 378 379 conclude that uncertainties in the measured data may have an impact on the results of inversion, as is the case with classical optimization techniques, but to reduce their impacts, it is necessary 380 to incorporate these uncertainties in the construction of the inversion model, i.e. in the learning 381 382 process.

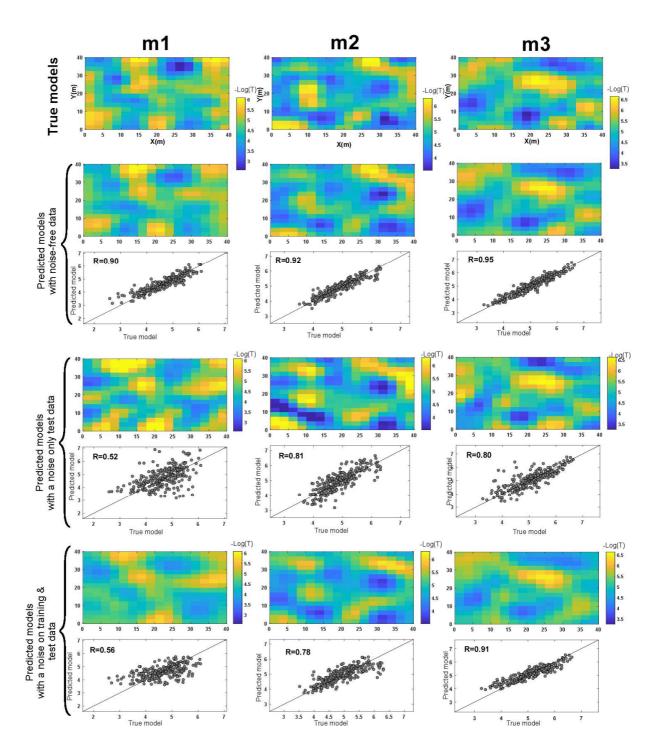


Figure 9: The purpose of these reconstructions is to analyze the effect of noisy data on the efficiency of the inversion with CNN-HT by comparing 3 types of predictions: 1) with the training and test data without noise; 2) with the training data without noise but the test models are noisy; 3) both the training and the test models are noisy. These comparisons show that the uncertainties in the measured data are affecting the quality of the reconstructions; however, their impacts could be attenuated by taking into account this noise in the training process.

390

391

4. Comparison of CNN-HT and Gauss-Newton algorithms

In this section, we compare the inversion results obtained with the CNN-HT algorithm 393 with those determined by the traditional Gauss-Newton (GN) method. To this end, we apply 394 both algorithms to the reconstruction of a transmissivity field with a mean of 10^{-5} m² s⁻¹ and 395 generated with a spherical covariance having 0.5 and 10 m as variance and range respectively. 396 However, this time, the models are highly discretized (32×32) compared to those used in the 397 398 previous sections. The Gauss-Newton algorithm, as described in the introductory section, is a deterministic approach that relies on the minimization of an objective function (Eq. 1) by 399 iteratively computing the Jacobian matrix until obtaining a local minimum. The calculation of 400 401 the Jacobian matrix remains the most demanding part of this process, especially when the number of unknowns is important. This calculation is usually done either by the finite difference 402 403 or the adjoint state technique. The finite difference method is simple to implement, but its computation is time consuming because it involves to solve the forward problem for each 404 405 unknown parameters (Cardiff & Kitanidis, 2008; Fischer et al., 2017). In our case study, where 406 the model is discretized into 32×32 cells, the construction of the Jacobian will therefore involve solving the forward problem 1024 times. On the other hand, the adjoint state technique requires 407 less computation, by iteratively solving an adjoint operator with the form of the forward 408 409 problem on the number of observation wells, but its implementation is quite complex (for more details see Sykes et al., 1985; Cardiff & Kitanidis, 2008). In this section, we test the 410 implementation of Gauss-Newton algorithm to invert the synthetic hydraulic data determined 411 with a hydraulic transmissivity field shown in figure 10, (called True model) and using the 412 piezometric coverage shown in figure 2. These data were then contaminated with noise with a 413 variance of 5cm². The inversion of these datasets was constrained, as explained in the 414 introduction, by geostatistical parameters such as covariance and mean. The results of the 415 inversions with the GN algorithm using the finite difference and adjoint state methods are 416

417 presented in Figure 10 where we can conclude that the GN algorithm succeeded in both cases 418 in identifying the main heterogeneities of the real field with certain smoothness. In the case 419 where the Jacobian is computed with the finite difference method, convergence was achieved 420 after 7 iterations in 4h10min and solving the forward problem $7\times(1024+1) = 7175$ times. In 421 contrast, the adjoint state method took only 1min24s.

Regarding the application of the CNN-HT algorithm, we geostatistically generate 11000 422 423 transmissivity models with the same variogram as the real model used for the comparison. We keep 10000 models for the training, and the rest of the models are equally split for validation 424 and testing. In this case study, we test two different ways of resizing the piezometric data, the 425 426 first method called projection method uses the Hessian matrix to reshape the input data into the same size as the transmissivity field (output) as described in the section 2. This method requires 427 the calculation of a Jacobian matrix with a homogeneous field for example the mean field value. 428 429 The second method is based on a linear interpolation of the hydraulic data recorded during each pumping test to form 5 maps of size 32×32 (i.e., one map for each pumping test), so the input 430 431 data has the form of a matrix of 5 channels $(32 \times 32 \times 5)$.

Both approaches were used to build the SegNet networks and then tested on the test 432 sample. The analysis of the predictions obtained on this sample shows that both methods 433 provide almost the same prediction qualities with a small advantage for the projection method 434 according to the correlation coefficients (see table 2). However, when we reduced the training 435 size to T=500 and T=2500 models, the network using the projection method was able to provide 436 accurate generalization while the network with interpolated data failed. We believe this is due 437 to the fact that the interpolation of piezometric data does not preserve all the information 438 masking the degree of variability in hydraulic responses, which makes learning difficult with 439 few training data. In addition, using the interpolation method increases the size of the input 440

442 to the projection technique that brings the data into a compact form.

441

Regarding the application of these two networks on the hydraulic data of the real model, 443 we find that the network built with compact data provides a better reconstruction of the spatial 444 variability of transmissivity field with a correlation coefficient R=0.80 (see figure 10). 445 Moreover, this prediction has almost the same accuracy as the result determined by the Gauss-446 Newton algorithm (R=0.83). We also mention that the time used in the GN with the finite 447 difference method is perhaps comparable to the time spent in the DL method mainly devoted 448 to the construction of the data set. In particular, the DL method becomes faster when we reduce 449 450 the amount of training data to T=500 and T=2500 (20 min and 1h41 min), because even with these small samples, the transmissivity field can be well reconstructed well in a reasonable 451 amount of time, as demonstrated by the comparison of real and inverted fields (figure 10). 452

453 At the end of this comparison, we conclude that both the CNN-HT and Gauss Newton codes are based on the forward problem for which the computation time depends on the 454 numerical method and the degree of mesh refinement used to solve the groundwater flow 455 equation. The machine learning technique uses this forward problem to generate a training 456 database to build the CNN network. The time spent in this construction can be equivalent or 457 458 less to the time spent in finding a local minimum with the Gauss Newton using the finite difference method to calculate the Jacobian matrix over a large number of unknown parameters. 459 However, it is difficult to have an idea in advance of the time that will be spent in building this 460 database because the size needed to get accurate results is related to the degree of complexity 461 of the input-output relationship and the type of network. The optimal size as well as the rest of 462 the network parameters (epochs, number of encoders) are only established by a trial-and-error 463 analysis, which can take a considerable amount of time. Therefore, the data size used in this 464 paper cannot be generalized to all cases treated by this network or others. 465

Both methods also share their dependence on prior information. In the Gauss-Newton 466 algorithm, this prior information is introduced as a constraint to guide the optimization towards 467 realistic models and reduce the uncertainties associated with the non-uniqueness of the solution. 468 On the other hand, the CNN-HT algorithm uses it to generate the training models. The CNN-469 HT offers the possibility of incorporating multiple prior models into the construction of the 470 training dataset, which is difficult to perform in a deterministic algorithm where the 471 mathematical forms of these constraints must be differentiable. In addition, both methods are 472 sensitive to the uncertainties associated with the measured data. The CNN-HT allows 473 simultaneous learning with data affected by several degrees of uncertainty by adding noise to 474 475 the training data, then combining this set with the original to double the size of the training data and take into account the impact of noise. Regarding the quantification of the prediction 476 uncertainties, in the Gauss Newton method, we use the a posteriori covariance formulation to 477 478 get an idea of the uncertainty interval concerning the local minimum (Tarantola and Valette, 1982). However, in the CNN-HT algorithm it is difficult to construct an uncertainty map of the 479 predicted model. We can only build an uncertainty map on the model used in the test. 480

481 Table 2: We report the metric evaluations (correlation coefficient R) and computational 482 time of inversions obtained with the CNN-HT and Gauss-Newton algorithms. The CNN-HT 483 network was trained with hydraulic data introduced either by the linear interpolation method or 484 by the projection method. We tested 3 different sizes of training data (500, 2500, 10000), but 485 the network with the interpolation method failed in the training process when the data size 486 became small (500 and 2500). We also present the evaluation of the Gauss-Newton metric using 487 the finite difference method and adjoint state on a test model shown in figure 10.

		Computation Time	R	R
			True Model (figure 10)	Test Sample min mean max
TH-NV	T1=500 Projection method	20min	0.73	0.54 0.69 0.79
	T2=2500 Projection method	1h41	0.79	0.65 0.77 0.86
	T3=10000 Projection method	6h	0.82	0.69 0.80 0.88
5	T3=10000 Interpolation method	6h38min	0.76	0.68 0.80 0.89
	Finite Difference	4h 10 min	0.83	
	Adjoint state method	1 min 24s	0.81	

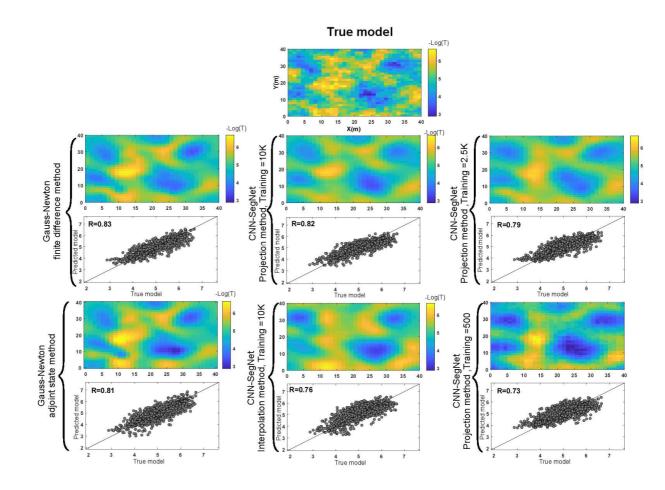


Figure 10: The results obtained when applying CNN-HT and Gauss-Newton algorithm in the reconstruction of the "True model". The results obtained with GN using the finite difference and adjoint state methods recovered the main heterogeneities of the field. The same reconstruction quality was obtained with CNN-HT, especially using the projection method for the preparation of the input data. Even with a small amount of data (T=500 and T=25000), this CNN-HT network provided satisfactory reconstructions.

496

497 **5. Conclusion and Summary**

In this work, we explore the relevance of a deep learning tool called the SegNet network in the mapping of spatial variability of hydraulic parameters by interpreting the piezometric data recorded during pumping tests. This approach relies on the approximation of the non-linear inverse operator connecting the hydraulic head data to the hydraulic transmissivity field. This approach is the outcome of an adaptation of the SegNet network that has been successfully applied in the realm of image segmentation by identifying the objects studied on the pixels of images. The SegNet is designed to establish a relationship between two images sharing the

same size and certain similarities, which is not the case for hydraulic head data, which are less 505 506 numerous than the hydraulic transmissivity values to be estimated. To meet this point, we test the projection and interpolation methods. Projection method is based on the classical iterative 507 508 formulation of a least-square method at the initial iteration to build a projection operator with the Hessian and Jacobian matrices derived from the mean value of transmissivity fields used in 509 510 the training data to resize the hydraulic data into the hydraulic transmissivity field size. 511 Interpolation method involves reshaping the data size by interpolating the hydraulic head data 512 to form a map with the same size as the output (transmissivity field) for each pumping test. As a result, the interpolation significantly increases the size of the input data, which can slow down 513 514 the numerical calculation without improving the quality of the prediction compared to the projection method, which better preserves the hydraulic information in compact form. 515

In the projection method, the operator is applied to the totality of the hydraulic pressure 516 data used in training which are determined numerically as responses to pumping tests carried 517 out on geostatistically generated hydraulic transmissivity fields with certain statistical 518 519 properties. Once the training data has been gathered, the inverse operator is established by identifying the different filters used in the SegNet, which is composed of encoders and decoders 520 521 networks. The encoder network allows, through multiple application of filters and sequential 522 calculations of convolution, Batch normalization, ReLU and max-pooling operations to extract the main features of the input data. This encoder process induces a drop in spatial resolution, 523 which will be recovered in the decoder network, which is designed to prepare the output with 524 the applications of up-sampling, convolution, Batch normalization and ReLU operations. This 525 network ends with a regression layer to evaluate the performance of the inverse operator 526 527 approximation.

528 The quality of this approximation is highly dependent on the amount of data used in the529 training with an accurate approximation when the dataset is large. This makes the assembly of

the training data the most expensive step in time, requiring several hours compared to the learning process that in this case does not exceed 30 minutes. In practice, it is preferable to launch the learning with few data and check whether or not the desired accuracy is achieved before generating large amounts of data unnecessarily. In terms of comparison, the time spent in forming a training database can be equivalent to the time spent to reach a local minimum with Gauss Newton using the finite difference method to compute the Jacobian matrix

536 In addition to the amount of training data, the nature of this data also controls the prediction results of the network. Indeed, the trained network is only intended to handle certain 537 types of models with similar characteristics to the models used in training. Thus, it is by no 538 539 means a universal inversion operator, which makes the choice of the nature of the training models a crucial step that must be based on realistic prior information. In real application case, 540 it should also be verified that the transmissivity fields used in the training allow to generate 541 hydraulic data with the magnitude of the piezometric data observed in the field. It is also 542 necessary to ensure that the degree of heterogeneity of the generated models can be easily 543 544 mapped to the piezometric data during training process. Because if the piezometric coverage does not allow to cover transmissivity models with high heterogeneity, the learning process will 545 fail to establish this relation. The result of the inversion with CNN-HT depends on the prior 546 547 information and this particularity is also found with classical inversion techniques such as Gauss-Newton. 548

Noise can alter the information carried by the piezometric data and is a frequent problem in inversion algorithms. The deep learning algorithms are no exception, but its impact can be reduced by integrating it into the learning step by contaminating the input of training data. The number of wells and their spatial configurations used in the study also had a determining effect on the effectiveness of the CNN-SegNet inversion tool, with more reliable reconstructions when these piezometers are well distributed over the study area, which was also observed in the useof conventional deterministic and stochastic methods.

556 At the end of this conclusion, we enumerate and discuss the main advantages and 557 limitations of the proposed method. Regarding the advantages, we point out that:

558 ✓ The CNN-HT method uses the forward problem as a black box to generate the training 559 data with user-selected transmissivity fields, thus avoiding numerical instability and outliers in 560 hydraulic head simulations related to its solving. These numerical problems can be encountered 561 when using conventional inversion methods (stochastic and deterministic) if they are ill-562 constrained.

563 ✓ Once the inverse operator is trained, the network permits to interpret the piezometric
564 data in a few seconds without the user having to add new parameters.

565 ✓ The CNN-HT code provides predictions that require any initial hydraulic transmissivity 566 models or computation of sensitivity matrix, as is the case with deterministic methods. This 567 advantage makes CNN-HT an effective inversion tool for dealing with highly nonlinear inverse 568 problems with large-scale parameterization that are difficult to handle with deterministic codes, 569 especially when the computation of sensitivity is complex.

570 ✓ The CNN-HT code performs reconstructions of the transmissivity field that depend on 571 the a priori model used in the generation of transmissivity field for the training data. This 572 dependence to the a priori model is a property shared with other inversion methods. However, 573 in this CNN-HT tool, the mathematical form of this a priori model is not subject to any 574 condition, which is not the case with deterministic methods where the model must have a 575 quadratic and differentiable form.

576

577 ✓ The CNN-HT algorithm also provides great flexibility for the user to incorporate
578 multiple prior models, especially when in doubt about which a priori information to use.

 \checkmark The proposed method is based on a simple network that can be generalized and applied 579 to handle hydraulic tomography in 3D or in time domain in order to identify the storativity 580 coefficient as well, by only modifying the training dataset. It is also easily adaptable to deal 581 with hydraulic tomographies with complex heterogeneities having for example multiple 582 583 discontinuous hydrofacies with contrasting transmissivities, or parameterized in discrete fracture network mode. These types of problems can be treated as a classification task by 584 replacing the regression layer of the network with the Softmax layer. In this way, the network 585 586 can identify fracture geometries or hydrofacies shapes.

587 ✓ The code allows the incorporation of uncertainties on the boundary conditions imposed
588 in the numerical solving of the forward problem by randomly choosing their values within a
589 confidence interval when building the training data set.

✓ This code, like other deep learning algorithms that use the convolutional neural network technique are able to better reduce the effects of noise in the data. This is thanks to the convolution concept that processes the input data with some local connectivity and not in a pointwise manner as in other conventional inversion methods (Vu and Jardani, 2021). The impact of noise could be minimized and accounted for in the inversion results by contaminating the hydraulic data used in the training data with a noise signal.

596 Regarding the disadvantages of this method, we mention that

597 \checkmark It is based on a repetitive and time-consuming numerical resolution of the forward 598 problem, which is involved in the construction of the training database. The size of this database 599 conditions the quality of the predictions, and its optimal size can only be obtained by a long 600 trial-and-error procedure.

601

 \checkmark The method requires important computing resources to carry out the learning.

✓ The code also uses a large number of parameters that intervene in the construction of
 the networks (such as: size of the decoder-encoder, number of filters and their size), and in the
 optimization algorithm used in the training phase (number of epochs, regularization parameters,
 initialization of filters, Batchsize, Dropout). The determination of all these parameters implies
 a tedious sensitivity analysis with several training operations to choose the best parameters.

607 ✓ The method does not provide an uncertainty assessment of the predicted transmissivity 608 fields. This is a major weakness of this method compared to classical approaches. There are 609 some attempts in the literature to estimate the uncertainty of the prediction determined with DL 610 algorithms by performing the learning in a probabilistic way (Kendall et al., 2016; Gal and 611 Ghahramani, 2016). However, this type of approach requires a lot of computational time and it 612 does not take into account all the uncertainties that can come from all the parameters used in 613 the learning.

614 \checkmark The method is highly dependent on prior information. This dependence can be an obstacle to its application, especially when the network fails in the learning phase to relate the 615 616 transmissivity fields to their piezometric data. This can occur when the transmissivity fields 617 have complex heterogeneities that are not captured by low piezometric resolution. This dependence can also lead to a misinterpretation, even with a well-trained inversion operator, 618 when the magnitudes of the piezometric data used in the training are so different from those we 619 desire to predict. To avoid both of these problems, it is necessary to ensure that generated 620 transmissivity fields have a degree of heterogeneity that can be mapped with the piezometer 621 coverage and that they produce hydraulic head data that include the magnitude of the hydraulic 622 observations to be interpreted. 623

In conventional inversion methods, the predicted transmissivity distribution is the result of acompromise between the a priori model and the piezometric data. However, here, the result is

strongly linked to the a priori model and the piezometric data to be interpreted come only at theend of the learning process.

628

629

Acknowledgements : We thank Normandy region for its financial support to our consortiumon the hydraulic characterization of aquifers.

632 Appendix

•

Here we present the mathematical formulation to analytically calculate the Jacobian matrix of
a homogeneous transmissivity field. This matrix is used in the projection method to transform
the hydraulic data to the same size as the transmissivity field to form the Segnet network.

We recall that in the joint state method, the sensitivity matrix can be expressed as an integral(Sykes et al.,1985):

638
$$\frac{\partial h_i}{\partial m_j} = -\log 10 \iint_{\Omega_j} \left[T \nabla h \nabla \phi \right] d\Omega \quad ; \qquad A1$$

639 Where *h* is the hydraulic head measured at location $O(x_{obs}, y_{obs})$ due to a pumping source 640 placed at $P(x_p, y_p)$ with a rate q_p . ϕ is the adjoint operator with a source term $q_{adj} = -1$ 641 placed at the observation point $O(x_{obs}, y_{obs})$. *T* is transmissivity hydraulic with *m*=-log(*T*). 642 When the transmissivity is constant, *h* and ϕ can be determined analytically with Thiem's 643 equation for a confined aquifer

644
$$h = h_0 + \left(\frac{q_p}{2\pi T}\right) \cdot \ln\left(\frac{R}{r_p}\right) \text{ and } \phi = \phi_0 + \left(\frac{q_{adj}}{2\pi T}\right) \cdot \ln\left(\frac{R}{r_{obs}}\right); \quad A2$$

645 with
$$r_{obs} = \sqrt{(x_{obs} - x)^2 + (y_{obs} - y)^2}$$
 and $r_p = \sqrt{(x_p - x)^2 + (y_p - y)^2}$

646 h0 and \notin are the values imposed as the boundary of the domain located at a distance R, away 647 from the study area. *r* is radial distance. Then the sensitivity of hydraulic head at point O to a 648 variation of the transmissivity at cell Ω_{i} can be approximated with this integral:

649
$$\frac{\partial h_i}{\partial m_j} = -\log(10) \frac{q_p q_{adj}}{(4\pi^2 T)} \iint_{\Omega_j} \frac{(x_{obs} - x)(x_p - x) + (y_{obs} - y)(y_p - y)}{r_{obs}^2 \cdot r_p^2} d\Omega$$
 A3.

650

651 **References**:

Alcolea A., Castro E., Barbieri M., Carrera J., Bea S., 2007. Inverse modeling of coastal
aquifers using tidal response and hydraulic tests, *Groundwater*, 45(6), 711-722.

Badrinarayanan V., Kendall A., Cipolla R., 2017. Segnet: A deep convolutional encoderdecoder architecture for image segmentation, *IEEE transactions on pattern analysis and machine intelligence*, 39(12), 2481-2495.

Berg S.J., Illman W.A., 2011. Three-dimensional transient hydraulic tomography in a highly
heterogeneous glaciofluvial aquifer-aquitard system. *Water Resources Research*, 47, W10507.

Bohling G.C., Zhan X., Butler J.J.J., Zheng L., 2002. Steady shape analysis of tomographic
pumping tests for characterization of aquifer heterogeneities, *Water Resources Research*,
38(12), 1-15.

Cardiff M., Kitanidis P., 2008. Efficient solution of nonlinear, underdetermined inverse
problems with a generalized PDE model. *Computers & Geosciences*, 34(11), 1480-1491.

Cardiff M., Barrash W., Kitanidis P., 2013. Hydraulic conductivity imaging from 3-D transient
hydraulic tomography at several pumping/observation densities, *Water Resources Research*,
49(11), 7311-7326.

Castagna M., Bellin A., 2009. A Bayesian approach for inversion of hydraulic tomographic
data, *Water Resources Research*, 45(4), W04410.

- Elsheikh A.H., Jackson M.D., Laforce T.C., 2012. Bayesian Reservoir History Matching
 Considering Model and Parameter Uncertainties, *Mathematical Geosciences*, 44, 515-543.
- 671 Fernández-Martínez J.L., Mukerji T., García-Gonzalo E., Fernández-Muñiz Z., 2011.
- 672 Uncertainty assessment for inverse problems in high dimensional spaces using particle swarm
- optimization and model reduction techniques. *Mathematical and Computer Modelling*, 54(11-
- **674** 12), 2889-2899.
- Fischer P., Jardani A., Soueid Ahmed A., Abbas M., Wang X., Jourde H., Lecoq N., 2017.
- Application of large-scale inversion algorithms to hydraulic tomography in an alluvial aquifer. *Groundwater*, 55(2), 208-218.
- 678 Fischer P., Jardani A., Lecoq N., 2018. Hydraulic tomography of discrete networks of conduits
- and fractures in a karstic aquifer by using a deterministic inversion algorithm, *Advances in Water Resources*, 112, 83-94.
- Fu J., Gómez-Hernández J., 2009a. A blocking Markov Chain Monte Carlo method for inverse
 stochastic hydrogeological modeling. *Mathematical Geosciences*, 41, 105–128.
- Fu J., Gómez-Hernández J., 2009b. Uncertainty assessment and data worth in groundwater flow
 and mass transport modeling using a blocking Markov chain Monte Carlo method. *Journal of Hydrology*, 364(3-4), 328-341.
- Gal Y., Ghahramani Z., 2016. Dropout as a Bayesian approximation: Representing model
 uncertainty in deep learning. Proceedings of the 33rd International Conference on Machine
 Learning.
- Gernez S., Bouchedda A., Gloaguen E., Paradis D., 2019. Comparison between hydraulic
 conductivity anisotropy and electrical resistivity anisotropy from tomography inverse
 modeling, *Frontiers in Environmental Science*, 7, 67.
- Gottlieb J., Dietrich P., 1995. Identification of the permeability distribution in soil by hydraulic
 tomography, *Inverse Problems*, 11(2), 353–360.
- Indolia S., Goswami A.-K., Mishra S.P., Asopa P., 2018. Conceptual understanding of
 convolutional neural network a deep learning approach, *Procedia computer science*, 132, 679688.

- 697 Ioffe S., Szegedy C., 2015. Batch normalization: Accelerating deep network training by
 698 reducing internal covariate shift, Proceedings of the 32nd International Conference on Machine
 699 Learning, 37, 448-456.
- Jimenez S., Mariethoz G., Brauchler R., Bayer P., 2016. Smart pilot points using reversiblejump Markov-chain Monte Carlo, *Water Resources Research*, 52, 3966-3983.
- Jardani A., Dupont J.-P., Revil A., Massei N., Fournier M., Laignel B., 2012. Geostatistical
- inverse modeling of the transmissivity field of a heterogeneous alluvial aquifer under tidal
- influence, *Journal of Hydrology*, 472, 287-300.
- Kendall A., Cipolla R., 2016. Modelling uncertainty in deep learning for camera relocalization,
- 2016 IEEE international conference on Robotics and Automation (ICRA), 4762-4769.
- Kingma D.P., Ba J., 2014. Adam: A method for stochastic optimization, InternationalConference on Learning Representations (ICLR).
- 709 Kitanidis P.K., 1995. Quasilinear geostatistical theory for inversing, *Water Resources*710 *Research*, 31(10), 2411-2419.
- Kitanidis P.K, 1997. *Introduction to geostatistics: applications in hydrogeology*, Cambridge
 University Press.
- Laloy E., Hérault R., Jacques D., Linde N., 2018. Training-image based geostatistical inversion
- using a spatial generative adversarial neural network, *Water Resources Research*, 54(1), 381406.
- LeCun Y., Bottou L., Bengio Y., Haffner P., 1998. Gradient-based learning applied to document
 recognition, *Proceedings of the IEEE*, 86(11), 2278-2324.
- Li W., Nowak W., Cirpka O., 2005. A Geostatistical inverse modeling of transient pumping
 tests using temporal moments of drawdown, *Water Resources Research*, 41(8).
- Lochbühler T., Doetsch J., Brauchler R., Linde N., 2013. Structure-coupled joint inversion of
 geophysical and hydrological data, *Geophysics*, 78(3).
- Neuman S.P., 1987. Stochastic continuum representation of fractured rock permeability as an
- alternative to the R.E.V. and fracture network concepts, *Rock Mechanics: Proceedings of the*
- 724 *28th*, 533-561.

- Oliver D.S., Cunha L.B., Reynolds, A.C., 1997. Markov chain Monte Carlo methods for
 conditioning a permeability field to pressure data, *Mathematical Geology*, 29, 61-91.
- Puzyrev V., Swidinsky A., 2021. Inversion of 1D frequency-and time-domain electromagnetic
 data with convolutional neural networks, *Computers & Geosciences*, 149, 104681.
- Remy N., Boucher A., Wu J., 2009. Applied Geostatistics with SGeMS: A User's Guide,
- 730 Cambridge University Press.
- Rizzo D.M., Dougherty D.E., 1994, Characterization of aquifer properties using artificial neural
- networks: Neural kriging, *Water Resources Research*, 30(2), 483-497.
- 733 Scales J.A., Smith M.L., Fischer T.L., 1992. Global optimization methods for multimodal
- inverse problems, *Journal of Computational Physics*, 103(2), 258-268.
- Shen C., 2018. A transdisciplinary review of deep learning research and its relevance for water
 resources scientists, *Water Resources Research*, 54(11), 8558-8593.
- Shigidi A., Garcia L.A., 2003. Parameter estimation in groundwater hydrology using artificial
 neural networks, *Journal of Computing in Civil Engineering*, 17(4), 281-289.
- 739 Soueid A., Jardani A., Revil A., Dupont J.-P., 2016. Specific storage and hydraulic conductivity
- tomography through the joint inversion of hydraulic heads and self-potential data, *Advances in*
- 741 *Water Resources*, 89, 80-90.
- Sun A.Y., 2018. Discovering State-Parameter Mappings in Subsurface Models Using
 Generative Adversarial Networks, *Geophysical Research Letters*, 45(20), 11137-11146.
- 744 Sykes J.F., Wilson J.L., Andrews R.W., 1985. Sensitivity analysis for steady state groundwater
- flow using adjoint operators, *Water Resources Research*, 21(3), 359-371.
- Tarantola A., Valette B., 1982. Generalized nonlinear inverse problems solved using the least
- squares criterion, *Reviews of Geophysics*, 20(2), 219-232.
- Tso C.H.M., Za Y., Yeh T.C.J., Wen J.C., 2016. The relative importance of head, flux, and prior
- information in hydraulic tomography analysis, *Water Resources Research*, 52(1), 3-20.
- 750 Vu M.T, Jardani A, 2021. Convolutional neural networks with SegNet architecture applied to
- three-dimensional tomography of subsurface electrical resistivity: CNN-3D-ERT, *Geophysical*
- 752 Journal International, 225 (2), 1319–1331, https://doi.org/10.1093/gji/ggab024

- Wang X., Jardani A., Jourde H., 2017. A hybrid inverse method for hydraulic tomography in
 fractured and karstic media, *Journal of Hydrology*, 551, 29-46.
- Wu Y., Lin Y., 2018. InversionNet: A Real-Time and Accurate Full Waveform Inversion with
 CNNs and continuous CRFs, arXiv preprint arXiv:1811.07875.
- Yeh T.C.J., Liu S., 2000. Hydraulic tomography: Development of a new aquifer test
 method. *Water Resources Research*, *36*(8), 2095-2105.
- 759 Zhao Z., Illman W.A., Berg S.J., 2016. On the importance of geological data for hydraulic
- tomography analysis: laboratory sandbox study, *Journal of Hydrology*, 542, 156-171.
- Zhu Y., Zabaras N., 2018, Bayesian deep convolutional encoder-decoder networks for
 surrogate modeling and uncertainty quantification, *Journal of Computational Physics*, 366,
 415-447.
- Zio E., 1997, Approaching the inverse problem of parameter estimation in groundwater models
- by means of artificial neural networks, *Progress in Nuclear Energy*, 31(3), 303-315.