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1 Mica trace-element signatures: highlighting superimposed W -

2 Sn mineralizations and fluid sources

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9 Highlights

- 10 • Mica generations are characterized by unique trace-element signatures.
- 11 • Igneous lepidolite is the main host of W in the Beauvoir granite body.
- 12 • Alteration of igneous micas can lead to economic W deposition.
- 13 • Greisen micas can incorporate significant Sn, inhibiting cassiterite deposition.

14 Abstract

15 In the Echassières district of the French Massif Central, occur several outstanding magmatic/hydrothermal
16 systems enriched in strategic metals, such as the Beauvoir rare-metal granite. In this contribution we
17 propose a systematic approach, based on mica trace chemistry, to decipher the different events leading to
18 mineralization. Twelve groups of micas were defined by their specific petrographic features and/or location
19 in the district. Their trace element composition, obtained by LA-ICP-MS, varies widely from one mica group
20 to another, although homogeneous signatures within groups could be distinguished. Some of the trace
21 elements are remarkably enriched, such as W in igneous lepidolite and Sn in greisen muscovite, both of
22 which occur in the Beauvoir granite. A statistical approach based on a set of multivariate analyses highlights
23 that the trace chemistry of micas is inherited from their source, whether hydrothermal or igneous, thus
24 providing a signature for their origin. This approach also shows that differences in major element

25 composition (i.e., different mica species) impact only slightly the trace-element signature. For instance,
26 muscovite and zinnwaldite from one granite have a coherent signature, but they contrast with same mica
27 species in another granite or hydrothermal veins. It is thus possible to genetically link two different mica
28 species from remote locations, or inversely, to recognize different origins for a same mica species in the
29 same sample (e.g., superposed alterations). A second, important implication is that trace-element
30 signatures of micas provide a record of metal remobilization and transport. In the Echassières district,
31 greisen alteration of the Beauvoir granite caused dissolution of W-rich (ca. 290 ppm in average) lepidolite
32 and cassiterite (SnO_2). Newly-formed greisen muscovite incorporated most remobilized Sn (ca. 1000 ppm in
33 average), while W precipitated in distal quartz veins as wolframite. As a consequence, Sn is concentrated in
34 the granite, while W occurs outside of it. This is also indicated by the gradual Sn decrease and W increase
35 recorded in micas from distal veins. Finally, wolframite-bearing veins do not contain cassiterite, validating
36 mica trace-element chemistry as a powerful tool to decipher Sn-W ore forming hydrothermal processes.

37 **Keywords**

38 Micas, trace-element signature, LA-ICP-MS, Beauvoir granite, Echassières district.

39 **1. Introduction**

40 Minerals of the mica group can incorporate a wide variety of elements due to their sheet structure and
41 countless substitutions permitted by complex multiple solid solutions (Foster, 1960; Hazen and Wones,
42 1972; Iiyama and Volfinger, 1976; see also Brigatti et al., 2002 and references therein). Evolved granitic
43 systems and related hydrothermal assemblages commonly display complex parageneses that include a
44 variety of mica species (more or less rich in F, Li, Rb, Cs, Fe, Mg; Tischendorf et al., 2001). These micas can
45 exhibit a remarkable enrichment in incompatible elements, to the point that they are considered important
46 reservoirs of incompatible LILE (Large Ion Lithophile Elements) in various rocks (Beswick, 1973; Nash and
47 Crecraft, 1985; Adam and Green, 2006; Bebout, et al., 2007; Hulsbosch et al., 2014). Moreover, it has
48 recently been shown that micas are good competitors for the HFSE (High Field-Strength Elements; also
49 highly incompatible) against element-specific accessory phases (e.g., see Zhu et al., 2018 and Yang et al.,
50 2020, for Nb-Ta and **Breiter et al.**, Guimarães et al., 2021, for Sn). Consequently, crystallization (or
51 inversely, dissolution) of a specific mica generation during hydrothermal alteration can drastically affect the
52 fate of these metals and the potential for mineralization.

53 In granitic systems, W mineralization is usually related to acidic alteration zones (e.g., greisen in
54 peraluminous granites, sericitization in porphyries) that are rich in micas. Because of this, a growing
55 number of researchers are turning to the study of the trace composition of mica from Sn-W mineralized
56 systems (e.g., Li et al., 2015, 2018; Xie et al., 2015, 2019; Legros et al., 2016, 2018; Breiter et al., 2017,
57 2019; Gao et al., 2019; Sun et al., 2019; Yin et al., 2019; Azadbakht and Lentz, 2020; Azadbakht et al., 2020;
58 Chen et al., 2020; Xie et al., 2020; Wei et al., 2020; Codeço et al., 2021; **Duan et al., 2021**; Guimarães et al.,
59 2021; Shi et al., 2021; Li et al., 2021). An interesting outcome of these studies is that most micas have much
60 higher W, Sn and Nb-Ta contents than micas in barren magmatic-hydrothermal systems. However, it is also
61 reported that several trace elements common in these micas can vary by few orders of magnitude in the
62 different deposits (Breiter et al., 2017; Codeço et al., 2021; Wei et al., 2020), and even within a single
63 deposit (e.g., Legros et al., 2018; Breiter et al., 2019). As a consequence, these studies contain as many

64 specific trace-element signatures as there are analyzed mica generations. To enhance our knowledge on
65 the potential applications of trace element distribution in micas, we investigate the Echassières

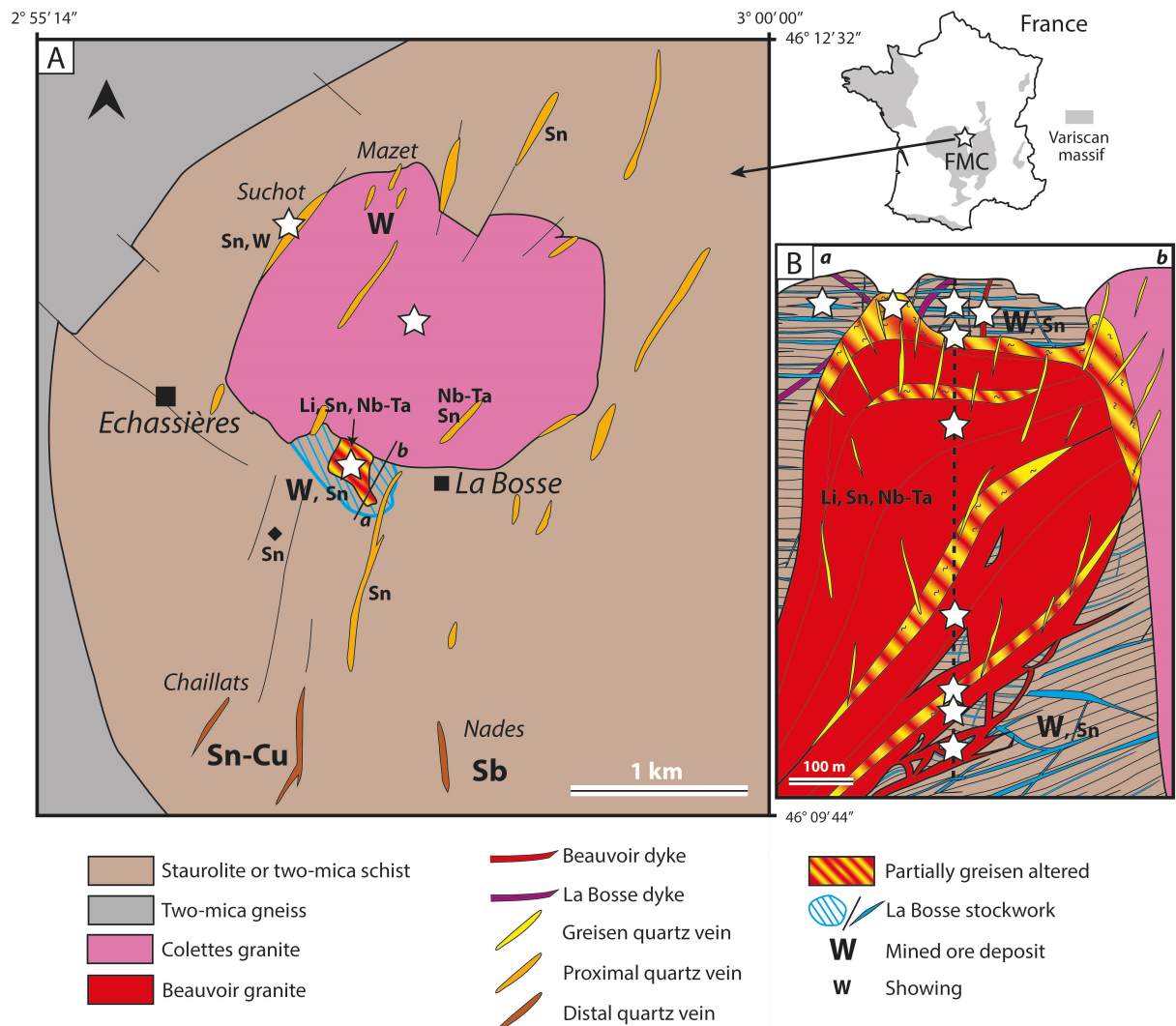
66 district (France), a particularly complex system that provides a unique opportunity to monitor the evolution
67 of mica composition across a protracted ore depositing event, characterized by several mineralizing
68 episodes.

69 **2. Geological settings**

70 The Echassières district (Fig.1.A) is a long-lived polymetallic (Li, W, Sn, Nb-Ta, Sb) deposit located in the
71 Sioule nappe of the northern French Massif Central. The architecture of the Echassières district consists of
72 an antiform where three metasedimentary units are stacked, intruded in their core by the Colettes two-
73 micas granite and to the south by the smaller Beauvoir rare-metal granite. The bulk of the mineralization is
74 restricted to within a radius of 3 km around the Beauvoir granite (Fig.1.B). Ore deposition took place during
75 different stages of the Variscan belt evolution, likely ranging from subduction and arc magmatism (>360
76 Ma) to phases of orogenic collapse (310 Ma; Monnier et al., 2021a). The peak of the Barrovian
77 metamorphism in the Sioule nappe occurred at 350-360 Ma (Do Couto et al., 2016), with temperature and
78 pressure estimated at ca. 600 °C and 7 kb for the lower unit of the antiform (Schulz, 2009).

79 At least three generations of wolframite (Monnier et al., 2019, 2020) and two generations of stibnite
80 (Monnier et al., 2021b) occur in the Sioule nappe. Some of these mineralizations were mined in the 20th
81 century (Mazet and La Bosse localities for the wolframite and Nades for the stibnite; Aubert and
82 Contamine, 1958; Aubert, 1969), while columbo-tantalite is presently recovered as by-product of kaolinite
83 extraction at the Beauvoir granite quarry. Intrusion of Beauvoir/Colettes complex was the most recent, and
84 consequently the best preserved magmatic/hydrothermal event in the zone. The Beauvoir rare metal
85 granite (RMG) has been intensively investigated in the past by a number of authors (Fouillac and Rossi,
86 1991; Cuney et al., 1992, Raimbault et al., 1995), who concord to evoke extreme fractionation processes
87 coupled to massive meteoric fluid incorporation to explain the outstanding enrichment of incompatible
88 elements at the magmatic stage. More recently, the hydrothermal events were investigated in depth to
89 unravel the several alteration episodes superimposed in the host schist. Stockwork quartz veins (>360 Ma)
90 and topaz veins (ca. 335 Ma) were recognized as distinct igneous-related hydrothermal episodes that took
91 place well before crystallization of the Beauvoir granite (ca. 310 Ma; Harlaux et al., 2017; Monnier et al.,

92 2019, 2020). Three alteration episodes affected the Beauvoir granite and surrounding rocks, to variable
93 extent. An early orthomagmatic stage, initiated by exsolution of a hot (> 450 °C) fluid during granite
94 crystallization, induced acidic F-rich and alkali alterations, which are manifested mostly in contact zones
95 between the granite and the host rock (Cuney et al., 1992; Monnier et al., 2020). A second stage is
96 characterized by widespread greisenization (Monnier et al., 2018; Monnier et al., 2020). Key features of the
97 greisen fluid include F-poor and P-rich sericitization, marked by the occurrence of abundant OH-muscovite
98 and apatite (Monnier et al., 2019). Greisen alteration was a continuous process, initiated within the granite
99 at ca. 450°C and evolving towards the host rock where it waned at about 200°C. Peak greisen activity in the
100 granite was contemporaneous to fracturing episodes in the area, promoting formation of proximal veins
101 around the granitic system, by percolation of the Beauvoir greisen fluid to these sites (Monnier et al., 2018,
102 2020). Disparity in W ore deposition in these veins was a function of the intensity of fracturing, which
103 induced more or less important boiling/flashing of the fluid (Monnier et al., 2020). Although it is recognized
104 that the high temperature greisen fluids were enriched in Sb (Monnier et al., 2018), mineralization did not
105 form in the granite and Sb was transported by the greisen fluid and precipitated in distal veins (Nades)
106 when the fluid cooled (Monnier et al., 2020; 2021b). The third and last stage corresponds to widespread
107 supergene alteration. This consisted in replacement of greisen muscovite and igneous feldspar by kaolinite,
108 partial replacement of wolframite by W-rich goethite, and formation of quartz veins, accompanied by low-
109 temperature minerals such as the phosphate goyazite-gorceixite, fluorite, allophane and Mn-oxides
110 (Monnier et al., 2019).



111 *Fig. 1. Simplified geology of the study area. Samples are localized by star symbols. [A]: Geological map of the*
 112 *Echassières district. The a-b segment indicates the cross-section depicted in [B]. [B]: A sketch of the Beauvoir*
 113 *granite along the a-b cross-section, constructed using information from surface outcrops and from a GPF bore-*
 114 *hole projection (deep drilling project of France; located by the dashed line in [B]). Vein and dyke thicknesses are*
 115 *not to scale.*

116 3. Material and methods

117 Samples consist of outcrop and drill core material that was collected on site or from storage at the French
 118 Geological Survey (BRGM, Orleans, France). Thin and thick sections (30 μm and ca. 200 μm , respectively)
 119 were prepared from each sample.

120 The trace-element content of micas was analyzed by LA-ICP-MS using a UV laser-ablation micro-sampler
 121 CETAC® Excite (193 nm) coupled to a ThermoScientific® X serie II quadrupole plasma mass spectrometer, at
 122 the BRGM laboratory in Orleans. The laser configuration was optimized by performing multiple trial
 123 measurements on several magmatic and hydrothermal micas with different resistance to the laser

124 impulsion. The laser was equipped with a HelEx[®] 2 volume ablation chamber. The frequency was 8 Hz,
125 fluency 6 J.cm⁻², and spot diameter 50 μm. After ablation, the material is transported by a He flux of
126 0.3L/min and mixed with an Ar flux of 1L/min, before entering the plasma. The ICP RF power is 1400 W. All
127 micas crystals were analyzed for Li, Na, Mg, Al, K, Ti, Fe, Ge, Rb, Sr, Nb, Sn, Cs, Sb, Ba, Ta, W. Acquisition
128 consisted of 40 s runs (320 laser pulses), separated, before and after, by 20 s of background recording.
129 Presence of mineral inclusions was verified by monitoring the presence of anomalous peaks on the flat part
130 of the signal, and excluded from the integration of the signal. To avoid background noise from analyzes of
131 Rb-rich micas, these micas groups were studied after the Rb-poor micas. Detection limits (d.l.) are given in
132 Table A2 (Appendices) for each element. Acquisition errors are lower than 10 % for concentrations above
133 d.l. by at least 1 order of magnitude. Data reduction was undertaken using Glitter[®] software, calibrated
134 using NIST SRM 612 as external standard (concentration from Pearce et al., 1997). Silica was used as
135 internal standard reported to values measured by Electron Probe Micro-Analyser (EPMA) for all mica
136 groups considered (Table A1 in Appendices). EPMA were performed at the Centre Castaing - University Paul
137 Sabatier (Toulouse, France) using a CAMECA SXFive. Surface areas of 2x2 μm were analyzed using
138 accelerating voltage 15kV and a beam current of 10nA. The following standards were used for calibration:
139 topaz (F), albite (Na), periclase (Mg), corundum (Al), tugtupite (Cl), sanidine (K), wollastonite (Ca, Si),
140 pyrophanite (Mn), hematite (Fe), Ti (pyrophanite) chromium oxide (Cr), barite (Ba), and Glass Ge-Al-Ca-Rb
141 (Rb).

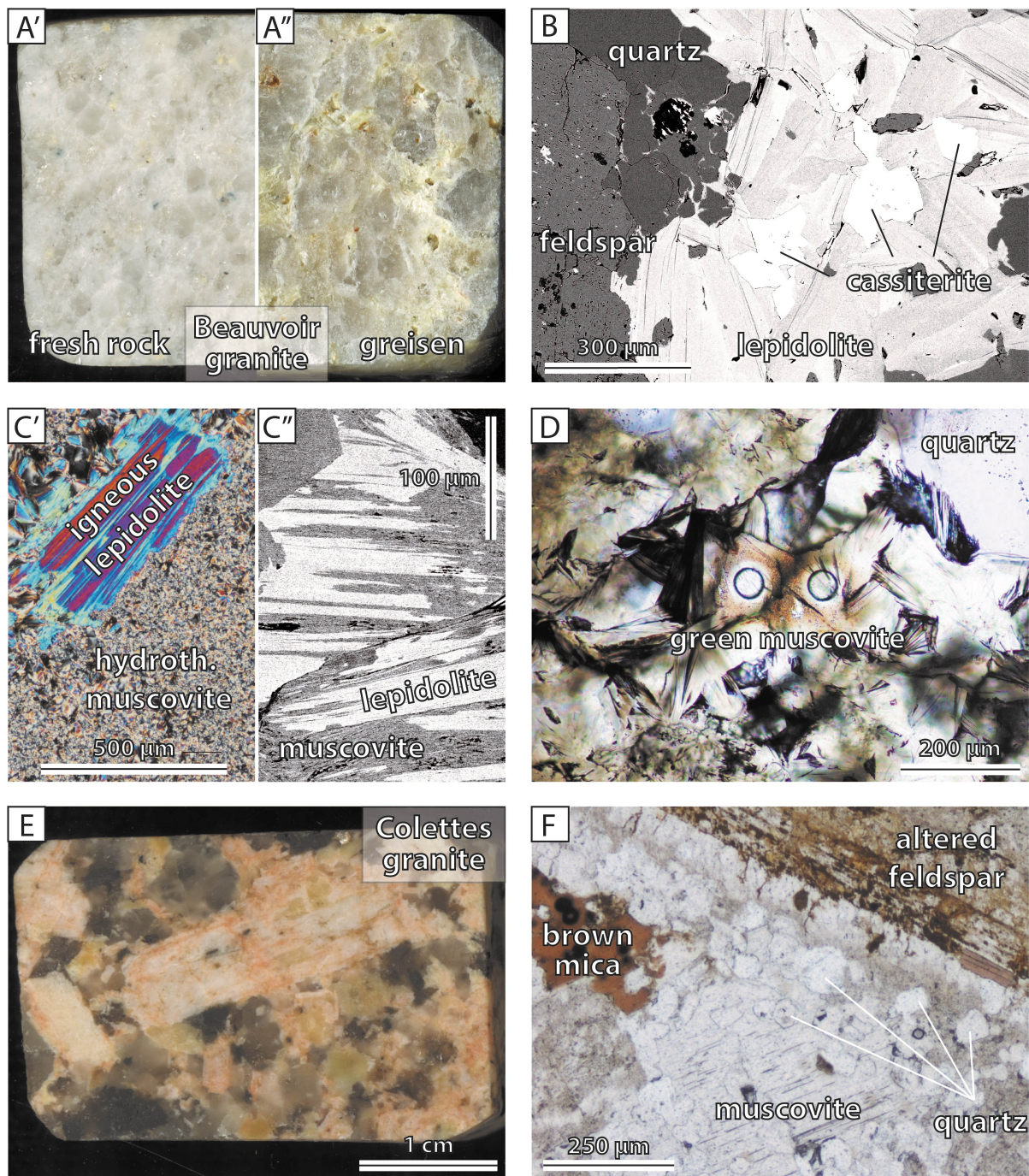
142 **4. Results**

143 **4.1. Petrography**

144 In this section, we describe the main petrographic features of the different magmatic facies and
145 hydrothermal episodes considered in this study, focusing in particular on the characteristics of each mica
146 generation. For specific information on the entire paragenesis and the diverse variety of facies and textures
147 occurring in the area, the reader is referred to Monnier et al. (2018, 2019, 2020), where igneous and
148 hydrothermal textures are fully detailed.

149 4.1.1. *Beauvoir and Colettes granites.*

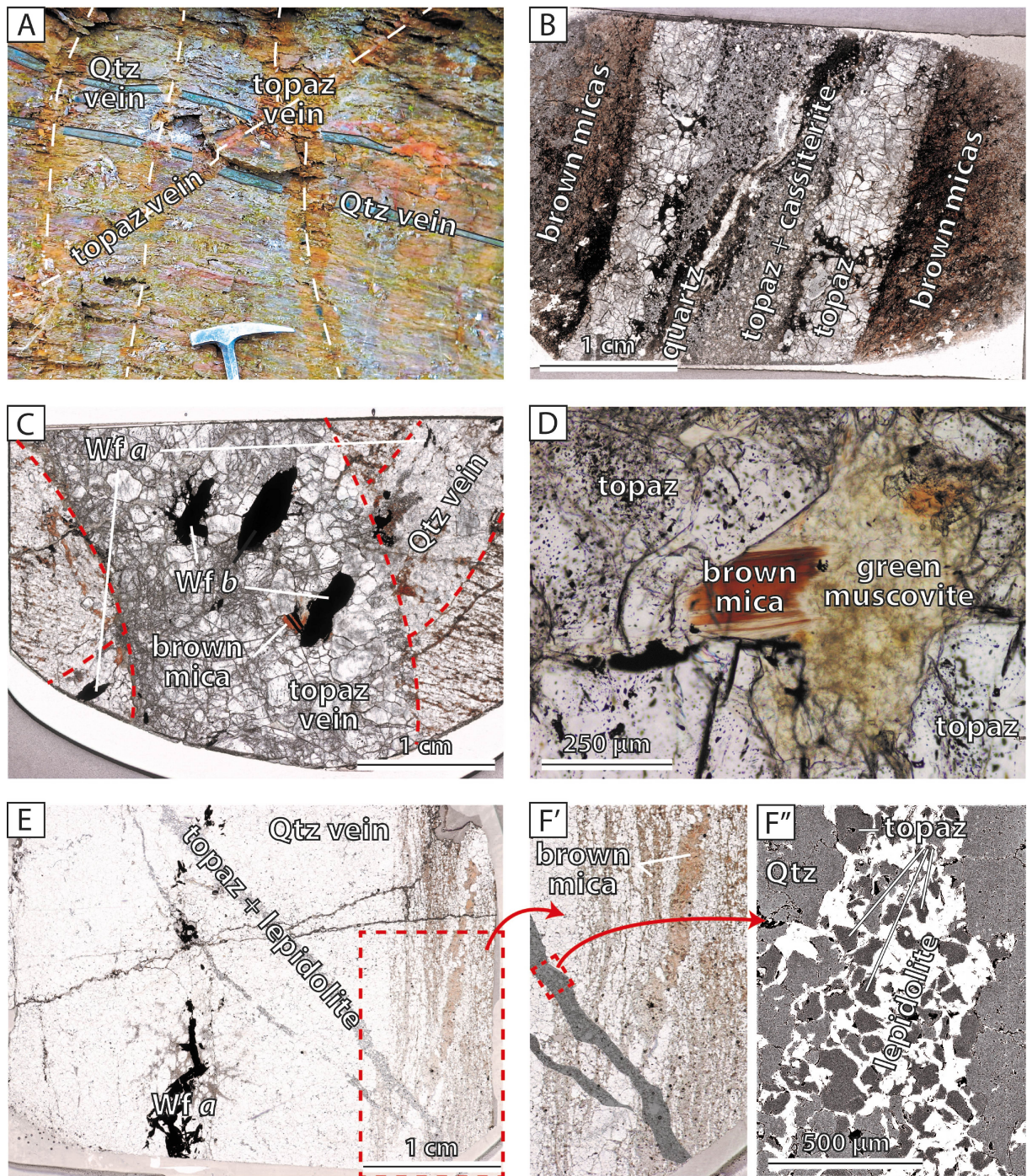
150 The hololeucocratic equigranular Beauvoir granite is made up of albite, quartz, lepidolite, subordinate
151 variable amounts of topaz and potassic feldspar, plus a wide diversity of accessory minerals (Fig. 2.A').
152 Among these, cassiterite occurs in the entire granite body (Fig. 2.B), while columbo-tantalite and
153 pyrochlore are mostly concentrated in the apex of the granite body. Greisen alteration caused dissolution
154 of the magmatic minerals (Fig. 2.A''), replacing them by various proportions of quartz, muscovite and, to a
155 lesser extent, apatite. In most of the greisen zones, igneous quartz is only partially altered while feldspar
156 and lepidolite were more efficiently removed. In some samples, however, igneous as well as hydrothermal
157 quartz is totally absent and only greisen muscovite replaces magmatic minerals. In this particular case, it is
158 possible to observe relicts of igneous lepidolite, even if all the igneous quartz was dissolved (Fig. 2.C'.C'').
159 Greisen muscovite can be green in color near the edges of quartz veins and at the borders between greisen
160 and fresh rock (Fig. 2.D), while pale-yellow muscovite mainly occurs in the central parts of greisenized
161 areas. The Colettes granite corresponds to a porphyritic two-mica intrusive (Fig. 2.E). The phenocrysts
162 consist of euhedral feldspar and quartz, with subordinate amount of cordierite. Muscovite, brown mica and
163 quartz form the phaneritic matrix (Fig. 2.F).



164 Fig. 2. Petrography of the Beauvoir/Colettes granitic bodies. [A']: high resolution scan of a thin section block
 165 showing typical unaltered Beauvoir granite texture. [A'']: High resolution scan of a thin section block depicting
 166 quartz-rich greisen. [B]: Scanning Electron Microscopy (SEM) backscattered electron (BSE) image of the typical
 167 igneous assemblage of the Beauvoir granite, i.e., feldspar, quartz, lepidolite and cassiterite. [C']: Cross-polarized
 168 light (XPL) thin section image of an uncommon greisen texture with lepidolite as the only igneous relict in a
 169 sericitized matrix composed exclusively of muscovite (no quartz). [C'']: SEM BSE image illustrating partial
 170 replacement of igneous lepidolite by greisen muscovite. In this sample, lepidolite is more strongly altered than
 171 other igneous minerals. [D]: Plane-polarized light (PPL) thin section image of greisen samples exhibiting sprays
 172 of hydrothermal green muscovite crystals plus quartz. Two laser ablation holes are visible. [E]: High resolution
 173 scan of a thin section block from the two-mica Colettes granite. [F]: PPL thin section image of Colettes granite
 174 showing textural relationships between zinnwaldite, muscovite and other igneous phases.

175 4.1.2. The La Bosse stockwork.

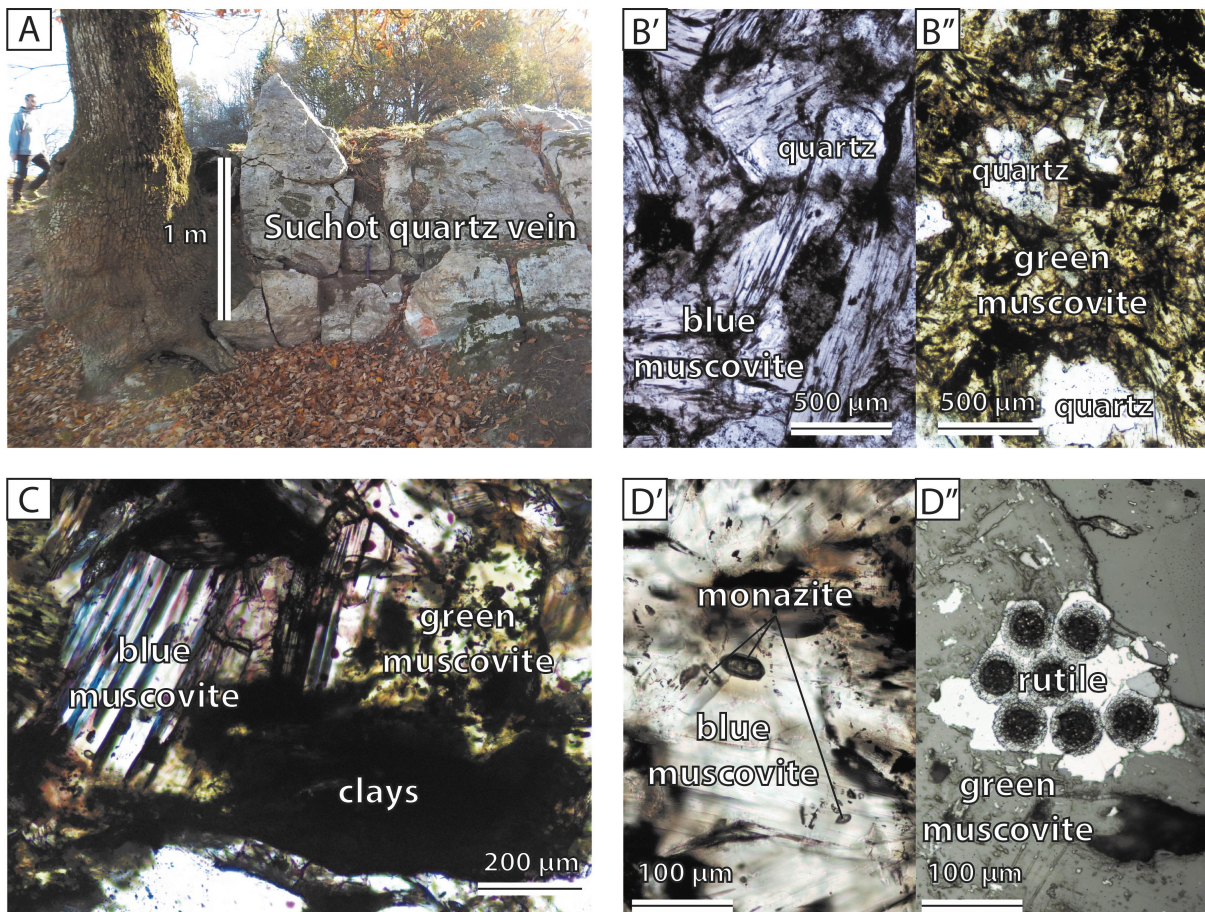
176 A swarm of sub-horizontal quartz veins, known as the La Bosse stockwork (Fig. 1.A,B), is emplaced in the
177 schist and forms the earliest hydrothermal event in the region. These veins exhibit all the later alteration
178 types that have been recorded in the Sioule nappe. The first hydrothermal episode is evidenced by the
179 formation of wolframite (generation *a*) in the quartz veins (Fig. 3.A,C,E), accompanied by formation of
180 alteration haloes in the schist, consisting of brown mica (Fig. 3.F), tourmaline and ilmenite. The second
181 alteration event affected the schist and the quartz veins; it was manifested by formation of topaz veins (Fig.
182 3.A,B) and consequent pervasive alteration where quartz was replaced by topaz plus a second generation
183 of wolframite (*b*). Other minerals include brown mica, occurring in small amounts within the veins (Fig. 3.D)
184 but abundantly in the alteration halos (Fig. 3.A,B). This alteration was considered as an early F-rich greisen
185 episode of unknown magmatic origin (cf. Monnier et al., 2020). The last magmatic/hydrothermal episode to
186 affect the stockwork was coeval to the crystallization of the Beauvoir granite. Two alteration episodes can
187 be linked to this granite: 1) an F-rich orthomagmatic alteration expressed by crystallization of lepidolite, a
188 second generation of topaz, and columbo-tantalite plus cassiterite, 2) an OH-greisen, marked by the
189 occurrence of muscovite, quartz, apatite, wolframite, minor topaz and cassiterite (Fig 3.B).



190 Fig. 3. Petrography of the La Bosse stockwork. Qtz: Quartz; Wf: Wolframite. [A]: Outcrop image showing a
 191 detail of The La Bosse stockwork. Quartz veins are transposed along schistosity and intersected by subvertical
 192 topaz veins (dashed lines), visible mostly by their alteration halos. [B]: High resolution scan of a thick section
 193 depicting a polyphased topaz veins sealed firstly by an early F-rich greisen fluid (topaz) and, at some later time,
 194 by a later greisen fluid issued from the Beauvoir granite (topaz and cassiterite). The alteration halo consists
 195 mostly of hydrothermal brown micas, accompanied by rutile and monazite. [C]: High resolution scan of a thick
 196 section illustrating cross-cutting relationships between an early quartz vein carrying wolframite a and a later
 197 topaz vein carrying wolframite b and brown mica. [D]: PPL image of a brown mica crystal partially replaced by
 198 green muscovite in the core of a topaz vein. [E]: High resolution scan of a thick section depicting a wolframite a-
 199 bearing quartz vein intersected by topaz-lepidolite veinlets. [F]: Zoom on part of image [E] detailing the selvage
 200 zone between quartz vein and host schist which is partially obliterated by later metamorphism. The large brown
 201 micas crystals containing dark specks of ilmenite and tourmaline differ from the smaller metamorphic biotite in
 202 the schist groundmass and are part of the vein's alteration halo. [F'']: SEM BSE image detailing the texture of a
 203 topaz-lepidolite veinlet.

204 4.1.3. The Suchot vein.

205 A ca. one-meter-thick quartz vein occurs in the Suchot area (Fig. 4.A). Parts of this vein contain minor
206 muscovite (Monnier et al., 2019) and traces of cassiterite (Aubert, 1969). The schist within the vein
207 selvages is strongly affected by alteration halos of several meters in width. Alteration minerals include
208 quartz and two generations of muscovite (a euhedral pale-blue muscovite, Fig. 4.B', replacing anhedral
209 green variety, Fig. 4.B''). Both muscovite generations coexist near the main quartz vein (Fig. 4.C), although
210 blue muscovite fades out as one moves outwards in the selvage. Monazite and rutile crystals are commonly
211 included in the pale-blue (Fig. 4.D') and green muscovite (Fig. 4.D'') varieties, respectively.



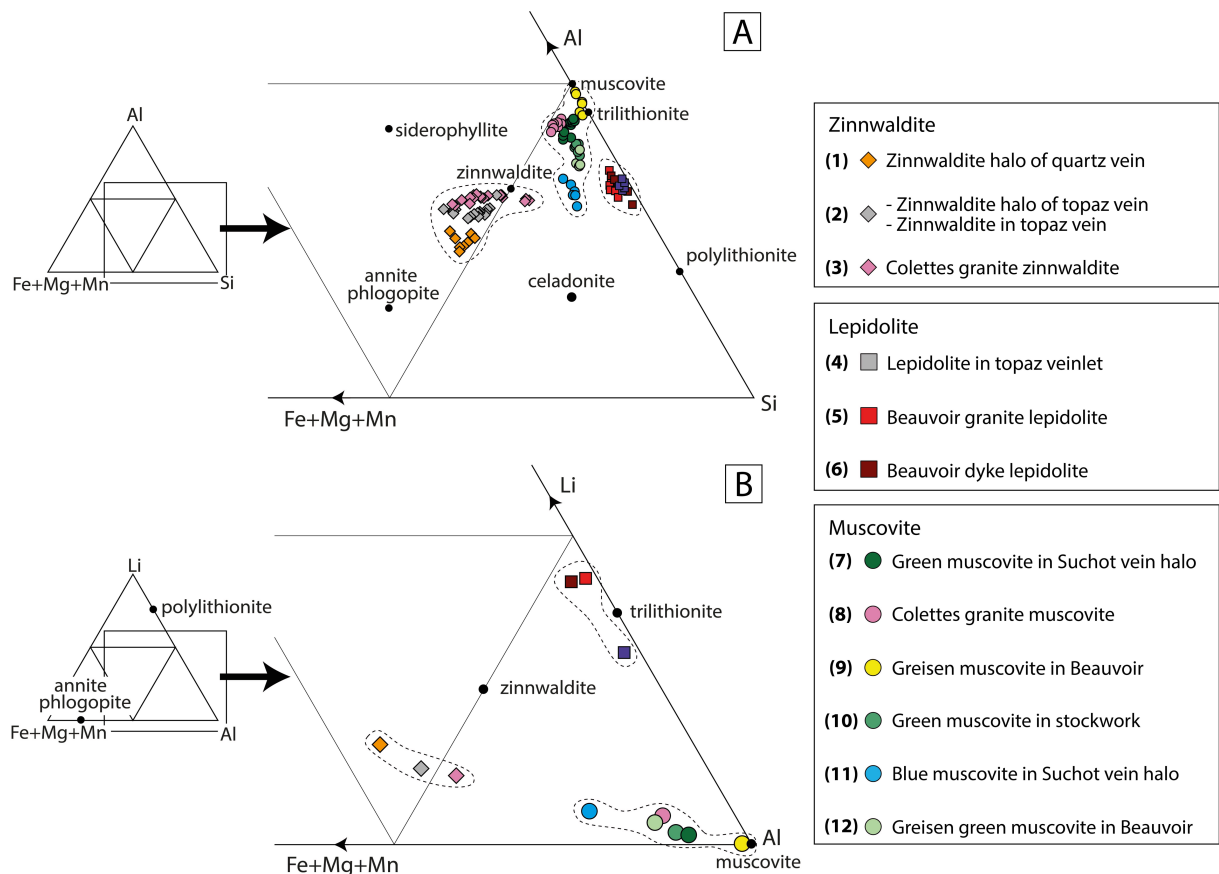
212 Fig. 4. Petrography of the selvage of the Suchot quartz vein. [A]: Vue of an outcrop showing the meter-scale
213 Suchot vein. [B']: PPL thin-section image of an assemblage of quartz and pale blue muscovite replacing
214 metamorphic minerals in the Suchot vein's selvage. [B'']: PPL thin-section image showing an assemblage of
215 quartz and green muscovite replacing metamorphic minerals in the Suchot vein's selvage. [C]: PPL thin-section
216 image illustrating coexisting pale-blue and green muscovite types. [D']: PPL thin-section image detailing
217 monazite inclusions in pale-blue muscovite. U-Pb dating allowed recognition of two monazite generations,
218 dated at 331 ± 5 and 307 ± 3 Ma (Monnier et al., 2020). [D'']: Reflected light thin-section image of a W-rich rutile
219 grain included in green muscovite. Rutile crystals have been dated at 328 ± 5 Ma (Monnier et al., 2020). Ablation
220 craters are visible.

221 **4.2. Mica characterization**

222 Based on petrographic features, including textural and optical properties as well as occurrences in different
223 assemblages, micas samples were organized into 12 groups. These are listed in Table 1, together with their
224 characteristic features. In the lines below, we classify these groups according to standard parameters based
225 on their major chemical composition. Using the Al-R²⁺-Si and Li-R²⁺-Al ternary diagrams and the
226 nomenclature of Monier and Robert (1986) and Foster (1960), a plot of the 12 groups of micas recognized
227 in the Echassières district highlights roughly 3 clusters (Fig. 5). The first cluster, formed by brown mica
228 groups (1), (2) and (3), is located within the biotite-zinnwaldite series, close to the latter end-member. The
229 cluster formed by these 3 groups are hereon referred to as zinnwaldite type. The second cluster,
230 corresponding to micas groups (4), (5) and (6), chemically consists of trilithionite with some polyolithionite
231 component, as shown by the Al-R²⁺-Si ternary diagram (Fig. 5.A). In the rest of the text, we refer to this
232 cluster as the lepidolite type (solid solution between the zinnwaldite, polyolithionite and trilithionite poles).
233 Finally, the third cluster is formed by the groups (7), (8), (9), (10), (11) and (12) and is classified as
234 muscovite type, with possibly a minor amount of celadonite component.

Petrographic groups	(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)	(9)	(10)	(11)	(12)
Micas species	Brown micas (zinnwaldite, see Fig. 5)	Brown micas (zinnwaldite, see Fig. 5)	Brown micas (zinnwaldite, see Fig. 5)	Lepidolite	Lepidolite	Lepidolite	Muscovite	Muscovite	Muscovite	Muscovite	Muscovite	Muscovite
Nature	Hydrothermal I	Hydrothermal	Igneous	Hydrothermal	Igneous	Igneous	Hydrothermal	Igneous	Hydrothermal	Hydrothermal	Hydrothermal	Hydrothermal
Picture	Fig. 3.F'	Fig. 3.C,D	Fig. 2.F	Fig. 3.F''	Fig. 2.B,C',C''	Fig. 6.F in Monnier et al., 2019	Fig. 4.B'',C,D''	Fig. 2.F	Fig. 2.C',C''	Fig.3.D	Fig. 4.B,C,D	Fig. 2.D
Sample(s)	001c	P50; 004f	Col-2	001c	P48; P4; BE05; 0003a; BE49; BE54	005f	012a	Col-2	003a	005b; P50	012a	P48
Main petrographic features	Large anhedral crystal in core of quartz veinlets. Abundant ilmenite and tourmaline inclusions	Large subhedral to euhedral crystals. Alteration of schist in vein halo. W-rutile and monazite inclusions	Phenocryst, mm-sized subhedral crystals. Rich in small quartz crystals inclusions	Anhedral small crystal, precipitating around topaz grain.	mm-sized euhedral lath, scarce spray crystals. Common cassiterite inclusions.	mm-sized euhedral lath. Common cassiterite inclusions	Anhedral and disorganized small green crystal, common spray crystal. Abundant W-zoned rutile inclusions	Phenocryst, mm-sized subhedral crystals. Rich in small quartz crystals inclusions.	Disorganized small to large anhedral pale yellow crystal, various cleavage orientation, common spray crystal	Disorganized small green crystal, various cleavage orientation. Alteration of topaz or biotite (2)	Large euhedral to subhedral crystals, pale blue coloured, parallel sheets. Rich in monazite inclusions	Disorganized small to large green, spray, crystals. Alteration of quartz, feldspar and lepidolite (5)
Associated paragenesis (more details in Monnier et al., 2018; Monnier et al., 2019)	Quartz, ilmenite, tourmaline, rutile, native bismuth wolframite a ¹	Topaz I, rutile, wolframite b ¹ , monazite	K-feldspar, plagioclase, quartz, rutile	Topaz II, cassiterite	Albite, quartz, topaz, cassiterite, columbo-tantalite series, pyrochlore	Albite, quartz, topaz, cassiterite, columbo-tantalite series,	Rutile, monazite,	K-feldspar, plagioclase, quartz, rutile	Apatite, quartz, columbo-tantalite series, pyrochlore	Scarce apatite, Wolframite c ¹ .	Quartz, monazite, xenotime	Apatite, quartz, columbo-tantalite series, pyrochlore
Location	La Bosse stockwork	La Bosse stockwork	Colettes granite	La Bosse stockwork	Beauvoir granite	La Bosse stockwork	Suchot area	Colettes granites	Beauvoir granite	La Bosse stockwork	Suchot area	Beauvoir granite
Sub location	Quartz vein selvage	Topaz veins; Topaz vein halos	-	Micro veins or pervasive alteration	-	dyke	Suchot quartz vein halo and schist	-	Greisen altered zone	Topaz veins, quartz veins	Suchot quartz vein halo	Greisen altered zone
Age (Monnier et al., 2021a)	>360 Ma	~330-335 Ma	~310 Ma	<360 Ma	~310 Ma	<335 Ma	~330 Ma	~310 Ma	~310 ²	<330 Ma	~310 Ma	~310 ²

236 Table 1: Main characteristics of micas groups analyzed by LA-ICP-MS. ¹ Three types of wolframite (a, b and c) occurs in the Echassières district, see Monnier et al., 2019 and Monnier et al.,
237 2020, for further information about W ore. ² Beauvoir greisen muscovite or associated minerals has not been directly dated, nevertheless greisen alteration initiated at hot temperature
238 (Monnier et al., 2020), during cooling of the Beauvoir granite which crystallized at ca. 310 Ma (Monnier et al., 2021a and references within).



239 Fig. 5. Mica classification ternary diagrams (Foster, 1960; Monier and Robert, 1986). Diagram [A] is constructed
 240 using EPMA data while diagram [B] is constructed using mean Li values obtained by LA-ICP-MS (undetectable
 241 with the EPMA protocol), together with EPMA data. Each group is identified by a number (in brackets) and a
 242 colored symbol, consistent through the following text and Figures. Diamonds, squares and circles correspond
 243 respectively to symbols for brown micas, lepidolite and muscovite series.

244 The three types of micas defined above (zinnwaldite, lepidolite, muscovite) are also well distinguished by
 245 their respective Li contents (Fig. 6). Lepidolite exhibits the highest Li concentrations (ca. 10,000 to 30,000),
 246 zinnwaldite has intermediate contents (ca. 4000 to 10,000 ppm), while muscovite is relatively Li-poor (few
 247 ppm to ca. 3000 ppm).

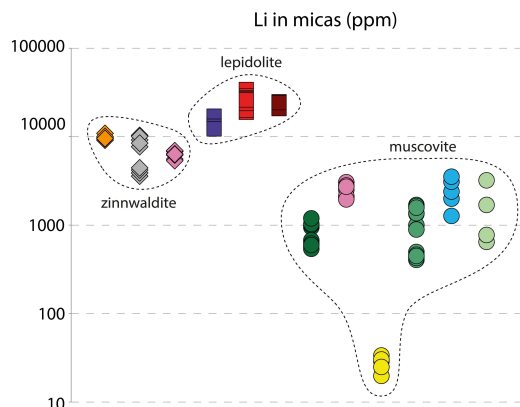
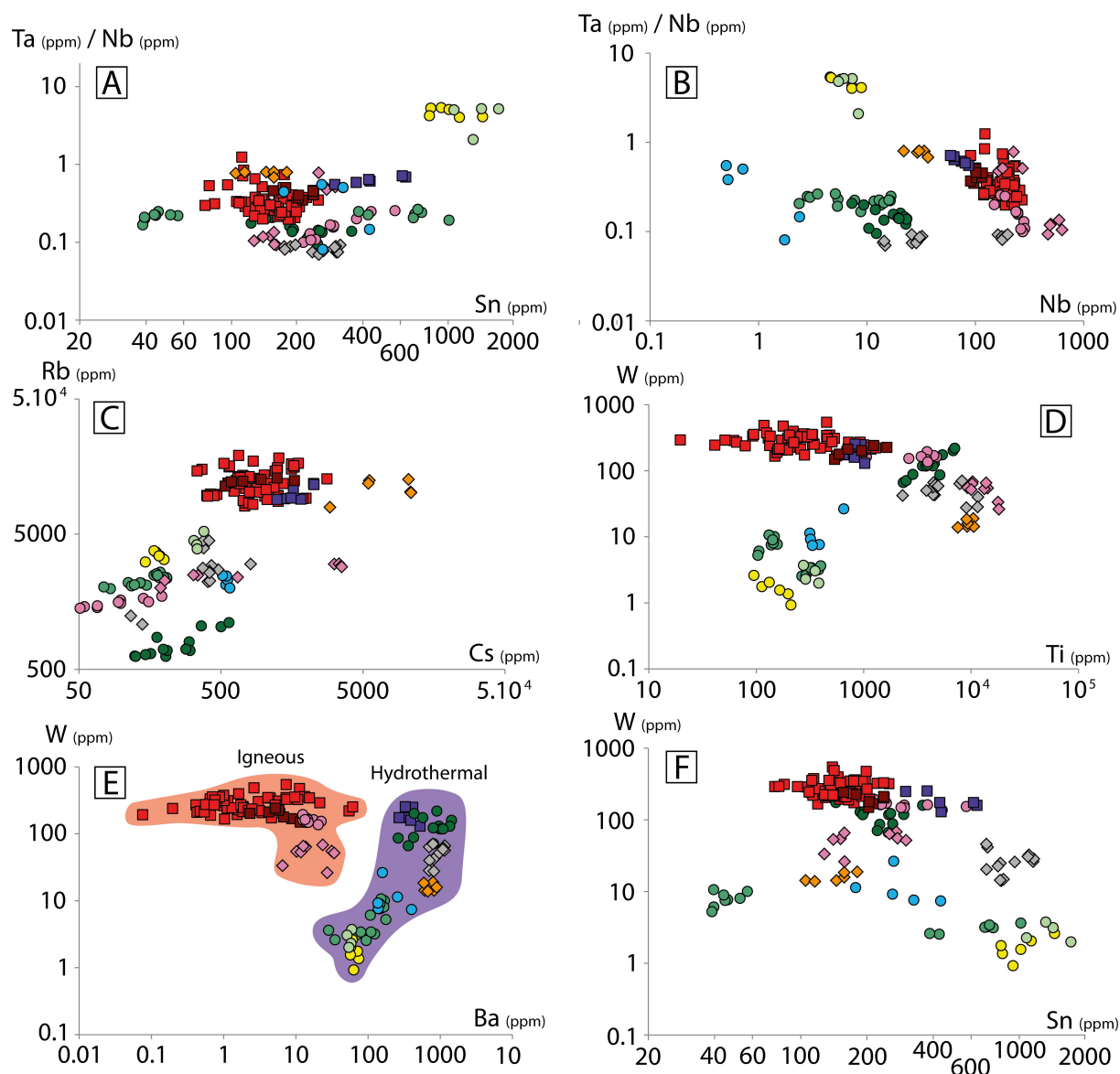


Fig. 6. Univariate diagram depicting the range of Li concentrations (measured by LA-ICP-MS) for the different mica petrographic groups. Symbols as in Fig. 5.

253 Regarding other elements, below are described only the most remarkable results, and the reader is
254 referred to Table 2 for average concentrations of trace elements for each petrographic group, and to Table
255 A2 (Appendices) for the full dataset. Tin is relatively enriched (tens to ca one thousand ppm) in all mica
256 groups compared to other micas from W-Sn deposits (Codeço et al., 2021), and is strongly enriched in
257 muscovite (9) and (12) from the Beauvoir greisen (ca. 1000 and 1400 ppm in average for yellow and green
258 muscovite, respectively; Fig. 7.A). Such concentrations (more than 1000 ppm) have only been recorded in
259 muscovite from the Cinovec granite during late greisen alteration, called “muscovitization” by Breiter et al.
260 (2019). The Ta/Nb ratio is highest (> 4) for the (9) and (12) muscovite groups from the Beauvoir greisen and
261 lowest (< 0.1) for zinnwaldite from the topaz veins (2). Ta and Nb are globally enriched ($>$ to 20 ppm and
262 100 ppm, respectively) in igneous micas (Fig. 7.B). Rubidium is strongly enriched in lepidolite and
263 zinnwaldite from quartz vein halos (1), with concentrations mostly superior to 10 000 ppm (Fig. 7.C).
264 Cesium content is also particularly high (ca. 7000 ppm in average) in these micas (1), and low in the
265 muscovite groups and in zinnwaldite from topaz veins (< 600 ppm; Fig. 7.C). Tungsten is remarkably
266 enriched, with more than 100 ppm in average in lepidolite (4), (5), (6), in muscovite from the Colettes
267 granite (8) and in green muscovite (7) from the Suchot area (Fig. 7.D). In some samples of lepidolite from
268 the Beauvoir granite (5), W concentrations can reach 300 ppm, which, to our knowledge, is the highest W
269 content in mica ever reported in the literature. Titanium contents are low in lepidolite and muscovite
270 groups, except for muscovite in the Colettes granite (8) and green muscovite of the Suchot area (7), which
271 display concentrations of ca. 4000 ppm, while zinnwaldite are the most Ti-enriched mica species, with ca.
272 10,000 ppm in all the three groups (Fig. 7.D). We noticed that all magmatic or hydrothermal micas with Ti
273 concentrations higher than ca. 1000 ppm are associated with a Ti-phase, either rutile or ilmenite.
274 Conversely, micas with subordinate Ti contents are not coeval with Ti-bearing minerals. Barium is enriched
275 in hydrothermal micas (from ca. 50 to 900 ppm in average) and depleted in magmatic micas (< 20 ppm in
276 average; Fig. 7.E). Magmatic micas are also strongly depleted in Sr (< 0.3 ppm in average), which can exceed
277 10 ppm in zinnwaldite from topaz veins (2) and in green muscovite from the Suchot area (7). Germanium
278 varies only slightly, averaging 6 ppm in zinnwaldite from the Colettes granite (3), and 1.5 ppm in greisen

279 muscovite from the Beauvoir granite (9), (12). A noteworthy observation is that, with few exceptions, in
 280 the Echassières district, W-rich micas are relatively depleted in Sn, and vice versa (Fig. 7.F).



281 *Fig. 7. Binary diagrams depicting LA-ICP-MS data for selected element pairs, for the different petrographic*
 282 *groups. The red and violet shapes in E correspond respectively to igneous and hydrothermal mica groups.*
 283 *Symbols as in Fig. 5.*

284 A remarkable feature emerging from these results is that groups that contain micas originating from
 285 different samples, such as (2), (5) and (10) have homogeneous signatures (cf. Table 1), with the exception
 286 of a few elements such as Sn for group (10) and Nb for group (2) that vary strongly across samples. Another
 287 notable feature is the homogeneous signature of group (2) mica, given that this group contains zinnwaldite
 288 from both core and alteration halos of topaz veins, located above and below the Beauvoir granite.
 289 Signatures are also homogeneous within single-sample groups, excepted few elements which display
 290 significant variations, such as Cs and Sn in zinnwaldite from the Colettes granite (3).

group	(1) n=6		(2) n=13		(3) n=9		(4) n=6		(5) n=58		(6) n=7	
	mean	S.D.	mean	S.D.	mean	S.D.	mean	S.D.	mean	S.D.	mean	S.D.
Li	9763	564	7209	2590	6204	616	14782	1873	24619	3420	23104	1320
Na	1223	171	2201	760	2395	134	1458	154	2647	415	2191	246
Ti	9391	1022	6305	2895	13958	0.34	827	129	351	305	174	23
Ge	4.13	0.79	4.05	1.16	6.13	1.91	3.21	1.21	3.44	1.16	3.39	0.74
Rb	10884	1672	2798	1000	2516	359	9756	974	12333	2810	12263	503
Sr	1.6	0.5	13.1	2.5	0.1	0.0	Bdl	-	Bdl	-	Bdl	-
Nb	30.9	4.7	26.2	9.4	419.1	169.6	71.3	10.0	177.3	40.1	103.3	8.6
Sn	143	26	246	57	193	64	467	122	152	37	205	28
Cs	7704	3174	397	158	1427	1500	1653	326	986	489	893	388
Ba	732	113	899	166	18	9	376	82	5	10	6	2
Ta	23.8	3.2	2.2	0.9	81.8	38.8	44.5	2.6	63.2	26.8	42.9	6.7
W	15.9	2.0	49.9	13.4	48.2	15	189.4	45.5	286.3	75.8	183.5	37.12

group	(7) n=14		(8) n=10		(9) n=6		(10) n=16		(11) n=5		(12) n=4	
	mean	S.D.	mean	S.D.	mean	S.D.	mean	S.D.	mean	S.D.	mean	S.D.
Li	734	203	2407	277	27	4	871	481	2451	795	1578	1021
Na	6385	484	6211	833	564	137	364	107	637	54	325	70
Ti	4364	1364	3582	1010	153	43	233	104	402	128	325	45
Ge	3.68	0.35	3.99	1.14	1.49	0.48	1.97	0.73	4.27	0.49	2.38	0.52
Rb	771	166	1548	102	3445	228	2277	193	2263	185	4435	500
Sr	18.8	6.8	0.2	0.1	0.3	0.1	0.4	0.3	6.7	4.8	0.2	0.1
Nb	16.7	5.4	234.0	45.5	6.4	1.5	9.1	5.1	1.2	0.8	6.8	1.1
Sn	223	61	319	107	1031	214	362	343	293	85	1385	233
Cs	261	132	102.0	44	176	16	152	41	548	21	344	21
Ba	817	352	15	3	64	6	113	46	219	101	54	3
Ta	2.4	0.7	36.1	6.4	29.1	3.9	2.0	1.2	0.3	0.1	28.3	7.6
W	128.3	44.1	159.2	14.8	1.7	0.5	5.6	2.7	12.4	7.1	2.8	0.7

Table 2: Mean and standard deviation of element concentrations for the petrographic mica groups, analyzed by LA-ICP-MS.

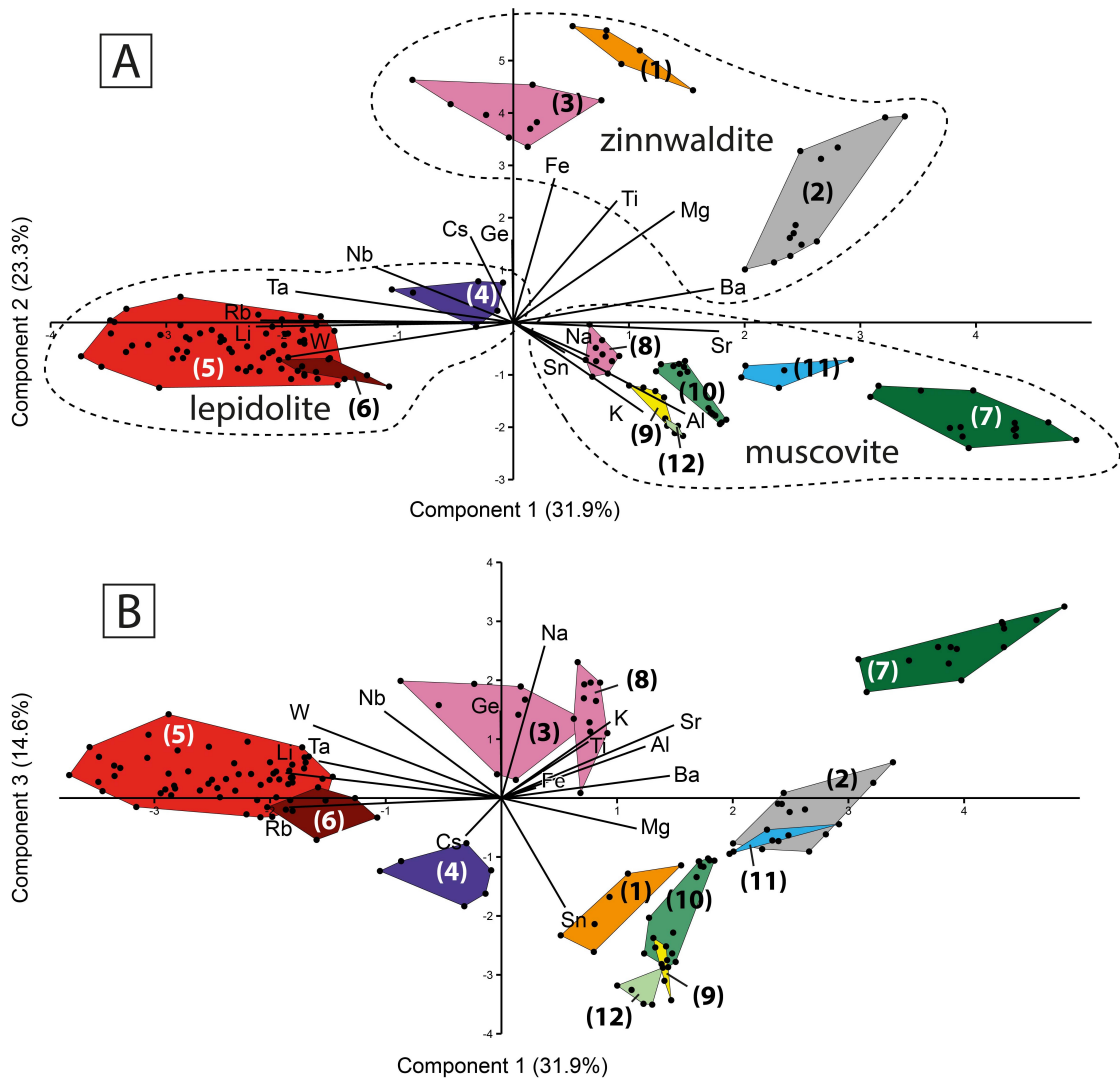
292 **4.3. Chemical signatures and identification of mica generations**

293 In order to properly examine the large trace-element dataset for all mica groups, we implemented a panel
294 of multivariate statistical analyses, namely, Principal Component Analysis (PCA) ordination, Hierarchical
295 Cluster Analysis (HCA), and Permutational Analysis of Variance (PERMANOVA). All these multivariate
296 analyses were monitored with the software Past 3.17 (Hammer et al., 2001). The few data that fell below
297 the limits of detection (available in table A2, Appendices) were considered as equal to the given limit
298 (almost exclusively Sr measurement for igneous micas). Statistical distance of similarity were calculated
299 with the Bray Curtis similarity index (Bray and Curtis, 1957). A similarity distance equal to 1 means that the
300 two analyses considered are strictly equal for all elements measured, and as the distance decreases, the
301 difference increases. This protocol was also used to characterized trace-element signatures in quartz from
302 the Echassières and other Sb districts in France (Monnier et al., 2018, 2021b). HCA and PCA are statistical
303 tools that are becoming of common use in the geochemical literature (cf., Lin et al., 2014 and Codeço et al.,
304 2021, for other applications in the field of economic geology). Anderson (2001) provides a detailed
305 description of the PERMANOVA test. See Supplementary Material for a brief explanation of process and
306 objectives for each specific multivariate analysis.

307 *4.3.1. Principal Component Analysis.*

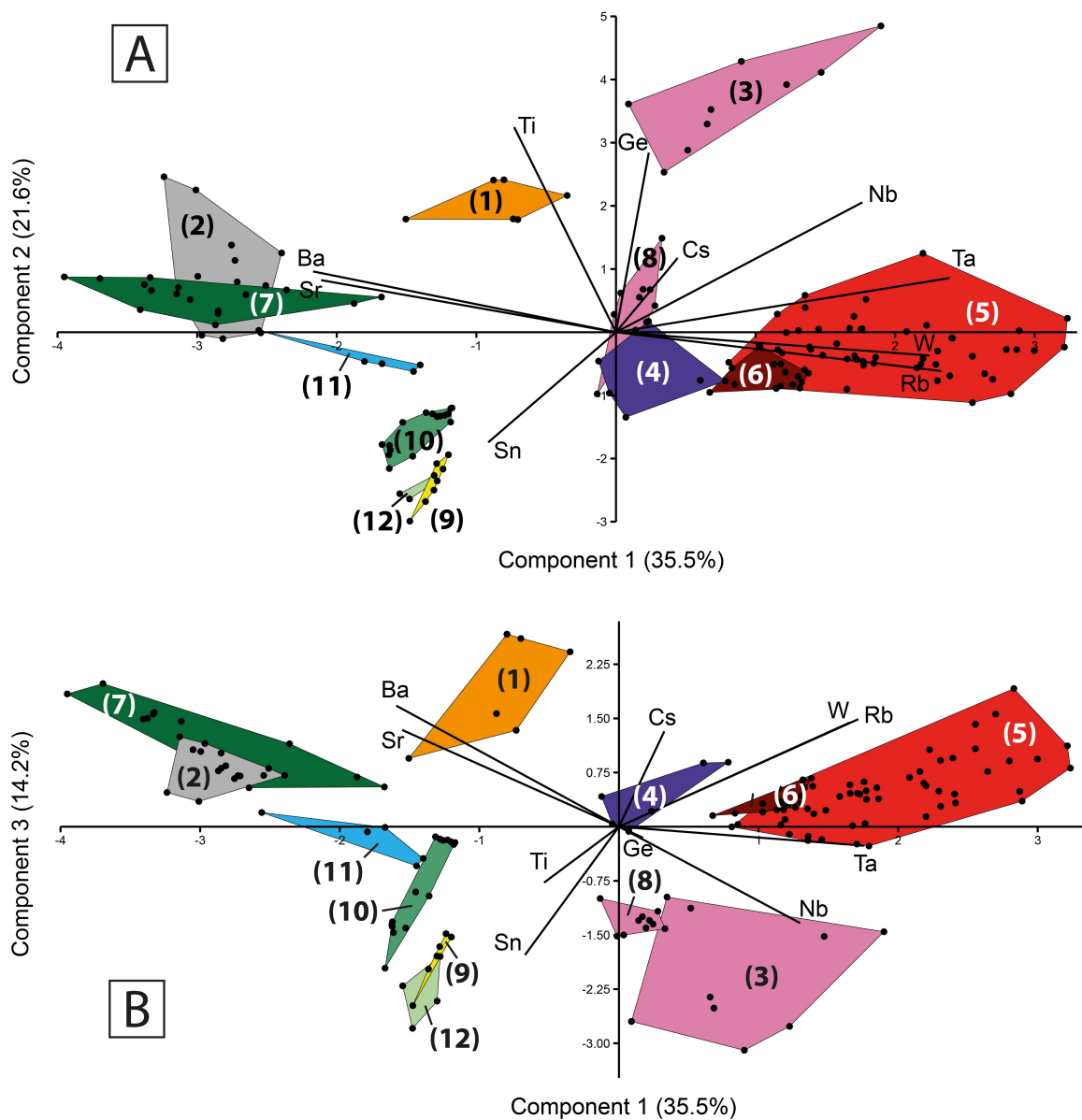
308 The first PCA ordination (Fig. 8) was calculated using all elements analyzed (i.e., major and trace elements).
309 As expected, the first and the second components (which best represent the entire data) discriminate well
310 the different types of mica (lepidolite, zinnwaldite and muscovite groups), notably by their enrichment in
311 Na, Al and K for the muscovite groups and in Fe and Mg for the zinnwaldite groups (Fig. 8.A). Only the third
312 component seems to group different mica species, in particular muscovite (3) and zinnwaldite (8) from the
313 Colettes granite on one hand, and the majority of hydrothermal micas on the other (Fig. 8.B). To obviate
314 artifacts owing to a specific mineral's nature, we repeated the PCA excluding major elements (K, Al, Na, Fe,
315 Mg) and Li, which are variably enriched depending on the type of mica (Fig. 6). The second PCA ordination
316 (Fig. 9) forms several clusters that are characterized by the same positive and negative correlations. Ba and
317 Sr are particularly enriched in zinnwaldite of topaz veins (2) and in green muscovite in Suchot vein halos (7);

318 they are depleted in igneous micas (3), (5), (6), (8) and hydrothermal lepidolite from the La Bosse stockwork
 319 (4) for the principal component. Niobium, Ta, W and Rb enrichment characterizes the group of igneous
 320 micas. Germanium is enriched in Colettes Zinnwaldite (3), and Sn is particularly enriched in greisen
 321 muscovite of the Beauvoir granite. The remaining muscovite groups (10) and (11) are not so much
 322 characterized by an Sn enrichment as by the Ba-Sr couple. On a first order, Ti is more enriched in
 323 zinnwaldite (1); (2); (3), in agreement with numerous studies which report the presence of up to 7 w.% of
 324 TiO₂ in biotite from various origin (e.g., Thompson et al., 1997; Cesare et al., 2008; Wu and Chen, 2015.
 325 Nevertheless, some muscovite, (7) and (8), displays Ti concentrations (ca. 4000 ppm) only slightly inferior
 326 then those in zinnwaldite (2), roughly equals to 6000 ppm.



327 Fig. 8. Principal component analysis (PCA) diagrams projecting components 1 and 2 [A] and 1 and 3 [B] for all
 328 trace and major elements analyzed by LA-ICP-MS. The different mica species are identified and clustered in [A]
 329 by the dashed lines. Shape colors and numbers refer to mica groups (cf. Fig. 5).

330 In brief, PCA ordination of trace elements discriminates three distinct clusters, mainly by the first
331 component: i) igneous micas and hydrothermal lepidolite (3), (4), (5), (6) (8), ii) zinnwaldite in quartz vein
332 halos (1), iii) and a last cluster including the other hydrothermal micas (2), (7), (9), (10), (11), (12). We can
333 subdivide the first and the third clusters using the second and third component: micas from the Colettes
334 granite (3) and (8) are distinguished from the lepidolite groups (4), (5), (6), and the (2) and (7) mica groups
335 are distinguished from muscovite (9), (10), (11), (12), particularly for greisen muscovite in the Beauvoir
336 granite. Remarkable elemental associations include the Ba-Sr and W-Rb couples, which are strongly
337 correlated in micas from the Echassières district. Conversely, Cs and Rb are only slightly correlated,
338 although these elements occupy the same site in the lepidolite micas series.

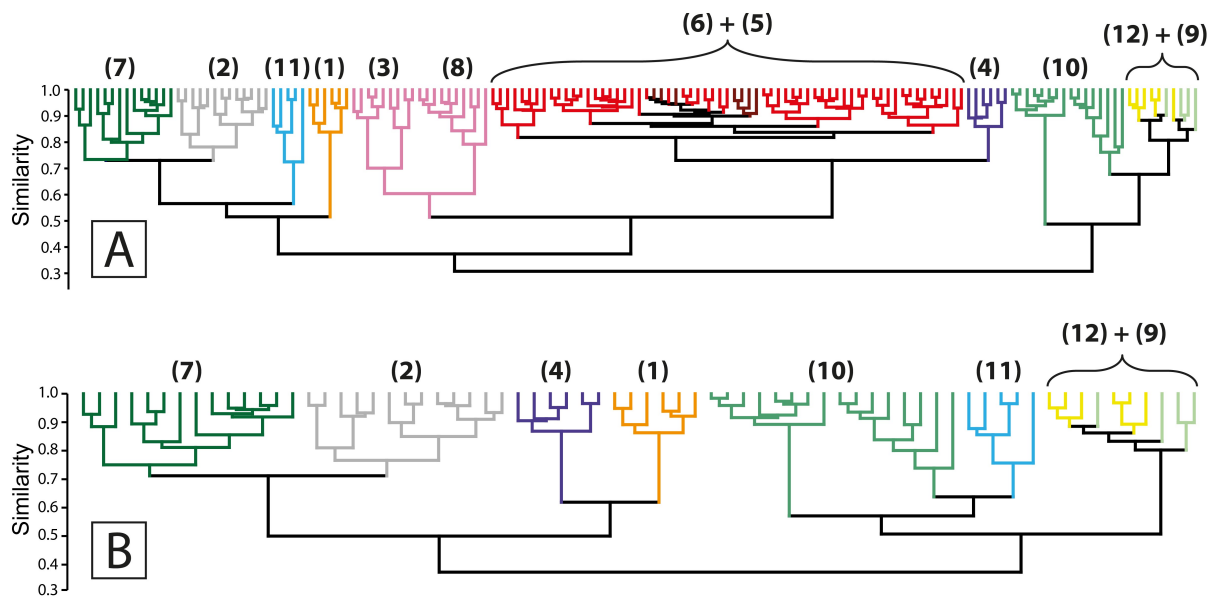


339 Fig. 9. PCA diagrams projecting components 1 and 2 [A] and 1 and 3 [B] for selected trace elements (Ti, Ge, Rb,
 340 Sr, Nb, Sn, Cs, Ba, Ta, W) analyzed by LA-ICP-MS. Shape colors and numbers refer to mica groups (cf. Fig. 5).

341 4.3.2. Hierarchical Cluster Analysis.

342 At a first glance, HCA for all groups (Fig. 10.A) confirms that each petrographic mica group (with the
 343 exception of the (6), (5) and (12), (9) couples, discussed below) displays a unique and homogeneous trace-
 344 element signature, as each analysis assembles mostly the same clusters that already belonged to the
 345 different petrographic groups. Similarly, micas groups that include several samples, i.e. (2), (5), (10) are
 346 characterized by unique, coherent signatures, as previously noted. This is not entirely the case for the two
 347 samples of stockwork green muscovite (10), which are clearly separated but occupy nearby clusters in the
 348 dendrogram, suggesting limited differences in their chemical signatures. Some petrographic groups, (5) and
 349 (6) on one hand and (9) and (12) on the other, display very similar chemistry and form indiscernible

350 clusters. These results imply that chemical signatures of lepidolite from dykes (6) and from the main body
351 of the Beauvoir granite (5) are almost undistinguishable, as for green (12) and pale-yellow muscovite (9) in
352 the Beauvoir granite, in accordance with obvious genetic links which connect these two groups. We can
353 highlight the resemblance of some mica groups by clustering; for example, groups (2) and (7) form a cluster
354 with a similarity of ca. 0.7, and mica groups from the Colettes granite (3) and (8) have a similarity of ca. 0.6.
355 Hydrothermal lepidolite (4) displays a similarity of ca. 0.7 with the igneous micas from Beauvoir granite (5)
356 and (6), in the same way that one sample of greisen muscovite from stockwork (10) and the Beauvoir
357 greisen muscovites (12) and (9). The second sample of muscovite (10) is clearly more distant, sharing a
358 similarity of less than 0.5 with the latter cluster. Two other mica groups that display relatively low similarity
359 values with other groups are zinnwaldite in quartz vein halos (1) and greisen blue muscovite at Suchot (11).
360 Zinnwaldite in quartz vein halos (1) is not closely associated with other groups. The group of blue muscovite
361 (11) is close to the cluster (2) and (7), as it has been previously highlighted by similar elemental correlation,
362 notably Ba and Sr, in PCA (Fig. 9). However, when we consider only the hydrothermal micas group (Fig.
363 10.B) in a novel HCA, we observe a major difference which consists of the location of the blue muscovite
364 groups, closely associated (similarity of ca. 0.6) to one sample of greisen muscovite from stockwork (10).
365 This outcome implies that blue muscovite (11) and Beauvoir greisen muscovite (12) and (9) are relatively
366 analogous to green muscovite from the stockwork (10), while mica groups (11) on one hand and (12) and
367 (9) on the other, are more different from each other. In other terms, green muscovite from the stockwork
368 (10) seems to have intermediate composition between two chemical poles consisting of blue muscovite
369 (11) and by Beauvoir greisen muscovite groups (12) and (9). Also, due to the absence of igneous lepidolite
370 in the Beauvoir granite (6), (5), hydrothermal lepidolite disposition differs from the first HCA and forms,
371 with zinnwaldite in quartz vein halos (1), a cluster of similarity ca. equal to 0.6.



372 Fig. 10. [A]: Dendrogram resulting from hierarchical cluster analysis (HCA) of the same panel of trace element
 373 data as in Fig. 10 (i.e., Ti, Ge, Rb, Sr, Nb, Sn, Cs, Ba, Ta, W), considering all the petrographic mica groups. [B]:
 374 Dendrogram resulting from HCA of the above elements, excluding the (3), (5), (6) and (8) petrographic groups
 375 (igneous micas). Shape colors and numbers refer to mica groups (cf. Fig. 5).

376 In summary, the most significant results of HCA are similar to those obtained by PCA, i.e, highlighting a
 377 strong affinity between mica groups (7) and (2), as well as for hydrothermal lepidolite (4) and igneous
 378 lepidolite from the Beauvoir granite (5) and (6), and a clear similarity between zinnwaldite and muscovite
 379 from the Colettes granite. Muscovite (9), (10), (11) and (12) also display similar chemistry, whereby the
 380 greatest difference was found between blue muscovite from Suchot veins (11) and greisen muscovite from
 381 the Beauvoir granite (9) and (12). Finally, zinnwaldite in quartz vein halos (1) shares limited similarities with
 382 the hydrothermal lepidolite (4).

383 4.3.3. Permutational Analysis of Variance.

384

group	(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)	(9)	(10)	(11)	(12)
(1)	0.16	0.01	0.03	0.16	0.01	0.02	0.01	0.02	0.09	0.01	0.13	0.4
(2)	0.01	0.16	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.03
(3)	0.03	0.01	0.16	0.02	0.01	0.01	0.01	0.01	0.01	0.01	0.03	0.13
(4)	0.16	0.01	0.02	0.16	0.03	0.05	0.01	0.03	0.15	0.01	0.1	0.33
(5)	0.01	0.01	0.01	0.03	1	1	0.01	0.01	0.01	0.01	0.01	0.01
(6)	0.02	0.01	0.01	0.05	1	0.16	0.01	0.01	0.04	0.01	0.06	0.25
(7)	0.01	0.01	0.01	0.01	0.01	0.01	0.16	0.01	0.01	0.01	0.01	0.02
(8)	0.02	0.01	0.01	0.03	0.01	0.01	0.01	0.16	0.02	0.01	0.03	0.07
(9)	0.09	0.01	0.01	0.15	0.01	0.04	0.01	0.02	0.16	0.01	0.09	0.32
(10)	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.16	0.55	0.01
(11)	0.13	0.01	0.03	0.1	0.01	0.06	0.01	0.03	0.09	0.55	0.16	0.48
(12)	0.4	0.03	0.13	0.33	0.01	0.25	0.02	0.07	0.32	0.01	0.48	0.16

≤0.05 >0.05

386 *Table 3: Permanova p-values between all groups considering the same panel of trace element data as in Fig. 9*
 387 *and Fig. 10 (i.e., Ti, Ge, Rb, Sr, Nb, Sn, Cs, Ba, Ta, W). The f-values are available in Table 4.*

gr.	(1)	(2)	(3)	(4)	(5)	(6)	(7)	(8)	(9)	(10)	(11)	(12)
(1)		30.74	29.13	61.23	72.71	98.42	84.52	119.1	232.9	217.8	185.8	126.3
(2)	30.74		18.09	52.75	260.8	87.66	13.07	19	48.25	95.63	32.39	30.59
(3)	29.13	18.09		110.7	271	167.3	64.65	62.27	120.9	178.9	90.54	77.6
(4)	61.23	52.75	110.7		9.17	36.57	125.4	150	246	152.6	201	109.7
(5)	72.71	260.8	271	9.17		1.09	473.3	313.9	132.5	391.1	131.7	63.37
(6)	98.42	87.66	167.3	36.57	1.09		203.4	240.4	530.3	220	432.7	261.6
(7)	84.52	13.07	64.65	125.4	473.3	203.4		17.66	133.4	196.3	74.83	86.96
(8)	119.1	19	62.27	150	313.9	240.4	17.66		118.4	118	69.53	84.41
(9)	232.9	48.25	120.9	246	132.5	530.3	133.4	118.4		21.86	70.71	14.23
(10)	217.8	95.63	178.9	152.6	391.1	220	196.3	118	21.86		10.11	32.86
(11)	185.8	32.39	90.54	201	131.7	432.7	74.83	69.53	70.71	10.11		61.21
(12)	126.3	30.59	77.6	109.7	63.37	261.6	86.96	84.41	14.23	32.86	61.21	

388

≤ 2 ≤ 2; > 15 ≤ 15; > 50 ≤ 50; > 100 ≤ 100

389 *Table 4: Permanova f-values between all groups considering the same trace elements as in Fig. 9 and Fig. 10*
 390 *(i.e., Ti, Ge, Rb, Sr, Nb, Sn, Cs, Ba, Ta, W). The p-values are available in Table 3.*

391 We are not going to comment on single Permanova results (Tables 3, 4), but we will highlight remarkable
 392 results with respect to the previous multivariate analyses. Note that groups with the smallest population,
 393 e.g. (1), (4), (9), (11), (12) are subject to higher p-value data, due to possible high variations in
 394 concentration coupled to possible low replicates. F-values are less affected by such bias. Zinnwaldite in
 395 quartz vein halos (1) exhibits marked differences from most of the other groups, even if their chemistry is
 396 closer to that of the other zinnwaldite groups (f-value ≈ 30). Zinnwaldite (2) is also well distinguished by p-
 397 value from the other groups, although the difference is less substantial (f-value = 13) for the green
 398 muscovite (7). Much as the previous zinnwaldite groups, the Colettes zinnwaldite (3) differs clearly from
 399 the other mica groups. The least difference is with zinnwaldite (2), as revealed by the smallest f-value
 400 between the two groups. This contrasts with the PCA and HCA results, which had suggested a stronger
 401 similarity of Colettes zinnwaldite (3) with the Colettes muscovite (8). The chemistry of hydrothermal
 402 lepidolite (4) is close to that of igneous lepidolite (5) and (6). Igneous lepidolite from the Beauvoir main
 403 body (5) and dyke (6) displays an indistinguishable trace element signature (p-value = 1; f-value = 1). Green
 404 muscovite from the Suchot area shares the lowest f-value with zinnwaldite from topaz veins (2). Colettes
 405 muscovite (8) chemistry is close to that of the green muscovite from the Suchot area (7) as well as to that of
 406 zinnwaldite from topaz vein (2). As for the Colettes zinnwaldite, results of PERMANOVA on Colettes

407 muscovite disagree with the previous multivariate analyses. Micas originating from Beauvoir greisen, i.e., (9)
408 and (12), display a similar chemistry, (p -value = 0.32 ; f -value = 14) and are both relatively close (f -value of
409 22 and 33) to green muscovite from La Bosse stockwork (10). Finally, blue muscovite from quartz halos in
410 Suchot veins (11) exhibits a strong chemical affinity (p -value = 0.55; f -value = 10) with the green muscovite
411 from La Bosse stockwork (10).

412 To resume, PERMANOVA confirms most of the relationship highlighted by the previous statistical analyses,
413 i.e., i) an affinity between igneous lepidolite from Beauvoir (5) and (6), and hydrothermal lepidolite (4) from
414 the La Bosse stockwork, ii) an affinity between zinnwaldite from topaz veins (2) and green muscovite from
415 Suchot area (7), iii) a similarity between muscovite (9), (10), (11) and (12), iv) as well as the singular
416 signature of zinnwaldite (1). Nonetheless, the similarity between micas from Colettes granite suggested by
417 the PCA and HCA is not verified by PERMANOVA. We can potentially attribute the large f -value of ca. 62 for
418 Colettes' zinnwaldite and muscovite as the result of chemical differences induced by different partition
419 coefficients between the muscovite and Fe-Mg rich species (Stepanov et al., 2014). Indeed, PERMANOVA is
420 more affected by absolute concentration variations while PCA relationships are best constrained by the
421 correlation, positive or negative, of trace-element for the different groups, which remains essentially the
422 same for muscovite and zinnwaldite from Colettes.

423 4.3.4. *Key results of multivariate analyses.*

424 Considering all results from the different multivariate analyses, we obtain several group connections based
425 on trace-element similarities: i) igneous mica from the Colettes granite (3) and (8); ii) igneous and
426 hydrothermal lepidolite (4), (5) and (6); iii) zinnwaldite from quartz vein halos (1) that are well distinguished
427 from the other groups; iv) green muscovite from the Suchot area (7) and zinnwaldite from topaz veins (2);
428 v) all remaining muscovite groups (9), (10), (11), (12). Nevertheless, the link between the chemical
429 signatures of the blue muscovite from the Suchot vein (11) and greisen muscovite from the Beauvoir
430 granite (9) and (12) would have been hard to pinpoint without the group of green muscovite from the
431 stockwork (10), which exhibits intermediate chemistry.

432 5. Discussion

433 **5.1. *Micas signatures***

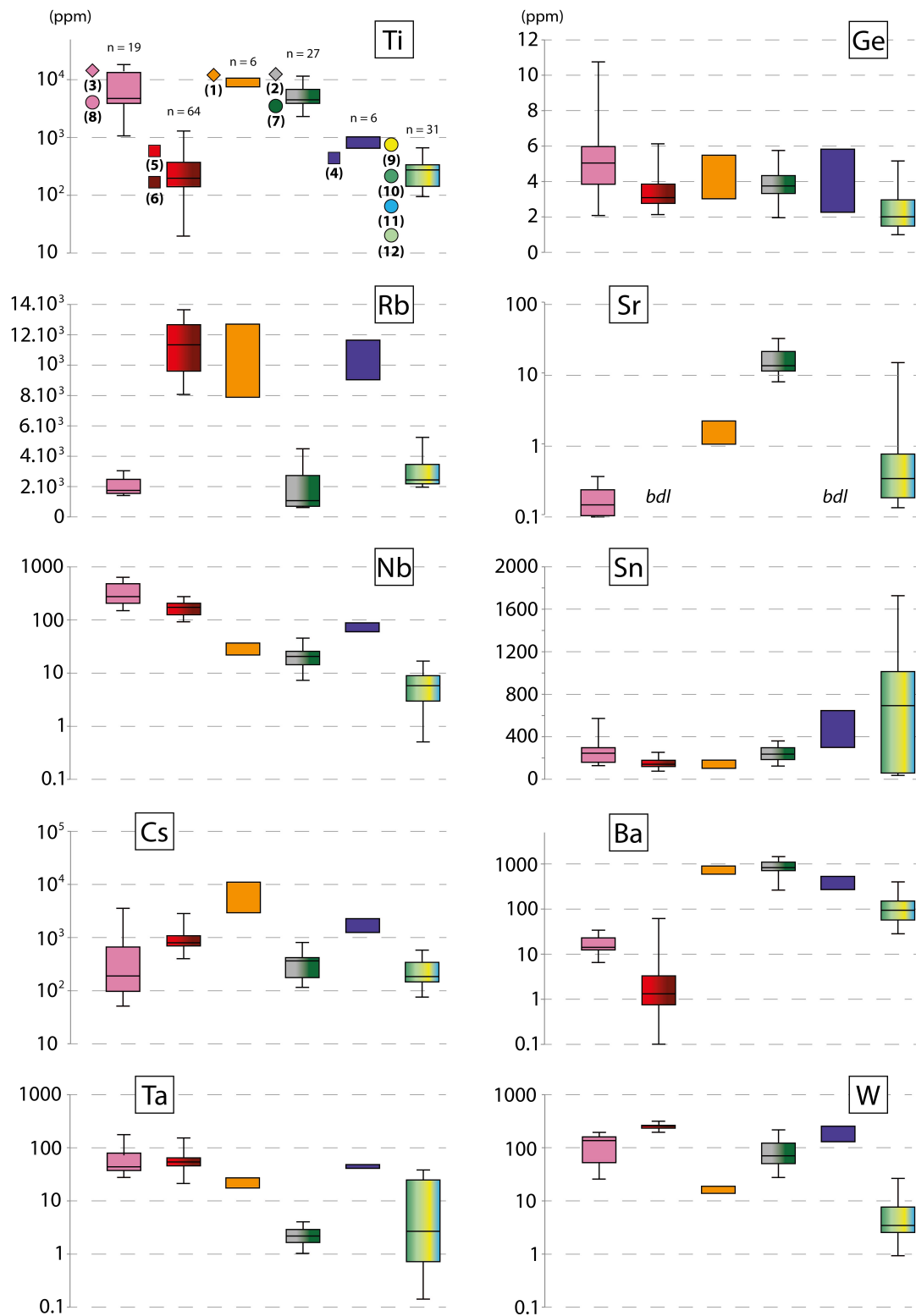
434 *5.1.1. Matching geochronological constrains to W-Sn mineralizations.*

435 In the Echassières district, the La Bosse stockwork event preceded the Barrovian metamorphism dated at
436 350-360 Ma (Do Couto et al., 2016, Monnier et al., 2021a). Zinnwaldite (1) from quartz vein halos,
437 considering that quartz vein emplacement and selvage alteration were coeval, may originate from the older
438 hydrothermal event recognized in the Echassières district. This is consistent with the particular signature of
439 zinnwaldite (1), which contrasts clearly with that of the other micas. The topazification episode affected the
440 La Bosse stockwork at ca. 330-335 Ma (Monnier et al., 2021a). Strong textural evidence links zinnwaldite (2)
441 to this F-rich alteration. Moreover, minerals dated at ca. 330 Ma in the Suchot vein area occur as inclusions
442 in the green muscovite group (7), which displays chemical similarities with the zinnwaldite (2). Based on
443 these chemical and chronological evidences, we interpret zinnwaldite (2) and muscovite (7) to originate
444 from the same mineralizing fluid that was responsible for the topazification of the La Bosse stockwork.
445 Despite green muscovite (7) being located in the alteration halo of the Suchot vein, our data and previous
446 ages (Monnier et al., 2021a) show that this mica group should be considered as an inherited generation
447 with respect to the formation of the Suchot vein, i.e., it was present before vein formation. Consequently,
448 the topazification that we observed at the La Bosse stockwork seems to have had a regional influence (or at
449 least was part of a more widespread episode). The Colettes and Beauvoir granites crystallized
450 contemporaneously, at ca. 310 Ma. Expectedly, muscovite and zinnwaldite from the Colettes granite share
451 a roughly homogeneous chemical signature. Yet, lepidolite from the different facies of the Beauvoir granite
452 (5) and (6), exhibits a very consistent signature that differs clearly from that of Colettes' micas. Remarkably,
453 hydrothermal lepidolite (4) associated with the second generation of topaz in the stockwork, and igneous
454 lepidolite (5) and (6), show analogous signatures, implying a roughly similar chemistry for the Beauvoir
455 granitic melt and exsolved fluid during late crystallization stage. The late Sn-W mineralizing episode consists
456 of OH-rich greisen alteration occurring during cooling of the Beauvoir granite (ca. 310 Ma). Hydrothermal
457 muscovite in the different greisen facies of the Beauvoir granite (12) and (9) are chemically related to green
458 muscovite (10) that replaced minerals in the topaz vein, which formed at ca. 330 Ma. Blue muscovite of the

459 greisen-related Suchot vein halo (11), which exhibits monazite inclusions dated at 307 ± 3 Ma, shows
460 chemical similarities with green muscovite (10) of the stockwork. Based on the strong trace-element affinity
461 with other greisen muscovite (constrained at ca. 310 Ma), and on the relative timing of green muscovite
462 (10) in the stockwork (postdates the 330 Ma topazification episode, see above), we can reasonably
463 associate this mica group to the Beauvoir greisen alteration. To conclude, all genetic links highlighted by the
464 trace element chemistry are supported by geochronological data, validating the multivariate analytical
465 protocol.

466 *5.1.2. Typical signature enrichment for each mica-forming episode.*

467 Combining petrological observations, multivariate analysis and chronological data, we have clustered the
468 different mica groups as a function of the different mineralizing episodes, in order to characterize each
469 magmatic and hydrothermal episode by their specific trace-element concentration (Fig. 11). Micas from the
470 Colettes granite are characterized by high Ti, Nb, Ta, Ge contents, are slightly enriched in W, and have a
471 medium Sn and low Rb, Sr, Cs and Ba contents. Igneous micas from the Beauvoir magmatic episode are
472 characterized by high Rb, Nb, Ta, W contents, intermediate Cs and Ge contents, and low Ti, Sr (bdl), Sn, and
473 Ba. Zinnwaldite (1) is considered as originating from the La Bosse stockwork formation. Although constrains
474 on this episode are weak due to the several metamorphic and magmatic/hydrothermal episodes that
475 affected the area, zinnwaldite (1) records a singular signature, suggesting a specific hydrothermal event.
476 Notably, Ti, Cs, Rb and Ba are relatively enriched, Nb, Ta, Ge and Sr are present in intermediate
477 proportions, while Sn and W are relatively depleted. The topazification episode led to crystallization of
478 micas that incorporated high Ti, Sr, Ba, intermediate Nb, Ge, Sn W, and low Rb, Cs, Ta contents. Lepidolite
479 (4) from the orthomagmatic Beauvoir episode is characterized by high Rb, Cs, Ta, Sn, Ba and W contents,
480 intermediate Ti, Nb and Ge, and low Sr (bdl) contents. Micas precipitating from the Beauvoir greisen fluid
481 display high Sn, intermediate Rb, Sr, Ba, and low Ti, Nb, Cs, Ta, Ge, W contents, although some of these
482 elements, i.e., Ta, Sr, Sn, vary widely.



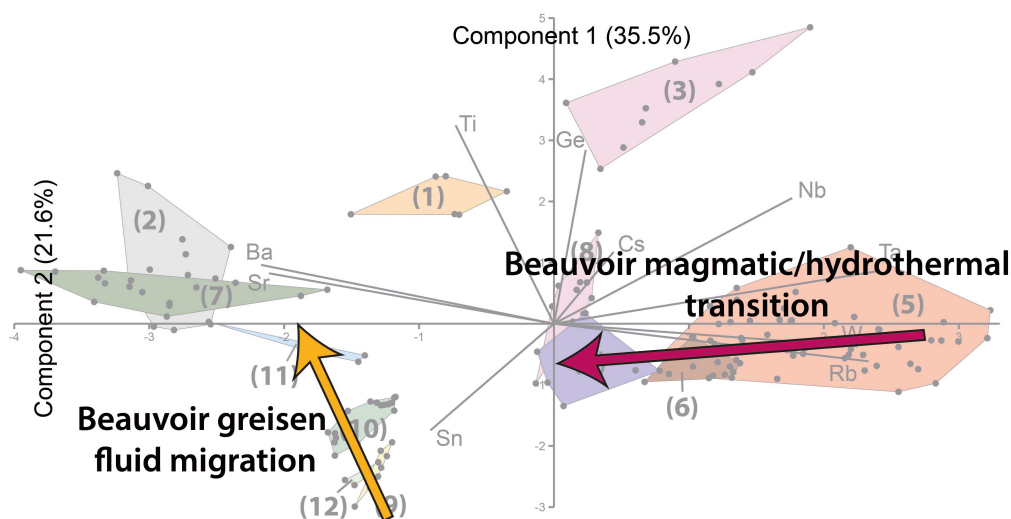
483 Fig. 11. Boxplot diagrams depicting trace-element data for the petrographic mica groups (cf., Fig. 5 for legend),
 484 clustered according to igneous and hydrothermal episodes recognized in the Echassières district by a
 485 combination of petrological, geochronological and trace-element data. The boxes for zinnwaldite (1, orange)
 486 and lepidolite (4, violet) are formed only by minimum and maximum values due to a low number of analyses.

487 5.1.3. Global implication.

488 It was expected that incorporation of trace elements in micas would be dictated mostly by the nature of
 489 the mica species (i.e., major elements) and the corresponding trace element partition coefficients (e.g.,

490 Stepanov et al., 2014). However, the results of the multivariate analysis reveal groups that are
491 heterogeneous in terms of major compositions, but share a similar origin/source. This implies that trace
492 element incorporation in micas is in fact dictated mainly by the chemistry of the melt or fluid from which
493 they precipitate. Different mica species can therefore be compared and their trace-element signatures used
494 as a powerful tool to unravel intricate paragenesis, ranging from magmatic to hydrothermal conditions, of
495 which the polyphased ore system investigated herein is one example. With sufficient samples, it is even
496 possible to trace the evolution of a fluid that underwent cooling and mixing, leading to significant chemical
497 changes, by recognizing different snapshots of its passage in hydrothermal mineral assemblages that could
498 otherwise appear petrographically unrelated (e.g., evolution path of Beauvoir greisen fluid; yellow arrow on
499 Fig. 12). Moreover, orthomagmatic micas can display a roughly similar chemistry to that of primary igneous
500 micas, where only very few elements vary (e.g. Beauvoir magmatic/hydrothermal path, red arrow on Fig.
501 12). It can be deduced that the trace chemistry of an orthomagmatic fluid remains largely dominated by the
502 chemical nature of the melt that sourced it, providing an opportunity to use the trace chemistry of
503 alteration mica as a proxy for strongly evolved rare-metal granite. It should be kept in mind that these
504 concentrations can vary within the order of magnitude scale, given that partition coefficient between fluid
505 and melt as well as between various micas and the crystallizing medium are a function of different
506 parameters. Nevertheless, the chemical variation of magmas and fluids from radically different sources will
507 vary to a much wider scale (several orders of magnitude for key elements) thus allowing definition of
508 specific signatures. Similarly to major element chemistry, petrographic features, such as mineral texture
509 and color, can vary widely during a single mineralizing episode, remaining poor indicators for deciphering
510 polyphased mineralization. Moreover, considering that Ar-Ar dating of mica is often partially, if not entirely,
511 reset in these complex systems and does not reflect ore deposition age (e.g., Zhang et al., 2017;
512 Tichomirowa et al., 2019; Legros et al., 2020; Carocci et al., 2020), trace-element signature determination
513 coupled to multivariate analysis could be a powerful alternative to provide relative chronology and
514 distinguish successive hydrothermal events.

515



516 Fig. 12 superposition of Beauvoir magmatic/hydrothermal transition and Beauvoir greisen fluid transport trends
 517 on PCA (component 1 and 2) from Figure 9. The Beauvoir magmatic hydrothermal transition is characterized by
 518 enrichment in Sn, Ba, Sr and depletion in Nb, Ta, W, Rb. Concerning the greisen fluid, Ba, Sr and Ti increase all
 519 along the transport, on the contrary of Sn which drastically decreases, among others.

520

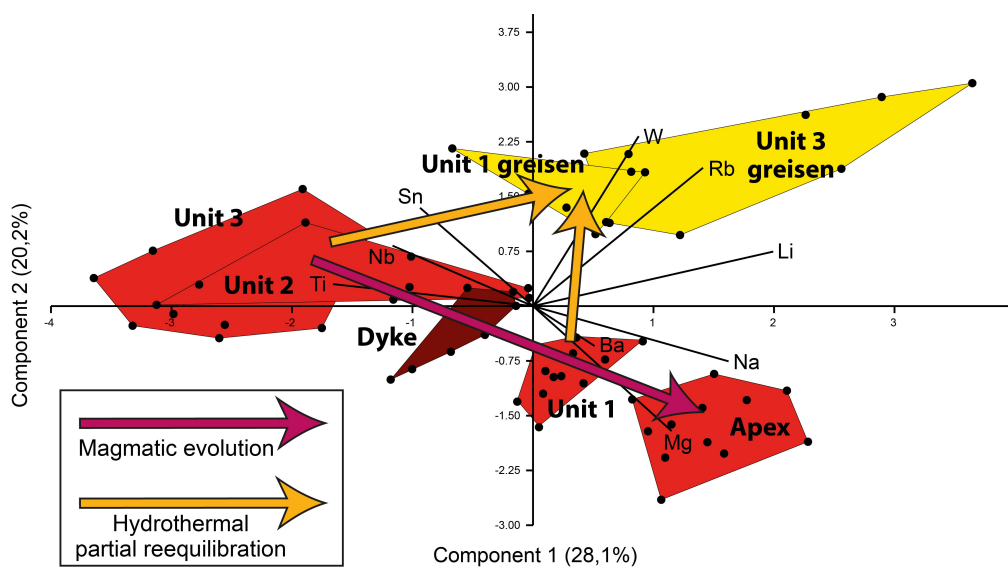
521 5.2. Genetic considerations on the Colettes/Beauvoir system

522 5.2.1. Magmatic evolution of the Colettes and Beauvoir granites.

523 Considering that the Beauvoir granite is one of the most evolved peraluminous melt occurrences known
 524 (Linnen and Cuney, 2005), the exceptional enrichment in several trace elements in its primary igneous mica
 525 (5) is not surprising. On the other hand, the strong enrichment in some elements (Nb, Ta, W, Sn) measured
 526 in zinnwaldite and muscovite from the Colettes granite was unexpected. The amount of several of these
 527 elements is comparable to those measured in micas from a number of RMG (e.g., Li et al., 2015), indicating
 528 that the Colettes granite, albeit barren of mineralization *per se*, is probably related to the evolution of the
 529 Beauvoir granite, as also suggested by their contemporaneous emplacements (Monnier et al., 2021a and
 530 references therein). Most RMG occurrences that have been drilled at depth consist of a well delimited root
 531 of zinnwaldite or two-micas granite, and an apex of albite-lepidolite/zinnwaldite granite (e.g., Yashan,
 532 Cínovec and Panasqueira granites, respectively, Li et al., 2015, Breiter et al., 2017, and Marignac et al.,
 533 2020). In contrast, deep drilling of the Beauvoir RMG revealed only the presence of the schist host rock at
 534 its base. Thus, if we assume that the Colettes granite represents the equivalent of a barren granitic root for
 535 other RMG, we could conclude that the final crystallization of the Beauvoir and Colettes granites took place

536 after melt segregation and transport of the resulting melts, contrasting with emplacement models
 537 proposed for most RMG, worldwide. The magmatic source of Beauvoir and Colettes granites could be
 538 related to the porphyry enclave (Monnier et al., 2018) commonly found in the root of the Beauvoir granite.
 539 The less evolved nature of the Colettes magma compared to Beauvoir is not recorded by the rare element
 540 signatures of micas. Indeed, the contents of some elements are higher in micas from Colettes than
 541 Beauvoir, i.e., Nb, Ta, and Sn in both muscovite (8) and zinnwaldite (3).

542 Concerning the Beauvoir intrusion, variations in the content of some elements in lepidolite (5) provide a
 543 nice record of its magmatic evolution, from the least evolved unit 3, at the root of the granite, to the highly
 544 evolved apex, illustrating a vertical zonation through the Beauvoir intrusion (Fig. 13). Surprisingly, however,
 545 in the least evolved unit, lepidolite concentrates the highest amounts of some incompatible elements such
 546 as Nb and Sn, possibly because of the relatively low abundance of competitor ore minerals in this unit.
 547 Conversely, lepidolite from the more evolved magmatic facies is enriched in Ba, Mg and Na, which mostly
 548 reflects the input of exotic fluids during the last magmatic stage, as also documented by isotopic evidence
 549 (Fouillac and Rossi, 1991) and whole-rock chemistry (Raimbault et al., 1995).



550 Fig. 13. PCA diagram projecting components 1 and 2 for LA-ICP-MS data for a selected panel of elements (Li,
 551 Na, Mg, Ti, Rb, Nb, Sn, Ba, W) from all igneous lepidolite analyzed in this study.

552 The unprecedented concentration of W in igneous micas (5) recorded in this study signifies that an
 553 important quantity of igneous-sourced W remained in the Beauvoir granite during crystallization. Also,
 554 although igneous columbo-tantalite and pyrochlore host significant amounts of W (ca. 1 and 2 wt % of

555 WO₃, respectively, see Table A3 in Appendices), a simple model based on micas chemistry, modal
556 composition and whole rock values (ca. 35 ppm; Raimbault et al., 1995) indicates that the majority of the
557 Beauvoir granite W stock is hosted in lepidolite (Fig. 14). This result has far-reaching implications:
558 incorporation of significant concentrations of W in lepidolite from the Beauvoir granite disagrees with the
559 generally accepted paradigm that W fractionates into the fluid during fluid/melt interaction (Keppler et al.,
560 1991; Zajacz et al., 2008). Furthermore, incorporation of W in igneous micas can potentially inhibit
561 subsequent hydrothermal W mineralization. In comparison, at Cínovec, where important hydrothermal W
562 mineralization occurs, magmatic zinnwaldite contains only between about 3 to 30% of the total W stock
563 (Breiter et al., 2019). As expected, Li and Rb behave similarly, i.e., they are hosted in trioctahedral micas
564 (Tischendorf et al., 2001) while Ta, and especially Sn, are mainly hosted in other minerals, undoubtedly
565 columbo-tantalite and cassiterite, respectively.

566 Only negligible variations in lepidolite Nb/Ta ratios were observed from the lowermost unit to the apex of
567 the granite. These range from a minimum of 2.7 in unit 2 to a maximum of 3.8 in unit 3. Still, these values
568 are markedly higher than the whole-rock Nb/Ta ratios (<1 in the upper units and ~1 in unit 3; Raimbault et
569 al., 1995), indicating that columbo-tantalite or pyrochlore should be expected to be an important stock of
570 Ta at Beauvoir, as in other extremely evolved melts. This process, and specifically the role of mica
571 crystallization in the decrease of the Nb/Ta ratio in granitic melts, has been thoroughly investigated by
572 Stepanov et al. (2014).

573

574 *5.2.2. Hydrothermal alteration: orthomagmatic, hybrid or exotic fluid origin.*

575 The source of fluids that precipitate W and Sn is a highly debated topic with implications for the
576 comprehension and recognition of these deposits. The Beauvoir granite and its proximal host rock
577 experienced several types of alterations (Monnier et al., 2018, 2019, 2020, 2021a), to variable intensities.
578 In chronological order, these consist mainly of: orthomagmatic alteration (localized), alkali alteration
579 (localized), and greisenization (widespread). Long after these events, a widespread kaolinisation affected
580 the whole area. As mentioned above, the chemical signature of lepidolite (4) that crystallizes from the

581 orthomagmatic fluid is strongly similar to that of igneous lepidolite, suggesting that the orthomagmatic
582 fluid and the Beauvoir melt have similar trace-element signatures. One can also deduce that the
583 orthomagmatic fluid was not subject to drastic chemical variations before precipitation of lepidolite (4)
584 and, by extrapolation, did not undergo mixing nor fluid immiscibility or segregation, at least at an early
585 stage. This also implies that a massive hydrothermal precipitation of ore minerals before formation of this
586 lepidolite (4) is to be excluded, as it would have drastically decreased the contents of Nb, Ta, W and Sn in
587 micas. Conversely, micas precipitating from the greisen episode display a markedly different signature than
588 that of igneous lepidolite. Such disparity can be explained by repeated boiling episodes of the greisen fluid
589 and its mixing with other fluids. This, as evidenced by Monnier et al. (2020), led to phase segregation and
590 dissolution of igneous minerals with coeval muscovite precipitation, efficiently altering the pristine
591 signature of the fluid. In the Echassières district, Sr and Ba incorporation in micas correlates roughly with
592 the distance from the magmatic source, reflecting a potential contamination from the schist during fluid-
593 rock interaction. Similar findings are also reported by Codeço et al. (2021) at Panasqueira (Portugal) for
594 both elements, and by Legros et al. (2018) for Ba at Jiangxi (China; Sr was not analyzed). Consequently,
595 variations in the concentration of these elements in mica could be used as a proxy for the distance traveled
596 by a fluid from its granitic source. Monitoring in parallel distal hydrothermal micas for their content in W,
597 Sn, Nb-Ta, should prove mica trace-element signature as a robust tool for targeting potentially mineralized
598 granitic body.

599 *5.2.3. Metals behavior during greisen alteration of the Beauvoir granite.*

600 Recently, an experimental study by Schmidt et al. (2020) highlighted the contrasting behavior of W and Sn
601 during melt/fluid partitioning, replicating the distal hydrothermal deposition of W and the conservation of
602 Sn in the magma at late stages of crystallization, as is generally inferred for natural systems. Nevertheless,
603 their study indicates that W preferentially fractionates in the melt phase for strongly peraluminous
604 compositions, as is the case for the Beauvoir granite. At first sight, this remarkable result contradicts the
605 common observation that, in some peraluminous RMG, most cassiterite is igneous while wolframite
606 crystallizes in quartz veins (e.g., Argemela, Michaud et al., 2020; Abu Dabbab, Zoheir et al., 2020; Beauvoir,

607 this study). Moreover, it has been shown experimentally that wolframite cannot reach saturation in a
608 halogen-rich melt above 500 °C (Che et al., 2013), implying that, in these peraluminous melts, W is hosted
609 in (an) igneous phase(s) other than wolframite. In the case of the Beauvoir granite, and likely for other
610 peraluminous RMG, igneous mica incorporates significant quantities of W, and fluid exsolution during
611 magmatic crystallization will not lead to W mineralization, in agreement with the high melt/fluid
612 partitioning for W highlighted by Schmidt et al. (2020). Monnier et al. (2020) demonstrated that, in the
613 Echassières district, the mineralization of wolframite *c*, found outside of the Beauvoir granite, clearly
614 postdated crystallization of the latter. They also showed that the tungsten could not have derived from the
615 early orthomagmatic fluid. If its source is unclear, the most likely hypothesis would be remobilization
616 during greisen alteration. Dissolution of the early wolframite generations (*a* and *b*) may have provided the
617 W to form wolframite *c*, but textural evidences for replacement of early wolframite generations are scarce
618 at best and do not support this postulate. In the light of the findings discussed so far, it would seem
619 worthwhile to seek answers by investigating the composition of associated mica groups. Surprisingly, W
620 concentrations measured in greisen muscovite from Beauvoir greisen are particularly low compared to
621 igneous lepidolite. Indeed, our data (Fig. 14) indicate that lepidolite stocked the majority of the W
622 contained in the Beauvoir

623 granite, as it is generally the case for white micas in peraluminous granites (e.g, Neves, 1997). It is
 624 therefore a likely scenario that alteration of lepidolite into muscovite, occurring during greisen alteration of
 625 the granitic stock, allowed remobilization of important quantities of W, which did not incorporate in the
 626 newly formed, W-depleted, muscovite. A quick calculation can provide an estimate of the mass of W
 627 released to the greisen fluid (Fig. 15, see the caption for more details). For a volume of ca. 5 % of greisen
 628 altered granite, the quantity of W released corresponds to 5000 T of WO_3 , i.e., an amount significantly
 629 higher than that of the W stock assessed for the Mazet veins prospect, where most wolframite c was found
 630 (Aubert, 1969). This result is provocative and implies that if a similar/higher volume of a rare-metal granite
 631 with an analogous chemistry to Beauvoir were to be altered to a similar/higher intensity, the greisen fluid
 632 could remobilize a sufficient amount of W to form a major deposit. A similar approach, using trace element
 633 differences between igneous and replacing hydrothermal mica, reached the same conclusions for the
 634 Cínovec granite (Johan et al., 2012), though these are subject of debate (Breiter et al.,).

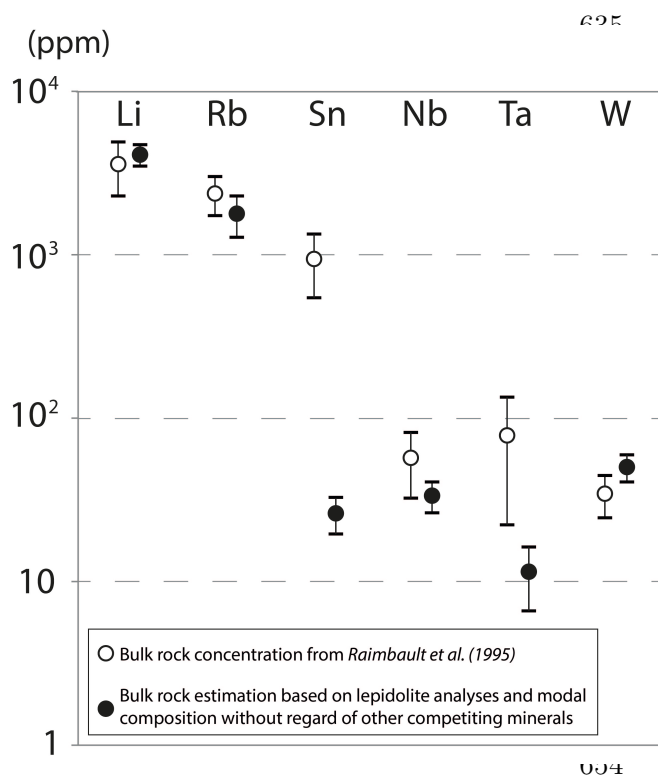
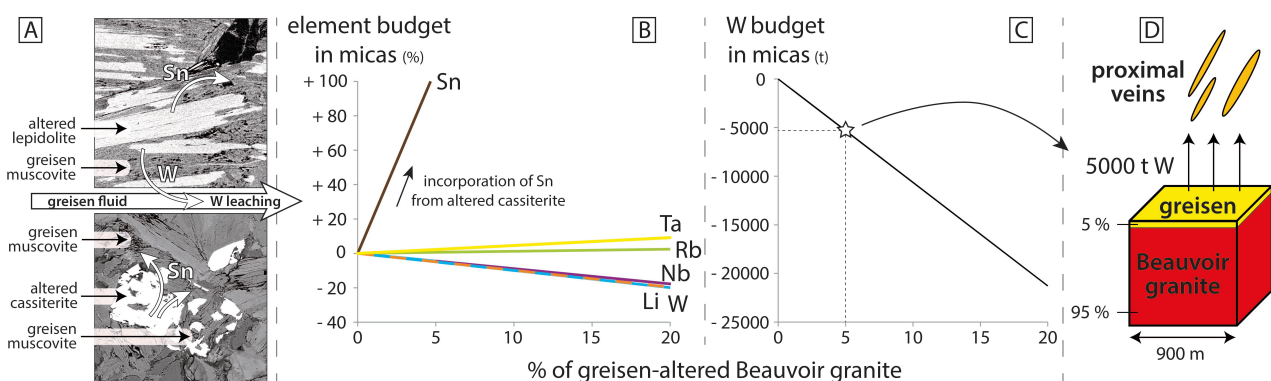


Fig. 14. Comparison between bulk rock concentrations in rare metals and their estimated stock in micas for the Beauvoir granite. The latter value was obtained combining the LA-ICP-MS analyses in this study with modal mica compositions. Mean bulk rock data was calculated from concentrations measured in drill core by listed in Raimbault et al. (1995). Altered samples, uncommon rock facies and samples displaying a high number of outlier concentrations were not considered. Intervals used are (in meters below surface): Unit 1: -129.4; -208.9; -269.7; -354.6. Unit 2: -445.2; -507.6; -517.4; -521.6; -561.9; -632.7; -656.8; -716.6. Unit 3: -751.1; -771.0; 785.6. Error bars correspond to the standard deviations of lepidolite and whole rock analyses. The similar bulk and mica Li and Rb stocks displayed validate this approach, considering the fact that these elements occur mostly in mica.

655

656 In a few greisen samples, igneous lepidolite relicts were relatively well preserved (Fig. 2.C). Their
 657 composition differs from that of the igneous lepidolite population from unaltered samples, which has a

658 relatively consistent chemistry, by an increase in Rb, Li and W (Fig. 13). These chemical changes are clearly
 659 minor in regard of the differences segregating the mica groups sourced by distinct magma or fluids.
 660 Nevertheless, this observation suggests a partial re-equilibration of the lepidolite relicts during greisen fluid
 661 percolation, confirming its particularly high W content. Other similar outcomes, such as the significant
 662 increase of W concentration monitored in hydrothermal pyrochlore and columbo-tantalite (ca. 3 wt % of
 663 WO_3) compared to the igneous species (Table A3), confirm the global W enrichment in the Beauvoir greisen
 664 fluid. We assume that this enrichment is mainly due to release of W during alteration of lepidolite into
 665 muscovite, as discussed in section 5.2.3 (Fig. 15). It follows that the conditions necessary for the
 666 breakdown of lepidolite, or other W-rich micas, are key factors for releasing potential economic quantities
 667 of this metal. Observations at a global scale reveal that particularly evolved peraluminous granites have a
 668 better chance of developing quartz vein-hosted W mineralization. Yet, the experimental investigation of
 669 Schmidt et al. (2020) indicates that, in strongly peraluminous systems, most of the W is retained in the
 670 magma throughout granite crystallization. Their results are inconsistent with precipitation of abundant
 671 hydrothermal wolframite, given that a fluid exsolving from a crystallizing granite – the generally evoked
 672 process to account for these mineralizations – would probably carry only little W in solution. An
 673 explanation to this paradox would be that, sometime after crystallization of the granite, metals are release
 674 into the fluid that is altering the W-bearing igneous minerals. Such a fluid could be a moderate-
 675 temperature evolved greisen fluid, similar to what we have documented at Beauvoir (Monnier et al., 2020).



676 *Fig. 15. A conceptual model summarizing the transferring process of metals from the source rock to the*
 677 *depositional site via fluid transport. [A]: During greisen alteration, primary lepidolite and cassiterite are*
 678 *replaced by muscovite. Tungsten in lepidolite is leached by the fluid while Sn from both altered minerals is*
 679 *incorporated in newly formed muscovite. [B]: Based on modal proportions and average compositions of micas*
 680 *from this study, this diagram monitors the estimated budget of metals in micas during greisen alteration of the*
 681 *Beauvoir granite. The steep increase of the Sn budget in mica is only possible for a case where Sn is provided by*

682 (an) other source(s), such as cassiterite; if all Sn were to come from lepidolite, even if entirely incorporated in
683 muscovite (100%), there would be a steady state in the total budget (0%), as is approximately the case for Rb
684 and Ta. Conversely, the negative slopes for Nb, Li and W indicate that these elements are not incorporated in
685 greisen mica. This may imply that they are fractionated into the fluid, although Nb-Ta behavior is also
686 constrained by formation/dissolution of columbo-tantalite and pyrochlore. [C]: This diagram focusses on the W
687 budget (in metric tons) in mica, for the same amount of rock altered as in B. It indicates that, for 5% altered
688 granite, $5 \cdot 10^3$ t of tungsten are removed from the altered rock. This amount of metal is transported by the fluid
689 and reprecipitated in form of wolframite in proximal veins [D]. The model considers that lepidolite is the main
690 primary host of W (cf., Fig. 14) and that secondary minerals other than wolframite do not incorporate
691 significant amounts of this metal. In this simulation, unaltered granite is considered to contain 18 % modal
692 lepidolite and have a whole-rock average density of 2.6, for a total granite volume of 7.29×10^8 m³. The modal
693 proportion of greisen muscovite (9) is 66%, with average bulk density of 2,3 for the greisen. The simulation
694 assumes that the altered portion of the primary rock is entirely transformed to greisen.

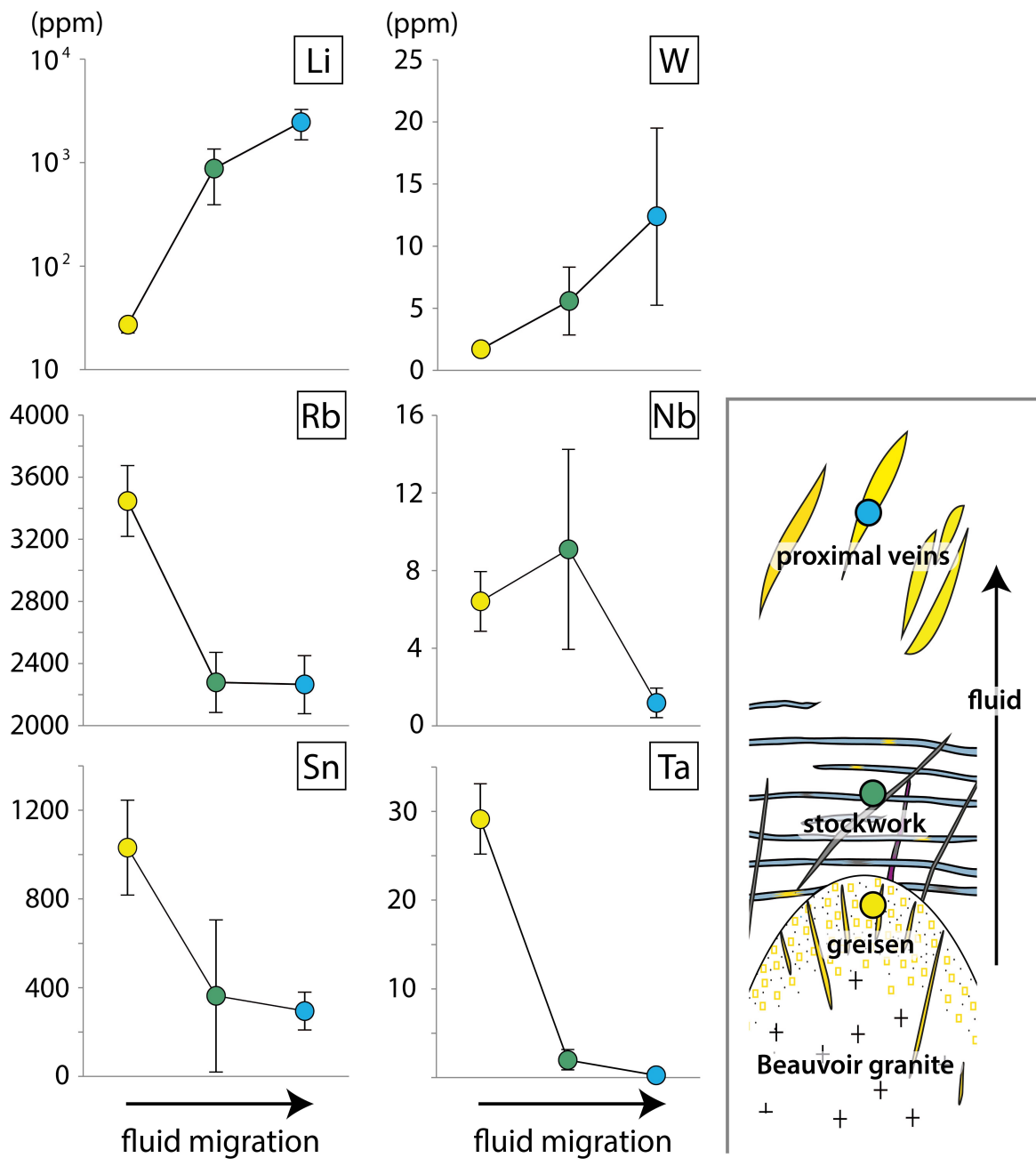
695 Cassiterite is the main Sn-bearing igneous mineral in the Beauvoir granite (Fig. 15). Considering the fact
696 that cassiterite was strongly altered during greisen alteration (Wang, 1988; Monnier et al., 2019), it could
697 be expected that, as was the case for W, Sn should undergo important remobilization away from the
698 granite during alteration by the greisen fluids (e.g., Linnen and Cuney, 2004; Chen et al., 2019). However,
699 only minor cassiterite occurs outside of the granite, far too little to account for the high quantity of igneous
700 cassiterite altered. However, greisen muscovite (9) that replaced igneous minerals contains exceptional Sn
701 contents, largely superior to the amount recorded in igneous lepidolite. Also, the modal amount of
702 muscovite in greisen is much higher than that of lepidolite in the fresh granite. It is thus quite likely that Sn
703 released from cassiterite and lepidolite during alteration was mostly trapped in greisen muscovite, thus
704 inhibiting Sn migration outside of the Beauvoir granite. A rough estimate of the muscovite capacity to stock
705 Sn issued from altered cassiterite can be obtained using the same simple model used above to calculate W
706 remobilization. An alteration of ca. 5 % of the granite volume would lead to incorporation of ca. 50,000
707 metric tons of Sn in muscovite. This represents an important amount, considering that the estimated Sn
708 stock for the 300 upper meters of the granite was of ca. 150,000 metric tons (Cuney et al., 1992). Niobium,
709 Ta, Li and Rb form two element couples with similar behavior in geological systems, so that they are
710 commonly incorporated in the same minerals in evolved peraluminous granites, i.e., columbo-tantalite for
711 Nb-Ta and micas for Li-Rb. The latter pair are particularly enriched in lepidolite from the Beauvoir granite
712 (5), however, only Rb was efficiently incorporated in the greisen muscovite (9), inducing a roughly steady
713 stock of Rb in Beauvoir micas during greisen alteration. Lithium, on the other hand, is very low in greisen
714 muscovite, suggesting that, in contrast to Rb, it remained mostly in solution after lepidolite dissolution.

715 Again, if we use the above model, altering 5 % of Beauvoir granite to greisen and considering that
716 muscovite was the only mineral efficiently incorporating Li, would induce an estimate of ca. $4 \cdot 10^5$ T of Li to
717 be removed from the granite. This observation confirms that the Li anomaly marking a well-defined
718 geochemical halo in the schist around the Beauvoir granite (Cuney et al., 1992) can be related to greisen
719 alteration of the granite. Similarly to Li and Rb, Nb and Ta were unequally incorporated in greisen
720 muscovite (9). The Nb/Ta ratio is lower in greisen muscovite than in igneous lepidolite, suggesting that the
721 solubility of Nb species exceeds that of Ta in the greisen fluid. A similar process has been proposed for
722 zoned micas of the Dahutang deposit (Yin et al., 2020), confirming that Nb has a higher affinity for greisen
723 fluids than Ta does. Nevertheless, in the Beauvoir granite, columbo-tantalite and pyrochlore seem to be
724 major reservoirs for Ta and possibly Nb (Fig. 14) both at the magmatic and greisen stages, complicating the
725 interpretation of variations in Nb and Ta concentrations in micas. It would therefore be advisable to
726 envisage a paired investigation on these ore minerals in the Echassières district, and, in a more general
727 manner, other occurrences elsewhere.

728 *5.2.4. Transport of Sn, W, Nb, Ta, Li and Rb during the greisen episode.*

729 The various micas that crystallized from the greisen fluid form a group with a coherent chemical signature
730 that differentiates them from the other mica groups. Nonetheless, several minor differences do occur in
731 each subgroup (Fig. 9 and 10, Table 3 and 4), particularly as a function of the distance from the Beauvoir
732 granite. These differences illustrate the ability of the greisen fluid to transport some of these key elements.
733 They also provide a vectoring tool, assuming that the increased content of metals that were strongly
734 remobilized from the Beauvoir granite during greisen alteration (i.e., Li, W and possibly Nb) reflects their
735 respective decrease in solubility in the greisen fluid as the latter evolves during migration. On the contrary,
736 the contents of elements that mostly remained in the granite after greisen alteration (i.e., Rb, Sn and Ta)
737 decrease markedly in micas with distance from the granite, corroborating that the fluid is not able to
738 mobilize these metals (Fig. 16). These observations are consistent with the estimated metal budget in
739 Beauvoir micas during greisen alteration: rare elements that strongly incorporated in greisen muscovite (9)
740 are very low in muscovite (11) from the Suchot vein (located outside the granite), while elements that were

741 easily remobilized in the greisen fluid are relatively concentrated in micas outside of the granite. This
 742 process led to the occurrence of wolframite *c* outside of the Beauvoir granite (increase of W in micas after
 743 fluid transport) while cassiterite is scarce in these same rocks (decrease of Sn in micas after fluid transport).
 744 The above observations confirm that the mica trace-element signatures provide a suitable record for ore
 745 metal remobilization and transport. Additional studies involving ore element mobility in greisen fluids are
 746 clearly needed in order to better constrain the genesis of greisen-related deposits, globally.



747 *Fig. 16. On the left-hand side are depicted mean and standard deviation concentrations of a number*
 748 *of critical metals analyzed in the greisen muscovite from granite, stockwork and proximal veins*
 749 *(respectively, yellow, green and blue circles in the sketch to the right). The data trends record the*
 750 *evolution of the fluid during transport of these metals (fluid migration).*

751 In summary, it is clear that crystallization and dissolution of micas played a key role in determining the fate
752 of at least some critical metals throughout the magmatic/hydrothermal evolution of the Beauvoir granitic
753 system. In particular, remobilization of W, Sn, Li and Rb was strongly influenced by dissolution of igneous
754 lepidolite and crystallization of greisen muscovite, while the process of Nb and Ta remobilization seems
755 more complex, possibly controlled by the behavior of columbo-tantalite and pyrochlore
756 precipitation/alteration. Metal incorporation into greisen muscovite (W-poor and Sn-rich) controlled the
757 leaching of W away from the granite, while the Sn remobilized from cassiterite remained sequestered in
758 the granite (Table 5). A similar process (alteration of metal-enriched igneous mica) was evoked as the
759 possible mechanism for the origin of the Nb-Ta hydrothermal mineralization and peripheral W vein
760 formation in the Yashan granite (China, Li et al., 2015). This process must therefore be kept in mind when
761 investigating greisen-related W deposits, particularly those associated to RMG peraluminous magmatism.

Metal	Main igneous mineral host	Behavior during greisen alteration	Fate after transport away from the granite
W	Lepidolite	Mainly remobilized by fluids	Wolframite c
Sn	Cassiterite	Mostly incorporated in greisen muscovite	Limited occurrence of cassiterite
Li	Lepidolite	Mainly remobilized by fluids	Incorporation in metamorphic mica
Rb	Lepidolite	Mostly incorporated in greisen muscovite	Limited occurrence
Nb	Columbo-tantalite (lepidolite)	Mainly remobilized by fluids	Lack of mineralization evidence. Incorporation in rutile or micas ?
Ta	Columbo-tantalite, pyrochlore	Mostly incorporated in hydrothermal pyrochlore	Limited occurrence

762 *Table 5: Ore metal behaviors during Beauvoir greisen alteration.*

763 **6. Conclusions**

764 A multivariate statistical approach on trace element data obtained from micas by LA-ICP-MS provides
765 important insights on the magmatic/hydrothermal processes that can lead to ore accumulation. Key
766 findings of this study are hither condensed:

- 767 • Micas chemical signatures provide a direct vector to W, Sn, Nb-Ta mineralizations.

- 768 • During a single hydrothermal episode, micas precipitating in different hydrothermal parageneses
769 maintain a roughly similar trace-element signature. This also applies to micas formed within (in the
770 quartz or topaz matrix) and outside a vein (in alteration halo). It follows that fluid-rock interaction
771 has a low impact on the chemical compositions of newly formed micas.
- 772 • The trace element chemistry of micas can be used as a powerful tool to differentiate or link
773 different alteration parageneses, more efficient than major elements and/or petrographic features.
774 It can also relate distal dykes to a main magmatic body.
- 775 • Igneous micas can host significant concentrations of W and Sn which can be remobilized during
776 hydrothermal alteration to potentially form ore deposit. Conversely, hydrothermal mica
777 precipitation can inhibit ore formation by incorporating substantial quantities of metals.

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786

787 **Appendices**

788 Appendices contains the following research data:

789 Table A1: EPMA analyses of micas from Echassières district (in wt. %)

790 Table A2: LA-ICP-MS analyses of micas from Echassières district (in ppm)

791 Table A3: EPMA analyses of columbo-tantalite and pyrochlore from the Beauvoir granite (in wt. %)

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