

# $\begin{array}{c} \mbox{Prediction of nanomagnetite stoichiometry} \\ ({\rm Fe}({\rm II})/{\rm Fe}({\rm III})) \mbox{ under contrasting pH and redox} \\ \mbox{ conditions} \end{array}$

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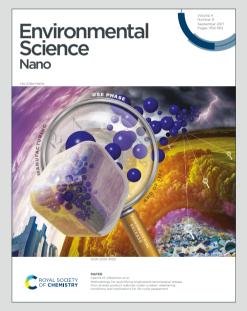
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# Prediction of nanomagnetite stoichiometry (Fe(II)/Fe(III)) under contrasting pH and redox conditions

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#### Abstract

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Magnetite (Fe(III)<sub>2</sub>Fe(II)O<sub>4</sub>) nanoparticles are fascinating nanoparticulate minerals for their electronic, magnetic and chemical properties. Ubiquitous, in the environment, they are also among the most used ferromagnetic nanomaterials in environmental, industrial and biomedical applications. Their intriguing structural and reactivity features do not only arise from the "nano-effect" but also, from the occurrence of Fe<sup>2+</sup> ions in their structure. Previous studies showed that partial oxidation of (nano)magnetite may occur. However, such transformations were only monitored under either oxidizing or very acidic conditions. Here, we report that 10 nm-sized stoichiometric magnetite particles (Fe(II)/Fe(III) = 0.5) are in fact not stable in aqueous solutions over a biologically- and environmentally- relevant pH range (4-7). In the absence of O<sub>2</sub>, an H<sup>+</sup>-promoted dissolution process is responsible for the preferential release of Fe(II) into solution, which leads to partial oxidation of magnetite to a magnetite-maghemite solid-solution. Long-term kinetic investigations combined with XMCD measurements reveal that the dynamic exchange of Fe(II) between the surface and the solution is key in determining the magnetite stoichiometry even at circumneutral pH. Based on this finding, we developed a thermodynamic model for the magnetite-maghemite solid-solution able to predict the chemical stability of the 10 nm-sized magnetite. This model enables to rationalize and predict the behavior and transformation of magnetite nanoparticles in aqueous solutions, which is crucial for a broad range of applications (medicine, biology, chemistry, environment, etc.).

#### **Environmental significance**

Magnetite nanoparticles are highly reactive ubiquitous nanominerals in natural systems, which control the behavior and fate of many contaminants. However, their reactivity depends on the composition, especially their Fe(II) and Fe(III) content. Therefore, it is crucial to understand and predict the effect of various factors on magnetite stoichiometry (Fe(II)/Fe(III)), as it is known to vary greatly in aquatic environments. This work focuses on two of the most important abiotic processes controlling magnetite stoichiometry, namely the oxidation of Fe(II) to Fe(III) and the preferred release of Fe(II) over Fe(III) to the solution at acidic pH. It provides the first predictive model of nanomagnetite stoichiometry under various pH and redox conditions, which could be used to unravel the complex behavior and reactivity of these nanomaterials in contrasted environmental settings.

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## **1** Introduction

Magnetite is a mixed-valence iron oxide with an inverse spinel structure, which is commonly found at a nanometric size in soils and sediments. Its formula is  $[Fe^{2+} Fe^{3+}]_{Oh} Fe^{3+}_{Td}O^{2-}_{4}$ , where high spin  $Fe^{2+}$  and  $Fe^{3+}$  can be found in octahedral (Oh) and tetrahedral (Td) coordination. However, magnetite can be partially or entirely oxidized into the isostructural mineral-phase, maghemite (Fe<sub>2</sub>O<sub>3</sub> or  $[Fe^{3+}_{5/3}\Box_{1/3}]_{Oh}Fe^{3+}_{Td}O^{2-}_{4}$ ), containing vacant Oh sites ( $\Box$ ) instead of  $Fe^{2+}$ . Magnetite, maghemite and intermediate phases are characterized by the Fe(II)/Fe(III) ratio, denoted R, ranging from 0 (pure maghemite) to 0.5 (pure magnetite).<sup>1–3</sup> The occurrence of Fe<sup>2+</sup> ions in the magnetite structure gives rise to its unique electronic and magnetic properties, and makes it a key-ingredient in a wide range of applications in catalysis, biomedical and environmental chemistry.

On one hand side, the redox sensitivity of magnetite is problematic for many applications, because it depends on the mixing of iron valence states in the crystal structure. For instance, the high bulk electron conductivity of magnetite is only possible due to the fast electron hopping between  $Fe^{2+}$  and  $Fe^{3+}$  ions on Oh sites,<sup>4</sup> or the strong magnetic moment of magnetite fully depends on the presence of  $Fe^{2+}$ , because the magnetic moments of both types of  $Fe^{3+}$  ions cancel each other.<sup>1</sup> In such cases,  $Fe^{2+}$  release or its oxidation to  $Fe^{3+}$  may impede magnetite applicability.

On the other hand side, the redox properties make magnetite excellent "batteries" for both human activities<sup>5</sup> and microorganisms,<sup>6</sup> and are responsible for the reduction or oxidation of redox-sensitive chemical compounds at the magnetite surface in industrial and environmental processes.<sup>2,7–12</sup> or in medical applications.<sup>13,14</sup> For instance, the impact of the stoichiometry of nanomagnetite on the thermodynamics and kinetics of the reduction reactions of inorganic and organic compounds has been investigated for U(VI) reduction to U(IV)<sup>8</sup> and for nitrobenzene reduction to aniline<sup>7</sup>. The adsorption of organic contaminants (e.g. quinolone antibiotics)<sup>15</sup> was also shown to increase with R, which could have a drastic effect on the fate and behavior of such contaminants in natural systems. For these reasons, it appears crucial to understand and predict the impact of physico-chemical conditions on magnetite stoichiometry.

The oxidation of magnetite to maghemite is well documented and, in the presence of  $O_2$ , can be formulated as:

$$4 \operatorname{Fe_3O_{4(s)}} + \operatorname{O_{2(g)}} \rightleftharpoons 6 \operatorname{Fe_2O_{3(s)}} \tag{1}$$

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$$Fe_3O_{4(s)} + 2 H^+_{(aq)} \rightleftharpoons Fe_2O_{3(s)} + Fe^{2+}_{(aq)} + H_2O_{(l)}$$
 (2)

This reaction is highly relevant for many applications of magnetite nanoparticles in aqueous suspensions, e.g. medicine,<sup>13,14</sup> biology,<sup>6,16</sup> chemistry,<sup>17,18</sup> environment,<sup>19,20</sup> Unfortunately, the transformation of nanomagnetite into non-stoichiometric nanomagnetite via equation 2 has only been characterized in acidic solutions (e.g. pH < 3),<sup>21,22</sup> with large magnetite particles (> 100 nm),<sup>23</sup> at high temperature,<sup>24</sup> or in presence of O<sub>2</sub>.<sup>19</sup> Knowledge on the magnetite stability and reactivity under low O<sub>2</sub> content and slightly acidic to circumneutral pH is scarce, even though these conditions are very often encountered in biological media or subsurface environments. Previous reports have either limited investigations on the magnetite reactivity to neutral and alkaline conditions under O<sub>2</sub>-free atmosphere (e.g.<sup>7,8,10,15,25</sup>), or overlooked Fe<sup>2+</sup> release at low pH through the chemical transformation of magnetite, which occurs under these conditions (e.g. 9,12,26). In addition, because slow processes can be observed, only the kinetic aspects of equation 2 were treated.<sup>23</sup> This is problematic, especially in natural systems experiencing large fluctuations of the local pH and redox conditions, because equilibrium data are also required for the prediction of magnetite stoichiometry. Given the extensive use of magnetite nanoparticles and their importance in a biological and environmental context, it becomes urgent to fully characterize the impact of physico-chemical conditions, and especially the pH on magnetite stoichiometry, and the consequences on magnetite redox and surface reactivity in aqueous systems.

In the present work, an extensive dataset on the solubility of magnetite nanoparticles (10 nm-sized) in aqueous suspension is presented. Long-term kinetic investigations (up to 560 days) shed light on overlooked aspects of the dissolution processes and whether chemical equilibrium can be reached under controlled laboratory conditions. Based on these data the first chemical thermodynamic model for nanosized magnetite stoichiometry in aqueous solutions is developed, capable of predicting the composition of the investigated nanomagnetites as a function of solution chemistry (i.e. dissolved Fe<sup>2+</sup> concentration, pH, and redox potential). Given the extensive use of magnetite nanoparticles in biological and environmental sectors, this model will help in examining and predicting the reactivity of either natural magnetite nanoparticles or engineered ones prior to technological applications.

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59 60 2 Materials and methods

#### 2.1 Chemicals.

Iron(III) chloride hexahydrate (FeCl<sub>3</sub>·6H<sub>2</sub>O), Iron(II) chloride tetrahydrate (FeCl<sub>2</sub>·4H<sub>2</sub>O), and NaCl were purchased from AnalaR NORMAPUR. Hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>), hydrochloric acid (HCl), sodium hydroxide (NaOH), and sodium citrate were obtained from Sigma-Aldrich. The sample solutions were prepared with ultrapure "MilliQ" water (specific resistivity 18.2 M $\Omega$ cm<sup>-1</sup>). All experiments were carried out in an anaerobic chamber (N<sub>2</sub>-glovebox, JACOMEX) and all solutions were purged from O<sub>2(g)</sub> prior to use. All pH adjustments were done by HCl and NaOH addition (no buffer was used).

#### 2.2 Synthesis of magnetites with various stoichiometries.

Magnetite nanoparticles (10 nm) preparation was described previously.<sup>27</sup> Samples used in the present study come from the same batch. The synthesis procedure is repeated for clarity. The procedure involves a room temperature aqueous precipitation.<sup>27,28</sup> Briefly, a 0.5 M HCl solution (40 mL) containing a 0.5 M FeCl<sub>2</sub> and 1 M FeCl<sub>3</sub> (1:2 molar ratio) was added dropwise into a 0.5 M NaOH solution (250 mL) with continuous stirring, leading to instantaneous precipitation of magnetite particles. After the synthesis, the pH of the suspension was adjusted to 8.5. Then, the solid phase was washed three times with ultrapure water at pH 8.5 (adjusted using NaOH) to avoid any Fe<sup>2+</sup> release, as observed in previous work,<sup>15,29</sup> and thus, to guarantee the stoichiometry of R0.5. Known amounts of H<sub>2</sub>O<sub>2</sub> were added to R0.5 to produce defined non-stoichiometric magnetite (R0.1, R0.2, R0.3, R0.4) sets.<sup>15,27</sup> Then, mother solutions were replaced with ultrapure water (at pH = 8.5) in order to remove residual H<sub>2</sub>O<sub>2</sub>. Initial Fe(II)/Fe(III) ratios (R<sub>ini</sub>) of these (non)-stoichiometric magnetites were determined by acid digestion in 0.6 M HCl during 3 days, followed by spectrophotometric determination of dissolved [Fe(II)] and total [Fe] (= [Fe(III)] + [Fe(II)]) using the 1-10 phenantroline colorimetric method,<sup>30</sup> as well as by X-ray Magnetic Circular Dichroism (XMCD).<sup>27</sup> Results were in excellent agreement with expected values from the amount of H<sub>2</sub>O<sub>2</sub> added, as expected from previous studies.<sup>15,27,31</sup> Acid digestion followed by spectrophotometry was repeated three times for R0.5 and each non-stoichiometric magnetite, which led to an error of  $\pm 0.01$  on the determination of R<sub>ini</sub>. This error is consistent with previous work.32

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#### 2.3 Batch studies.

All experiments were conducted with 10 mM NaCl as background electrolyte. Fe dissolution kinetic studies were at pH 4, 5.5 and 8 for  $R_{ini} = 0.5$ , and at pH 4 for  $R_{ini} = 0.10$  and 0.30. Iron total concentration in the suspensions was 6.5 mM, corresponding to ~0.5 g L<sup>-1</sup> of magnetite, in 100 mL suspensions in 250 mL HDPE bottles. pH was controlled and regularly adjusted over a 560-days period using HCl and NaOH. Equilibrium studies were conducted in 10 mM NaCl solutions and in 15 mL polypropylene tubes containing 10 mL of solution. Reaction time was 20 days (for  $R_{ini} = 0.50$ ), 60 days ( $R_{ini} = 0.30$ , 0.40) or 320 days ( $R_{ini} = 0.10$ , 0.20). The impact of dissolved Fe(II) addition on the recharge of R0.1 was investigated by adding small amounts of a 100 mM FeCl<sub>2</sub> solution. The impact of the presence of an excess of Fe(II) or citrate on R0.5 was investigated by adding 0.25, 0.5 or 1 mM Fe<sup>2+</sup>, or 1 mM citric acid to R0.5. After a given reaction time, pH and redox potential ( $E_H$ ) were measured and the suspension was passed through a 0.2 µm cellulose acetate filter (Sartorius Minisart), followed by dissolved [Fe(II)] and [Fe(III)] determination by spectrophotometry. Citrate-containing samples were ultrafiltered at 5kDa (Vivaspin 15RH12, Sartorius). The absence of Fe nanoparticles in the filtrate was checked by Dynamic Light Scattering (VASCO Flex).

The effective stoichiometry of magnetite ( $R_{eff}$ ) was calculated from the initial stoichiometry ( $R_{ini}$ ), the total concentration of Fe(III) in both the aqueous and solid phases ([Fe(III)]<sub>tot</sub>) and the measured aqueous Fe<sup>2+</sup> concentration ([Fe(II)]<sub>aq</sub>):

$$R_{eff} = R_{ini} - \frac{[Fe(II)]_{aq}}{[Fe(III)]_{tot}}$$
(3)

By assuming 5% uncertainty on  $[Fe(II)]_{aq}$  determination by spectrophotometry, and accounting for the uncertainty on R<sub>ini</sub>, we determine that the error on R<sub>eff</sub> was comprised between 0.01 and 0.03.

pH and  $E_H$  were recorded in both magnetite and Fe(II)-amended magnetite suspensions using a multiparameter electrode (pH,  $E_H$  and T; Hach, sensION+5045). The pH electrode was calibrated with 3 standard buffers (pH 4, 7, and 10). The Pt electrode combined with a Ag/AgCl reference electrode, used for redox potential measurements, was calibrated using a commercial redox buffer (220 mV vs Ag/AgCl). Raw data were converted into  $E_H$  vs standard hydrogen electrode (SHE) by correcting for the potential of the reference electrode. An equilibration time of 15 minutes was applied for all  $E_H$  measurements. The suspension was stirred prior to the  $E_H$  measurements. The electrode surface was periodically cleaned by exposing it to 0.1 M HCl for 1 h.

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#### 2.4 Characterization by TEM and XMCD.

Transmission Electron Microscopy (TEM; Jeol JEM 1230 microscope) was used for characterization of magnetite nanoparticles. Briefly, A small aliquot of magnetite suspension was diluted with ultrapure water and sonicated for 20 min. A droplet of the diluted suspension was deposited on a carbon-coated 200-mesh copper grid and dried inside the anaerobic chamber. Samples were transported to the microscope in an N<sub>2</sub> atmosphere using a hermetic holder and the samples were analyzed at an acceleration voltage of 200 kV. Average particle diameters were determined by measuring 100 particles.

XMCD signals were recorded at the Fe  $L_{2,3}$  edges (700 – 730 eV) on the DEIMOS beamline at the synchrotron light source SOLEIL as previously described.<sup>27,33,34</sup> Briefly, samples were transported to the SOLEIL Synchrotron facility at the DEIMOS beamline in 1 mL tubes, placed in airtight bottles that had been closed in the N<sub>2</sub>-glovebox. Colloidal suspensions of nanoparticles were drop-casted on silicon substrates and dried at room temperature, in an Ar-glove box ( $O_{2(g)} < 1$  ppm) connected to the end station. The silicon substrates were fixed on a sample holder and transferred into the cryomagnet, under ultrahigh vacuum (UHV-10<sup>-10</sup> mbar). All spectra were measured in Total Electron Yield mode (TEY) at 4.2 K under UHV conditions and an applied magnetic field H ( $H^+ = +6$  Tesla and  $H^{-}$  = -6 Tesla). The beam size was 800\*800  $\mu$ m<sup>2</sup> and the resolution was 100 meV. XMCD spectra were plotted by considering the absorption cross-section measured with left ( $\sigma_L$ ) and right ( $\sigma_R$ ) circularly polarized X-rays. XMCD spectra were plotted as  $\sigma_{XMCD} = (\sigma_+ - \sigma_-)$  where  $\sigma_{+} = [\sigma_{L}(H^{+}) + \sigma_{R}(H^{-})]/2$  and  $\sigma_{-} = [\sigma_{L}(H^{-}) + \sigma_{R}(H^{+})]/2$ . The circularly polarized X-rays are provided by an Apple-II HU-52 helical undulator for XMCD measurements and by sweeping the magnetic field from +6T to -6T. XMCD signals were normalized by dividing the raw signal by the height of the maximum.

#### 2.5 Aqueous speciation modeling

PHREEQC (version 2)<sup>35</sup> is a computer code based on an ion-association aqueous model, which was designed to perform speciation and saturation-index calculations in water. Davies equation was used for activity coefficients calculation, being valid up to an ionic strength of 0.1 M. PhreePlot<sup>36</sup> contains an embedded version of the geochemical speciation program PHREEQC. It includes a parameter optimization procedure, which automatically adjusts a model to experimental data by minimizing the weighted sum of squares of the residuals.

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#### **3** Results

# **3.1** Time dependent Fe<sup>2+</sup> release from magnetite.

When maintaining stoichiometric magnetite nanoparticles (0.5 g L<sup>-1</sup>) at pH = 8 in 10 mM NaCl solution, a marginal release of Fe into solution is observed, evidencing the stability of magnetite at these conditions. By contrast, at pH = 5.5, a fast release of Fe<sup>2+</sup> occurs already after 1 day, while aqueous Fe<sup>3+</sup> is not detectable. This incongruent dissolution process impacts the effective stoichiometry of magnetite (R<sub>eff</sub>; Fig. 1a). Accordingly, R<sub>eff</sub> drops from 0.50 to 0.36 within 1 day, and then remains constant over time, suggesting the establishment of a chemical equilibrium.

Exposing nanomagnetite to a pH 4 solution leads to a rapid release of Fe<sup>2+</sup> within the first two days, with a concomitant decrease of R<sub>eff</sub> from 0.50 to ca. 0.20. The dissolution process slows down when R<sub>eff</sub> approaches 0.10, but it seems to tend to 0 over time scales much longer than that of this study ~1.5 year (540 days) (Fig. 1a). The experiments on Fe(II) release from stoichiometric magnetite are complemented by experiments using nanomagnetites of defined stoichiometries ( $R_{ini} = 0.10, 0.30$  and 0.50) (Fig. 1b). Here, a fast initial  $Fe^{2+}$  release was observed for  $R_{ini} = 0.30$  then, the reaction slowed down when  $R_{eff}$ approached 0.10, whereas the process was slow from the beginning in the experiment for R<sub>ini</sub> = 0.10. As previously suggested in the literature<sup>37</sup> the initial release of Fe(II) from the near surface region may be fast due to electron transfer from the interior of the particle to the surface. This, however, causes a charge imbalance inside the particle, which must be compensated by the movement of Fe-ions and concomitant generation of octahedral vacancies. Recent diffraction investigations<sup>22</sup> demonstrated oxidation-induced strain in magnetite particles, an effect that likely correlates with this interpretation. This process slows down further with time, also in agreement with an alternative conceptual model, where a Fe(II)-depleted outer rim grows thicker while the core shrinks.<sup>19,23</sup>

#### 3.2 Characterization of the reaction products of magnetite with H<sup>+</sup>, H<sub>2</sub>O<sub>2</sub> or Fe<sup>2+</sup>.

As previously shown for the same magnetite batch,<sup>27</sup> the particle size and shape did not evolve much between the pristine stoichiometric magnetite maintained at pH = 8 during 20 days ( $R_{ini} = 0.50$ ; Fig. 2a), oxidized by  $H_2O_2$  (e.g.  $R_{ini} = 0.10$ ; Fig. 2c) and subsequently recharged with dissolved Fe<sup>2+</sup> (Fig. 2d). Maintaining pristine stoichiometric magnetite at pH = 4 during 20 days ( $R_{eff} = 0.13$ ; Fig. 2b) also gave similar particle size and shape (see Table S2). The structure of these nanomagnetites were probed by synchrotron X-ray magnetic

circular dichroism (XMCD) spectroscopy, which enables to distinguish  $Fe^{2+}_{Oh}$ ,  $Fe^{3+}_{Td}$  and  $Fe^{3+}_{Oh}$  according to their distinct characteristic peaks (Fig. 3).<sup>38</sup> The intensity of the  $Fe^{2+}_{Oh}$  peak relative to that of  $Fe^{3+}_{Td}$  and  $Fe^{3+}_{Oh}$  decreases with decreasing pH or after oxidation of the pristine magnetite with H<sub>2</sub>O<sub>2</sub> at pH 8. XMCD spectra of magnetite exposed to H<sub>2</sub>O<sub>2</sub> and showing  $R_{ini} = 0.10$  and 0.30 are quite similar to those of stoichiometric magnetite equilibrated at pH 4 and 5.5 during 20 days. As observed in previous studies, a complete  $Fe^{2+}$  uptake proceeds within less than 1 day when equilibrating a non-stoichiometric magnetite ( $R_{ini} = 0.10$ ) at pH = 7 with an appropriate amount of dissolved  $Fe^{2+}$  to fully restore  $R_{eff} = 0.50.^{7,15,31}$  Although  $Fe^{2+}$  uptake by oxidized magnetite leads to surface recrystallization,<sup>27</sup> it was previously shown that nanomagnetites experiencing oxidation and subsequent  $Fe^{2+}$  recharge reacted similarly to the pristine nanomagnetite concerning contaminant reduction<sup>7</sup> and surface complexation.<sup>15</sup> All these observations suggest that the release-recharge process of  $Fe^{2+}$  from magnetite is reversible.

#### 3.3 Combined effects of pH and oxidation on magnetite stoichiometry.

The apparent establishment of an equilibrium at  $pH \ge 5.5$  after one day (which might be generalized to  $pH \ge 5$  after 20 days)<sup>39</sup> and the similar reaction products of magnetite with H<sup>+</sup>, H<sub>2</sub>O<sub>2</sub> or Fe<sup>2+</sup> further motivated investigation of the magnetite phase-stability at chemical equilibrium. Stoichiometric nanomagnetite suspensions (0.5 g L<sup>-1</sup>) were equilibrated at different pH values during 20 days (Fig. 4a, in black). In agreement with equation 2, the concentration of Fe(II) released in solution increases with decreasing pH, and R<sub>eff</sub> decreases from 0.50 at pH = 7 to 0.13 at pH = 5.3. The XMCD spectra were treated using a linear combination fit involving magnetite and maghemite references<sup>27</sup> to calculate R<sub>eff</sub>. Spectroscopic results were found in excellent agreement with wet chemical analysis (large diamonds in Fig. 4a). These results further confirmed that magnetite nanoparticles cannot be considered as stoichiometric after contact with aqueous solution under acidic conditions, even at moderate acidity (5.3 < pH < 7). Attempts to maintain magnetite stoichiometry at a value of 0.50 by supersaturating the system failed:

(i) working with 5 times more concentrated magnetite suspensions (2.5 g L<sup>-1</sup>), similar R<sub>eff</sub> values are obtained although about 5 times larger [Fe(II)]<sub>aq</sub> are measured (Fig. 4a, in red). If this result is unexpected for a pure phase, it supports the hypothesis that magnetite-maghemite system should be treated as a solid solution. In the first case, [Fe(II)]<sub>aq</sub> is not supposed to increase with the solid content because the activity

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coefficient of a pure phase equals one. Whereas, the activity coefficients or solid solution differs from one, which allows [Fe(II)]<sub>aq</sub> to increase with the solid content.<sup>40</sup>

(ii) addition of an excess of dissolved Fe(II) (0.25-1 mM) to a stoichiometric magnetite does not significantly affect R<sub>eff</sub> at pH < 7 (Fig. S1). In fact, much larger [Fe(II)]<sub>aq</sub> would be required to shift equilibria, but it could also lead to Fe(II)-(hydr)oxide precipitation, as previously observed.<sup>27</sup>

In addition, attempts to coat the surface with an organic ligand such as citrate (1 mM), often used as surfactant to increase the colloidal stability of magnetite, to protect it from oxidation or to reduce its cytotoxicity,<sup>14,41</sup> did not inhibit Fe<sup>2+</sup> release. Although citrate adsorption occurred, Fe<sup>2+</sup> complexation by the remaining fraction of citrate in solution leads to a slightly enhanced Fe<sup>2+</sup> release (Fig. S2). One of the most important conclusions from these experiments is that in all the experiments at pH < 7, suspensions are composed of non-stoichiometric magnetites nanoparticles in solutions containing significant concentrations of dissolved Fe(II).

We further investigated the behavior of magnetites exhibiting different initial stoichiometries ( $R_{ini} = 0.10$ , 0.20, 0.30, 0.40 and 0.50) obtained from the original stoichiometric magnetite sample by addition of H<sub>2</sub>O<sub>2</sub>. As for  $R_{ini} = 0.50$ , H<sup>+</sup>-promoted Fe<sup>2+</sup> release leads to a decrease of  $R_{eff}$  with decreasing pH (Fig. 4b). Interestingly,  $R_{eff}$  values of all initial magnetites follow the same trend, once pH falls below a certain value (e.g. at pH < 5.5,  $R_{eff}$  values are equal for magnetites with  $R_{ini} = 0.30$ , 0.40 and 0.50), evidencing the strong relationship between  $R_{eff}$ , pH, and aqueous Fe<sup>2+</sup> concentrations. This also suggests fully reversible release/uptake of Fe(II) by magnetite, which was further confirmed over a large pH range by adding Fe<sup>2+</sup> to a suspension of non-stoichiometric magnetite ( $R_{ini} = 0.10$ ). Consistently, final  $R_{eff}$  values did not significantly differ from the original stoichiometric magnetite (Fig. 4b).

#### 3.4 Chemical thermodynamic modeling of the magnetite-maghemite system.

If variations of magnetite stoichiometry with pH and redox conditions cannot be avoided, quantitative prediction of this phenomenon becomes crucial. For this purpose, we consider non-stoichiometric magnetite as a binary solid-solution between maghemite and magnetite. The Gibbs free energies of formation of the solid-solution ( $\Delta G_{ss,nano}$ ) of magnetite (mt)-maghemite (mm) nanoparticles is expressed by:

$$\Delta G_{\rm ss,nano} = X \Delta G_{\rm mt,nano} + (1 - X) \Delta G_{\rm mm,nano} + \Delta G_{\rm mix}$$
<sup>(4)</sup>

Page 11 of 21

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58 59 60 The excess free energy of mixing ( $\Delta G_{mix}$ ) can be expressed according to Guggenheim's expansion series:

$$\Delta G_{\text{mix}} = X(1-X)RT(a_0 + a_1(2X-1) + \dots)$$
(5)

Where a<sub>0</sub>, a<sub>1</sub>, etc. are adjustable non-dimensional Guggenheim parameters.<sup>40</sup> However, in order to describe the thermodynamic equilibrium between a magnetite-maghemite solidsolution and an aqueous solution with the Guggenheim equation, it is necessary to ensure that (i) one ion on estructural position is exchanged in the mixing process and (ii) the number of anions (i.e. O<sup>2-</sup>) is the same in the stoichiometric formula of both end-members. As this is not the case in equation 2, the magnetite oxidation reaction was reformulated as:

$$[Fe_{3}^{2+} Fe_{3}^{3+}]_{Oh} Fe_{3 Td}^{3+}O_{12}^{2-} + 2 H^{+} \rightleftharpoons [Fe_{5}^{3+} \Box]_{Oh} Fe_{3}^{3+}T_{d}O_{12}^{2-} + Fe^{2+} + H_{2}$$
(6)

It appears that equation 6 does not only assume a H<sup>+</sup>-promoted dissolution process, but it combines it with a redox reaction, as an oxidative dissolution process.<sup>22</sup> Note that this reaction could also be formulated differently, e.g. by adding 0.5 O<sub>2</sub> to the left side, to form 1 H<sub>2</sub>O on the right-hand side.

Tabulated Gibbs free energies of the bulk iron oxides ( $\Delta G_{mm,bulk}$  and  $\Delta G_{mt,bulk}$ , Table S1) were used.<sup>42</sup> Additionally, surface free energies (referring to hydrated surfaces,  $\gamma_{mm}$  and  $\gamma_{mt}$ in J m<sup>-2</sup>) of the end-member phases were adjusted to determine the  $\Delta G$  of the maghemite and magnetite nanoparticles ( $\Delta G_{mm,nano}$  and  $\Delta G_{mt,nano}$ ), in order to precisely match the experimental solubility data.<sup>43</sup> Because particle size determined by TEM (Fig. 2a-d) and crystallite size previously determined by XRD<sup>27</sup> did not significantly vary in all our experiments, the surface area was assumed to remain constant in the calculations. Previous attempts to measure the surface area of 10 nm-size nanomagnetites were in excellent agreement with surface areas calculated from TEM average size.<sup>15,28,29</sup> Such observation was also made for other minerals such as calcite.<sup>44</sup> Presently, the surface area was calculated equal to 101 m<sup>2</sup> g<sup>-1</sup>, using the average particle size determined by TEM (10.3 nm) and the assumption that all magnetite particles are spherical in shape. The large Fe<sup>2+</sup> solubility dataset produced in our study at various pH and Rini was used to calibrate this model, discarding the results obtained at pH < 5 that are supposed to not reflect chemical equilibrium. The fitting exercise using the aqueous chemical speciation code Phreeplot (PHREEQC-Phreeplot coupling)<sup>35,36</sup> gave  $\gamma_{mt} = 0.52 \pm 0.10$  J m<sup>-2</sup> and  $\gamma_{mm} = 0.57 \pm 0.29$  J m<sup>-2</sup> consistent with literature data  $(0.79 \pm 0.28 \text{ and } 0.57 \pm 0.10 \text{ J m}^{-2}$ , respectively),<sup>45</sup> and we found that a single

Guggenheim parameter was sufficient to obtain a satisfactory match between model and data (Fig. 4). A negative value of  $a_0 = -5.49 \pm 0.50$  confirms the negative  $\Delta G_{mix}$  previously reported from oxide melt solution calorimetry for 100-300 nm sized magnetites.<sup>3</sup> It demonstrates that non-stoichiometric magnetites are a very special type of solid-solution, in comparison to many other solid-solutions exhibiting positive  $\Delta G_{mix}$ .<sup>40,46</sup> This may be explained by the fact that mixing of magnetite and maghemite involves generation of vacancies and electron transfer processes, rather than mixing of ionic species on a given crystallographic position.

#### **3.5** Prediction of redox potential in magnetite suspensions.

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In environmental systems, magnetite plays an important role on biogeochemical cycling of redox sensitive contaminants and, hence, their fate, bioavailability and toxicity. For instance, magnetite can (i) immobilize uranium by reducing the rather mobile U(VI) to less mobile  $U(IV)^8$ , (ii) reduce Cr(VI) or nitrobenzene to less toxic Cr(III) or aniline<sup>7,10</sup> or, conversely, (iii) reduce As(V) to a more toxic As(III)<sup>9</sup>. Therefore, the redox potential ( $E_{H}$ , reported versus the standard hydrogen electrode) in magnetite containing environmental systems is essential to predict the behavior and fate of contaminants. For this reason, E<sub>H</sub> values were experimentally recorded in all experiments using a Pt electrode and our model was used to predict the corresponding  $E_H$  values. At pH < 7, when pH decreases, all nanomagnetites tend to exhibit the same stoichiometry (Fig. 4b), which leads to similar calculated E<sub>H</sub> values (Fig. 5a), in agreement with experimental results. At these conditions, accurate  $E_H$  might be achieved thanks to the large amount of dissolved  $Fe^{2+}$  and the small size of the nanoparticles that could react with the electrode.<sup>47</sup> At pH > 7, the model generally predicts lower  $E_H$  values than measured (open symbols in Fig. 5a). At these conditions  $Fe^{2+}$ is retained in the solid phase and experimental E<sub>H</sub> determination is less reliable.<sup>48</sup> Nevertheless, calculated  $E_H$ , from +30 mV ( $R_{ini} = 0.10$ ) to -154 mV ( $R_{ini} = 0.50$ ) at pH = 7.2, compared rather well with expected E<sub>H</sub> values according to the U(VI)/U(IV) redox chemistry.<sup>8</sup> This further supports the reliability of the present model. Therefore, our results allow proposing a new pH-E<sub>H</sub> predominance diagram of Fe in the magnetite-maghemite-H<sub>2</sub>O system ( $[Fe_3O_4] = 0.5 \text{ g L}^{-1}$ , 10 nm particles, in 10 mM NaCl) accounting for the formation of the solid-solution (Fig. 5b). The predominance field of quasi-stoichiometric magnetite  $(R_{eff} > 0.495)$  is rather narrow in the simulated conditions, restricted to circumneutral to alkaline pH (> 6-7 depending on the E<sub>H</sub>).

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#### 4 Conclusion

This study investigates the solubility and the stoichiometry of 10 nm-sized magnetite under a wide range of environmentally relevant pH (4-9) and redox conditions. It demonstrates that nanosized magnetite particles are in fact not stable at pH < 7, even in the absence of O<sub>2</sub>. Magnetite-maghemite solid solutions form due to the H<sup>+</sup>-promoted dissolution of Fe(II) from magnetite. Our long-term kinetic investigations and XMCD measurements of the magnetite stoichiometry reveal two stages concerning Fe(II) release from magnetite. The first one is rather fast for 10 nm-sized particles and proceeds within less than 24h, if the pH is moderately acidic (pH  $\geq$  5) and the stoichiometry of the final magnetite-maghemite solid solution exceeds 0.1. The Fe(II) release process becomes much slower for R<sub>eff</sub> < 0.1, and even slows down further as the Fe(II)-depleted outer rim grows thicker while the core shrinks.<sup>21,39</sup> Accordingly, we developed the first equilibrium model that could predict the stoichiometry of nanomagnetites exposed to various pH conditions, partially oxidized by H<sub>2</sub>O<sub>2</sub> and subsequently recharged by dissolved Fe(II). The model was also shown to provide estimations of the E<sub>H</sub> consistent with experimental results.

Because magnetite nanoparticles play a key role in the biogeochemical cycling of various organic and inorganic pollutants, affecting the mobility, redox transformation, and toxicity of various organic and inorganic pollutants, this model could be a powerful decision support tool for environmental remediation using magnetite nanoparticles, or to unravel their complex behavior in natural systems. This model could also be used for tuning the key properties of magnetite nanoparticles, and thus provide theoretical guidance for analyses and design of environmental, catalytic and biomedical reactions.

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## **Figures and Tables**

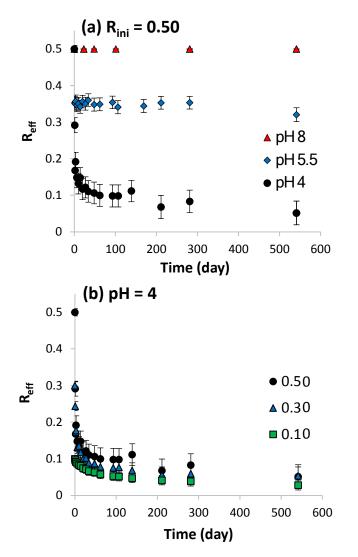
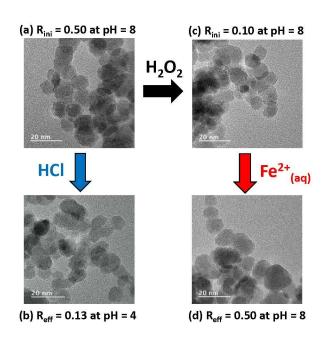


Figure 1. Long term kinetic experiments. (a) Transformation kinetics of stoichiometric magnetite ( $R_{ini} = 0.50$ ) to magnetite-maghemite solid-solution at different pH values (4, 5.5 and 8). (b) Impact of magnetite initial stoichiometry ( $R_{ini} = 0.10, 0.30$  and 0.50) on its transformation kinetics at pH 4.



# Figure 2. TEM characterization of magnetite reaction products with H<sup>+</sup>, H<sub>2</sub>O<sub>2</sub> or Fe<sup>2+</sup>. (a-d) TEM images of the pristine magnetite ( $R_{ini} = 0.50$ ) equilibrated at pH = 8 (a) or pH = 4 ( $R_{eff} = 0.13$ ; b) during 20 days, or oxidized with H<sub>2</sub>O<sub>2</sub> ( $R_{ini} = 0.10$ ; c) followed by the addition of Fe<sup>2+</sup><sub>(aq)</sub> ( $R_{eff} = 0.50$ ; d).

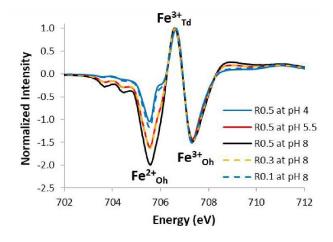


Figure 3. Normalized XMCD spectra of stoichiometric magnetite (R0.5) equilibrated at different pH values (4, 5.5 and 8) during 20 days as compared with stoichiometric magnemite oxidized by H<sub>2</sub>O<sub>2</sub> (R0.1 and R0.3) at pH 8.

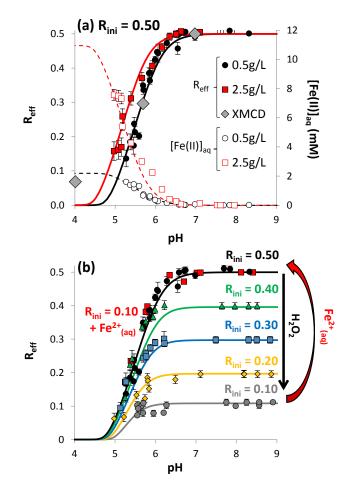


Figure 4. Magnetite stoichiometry under the influence of pH and redox conditions. (a) Effect of pH (4-9) and magnetite concentration (0.5 and 2.5 g L<sup>-1</sup>) on Fe(II) solubility (open symbols) and corresponding magnetite effective stoichiometry (for R<sub>ini</sub> = 0.50; full symbols). Large diamonds correspond to R<sub>eff</sub> determined by X-ray magnetism circular dichroism (XMCD). (b) Effect of pH on R<sub>eff</sub> for magnetites exhibiting different initial stoichiometries ( $0.10 \le R_{ini} \le 0.50$ ), and of a non-stoichiometric magnetite ( $R_{ini} = 0.10$ ) to which dissolved Fe(II) was added in order to reach an overall (solid+solution) Fe(II)/Fe(III) ratio of 0.50. Lines are modeling results, with the same color as the corresponding experimental data.

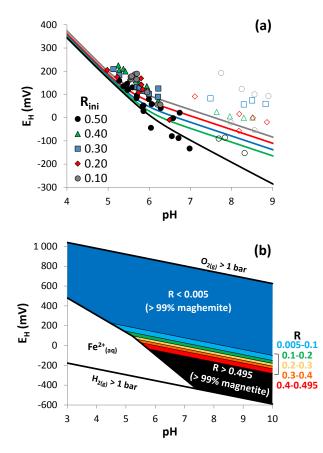


Figure 5. Relationships between pH,  $E_H$  and magnetite stoichiometry. (a) Effect of pH on measured and calculated redox potential in nanomagnetite suspensions ( $0.10 \le R_{ini} \le 0.50$ ). (b) pH- $E_H$  predominance diagram of Fe in the magnetite-maghemite-H<sub>2</sub>O system (10-nm-sized nanoparticles, [Fe<sub>3</sub>O<sub>4</sub>] = 0.5 g L<sup>-1</sup> corresponding to 6.5 mM Fe, in 10 mM

NaCl).