

Investigation of magnetic field signals during vortex-induced pressure drops at InSight

Shea N. Thorne, Catherine L. Johnson, Anna Mittelholz, Benoît Langlais, Ralph Lorenz, Naomi Murdoch, Aymeric Spiga, Suzanne E. Smrekar, W. Bruce Banerdt

► To cite this version:

Shea N. Thorne, Catherine L. Johnson, Anna Mittelholz, Benoît Langlais, Ralph Lorenz, et al.. Investigation of magnetic field signals during vortex-induced pressure drops at InSight. Planetary and Space Science, 2022, 217, pp.105487. 10.1016/j.pss.2022.105487. insu-03682334

HAL Id: insu-03682334 https://insu.hal.science/insu-03682334

Submitted on 14 Oct 2022

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers. L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

1 Investigation of Magnetic Field Signals During Vortex-Induced Pressure Drops at InSight

- 23 List of Authors:
- 4 Shea N. Thorne¹, Catherine L. Johnson^{1,2}, Anna Mittelholz³, Benoit Langlais⁴, Ralph Lorenz⁵,
- 5 Naomi Murdoch⁶, Aymeric Spiga^{7,8}, Suzanne E. Smrekar⁹, W. Bruce Banerdt⁹
- 6
- 7 Affiliations:
- 8 ¹Department of Earth, Ocean and Atmospheric Sciences, The University of British Columbia,
- 9 Vancouver, Canada.
- ¹⁰ ²Planetary Science Institute, Tucson, Arizona, USA.
- ³Department of Earth and Planetary Sciences, Harvard University, Cambridge, MA, USA.
- ⁴ Laboratoire de Planétologie et Géosciences, UMR 6112, Nantes Université, Université
- 13 d'Angers, Le Mans Université, CNRS, Nantes, France.
- ⁵Johns Hopkins Applied Physics Laboratory, Laurel, MD, USA.
- ⁶ISAE-SUPAERO, Université de Toulouse, Toulouse, France.
- ⁷ Laboratoire de Météorologie Dynamique/Institut Pierre-Simon Laplace, Sorbonne Université,
- 17 Centre National de la Recherche Scientifique, École Polytechnique, École Normale Supérieure,
- 18 Campus Pierre et Marie Curie BC99, Paris, France.
- ⁸Institut Universitaire de France, Paris, France.
- 20 ⁹Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, USA
- 21
- 22 Corresponding author: Shea Thorne (sthorne@eoas.ubc.ca, 302 Earth and Oceans Science
- 23 6339 Stores Road V6T 1Z4 Vancouver, BC, Canada)
- 24
- 25 Abstract
- 26 The NASA InSight lander has recorded many pressure drops attributed to convective vortices
- 27 during its first full year of data collection. However, although dust-carrying vortices (dust devils)
- are a common phenomenon on Mars, they have not been observed in InSight images. On Earth,
- 29 magnetic signals associated with some dust devils have been reported. Data from the InSight
- 30 Fluxgate Magnetometer (IFG) provide the first opportunity for similar investigations on Mars.
- 31 Here, we evaluate whether magnetic signals are associated with daytime vortices. We
- 32 incorporate observations of environmental conditions, measurements of ground tilt from seismic
- data, and data from the lander's solar panels, and consider the potential for dust-laden vortices to
- 34 generate observable magnetic field signals. We find that 7.7% of pressure drop events greater
- 35 than 1 Pa show a resolvable magnetic field signal at the time of the pressure drops. The
- 36 resolvable magnetic signals, typically seen on the horizontal field components, are less than 1 nT
- in amplitude, and have no clear correlation with local time, duration, or pressure drop magnitude.
 During nine pressure drop events we found smoothly varying magnetic signals of at least 0.3 nT
- 39 on any one component. To investigate the origin of these magnetic signals we evaluated three
- 40 possible sources: solar array currents, ground and lander tilt, and triboelectric effects of lofted
- 41 dust. We find that SAC and tilt could contribute a change in the magnetic field but cannot solely
- 42 explain the observed signals. The observed changes in field strength could theoretically be
- 43 produced via triboelectric effects, but only in the case of exceptionally large dust devils that pass
- 44 close to the lander. The lack of imaged dust devils and the small number of observed magnetic
- 45 signatures despite numerous measured pressure drops is consistent with at most a small

46 proportion of dust laden convective vortices at InSight and associated predicted triboelectric

- 47 effects.
- 48

49 Keywords: Dust devils, Convective vortices, Magnetic Fields, Mars, InSight, Triboelectricity

- 50
- 51 1 Introduction

52 The Interior Exploration using Seismic Investigations, Geodesy and Heat Transport (InSight) 53 Mars Lander touched down on November 26, 2018 in Elysium Planitia. It carries the Auxiliary 54 Payload Sensor Suite (APSS) that includes the InSight FluxGate magnetometer (IFG). The IFG is 55 the first magnetometer on the surface of Mars, and thus provides a unique opportunity to study 56 the static crustal and time-varying external magnetic fields below the martian ionosphere 57 (Banfield et al., 2018). To date, the crustal magnetic field strength, ~2000 nT, and direction as 58 well as diurnal magnetic field variations measured by the IFG have been investigated (Banerdt et 59 al., 2020; Johnson et al., 2020; Mittelholz et al., 2020). However, the IFG has also measured 60 aperiodic magnetic signals that occur over timescales of seconds to minutes. Here we use IFG data, together with observations of environmental conditions around the lander by APSS 61 62 (Banfield et al., 2018), measurements of ground tilt from seismic data, and data from the 63 InSight's solar panels, to evaluate whether magnetic signals are associated with daytime 64 convective vortices. Of specific interest is whether potentially dust-laden convective vortices, 65 i.e., dust devils, generate observable magnetic field signals at the InSight landing site. 66 Dust devils (DDs) were first observed on Mars in Viking Orbiter images and have been subsequently imaged by many missions, either via direct imaging or by the tracks left from the 67 68 passage of DDs (Balme & Greeley, 2006; Thomas & Gierasch, 1985). Martian DDs provide an 69 important mechanism for dust injection into the atmosphere (Kahre, Murphy, & Haberle, 2006) 70 as well as for the vertical transport of heat (Renno et al. 2004), and have been proposed to 71 maintain current flow among atmospheric regions (Harrison et al., 2016). DDs also contribute to 72 global atmospheric conditions via their effects on convective circulation and a multi-year study of 73 martian DD heights found that they correlate seasonally with the thickness of the planetary 74 boundary layer, a descriptive measure of the vigor of convective atmospheric activity (Fenton & 75 Lorenz, 2015). The electrification of DDs could also affect martian habitability as they have been 76 purported to induce production of H_2O_2 which could be responsible for scavenging organic 77 material or reducing the residence time of methane gas in the martian atmosphere (Atreva et al., 78 2006). Although that process has been shown to have limited impact (Kok & Renno, 2009), the 79 electric and magnetic signatures of dusty atmospheric structures need to be better assessed. DDs 80 are specifically important to planetary missions because of the possibility for dust deposition on, 81 or removal from, solar panels that in turn can shorten or extend mission operations, respectively 82 (Lorenz & Reiss, 2015), and their electromagnetic noise has been hypothesized to affect 83 communication equipment (Harrison et al., 2016). 84 85 October 2020 marked a full martian year of data collection by the pressure sensor instrument on

86 InSight, providing a set of 11,251 pressure drops which were recently compiled (Spiga et al.,

87 2021; Chatain, Spiga, et. al. 2021). The InSight landing site is the most active site for vortices of 88 all landed martian missions to date (Banerdt et al., 2020). Convective vortices are classified as

89 DDs if they loft dust, but the fraction of DDs to dustless vortices is not well understood (Lorenz

90 & Jackson, 2015; Spiga et al., 2016). To date InSight has not imaged any DDs despite targeted

91 surveys (Banfield et al., 2020; Lorenz, Spiga, et al., 2021) but DD tracks around the lander have 92 been observed via satellite imagery (Perrin et al., 2020) and by InSight's cameras (Charalambous

93 et al., 2021). A recent study of parameters inferred for a subset of these vortices together with

94 advective wind speeds suggests that high wind speeds at the InSight landing site may suppress 95 the formation of DDs (Jackson et al., 2021).

96

97 The uplift of a passing vortex is known to also affect seismic measurements (Lorenz et al.,

98 2015). InSight's Seismic Experiment for Interior Structure (SEIS) instrument and the pressure

99 sensor have observed simultaneous seismic and pressure signals due to convective vortices

100 (Lognonné et al., 2020, Murdoch et al., 2021). One such observation also corresponds to a

101 newly-formed DD track that came within 19 m of the lander, and the track was imaged from the

102 lander's camera and from orbit (Banerdt et al., 2020). These joint observations permitted

103 estimates of the vortex properties (e.g., diameter), as well as the elastic rigidity (compliance) of

- 104 the ground near the landing site.
- 105

106 On Earth, there have been two observations of magnetic signals from dust carrying vortices in

107 the form of ultra-low frequency electromagnetic emission in conjunction with the passage of a

108 DD, at distances up to 100's of meters from the electromagnetic sensors (Farrell et al., 2004;

109 Houser, Farrell, & Metzger, 2003). The first such DD, reported by Houser et al. (2003) was

110 approximately 10 meters wide and 200 meters tall with continuous broadband electromagnetic

111 emission at frequencies below 50 Hz with peaks near 10 Hz, while the second reported by Farrell

et al., 2004 again detected magnetic emission between 1-10 Hz (see further Figures 3 and 4 in 112

Farrell et al., 2004), reporting a magnetic field of ~ 0.1 nT immediately outside the DD. One 113

114 possible origin for the magnetic field signals, discussed in the terrestrial literature is

115 cyclostrophic motion of triboelectrically charged lofted dust, behaving as a magnetic solenoid

(Farrell et al., 2006; Kurgansky, Baez, & Ovalle, 2007). 116

117

118 In this study we look for possible DD signatures in the magnetic field measurements recorded by

119 InSight. Assuming they can be tied to other meteorological information, they could help

120 constrain the subset of the convective vortices that are DDs (Houser et al., 2003). A few pressure

121 drop events have been found to exhibit magnetic signals; however the origin of these was unclear 122

(Charalambous et al., 2021; Johnson et al., 2020). A passing vortex, if a DD, could be observed 123

in IFG data via one or more of the following possible mechanisms: (1) Triboelectrically charged

124 lofted dust, as suggested for terrestrial DD magnetic emission (Farrell et al., 2006; Kurgansky, 125

Baez, & Ovalle, 2007), (2) Changing solar array currents as a result of entrained dust of a 126

passing DD obscuring the solar arrays (as observed at InSight during transits of Phobos (Stähler 127 et al., 2020), (3) Ground tilting resulting from the pressure lows at the center of convective

vortices (Lognonné et al., 2020, Murdoch et al., 2021) that changes the orientation of the IFG in 128

129 the ambient crustal magnetic field, and (4) similarly, wind-induced tilt of the lander as explored

130 previously for the seismometer on the lander deck (Murdoch et al, 2017, Panning et al., 2020).

131 The goal is therefore to establish, if, when, and why pressure drops at the InSight lander have

132 associated magnetic signals and in particular to evaluate whether any such signals can be

133 attributed to DDs. Here, we first identify the fraction of pressure drop events that have a

134 resolvable magnetic signature. These resolvable events and their characteristics are then used to

135 evaluate the four mechanisms through which a vortex or DD could create a magnetic signal.

136

- 137 2 Resolvable Magnetic Signatures at the Times of Pressure Drops
- 138 2.1 IFG Data during pressure drops
- 139 From the set of 11,251 pressure drops over the first martian year of InSight data collection
- 140 (Spiga et al., 2021; Chatain, Spiga, et. al. 2021), we selected the 1061 events that exhibit
- 141 pressure drops of at least 1 Pa (Figure 1). This excludes small events unlikely to loft dust
- 142 (Lorenz, Spiga, et al., 2021).
- 143
- 144



145

Figure 1. 1061 pressure drops greater than 1 Pa recorded by InSight during the first full Martian
year of observations. At sol 189 the IFG sampling rate for continuous data increased from 0.2 Hz
to 2 Hz (dark gray to light gray background transition) and at sol 426 the solar array current data

148 10.2 112 (dark gray to light gray background transmon) and a soi 420 the solar array current and 149 cadence increased from 0.0167 Hz to 0.25 Hz when recorded (light gray background to white

- background). All events (gray), events with either continuous 2 Hz IFG data or 20 Hz data down-
- sampled to 2 Hz (blue, N = 862), events with a standard deviation signal to noise ratio (SNR_{σ}) >=
- 152 2 (pink, N = 66), and the subset (N = 8) of $SNR_{\sigma} >= 2$ events with 0.25 Hz solar array current data
- 153 *(black stars).*
- 154

155 The continuous sampling rate of the IFG was initially 0.2 Hz, and on sol 189 it was increased to

156 2 Hz. For short timeframes of interest 20 Hz data can be requested and these are available for

- 157 some of the observed pressure drop times. Intermittent outages of the data acquisition electronics
- 158 for IFG (the Payload Auxiliary Electronics (PAE)) resulted in data gaps in the IFG and other
- 159 meteorological sensor data sets. For one 10-day interval, sols 261 269, the continuous sampling
- 160 rate was decreased again to 0.2 Hz because of downlink limitations just before solar conjunction
- and a spacecraft anomaly (sols 270 283) led to a subsequent data gap until communication with
- the spacecraft could be resumed. Further PAE anomalies occurred but were mostly addressed
- 163 within a time frame of a few hours.
- 164

- 165 We initially used 0.2 Hz data (recorded or decimated from higher frequency data) to assess the
- 166 overall shape and structure of magnetic signals at the times of the pressure drops (e.g., Figure 2),
- but found they lacked sufficient temporal resolution to characterize the magnetic field in detail
- during pressure drops. Therefore, we retained 879 events for which either 2 Hz or 20 Hz data
- 169 (V6; Joy et al., 2019) were available, and the 20 Hz data were down-sampled to 2 Hz to create a 170 consistent data set, using the same approach applied to the IFG data set (Joy et al., 2019). We
- inspected these 879 events by hand to identify any magnetic signals of artificial origin, i.e.,
- 172 lander activities, at the times of pressure drops. Such activities include the lander wake on/off
- 173 commands, communications, or movement of the lander's robotic arm, all of which generate
- magnetic signals that are not completely removed from the calibrated IFG data (Johnson et al.,
- 175 2020; Mittelholz et al., 2020; Thorne et al., 2020). 17 events had magnetic field signals with
- 176 clear artificial origins and were discarded, resulting in a final set of 862 events. Magnetic field
- 177 components are reported in the InSight lander level coordinate system in which X points north,
- 178 Y points east and Z points down (Johnson et al., 2020; Joy et al., 2019).
- 179
- 180 2.2 Event Identification and Resulting Statistics
- 181 Mars exhibits diurnal magnetic field changes as well as shorter period variations throughout a sol
- 182 (Mittelholz et al., 2020). As an example, we show the magnetic field for sols 529 539 (Figure
- 183 2); on one of these sols (534) a pressure drop with a resolvable magnetic field signal is seen
- 184 (Figure 2C, discussed more below). As discussed in Mittelholz et al. (2020), the diurnal cycle
- 185 (Figure 2A) has a peak in the early to mid-morning. This is attributable, at least in part, to wind-
- 186 driven ionospheric currents that in turn depend on atmospheric properties such as horizontal 187 wind speed and electron density (see review in Mittelholz et al., 2020). Increased noise in the
- 187 wind speed and electron density (see review in Mittelholz et al., 2020). Increased noise in the 188 data is clearly seen on all sols between $\sim 11:00 - 16:00$ hrs, and this is inferred to result from
- short time-scale variability in solar array currents (Johnson et al., 2020). Although the example
- 190 event occurs during a comparably "quiet" time (Figure 2A,B), during the event time frame
- 191 (Figure 2C) the background noise is substantial, requiring care in identifying any magnetic
- 192 signature of pressure drop events.
- 193



194 195

Figure 2. IFG magnitude |B|, detrended over a full sol, plotted for 10 sols (529-539) with sol
534 (red), on which a convective vortex studied here occurs. (a) |B| over a full sol. (b) LTST

198 ~11:00-16:00 with the time of event on sol 534 plotted in Figure 4A identified by vertical lines.

199 (c) |B| during the corresponding LTST for which an event was detected on sol 534 (shown in

200 Figure 4A), detrended over the time window shown to highlight the noise. Vertical lines show the

- 201 *duration of the event (28.8 seconds).*
- 202

We first assessed the typical variability in the field over the characteristic timescale of pressure drops at all local times (Figure 3). We compared this with the local magnetic field variability immediately before and after each of the 862 events, as well as with the variability during a pressure drop to assess whether magnetic signatures during the times of pressure drops can be distinguished from background fluctuations.

208

209 Pressure drop durations (twice the full-width-at-half-max, FWHM) were obtained via gaussian

- 210 fitting of a 200 second interval of detrended pressure data centered on the pressure drop peak,
- 211 following the methodology described in Murdoch et al. (2021). To capture the typical
- background variability in the field on the time scale of pressure drops, each sol with 2 Hz data
- 213 first had a 60s running mean removed and was then divided into sections the length of the
- average duration of pressure drop events (21.7 seconds). The standard deviation (σ) in each field
- 215 component for each section was computed, binned into half hour local time bins, and the median
- value of σ in each bin was computed (Figure 3). For pressure drop events, the local background variability in each magnetic field component was quantified by the average σ of the magnetic
- field in windows 1.5 times the duration of the individual event, directly before and after the



each event was characterized by the σ during the pressure drop (σ^{pd}) (the light grey shading in Figure 4).



223



Figure 3. Median standard deviation in each magnetic field component, σ , during (pink) and directly surrounding (green) pressure drop events, as well as the ambient variability in the field over the characteristic time scale of a pressure drop (black), all shown for the first year of the mission and binned in 30 min local true solar time (LTST) bins. The number of pressure drop events per bin is shown by gray bars. Blue stars are plotted below the x-axis at the local time for each of the 66 events with a $SNR_{\sigma} >= 2$ in at least one magnetic field component. Orange stars show the 9 events highlighted in section 2.3.



232 233

Figure 4. A-D show 4 examples of pressure drop events with sol number, local time of pressure

- 235 drop peak, and pressure drop (Pd) in Pa identified by Spiga et al., 2021. Top row: North (B_X -
- Red), East (B_Y Green) and Vertical Down (B_Z Blue) magnetic field components during each pressure drop (light grey shading) and the surrounding background period (darker grey shading,

1.5 times the event duration before and after the event). 2 Hz (colors) and decimated 0.2 Hz (blackline) data after subtraction of a 60s moving average are shown. Bottom row: pressure drop (blue) with gaussian fit (red). A and D have $SNR_{\sigma} > 2$ and $SNR_{P2P} > 1.8$ in the B_Y component, B has SNR_{σ} and SNR_{P2P} above their respective thresholds in all 3 components, and C has $SNR_{\sigma} > 2$ in B_Y.

243

244 Overall, the variability in the ambient magnetic field increases substantially between noon and 245 4pm local true solar time on each sol (Figures 2.3). This is in part related to varying current 246 draws from the solar array panels and other lander activities (Johnson et al., 2020; Mittelholz et 247 al., 2020; Lorenz et al., 2020). The σ immediately before and after an event, as well as during an 248 event, are similar in amplitude to the typical variability in the field. This suggests that, on 249 average, magnetic signals at times of pressure drops are not significantly different from those in 250 the immediate surrounding intervals or the typical background variability at that local time. 251 Furthermore, the typical variability in the vertical magnetic field component is much larger than 252 on the horizontal channels, typical ~0.5 nT between noon and 4pm LTST. This analysis also 253 shows that identification of signals in the IFG data is not trivial due to typical short-term

- variability in the data.
- 255

As we had no a priori expectations as to how a DD might manifest in the IFG data, we used two methods to quantify the signals during individual events. We first determined the signal to noise

ratio (SNR) for each of the 862 events to systematically identify those with a resolvable

259 magnetic signature. The signal for each event, as identified in the pressure data, was detrended

with a 60 second running mean (described above), to remove longer time scale variations unassociated with the pressure drop. We defined the standard deviation signal to noise ratio

262 SNR_{σ} as the ratio $\sigma^{pd} / \sigma^{bk}$. We found that 66 of the 862 events recorded at 2 Hz had a SNR_{σ}

263 greater than or equal to 2 on at least one magnetic field component (Figure 1).

264

We also implemented a selection scheme in which we replaced σ^{pd} and σ^{bk} with the maximum 265 266 excursion (i.e., the range or peak-to-peak amplitude) during and before/after the event 267 respectively and recomputed the SNR using these values, termed the SNR_{P2P}. We found that this method systematically resulted in the detection of fewer high SNR events c.f. the standard 268 269 deviation approach. To compare whether the two approaches identified similar populations of 270 events we first plotted empirical cumulative distribution functions for SNR_{P2P} and SNR_{σ} (for each 271 component). We used these to determine the SNR_{P2P} value that corresponds to the same number 272 of events with resolvable magnetic field signals as identified using a $SNR_{\sigma} = 2$. This resulted in a 273 $SNR_{P2P} = 1.8$. We found that 46 of the 66 events (70%) identified using the $SNR_{\sigma} = 2$ were also 274 identified using $SNR_{P2P} = 1.8$ indicating substantial, but not complete, overlap in the populations 275 (see details in Supplementary Table 1 and Catalogue). Although the SNR_{P2P} might seem a more 276 obvious approach, noise in the IFG data can result in spurious range measurements and is more 277 sensitive to outliers. We thus chose to use the events identified using SNR_{σ} for further analyses. 278 Examples for three events are shown in Figure 4, and all events identified with either the

standard deviation or range approach are catalogued in the Supplemental Information.

280

281 The 66 high SNR_{σ} events correspond to 6.2% of all pressure drops > 1 Pa and 7.7% of the 862

events with IFG data. In detail, we find that for most events only one magnetic field component 282 here SNP = 2. For P = P and P, this assumed for 2.0% (N=2C) 4.1% (N=2C) and 2.7%

283 has a SNR_{σ} >= 2. For B_X , B_Y , and B_Z this occurred for 3.0% (N=26), 4.1% (N=35), and 2.7%

284 (N=23) of the 862 pressure drops respectively (Figure 5). We further observe that resolvable

events have no clear correlation with local time, duration, and pressure drop magnitude (Figure

286 6). The local times of all 66 SNR $_{\sigma}$ events (shown with blue stars in Figure 3) are seen to have a

similar local time distribution to that of the pressure drop events overall.



Figure 5. Histograms of the standard deviation signal-to-noise ratios (SNR_{σ}) for the north (B_X), east (B_Y) and vertical down (B_Z) magnetic field components. The numbers of events with SNR_{σ} >=2 are given. Seven events with SNR_{σ} >4 are not shown (2 in B_X, 2 in B_Y and 3 in B_Z).



293

288

Figure 6. Pressure drop magnitude for events with $SNR_{\sigma} \ge 2$ in the B_X magnetic field component (colored dots), compared with the overall population of pressure drops greater than 1 Pa examined here. Pressure drop versus (a) event duration (2 x FWHM), and (b) local true solar time (LTST). The results for events with $SNR_{\sigma} \ge 2$ in B_X and B_Z are similar.

298

We evaluated whether this occurrence rate of high SNR_{σ} events occurs randomly in the IFG data set as follows. For every sol over the martian year with 2 Hz IFG, we computed a SNR for local time sections of the sol with the length of the average duration of pressure drop events, by dividing the σ during the short section by the average σ of the sections before and after it. We found that for B_X , B_Y , and B_Z , a $SNR_{\sigma} >= 2$ occurred for 2.5%, 2.0%, and 2.3% of the 469808 segments, less than the occurrence rate of a $SNR_{\sigma} >= 2$ during the times of pressure drops. This analysis confirms that our identification of events is statistically significant, and we reject the

306 hypothesis that our identifications are not associated with any events.

307

308 2.3 Magnetic Field Signals During Pressure Drops

- 309 Our event detection method identified 66 pressure drop events, 8 of which have concurrent solar
- 310 array current data (see Supplementary Figure 1 and Section 3.1). These events are reported
- 311 together with figures and their statistics in the supplementary information catalog which also 312 includes the events identified only by the SNR_{P2P} method and all events with pressure drop
- magnitudes >= 3.8 Pa (123 events total as listed in Supplementary Table 1). The set of resolvable
- events (as identified by both selection methods) and large pressure events was visually examined
- for magnetic signals that varied smoothly during the pressure drop duration while also showing a
- 316 clear peak-to-peak signal of at least 0.3 nT on any one component. We found that only 9 events
- 317 matched these criteria. Three such events are displayed in Figure 4 (A, B, & D) and the
- remaining 6 are shown in Figure 7. For these events, the components that show the smoothly
- 319 varying signals > 0.3 nT during the pressure drop are observed mostly B_X and B_Y (consistent with
- the lower noise level on these channels) and have an average peak-to-peak signal of 0.8 nT with a maximum observed range of 2.7 nT (B_Y in Figure 7F).
- 322





324 325

Figure 7. A-F show B_X (red), B_Y (green) and B_Z (blue) for 6 resolvable events with smoothly
 varying magnetic signals during the pressure drop in one or more component featuring a peak-

- 328 to-peak signal > .3 nT, plotted in the same style as top row of Figure 4. To highlight the
- 329 *identified signals, components that did not fit our selection criteria have increased transparency.*
- 330 The panels are titled with sol number, local time of pressure drop peak, and pressure drop (Pd)
- *in Pa as identified by Spiga et al., 2021.*
- 332

- 333 We consider these 9 events the best candidate examples of magnetic field signals during pressure
- drops of the 862 events examined. Figures 7E and 7F show the challenges associated with
- identifying signals in noisy IFG data. Although the remaining events with high SNR may contain

resolvable signals, the signals are either $< \sim 0.1$ nT or they occur during times of noisy magnetic

field measurements as illustrated by Figure 2. As shown in Figure 3, the 9 events cluster in local

time in the early part of the day and before the onset of the typical increase in variability in the ambient magnetic field. This is in contrast to the local times of all 66 SNR_{σ} events which have a

- ambient magnetic field. This is in contrast to the local times of all 66 SNR_{σ} eve similar local time distribution to that of all pressure drop events.
- 341
- 342 3 Discussion

We investigate three possible sources of the resolvable magnetic signals during the times of
pressure drops: solar array currents, ground and lander tilt, and triboelectric effects. We then
synthesize these results drawing on other InSight observations related to dust movement and DD.

- 346
- 347 3.1 Solar Array Currents

348 We investigated whether there is a correlation between changes in the solar array currents and 349 magnetic signals to assess the possibility that a change in current caused by a DD leads to a

350 change in the magnetic field. This could occur if lofted dust carried by DDs directly reduces

- 351 solar incidence on the panels and thus leads to a drop in the current. Alternatively, increases in
- 352 the solar array currents are possible as dust within a DD can scatter sunlight onto the arrays
- 353 (Lorenz & Jackson, 2015). The effect could occur on either or both solar panels, depending on
- the relative orientations of the passing DD, the sun, and the solar panels. The change in solar
- 355 array currents could have an attendant signal in the magnetic field measured by the IFG that 356 together could help discern dust-free vortices from DDs. Drops in solar flux due to obscuration
- of solar cells by DDs have been observed on Earth (Lorenz & Jackson, 2015). At InSight,
- magnetic signals have been observed on Earth (Lorenz & Jackson, 2015). At insight,
- arrays during the Phobos and Deimos transits and the associated temporary darkening (Stähler et
- 360 al., 2020).
- 361

362 The relevant solar array current (SAC) data for determining insolation are reported via 2

- 363 channels, 771 and 791. They indicate insolation on the east and west solar arrays respectively,
- the magnitude of which is relayed through currents from each solar array (Lorenz et al., 2020).
- 365 Solar array currents were recorded at a sampling interval of \sim 30 seconds until Sol 426, when the
- 366 sampling rate was increased to 4 seconds. However, InSight SAC data are only recorded when
- the lander is awake, i.e., not continuously (Lorenz et al., 2020) and long gaps in data coverage on any individual sol are present (see Figure 4 in Mittelholz et al., 2020). Because of the
- 369 any individual sol are present (see Figure 4 in Mitteinoiz et al., 2020). Because of the 369 intermittent SAC data, a model of daily solar array currents generated to fit data gaps is fit to the
- data points as discussed in Joy et al. (2019), and interpolated to the cadence of the IFG data to
- 371 correct IFG data for solar array currents. As a result, low frequency variations in the currents
- drawn by InSight's solar arrays are generally accounted for in the processing of magnetometer
- data, but variations on the time scales of pressure drops are not modeled and could contribute to
- 374 IFG signals (Joy et al., 2019; Mittelholz et al., 2020).
- 375
- To investigate whether the magnetic field at the time of a pressure drop is correlated with solar
- array current signals, we required SAC data throughout the pressure drop and corresponding flanking periods at a high enough ordenee to receive a change in current during the drop. We
- 378 flanking periods at a high enough cadence to resolve a change in current during the drop. We

- thus examined pressure drop events with durations greater than or equal to 8 seconds for which
- 380 0.25 Hz SAC data were available (i.e., on or after Sol 426), resulting in 73 events (see Figure 1).
- 381 This threshold for the pressure drop duration excludes the two DD events identified in the
- 382 Lorenz, Lemmon, & Maki, (2021) SAC survey.
- 383

Of the 73 pressure drop events that fit the SAC data criteria, only eight of them have magnetic

- field components with a $SNR_{\sigma} \ge 2$ and three of the eight also have a SNR_{P2P} that exceeds the
- 386 SNR_{P2P} = 1.8 detection threshold (including Supplementary Figure 1C). Of the eight events, the
- 387 magnetic signal is typically observed on only one of the X, Y, or Z IFG components, and not 388 consistently on a specific component (the first 8 events listed in Supplementary Table 1). IFG
- 389 signals are observed to be present both with and without concurrent SAC changes
- 390 (Supplementary Figure 1). Additionally, pressure drops with low IFG SNR_{σ} (unresolvable) are
- 391 commonly observed with and without SAC dips. We conclude that the SACs cannot be the only
- 392 source of magnetic field signal during pressure drop events.
- 393
- 394 3.2 Ground or Lander Tilt
- 395 3.2.1 Ground Tilt
- 396 Tilting of the InSight lander by ground motion could also be responsible for a change in
- 397 magnetic signal. As a DD or convective vortex is characterized by lower pressure in its center, its
- 398 passage can be observed as the negative load deforms the surface, generating ground tilt
- 399 (Banerdt et al., 2020; Murdoch et al., 2021).
- 400
- 401 The magnetic signal from a tilt of the lander due to ground motion caused by the low-pressure 402 field of a convective vortex can be approximated as a small rotation (α) in the magnetometer 403 orientation about a horizontal axis in the lander's reference frame. This results in a change in the 404 orientation of the IFG with respect to the background crustal magnetic field, producing a change 405 in the individual components, but leaving the magnitude of the overall signal unchanged, i.e. the 406 new field $\mathbf{B}' = \mathbf{B} + \Delta \mathbf{B}$, where $|\mathbf{B}'| = |\mathbf{B}|$ Some geometry and algebra yield an expression for $\Delta B = /$ 407 $\Delta \mathbf{B}/$
- 408

 $\Delta B = B 4 \sin\left(\frac{\alpha}{4}\right) \cos\left(\frac{\alpha}{4}\right)$

409 410

We calculate values for ΔB for ground tilts up to 10⁻² degrees assuming a surface field strength B 411 of 2000 nT at InSight (Johnson et al., 2020; Figure 8). Lorenz et al. (2015) reported terrestrial 412 observations of ground tilts of $\sim 10^{-7}$ ° correlated with long period (>9 second) convective 413 414 vortices identified by pressure drops. This terrestrial value, if applicable at Mars, would result in 415 a ΔB of 0.0002 nT, too small to be resolvable by the IFG. At InSight, seismic observations 416 concurrent with pressure drops have been made because the seismometer is able to measure the 417 resulting ground deformation (Banerdt et al., 2020; Lognonné et al., 2020; Murdoch et al., 2021). 418 Stott et al. (2021) investigated ground tilt during the initial phase of the mission when the 419 seismometer was still on the lander deck. They reported a maximum peak-to-peak tilt of ~0.002° 420 on a three-months timescale, an estimate which captures motion of the regolith and thermoelastic tilting of the lander, both expected to be much larger effects than ground tilts during vortex 421 events. This is supported by maximum ground tilt magnitudes between 10^{-9} ° and 10^{-6} ° for 492 422 423 vortex events (Figure 8 of Murdoch et al., 2021), all of which would have generated magnetic

424 signals too small to be resolvable by the IFG (Figure 8). The ground tilt required to cause a

425 measured magnetic signal of 1 nT is ~0.03°, much larger than the expected or measured ground 426 tilt values, and we therefore exclude this mechanism as the cause for magnetic signatures during 427 pressure drops.

428



429 430

431 *Figure 8.* Expected change in the measured magnetic field strength (ΔB) due to ground tilting of

432 *the lander (rotation around a horizontal direction, Section 3.2.1) and wind-induced lander tilt*

433 (Section 3.2.2). The reported terrestrial ground tilt (order of magnitude) from Lorenz et al.

434 (2015) with the associated predicted ΔB is shown (orange star), as are the 66 predicted wind-

435 induced tilts at the time of high-SNR_{σ} pressure drops (gray) with the maximum (25.0 m/s wind

436 speed; see section 3.2.2) highlighted in green. Histogram shows ground tilts measured by

437 InSight's seismometer (Murdoch et al., 2021).

438

439 3.2.2 Wind-Induced Lander Tilt

Another possibility is that the lander itself is slightly tilted due to wind loads. Considering, for purposes of exposition, only the drag force on the side of the lander, the moment of this force

442 must be reacted against by the forces on the feet. The feet on the downwind side of the lander 443 place a stronger load on the surface than on the upwind side, and if the ground deforms

elastically to provide these different forces, the lander tilts slightly away from the wind. These

- elastic loads on the ground caused by lander drag and lift are a significant source of noise for the
- seismometer (e.g. Murdoch et al., 2017), even on the ground a meter or two from the lander. For
- 447 a deck-mounted instrument, such as the magnetometer, (and in the first days of the landed
- 448 mission the seismometer), the compliance of the deck to lander wind loads in part results from
- elastic elements in the landing gear. The transfer function of wind speed to deck movement was estimated by Panning et al. (2020) to be $2.1 \times 10^{-7} \text{ (m/s)}^2$. We can crudely estimate the
- 450 estimated by Panning et al. (2020) to be 2.1 × 10⁻¹ (m/s) . We can crudely estimate the 451 lander tilt as being this quantity divided by local gravity 3.7 (m/s²) – in practice the lander tilt
- 451 will depend on the wind direction with respect to the lander feet. For our 66 resolvable events,
- 453 the maximum windspeeds during the pressure drops range from 3.6 to 25.0 m/s. Thus, for a

- maximum gust of 26.5 m/s associated with a dust devil, the tilt is $\sim 25.0^2 \times 2.1 \times 10^{-7} / 3.7 = 3.6 \times 10$ 454
- 455 10^{-5} radians or 0.002°. Per the foregoing discussion and computations of ΔB from tilts, such a
- 456 lander tilt can cause an apparent field change of 0.07 nT (Figure 8), and windspeeds would have
- 457 to reach ~95 m/s to cause the observed signals of $\Delta B \sim 1$ nT. Therefore, while wind-induced
- 458 lander tilt can affect IFG measurements, to date, windspeeds are not sufficiently large to be the
- 459 only source of the observed magnetic signals.
- 460
- 461 3.3 Triboelectric Effects
- 462 A magnetic signal could also be generated by swirling charged dust grains entrained in the DD. 463 These grains may become electrically charged via triboelectric effects (contact electrification) in
- 464 which collisions will charge the particles, and a characteristic charge-to-mass ratio develops (Farrell et al., 2004). This electrification, coupled with the differential movement of particles 465
- based on their mass, leads to a magnetic dipole moment and an electric field with an associated 466
- magnetic field emission (see Figure 1 in Jackson & Farrell, 2006). The spiraling motion of these 467
- 468 charged grains in cyclostrophic balance (in which the inward pressure gradient force balances the
- 469 outward centripetal force) can be modeled as current loops or a solenoid (Houser et al., 2003;
- 470 Lorenz, 2016). In this approximation, a DD with fixed parameters would produce a stable 471 magnetic field (see Figure 4 in Houser, Farrell, & Metzger, 2003). However, DDs have
- 472 constantly fluctuating cross-sectional areas, velocities, and inhomogeneous densities which
- 473 would result in dynamic magnetic fields (Farrell et al., 2004; Farrell et al., 2006; Lorenz, 2016).
- 474 Variations on this model have been invoked to explain the two terrestrial measurements of DD
- 475 magnetic signals (Farrell et al., 2004; Kurgansky et al., 2007), the largest of which was ~0.1 nT
- 476 (Farrell et al., 2004). Using this approximation, a large dust devil passing with a high density of
- 477 charged dust grains could be observed by InSight as a quasi-steady magnetic field mostly in the
- 478 vertical direction (Bz in the lander level frame), possibly with a horizontal field component
- 479 generated by azimuthal asymmetries, such as multiple vortex cores (Lorenz, 2016).
- 480 481 Recall that no DDs have been detected by InSight's cameras, and only a handful of faint shadows 482 on the solar panels have been detected (Lorenz et al., 2020). Furthermore, Jackson et al. (2021) 483 reported a maximum optical depth of $\tau \sim 0.01$ from InSight's cameras at the times of the dustless 484 vortices. Therefore, a dust-filled vortex would exceed this value of τ . We explore whether an as-485 yet unobserved vortex containing lofted dust could produce signals in the magnetic field. We
- 486 consider the case of a 10 m diameter vortex with a dusty annulus, such that the dust content is
- 487 equivalent to a layer of dust around the perimeter of the vortex that is one particle thick. Thus, the
- 488 effective area of dust is equal to the circumference of the DD multiplied by the diameter of the
- 489 dust grains. Assuming a diameter of 2 μ m for dust grains, with a grain density 2700 kg/m³, the
- 490 mass density per meter height, ρ_M , is ~0.17 kg/m. In reality, the dust is spread over a larger cross-491 sectional area, for an annulus of width 1 m in a 10 m diameter vortex, this would correspond to a
- number density of dust grains of $\sim 3 \times 10^{-12} \text{ m}^{-3}$. We further assume that dust has a charge-to-mass 492
- 493 ratio, Q/m, of 10⁻⁶ C/kg (Mendez Harper et al. 2017), the devil rotates once per second
- 494 (corresponding to a wall circumferential speed, v_w , of 10π m/s, typical of intense vortices (Balme
- 495 & Greeley, 2006) and at the upper limit as reported for InSight's vortex catalog by Lorenz et al.
- 496 (2021)). This corresponds to $\rho_M * Q/m^* v_w \sim 5.3 \ \mu A$ of current loop per meter height. The
- 497 solenoidal vertical magnetic field B inside the devil is then $B = \mu_0 * \rho_M * Q/m^* v_w = 6.7 \times 10^{-3} \text{ nT}$,
- 498 where $\mu_0 = 4\pi 10^{-7}$ Tm/A is the magnetic permeability of free space.
- 499



521 the event on sol 652; however, the simple point source model did not fit the seismic observations

522 well, likely due to the close approach of the vortex.



Figure 9. Trajectory of the 4.0 Pa convective vortex on Sol 534 (Figure 4A). The grey dashed
line modeled trajectory is the straight line fit to the pressure and seismic data inversion results in
blue.

528

529 We substitute the derived diameter of this close approach event into the triboelectric model

- 530 calculation above and the resulting DD could generate a magnetic field of ~0.016 nT. Therefore,
- for this candidate event, the magnetic signal is still an order of magnitude too small to generate
- 532 the observed signal of ~ 0.3 nT.
- 533

534 The following question arises: Can triboelectric effects cause a resolvable signature? To answer 535 this question, we explore the extent to which the parameters above - the DD diameter, wall 536 circumferential speed, v_w , dust grain diameter, and the charge-to-mass ratio, Q/m - could 537 reasonably be varied. The predicted magnetic field scales linearly with each of these parameters 538 and an order of magnitude increase in total could explain the observed signal amplitudes. 539 Jackson et al. (2021) used models of the pressure and wind speed profiles for their catalog of 540 InSight's vortices to infer intrinsic vortex parameters for a subset of events during which the 541 inferred closest approach distance was within a vortex diameter from the lander. While their 542 most extreme value for DD evewall velocity of \sim 37 m/s is similar to our v_w , they report 543 diameters reaching hundreds of meters (a maximum of ~517 m). Clearly this is not the case for 544 the event on sol 534, for which we can estimate the vortex diameter, but other vortices could be 545 larger. Furthermore, our estimate for dust grain diameter (2 μ m) is in line with observations for 546 fine atmospheric dust on Mars (Pollack et al., 1995); however, spectral observations during the 547 2018 dust storm by the Curiosity rover showed dust enhancement by particles ~3 times the 548 typical size (~8 μ m) (Lemmon et al. 2019). Terrestrial measurements of dust devil mean grain 549 diameter values range from $\sim 90-100 \,\mu\text{m}$ decreasing with height (Raack et al. 2018), although the 550 thin Martian atmosphere may poorly suspend larger grains comparatively (Merrison et al. 2004). 551 Dust grain diameters lifted by a large DD, especially closer to the ground, could therefore be ~ 5 552 times larger than we assume. For a constant optical depth, a dust devil comprised of larger grains could allow for a larger signal as their area increases by the square of the grain radius while their 553 554 mass (and therefore charge density) increases by the grain radius cubed. Alternatively, if large grains are not lifted far from the ground, a shallow toroid of swirling sand rather than a canonical 555 556 DD may yield the observed magnetic field signature signals. Terrestrial measurements for Q/mof sand and volcanic ash can reach maximum values of 10⁻⁵ C/kg (Tian-Li et al., 2014), while 557 558 experiments of the Mars dust simulant JSC Mars-1 tribocharged against stainless steel and 559 Teflon have generated Q/m ratios much larger (Sharma et al. 2008). The dust concentration could 560 also be higher. In conclusion, higher values for dust devil diameter, dust grain diameter, or Q/m561 are plausible and could lead to triboelectrically-driven magnetic fields on the order of 0.1 nT 562 (Section 2.3), for a close DD trajectory. 563

563 564

565 3.4 Synthesis

566 We examined several mechanisms that could lead to magnetic signatures during pressure drops.

- 567 Our investigation of solar array current drops at times of resolvable magnetic signatures shows
- that SAC changes are present both with and without magnetic signals (and vice versa) and are
- therefore not the sole contributor to IFG signals. Furthermore, tilt of either the lander or the
- 570 ground resulting in re-orientation of the IFG with respect to the background crustal field can be

571 ruled out because the expected and measured tilts reported at InSight are too small to create

- 572 magnetic signals of resolvable magnitude. Finally, although triboelectric effects are theoretically
- 573 possible, the estimates made in Section 3.3 indicate that these are not likely to frequently create
- 574 measurable magnetic fields unless the DD has a large diameter, a very close approach distance 575 and carries sufficient charged dust. A purely solenoidal (i.e. rotationally-symmetric) field from a
- 576 vertical dust devil would be a quasi-DC signal increasing slowly from zero as the devil
- 577 approached, reaching a maximum and then declining. Therefore, if the slow approach and signal
- 578 generation of a magnetic field-generating vortex was considerably longer than the duration of
- 579 our pressure drop, our resolvable event detection methodology may not positively detect it.
- 580

581 Convective vortices can also be detected by a corresponding rise in temperature with a core 582 temperature rise of up to 10 K, although elevated temperatures are observed if the sensor enters 583 the vortex core (Ringrose et al. 2007). Because IFG measurements are sensitive to temperature 584 fluctuations (Mittelholz et al. 2020, Thorne et al. 2020), IFG sensor temperature data for all 585 events in Figure 4 were investigated, and we found no corresponding temperature signals. In the 586 case of the event on sol 534 for which we have a derived trajectory, this is consistent with a dust 587 devil that passed close to but not directly over the lander. Furthermore, as shown in Joy et al. 588 (2019), if temperature were the cause, then we should see similar relative changes in the 3 589 components from one event to another, whose amplitude just depended on the amplitude of the 590 thermal perturbation. This is not the case as can be seen from Figures 4 and 7, e.g. in Figures 4A 591 and 4B, even the signs of the perturbations to B_X and B_Y are opposite. We thus rule out that

- 592 unmodeled temperature perturbations are responsible for the IFG signals
- 593

594 Although InSight has recorded a large number of pressure drops, there are relatively few with 595 resolvable magnetic signatures. Only 7.7% of pressure drops larger than 1 Pa correlate in time 596 with such magnetic signatures, which is only 0.6% of all recorded pressure drops. Furthermore, 597 given the frequency of pressure drops, it is remarkable that no characteristic DD has been 598 successfully imaged by the ongoing imagery surveys. While DD tracks have been observed 599 around the lander (Perrin et al., 2021), there have been no images of lofted columnar dust at the 600 InSight landing site to date. This is in stark contrast to other martian missions (Lorenz, Spiga, et 601 al., 2021). The dearth of imaged DDs and low proportion of observed magnetic signatures 602 despite numerous measured pressure drops is consistent with a small proportion of convective 603 vortices carrying lofted dust. In most cases with minimal entrained dust solar array currents are 604 largely unaffected with little to no contribution to IFG signals, and triboelectric charging likely 605 does not occur or is minimal.

606

607 The local environment at InSight may be a contributing factor to the lack of dust lifting

608 necessary for DD development. While aeolian activity has been reported at InSight including

- surface creep of particulate and dust removal (Charalambous et al., 2021; Baker et al. 2021), and
- 610 dust deposition occurs, including critically upon the solar panels (Lorenz, Lemmon, & Maki,
- 611 2021), the requirements for saltation leading to DDs may be rarely met. Lorenz, Lemmon, &
- 612 Maki (2021) estimate that the dust lifting pressure drop threshold at InSight is roughly between
- 4 and 10 Pa, though they also acknowledge much higher dust lifting thresholds from other Mars
 missions (6-40 Pa based on data from Pathfinder, Phoenix, and Spirit) and terrestrial
- 615 observations (20-80 Pa). Furthermore, based on the non-detection of DD's by the InSight
- 616 camera, Jackson et al. (2021) calculated a localized dust lifting threshold of ~7 Pa. As shown in

- 617 Figure 1, such a threshold would exclude most of the pressure drop events examined in this
- study, even if the observed pressure drop is a lower limit for the actual central pressure drop. By
- 619 comparing vortex encounters parameters with advective wind speeds Jackson et al. (2021) found
- 620 evidence that high wind speeds at InSight may have led to the suppression of dust lifting for the 621 most vigorous vortices.
- 622
- 623 4 Conclusions

624 We analyzed the magnetic field at InSight at times of pressure drops during the first year of 625 observations to identify if pressure drops might correlate with a magnetic response. We found 626 coincident magnetic field signals, albeit rarely, as 7.7% of data during drops greater than 1 Pa 627 had a magnetic signal in any component. The mechanism that causes these signals is as yet inconclusive, but is not indicative of ground or lander tilting in response to the localized pressure 628 629 drop (Murdoch et al., 2021). While dust laden convective vortices in close approach could 630 explain our resolvable events, especially those discussed in Section 2.3, our analysis indicates 631 that in general the magnetic signals expected from DDs are too small to be detected. In 632 summary, our observations show (a) no observed magnetic field signal for the vast majority of 633 pressure drops and (b) a few pressure drops, (~9, i.e. < 1% of the overall population of pressure 634 drops > 1Pa) with magnetic field signals $< \sim 1$ nT, possibly corresponding to more dust laden 635 vortices that would have had to have passed close to the lander. Modeling of pressure and 636 seismic data for one of these 9 pressure drops on sol 534, suggests a vortex diameter of ~23m, 637 and indicates that nominal parameters for dust grain diameter, charge-to mass density, and a dust 638 content corresponding to $\sim 3 \times 10^{-12}$ particles per m³, underpredicts the observed magnetic field 639 signal by an order of magnitude. The magnetic field signal depends linearly on all these 640 parameters, as well as on the dust-devil diameter and thus if triboelectric effects are indeed the 641 cause of the full signal during these few pressure drop events, larger values for one or more of 642 the dust parameters are required. The remaining population of events that appear to have an 643 enhanced magnetic signal (high SNR events, $\sim 5\%$ of the pressure drops > 1 Pa) compared with 644 the time intervals immediately preceding and after the pressure drop, are more difficult to assess: 645 some of these comprise very small signals (< -0.1 nT) that could potentially be consistent with 646 DD related signatures, however larger signals, especially those > 1 nT would require DD

- 647 parameters outside the limit of current studies
- 648

649 A larger number of events with 2 Hz IFG data and 0.2 Hz solar array current data could 650 strengthen this investigation, as would an unambiguous event detection by InSight's cameras to 651 provide a positive control. Analysis of the coincidence of solar array current signals and IFG 652 signals is limited by the intermittent nature of the SAC data. However, the DD solar panel 653 cleaning events which were anticipated have not yet transpired and the power constraints during 654 the current winter season resulting from the accumulation of dust on the solar panels have led to 655 IFG not collecting continuous data since late 2020. Further martian missions with meteorological 656 instruments and magnetometers, such as China's Tianwen Zhurong lander, could also address the magnetic signals associated with pressure drops and DDs (Du et al., 2020), as could 657 658 continued terrestrial studies like those of the MATADOR campaign in the Arizona desert (Renno 659 et al., 2004; Farrell et al., 2004). Further understanding of DDs can help explain why they have 660 not yet been imaged at the InSight landing site, while all other Mars surface missions with visual 661 cameras have imaged several cases of dust devils (Ferri et al., 2003, Ellehoj et al., 2010).

662

663 Acknowledgments

- 664
- 665 We acknowledge NASA, CNES, their partner agencies and Institutions (UKSA, SSO, DLR, JPL,
- 666 IPGP-CNRS, ETHZ, IC, MPS-MPG) and flight operations team at JPL, SIS-MOC (SEIS on
- 667 Mars Operations Center) and MSDS (Mars SEIS Data Service). We acknowledge support from
- the Natural Sciences and Engineering Research Council of Canada (S.T, C.L.J.), the Canadian
- 669 Space Agency and the InSight Mission (C.L.J.), CNES in the frame of the InSight mission (A.S,
- B.L.), as well as ETH 19-2 FEL-34 and funding through the Harvard Daly Fellowship (A.M.).
- All InSight data used in this study are publicly available in the Planetary Data System: InSight
- 672 IFG calibrated data https://doi.org/10.17189/1519202 and Calibration data
- https://doi.org/10.17189/1510486. This paper is InSight Contribution Number, ICN 231.
- 674

675 **References**

- Atreya, S. K., Wong, A. S., Renno, N. O., Farrell, W. M., Delory, G. T., Sentman, D. D., ...
 Catling, D. C. (2006). Oxidant enhancement in Martian dust devils and storms: Implications
- 678 for life and habitability. *Astrobiology*, 6(3), 439–450. https://doi.org/10.1089/ast.2006.6.439
- Baker, M., Newman, C., Charalambous, C., Golombek, M., Spiga, A., Banfield, D., et al. (2021).
- 680 Vortex-dominated aeolian activity at InSight's landing site, Part 2: Local meteorology,
- transport dynamics, and model analysis. *Journal of Geophysical Research: Planets, 126,*e2020JE006514. https://doi.org/10.1029/2020JE006514
- Balme, M., & Greeley, R. (2006). Dust devils on Earth and Mars. *Reviews of Geophysics*, 44(3),
 RG3003. https://doi.org/10.1029/2005RG000188
- Banerdt, W. B., Smrekar, S. E., Banfield, D., Giardini, D., Golombek, M., Johnson, C. L., ...
 Wieczorek, M. (2020). Initial results from the InSight mission on Mars. *Nature Geoscience*. https://doi.org/10.1038/s41561-020-0544-y
- Bo, T. L., Zhang, H., & Zheng, X. J. (2014). Charge-to-mass ratio of saltating particles in wind blown sand. *Scientific Reports*, 4. https://doi.org/10.1038/srep05590
- Banfield, D, Rodriguez-Manfredi, J. A., Russell, C. T., Rowe, K. M., Leneman, D., Lai, H. R.,
 ... Team, T. T. (2018). InSight Auxiliary Payload Sensor Suite (APSS). *Space Science Reviews*, *215*(1), 4. https://doi.org/10.1007/s11214-018-0570-x
- Banfield, D., Spiga, A., Newman, C., Forget, F., Lemmon, M., Lorenz, R., ... Banerdt, W. B.
 (2020). The atmosphere of Mars as observed by InSight. *Nature Geoscience*, *13*(3), 190–
 198. https://doi.org/10.1038/s41561-020-0534-0
- 697 Charalambous, C., McClean J., B., Baker, M., Pike, W. T., Golombek, M., Lemmon, M., ...
 698 Banerdt, W. B. (2021). Vortex-dominated aeolian activity at InSight's landing site, Part 1:
 699 Multi-instrument Observations, Analysis and Implications. *Journal of Geophysical*700 *Research: Planets*. https://doi.org/10.1029/2020je006757

701	Chatain, A. and Spiga, A. and the InSight team: Seasonal variability of Mars atmospheric
702	turbulence studied from InSight high-frequency pressure measurements, European
703	Planetary Science Congress 2021, online, 13–24 Sep 2021, EPSC2021-290,
704	https://doi.org/10.5194/epsc2021-290, 2021.
705	Du, A. M., Zhang, Y., Li, H. Y., Qiao, D. H., Yi, Z., Zhang, T. L., Dai, J. L. (2020). The
706	Chinese Mars ROVER Fluxgate Magnetometers. <i>Space Science Reviews</i> .
707	https://doi.org/10.1007/s11214-020-00766-8
708 709 710	Ellehoj M. D., Gunnlaugsson H. P., Taylor P. A. et al. (2010). Convective vortices and dust devils at the Phoenix Mars mission landing site. <i>Journal of Geophysical Research</i> . https://doi.org/10.1029/2009JE003413
711	Farrell, W. M., Smith, P. H., Delory, G. T., Hillard, G. B., Marshall, J. R., Catling, D.,
712	Johnson, B. (2004). Electric and magnetic signatures of dust devils from the 2000-2001
713	MATADOR desert tests. <i>Journal of Geophysical Research E: Planets</i> , 109(3).
714	https://doi.org/10.1029/2003je002088
715	Farrell, William M., Marshall, J. R., Cummer, S. A., Delory, G. T., & Desch, M. D. (2006). A
716	model of the ULF magnetic and electric field generated from a dust devil. <i>Journal of</i>
717	<i>Geophysical Research E: Planets</i> , 111(11), E11004. https://doi.org/10.1029/2006JE002689
718 719	Fenton, L. K., & Lorenz, R. (2015). Dust devil height and spacing with relation to the Martian planetary boundary layer thickness. <i>Icarus</i> . https://doi.org/10.1016/j.icarus.2015.07.028
720	Ferri F., Smith P. H., Lemmon, M., Renno, N. O. (2003) Dust devils as observed by Mars
721	Pathfinder. <i>Journal of Geophysical Research</i> . https://doi.org/10.1029/2000JE001421
722	Harrison, R. G., Barth, E., Esposito, F., Merrison, J., Montmessin, F., Aplin, K. L.,
723	Zimmerman, M. (2016). Applications of Electrified Dust and Dust Devil Electrodynamics
724	to Martian Atmospheric Electricity. <i>Space Science Reviews</i> , Vol. 203, pp. 299–345.
725	https://doi.org/10.1007/s11214-016-0241-8
726	Houser, J. G., Farrell, W. M., & Metzger, S. M. (2003). ULF and ELF magnetic activity from a
727	terrestrial dust devil. <i>Geophysical Research Letters</i> , 30(1).
728	https://doi.org/10.1029/2001GL014144
729	Jackson, T. L., & Farrell, W. M. (2006). Electrostatic fields in dust devils: An analog to Mars.
730	IEEE Transactions on Geoscience and Remote Sensing, 44(10), 2942–2949.
731	https://doi.org/10.1109/TGRS.2006.875785
732	Jackson, B., Crevier, J., Szurgot, M., Battin, R., Clement, P., Rodriguez, S. (2021) Inferring
733	Vortex and Dust Devil Statistics from InSight. <i>The Planetary Science Jorunal</i> , 2 206.
734	https://doi.org/10.3847/PSJ/ac260d
735	Johnson, C. L., Mittelholz, A., Langlais, B., Russell, C. T., Ansan, V., Banfield, D., Banerdt,

736 W. B. (2020). Crustal and time-varying magnetic fields at the InSight landing site on Mars. 737 Nature Geoscience, 13(3), 199–204. https://doi.org/10.1038/s41561-020-0537-x 738 Joy, S. P., Mafi, J. N., & Slavney, S. (2019). Interior Exploration Using Seismic Investigations, 739 Geodesy, and Heat Transport (InSight) Mission Insight Fluxgate Magnetometer (IFG) PDS 740 Archive Software Interface Specification. urn:nasa:pds:insight-ifg-mars:document:insight-741 ifg-sis. 742 Kahre, M. A., Murphy, J. R., & Haberle, R. M. (2006). Modelling the Martian dust cycle and 743 surface dust reservoirs with the NASA Ames general circulation model. Journal of 744 Geophysical Research E: Planets. https://doi.org/10.1029/2005JE002588 Kok, J. F., and Renno, N. O. (2009), Electrification of wind-blown sand on Mars and its 745 746 implications for atmospheric chemistry, Geophys. Res. Lett., 36, L05202, 747 doi:10.1029/2008GL036691. 748 Kurgansky, M. V, Baez, L., & Ovalle, E. M. (2007). A simple model of the magnetic emission 749 from a dust devil. Journal of Geophysical Research E: Planets, 112(11), 11008. 750 https://doi.org/10.1029/2007JE002952 751 Lemmon, M. T., Guzewich, S. D., McConnochie, T., de Vicente-Retortillo, A., Martínez, G., 752 Smith, M. D., ... Jacob, S. (2019). Large Dust Aerosol Sizes Seen During the 2018 Martian 753 Global Dust Event by the Curiosity Rover. *Geophysical Research Letters*, 46(16). 754 https://doi.org/10.1029/2019GL084407 755 756 Lognonné, P., Banerdt, W. B., Pike, W. T., Giardini, D., Christensen, U., Garcia, R. F., ... 757 Zweifel, P. (2020). Constraints on the shallow elastic and anelastic structure of Mars from 758 InSight seismic data. Nature Geoscience. https://doi.org/10.1038/s41561-020-0536-y 759 Lorenz, R. D. (2016). Heuristic estimation of dust devil vortex parameters and trajectories from 760 single-station meteorological observations: Application to InSight at Mars. *Icarus*, 271, 761 326-337. https://doi.org/10.1016/j.icarus.2016.02.001 762 Lorenz, R. D., & Jackson, B. K. (2015). Dust devils and dustless vortices on a desert playa 763 observed with surface pressure and solar flux logging. GeoResJ. 764 https://doi.org/10.1016/j.grj.2014.11.002 765 Lorenz, R. D., Kedar, S., Murdoch, N., Lognonné, P., Kawamura, T., Mimoun, D., & Banerdt, 766 W. B. (2015). Seismometer detection of dust devil vortices by ground tilt. Bulletin of the 767 Seismological Society of America, 105(6), 3015–3023. https://doi.org/10.1785/0120150133 768 Lorenz, R. D., Lemmon, M. T., & Maki, J. (2021). First mars year of observations with the 769 InSight solar arrays: Winds, dust devil shadows, and dust accumulation. Icarus. 770 https://doi.org/10.1016/j.icarus.2021.114468

771 Lorenz, R. D., Lemmon, M. T., Maki, J., Banfield, D., Spiga, A., Charalambous, C., ... Banerdt, 772 W. B. (2020). Scientific Observations With the InSight Solar Arrays: Dust, Clouds, and 773 Eclipses on Mars. Earth and Space Science. https://doi.org/10.1029/2019EA000992 774 Lorenz, R. D., & Reiss, D. (2015). Solar panel clearing events, dust devil tracks, and in-situ 775 vortex detections on Mars. Icarus, 248, 162-164. 776 https://doi.org/10.1016/J.ICARUS.2014.10.034 777 Lorenz, R. D., Spiga, A., Lognonné, P., Plasman, M., Newman, C. E., & Charalambous, C. 778 (2021). The whirlwinds of Elysium: A catalog and meteorological characteristics of "dust 779 devil" vortices observed by InSight on Mars. Icarus, 355, 114119. 780 https://doi.org/10.1016/j.icarus.2020.114119 781 Méndez Harper, J., McDonald, G., Dufek, J. Malaska, M. J., Burr D. M., Hayes, A. G., 782 McAdams, J., Wray, J. J. (2017). Electrification of sand on Titan and its influence on 783 sediment transport. Nature Geoscience 10, 260–265 https://doi.org/10.1038/ngeo2921 784 Merrison, J. P., Bertelsen, P., Frandsen, C., Gunnlaugsson, P., Knudsen, J. M., Lunt, S., ... 785 Uggerhøj, E. (2002). Simulation of the Martian dust aerosol at low wind speeds. Journal of 786 Geophysical Research E: Planets, 107(12). https://doi.org/10.1029/2001je001807 787 788 Mittelholz, A., Johnson, C. L., Thorne, S. N., Joy, S., Barrett, E., Fillingim, M. O., ... Banerdt, 789 W. B. (2020). The Origin of Observed Magnetic Variability for a Sol on Mars From 790 InSight. Journal of Geophysical Research: Planets, 125(9). 791 https://doi.org/10.1029/2020JE006505 792 Mittelholz, A., Johnson, C. L., Thorne, S. N., Yau, V., Joy, S., Barrett, E., ... Banerdt, W. B. 793 (2021). Magnetic Variations of a Sol Observed Over a Year on Mars with InSight. Lunar 794 and Planetary Science Conference. LPI Contributions 2548. 795 https://www.hou.usra.edu/meetings/lpsc2021/pdf/1941.pdf 796 Murdoch, N., Spiga, A., Lorenz, R., Garcia, R. F., Perrin, C., Widmer-Schnidrig, R., ... Banerdt, 797 W. B. (2021). Constraining Martian Regolith and Vortex Parameters From Combined 798 Seismic and Meteorological Measurements. Journal of Geophysical Research: Planets. 799 https://doi.org/10.1029/2020JE006410 800 Murdoch, N., Mimoun, D., Garcia, R. F., Rapin, W., Kawamura, T., Lognonné, P., ... Banerdt, 801 W. B. (2017). Evaluating the Wind-Induced Mechanical Noise on the InSight 802 Seismometers. Space Science Reviews. https://doi.org/10.1007/s11214-016-0311-y 803 Panning, M. P., Pike, W. T., Lognonné, P., Banerdt, W. B., Murdoch, N., Banfield, D., ... 804 Warren, T. (2020). On-Deck Seismology: Lessons from InSight for Future Planetary 805 Seismology. Journal of Geophysical Research: Planets. 806 https://doi.org/10.1029/2019JE006353

807 808 809 810	Pollack, J. B., Ockert-Bell, M. E., & Shepard, M. K. (1995). Viking Lander image analysis of Martian atmospheric dust. <i>Journal of Geophysical Research</i> , 100(E3). https://doi.org/10.1029/94JE02640
811	Raack, J., Reiss, D., Balme, M. R., Taj-Eddine, K., & Ori, G. G. (2018). In Situ Sampling of
812	Relative Dust Devil Particle Loads and Their Vertical Grain Size Distributions.
813	<i>Astrobiology</i> , 18(10). https://doi.org/10.1089/ast.2016.1544
814	Renno, N. O., Abreu, V. J., Koch, J., Smith, P. H., Hartogensis, O. K., De Bruin, H. A., Burose,
815	D., Delory, G. T., Farrell, W. M., Watts, C. J., & Garatuza, J. (2004). MATADOR 2002: A
816	pilot field experiment on convective plumes and dust devils. <i>Journal of Geophysical</i>
817	Research: Planets, 109(E7), E07001. https://doi.org/10.1029/2003JE002219
818 819 820 821	Ringrose, T. J., Patel, M. R., Towner, M. C., Balme, M., Metzger, S. M., & Zarnecki, J. C. (2007). The meteorological signatures of dust devils on Mars. <i>Planetary and Space Science</i> , 55(14). https://doi.org/10.1016/j.pss.2007.07.002
822 823 824 825	Sharma, R., Clark, D. W., Srirama, P. K., & Mazumder, M. K. (2008). Tribocharging characteristics of the Mars dust simulant (JSC Mars-1). <i>IEEE Transactions on Industry</i> <i>Applications</i> , 44(1). https://doi.org/10.1109/TIA.2007.912761
826	Spiga, A., Barth, E., Gu, Z., Hoffmann, F., Ito, J., Jemmett-Smith, B., Wei, W. (2016). Large-
827	Eddy Simulations of Dust Devils and Convective Vortices. <i>Space Science Reviews</i> .
828	https://doi.org/10.1007/s11214-016-0284-x
829	Spiga, A., Murdoch, N., Lorenz, R., Forget, F., Newman, C., Rodriguez, S., Banerdt, W. B.
830	(2021). A study of daytime convective vortices and turbulence in the Martian planetary
831	boundary layer based on half-a-year of InSight atmospheric measurements and large-eddy
832	simulations. <i>Journal of Geophysical Research: Planets</i> , 126, e2020JE006511.
833	https://doi.org/10.1029/2020JE006511
834	Stähler, S. C., Widmer-Schnidrig, R., Scholz, JR., van Driel, M., Mittelholz, A., Hurst, K.,
835	et al. (2020). Geophysical observations of Phobos transits by InSight. <i>Earth and Space</i>
836	<i>Science Open Archive</i> , 15. https://doi.org/10.1002/essoar.10503257.1
837 838 839 840 841	Stott, A. E., Garcia, R. F., Pinot, B., Murdoch, N., Mimoun, D., Spiga, A., Banfield, D., Navarro, S., Mora-Sotomayor, L., Charalambous, C., Pike, W. T., Lognonné, P., and Horleston, A.: Atmospherically driven ground motion at InSight: a machine learning perspective, <i>EGU General Assembly 2021</i> , online, 19–30 Apr 2021, EGU21-12344, https://doi.org/10.5194/egusphere-egu21-12344, 2021
842	Thomas, P., & Gierasch, P. J. (1985). Dust devils on Mars. <i>Science</i> .

843 https://doi.org/10.1126/science.230.4722.175

- 844 Thorne, S. N., Mittelholz, A., Johnson, C. L., Joy, S., Liu, X., Russell, C. T., ... Banerdt, W. B.
- 845 (2020). InSight fluxgate magnetometer data calibration assessment and implications. *LPI* 846 *Contributions 2089*.