



**HAL**  
open science

## Modelling wintertime Arctic Haze and sea-spray aerosols

Eleftherios Ioannidis, Kathy S. Law, Jean-Christophe Raut, Louis Marelle, Tatsuo Onishi, Rachel M. Kirpes, Lucia Upchurch, Andreas Massling, Henrik Skov, Patricia K. Quinn, et al.

### ► To cite this version:

Eleftherios Ioannidis, Kathy S. Law, Jean-Christophe Raut, Louis Marelle, Tatsuo Onishi, et al.. Modelling wintertime Arctic Haze and sea-spray aerosols. Atmospheric Chemistry and Physics Discussions, In press, 10.5194/egusphere-2022-310 . insu-03696235v1

**HAL Id: insu-03696235**

**<https://insu.hal.science/insu-03696235v1>**

Submitted on 15 Jun 2022 (v1), last revised 3 Jun 2023 (v2)

**HAL** is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.



Distributed under a Creative Commons Attribution 4.0 International License



## Modelling wintertime Arctic Haze and sea-spray aerosols

Eleftherios Ioannidis<sup>1</sup>, Kathy S. Law<sup>1</sup>, Jean-Christophe Raut<sup>1</sup>, Louis Marelle<sup>1</sup>, Tatsuo Onishi<sup>1</sup>, Rachel M. Kirpes<sup>2</sup>, Lucia Upchurch<sup>3</sup>, Andreas Massling<sup>4</sup>, Henrik Skov<sup>4</sup>, Patricia K. Quinn<sup>3</sup>, and Kerri A. Pratt<sup>2,5</sup>

<sup>1</sup>LATMOS/IPSL, Sorbonne Université, UVSQ, CNRS, Paris, France

<sup>2</sup>Department of Chemistry, University of Michigan, Ann Arbor, Michigan, USA

<sup>3</sup>Pacific Marine Environmental Laboratory, National Oceanic and Atmospheric Administration, Seattle, Washington 98115, United States

<sup>4</sup>Department of Environmental Science, iClimate, Aarhus University, Denmark

<sup>5</sup>Department of Earth and Environmental Sciences, University of Michigan, Ann Arbor, Michigan, USA

**Correspondence:** Eleftherios Ioannidis (eleftherios.ioannidis@latmos.ipsl.fr) and Kathy S. Law (Kathy.Law@latmos.ipsl.fr)

**Abstract.** Anthropogenic and natural emissions contribute to enhanced concentrations of aerosols, so-called Arctic Haze in the Arctic winter and early spring. Models still have difficulties reproducing available observations. Whilst most attention has focused on the contribution of anthropogenic aerosols, there has been less focus on natural components such as sea-spray aerosols (SSA), including sea-salt sulphate and marine organics, which can make an important contribution to fine and coarse mode aerosols, particularly in coastal areas. Models tend to underestimate sub-micron and overestimate super-micron SSA in polar regions, including in the Arctic region. Quasi-hemispheric runs of the Weather Research Forecast model, coupled with chemistry model (WRF-Chem) are compared to aerosol composition data at remote Arctic sites to evaluate the model performance simulating wintertime Arctic Haze. Results show that the model overestimates sea-salt (sodium and chloride) and nitrate and underestimates sulphate aerosols. Inclusion of more recent wind-speed and sea-surface temperature dependencies for sea-salt emissions, as well as inclusion of marine organic and sea-salt sulphate aerosol emissions leads to better agreement with the observations during wintertime. The model captures better the contribution of SSA to total mass for different aerosol modes, ranging from 20-93% in the observations. The sensitivity of modelled SSA to processes influencing SSA production are examined in regional runs over northern Alaska (United States) where the model underestimates episodes of high SSA, particularly in the sub-micron, that were observed in winter 2014 during field campaigns at the Barrow Observatory, Utqiagvik. A local source of marine organics is also included following previous studies showing evidence for an important contribution from marine emissions. Model results show relatively small sensitivity to aerosol dry removal with more sensitivity (improved biases) to using a higher wind speed dependence based on sub-micron data reported from an Arctic cruise. Sea-ice fraction, including sources from open leads, is shown to be a more important factor controlling modelled super-micron SSA than sub-micron SSA. The findings of this study support analysis of the field campaign data pointing out that open leads are the primary source of SSA, including marine organic aerosols during wintertime at the Barrow Observatory, Utqiagvik. Nevertheless, episodes of high observed SSA are still underestimated by the model at this site, possibly due to missing sources such as SSA production from breaking waves. An analysis of the observations and model results does not suggest an influence from blowing snow and frost flowers to SSA during the period of interest. Reasons for the high concentrations of sub-micron SSA observed at this site, higher than other Arctic sites, require further investigation.



## 25 1 Introduction

The Arctic region is warming faster than any other region on Earth (Allan, 2021). Greenhouse gases, in particular carbon dioxide, and short-lived climate forcers like methane, ozone and, aerosols have a significant impact on the environment, with a particularly strong warming effect in the Arctic region (AMAP, 2015; Allan, 2021). This region is influenced by enhanced concentrations of aerosols (including sulphate ( $\text{SO}_4^{2-}$ ), nitrate ( $\text{NO}_3^-$ ), black carbon (BC) and organic aerosols (OA) during winter and spring, a phenomenon called Arctic Haze (Rahn and McCaffrey, 1980; Barrie et al., 1994; Quinn et al., 2002). Transport of aerosols and their precursors from mid-latitudes anthropogenic emissions contribute to Arctic Haze (Heidam et al., 2004; Quinn et al., 2007; Law et al., 2014). Local within and near-Arctic anthropogenic and natural sources also contribute to Arctic Haze during wintertime and the winter-spring transition (Law et al., 2017; Schmale et al., 2018; Kirpes et al., 2019). During wintertime 14% of organic mass at Alert originated from gas flaring in northern Russia (Leitch et al., 2018). For example, gas flaring from Russia contributes to black carbon at Alert (northern Canada) and Utqiagvik (northern Alaska) (Stohl et al., 2013; Qi et al., 2017; Xu et al., 2017; Marelle et al., 2018). Metal industry and combustion sources, such as power generation, from Siberia (e.g. Kola peninsula) were identified as sources of pollution at Villum station, Greenland during winter and spring (Nguyen et al., 2013). Metal smelting from Siberia also contributes to  $\text{SO}_4^{2-}$  at Zeppelin during wintertime (Hirdman et al., 2010). A more recent study by Winiger et al. (2019) showed that during wintertime Arctic sites, such as Utqiagvik, Alert, Zeppelin, are influenced by fossil fuel combustion emissions. Petroleum extraction on the North Slope of Alaska, including Prudhoe Bay, was found to influence aerosol distributions, composition, and particle growth at Utqiagvik, with enhanced growth of ultrafine particles (Kolesar et al., 2017; Kirpes et al., 2018).

Natural aerosol sources also contribute to Arctic Haze such as dust, volcanic emissions and sea-spray aerosols (SSA) (Rahn et al., 1977; Barrie and Barrie, 1990; Quinn et al., 2002; Stone et al., 2014; Huang et al., 2015; Zwaafink et al., 2016; Kirpes et al., 2018). Dust is not only transported from mid-latitudes sources (Asia, Africa), but it is also produced within the Arctic, with local dust contributing up to 85% to total dust burden in the Arctic (Zwaafink et al., 2016). During wintertime, fresh SSA (including sodium ions ( $\text{Na}^+$ ), chloride ions ( $\text{Cl}^-$ ), sea-salt ( $\text{ss-SO}_4^{2-}$ ) and marine organics) can be a significant fraction of particulate matter, 40% of super-micron (1 to 10  $\mu\text{m}$  particle diameter) and 25% of sub-micron (up to 1  $\mu\text{m}$  particle diameter) (Quinn et al., 2002). While studies have largely focused on anthropogenic sources of Arctic Haze influencing, in particular BC and  $\text{SO}_4^{2-}$ , there have been fewer studies on the contribution of SSA, the focus of this study. The primary mechanism leading to the formation of SSA is bubble bursting (jet-drop and film-drop formations) on the sea surface due to wind stress during whitecap formation (Monahan et al., 1986). For this reason, wind speed is a significant parameter affecting SSA productivity (Russell et al., 2010; Saliba et al., 2019). Arctic warming is leading to a decrease of sea-ice during summertime and, as a result, less and thinner sea-ice is forming during wintertime (Stroeve et al., 2012). Thus, new SSA sources, such as open ocean and leads, may contribute more in the future to the total aerosol burden over Arctic coastal regions, impacting CCN concentrations and radiative forcing (Ma et al., 2008).

A detailed analysis of in-situ aerosol composition in Utqiagvik revealed that, due to long-range transport from the North Pacific (due to strong winds in source regions, such as in the Pacific Ocean), sub-micron SSA peaks in winter and early spring,



while super-micron SSA peaks in summer, due to sea-ice retreat (Quinn *et al.*, 2002). However in winter, super-micron SSA mass concentrations increase in the presence of open leads, while sub-micron SSA appear to be more influenced by long-range transport (May *et al.*, 2016; Kirpes *et al.*, 2019). Kirpes *et al.* (2018) analysed atmospheric particle samples collected in winter 2014 in Utqiagvik. They found that the samples were influenced by air masses from the Arctic Ocean to the north and Prudhoe Bay oilfields to the east. Aged SSA were always internally mixed with secondary  $\text{SO}_4^{2-}$ , or with both  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  and reduced chlorine, suggesting anthropogenic influence from background Arctic Haze or Alaskan oil field emissions. Kirpes *et al.* (2019) concluded that fresh SSA, based on the presence of  $\text{Na}^+$  and  $\text{Cl}^-$  in ratios similar to seawater, including marine organic aerosols, were produced locally from open leads, with indications of secretions from sea ice algae and bacteria based on observed enrichments in carbon/sodium (C: $\text{Na}^+$ ) ratios. Previous studies of the Arctic and North Atlantic during wintertime and the winter-spring transition also showed that the majority of sub-micron organic mass (OM) is highly correlated with  $\text{Na}^+$  concentrations (Russell *et al.*, 2010; Shaw *et al.*, 2010; Frossard *et al.*, 2011; Leaitch *et al.*, 2018). Frost flowers with organic-salt coatings have also been proposed as a possible source of wintertime SSA (Xu *et al.*, 2013), although Kirpes *et al.* (2019) found no evidence of frost flowers or blowing snow as a potential source, supporting the findings of older studies (Roscoe *et al.*, 2011).

Regional and global models have difficulties capturing wintertime Arctic Haze composition and often underestimate  $\text{SO}_4^{2-}$  and BC (Bond *et al.*, 2013; Eckhardt *et al.*, 2015; Sato *et al.*, 2016; Schwarz *et al.*, 2017; Whaley *et al.*, 2022), while the contribution of SSA to Arctic Haze remains poorly evaluated (Kirpes *et al.*, 2019). Representation of SSA concentrations in models has been improved over recent years, but with less focus on the Arctic winter. For example, SSA source functions with updated dependencies on wind speeds, sea surface temperatures (SSTs) or salinity (Revell *et al.*, 2019; Jaeglé *et al.*, 2011; Sofiev *et al.*, 2011) have led to improve simulation of super-micron SSA. However, sub-micron SSA is often still underestimated (Sofiev *et al.*, 2011; Huang and Jaeglé, 2017) and sub-micron emissions of SSA from frost flowers and blowing snow have been included in models (Xu *et al.*, 2013, 2016; Huang and Jaeglé, 2017; Rhodes *et al.*, 2017). Modelled SSA including a source of frost flowers captures better monthly SSA concentrations at Alert during wintertime, while a source of blowing snow overestimates observations (Huang and Jaeglé, 2017; Marelle *et al.*, 2021). At Utqiagvik during January and February a source of blowing snow improves modelled SSA; however it still cannot explain the high observed SSA, while the blowing snow explains high observed SSA in the Antarctic (Huang and Jaeglé, 2017).

In this study, the performance of the Weather Research Forecast model, coupled with chemistry (WRF-Chem), is examined with regard to its ability to simulate Arctic Haze composition as well as SSA components, including  $\text{ss-SO}_4^{2-}$  and marine organics. The model is first evaluated against available data over the wider Arctic, and the sensitivity to more recent treatments of SSA wind speed and SST dependencies, is investigated. Inclusion of a marine organic source is also examined (Fuentes *et al.*, 2010, 2011). The findings of Kirpes *et al.* (2019) are used as a basis for a more focused regional study to evaluate modelled Arctic wintertime aerosol composition in northern Alaska. The sensitivity of model results to processes influencing SSA production and concentrations are investigated including aerosol dry deposition, wind speed dependence and sea-ice fraction. Missing local sources of marine organics are also examined based on the findings of Kirpes *et al.* (2019). A companion paper,



*Ioannidis et al., (2022)* (in prep.), examines the contribution of remote and regional anthropogenic emissions to Arctic BC in northern Alaska and northern Russia during wintertime.

95 The model setup, including the emissions are described in Section 2. The observed aerosol composition used to evaluate the model performance are introduced in Section 3. The model runs, including sensitivity simulations, together with results are presented in Sections 4 and 5. First, in Section 4, simulated Arctic Haze, focusing on SSA, is evaluated at remote Arctic sites. Second, in Section 5, the results from the regional study over northern Alaska during wintertime and sensitivity of results to processes influencing SSA production in the model are presented. The implications of our findings for the simulation of Arctic  
100 Haze aerosols and conclusions are presented in Section 6.

## 2 WRF-Chem

### 2.1 Model Setup

WRF-Chem chemical transport model version 3.9.1.1 is used to simulate quasi-hemispheric and regional Arctic Haze aerosols and to examine local SSA sources over northern Alaska. WRF-Chem is a fully coupled, online meteorological and chemical  
105 transport mesoscale model (*Grell et al., 2005; Fast et al., 2006*). Recent improvements in the WRF-Chem model over the Arctic are included in the version used in this study (*Marelle et al., 2017*). The model setup, including meteorological and chemical schemes, is shown in Table 1. Briefly, Yonsei University (YSU - boundary layer), Model Version 5 similarity (MM5 - surface layer) and Noah-Multiparameterization Land Surface Model (NOAH MP, land surface model) are used. More details about the NOAH MP scheme are given in APPENDIX A.

110 All the various processes for aerosols in the atmosphere, like nucleation, evaporation, coagulation, condensation, dry deposition, aerosol/cloud interactions and aqueous chemistry, are included in the Model for Simulating Aerosol Interactions and Chemistry (MOSAIC, *Zaveri et al. (2008)*) scheme. MOSAIC treats all the major aerosol species, such as  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$ ,  $\text{Cl}^-$ , ammonium ( $\text{NH}_4^+$ ),  $\text{Na}^+$ , BC, and OA. The size distribution of each aerosol species is represented by eight bins, from 39 nm to 10  $\mu\text{m}$ . Each bin is assumed to be internally mixed, and both mass and number are simulated. The applied MOSAIC version  
115 includes secondary organic aerosol formation (SOA) from the oxidation of anthropogenic and biogenic species (*Shrivastava et al., 2011; Marelle et al., 2017*) and is combined with SAPRC-99 gas-phase chemistry. In the base model, OA is the sum of SOA and anthropogenic emissions of organic matter (OM). Aerosol sedimentation in MOSAIC is calculated throughout the atmospheric column based on the Stokes velocity scheme, as described in *Marelle et al. (2017)*.

### 2.2 Emissions

120 This section provides details about the emissions that are used in the simulations. More details are provided about SSA emissions since this is the focus of this study.



**Table 1.** WRF-Chem model setup.

Parametrization scheme	Options
Physics (WRF)	
Planetary boundary layer	Yonsei University (YSU) - ( <i>Hong et al., 2006</i> )
Surface layer	Pennsylvania State / NCAR Mesoscale Model Version 5 (MM5) similarity ( <i>Grell et al., 1994; Jiménez et al., 2012</i> )
Land surface	NOAH MP ( <i>Niu et al., 2011</i> )
Microphysics	Morrison ( <i>Morrison, 2009</i> )
SW & LW radiation	Rapid Radiative Transfer Model (RRTMG - <i>Iacono and D. (2008)</i> )
Cumulus parameterization	Kain-Fritsch with cumulus potential (KF-CuP) ( <i>Berg et al., 2013</i> )
Chemistry (WRF-Chem)	
Aerosol model	MOSAIC 8-bins ( <i>Zaveri et al., 2008</i> )
Gas-phase chemistry	Statewide Air Pollution Research Center SAPRC-99 modified with added dimethyl sulphide chemistry ( <i>Carter, 2000; Marelle et al., 2017</i> )
Photolysis	Fast-J ( <i>Wild et al., 2000</i> )
Sea-spray emissions	<i>Gong et al. (1997)</i>

### 2.2.1 Anthropogenic and natural emissions

Anthropogenic emissions are from the Evaluating the Climate and Air Quality Impacts of Short-Lived Pollutants version 6 (ECLIPSE v6b) inventory, with a resolution of  $0.5^\circ \times 0.5^\circ$  (*Whaley et al., 2022*), including emissions of OM. Emissions of  
 125 dimethyl sulfide (DMS), mineral dust, and lightning  $\text{NO}_x$  are calculated online in the model (see *Marelle et al. (2017)* and  
 references therein). Biogenic emissions for 2014 are calculated online using Model of Emissions of Gases and Aerosol from  
 Nature (MEGAN) model (*Guenther et al., 2012*).

### 2.2.2 Sea-spray emissions

In the control simulation, sea-salt emissions of  $\text{Na}^+$  and  $\text{Cl}^-$  are included. They are calculated per particle radius, with 1000  
 130 sub-bins per MOSAIC bin, using the density function  $dF/dr$  (in particles  $\text{m}^{-2} \text{s}^{-1} \mu \text{m}^{-1}$ ) from (*Gong et al., 1997*) (G97 from  
 now on) which represents the rate of seawater droplets form per unit area (sea surface) and per increase of particle radius and  
 its derived from the source function based on laboratory experiments described in *Monahan et al. (1986)* (MO86 from now  
 on):

$$\frac{dF}{dr} = 1.373 \times U_{10}^{3.41} \times r^{-3} (1 + 0.057 \times r^{1.05}) \times 10^{1.19e^B} \quad (1)$$

135 where  $F$  is a function of  $U$  and  $r$ ,  $r$  is the particle radius at relative humidity (RH) equal to 80%,  $U$  the 10m-elevation wind  
 speed and  $B = \frac{(0.380 - \log r)}{0.650}$ . The source function is applied for particles with dry diameters of  $0.45 \mu\text{m}$  or more. For particles



with dry diameters less than  $0.45 \mu\text{m}$ , a correction is applied to the formula based on *O'Dowd et al. (1997)*. This approach is based on the whitecap method, where the emission flux scales linearly with the fraction of ocean area covered by whitecaps. Over open ocean, the whitecap fraction,  $W(U)$ , is determined as a function of wind speed (*Monahan and Muircheartaigh (1980)*; MO80 from now on):

$$W(U) = 3.84 \times 10^{-6} \times U_{10}^{3.41} \quad (2)$$

This expression for  $W(U)$  is included implicitly in Equation (1) following details provided in MO80. In the base version of WRF-Chem SSA emissions are calculated for every grid cell, which is open ocean or salt-water lakes. In this study, the grid cell which is covered by sea-ice is considered and then the fraction of that ice-free grid is used. In this way, SSA emissions from open leads are taken into account.



**Figure 1.** WRF-Chem simulation domains: d1 is the 100km domain and d2 is the 20km domain.



## 2.3 Simulations

Two simulation domains on a polar stereo-graphic projection are used in this study, as shown in Figure 1. The first (parent) domain (d1) covers a large part of the Northern Hemisphere with  $100 \times 100$  km horizontal resolution. The boundary and initial conditions, are derived from National Centres for Environmental Prediction Final meteorological reanalysis data (NCEP FNL  $1^\circ \times 1^\circ$ ), (*National Centers for Environmental Prediction, National Weather Service, NOAA, U.S. Department of Commerce* 150 (2000) and Model for OZone And Related chemical Tracers (MOZART, *Emmons et al. (2010)*) for atmospheric trace gases and aerosols. The nested domain (d2), run at horizontal resolution of  $20 \times 20$  km, covers continental Alaska, a small area of northwest Canada, and the Chukchi and Beaufort Seas (see Figure 1). 50 vertical levels and grid nudging are used for the 100 km resolution domain, while calculating spectral nudging parameters as in *Hodnebrog et al. (2019)*, is implemented in the 155 nested domain. WRF-Chem temperatures and winds are nudged at each dynamical step to the reanalysis, which are updated every 6 hours, above the atmospheric boundary layer.

The simulations performed in this study are discussed in sections 4 and 5. Simulations at 100 km are run for 4 months from November 2013 until the end of February 2014, with the first two months considered as spin-up. The model is run at 20 km for two different periods (23–28 January 2014 and 24–28 February 2014) corresponding to the campaign which took place in 160 Utqiagvik, and described earlier (*Kirpes et al. (2018, 2019)*), KRP18 and KRP19 from now on, respectively, see also sub-section 3.2). For these simulations, the initial, and boundary conditions are derived from the quasi-hemispheric simulation. A series of sensitivity runs are performed to examine processes affecting SSA emissions over northern Alaska. They are summarized in Table 3 and discussed in detail in Section 5. At 20 km for all the simulations, 4 days prior to the beginning of the campaign considered spin up. In all runs, the model results are output every 3h.

## 165 3 Aerosols

### 3.1 Routine monitoring sites

Surface mass concentration data (for aerodynamic diameters (defined as  $d_a$ )  $\leq 2.5 \mu\text{m}$  and  $d_a < 10 \mu\text{m}$ ), from EMEP (European Monitoring and Evaluation Programme) dataBASE (EBAS - <http://ebas.nilu.no>) for Zeppelin, Ny-Ålesund, Norway (78.9N, 11.88W) and Alert, Canada (82.5N, -63.3W), are used to evaluate the quasi-hemispheric model simulations together with 170 data from Villum Research Station, Station Nord, Greenland (81.6N, -16.7W), referred to as Villum from now on (reporting total suspended particulates). The data are collected on a daily (Zeppelin) and weekly (Villum, Alert) basis. At Alert, and Zeppelin (*Aas et al., 2021*), observations for  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{NH}_4^+$ ,  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  measured with ion chromatography are used (*Sharma et al., 2019*). At Villum, the same observations are collected using a filter-pack over a week and analysed using by ion-chromatography ( $\text{Cl}^-$ ,  $\text{SO}_4^{2-}$ ), cat-ion ionchromatograph ( $\text{Na}^+$ ) and segmented flow analysis ( $\text{NH}_4^+$ ). For all the EBAS 175 stations, the units of inorganic aerosols ( $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$ ) are converted to model units ( $\mu\text{gm}^{-3}$ ), using the ratio of molar weights of  $\text{NH}_4^+$ ,  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$  to molar weights of nitrogen or sulphur, respectively. With regard to measurement uncertainties,



EBAS documentation notes, in the case of Alert only, that there are uncertainties of around 33% and 36% in  $\text{Na}^+$ ,  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$  and  $\text{Cl}^-$ , respectively, and higher uncertainties (43%) for  $\text{NH}_4^+$ .

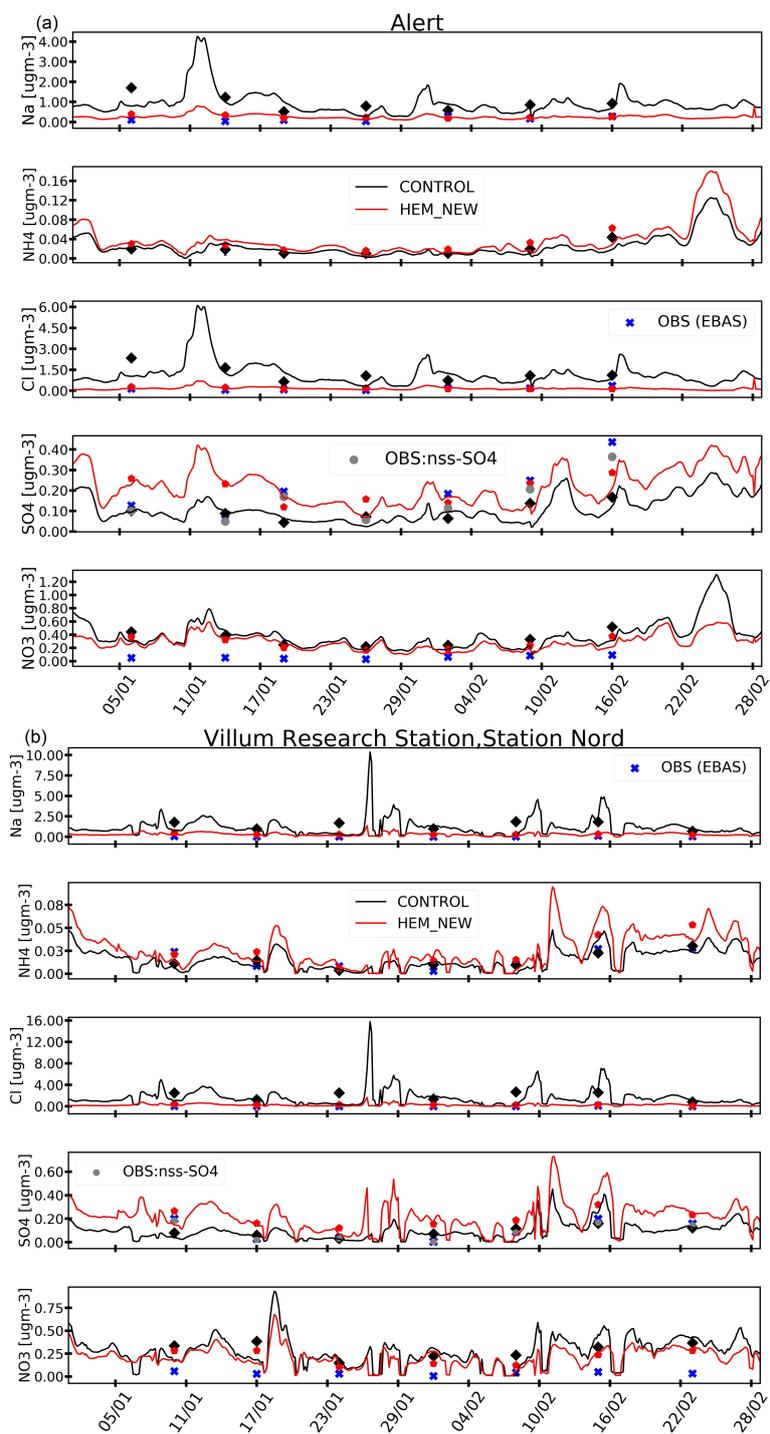
180 Surface mass concentration data, diameter less than  $2.5 \mu\text{m}$  ( $r_d \leq 2.5 \mu\text{m}$ ), from the Interagency Monitoring for Protected Visual Environments (IMPROVE) database is also used for model evaluation for Simeonof (55.3N, -160.5W), a sub-Arctic site on the Aleutians islands, south of Alaska and an inland site, Gates of the Arctic (66.9N, -151.5W) which is located south-east of Utqiagvik. The samples are collected on-site (e.g. Simeonof site) over 24 hours every three days (<http://views.cira.colostate.edu/fed/QueryWizard/Default.aspx>, *Malm et al. (1994)*). At these two sites observations of  $\text{Na}^+$ ,  $\text{Cl}^-$ , OC,  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  are used. To compare with the OC observations at the two Alaskan sites modelled OA is divided by 1.8, the reported ratio of  
185 OM/OC in the documentation for these two stations (*Malm et al., 1994*). In this study, mass concentration data with diameter  $\leq 2.5 \mu\text{m}$  are defined as fine mode aerosols, while diameter  $< 10 \mu\text{m}$  then are defined as coarse mode aerosols.

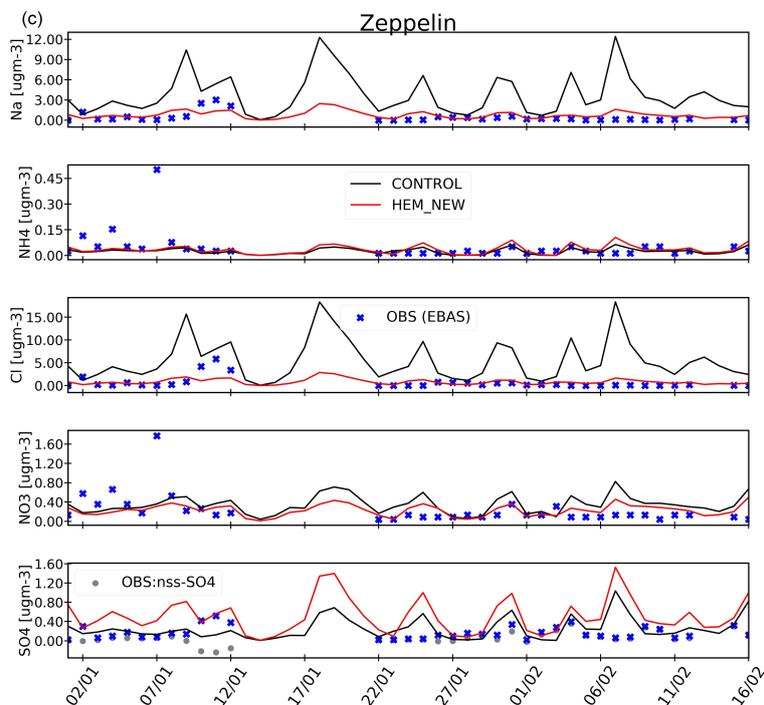
Sub-micron ( $d_a < 1.0 \mu\text{m}$ ) and super-micron ( $1.0 < d_a < 10 \mu\text{m}$ ) surface mass concentration data from the National Oceanic and Atmospheric Administration (NOAA) Barrow observatory (71.3N, -156.8W), near Utqiagvik town, is also used in this study, with daily and weekly temporal coverage, respectively. The sampling site is located 8 km northeast of Utqiagvik, 20 m  
190 above mean sea level (msl), with a prevailing, east-northeast wind off the Beaufort Sea. Concentration data ( $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{NH}_4^+$ ,  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$ ) are determined by ion chromatography (*Quinn et al., 1998*) and are sampled only for wind directions between 0 and 130 degrees (with 0 degrees indicating north). According to *Quinn et al. (2002)* measurement uncertainties of SSA components and  $\text{SO}_4^{2-}$  are below 1%, while for  $\text{NH}_4^+$  they are 7.8%. Observed ss- $\text{SO}_4^{2-}$  is calculated from observed  $\text{Na}^+$  concentrations and the mass ratio of  $\text{SO}_4^{2-}$  to  $\text{Na}^+$  in seawater of 0.252 (*Bowen et al., 1979; Calhoun et al., 1991*).

195 The model Stokes diameter ( $r_d$ ) is converted to aerodynamic diameter using the *Seinfeld and Pandis (1998)* formula. Thus, the diameter of modelled sub-micron particles is up to  $0.73 \mu\text{m}$  (including the first four MOSAIC bins and a fraction of the 5th bin), and super-micron particle diameters are between  $0.73$  to  $7.3 \mu\text{m}$  (fraction 5th bin, 6th and 7th bins and fraction 8th bin). Seven MOSAIC bins and a fraction of the 8th bin are used (modelled stokes  $r_d \leq 7.3 \mu\text{m}$ ) to compare with Alert and Zeppelin observations (aerodynamic  $d_a < 10 \mu\text{m}$ , coarse mode). All model aerosol bins are used to compare with observations  
200 at Villum, where the observations are reported as total suspended particulates (TSP), i.e. there is no cutoff. For each site, modelled aerosols are estimated at the same conditions (temperature, pressure) as the reported observations. Overall, particles at different size ranges (up to  $1.0 \mu\text{m}$ ,  $2.5 \mu\text{m}$ , and  $10 \mu\text{m}$ ) are used to validate the model performance in each domain.

### 3.2 Campaign data

205 Details about the field campaign (January 23–27 and February 24–28, 2014) measurements near Utqiagvik, Alaska can be found in KRP18 and KRP19. Briefly, atmospheric particles were collected using a rotating micro-orifice uniform deposition impactor located 2 m above the snow surface at a site located 5 km across the tundra from the NOAA Barrow Observatory and inland from the Arctic Ocean. The sampled particles were analysed by computer-controlled scanning electron microscopy with energy scattering X-ray spectroscopy (CCSEM-EDX) to determine the individual particle morphology and elemental composition. The analysed samples were collected either during daytime or nighttime and only when wind directions were





**Figure 2.** Evaluation of modelled aerosol composition (runs CONTROL and HEM\_NEW) against in-situ observations of (a) coarse mode aerosols ( $d_a < 10 \mu\text{m}$ ) at Alert, Canada (standard temperature pressure (STP) conditions), (b) TSP aerosols ( $d_a \leq 10 \mu\text{m}$ ) at Villum, Greenland and (c) coarse mode aerosols ( $d_a < 10 \mu\text{m}$ ) at Zeppelin, Norway in UTC. The black line shows model results from the CONTROL run; the red line shows the HEM\_NEW run, while observations are shown as blue crosses. Villum and Alert observations are weekly averages, and the corresponding model weekly averages are shown as black diamonds for CONTROL and red pentagons for HEM\_NEW. Zeppelin observations are daily 24h averages. Observations are shown only when they are available. See the text for details about the observations and model runs.

210 between 75 and 225 degrees, corresponding to minimise local pollution influence. Data analysis provided information about the different chemical components as a fraction of the total number of particles sampled.

#### 4 Processes influencing SSA over the wider Arctic and their contribution to wintertime Arctic aerosols

This section focuses on evaluating the capability of the model to simulate Arctic Haze aerosols during wintertime and improving model treatments of SSA. Briefly, in the base simulation (CONTROL), sea-salt emissions are calculated using the G97 parametrization scheme, including the MO80 whitecap method. All the updates described below are included in a new quasi-hemispheric simulation (HEM\_NEW) with the aim to improve the model. This includes addition of marine organics (*Fuentes et al., 2010*), using a more recent whitecap method (Salisbury et al., 2014), including the dependence of SSA emissions on SST (Jaeglé et al., 2011), and the addition of a  $\text{ss-SO}_4^{2-}$  component, based on Kelly et al. (2010). HEM\_NEW simulation is



then evaluated (sub-section 4.6) compared to CONTROL and the observations at the different sites followed by a discussion of  
220 the new results.

#### 4.1 Anthropogenic and natural aerosols in the Arctic

First considering the observations, at remote sites such as Alert (Fig 2a), observed  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{NO}_3^-$  coarse mode mass concentrations do not exceed 0.3, 0.5 and  $0.09 \mu\text{g m}^{-3}$ , respectively, during the study period. Total  $\text{SO}_4^{2-}$  (sum of  $\text{ss-SO}_4^{2-}$  and  $\text{nss-SO}_4^{2-}$ ) reach  $0.44 \mu\text{g m}^{-3}$ , which is mostly  $\text{nss-SO}_4^{2-}$ , as  $\text{ss-SO}_4^{2-}$  does not exceed  $0.09 \mu\text{g m}^{-3}$ , likely to be due to long-range  
225 transport from sources in north-central, western, northwest Russia and Europe (Leitch et al., 2018). Similar magnitudes have been reported in previous studies during winter months (Leitch et al., 2018).  $\text{NH}_4^+$  peaks at  $0.06 \mu\text{g m}^{-3}$  and originates from Russia and Europe during winter (Leitch et al., 2018). At Villum (Fig 2b), observed TSP  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{NO}_3^-$  reach up to 0.12, 0.13 and  $0.06 \mu\text{g m}^{-3}$ , respectively. These concentrations are lower than at Alert which could be explained by the fact that during winter the sea surrounding Villum station is frozen (Nguyen et al., 2013). Total  $\text{SO}_4^{2-}$  does not exceed  $0.2 \mu\text{g m}^{-3}$  and is  
230 mostly  $\text{nss-SO}_4^{2-}$  (up to  $0.18 \mu\text{g m}^{-3}$ ), while  $\text{ss-SO}_4^{2-}$  does not exceed  $0.03 \mu\text{g m}^{-3}$ . At Villum,  $\text{SO}_4^{2-}$  peaks during wintertime (Massling et al., 2015) and is the dominant component of Arctic Haze at this site (Lange et al., 2018).  $\text{NH}_4^+$  concentrations at Villum are up to  $0.1 \mu\text{g m}^{-3}$ . In the high Arctic,  $\text{Na}^+$  could potentially also originate from anthropogenic sources which could account for up to 35% of total  $\text{Na}^+$  (Barrie and Barrie, 1990). Note that this source is not included in the model, or in models generally. Higher  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{NO}_3^-$  concentrations are observed at Zeppelin (coarse mode) reaching up to 3.0, 5.9 and  $1.8$   
235  $\mu\text{g m}^{-3}$ , respectively (Fig. 2c). Total  $\text{SO}_4^{2-}$  does not exceed  $0.8 \mu\text{g m}^{-3}$  and might originate from metal smelting in Siberia (Hirdman et al., 2010).  $\text{ss-SO}_4^{2-}$  contributes up to  $0.8 \mu\text{g m}^{-3}$  of the total  $\text{SO}_4^{2-}$ . Note that, in some cases,  $\text{nss-SO}_4^{2-}$  has small negative concentrations, due to depletion of  $\text{ss-SO}_4^{2-}$  through fractionation processes (Quinn et al., 2002). Observed  $\text{NH}_4^+$  does not exceed  $0.5 \mu\text{g m}^{-3}$  during the study period.

At Simeonof, an ice-free sub-Arctic island in south western Alaska, high concentrations of fine mode  $\text{Na}^+$  and  $\text{Cl}^-$  are  
240 observed of up to 2.1 and  $1.0 \mu\text{g m}^{-3}$ , respectively (Fig. 3a), especially at the beginning of January 2014, with low values of  $\text{NO}_3^-$  (peaking at  $0.25 \mu\text{g m}^{-3}$ ). Total  $\text{SO}_4^{2-}$  reaches  $1.0 \mu\text{g m}^{-3}$  and is mostly  $\text{nss-SO}_4^{2-}$  ( $0.9 \mu\text{g m}^{-3}$ ), while the contribution of  $\text{ss-SO}_4^{2-}$  is smaller (up to  $0.3 \mu\text{g m}^{-3}$ ). Lower concentrations of fine mode  $\text{Na}^+$  and  $\text{Cl}^-$  (up to  $0.35 \mu\text{g m}^{-3}$ ) are observed at Gates of the Arctic (Fig. 3b), a non-coastal site located 404 km south-east of Utqiagvik in the Brooks Range Mountains, while  $\text{NO}_3^-$  peaks at  $0.45 \mu\text{g m}^{-3}$ . Total  $\text{SO}_4^{2-}$  peaks at  $0.64 \mu\text{g m}^{-3}$  and  $0.56 \mu\text{g m}^{-3}$  is  $\text{nss-SO}_4^{2-}$  possibly due to local anthropogenic  
245 emissions originating from the North Slope of Alaska oilfields which may affect the measurements although this site is located inland (391 km) south of the oilfields. The contribution of  $\text{ss-SO}_4^{2-}$  is insignificant (no more than  $0.08 \mu\text{g m}^{-3}$ ) at this site.

At Utqiagvik, observed super-micron ( $1.0 < d_a < 10.0 \mu\text{m}$ )  $\text{Na}^+$  and  $\text{Cl}^-$  concentrations reach  $1.2 \mu\text{g m}^{-3}$  (Fig. 4b), while  
 $\text{NO}_3^-$  peaks at  $0.2 \mu\text{g m}^{-3}$ . Super-micron  $\text{SO}_4^{2-}$  and  $\text{NH}_4^+$  do not exceed 0.16 and  $0.009 \mu\text{g m}^{-3}$ , respectively. Super-micron  
 $\text{NH}_4^+$  concentrations are insignificant (Quinn et al., 2002). However, there is more  $\text{ss-SO}_4^{2-}$  (up to  $0.18 \mu\text{g m}^{-3}$ ) than  $\text{nss-SO}_4^{2-}$ .  
250  $\text{SO}_4^{2-}$ . On the other hand, observed sub-micron  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{NO}_3^-$  at Utqiagvik peak at 2.0, 2.2, and  $0.9 \mu\text{g m}^{-3}$  respectively (Fig. 4a). Note that based on the findings of KRP18, only 1%, by number, of the particles across the  $0.15\text{-}1.0 \mu\text{m}$  size range corresponded to fly ash and dust, as compared to 50-90% from SSA across the same size range. This supports the assumption



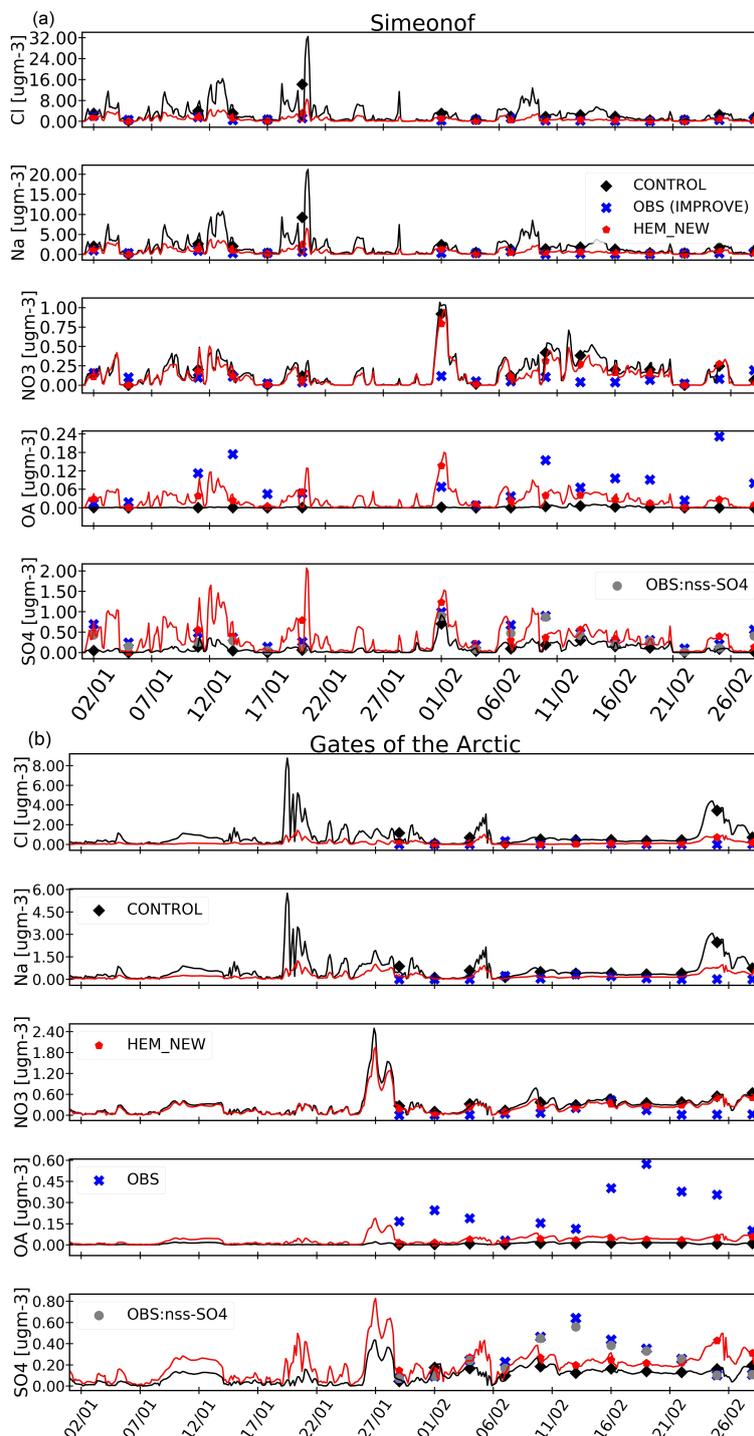
of  $\text{Na}^+$  being primarily from SSA during this study. High sub-micron observed total  $\text{SO}_4^{2-}$  (mostly nss- $\text{SO}_4^{2-}$ ) concentrations were measured at Utqiagvik and peak at  $2.4 \mu\text{g m}^{-3}$ , possibly due to local influence from Prudhoe Bay oil fields to the east (KRP18, KRP19), a magnitude much higher than super-micron  $\text{SO}_4^{2-}$ , also reported for Utqiagvik by Quinn *et al.* (2002). Enhanced nss- $\text{SO}_4^{2-}$  during this period at Utqiagvik could also be due to transport from mid-latitude sources, as well as due to transport and oxidation of  $\text{SO}_2$  to  $\text{SO}_4^{2-}$  near and within the Arctic region (Barrie and Hoff, 1984). Sub-micron ss- $\text{SO}_4^{2-}$  peaks at  $0.5 \mu\text{g m}^{-3}$ . Observed  $\text{NH}_4^+$  is higher compared to the other remote Arctic sites (up to  $0.34 \mu\text{g m}^{-3}$ ).  $\text{NH}_4^+$  temporal variation during January and February follows that of nss- $\text{SO}_4^{2-}$  due to  $\text{NH}_3$  reaction with acidic  $\text{SO}_4^{2-}$  aerosol near source regions outside of the Arctic (Quinn *et al.*, 2002) or due to regional sources of  $\text{NH}_3$ , e.g. combustion of fossil fuels (Whaley *et al.*, 2018).

Finally, only two sites provide total organic carbon (tOC) observations. Here, observed total organic carbon is assumed to include secondary organic aerosols, anthropogenic organic carbon emissions and marine organics. Thus, from now on it will be referred as tOC, to distinguish from OA and OM defined earlier. tOC ranges between  $0.15$  and  $0.3 \mu\text{g m}^{-3}$  at Simeonof and  $0.15$  and  $0.5 \mu\text{g m}^{-3}$  at Gates of the Arctic during January and February 2014.

Evaluation of the CONTROL simulation shows that the model overestimates observed fine/coarse mode, super-micron and TSP  $\text{Na}^+$  and  $\text{Cl}^-$  at most sites, and especially at Simeonof (by up to  $15 \mu\text{g m}^{-3}$ ), Zeppelin (Fig. 2c) (by up to  $5.0 \mu\text{g m}^{-3}$ ), Utqiagvik (by up to  $0.3 \mu\text{g m}^{-3}$ ) and Gates of the Arctic (Fig. 3b) (by up to  $4.0 \mu\text{g m}^{-3}$ ) site. The CONTROL simulation also overestimates  $\text{NO}_3^-$  by up to  $0.5 \mu\text{g m}^{-3}$  at each site. On the other hand, this simulation captures  $\text{NH}_4^+$  variability quite well at Alert, Villum and Utqiagvik (super-micron) (see also biases and RMSEs (Root Mean Square Error) in APPENDIX C and Tables C1, C2 and C7 respectively), whilst it overestimates  $\text{NH}_4^+$  at Zeppelin by up to  $0.4 \mu\text{g m}^{-3}$ . CONTROL includes only the nss- $\text{SO}_4^{2-}$  component, however it captures observed variability of total  $\text{SO}_4^{2-}$  at Zeppelin (coarse mode), Villum (TSP) and Utqiagvik (super-micron), but underestimates total  $\text{SO}_4^{2-}$  at Gates of the Arctic (fine mode) and Alert (coarse mode) by  $0.5$  and  $0.2 \mu\text{g m}^{-3}$ , respectively. In addition, the model underestimates sub-micron  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{SO}_4^{2-}$  and  $\text{NH}_4^+$  at Utqiagvik. It also underestimates OA at the two sites compared to the measurements. In the following sections, model improvements are described. Biases and RMSEs in  $\mu\text{g m}^{-3}$ , are given in APPENDIX C for all sites and available aerosol species at each location.

## 4.2 Marine organics

Recent data-analysis studies (Saliba *et al.*, 2019; Kirpes *et al.*, 2019), have suggested that marine organics contribute significantly to natural aerosol composition as ocean biomass can influence SSA number concentrations and diameter. In the CONTROL run, marine organics are not activated; however a source code is included in the model by Archer-Nicholls *et al.* (2014). For this reason, the parameterization, based on Fuentes *et al.* (2010, 2011) (F10 and F11 from now on, respectively) is activated in the MOSAIC scheme to include a source flux for marine organics with dry diameters up to  $0.45 \mu\text{m}$ . The scheme is based on an analysis of data from a cruise in mid-latitudes investigating the influence of dissolved organic matter on the production of sub-micron SSA. The F10 SSA source function also depends on MO80 whitecap coverage. In this study, organic fractions equal to 0.2 for the first and second MOSAIC bins, 0.1 for the third bin and 0.01 for the remaining bins are used following the high biogenic activity scenario which assumes high C:Chlorophyll-a (Chl-a) ratios. F11 found



**Figure 3.** Evaluation of modelled aerosol composition (runs CONTROL and HEM\_NEW) against in-situ aerosol observations of fine mode ( $\tau_d \leq 2.5 \mu\text{m}$ ) (both sites) at (a) Simeonof, Aleutians Islands, Alaska and (b) Gates of the Arctic, north of Alaska in local Alaskan time (AKST). The black line shows model results from the CONTROL run; the red line shows the HEM\_NEW run, while observations are shown as blue crosses. Simeonof and Gates of the Arctic observations are 24h averages every three days and the corresponding model daily averages are shown as black diamonds for CONTROL and red pentagons for HEM\_NEW. Observations are shown only when they are available. See the text for details about the observations and model runs.



that higher particle organic fractions are expected in algal bloom regions with high C:Chl-a ratios and Chl-a varying between 0.4-10  $\mu\text{g/L}$ . The use of the F11 high biogenic activity option in our simulations is justified since MODIS-Aqua satellite data ([https://neo.sci.gsfc.nasa.gov/view.php?datasetId=MY1DMW\\_CHLORA&date=2014-12-01](https://neo.sci.gsfc.nasa.gov/view.php?datasetId=MY1DMW_CHLORA&date=2014-12-01)) for January-February 2014 show that Chl-a south of Alaska and along the west coast of the United States varied between 0.3 and 3.0  $\mu\text{g/L}$ . Fujiki *et al.* (2009) also found that Chl-a varied between 0.4 and 1.0  $\mu\text{g/L}$  at six stations south of the Aleutian Islands, Alaska, during a sub-Arctic cruise in autumn 2005. Details about the F10 SSA source function are given in APPENDIX B. In the case of the model uses a source for marine organics, then OA is the sum of SOA, anthropogenic emissions of OM and marine organics. A more detailed analysis of marine organics, focusing on northern Alaska, is presented in 5.2.

### 4.3 Whitecap method

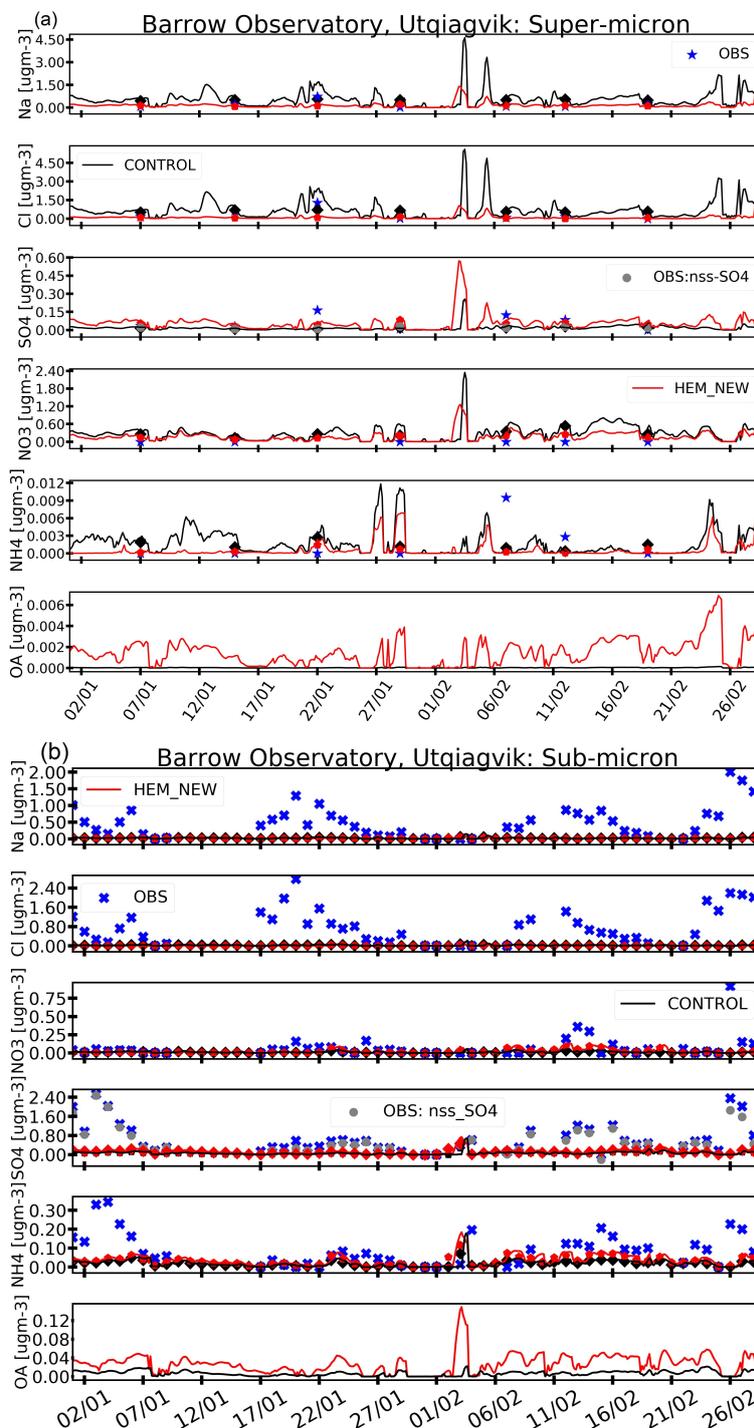
In agreement with previous modelling studies, e.g. (Jaeglé *et al.*, 2011), JA11 from now on and (Spada *et al.*, 2013; Revell *et al.*, 2019; Hartery *et al.*, 2020) the CONTROL simulation produces too much coarse mode and TSP  $\text{Na}^+$  and  $\text{Cl}^-$ . The G97 parametrization, which depends on the whitecap method and thus has a high wind speed dependence (see Eq. 1), has been widely adopted to simulate SSA emissions in global and regional models, e.g. JA11 and Barthel *et al.* (2019). Several studies tried to improve upon the whitecap method ( $W(U_{10})$ ), especially for super-micron SSA. Callaghan *et al.* (2008) used an automated whitecap extraction technique to derive two whitecap expressions that differ from MO80, which are based on cubed relationships for  $U_{10}$ . Other factors, such as the wave field (Salisbury *et al.*, 2013), surfactant (amphiphilic organic material) activity (Callaghan, 2013) and fetch-dependent threshold for breaking waves (Revell *et al.*, 2019; Hartery *et al.*, 2020), have also been shown to affect whitecap lifetime, with implications for SSA production. Goddijn-Murphy *et al.* (2011) analysed Marine Aerosol Production (MAP) whitecap data, in combination with analysis of in-situ and satellite data (from Quick Scatterometer, QuikSCAT) for winds and waves. The satellite data were used to derive an expression with a lower wind speed dependence compared to MO80 (Salisbury *et al.* (2014), SALI14 from now on). Here, the SALI14 parametrization is implemented instead of the MO80 whitecap fraction expression:

$$W(U) = 4.60 \times 10^{-3} \times U_{10}^{2.26} \quad (3)$$

Based on Figure 2 in SALI14, the seasonal mean of  $W(U_{10})$  using Eq. 3 is lower at high latitudes compared to MO80 during autumn and winter. By using this more recent whitecap fraction expression in the quasi-hemispheric simulation, super-micron SSA concentrations decrease overall within the Arctic (not shown here). More specifically, super-micron  $\text{Cl}^-$  and  $\text{Na}^+$  decrease more south of Alaska, by up to 20  $\mu\text{g m}^{-3}$  (Aleutians Islands) and less north of Alaska, by up to 0.5  $\mu\text{g m}^{-3}$ .  $\text{NO}_3^-$  also decreases slightly over continental Alaska, by up to 0.5  $\mu\text{g m}^{-3}$ , due to increased heterogeneous formation on SSA.

### 4.4 SST dependence

Recent data-analysis studies (Saliba *et al.*, 2019; Liu *et al.*, 2021) pointed out that wind speed alone cannot predict SSA variability, and it is important also to include a dependence on SSTs for SSA prediction. Recent modelling studies (Jaeglé *et al.*, 2011; Sofiev *et al.*, 2011; Spada *et al.*, 2013; Barthel *et al.*, 2019) showed that the application of SST dependence improves



**Figure 4.** Evaluation of modelled aerosol composition (runs CONTROL and HEM\_NEW) against in-situ observations at Barrow Observatory, near Utqiagvik, Alaska for (a) super-micron and (b) sub-micron in UTC and in STP conditions. The black line shows model results from the CONTROL run; the red line shows the HEM\_NEW run, while observations are shown as blue crosses. Sub-micron observations are daily averaged and super-micron observations are weekly averages. The corresponding model daily/weekly averages are shown as black diamonds for the CONTROL simulation and as red pentagons for the HEM\_NEW. Observations are shown only when there are available.



simulated SSA concentrations compared to observations. More specifically, previous studies (*Spada et al., 2013; Grythe et al., 2014; Barthel et al., 2019*) tested different SSA source functions, with and without SST dependence, and reported that including such a dependence improves model results, regardless of the SSA source function employed. However, uncertainties still remain about the role of SSTs on SSA production (*Revell et al., 2019*), including the role of other factors, such as seawater composition (*Callaghan et al., 2014*) or wave characteristics (e.g. wave speed and breaking wave type, *Callaghan et al. (2012)*), which might be more important than SSTs alone. In this study, the JA11 SST correction factor is applied when SSTs are between  $-2^{\circ}\text{C}$  and  $30^{\circ}\text{C}$  to evaluate the effect of SST on sub- and super-micron SSA emissions in the Arctic. In our simulations, SSTs are provided by reanalyses data, in this case, FNL, and in the presence of sea-ice, SST is set equal to  $-1.75^{\circ}\text{C}$ .

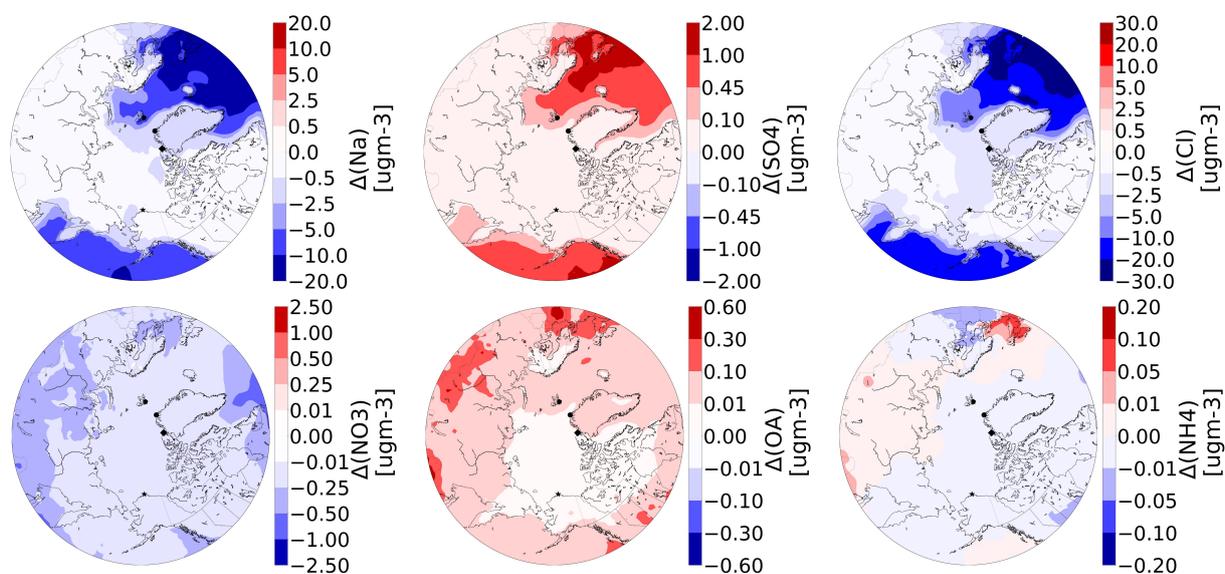
#### 4.5 Sea-salt sulphate

Standard versions of the WRF-Chem model do not include  $\text{ss-SO}_4^{2-}$ . The Community Multiscale Air Quality (CMAQ) model includes a  $\text{ss-SO}_4^{2-}$  component estimating it to be 7% of the total SSA emissions. The mass fraction of  $\text{ss-SO}_4^{2-}$  can be estimated to be 0.25 of the  $\text{Na}^+$  mass fraction (*Kelly et al., 2010; Neumann et al., 2016*) and applied in WRF-Chem to calculate  $\text{ss-SO}_4^{2-}$ . Note that the total fraction of  $\text{Na}^+$ ,  $\text{Cl}^-$ , marine organics and  $\text{ss-SO}_4^{2-}$  is equal to 1.0, and additional mass is not added. The mass fraction of  $\text{ss-SO}_4^{2-}$  is estimated to be 9.9% of the total SSA emissions in our simulations.

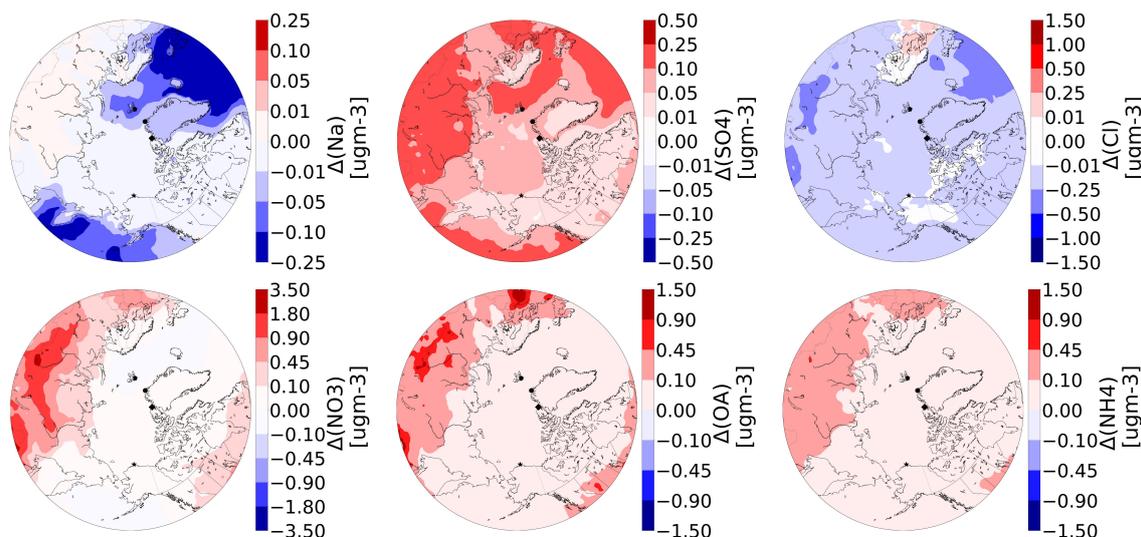
#### 4.6 Discussion

Average differences in aerosol concentrations between the HEM\_NEW and CONTROL simulations are shown in Figures 5 and 6 for January and February 2014 super-micron and sub-micron aerosols, respectively. The updated model simulates less super-micron  $\text{Na}^+$  by up to  $20 \mu\text{g m}^{-3}$ , and  $\text{Cl}^-$  by up to  $30 \mu\text{g m}^{-3}$ , especially south of Alaska and north of the Atlantic Ocean (Fig. 5). These decreases lead to an overall decrease (up to  $2.5 \mu\text{g m}^{-3}$ ) in super-micron  $\text{NO}_3^-$ , over continental and coastal regions and the North Atlantic. This is in agreement with *Chen et al. (2016)* who examined the influence of SSA on  $\text{NO}_3^-$  and reported that overestimation of SSA can lead to an overestimation of super-micron  $\text{NO}_3^-$ , due to formation of  $\text{NO}_3^-$  via heterogeneous uptake of nitric acid ( $\text{HNO}_3$ ) on SSA. Furthermore, due to the addition of  $\text{ss-SO}_4^{2-}$  component in the model, there is more super-micron  $\text{SO}_4^{2-}$ , of up to  $2 \mu\text{g m}^{-3}$ , over marine regions. Super-micron  $\text{NH}_4^+$  also increases (by up to  $0.2 \mu\text{g m}^{-3}$ ) in regions, such as Siberia and North of Europe, coinciding with decreases and increases in  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$ , respectively. Super-micron OA increases by up to  $0.6 \mu\text{g m}^{-3}$  due to the inclusion of marine organics. During winter, the Beaufort Sea, located north of Alaska is covered by sea-ice. Here, the implemented changes lead to smaller decreases in super-micron  $\text{Na}^+$  and  $\text{Cl}^-$  compared to ice-free regions such as the Aleutians islands, e.g., Simeonof site (Fig. 2a) further south. The local effect of sea-ice fraction and open leads on SSA production is examined further in 5.4.

On the other hand, the effect of model updates on sub-micron  $\text{Na}^+$  is smaller, with decreases of up to  $0.25 \mu\text{g m}^{-3}$  south of Alaska and the North Atlantic (Fig. 6) due to use of lower wind speed dependence (SALI14 instead of MO80). The lifetime of SSA, estimated to be between 1 to 4 days over open ocean, in the Arctic and during wintertime (*Rhodes et al., 2017; Xu et al., 2016; Huang and Jaeglé, 2017; Hoppel et al., 2002*), could explain the small decrease of sub-micron  $\text{Cl}^-$  over continental coastal areas (e.g. south of Alaska) in HEM\_NEW. This could also affect long range transport of sub-micron SSA



**Figure 5.** Average differences in super-micron aerosol mass concentrations ( $\mu\text{g m}^{-3}$ ) at the surface between HEM\_NEW and CONTROL during January and February 2014. The black star in northern Alaska shows where Utqiagvik is located. The black circle shows Alert, Canada, the black diamond shows Villum in Greenland, while the black pentagon shows Zeppelin, Svalbard.



**Figure 6.** Average differences in sub-micron aerosol mass concentrations ( $\mu\text{g m}^{-3}$ ) and at the surface between HEM\_NEW and CONTROL during January and February 2014. The black star in northern Alaska shows where Utqiagvik is located. The black circle shows Alert, Canada, the black diamond shows Villum in Greenland, while the black pentagon shows Zeppelin, Svalbard.



from oceanic regions leading to decreases over continental regions, such as northeast United States of America (USA) and Siberia. Sub-micron OA increases by up to  $1.5 \mu\text{g m}^{-3}$  due to inclusion of the F10 parametrization. Note that including ss- $\text{SO}_4^{2-}$  leads to a decrease in  $\text{Na}^+$  and  $\text{Cl}^-$  fractions per bin since no additional mass is added. In contrast to super-micron  $\text{NO}_3^-$ , sub-micron  $\text{NO}_3^-$  increases by  $3.5 \mu\text{g m}^{-3}$  over sources regions and total  $\text{SO}_4^{2-}$  increases due to ss- $\text{SO}_4^{2-}$  component. Also, sub-micron  $\text{NH}_4^+$  slightly increases, showing similar patterns to sub-micron  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$ , probably due to a potential shift in the balance between  $(\text{NH}_4)_2\text{SO}_4$  and  $\text{NH}_4\text{NO}_3$ .

To investigate the variations in modelled  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$  and  $\text{NH}_4^+$ , the mean neutralized factor ( $f$ ) is calculated (not shown here) as the ratio of  $\text{NH}_4^+$  to the sum of  $(2\text{SO}_4^{2-} + \text{NO}_3^-)$ , in molar concentrations, following, for example *Fisher et al. (2011)*, for sites in the Arctic with available observations of these aerosols. When  $f$  is equal to 1 aerosols are more neutralized, while when  $f < 1$  then aerosols are acidic, and more acidic when  $f$  is closer to 0 (*Fisher et al., 2011*). At all sites, except Zeppelin, higher molar concentrations were observed for  $\text{SO}_4^{2-}$  compared to  $\text{NO}_3^-$  and  $\text{NH}_4^+$ . At Utqiagvik, the average observed value of  $f$  is equal to 0.15 for super-micron aerosols, whilst in the model  $f$  decreases from 0.7 to 0.66. This implies that observed super-micron aerosols are more acidic while in the model they are more neutralized (*Fisher et al., 2011*), probably because modelled  $\text{NH}_4^+$  decreased more than  $\text{SO}_4^{2-}$  and  $\text{NO}_3^-$  (Fig. 4a) between the two simulations. There is less super-micron  $\text{NH}_4^+$  in the model than the sum  $2\text{SO}_4^{2-} + \text{NO}_3^-$ , as in observations, however observed  $2\text{SO}_4^{2-} + \text{NO}_3^-$  is much higher than modelled. Observed super-micron aerosols at Utqiagvik are more acidic compared to sub-micron aerosols for which  $f$  equals 0.34. For sub-micron aerosols, HEM\_NEW has an average  $f$  value of 0.08 compared to 0.01 in the CONTROL run. The increase in modelled sub-micron  $f$  could be due to the bigger increase in modelled  $\text{NH}_4^+$  between the two simulations (Fig. 4b). However, in the observations, the higher sub-micron  $f$  is because the sum  $2\text{SO}_4^{2-} + \text{NO}_3^-$  is much higher than  $\text{NH}_4^+$ . At Alert (coarse mode), model  $f$  increases from 0.14 (CONTROL) to 0.19 (HEM\_NEW), with observed  $f$  equal to 0.21, implying that model and observations are acidic, in contrast to Utqiagvik modelled super-micron aerosols. Similar values of  $f$  are found for Zeppelin (coarse mode) and Villum (TSP) (0.12 for CONTROL, 0.13 and 0.18 for HEM\_NEW, respectively) with observed aerosols (0.34 and 0.36 respectively) being less acidic at these sites. Overall the model inorganic aerosols are mostly too acidic compared to the observations. This could be due to underestimation of anthropogenic sources of  $\text{NH}_3$  on the above sites, originating from mid-latitudes. It can be noted that in the model is assumed that all of the aerosol species are internally mixed. However, in reality some of the  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  are observed to be mixed with SSA (KRP18). Based on that, the calculated  $f$  for observations would be biased (too acidic), as some of the  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  are present as  $\text{Na}_2\text{SO}_4$  and  $\text{NaNO}_3$ .

Figures 2 and 3 show the effect of all the modifications (HEM\_NEW) compared to CONTROL and the observations at four Arctic and one sub-Arctic sites. At the two remote high-Arctic sites surrounded by sea-ice (Alert and Villum, Figure 2a,b), HEM\_NEW captures better  $\text{Na}^+$  and  $\text{Cl}^-$  variability, with a small overestimation at Villum (maximum  $0.2 \mu\text{g m}^{-3}$ ). Biases, in  $\mu\text{g m}^{-3}$ , at Alert for  $\text{Na}^+$  and  $\text{Cl}^-$  decrease from 0.81 to 0.12 and from 1.05 to -0.03, respectively. Model results also improve at Villum for  $\text{Na}^+$  and  $\text{Cl}^-$  with biases reduced from 1.3 to 0.25 and from 1.9 to 0.22  $\mu\text{g m}^{-3}$ , respectively. The high variability in SSA at Villum at the end of January and the middle of February is likely to be due to fluctuations in sea-ice fraction around the site, as seen in the FNL sea-ice reanalysis product (varies between 0.93 and 1.0-fully covered, in January and February). Also, HEM\_NEW captures better  $\text{NO}_3^-$  and  $\text{NH}_4^+$  at Alert while slightly overestimates total  $\text{SO}_4^{2-}$  (see APPENDIX C, Table C1). At



Villum, HEM\_NEW captures better  $\text{SO}_4^{2-}$  compared to CONTROL run, slightly underestimates  $\text{NH}_4^+$  end of February, but still overestimates  $\text{NO}_3^-$  (see APPENDIX C, Table C2). Similar results are found for Zeppelin where HEM\_NEW simulates better  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{NO}_3^-$  and  $\text{NH}_4^+$ , but overestimates  $\text{SO}_4^{2-}$ . More specifically, HEM\_NEW slightly underestimates observed  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{NH}_4^+$ , but the model results improve in this site. SSA updates also improve modelled  $\text{NO}_3^-$  (see APPENDIX C, Table C3).

At Simeonof, HEM\_NEW captures better  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{NO}_3^-$  variability during winter 2014 and, due to the inclusion of marine organics, the model simulates more tOC, although it still underestimates observed variability. Calculated biases decrease from 1.4 to 0.3, 2.0 to 0.1, 0.12 to 0.08, -0.08 to -0.05  $\mu\text{g m}^{-3}$  for  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{NO}_3^-$  and tOC, respectively. Also, the addition of ss- $\text{SO}_4^{2-}$  in HEM\_NEW leads to improvements (biases, RMSEs) in simulated  $\text{SO}_4^{2-}$  even if the model occasionally underestimates by up to 0.6  $\mu\text{g m}^{-3}$ . Similar patterns are found for the Gates of the Arctic in northern Alaska.  $\text{Na}^+$  and  $\text{Cl}^-$  are lower in HEM\_NEW while modelled  $\text{NO}_3^-$  and tOC also improve, with biases decreasing for all the four species (0.56 to 0.16, 0.70 to 0.09, 0.26 to 0.18 and -0.24 to -0.21  $\mu\text{g m}^{-3}$  for  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{NO}_3^-$  and tOC, respectively, see also APPENDIX C for RMSEs). HEM\_NEW simulates more total  $\text{SO}_4^{2-}$  at this site but still underestimates the observations, in particular nss- $\text{SO}_4^{2-}$ . Here, the contribution of ss- $\text{SO}_4^{2-}$  is minimal, as shown in Fig. 3b. Thus, the model underestimation could be due to issues related to long-range transport of nss- $\text{SO}_4^{2-}$ , such as wet deposition, or to missing local anthropogenic sources (e.g. Prudhoe Bay oil-fields). Additional wintertime production of  $\text{SO}_4^{2-}$  via mechanisms not requiring sunlight may also contribute. For example, McCabe et al. (2006) suggested that there is secondary  $\text{SO}_4^{2-}$  at Alert during wintertime from metal catalyzed  $\text{O}_2$  oxidation of S(IV) (10–18%). Results from HEM\_NEW also underestimate tOC at Gates of the Arctic, possibly due to underestimation of marine organics (see discussion in next section) or missing regional or remote sources.

Figure 4 compares results from CONTROL and HEM\_NEW with observations for super-micron (weekly averages) and sub-micron (daily averages) aerosols at Utqiagvik. While CONTROL overestimates SSA and  $\text{NO}_3^-$  and underestimates  $\text{SO}_4^{2-}$  (only non-ss- $\text{SO}_4^{2-}$ ), in general, HEM\_NEW captures better observed super-micron  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{NO}_3^-$  and  $\text{NH}_4^+$  aerosols during the simulation period (Fig. 4a) (see also Appendix C). The  $\text{Na}^+$  bias decreases from 0.3 to -0.07  $\mu\text{g m}^{-3}$  but  $\text{Cl}^-$  is now underestimated (bias decreases from 0.27 to -0.26  $\mu\text{g m}^{-3}$ ), due to the introduction of the SST dependence (not shown). Also, there is more super-micron  $\text{SO}_4^{2-}$  in HEM\_NEW and the model slightly underestimates observed  $\text{SO}_4^{2-}$  by about 0.1  $\mu\text{g m}^{-3}$ . Super-micron OA is smaller in magnitude compared to the other aerosol components. However, super-micron OA mass concentration measurements are not available in winter 2014 to evaluate the model. Overall, modelled super-micron SSA concentrations decrease in HEM\_NEW, as at other remote sites (Fig. 2 & 3) in better agreement with the observations compared to the CONTROL run.

On the other hand, while HEM\_NEW (Fig. 4b) represents better periods with low concentrations of sub-micron  $\text{Na}^+$  and  $\text{Cl}^-$  at Utqiagvik in January and February 2014 (up to 0.3  $\mu\text{g m}^{-3}$ ), it still underestimates episodes with very high observed  $\text{Na}^+$  and  $\text{Cl}^-$ , especially at the end of February 2014. The model simulates better  $\text{NO}_3^-$  but underestimates  $\text{NH}_4^+$  and  $\text{SO}_4^{2-}$ , especially at the beginning of January and end of February 2014. Sub-micron OA at 100 km ranges between 0.01 and 0.15  $\mu\text{g m}^{-3}$ . However, observations of OA at Utqiagvik during this period are not available with which to validate the model. Barrett et al. (2015) and Barrett and Sheesley (2017) showed that OC at Utqiagvik is influenced by primary and secondary biogenic carbon and fossil fuel carbon, with air masses originating from the Arctic Ocean, Russian and Canadian Arctic. More



**Table 2.** Calculated fractions of observed and modelled (HEM\_NEW) SSA to total aerosol mass concentrations (summed from available observations at each site). For each site SSA are defined as the sum of  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{ss-SO}_4^{2-}$ . Total is defined as the sum of SSA and inorganic aerosols. Inorganic is the sum of  $\text{nss-SO}_4^{2-}$ ,  $\text{NH}_4^+$  and  $\text{NO}_3^-$  for each station except for Simeonof and Gates of the Arctic where inorganic is the sum of  $\text{nss-SO}_4^{2-}$  and  $\text{NO}_3^-$ . Note that  $\text{NH}_4^+$  is rarely internally mixed within SSA aerosol, because most  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  forms via  $\text{Cl}^-$  (e.g.  $\text{NaCl} + \text{HNO}_3 \rightarrow \text{NaNO}_3 + \text{HCl}$ ). Total\_all below is defined as the sum of SSA,  $\text{nss-SO}_4^{2-}$ ,  $\text{NH}_4^+$ ,  $\text{NO}_3^-$ , BC, OA and dust (model only). The aerosol size for SSA, Total and Total\_all varies per station and corresponds to observed aerosol sizes as described in Section 3.

Sites	SSA/Total [obs]	SSA/Total [HEM_NEW]	SSA/Total_all [HEM_NEW]
Simeonof (fine mode)	0.73	0.84	0.74
Gates of the Arctic (fine mode)	0.20	0.44	0.33
Utqiagvik-sub-micron	0.60	0.22	0.13
Utqiagvik-super-micron	0.93	0.57	0.54
Alert (coarse mode)	0.59	0.54	0.45
Villum (TSP)	0.32	0.63	0.52
Zeppelin (coarse mode)	0.56	0.75	0.62

specifically, *Barrett and Sheesley (2017)* made measurements of OC (diameter less than  $10 \mu\text{m}$ ) collected during winter 2012-2013 northeast of Utqiagvik and reported average OC of  $0.22 \mu\text{g m}^{-3}$ . To compare directly with the model results we divide the modeled value by 1.4 (Fig. 5). In that case, modelled super-micron tOC at Utqiagvik is three times less than the observed OC, showing that the model lack sources of OC. *Shaw et al. (2010)* reported sub-micron OM equal to  $0.43 \mu\text{g m}^{-3}$  during winter 2008 (November to February) at Utqiagvik, almost double the simulated OA at Utqiagvik. Their analysis showed that OM was correlated with organic and inorganic seawater components with the air masses originating along the coastal regions of the Northwest Territories of Canada. Also, the model results can be compared with weekly average sub-micron OM data collected at Alert (*Leitch et al. (2018)* and Fig. 2). At Alert, OM reaches up to  $0.25 \mu\text{g m}^{-3}$  in February 2014, which is almost double compared to the model results for Utqiagvik, North Slope of Alaska and Alert (Fig. 6). At Villum, a recent study by *Nielsen et al. (2019)* showed that OA peaks at  $2.2 \mu\text{g m}^{-3}$  at the beginning (21 to 28 February 2015) of their study period. Their study shows that the majority of OA is mostly due to Arctic Haze influence (up to  $1.1 \mu\text{g m}^{-3}$ ) with secondary influence, due to hydrocarbon-like organics (up to  $1.0 \mu\text{g m}^{-3}$ ) and a marine influence (up to  $0.2 \mu\text{g m}^{-3}$ ). Reasons for these differences on modelled and observed OA are investigated in the next section focusing on regional processes affecting SSA near northern Alaska.

Previous studies (*Quinn et al., 2002; Quinn and Bates, 2005; May et al., 2016; Kirpes et al., 2018, 2019*), pointed out that SSA are an important contributor in the total sub-micron and super-micron mass fraction in the Arctic during wintertime. A recent study by *Moschos et al. (2022)* showed that during wintertime SSA dominates  $\text{PM}_{10}$  (particulate matter with diameter  $\leq 10 \mu\text{m}$ ) mass concentrations at remote Arctic sites, including Alert (56%), Baranov (41%) (Russia), Gruvebadet (44%) (Norway), Utqiagvik (66%), Villum (32%), and Zeppelin (65%). In contrast, at sites such as Tiksi (Russia) and Pallas (Finland),  $\text{SO}_4^{2-}$  and OA dominate (70% and 55%, respectively). To investigate the contribution of SSA to total mass concentrations



during the period of this study, the observed and modelled fraction of SSA to "total" (SSA plus inorganic) aerosols are estimated (see Table 2). However, it should be noted that this fraction varies between sites since not all components were measured.

Overall, taking into account the observations available at each site, the fraction of SSA to total SSA+inorganics is higher at all the coastal sites (Utqiagvik, Alert, Simeonof, Villum) and Zeppelin ranging from 54 to 93%. Only at the Gates of the Arctic and Villum stations the fraction of SSA is smaller (20% and 32%). The modelled HEM\_NEW SSA fraction shows similar patterns (fraction ranges between 44% and 84%) compared to the observations. An exception is sub-micron modelled SSA at Utqiagvik due to low modelled concentrations. When taking into account all aerosol components in the model, including OA, BC and dust, SSA is dominant at Simeonof, Utqiagvik (super-micron), Zeppelin and Villum (more than 54%), whereas at Alert, SSA contributes about 45%. This analysis shows that SSA is an important fraction of total fine mode, super-micron, coarse mode and TSP aerosols in the most Arctic coastal sites during wintertime.

## 5 Regional processes influencing SSA over northern Alaska

In this section, processes which could affect SSA emissions on a regional scale over northern Alaska are examined. In general, the improved model simulates better observed super-micron, TSP, fine and coarse mode SSA,  $\text{NO}_3^-$  and  $\text{SO}_4^{2-}$  aerosols at different sites in the Arctic but the model has difficulties capturing sub-micron SSA during wintertime at Utqiagvik. Possible reasons for these discrepancies are investigated in model runs at 20 km resolution during the campaign periods in January and February 2014 with boundary and initial conditions from HEM\_NEW. The sensitivity of modelled SSA to various processes is examined including aerosol dry deposition over snow/ice, inclusion of local marine organic aerosols, higher wind speed dependence for sub-micron SSA and representation of sea-ice fraction. The possible role of blowing snow and frost flowers is also discussed. Details about the simulations are provided in Table 3.

**Table 3.** WRF-Chem model simulations including details about SSA treatments in the regional runs.

Simulation Name	Description
Regional simulations [20km]	
ALASKA_CONTROL_FEB	Control run for February 2014
DRYDEP_FEB	+ Updated dry deposition velocities over snow-ice and open water on (Zhang et al., 2001) based on (Nilsson et al., 2001)
LOC_ORG_FEB	+ Local source marine organics (Kirpes et al., 2019)
SSA_WS_DEP_FEB	+ Sub-micron SSA wind-speed dependence (Russell et al., 2010)
NEW_ALASKA_FEB	DRYDEP_FEB + LOC_ORG_FEB + SSA_TEST_FEB + ERA5 sea-ice fraction
ALASKA_CONTROL_JAN	Control run for January 2014 (same setup as ALASKA_CONTROL_FEB)
NEW_ALASKA_JAN	including all updates in NEW_ALASKA_FEB

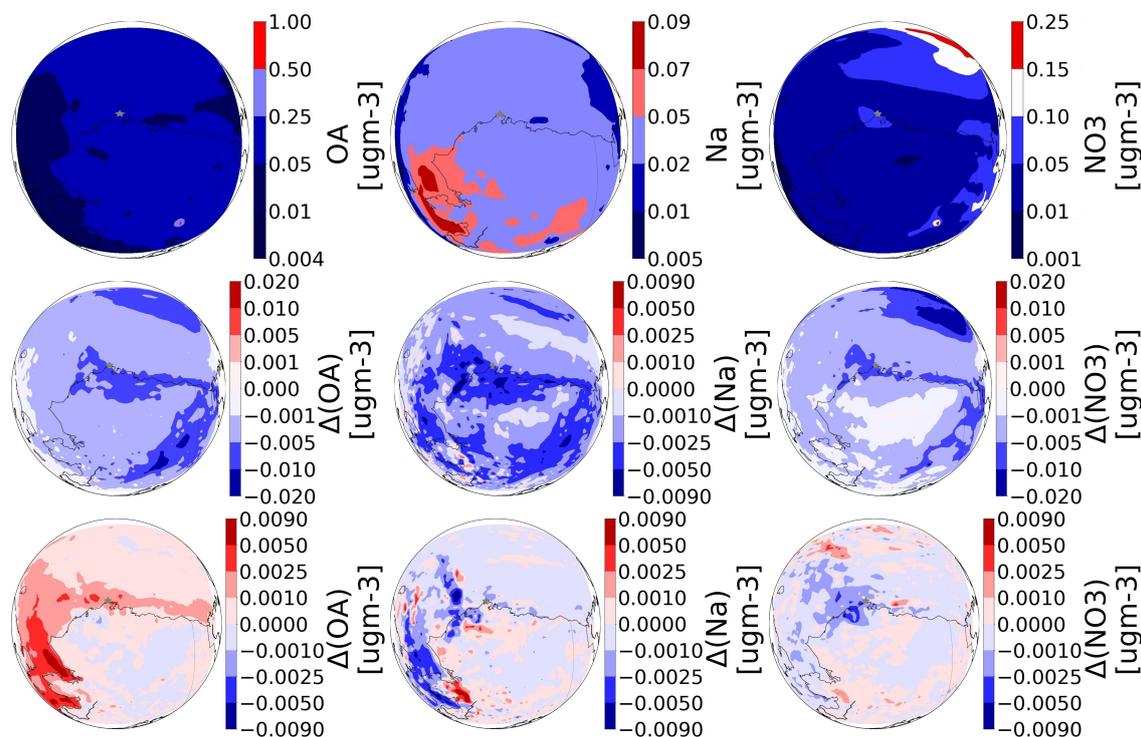


## 5.1 Aerosol dry deposition

Previous studies have shown the importance of including wet and dry removal treatments in models (*Witek et al., 2007; Eckhardt et al., 2015; Whaley et al., 2022*). *Sofiev et al. (2011)* estimated that dry deposition, including sedimentation, could contribute more than 50% to SSA removal, especially for super-micron SSA. JA11, using model treatments for dry deposition  
465 from *Zhang et al. (2001)* over land and *Slinn (1982)* over ocean, reported that the loss of super-micron SSA is dominated by dry deposition. In the quasi-hemispheric simulations, dry deposition velocities are calculated in MOSAIC based on the *Binkowski and Shankar (1995)* parametrization. Here, the *Zhang et al. (2001)* scheme is applied over Alaska, in which the dry deposition velocities are calculated taking into account the different land categories, in contrast to MOSAIC scheme, which uses universal values for processes such as Brownian diffusion and Schmidt number (*Slinn and Slinn, 1980; Slinn, 1982*). *Zhang et al.*  
470 (*2001*) has been used in previous studies, for example, by *Fisher et al. (2011)* and *Huang and Jaeglé (2017)*. These studies applied aerosol dry deposition velocities of  $3.0 \times 10^{-4} \text{ m s}^{-1}$  over snow and ice surfaces for all aerosol diameters and the dry deposition velocity is calculated as a function of aerosol diameter. *Zhang et al. (2001)* includes detailed treatments of deposition processes, such as Brownian diffusion, impaction, interception, gravitational settling and particle rebound, which highly vary depending on land surface type. Certain parameters link to interception, such as collection efficiency by interception, or  
475 impaction processes (e.g. Stokes number) over specific land use categories (such as ice/snow and open ocean), are calculated without considering the radius of surface collectors (*Giorgi, 1988*), but using kinematic viscosity of air, gravitational settling velocity of particle, friction velocity (*Slinn, 1982; Seinfeld, 1986*). Thus, dry deposition velocities over ice/snow and open ocean are set equal to  $3.0 \times 10^{-4}$  and  $1.9 \times 10^{-3} \text{ m s}^{-1}$ , respectively, for both sub- and super-micron aerosols, following *Nilsson and Rannik (2001)*, who reported dry deposition velocity measurements from an Arctic Ocean expedition in 1999. In that way,  
480 the influence of more realistic dry deposition velocities on SSA aerosols is examined during wintertime.

Figures 7 and 8 show the effect of this modification for sub- and super-micron SSA and  $\text{NO}_3^-$ , respectively (differences between DRYDEP\_FEB and ALASKA\_CONTROL\_FEB runs). Sub-micron  $\text{Na}^+$ , OA and  $\text{NO}_3^-$  decrease very slightly, whereas super-micron  $\text{Na}^+$ , OA and  $\text{NO}_3^-$  increase by up to 0.6, 0.02 and  $0.3 \mu\text{g m}^{-3}$ , respectively, with the largest increase over sea-ice areas or regions with snow cover. These changes in modelled sub- and super- micron aerosols are due to differences between  
485 the dry deposition velocities in the two schemes. Over model grids covered with snow or ice and open ocean MOSAIC dry deposition velocities are smaller (larger) for sub-micron (super-micron) in magnitude compared to reported velocities by *Nilsson and Rannik (2001)*. During wintertime over northern (in-land) Alaska, all the grid cells during the simulations are snow covered. Based on these results, and the fact that super-micron  $\text{Na}^+$  and  $\text{Cl}^-$  are slightly underestimated at 100 km and Utqiagvik (see section 4.5), the following simulations use the observed dry deposition velocities reported by *Nilsson and Rannik (2001)*.

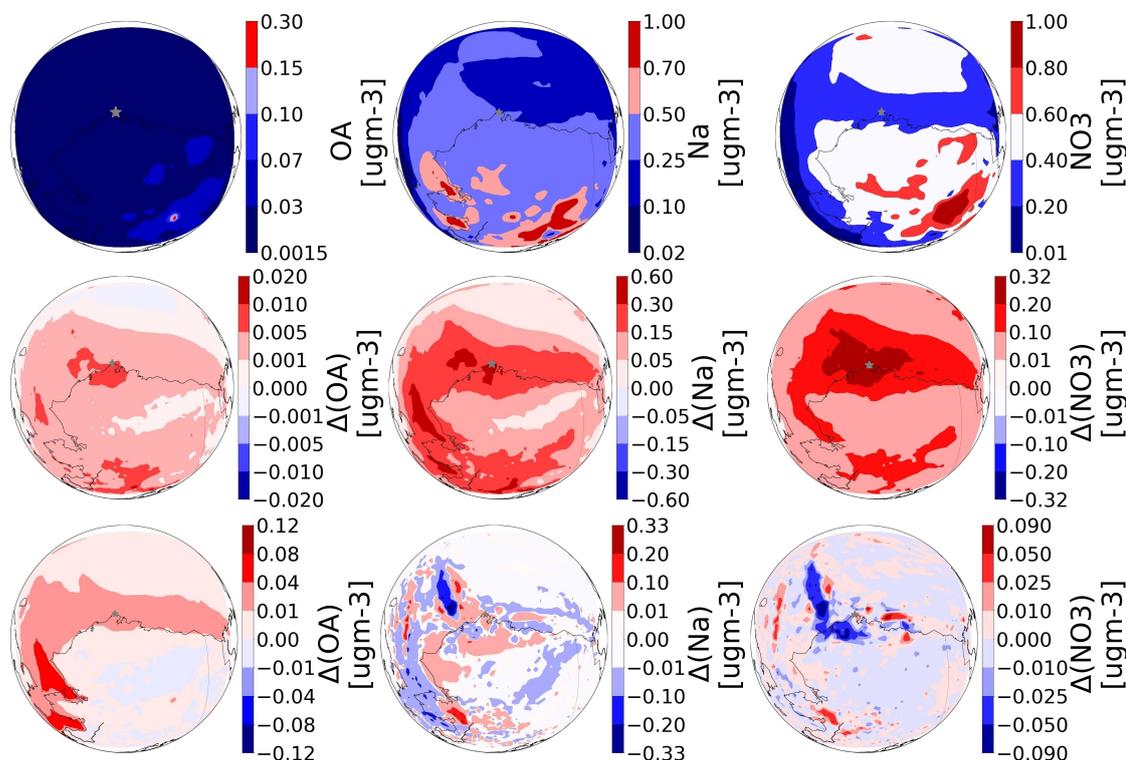
490 These results show that sub- and super-micron (mostly) SSA are sensitive to different dry deposition parametrization in WRF-Chem. To address potential uncertainties in dry removal treatments and their influence on SSA regionally, a series of sensitivity tests are also performed. Firstly, correct modelling of aerosol dry deposition depends on the ability of the model to capture the structure of the Arctic boundary layer including vertical temperatures and winds. Model results at 20 km and 100 km horizontal resolution are compared against hourly in-situ 2 m, 10 m temperatures and 10 m wind speeds or temperature



**Figure 7.** First row shows the average values of aerosol mass concentrations for sub-micron during February campaign. Average differences at the surface between DRY\_DEP\_FEB and ALASKA\_CONTROL\_FEB (second row), and between LOC\_ORG\_FEB and DRY\_DEP\_FEB (third row) during February campaign for sub-micron  $\text{Na}^+$ , OA,  $\text{NO}_3^-$  ( $\mu\text{g m}^{-3}$ ). See text and Table 2 for detailed description of the model runs.

495 and wind speed profiles up to 4 km for January and February 2014 (see figures in APPENDIX D with calculated bias and RMSE for the two periods). Observed wind speeds during January ranged between 4.7 and 14.1  $\text{ms}^{-1}$  and wind directions were mostly easterly (77 to 135 degrees). During February, wind speeds ranged between 0.4 and 13.3  $\text{ms}^{-1}$  and wind directions were mostly easterly, except from 22 UTC 25 February to 11 UTC 26 February when the winds were westerly. In general, the model performs well at 20 km, and better than at 100 km, in terms of temperature and winds, although it slightly underestimates  
 500 observations at the surface. On the other hand, there are small discrepancies, of up to 10 degrees, between modelled (at 20 km) and observed wind direction at the Barrow site, near Utqiagvik town, except at 26 February when these discrepancies are up to 70 degrees.

To examine further causes of variability in modelled dry removal of SSA, a sensitivity test is carried out where aerosol dry deposition and gravitational settling are switched off during a windy day, 28 February 2014. On this day, 10 m wind speeds at  
 505 Utqiagvik varied between 7 and 13.5  $\text{m s}^{-1}$  and were easterly (104 to 130 degrees). This corresponds to a period when observed sub-micron  $\text{Na}^+$  and  $\text{Cl}^-$  concentrations were high, around 1.4 and 2.0  $\mu\text{g m}^{-3}$ , respectively (see Figure 10b). During this day the model captures quite well observed wind speeds and directions, with small differences of up to 2  $\text{m s}^{-1}$  and up to 10 degrees



**Figure 8.** First row shows the average values of aerosol mass concentrations of super-micron during February campaign. Average differences at the surface between DRY\_DEP\_FEB and ALASKA\_CONTROL\_FEB (second row), and between LOC\_ORG\_FEB and DRY\_DEP\_FEB (third row) in super-micron  $\text{Na}^+$ , OA,  $\text{NO}_3^-$  ( $\mu\text{g m}^{-3}$ ) during February campaign. See text and Table 2 for detailed description of the model runs.

differences in wind direction (not shown here). In this sensitivity run, the model simulates more super-micron SSA (an increase up to  $0.8 \mu\text{g m}^{-3}$ ), which is expected due to the influence of gravitational settling on super-micron particles. The increase on sub-micron SSA is smaller. However, observations of dry deposition of different aerosols are needed to better constrain the model in the Arctic.

## 5.2 Local source of marine organics

For the simulations at 100 km, the F10 parametrization is used based on C:Chl-a from a cruise at mid-latitudes. Whilst phytoplankton blooms may not be expected in the high Arctic winter, previous studies have shown evidence of sea ice biological activity under low light conditions in the Arctic (Krembs et al., 2002; Lovejoy C., 2007; Hancke et al., 2018). In addition, Russell et al. (2010) (from now on RUS10) analysed samples from the International Chemistry Experiment in the Arctic Lower Troposphere (ICEALOT) cruise and found that most organic mass in clean regions of the North Atlantic and the Arctic is composed of carbohydrate-like compounds containing organic hydroxyl groups from primary ocean emissions. Frossard et al.



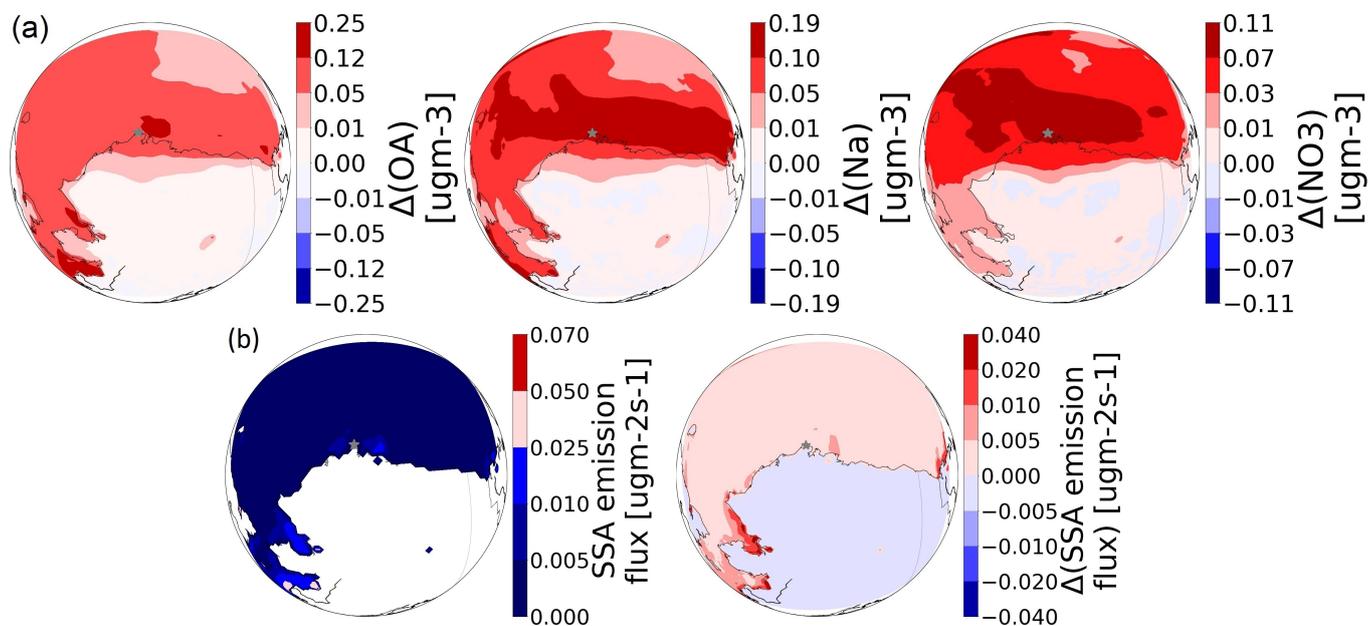
(2014) (FRSS14 from now on) investigated the sources and composition of atmospheric marine aerosol particles based on the  
520 analysis of samples, including those from ICEALOT, reporting that ocean-derived organic particles include primary marine or-  
ganic aerosols. In particular, they calculated the ratio of OC:Na<sup>+</sup> as a metric for comparing the composition of model-generated  
primary marine aerosol and seawater, previously used by *Hoffman and Duce (1976)*, and reported OC:Na<sup>+</sup> ratios of 0.45 for  
atmospheric marine aerosol particles. KRP19 also reported that during their campaigns in 2014 almost all the individual SSAs  
525 had thick organic coatings (average C:Na mole ratios of 0.5 and 0.3 for sub-micron and super-micron SSA, respectively) made  
up of marine saccharides. They also identified open sea ice leads enriched with exopolymeric substances as contributing to  
organics in winter SSA.

Here, elemental fractions for sub- and super-micron aerosols sampled during the KRP19 campaigns are used to better con-  
strain modelled organic marine emissions (mOC). More specifically, the ratio of sub- and super-micron OC:Na<sup>+</sup> is calculated,  
following FRSS14 and using the elemental fractions from KRP19, as an indicator of the presence of a local source of marine  
530 organics. The organic fraction of the total SSA for the high organic activity scenario in WRF-Chem is increased to 0.4 for sub-  
micron (1st to 5th MOSAIC 8-bins) and to 0.11 for super-micron (6th to 8th MOSAIC 8-bins). Note again that no additional  
SSA mass is added. Figures 7 & 8 show the sensitivity of the model results to including a larger marine organic fraction. Sub-  
and super-micron OA concentrations increase on average by up to 0.009 and 0.12  $\mu\text{g m}^{-3}$ , respectively, especially south-west  
of Alaska and along coastal areas, including Utqiagvik. Sub-micron Na<sup>+</sup> and NO<sub>3</sub><sup>-</sup> slightly decrease (0.005  $\mu\text{g m}^{-3}$ ) around  
535 Utqiagvik region, and super-micron Na<sup>+</sup> and NO<sub>3</sub><sup>-</sup> decrease north-west of Utqiagvik.

KRP19 reported measured sub-micron organic carbon volume fractions based on analysis of 150 SSA particles between 0.3  
 $\mu\text{m}$  and 0.6  $\mu\text{m}$  comparable to organic carbon volume fractions observed in SSA produced in mid-latitude algal bloom regions.  
This suggests the presence of significant organic carbon associated with locally produced SSA on the coast of northern Alaska.  
There are two available daily observations at Gates of the Arctic during the February campaign. The model captures better  
540 observed tOC at the end of February in the run LOC\_ORG\_FEB with higher organic fraction. However, it underestimates tOC  
on 25 February when the observed tOC reached 0.33  $\mu\text{g m}^{-3}$  (see APPENDIX F). As mentioned previously, this discrepancy  
could also be due to missing local anthropogenic sources related to North Slope oil field emissions (*Gunsch et al., 2020*).  
In the following runs, marine organics based on the calculated ratio of OC:Na<sup>+</sup> are included instead of F10 considering the  
importance of local SSA marine sources at Utqiagvik (KRP19). By including a local source of marine organics in the model,  
545 this leads to a better agreement with the findings of the previous studies discussed in section 4.6. It can be noted that there  
are only sporadic measurements of OA/OC at remote Arctic sites and detailed long-term observations are not available which  
might help to better constrain model simulations.

### 5.3 Wind-speed sensitivity to sub-micron SSA emissions

In the regional runs presented so far, the lower wind speed dependence based on satellite data is used (SALI14) since it  
550 improves modelled SSA compared to observations in the 100 km runs (see in section 4.2). However, RUS10 found evidence  
for higher wind speed dependence during the ICEALOT cruise in the Arctic. They found that wind speed is a good predictor of  
a marine factor, calculated using positive matrix factorization, for sub-micron organic mass (OM1<sub>sea</sub>). Their analysis showed a



**Figure 9.** Average differences in mass concentrations of (a) sub- micron  $\text{Na}^+$ , OA,  $\text{NO}_3^-$ , in  $\mu\text{gm}^{-3}$ , at the surface between SSA\_WS\_DEP\_FEB and LOC\_ORG\_FEB. Grey star indicates the location of Utqiagvik. (b) The map on the left shows the average value of SSA emission fluxes in  $\mu\text{gm}^{-2}\text{s}^{-1}$  during February campaign and the map on the right shows average differences between SSA\_WS\_DEP\_FEB and LOC\_ORG\_FEB in  $\mu\text{gm}^{-2}\text{s}^{-1}$ .

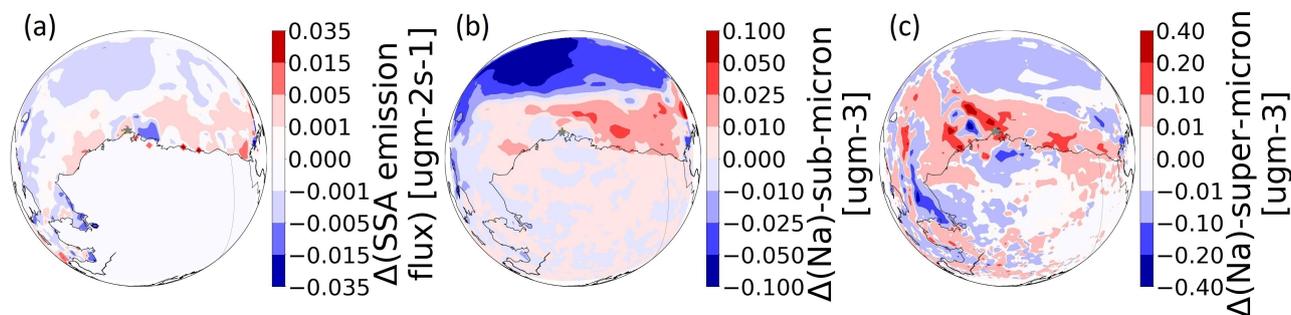
high correlation between  $\text{OM1}_{\text{sea}}$ , sub-micron sodium ( $\text{Na}^+$ ) concentration and wind speed at 18 m (correlation  $r$  equal to 0.90 for the North Atlantic and Arctic region, see Table S3 Supplementary Material in RUS10). Average  $\text{OM1}_{\text{sea}}$  concentrations (0.2  $\mu\text{g m}^{-3}$ ) reported by RUS10 for the eastern Arctic Ocean are about half those reported at Utqiagvik by *Shaw et al. (2010)* during wintertime.

In a sensitivity simulation, the results from RUS10 are used to include a higher wind speed dependence for sub-micron SSA. Equations (5) and (6) from the RUS10 analysis for the Arctic legs of their cruise are applied to the model as a correction factor:

$$\text{Na}^+1 = 0.022 \times U_{18} - 0.012 \quad (4)$$

$$\text{OM1}_{\text{sea}} = 0.025 \times U_{18} - 0.049 \quad (5)$$

where  $U_{18}$  is wind speed at 18 m in  $\text{ms}^{-1}$ , ranging between 2 and 14  $\text{m s}^{-1}$  (Figure 2, RUS10). RUS10 used  $\text{Na}^+1$  as a proxy for sub-micron NaCl, and subsequently SSA, because  $\text{Na}^+1$  equalled sub-micron  $\text{Cl}^-1$  on a molar basis for the North Atlantic and Arctic sampling regions. Thus, Equation (5) is also used to estimate a correction factor for  $\text{Cl}^-$ . Here, wind speeds in the first model layer are used, i.e. around 26 m. Differences in  $U_{18\text{m}}$  and  $U_{26\text{m}}$  reach a maximum of 1  $\text{m s}^{-1}$  (see Fig.D1 in APPENDIX D). Following RUS10, correction factors are only applied in the model to the number and mass of the SSA emissions when



**Figure 10.** Average differences between ALASKA\_NEW\_FEB and SSA\_WS\_DEP\_FEB showing the effect of switching from FNL to ERA5 sea-ice fractions during February for (a) SSA emission fluxes ( $\mu\text{g m}^{-2}\text{s}^{-1}$ ), (b) sub-micron mass concentration of  $\text{Na}^+$  and (c) super-micron mass concentration of  $\text{Na}^+$  in  $\mu\text{g m}^{-2}\text{s}^{-1}$ . The grey star shows the location of Utqiagvik.

modelled wind speeds are between 2 and 14  $\text{ms}^{-1}$ , and when RUS10-calculated sub-micron SSA concentrations are greater than model calculated SSA. In this way, SSA emissions are enhanced during periods of higher wind speeds.

To illustrate the sensitivity of the results to applying this correction, Fig. 9 shows sub-micron aerosol concentrations and SSA emission fluxes, the latter being the sum of dry mass emissions calculated in the model. Overall, this leads to an increase of 0.25, 0.19 and 0.11  $\mu\text{g m}^{-3}$  in sub-micron  $\text{Na}^+$ ,  $\text{NO}_3^-$  and OA, respectively, over the Utqiagvik region and southwest Alaska during the February campaign. The SSA emission flux is influenced directly by the area in the model grid which is ice-free. This leads to SSA production east and west of Utqiagvik while the highest values are southwest of Alaska. By adding the RUS10 correction, SSA emission fluxes increase slightly by up to 0.035  $\mu\text{g m}^{-2}\text{s}^{-1}$  along the southwest Alaskan coast, and by up to 0.015  $\mu\text{g m}^{-2}\text{s}^{-1}$  around Utqiagvik. RUS10 showed that sub-micron SSA and wind speed are well correlated over open ocean. Thus, a correction factor to sub-micron SSA, based on is-situ data, improves sub-micron model SSA and could be included in future simulations for studies focusing on the Arctic region.

#### 5.4 Sea-ice fraction

The sensitivity of modelled SSA to prescribed sea-ice fractions during wintertime and the role of leads, is investigated since KRP19 already pointed out the importance of using realistic sea-ice to simulate marine aerosols. High spatial resolution images of sea-ice cover are available, including during the Polar Night, from a marine radar operating on top of a building in downtown Utqiagvik (71°17'13" N, 156°47'17" W), 22.5 m above sea level, with a range of up to 11 km to the northwest ([http://sealice.alaska.edu/gi/data/barrow\\_radar](http://sealice.alaska.edu/gi/data/barrow_radar)) (Druckenmiller et al., 2009; Eicken et al., 2011). May et al. (2016) showed increased super-micron  $\text{Na}^+$  mass concentrations during periods of elevated wind speeds and lead presence, in a multiyear study using the sea ice radar data at Utqiagvik. Between 23-28 January 2014, when the winds at Barrow observatory were easterly, the radar showed that the coastal area east of Utqiagvik featured leads (KRP19). From 24-28 February 2014, the west coastal area also featured leads as shown by Moderate Resolution Imaging Spectroradiometer (MODIS) satellite images (KRP19). To examine the sensitivity of SSA emissions to sea-ice cover, ERA5 sea-ice fractions with a resolution of  $0.25^\circ \times 0.25^\circ$  are used instead



of FNL fraction at  $1.0^\circ \times 1.0^\circ$  resolution. Note that only sea-ice fraction field is different, while the rest of the meteorological fields are from FNL.

590 Results for February are shown in Fig. 10. The SSA emission flux (Fig. 10a) increases over a small region around Utqiagvik and across the North Slope of Alaska due to decreased sea-ice fraction but decreases east of Utqiagvik and southwest of Alaska (e.g. Selawik Lake and Norton Bay) due to increased sea-ice fraction. Sub-micron  $\text{Na}^+$  slightly increases along the north coast of Alaska by up to  $0.1 \mu\text{g m}^{-3}$  and around Utqiagvik (see Fig. 10b) and super-micron  $\text{Na}^+$  increases by up to  $0.4 \mu\text{g m}^{-3}$  around Utqiagvik and decreases by up to  $0.4 \mu\text{g m}^{-3}$  southwest of Alaska (Fig. 10c). The model results for January indicate that there  
595 is less sea-ice in the region around Utqiagvik and south west of Alaska. Therefore, higher SSA emission fluxes were simulated for February ( $0.035 \mu\text{g m}^{-2} \text{s}^{-1}$ ) compared to January ( $0.015 \mu\text{g m}^{-2} \text{s}^{-1}$ ) (maps not shown here).

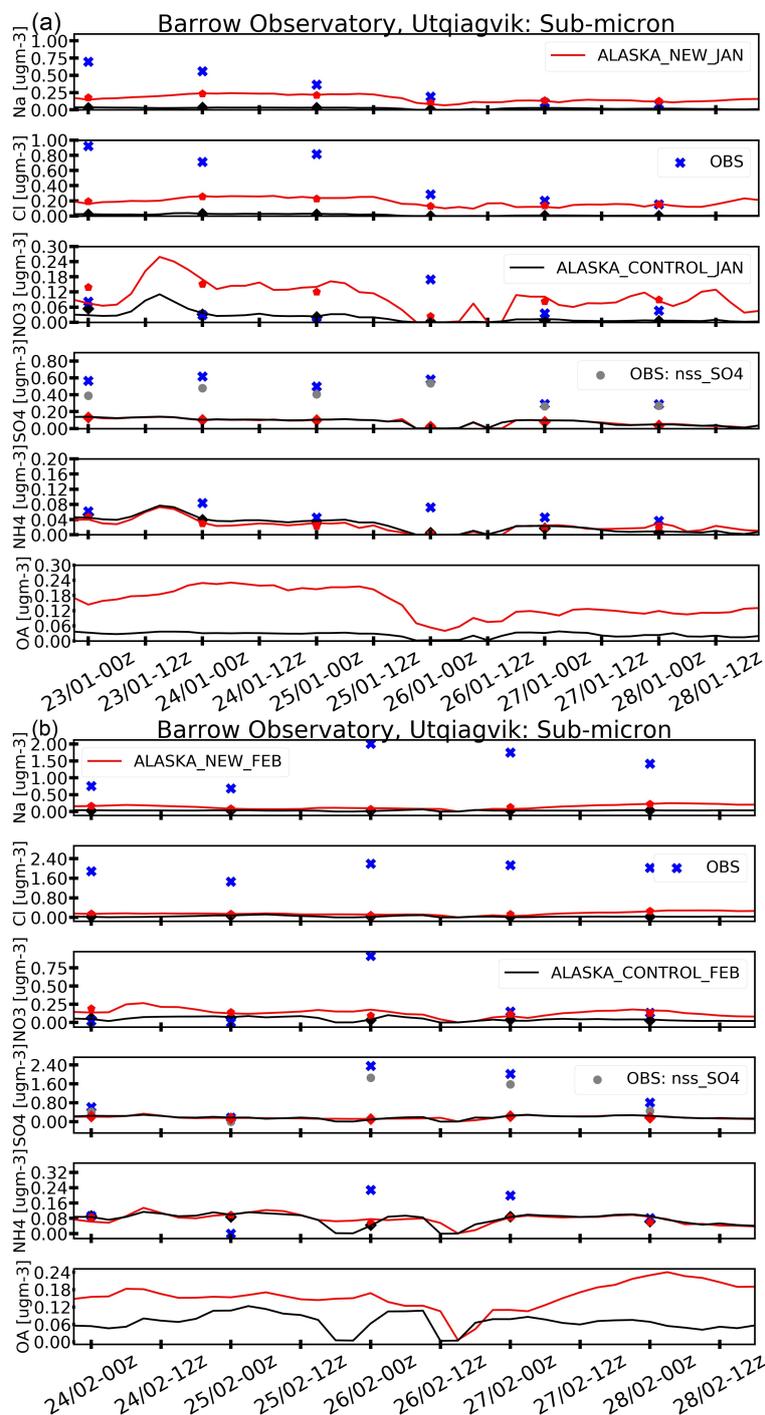
Two further simulations are performed to examine model sensitivity to sea-ice fraction. First, ERA5 sea-ice fractions are set equal to 0 (ice-free conditions) to the north, west, and east of Utqiagvik to examine the effect of having ice-free conditions and the presence of open leads locally (as seen by the radar). Second, ERA5 sea-ice fractions are set equal to 0.75 north, west, east  
600 of Utqiagvik and northwest of Alaska. In both cases, the model is run on a windy day (28 February 2014). The first sensitivity test leads to an increase in SSA emission fluxes by up to  $0.2 \mu\text{g m}^{-2} \text{s}^{-1}$  where sea-ice fraction equals zero (not shown) and to an increase of up to  $1.2 \mu\text{g m}^{-3}$  and  $0.05 \mu\text{g m}^{-3}$  in super-micron and sub-micron  $\text{Na}^+$  respectively. The second sensitivity test yields similar results. This is because ERA5 sea-ice fractions are higher than the test case (0.75) leading to an overall increase in the SSA emission flux of up to  $0.02 \mu\text{g m}^{-2} \text{s}^{-1}$ , especially east of Utqiagvik, affecting primarily super-micron SSA  
605 (increases of up to  $1.5 \mu\text{g m}^{-3}$ ) rather than sub-micron SSA, probably due to the short simulation period.

These results illustrate the sensitivity of super-micron SSA, in particular, to the prescribed sea-ice fraction and point out the need of improving this in models. Regarding sub-micron SSA, which is less sensitive to local sea-ice in these model simulations, there is the possibility that missing mechanisms influencing sub-micron SSA emissions need to be included such as SSA production from breaking waves in the surf zone for particles with diameters between  $1.6$  and  $20 \mu\text{m}$  (De Leeuw *et al.*,  
610 2000) or diameters ranging between  $0.01$ - $0.132$  (ultrafine),  $0.132$ - $1.2$  and  $1.2$ - $8.0 \mu\text{m}$  (Clarke *et al.*, 2006), which would be important and in the ice free ocean.

## 5.5 Evaluation against observations in northern Alaska

The model is also run for January 2014 including all the updates described above (see Table 2 and section 2.3). Fig. 11 shows the comparison between runs with and without all these updates compared to sub-micron aerosol observations at Utqiagvik  
615 during the January and February campaigns. Note that the focus of this section is on sub-micron SSA, as there are not detailed super-micron observations during the periods of the simulations due to their weekly temporal variation and due to the fact that the model still underestimates observed sub-micron SSA at Utqiagvik.

There are differences in the observations between the two periods. While sub-micron observed  $\text{Na}^+$  and  $\text{Cl}^-$  did not exceed  $1 \mu\text{g m}^{-3}$  during January, observed sub-micron  $\text{Na}^+$  and  $\text{Cl}^-$  concentrations reached up to  $2.5 \mu\text{g m}^{-3}$  in February. As noted earlier  
620 such high concentrations of  $\text{Na}^+$  and  $\text{Cl}^-$  were not observed at Alert and Villum during January and February 2014. This could be explained by the fact that these two sites are entirely surrounded by more sea-ice in winter. Overall, the model simulates



**Figure 11.** Time-series during a) January and b) February 2014 of sub-micron mass concentrations of  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{NO}_3^-$ ,  $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$ , in  $\mu\text{g m}^{-3}$ , simulation period. Model simulations are validated against in-situ sub-micron aerosols at Utqiagvik, Alaska, in UTC (every 12h; 00z,12z). The black line shows model results from the CONTROL run; the red line shows the ALASKA\_NEW run, while the daily observations are shown in blue crosses. The corresponding model daily averages are shown as black diamonds for the control simulation and as red pentagons for the ALASKA\_NEW runs. See the text for details about the observations and model runs.



better observed sub-micron  $\text{Na}^+$  and  $\text{Cl}^-$  in January but still underestimates concentrations by up to 0.3 and  $0.6 \mu\text{g m}^{-3}$ , respectively, while sub-micron  $\text{NO}_3^-$  is slightly overestimated. Biases in January decrease from -0.31 to -0.16, -0.50 to -0.33 and -0.04 to  $0.039 \mu\text{g m}^{-3}$ , respectively. On the other hand, sub-micron  $\text{Na}^+$ ,  $\text{Cl}^-$  are still underestimated in the run including all the updates (ALASKA\_NEW\_FEB) by up to  $2.0 \mu\text{g m}^{-3}$  in February indicating that there are missing processes in the model linked to sub-micron SSA emissions, as discussed earlier. However, overall the results at Utqiagvik in February, including all the updates (ALASKA\_NEW\_FEB), are better compared to the control simulation (ALASKA\_CONTROL\_FEB). Biases for  $\text{Na}^+$ ,  $\text{Cl}^-$  and  $\text{NO}_3^-$  decrease from -1.29 to -1.18, -1.90 to -1.78 and -0.20 to  $-0.11 \mu\text{g m}^{-3}$ , respectively. During both months, the model lacks  $\text{SO}_4^{2-}$  due to missing local anthropogenic sources, as discussed in section 4.6 and due to small contribution from ss- $\text{SO}_4^{2-}$  as is shown also in section 4.1 for different Arctic sites. Missing aqueous phase reactions, such as the oxidation of  $\text{SO}_2$  by ozone in alkaline SSA aerosols (Alexander *et al.*, 2005) and  $\text{SO}_4^{2-}$  production from metal catalyzed  $\text{O}_2$  oxidation (McCabe *et al.*, 2006) are missing from the model and might explain these high discrepancies compared to sub-micron observations at Utqiagvik. Also, the variations in modelled  $\text{NH}_4^+$  between the ALASKA\_CONTROL and ALASKA\_NEW for January and February simulations are small. The model underestimates observed  $\text{NH}_4^+$ , which peaks at  $0.2 \mu\text{g m}^{-3}$  during February. Calculated biases and RMSEs for all aerosol species and for January and February campaigns are given in APPENDIX F.

Comparison with data from Gates of the Arctic (see APPENDIX E) shows that there are not significant differences between the control run and including all the updates in February 2014. The model still underestimates observed tOC due to missing local anthropogenic sources and overestimates  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$ ,  $\text{Na}^+$  and  $\text{Cl}^-$ . However, due to short period of the simulation only two observations are available, thus more detailed observations are needed to examine further the reason why the model differs from the observations at this site.

## 5.6 Are blowing snow and/or frost flowers a source of sub-micron SSA during wintertime at Utqiagvik?

Lastly, we consider whether enhanced SSA at Utqiagvik could be due to blowing snow or frost flower sources. We noted earlier that KRP19 found no evidence of blowing snow or frost flowers at this site but that SSA originated from open leads during wintertime. The findings of KRP19 are supported by the earlier laboratory study of Roscoe *et al.* (2011) who reported that frost flowers are not an efficient source of SSA. However, an older study by Shaw *et al.* (2010) found that during winter at Utqiagvik surface frost flowers formed on the sea and lake ice are a source of ocean derived OM. Modelling studies that have included a source of blowing snow and frost flowers suggest that they are contributing to SSA at this time of year at Utqiagvik, Alert and Zeppelin (Xu *et al.*, 2013, 2016; Huang and Jaeglé, 2017; Rhodes *et al.*, 2017).

To investigate whether blowing snow or frost flowers could also be a source of SSA during the campaigns at Utqiagvik, depletion factors are estimated following Frey *et al.* (2020). Frey *et al.* (2020) reported that blowing snow was the main source of SSA rather than frost flowers and open-leads in Antarctic wintertime, based on  $\text{SO}_4^{2-}$  and  $\text{Br}^-$  depletion in SSA being indicative of blowing snow origin, and not sea water. Here, depletion factors are calculated using modelled and observed sub-micron aerosol concentrations, during the campaign periods. More specifically,  $\text{SO}_4^{2-}$  depletion relative to  $\text{Na}^+$  ( $\text{DF}_{\text{SO}_4^{2-}}$ ),  $\text{Na}^+$  depletion relative to  $\text{Cl}^-$  ( $\text{DF}_{\text{Na}^+}$ ) and  $\text{Br}^-$  depletion relative to  $\text{Na}^+$  ( $\text{DF}_{\text{Br}^-}$  only for observations in this case) are calculated using the following equation:



**Table 4.** Average sub-micron modelled and observed depletion factors, following *Frey et al. (2020)*, during the January and February campaigns 2014 in Utqiagvik. Model results for ALASKA\_NEW\_JAN and ALASKA\_NEW\_FEB simulations are shown here, respectively. Observations refer to sub-micron data from NOAA. See text for details.

Depletion Factors	Model	Observations
January campaign		
$DF_{SO_4^{2-}}$	-0.77	-7.56
$DF_{Na^+}$	-1.05	-0.09
February campaign		
$DF_{SO_4^{2-}}$	-4.8	-2.15
$DF_{Na^+}$	-1.1	-0.19
$DF_{Br^-}$	-	0.063

$$DF_x = 1 - \frac{R_{\text{smpl}}}{R_{\text{RSW}}} \quad (6)$$

where, R is the mass ratio (x:y) in the model or in the sample (smpl) and in reference seawater (RSW) (*Millero et al., 2008*). A depletion factor ( $DF_x$ ) between 0 (small) and 1 (strong) indicates 0–100% depletion, whereas  $DF_x$  smaller than 0 indicates enrichment. *Frey et al. (2020)* suggested, based on depletion of  $SO_4^{2-}$  relative to  $Na^+$ , that most SSA originates from blowing snow on sea-ice with minor contributions from frost flowers, and not from open leads.

Average values of modelled and observed DFs are shown in Table 4. In January, observed  $SO_4^{2-}$  concentrations are 8.56 times more than in reference seawater, possibly due to internal mixing with anthropogenic  $SO_4^{2-}$  from NSA oilfield emissions (KRP18), whilst in the model,  $SO_4^{2-}$  concentrations are 1.77 times higher than in reference seawater, showing enrichment in both cases (Table 4). Modelled and observed depletion factors also show enrichment in February. This is in contrast with results from *Frey et al. (2020)* who reported substantial depletion. They also reported a case of enrichment due to possible contamination from the ship, an anthropogenic source. Our modelling results and the observations at Utqiagvik indicate enrichment of  $SO_4^{2-}$  relative to  $Na^+$ , suggesting that blowing snow and frost flowers are not a source of SSA, at least during these campaigns. Previous studies (*Douglas et al., 2012; Jacobi et al., 2012*) suggested that blowing snow and frost flowers near Utqiagvik are characterised by  $SO_4^{2-}$  depletion compared to seawater.  $Na^+$  depletion relative to  $Cl^-$  during both campaigns also shows enrichment, albeit more negligible in the observations than in the model. Observed  $Na^+$  depletion relative to  $Cl^-$  is 1.09 or 1.19 times more than in reference seawater, during January and February, respectively.

SSA can also play an important role in polar tropospheric ozone and halogen chemistry through the release of active bromine during spring (*Fan and Jacob, 1992; Simpson et al., 2007; Peterson et al., 2017*). Reactions involving bromine are an important sink of ozone ( $O_3$ ) (*Barrie, 1986; Barrie et al., 1988; Wang et al., 2019a; Marelle et al., 2021*) and also cause mercury oxidation (*Schroeder et al., 1998*).  $Br^-$  depletion relative to  $Na^+$  is calculated only during February, since observed  $Br^-$  was zero during the period of January campaign, and indicates a small depletion in reference seawater. The calculated observed mass ratio of



**Table 5.** Average modelled and observed molar ratios for sub-micron SSA, following *Kirpes et al. (2019)*, during January and February campaign 2014 in Utqiagvik. Model results from ALASKA\_NEW\_JAN and ALASKA\_NEW\_FEB simulations are used. Observations refer to sub-micron data from NOAA.

Molar ratios	Model	Observations
January campaign		
$\text{SO}_4^{2-}:\text{Na}^+$	0.11	0.55
$\text{Cl}^+:\text{Na}^+$	0.74	1.1
February campaign		
$\text{SO}_4^{2-}:\text{Na}^+$	0.37	0.2
$\text{Cl}^+:\text{Na}^+$	0.8	1.08

Br<sup>-</sup> to Na<sup>+</sup>, based on the available observations of Br<sup>-</sup> during February, indicates a seawater origin. The observed mass ratio of Br<sup>-</sup> to Na<sup>+</sup> ranges between 0.0057 and 0.0059, while the mass ratio of Br<sup>-</sup> to Na<sup>+</sup> in reference seawater is equal to 0.006. On the other hand, *Frey et al. (2020)* reported no or little Br<sup>-</sup> depletion relative to Na<sup>+</sup> due to Br<sup>-</sup> loss at the surface and small  
680 depletion further aloft. For a more comprehensive analysis, observations are required at different locations and altitudes across coastal northern Alaska.

We note that the version of WRF-Chem used in this study does not include halogen chemistry. It has since been implemented in a later version by *Marelle et al. (2021)* to examine ozone depletion events during March-April 2012 at Utqiagvik. Heterogeneous reactions on sea salt aerosols emitted from the sublimation of lofted blowing snow were included. Their results  
685 suggested that blowing snow could be a source of SSA during spring although it should be noted that this version of the model overestimated SSA at remote Arctic sites, such as Alert and Villum, when blowing snow was included as a source of SSA. Also, they did not examine wintertime conditions.

Finally, following *KRP19*, modelled and observed molar ratios of sub-micron Cl<sup>-</sup>:Na<sup>+</sup> and SO<sub>4</sub><sup>2-</sup>:Na<sup>+</sup> are estimated to further examine the origins of SSA and to compare our findings with *KRP19* (see Table 5). Observed molar ratios of Cl<sup>-</sup>:Na<sup>+</sup>  
690 and SO<sub>4</sub><sup>2-</sup>:Na<sup>+</sup> for January and February campaign periods agree with *KRP19* (Cl<sup>-</sup>:Na<sup>+</sup> equal to 1.08, see *KRP19* supplement - Table S3 and text). This indicates a seawater origin (following *Pilson (2012)*), and confirms the findings of *KRP19* that there was no evidence for blowing snow and frost flowers as a source of SSA during the campaigns. Model averaged molar ratios are smaller in magnitude than the observations. Observed and modelled ratios differences in magnitude could be altered by the fact that the model underestimates sub-micron SSA and SO<sub>4</sub><sup>2-</sup>, due to missing mechanisms for sub-micron SSA emissions  
695 and local/regional anthropogenic sources of SO<sub>4</sub><sup>2-</sup>. Differences between observed and modelled Cl<sup>-</sup>:Na<sup>+</sup> ratios could also be due related to issues with modelled SSA lifetime and chemical processing during long range transport. Previous studies found that sub-micron SSA have larger chloride depletion than super-micron SSA (*Barrie et al., 1994; Hara et al., 2002; Leck et al., 2002*). *May et al. (2016)* used molar ratio enrichment factors of Cl<sup>-</sup>:Na<sup>+</sup> as an indicator of long-range transport influence on SSA at Utqiagvik. They reported that Cl<sup>-</sup> depletion was larger for sub- than super-micron SSA due to a longer lifetime. On  
700 average during the simulation periods in January and February 2014, the results indicate that modelled Cl<sup>-</sup> has undergone



significant atmospheric processing. This is consistent with KRP18 observing the presence of both nascent (locally-produced) SSA and aged (partially chloride-depleted) SSA. Based on this analysis of depletion factors and molar ratios, little evidence suggest a blowing snow influence on SSA during the campaigns at Utqiagvik is found. Rather, the presence of predominantly easterly winds (s.s. 5.3) and the presence of leads east of Utqiagvik (especially during February), suggests that the primary source of SSA was marine from open-leads, in agreement with the findings of KRP19.

## 5.7 Conclusions

In this study WRF-Chem is used to investigate Arctic Haze composition at remote Arctic sites during wintertime with a particular focus on SSA, processes influencing SSA emissions and the contribution of SSA to Arctic Haze. Model performance is evaluated first in terms of reproducing aerosol composition in the Arctic before focusing on processes influencing SSA at regional scales over northern Alaska during winter 2014.

The control version of WRF-Chem overestimates super-micron, coarse mode and TSP SSA due to missing and out of date SSA emission treatments in the model. In particular, the addition of a more realistic wind speed dependence for SSA, based on satellite data, and inclusion of a dependence of SSA emissions on SSTs leads to improved results for super-micron, coarse mode and TSP SSA and  $\text{NO}_3^-$  over the Arctic. The latter has already been included in certain modelling studies. Also, recent data analysis studies in the Arctic have pointed out that wind speed alone cannot predict SSA production and that other mechanisms, such as SST dependence, are needed. However, there are still uncertainties regarding the role of SSTs in SSA production. Other factors such as seawater composition, wave characteristics, fetch model and salinity need to be considered in future versions of WRF-Chem. In this study, marine organic aerosol emissions are also activated in the model since they are an important component of SSA in the Arctic and globally. Inclusion of all these updates leads to improved representation of SSA over the wider Arctic. Modelled super-micron, coarse mode and TSP SSA are reduced at all Arctic sites in better agreement with the observations. Results for  $\text{NO}_3^-$  are also improved overall due to less formation via heterogeneous uptake of  $\text{HNO}_3$ . Inclusion of the SST dependence only has a small effect on sub-micron SSA in the Arctic. In the future, other SST dependencies could be considered such as that proposed by *Sofiev et al. (2011)* which could increase sub-micron SSA at low temperatures (*Salter et al., 2014, 2015; Barthel et al., 2019*). However, further field data studies are needed to confirm such dependencies in the Arctic.

A source of  $\text{ss-SO}_4^{2-}$  is also added to the model leading to improved modeled  $\text{SO}_4^{2-}$  in the high Arctic (e.g. Alert) and Alaskan (e.g. Gates of the Arctic, Simeonof) sites. However, at sites such as Utqiagvik, which may be influenced by the Prudhoe Bay oilfields, the model still underestimates sub-micron  $\text{SO}_4^{2-}$  possibly due to missing anthropogenic emissions. Missing aqueous chemical formation of  $\text{SO}_4^{2-}$  in dark conditions may also explain these discrepancies (e.g.  $\text{SO}_4^{2-}$  production from metal catalyzed  $\text{O}_2$  oxidation of S(IV), *McCabe et al. (2006)*). Results from the improved quasi-hemispheric run indicate a shift in the balance between  $(\text{NH}_4)_2\text{SO}_4$  and  $\text{NH}_4\text{NO}_3$ , with aerosols being less acidic than the base version of the model.

Overall, super-micron, coarse mode and TSP SSA, OA,  $\text{SO}_4^{2-}$ ,  $\text{NH}_4^+$  and  $\text{NO}_3^-$  are improved in the HEM\_NEW quasi-hemispheric simulation compared to observations at Arctic sites, based on biases and RMSEs. However, the model underestimates sub-micron SSA at Utqiagvik where there are episodes with significantly higher SSA compared to other Arctic sites.



735 Model sensitivity to different processes affecting SSA over northern Alaska during winter is explored. KRP19 pointed out  
that there is sea ice biology influence at Utqiagvik during wintertime and that marine emissions are an important source of  
organic aerosols at this location. In order to include local sources of marine organics, the ratio of OC:Na<sup>+</sup> is used leading to  
higher modelled OA, in better agreement with previous measurements at this site (and at Alert) and its advised to be included  
in future WRF-Chem simulations in the Arctic region. To further explore the uncertainties on sub-micron SSA, ERA5 sea-ice  
740 fraction is tested in the model. The results, in combination with different sensitivities changing sea-ice fraction, show that  
super-micron SSA are more sensitive to sea-ice treatments than sub-micron SSA in the model. The use of satellite sea-ice data,  
combined with higher resolution simulations over Utqiagvik and coastal Arctic sites, will help to gain more detailed insights  
into the influence of open-leads on SSA production during wintertime. The results of this study also highlight that SSA dry  
removal is less important than the role of open leads in the Arctic during wintertime. The role of wet deposition on SSA is also  
745 examined. In that case, the precipitation flux is doubled and as result super-micron SSA decreased, but the sensitivity did not  
affect sub-micron SSA. Wet deposition is not addressed further in this study, because according to NOAA climate data recorded  
precipitation and snowfall was the lowest during February 2014. Wet deposition is addressed in details in the companion paper  
for BC. Our results suggest that further investigation is needed to determine more realistic dry deposition velocities over snow,  
ice and ocean in the Arctic and to derive more realistic sea-ice fractions, including the presence of open leads, which can vary  
750 over periods of days. The sensitivity of model results to using a higher wind speed dependence, based on data from *Russell  
et al. (2010)*, is investigated for sub-micron SSA. This leads to small improvements in the model sub-micron SSA, with the  
model performing better during January than February period of the campaign.

Further analysis is required to understand the origins of, in particular sub-micron SSA in northern Alaska, and to improve  
their representation in the WRF-Chem model. For example, missing sources of sub-micron SSA, such as a source function for  
755 ultrafine SSA particles due to breaking waves (*Clarke et al., 2006*) could be included. Also, anthropogenic sources of Cl<sup>-</sup> may  
need to be considered, such as road salt in urban areas (*McNamara et al., 2020; Denby et al., 2016*) or coal combustion, waste  
incineration, and industrial activities (*Wang et al., 2019b*) which are not included in current global inventories. The model  
also lacks anthropogenic emissions of Na<sup>+</sup>. Anthropogenic sources of Na<sup>+</sup> could be wastewater and sewage treatment systems,  
contamination from landfills and salt storage areas (e.g. *Panno et al. (2006)*). However, detailed analysis of depletion factors  
760 and molar ratios at Utqiagvik, Alaska showed that during the simulation period the main source of SSA are from marine  
emissions including open ocean or leads and there is no evidence of frost flowers or blowing snow as a source of SSA, at  
least during the periods considered in this study, in agreement with the findings of KRP19. Further observations from field  
measurements are needed to better understand SSA emissions and their dependencies.

This model study supports recent findings based on observations that SSA make an important contribution to super-micron  
765 (coarse mode, TSP) mass concentrations during wintertime at remote Arctic sites. Future work has to consider carefully possi-  
ble sources of sub-micron SSA and their inclusion in models, in order to explain observed SSA during wintertime. Processes  
linked to the open ocean are likely to become more important with decreasing sea-ice cover in the Arctic due to climate warm-  
ing. Observations of SSA components including organic aerosols (often missing) are needed to improve understanding about  
processes and their treatments in models, and in order to reduce uncertainties in estimation of aerosol radiative effects.

<https://doi.org/10.5194/egusphere-2022-310>

Preprint. Discussion started: 11 May 2022

© Author(s) 2022. CC BY 4.0 License.



770 *Code availability.* Available upon request.

*Data availability.* All data used in the present paper for Zeppelin and Alert are open access and are available at the EBAS database infrastructure at NILU - Norwegian Institute for Air research: <http://ebas.nilu.no/>. Observations for Villum are obtained after personal communication with Henrik Skov. Observations from IMPROVE database can be obtained from: <http://views.cira.colostate.edu/fed/QueryWizard/>. Sub- and super-micron aerosol mass concentrations at Utqiagvik, Alaska can be obtained from the follow link: <https://saga.pmel.noaa.gov/data/stations/>.

775



**Table A1.** Land Surface model's (NOAH MP) parametrization. "Opt\_" indicates the namelist option for NOAH MP.

NOAH MP Parametrization	
Dynamic Vegetation (DVEG)	On
Stomatal Resistance	Ball-Berry <i>Ball et al. (1987)</i> , <i>Collatz et al. (1991)</i> , <i>Collatz et al. (1992)</i> , <i>Bonan (1996)</i> , <i>Sellers et al. (1996)</i>
Surface layer drag coefficient (opt_sfc)	Original Noah <i>Chen et al. (1997)</i>
Soil moisture for stomatal resistance (opt_btr)	Noah (soil moisture)
Runoff (opt_run)	TOPMODEL with groundwater <i>Niu et al. (2007)</i>
Supercooled liquid water (opt_frz)	no iteration <i>Niu and Yang (2004)</i>
Soil permeability (opt_inf)	linear effects, more permeable <i>yue Niu and liang Yang (2006)</i>
Radiative transfer (opt_rad)	modified two-stream (gap = F(solar angle, 3D structure ...) < 1-FVEG) <i>Yang and Friedl (2003)</i> , <i>Niu and Yang (2004)</i>
Ground surface albedo (opt_alb)	BATS <i>Yang Z.-L. and Vinnikov. (1997)</i>
Precipitation (snow/rain) partitioning (opt_snf)	<i>Jordan (1991)</i>
Soil temperature lower boundary (opt_tbot)	TBOT at ZBOT (8m) read from a file (original Noah)
Soil/snow temperature time scheme (opt_sfc)	semi-implicit; flux top boundary condition
Surface resistance to evaporation/sublimation (opt_rsf)	<i>Sakaguchi and Zeng (2009)</i>
Glacier treatment (opt_gla)	include phase change of ice

## Appendix A

Following *Monaghan et al. (2018)*, NOAH-MP parameter file MPTABLE.TBL has been modified, and it can be used for simulations over Alaska. These modifications improved the model's capability to capture cold surface temperature and meteorological profiles (e.g. wind speed, relative humidity, temperature) over Alaska.

## 780 Appendix B

Fuentes size-resolved sea-spray source flux



$$\frac{dF_o}{d\log D_{p0}} = \frac{dF_p}{d\log D_{p0}} \times W = \frac{Q}{A_b} \times \frac{dN_T}{d\log D_{p0}} \times W \quad (\text{B1})$$

where  $W(U)$  is Monahan and O’Muircheartaigh whitecap coverage,  $dF_p/d\log D_{p0}$  is the size-resolved particle flux per unit time and water surface covered by bubbles,  $D_{p0}$  is the dry diameters,  $Q$  is the sweep air flow,  $A_b$  is the total surface area covered by bubbles,  $dN_T/d\log D_{p0}$  is the particle size distribution (the sum of four log-normal modes) and is equal to:

$$\frac{dN_T}{d\log D_{p0}} = \sum_{i=1}^4 \frac{dN_{T,i}}{d\log D_{p0}} = \sum_{i=1}^4 \frac{N_{T,i}}{\sqrt{2\pi} \times \log \sigma_i} \times \exp\left[-\frac{1}{2} \times \left(\frac{\log \frac{D_{p0}}{D_{p0g,i}}}{\log \sigma_i}\right)^2\right] \quad (\text{B2})$$

where  $i$  is the sub-index for the mode number and  $N_i$ ,  $D_{p0g,i}$  and  $\sigma_i$  are the total particle number, geometric mean and geometric standard deviation for each log-normal mode.  $N_{T,i}$  and  $D_{p0g,i}$  are depending on parameters  $a_i$  and  $\beta_i$  derived from polynomial and exponential regressions defining the total particle number and geometric mean diameter of the log-normal modes, and can be found in Table 5 *Fuentes et al. (2010)*.

## Appendix C

In this APPENDIX, the biases and RMSEs are calculated for each site, as shown in Fig. 1, and are shown in the tables below. Each table corresponds to a site and for the available observed aerosol concentrations, such as  $\text{Na}^+$ ,  $\text{Cl}^-$ ,  $\text{SO}_4^{2-}$  (total and non-sea component),  $\text{NO}_3^-$ ,  $\text{NH}_4^+$  and OC. Bias is calculated as the difference between model simulation and observation.



**Table C1.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for aerosols at the Alert, Canada, during January and February 2014 and for CONTROL and HEM\_NEW simulations at 100km.

	CONTROL		HEM_NEW	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	0.81	0.91	0.12	0.18
<b>Cl<sup>-</sup></b>	1.05	1.2	-0.03	0.19
<b>NO<sub>3</sub><sup>-</sup></b>	0.28	0.30	0.25	0.22
<b>NH<sub>4</sub><sup>+</sup></b>	-0.003	0.01072	0.007	0.01079
<b>SO<sub>4</sub><sup>2-</sup></b>	0.06	0.1	0.02	0.11

**Table C2.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for aerosols at Villum Research station, Greenland, during January and February 2014 and for CONTROL and HEM\_NEW simulations at 100km.

	CONTROL		HEM_NEW	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	1.3	1.4	0.25	0.26
<b>Cl<sup>-</sup></b>	1.9	2.1	0.22	0.24
<b>NO<sub>3</sub><sup>-</sup></b>	0.25	0.26	0.17	0.19
<b>NH<sub>4</sub><sup>+</sup></b>	-0.001	0.006	0.01	0.01
<b>SO<sub>4</sub><sup>2-</sup></b>	0.05	0.1	0.08	0.1

795

**Table C3.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for aerosols at Zeppelin, Norway, during January and February 2014 and for CONTROL and HEM\_NEW simulations at 100km.

	CONTROL		HEM_NEW	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	3.31	4.4	0.36	0.78
<b>Cl<sup>-</sup></b>	4.86	6.48	0.22	0.73
<b>NO<sub>3</sub><sup>-</sup></b>	0.13	0.36	0.01	0.29
<b>NH<sub>4</sub><sup>+</sup></b>	-0.03	0.077	-0.02	0.076
<b>SO<sub>4</sub><sup>2-</sup></b>	0.16	0.25	0.32	0.45



**Table C4.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for aerosols at Simeonof, south of Alaska, during January and February 2014 and for CONTROL and HEM\_NEW simulations at 100km.

	CONTROL		HEM_NEW	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	1.4	2.5	0.3	0.6
<b>Cl<sup>-</sup></b>	2.0	3.7	0.1	0.7
<b>NO<sub>3</sub><sup>-</sup></b>	0.12	0.23	0.08	0.20
<b>SO<sub>4</sub><sup>2-</sup></b>	-0.2	0.25	-0.05	0.26
<b>OA</b>	-0.08	0.1	-0.05	0.08

**Table C5.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for aerosols at Gates of the Arctic, south of Alaska, during January and February 2014 and for CONTROL and HEM\_NEW simulations at 100km.

	CONTROL		HEM_NEW	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	0.6	0.9	0.2	0.3
<b>Cl<sup>-</sup></b>	0.7	1.2	0.1	0.3
<b>NO<sub>3</sub><sup>-</sup></b>	0.3	0.3	0.2	0.2
<b>SO<sub>4</sub><sup>2-</sup></b>	-0.04	0.2	-0.1	0.2
<b>OA</b>	-0.24	0.28	-0.21	0.26

**Table C6.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for super-micron aerosols at Utqiagvik, north of Alaska, during January and February 2014 and CONTROL and HEM\_NEW simulations at 100km.

	CONTROL		HEM_NEW	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	0.3	0.37	-0.07	0.25
<b>Cl<sup>-</sup></b>	0.27	0.48	-0.26	0.51
<b>NO<sub>3</sub><sup>-</sup></b>	0.26	0.3	0.13	0.17
<b>NH<sub>4</sub><sup>+</sup></b>	-0.0004	0.00368	-0.001	0.0037
<b>SO<sub>4</sub><sup>2-</sup></b>	0.005	0.005	-0.01	0.06



**Table C7.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for sub-micron aerosols at Utqiagvik, north of Alaska, during January and February 2014 and CONTROL and HEM\_NEW simulations at 100km.

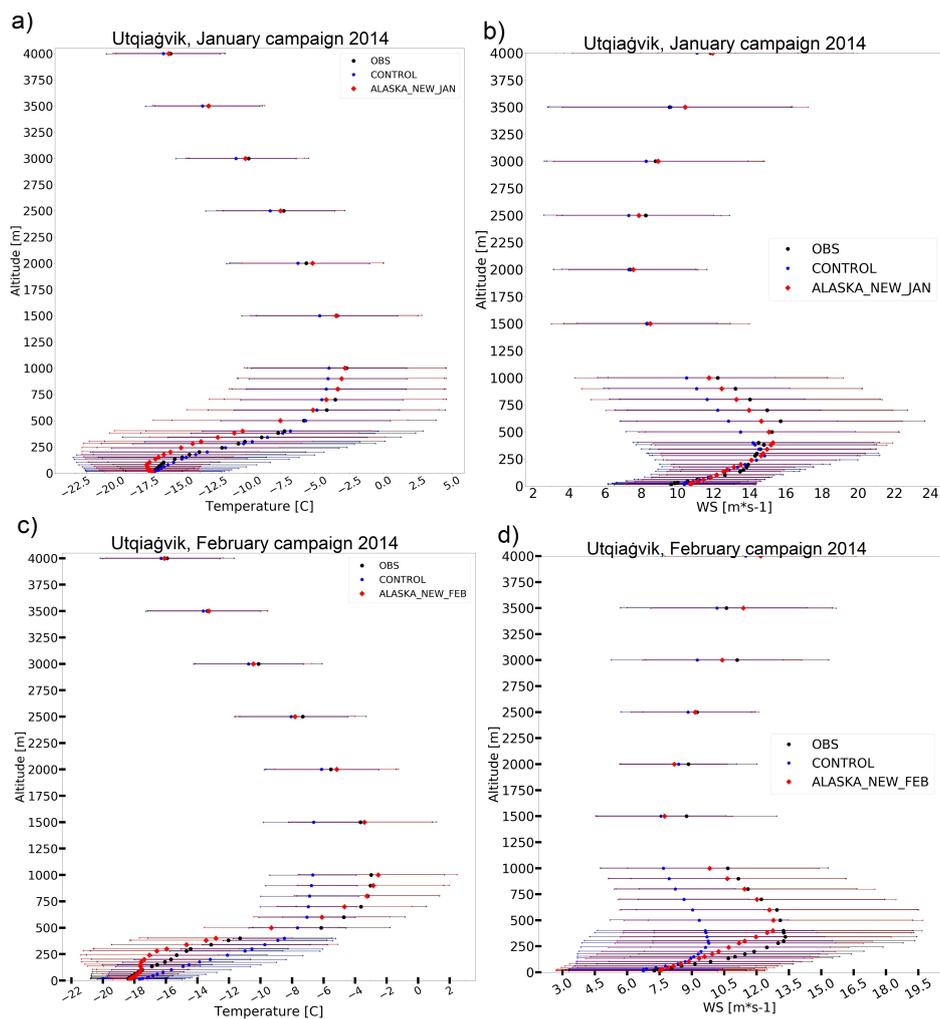
	CONTROL		HEM_NEW	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	-0.485	0.66	-0.489	0.67
<b>Cl<sup>-</sup></b>	-0.116	0.361	-0.124	0.364
<b>NO<sub>3</sub><sup>-</sup></b>	-0.065	0.162	-0.054	0.158
<b>NH<sub>4</sub><sup>+</sup></b>	-0.069	0.106	-0.057	0.100
<b>SO<sub>4</sub><sup>2-</sup></b>	-0.621	0.875	-0.591	0.853

**Table D1.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated between ALASKA\_NEW\_JAN, ALASKA\_NEW\_FEB and in-situ meteorological parameters derived from NOAA Baseline Observatories during the campaign's periods in January and February 2014. Bias was calculated as the difference between model simulation and observations.

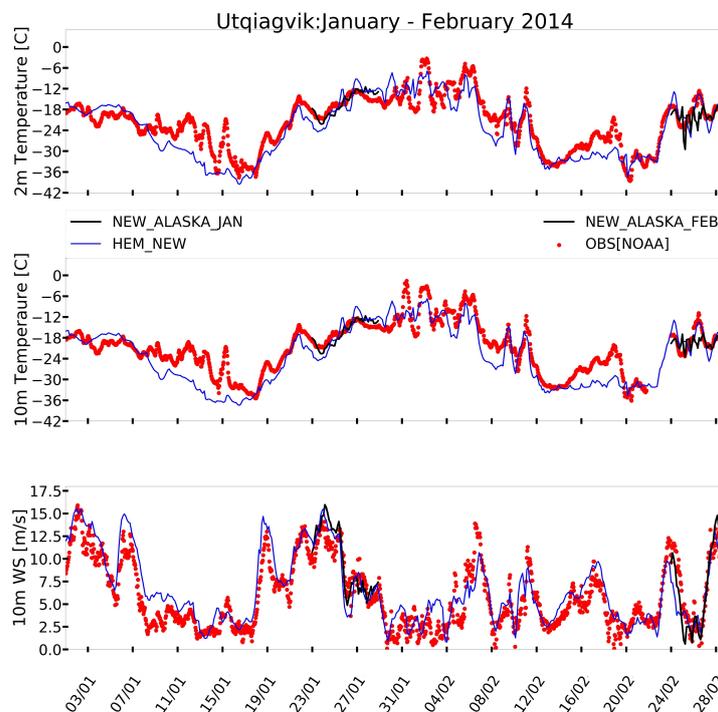
	January campaign		February campaign	
	Bias	RMSE	Bias	RMSE
2m Temperature	0.1	1.9	-1.0	3.2
10m Temperature	-0.03	1.8	-0.66	2.7
10m Wind speed	0.08	1.4	-0.33	1.7
10m Wind direction	-11.2	13.2	-11.2	39.0

## Appendix D

Surface observations are used to validate the meteorological conditions that occur over Utqiagvik and Alaska in wintertime. See also discussion in sub-section 5.1 in the main text. The model is validated against the surface (hourly) observations obtained from National Oceanic and Atmospheric Administration / Earth System Research Laboratory / Global Monitoring Division (NOAA/ESRL/GMD) Baseline Observatories. Also, radiosondes data are used to evaluate the model's performance at different altitudes. Radiosonde data (every 12h) are derived from Integrated Global Radiosonde Archive version 2 (IGRA 2) (Durre *et al.* (2018)). Site is located at latitude: 71.2889 and longitude: -156.7833.



**Figure D1.** Average temperatures, in degrees C, and wind speeds, in  $\text{ms}^{-1}$ , as a function of altitude (m), up to 4km, during (a,b) January and (c,d) February campaign in 2014, at Utqiagvik, Alaska. The observations are shown in black (cicle). The blue pentagon shows the model results for the CONTROL simulation (at 100km) and the red diamond shows the model results for the NEW\_ALASKA\_JAN and NEW\_ALASKA\_FEB simulation. Observations are derived from IGRA2 and are available every 12h (0Z and 12Z, UTC). For the comparison, model output at 0 and 12Z UTC are used. The corresponding horizontal lines show the standard deviation.



**Figure D2.** Time series of observed and modelled 2m and 10m temperature, and 10m wind speed, at Utqiagvik, Alaska, in UTC. The observations are shown in red and derived from the NOAA observatory. The blue line shows the results for the HEM\_NEW simulation at 100km, while the black line shows the results for ALASKA\_NEW\_JAN and ALASKA\_NEW\_FEB simulations at 20km. The observations are hourly, while the model output is every 3h.

## Appendix E

Here the bias and RMSE are shown between ALASKA\_NEW\_JAN and ALASKA\_NEW\_FEB and the observations for  
805 Utqiagvik at 20km.

## Appendix F

This APPENDIX shows the comparison for the Gates of the Arctic site at 20 km, for ALASKA\_NEW\_FEB and  
ALASKA\_CONTROL\_FEB. The observations are only available for the February campaign, daily averaged in local Alaskan  
time every three days.

810 Also, the table below shows biases and RMSEs, in  $\mu\text{g m}^{-3}$ , for all available aerosol species at the Gates of the Arctic.



**Table E1.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for aerosols at Utqiagvik, north of Alaska, during January 2014 and for ALASKA\_CONTROL\_JAN and ALASKA\_NEW\_JAN simulations at 20km.

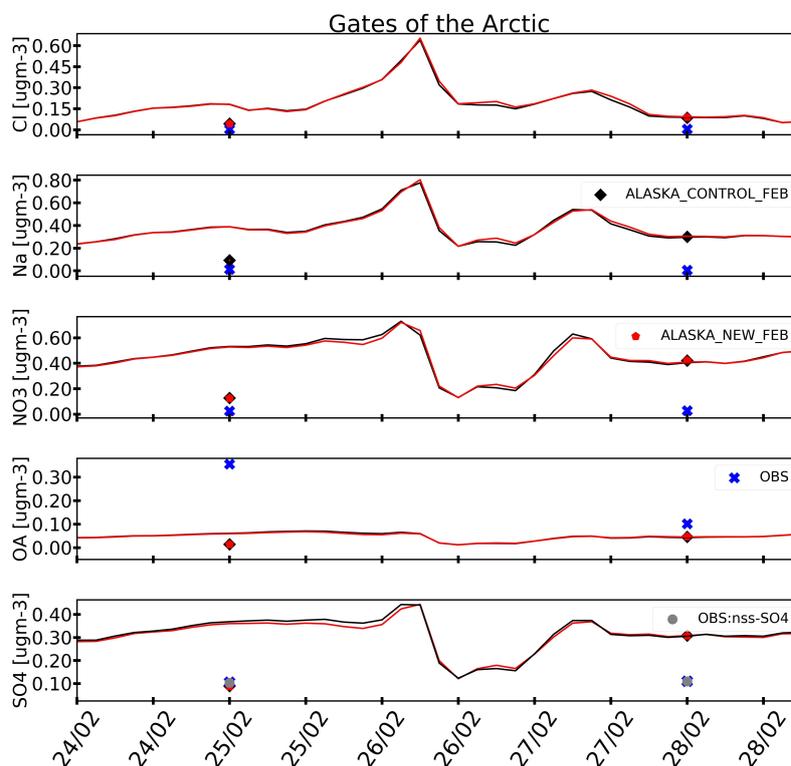
	ALASKA_CONTROL_JAN		ALASKA_NEW_JAN	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	-0.31	0.38	-0.16	0.26
<b>Cl<sup>-</sup></b>	-0.50	0.59	-0.33	0.43
<b>NO<sub>3</sub><sup>-</sup></b>	-0.040	0.07	0.039	0.09
<b>SO<sub>4</sub><sup>2-</sup></b>	-0.396	0.414	-0.398	0.417
<b>NH<sub>4</sub><sup>+</sup></b>	-0.033	0.038	-0.035	0.040

**Table E2.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , calculated for aerosols at Utqiagvik, north of Alaska, during February 2014 and for ALASKA\_CONTROL\_FEB and ALASKA\_NEW\_FEB simulations at 20km.

	ALASKA_CONTROL_FEB		ALASKA_NEW_FEB	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	-1.29	1.40	-1.18	1.30
<b>Cl<sup>-</sup></b>	-1.90	1.92	-1.78	1.80
<b>NO<sub>3</sub><sup>-</sup></b>	-0.20	0.40	-0.11	0.38
<b>SO<sub>4</sub><sup>2-</sup></b>	-1.019	1.322	-1.020	1.326
<b>NH<sub>4</sub><sup>+</sup></b>	-0.045	0.097	-0.043	0.10

**Table F1.** Biases and RMSEs, in  $\mu\text{g m}^{-3}$ , are calculated for aerosols at Gates of the Arctic, north of Alaska, during February campaign and for ALASKA\_CONTROL\_FEB and ALASKA\_NEW\_FEB simulations at 20km.

	ALASKA_CONTROL_FEB		ALASKA_NEW_FEB	
	Bias	RMSE	Bias	RMSE
<b>Na<sup>+</sup></b>	0.344	0.346	0.342	0.341
<b>Cl<sup>-</sup></b>	0.155	0.1578	0.154	0.1573
<b>NO<sub>3</sub><sup>-</sup></b>	0.47	0.48	0.42	0.47
<b>SO<sub>4</sub><sup>2-</sup></b>	0.24	0.24	0.23	0.23
<b>OA</b>	-0.198	0.2446	-0.197	0.2445



**Figure F1.** Model inter-variability during February campaign. Model simulations are validated against aerosols at the gates of the Arctic site, north of Alaska. The black line shows ALASKA\_CONTROL\_FEB simulation and the black symbol the daily averaged values. The red line shows ALASKA\_NEW\_FEB simulation and the red pentagon the daily averaged values. The blue star indicates averaged daily observations. Observations and model are in local Alaskan time. Observed and modelled  $\text{SO}_4^{2-}$  is total  $\text{SO}_4^{2-}$ .

*Author contributions.* The first author (EI) implemented the updates, performed the simulations and the analysis, and drafted the paper. KSL designed the study and contributed to the interpretation of results and the analysis, and to the writing of the paper. JCR, LM and TO contributed to discussions about the model setup and simulations. JCR, LM, KP, RMK, PKQ and LU contributed to the analysis and interpretation of the results. AM and HS contributed to the interpretation of the results. All co-authors contributed to the paper and the discussions about the results.

*Competing interests.* The authors declare that they have no conflict of interest.

*Acknowledgements.* This study is supported by the PhD programme of the Make our planet great again (MOGPA) French initiative. We also acknowledge support from French ANR CASPA (Climate-relevant Aerosols and Sources in the Arctic). Computer modeling performed by Eleftherios Ioannidis, benefited from access to IDRIS HPC resources (GENCI allocations A009017141 and A011017141) and



820 the IPSL mesoscale computing center (CICLAD: Calcul Intensif pour le Climat, l' Atmosphère et la Dynamique). We acknowledge the use  
of MOZART-4 global model output available at  
825 <http://www.acom.ucar.edu/wrf-chem/mozart.shtml>. We acknowledge use of the WRF-Chem preprocessor tool mozbc provided by the Atmo-  
spheric Chemistry Observations and Modeling Lab (ACOM) of NCAR. We would like to thank Arctic Monitoring & Assessment Programme  
(AMAP) and Zbigniew Klimont for ECLIPSE v6b anthropogenic emissions. K.A.P. acknowledges funding from the US National Science  
830 Foundation (OPP-1724585). This is PMEL contribution number 5366. Also, we would like to thank the IMPROVE database for the aerosol  
observations in Alaska. IMPROVE is a collaborative association of state, tribal, and federal agencies, and international partners. US Envi-  
ronmental Protection Agency is the primary funding source, with contracting and research support from the National Park Service. The Air  
Quality Group at the University of California, Davis is the central analytical laboratory, with ion analysis provided by Research Triangle  
Institute, and carbon analysis provided by Desert Research Institute. Meteorological data provided by NOAA Global Monitoring Laboratory,  
835 Boulder, Colorado, USA (<https://esrl.noaa.gov/>). We would like to thank Josh Jones for providing us with detailed maps of sea-ice fraction  
plots from the radar at Utqiagvik. The Villum Foundation is gratefully acknowledged for financing the establishment of the Villum Research  
Station. Finally, we would like to thank Aas Wenche (Norwegian Institute for Air Research) and Sharma Sangeeta (Environment and Cli-  
mate Change Canada, Science and Technology Branch, Toronto, Canada) for providing us with EBAS observations at Zeppelin and Alert,  
respectively. Authors also want to thank Canadian Forces Services, Alert, NU for maintaining the site. The observations at Zeppelin have  
been supported by the Norwegian Environment Agency (grant no. 17078061).



## References

- Aas, W., Eckhardt, S., Fiebig, M., Platt, S. M., Solberg, S., Yttri, K. E., and Zwaafink, C. G.: Monitoring of long-range transported air pollutants in Norway. Annual Report 2020., NILU rapport, 2021.
- Alexander, B., Park, R. J., Jacob, D. J., Li, Q., Yantosca, R. M., Savarino, J., Lee, C., and Thiemens, M.: Sulfate formation in sea-salt aerosols: Constraints from oxygen isotopes, *Journal of Geophysical Research: Atmospheres*, 110, 2005.
- Allan, R. P.: Climate Change 2021: The Physical Science Basis: Working Group I Contribution to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change, WMO, IPCC Secretariat, 2021.
- AMAP: AMAP Assessment 2015: Black carbon and ozone as Arctic climate forcers, 2015.
- Archer-Nicholls, S., Lowe, D., Utembe, S., Allan, J., Zaveri, R. A., Fast, J. D., Hodnebrog, Ø., Denier van der Gon, H., and McFiggans, G.: Gaseous chemistry and aerosol mechanism developments for version 3.5.1 of the online regional model, WRF-Chem, *Geoscientific Model Development*, 7, 2557–2579, <https://doi.org/10.5194/gmd-7-2557-2014>, 2014.
- Ball, J. T., Woodrow, I. E., and Berry, J. A.: A model predicting stomatal conductance and its contribution to the control of photosynthesis under different environmental conditions, in: *Progress in photosynthesis research*, pp. 221–224, Springer, 1987.
- Barrett, T. and Sheesley, R.: Year-round optical properties and source characterization of Arctic organic carbon aerosols on the North Slope Alaska, *Journal of Geophysical Research: Atmospheres*, 122, 9319–9331, 2017.
- Barrett, T., Robinson, E., Usenko, S., and Sheesley, R.: Source contributions to wintertime elemental and organic carbon in the western Arctic based on radiocarbon and tracer apportionment, *Environmental science & technology*, 49, 11 631–11 639, 2015.
- Barrie, L. and Barrie, M.: Chemical components of lower tropospheric aerosols in the high Arctic: Six years of observations, *Journal of Atmospheric Chemistry*, 11, 211–226, 1990.
- Barrie, L. and Hoff, R.: The oxidation rate and residence time of sulphur dioxide in the Arctic atmosphere, *Atmospheric Environment (1967)*, 18, 2711–2722, 1984.
- Barrie, L., Bottenheim, J., Schnell, R., Crutzen, P., and Rasmussen, R.: Ozone destruction and photochemical reactions at polar sunrise in the lower Arctic atmosphere, *Nature*, 334, 138–141, 1988.
- Barrie, L., Staebler, R., Toom, D., Georgi, B., Den Hartog, G., Landsberger, S., and Wu, D.: Arctic aerosol size-segregated chemical observations in relation to ozone depletion during Polar Sunrise Experiment 1992., *Journal of Geophysical Research: Atmospheres*, 99, 25 439–25 451, 1994.
- Barrie, L. A.: Arctic air pollution: An overview of current knowledge, *Atmospheric Environment (1967)*, 20, 643–663, 1986.
- Barthel, S., Tegen, I., and Wolke, R.: Do new sea spray aerosol source functions improve the results of a regional aerosol model?, *Atmospheric Environment*, 198, 265–278, 2019.
- Berg, L. K., Gustafson, W. I., Kassianov, E. I., and Deng, L.: Evaluation of a modified scheme for shallow convection: Implementation of CuP and case studies, *Monthly weather review*, 141, 134–147, 2013.
- Binkowski, F. S. and Shankar, U.: The regional particulate matter model: 1. Model description and preliminary results, *Journal of Geophysical Research: Atmospheres*, 100, 26 191–26 209, 1995.
- Bonan, G. B.: Land surface model (LSM version 1.0) for ecological, hydrological, and atmospheric studies: Technical description and users guide. Technical note, Tech. rep., National Center for Atmospheric Research, Boulder, CO (United States . . . , 1996.
- Bond, T. C., Doherty, S. J., Fahey, D. W., Forster, P. M., Berntsen, T., DeAngelo, B. J., Flanner, M. G., Ghan, S., Kärcher, B., Koch, D., Kinne, S., Kondo, Y., Quinn, P. K., Sarofim, M. C., Schultz, M. G., Schulz, M., Venkataraman, C., Zhang, H., Zhang, S., Bellouin, N.,



- Guttikunda, S. K., Hopke, P. K., Jacobson, M. Z., Kaiser, J. W., Klimont, Z., Lohmann, U., Schwarz, J. P., Shindell, D., Storelvmo, T., Warren, S. G., and Zender, C. S.: Bounding the role of black carbon in the climate system: A scientific assessment, *Journal of Geophysical Research: Atmospheres*, 118, 5380–5552, <https://doi.org/10.1002/jgrd.50171>, 2013.
- 875 Bowen, H. J. M. et al.: Environmental chemistry of the elements., Academic Press., 1979.
- Calhoun, J. A., Bates, T. S., and Charlson, R. J.: Sulfur isotope measurements of submicrometer sulfate aerosol particles over the Pacific Ocean, *Geophysical Research Letters*, 18, 1877–1880, 1991.
- Callaghan, A., de Leeuw, G., Cohen, L., and O’Dowd, C. D.: Relationship of oceanic whitecap coverage to wind speed and wind history, *Geophysical Research Letters*, 35, 2008.
- 880 Callaghan, A. H.: An improved whitecap timescale for sea spray aerosol production flux modeling using the discrete whitecap method, *Journal of Geophysical Research: Atmospheres*, 118, 9997–10, 2013.
- Callaghan, A. H., Deane, G. B., Stokes, M. D., and Ward, B.: Observed variation in the decay time of oceanic whitecap foam, *Journal of Geophysical Research: Oceans*, 117, 2012.
- 885 Callaghan, A. H., Stokes, M., and Deane, G.: The effect of water temperature on air entrainment, bubble plumes, and surface foam in a laboratory breaking-wave analog, *Journal of Geophysical Research: Oceans*, 119, 7463–7482, 2014.
- Carter, W. P.: Documentation of the SAPRC-99 chemical mechanism for VOC reactivity assessment, Contract, 92, 95–308, 2000.
- Chen, F., Janjić, Z., and Mitchell, K.: Impact of atmospheric surface-layer parameterizations in the new land-surface scheme of the NCEP mesoscale Eta model, *Boundary-Layer Meteorology*, 85, 391–421, 1997.
- 890 Chen, Y., Cheng, Y., Ma, N., Wolke, R., Nordmann, S., Schüttauf, S., Ran, L., Wehner, B., Birmili, W., Gon, H. A., et al.: Sea salt emission, transport and influence on size-segregated nitrate simulation: a case study in northwestern Europe by WRF-Chem, *Atmospheric Chemistry and Physics*, 16, 12 081–12 097, 2016.
- Clarke, A. D., Owens, S. R., and Zhou, J.: An ultrafine sea-salt flux from breaking waves: Implications for cloud condensation nuclei in the remote marine atmosphere, *Journal of Geophysical Research: Atmospheres*, 111, 2006.
- 895 Collatz, G. J., Ball, J. T., Grivet, C., and Berry, J. A.: Physiological and environmental regulation of stomatal conductance, photosynthesis and transpiration: a model that includes a laminar boundary layer, *Agricultural and Forest meteorology*, 54, 107–136, 1991.
- Collatz, G. J., Ribas-Carbo, M., and Berry, J.: Coupled photosynthesis-stomatal conductance model for leaves of C4 plants, *Functional Plant Biology*, 19, 519–538, 1992.
- De Leeuw, G., Neele, F. P., Hill, M., Smith, M. H., and Vignati, E.: Production of sea spray aerosol in the surf zone, *Journal of Geophysical Research: Atmospheres*, 105, 29 397–29 409, 2000.
- 900 Denby, B. R., Ketzler, M., Ellermann, T., Stojiljkovic, A., Kupiainen, K., Niemi, J., Norman, M., Johansson, C., Gustafsson, M., Blomqvist, G., et al.: Road salt emissions: A comparison of measurements and modelling using the NORTRIP road dust emission model, *Atmospheric environment*, 141, 508–522, 2016.
- Douglas, T. A., Domine, F., Barret, M., Anastasio, C., Beine, H. J., Bottenheim, J., Grannas, A., Houdier, S., Netcheva, S., Rowland, G., et al.: Frost flowers growing in the Arctic ocean-atmosphere–sea ice–snow interface: 1. Chemical composition, *Journal of Geophysical Research: Atmospheres*, 117, 2012.
- 905 Druckenmiller, M. L., Eicken, H., Johnson, M. A., Pringle, D. J., and Williams, C. C.: Toward an integrated coastal sea-ice observatory: System components and a case study at Barrow, Alaska, *Cold Regions Science and Technology*, 56, 61–72, 2009.
- Durre, I., Yin, X., Vose, R. S., Applequist, S., and Arnfield, J.: Enhancing the data coverage in the Integrated Global Radiosonde Archive, *Journal of Atmospheric and Oceanic Technology*, 35, 1753–1770, 2018.
- 910



- Eckhardt, S., Quennehen, B., Olivie, D. J. L., Berntsen, T. K., Cherian, R., Christensen, J., Collins, W., Crepinsek, S., Daskalakis, N., Flanner, M., et al.: Current model capabilities for simulating black carbon and sulfate concentrations in the Arctic atmosphere: a multi-model evaluation using a comprehensive measurement data set, *Atmospheric Chemistry and Physics*, 15, 9413–9433, 2015.
- 915 Eicken, H., Jones, J., Meyer, F., Mahoney, A., Druckenmiller, M. L., Rohith, M., and Kambhamettu, C.: Environmental security in Arctic ice-covered seas: From strategy to tactics of hazard identification and emergency response, *Marine Technology Society Journal*, 45, 37–48, 2011.
- Emmons, L. K., Walters, S., Hess, P. G., Lamarque, J.-F., Pfister, G. G., Fillmore, D., Granier, C., Guenther, A., Kinnison, D., Laepple, T., et al.: Description and evaluation of the Model for Ozone and Related chemical Tracers, version 4 (MOZART-4), *Geoscientific Model Development*, 3, 43–67, 2010.
- 920 Fan, S.-M. and Jacob, D. J.: Surface ozone depletion in Arctic spring sustained by bromine reactions on aerosols, *Nature*, 359, 522–524, 1992.
- Fast, J. D., Gustafson Jr, W. I., Easter, R. C., Zaveri, R. A., Barnard, J. C., Chapman, E. G., Grell, G. A., and Peckham, S. E.: Evolution of ozone, particulates, and aerosol direct radiative forcing in the vicinity of Houston using a fully coupled meteorology-chemistry-aerosol model, *Journal of Geophysical Research: Atmospheres*, 111, 2006.
- 925 Fisher, J. A., Jacob, D. J., Wang, Q., Bahreini, R., Carouge, C. C., Cubison, M. J., Dibb, J. E., Diehl, T., Jimenez, J. L., Leibenberger, E. M., et al.: Sources, distribution, and acidity of sulfate–ammonium aerosol in the Arctic in winter–spring, *Atmospheric Environment*, 45, 7301–7318, 2011.
- Frey, M. M., Norris, S. J., Brooks, I. M., Anderson, P. S., Nishimura, K., Yang, X., Jones, A. E., Nerentorp Mastromonaco, M. G., Jones, D. H., and Wolff, E. W.: First direct observation of sea salt aerosol production from blowing snow above sea ice, *Atmospheric Chemistry and Physics*, 20, 2549–2578, 2020.
- 930 Frossard, A. A., Shaw, P. M., Russell, L. M., Kroll, J. H., Canagaratna, M. R., Worsnop, D. R., Quinn, P. K., and Bates, T. S.: Springtime Arctic haze contributions of submicron organic particles from European and Asian combustion sources, *Journal of Geophysical Research: Atmospheres*, 116, 2011.
- Frossard, A. A., Russell, L. M., Burrows, S. M., Elliott, S. M., Bates, T. S., and Quinn, P. K.: Sources and composition 935 of submicron organic mass in marine aerosol particles, *Journal of Geophysical Research: Atmospheres*, 119, 12,977–13,003, <https://doi.org/https://doi.org/10.1002/2014JD021913>, 2014.
- Fuentes, E., Coe, H., Green, D., Leeuw, G. d., and McFiggans, G.: On the impacts of phytoplankton-derived organic matter on the properties of the primary marine aerosol–Part 1: Source fluxes, *Atmospheric Chemistry and Physics*, 10, 9295–9317, 2010.
- Fuentes, E., Coe, H., Green, D., and McFiggans, G.: On the impacts of phytoplankton-derived organic matter on the properties of the primary 940 marine aerosol–Part 2: Composition, hygroscopicity and cloud condensation activity, *Atmospheric Chemistry and Physics*, 11, 2585–2602, 2011.
- Fujiki, T., Matsumoto, K., Honda, M. C., Kawakami, H., and Watanabe, S.: Phytoplankton composition in the subarctic North Pacific during autumn 2005, *Journal of plankton research*, 31, 179–191, 2009.
- Giorgi, F.: Dry deposition velocities of atmospheric aerosols as inferred by applying a particle dry deposition parameterization to a general 945 circulation model, *Tellus B*, 40, 23–41, 1988.
- Goddijn-Murphy, L., Woolf, D., and Callaghan, A. H.: Parameterizations and Algorithms for Oceanic Whitecap Coverage, *Journal of Physical Oceanography*, 41, 742–756, 2011.



- Gong, S., Barrie, L., and Blanchet, J.-P.: Modeling sea-salt aerosols in the atmosphere: 1. Model development, *Journal of Geophysical Research: Atmospheres*, 102, 3805–3818, 1997.
- 950 Grell, G. A., Dudhia, J., Stauffer, D. R., et al.: A description of the fifth-generation Penn State/NCAR Mesoscale Model (MM5), 1994.
- Grell, G. A., Peckham, S. E., Schmitz, R., McKeen, S. A., Frost, G., Skamarock, W. C., and Eder, B.: Fully coupled “online” chemistry within the WRF model, *Atmospheric Environment*, 39, 6957–6975, 2005.
- Grythe, H., Ström, J., Krejci, R., Quinn, P., and Stohl, A.: A review of sea-spray aerosol source functions using a large global set of sea salt aerosol concentration measurements, *Atmospheric Chemistry and Physics*, 14, 1277–1297, 2014.
- 955 Guenther, A., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L., and Wang, X.: The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2. 1): an extended and updated framework for modeling biogenic emissions, *Geoscientific Model Development*, 5, 1471–1492, 2012.
- Gunsch, M. J., Liu, J., Moffett, C. E., Sheesley, R. J., Wang, N., Zhang, Q., Watson, T. B., and Pratt, K. A.: Diesel Soot and Amine-Containing Organic Sulfate Aerosols in an Arctic Oil Field, *Environmental Science & Technology*, 54, 92–101, 2020.
- 960 Hancke, K., Lund-Hansen, L. C., Lamare, M. L., Højlund Pedersen, S., King, M. D., Andersen, P., and Sorrell, B. K.: Extreme low light requirement for algae growth underneath sea ice: A case study from Station Nord, NE Greenland, *Journal of Geophysical Research: Oceans*, 123, 985–1000, 2018.
- Hara, K., Osada, K., Matsunaga, K., Iwasaka, Y., Shibata, T., and Furuya, K.: Atmospheric inorganic chlorine and bromine species in Arctic boundary layer of the winter/spring, *Journal of Geophysical Research: Atmospheres*, 107, AAC–4, 2002.
- 965 Hartery, S., Toohey, D., Revell, L., Sellegri, K., Kuma, P., Harvey, M., and McDonald, A. J.: Constraining the surface flux of sea spray particles from the Southern Ocean, *Journal of Geophysical Research: Atmospheres*, 125, e2019JD032026, 2020.
- Heidam, N., Christensen, J., Skov, H., and Wählin, P.: Monitoring and modelling of the atmospheric environment in Greenland. A review, *Sci Total Environ*, 331, 5–28, 2004.
- Hirdman, D., Sodemann, H., Eckhardt, S., Burkhardt, J. F., Jefferson, A., Mefford, T., Quinn, P. K., Sharma, S., Ström, J., and Stohl, A.: 970 Source identification of short-lived air pollutants in the Arctic using statistical analysis of measurement data and particle dispersion model output, *Atmospheric Chemistry and Physics*, 10, 669–693, 2010.
- Hodnebrog, Ø., Marelle, L., Alterskjær, K., Wood, R. R., Ludwig, R., Fischer, E. M., Richardson, T., Forster, P., Sillmann, J., and Myhre, G.: Intensification of summer precipitation with shorter time-scales in Europe, *Environmental Research Letters*, 14, 124050, 2019.
- Hoffman, E. J. and Duce, R. A.: Factors influencing the organic carbon content of marine aerosols: A laboratory study, *Journal of Geophysical 975 Research*, 81, 3667–3670, 1976.
- Hong, S.-Y., Noh, Y., and Dudhia, J.: A new vertical diffusion package with an explicit treatment of entrainment processes, *Monthly weather review*, 134, 2318–2341, 2006.
- Hoppel, W. A., Frick, G. M., and Fitzgerald, J. W.: Surface source function for sea-salt aerosol and aerosol dry deposition to the ocean surface, *Journal of Geophysical Research: Atmospheres*, 107, AAC 7–1–AAC 7–17, <https://doi.org/https://doi.org/10.1029/2001JD002014>, 2002.
- 980 Huang, J. and Jaeglé, L.: Wintertime enhancements of sea salt aerosol in polar regions consistent with a sea ice source from blowing snow, *Atmospheric Chemistry and Physics*, 17, 3699–3712, <https://doi.org/10.5194/acp-17-3699-2017>, 2017.
- Huang, Z., Huang, J., Hayasaka, T., Wang, S., Zhou, T., and Jin, H.: Short-cut transport path for Asian dust directly to the Arctic: a case study, *Environmental Research Letters*, 10, 114018, 2015.
- Iacono, M. J., D. J. S. M. E. J. S. M. W. C. S. A. and D., C. W.: Radiative forcing by long-lived greenhouse gases: Calculations with the 985 AER radiative transfer models, *J. Geophys. Res.*, 113, <https://doi.org/doi:10.1029/2008JD009944>, 2008.



- Jacobi, H., Voisin, D., Jaffrezo, J., Cozic, J., and Douglas, T.: Chemical composition of the snowpack during the OASIS spring campaign 2009 at Barrow, Alaska, *Journal of Geophysical Research: Atmospheres*, 117, 2012.
- Jaeglé, L., Quinn, P. K., Bates, T. S., Alexander, B., and Lin, J.-T.: Global distribution of sea salt aerosols: new constraints from in situ and remote sensing observations, *Atmospheric Chemistry and Physics*, 11, 3137–3157, <https://doi.org/10.5194/acp-11-3137-2011>, 2011.
- 990 Jiménez, P. A., Dudhia, J., González-Rouco, J. F., Navarro, J., Montávez, J. P., and García-Bustamante, E.: A revised scheme for the WRF surface layer formulation, *Monthly weather review*, 140, 898–918, 2012.
- Jordan, R.: A One-dimensional temperature model for a snow cover : technical documentation for SNThERM.89, 1991.
- Kelly, J. T., Bhave, P. V., Nolte, C. G., Shankar, U., and Foley, K. M.: Simulating emission and chemical evolution of coarse sea-salt particles in the Community Multiscale Air Quality (CMAQ) model, *Geoscientific Model Development*, 3, 257–273, <https://doi.org/10.5194/gmd-3-257-2010>, 2010.
- 995 Kirpes, R. M., Bondy, A. L., Bonanno, D., Moffet, R. C., Wang, B., Laskin, A., Ault, A. P., and Pratt, K. A.: Secondary sulfate is internally mixed with sea spray aerosol and organic aerosol in the winter Arctic, *Atmospheric Chemistry and Physics*, 18, 3937–3949, <https://doi.org/10.5194/acp-18-3937-2018>, 2018.
- Kirpes, R. M., Bonanno, D., May, N. W., Fraund, M., Barget, A. J., Moffet, R. C., Ault, A. P., and Pratt, K. A.: Winter-time Arctic Sea Spray Aerosol Composition Controlled by Sea Ice Lead Microbiology, *ACS Central Science*, 5, 1760–1767, <https://doi.org/10.1021/acscentsci.9b00541>, 2019.
- 1000 Kolesar, K. R., Cellini, J., Peterson, P. K., Jefferson, A., Tuch, T., Birmili, W., Wiedensohler, A., and Pratt, K. A.: Effect of Prudhoe Bay emissions on atmospheric aerosol growth events observed in Utqiagvik (Barrow), Alaska, *Atmospheric Environment*, 152, 146–155, <https://doi.org/https://doi.org/10.1016/j.atmosenv.2016.12.019>, 2017.
- 1005 Krembs, C., Eicken, H., Junge, K., and Deming, J.: High concentrations of exopolymeric substances in Arctic winter sea ice: implications for the polar ocean carbon cycle and cryoprotection of diatoms, *Deep Sea Research Part I: Oceanographic Research Papers*, 49, 2163–2181, [https://doi.org/https://doi.org/10.1016/S0967-0637\(02\)00122-X](https://doi.org/https://doi.org/10.1016/S0967-0637(02)00122-X), 2002.
- Lange, R., Dall’Osto, M., Skov, H., Nøjgaard, J., Nielsen, I., Beddows, D., Simó, R., Harrison, R. M., and Massling, A.: Characterization of distinct Arctic aerosol accumulation modes and their sources, *Atmospheric Environment*, 183, 1–10, 2018.
- 1010 Law, K. S., Stohl, A., Quinn, P. K., Brock, C. A., Burkhardt, J. F., Paris, J.-D., Ancellet, G., Singh, H. B., Roiger, A., Schlager, H., et al.: Arctic air pollution: New insights from POLARCAT-IPY, *Bulletin of the American Meteorological Society*, 95, 1873–1895, 2014.
- Law, K. S., Roiger, A., Thomas, J. L., Marelle, L., Raut, J.-C., Dalsøren, S., Fuglestedt, J., Tuccella, P., Weinzierl, B., and Schlager, H.: Local Arctic air pollution: Sources and impacts, *Ambio*, 46, 453–463, 2017.
- Leaich, W. R., Russell, L. M., Liu, J., Kolonjari, F., Toom, D., Huang, L., Sharma, S., Chivulescu, A., Veber, D., and Zhang, W.: Organic functional groups in the submicron aerosol at 82.5° N, 62.5° W from 2012 to 2014, *Atmospheric Chemistry and Physics*, 18, 3269–3287, <https://doi.org/10.5194/acp-18-3269-2018>, 2018.
- 1015 Leck, C., Norman, M., Bigg, E. K., and Hillamo, R.: Chemical composition and sources of the high Arctic aerosol relevant for cloud formation, *Journal of Geophysical Research: Atmospheres*, 107, AAC 1–1–AAC 1–17, <https://doi.org/https://doi.org/10.1029/2001JD001463>, 2002.
- 1020 Liu, S., Liu, C.-C., Froyd, K. D., Schill, G. P., Murphy, D. M., Bui, T. P., Dean-Day, J. M., Weinzierl, B., Dollner, M., Diskin, G. S., et al.: Sea spray aerosol concentration modulated by sea surface temperature, *Proceedings of the National Academy of Sciences*, 118, 2021.
- Lovejoy C., W. F. Vincent, S. B. S. R. M.-J. M. R. T. M. P. R. M. C. P.: Distribution, Phylogeny, and Growth of Cold-Adapted Picoprasinophytes in Arctic Seas, *J. Phycol.*, 43, 78–89, <https://doi.org/https://doi.org/10.1111/j.1529-8817.2006.00310.x>, 2007.



- 1025 Ma, X., von Salzen, K., and Li, J.: Modelling sea salt aerosol and its direct and indirect effects on climate, *Atmospheric Chemistry and Physics*, 8, 1311–1327, <https://doi.org/10.5194/acp-8-1311-2008>, 2008.
- Malm, W. C., Sisler, J. F., Huffman, D., Eldred, R. A., and Cahill, T. A.: Spatial and seasonal trends in particle concentration and optical extinction in the United States, *Journal of Geophysical Research: Atmospheres*, 99, 1347–1370, <https://doi.org/https://doi.org/10.1029/93JD02916>, 1994.
- 1030 Marelle, L., Raut, J.-C., Law, K. S., Berg, L. K., Fast, J. D., Easter, R. C., Shrivastava, M., and Thomas, J. L.: Improvements to the WRF-Chem 3.5.1 model for quasi-hemispheric simulations of aerosols and ozone in the Arctic, *Geoscientific Model Development*, 10, 3661–3677, <https://doi.org/10.5194/gmd-10-3661-2017>, 2017.
- Marelle, L., Raut, J.-C., Law, K. S., and Duclaux, O.: Current and future arctic aerosols and ozone from remote emissions and emerging local sources—Modeled source contributions and radiative effects, *Journal of Geophysical Research: Atmospheres*, 123, 12,942–12,963, <https://doi.org/https://doi.org/10.1029/2018JD0288637>, 2018.
- 1035 Marelle, L., Thomas, J. L., Ahmed, S., Tuite, K., Stutz, J., Dommergue, A., Simpson, W. R., Frey, M. M., and Baladima, F.: Implementation and Impacts of Surface and Blowing Snow Sources of Arctic Bromine Activation Within WRF-Chem 4.1.1, *Journal of Advances in Modeling Earth Systems*, 13, e2020MS002391, <https://doi.org/https://doi.org/10.1029/2020MS002391>, e2020MS002391 2020MS002391, 2021.
- 1040 Massling, A., Nielsen, I., Kristensen, D., Christensen, J., Sørensen, L., Jensen, B., Nguyen, Q., Nøjgaard, J., Glasius, M., and Skov, H.: Atmospheric black carbon and sulfate concentrations in Northeast Greenland, *Atmospheric Chemistry and Physics*, 15, 9681–9692, 2015.
- May, N. W., Quinn, P. K., McNamara, S. M., and Pratt, K. A.: Multiyear study of the dependence of sea salt aerosol on wind speed and sea ice conditions in the coastal Arctic, *Journal of Geophysical Research: Atmospheres*, 121, 9208–9219, <https://doi.org/https://doi.org/10.1002/2016JD025273>, 2016.
- 1045 McCabe, J. R., Savarino, J., Alexander, B., Gong, S., and Thiemens, M. H.: Isotopic constraints on non-photochemical sulfate production in the Arctic winter, *Geophysical research letters*, 33, 2006.
- McNamara, S. M., Kolesar, K. R., Wang, S., Kirpes, R. M., May, N. W., Gunch, M. J., Cook, R. D., Fuentes, J. D., Hornbrook, R. S., Apel, E. C., et al.: Observation of road salt aerosol driving inland wintertime atmospheric chlorine chemistry, *ACS central science*, 6, 684–694, 2020.
- 1050 Millero, F. J., Feistel, R., Wright, D. G., and McDougall, T. J.: The composition of Standard Seawater and the definition of the Reference-Composition Salinity Scale, *Deep Sea Research Part I: Oceanographic Research Papers*, 55, 50–72, <https://doi.org/https://doi.org/10.1016/j.dsr.2007.10.001>, 2008.
- Monaghan, A. J., Clark, M. P., Barlage, M. P., Newman, A. J., Xue, L., Arnold, J. R., and Rasmussen, R. M.: High-Resolution Historical Climate Simulations over Alaska, *Journal of Applied Meteorology and Climatology*, 57, 709–731, <https://www.jstor.org/stable/26501015>, 2018.
- 1055 Monahan, E. and Muircheartaigh, I.: Optimal Power-Law Description of Oceanic Whitecap Coverage Dependence on Wind Speed, *Journal of Physical Oceanography*, 10, 2094–2099, 1980.
- Monahan, E., Spiel, D. E., and Davidson, K.: A Model of Marine Aerosol Generation Via Whitecaps and Wave Disruption, 1986.
- Morrison, H., T. G. T. V.: Impact of Cloud Microphysics on the Development of Trailing Stratiform Precipitation in a Simulated Squall Line: Comparison of One- and Two-Moment Schemes, *Mon. Wea. Rev.*, 137, 991—1007, <https://doi.org/10.1175/2008MWR2556.1>, 2009.



- 1060 Moschos, V., Schmale, J., Aas, W., Becagli, S., Calzolari, G., Eleftheriadis, K., Moffett, C. E., Schnelle-Kreis, J., Severi, M., Sharma, S., et al.: Elucidating the present-day chemical composition, seasonality and source regions of climate-relevant aerosols across the Arctic land surface, *Environmental Research Letters*, 17, 034 032, 2022.
- National Centers for Environmental Prediction, National Weather Service, NOAA, U.S. Department of Commerce: NCEP FNL Operational Model Global Tropospheric Analyses, continuing from July 1999, <https://doi.org/10.5065/D6M043C6>, 2000.
- 1065 Neumann, D., Matthias, V., Bieser, J., Aulinger, A., and Quante, M.: Sensitivity of modeled atmospheric nitrogen species and nitrogen deposition to variations in sea salt emissions in the North Sea and Baltic Sea regions, *Atmospheric Chemistry and Physics*, 16, 2921–2942, <https://doi.org/10.5194/acp-16-2921-2016>, 2016.
- Nguyen, Q., Skov, H., Sørensen, L., Jensen, B., Grube, A., Massling, A., Glasius, M., and Nøjgaard, J.: Source apportionment of particles at Station Nord, North East Greenland during 2008–2010 using COPREM and PMF analysis, *Atmospheric Chemistry and Physics*, 13, 35–49, 2013.
- 1070 Nielsen, I. E., Skov, H., Massling, A., Eriksson, A. C., Dall’Osto, M., Junninen, H., Sarnela, N., Lange, R., Collier, S., Zhang, Q., et al.: Biogenic and anthropogenic sources of aerosols at the High Arctic site Villum Research Station, *Atmospheric Chemistry and Physics*, 19, 10 239–10 256, 2019.
- Nilsson, E. D. and Rannik, : Turbulent aerosol fluxes over the Arctic Ocean: 1. Dry deposition over sea and pack ice, *Journal of Geophysical Research: Atmospheres*, 106, 32 125–32 137, <https://doi.org/https://doi.org/10.1029/2000JD900605>, 2001.
- 1075 Nilsson, E. D., Rannik, , Swietlicki, E., Leck, C., Aalto, P. P., Zhou, J., and Norman, M.: Turbulent aerosol fluxes over the Arctic Ocean: 2. Wind-driven sources from the sea, *Journal of Geophysical Research: Atmospheres*, 106, 32 139–32 154, <https://doi.org/https://doi.org/10.1029/2000JD900747>, 2001.
- Niu, G.-Y. and Yang, Z.-L.: Effects of vegetation canopy processes on snow surface energy and mass balances, *Journal of Geophysical Research: Atmospheres*, 109, <https://doi.org/https://doi.org/10.1029/2004JD004884>, 2004.
- 1080 Niu, G.-Y., Yang, Z.-L., Dickinson, R. E., Gulden, L. E., and Su, H.: Development of a simple groundwater model for use in climate models and evaluation with Gravity Recovery and Climate Experiment data, *Journal of Geophysical Research: Atmospheres*, 112, <https://doi.org/https://doi.org/10.1029/2006JD007522>, 2007.
- Niu, G.-Y., Yang, Z.-L., Mitchell, K. E., Chen, F., Ek, M. B., Barlage, M., Kumar, A., Manning, K., Niyogi, D., Rosero, E., Tewari, M., and Xia, Y.: The community Noah land surface model with multiparameterization options (Noah-MP): 1. Model description and evaluation with local-scale measurements, *Journal of Geophysical Research: Atmospheres*, 116, <https://doi.org/https://doi.org/10.1029/2010JD015139>, 2011.
- 1085 O’Dowd, C. D., Smith, M. H., Consterdine, I. E., and Lowe, J. A.: Marine aerosol, sea-salt, and the marine sulphur cycle: a short review, *Atmospheric Environment*, 31, 73–80, [https://doi.org/https://doi.org/10.1016/S1352-2310\(96\)00106-9](https://doi.org/https://doi.org/10.1016/S1352-2310(96)00106-9), 1997.
- 1090 Panno, S., Hackley, K. C., Hwang, H., Greenberg, S., Krapac, I., Landsberger, S., and O’kelly, D.: Characterization and identification of Na-Cl sources in ground water, *Groundwater*, 44, 176–187, 2006.
- Peterson, P. K., Pöhler, D., Sihler, H., Zielcke, J., General, S., Frieß, U., Platt, U., Simpson, W. R., Nghiem, S. V., Shepson, P. B., et al.: Observations of bromine monoxide transport in the Arctic sustained on aerosol particles, *Atmospheric Chemistry and Physics*, 17, 7567–7579, 2017.
- 1095 Pilson, M. E. Q.: *An Introduction to the Chemistry of the Sea*, Cambridge University Press, 2 edn., <https://doi.org/10.1017/CBO9781139047203>, 2012.



- Qi, L., Li, Q., Henze, D. K., Tseng, H.-L., and He, C.: Sources of springtime surface black carbon in the Arctic: an adjoint analysis for April 2008, *Atmospheric Chemistry and Physics*, 17, 9697–9716, <https://doi.org/10.5194/acp-17-9697-2017>, 2017.
- 1100 Quinn, P. K. and Bates, T. S.: Regional aerosol properties: Comparisons of boundary layer measurements from ACE 1, ACE 2, Aerosols99, INDOEX, ACE Asia, TARFOX, and NEAQS, *Journal of Geophysical Research: Atmospheres*, 110, 2005.
- Quinn, P. K., Coffman, D. J., Kapustin, V. N., Bates, T. S., and Covert, D. S.: Aerosol optical properties in the marine boundary layer during the First Aerosol Characterization Experiment (ACE 1) and the underlying chemical and physical aerosol properties, *Journal of Geophysical Research: Atmospheres*, 103, 16 547–16 563, <https://doi.org/https://doi.org/10.1029/97JD02345>, 1998.
- 1105 Quinn, P. K., Miller, T. L., Bates, T. S., Ogren, J. A., Andrews, E., and Shaw, G. E.: A 3-year record of simultaneously measured aerosol chemical and optical properties at Barrow, Alaska, *Journal of Geophysical Research: Atmospheres*, 107, AAC 8–1–AAC 8–15, <https://doi.org/https://doi.org/10.1029/2001JD001248>, 2002.
- Quinn, P. K., SHAW, G., ANDREWS, E., DUTTON, E. G., RUOHO-AIROLA, T., and GONG, S. L.: Arctic haze: current trends and knowledge gaps, *Tellus B*, 59, 99–114, <https://doi.org/https://doi.org/10.1111/j.1600-0889.2006.00238.x>, 2007.
- 1110 Rahn, K. A. and McCaffrey, R. J.: On the origin and transport of the winter Arctic aerosol, *Annals of the New York Academy of Sciences*, 338, 486–503, 1980.
- Rahn, K. A., Borys, R. D., and Shaw, G. E.: The Asian source of Arctic haze bands, *Nature*, 268, 713–715, 1977.
- Revell, L. E., Kremser, S., Hartery, S., Harvey, M., Mulcahy, J. P., Williams, J., Morgenstern, O., McDonald, A. J., Varma, V., Bird, L., and Schuddeboom, A.: The sensitivity of Southern Ocean aerosols and cloud microphysics to sea spray and sulfate aerosol production in the HadGEM3-GA7.1 chemistry–climate model, *Atmospheric Chemistry and Physics*, 19, 15 447–15 466, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-19-15447-2019)
- 1115 [19-15447-2019](https://doi.org/10.5194/acp-19-15447-2019), 2019.
- Rhodes, R. H., Yang, X., Wolff, E. W., McConnell, J. R., and Frey, M. M.: Sea ice as a source of sea salt aerosol to Greenland ice cores: a model-based study, *Atmospheric Chemistry and Physics*, 17, 9417–9433, <https://doi.org/10.5194/acp-17-9417-2017>, 2017.
- Roscoe, H. K., Brooks, B., Jackson, A., Smith, M., Walker, S., Obbard, R. W., and Wolff, E. W.: Frost flowers in the laboratory: Growth, characteristics, aerosol, and the underlying sea ice, *Journal of Geophysical Research: Atmospheres*, 116, 2011.
- 1120 Russell, L. M., Hawkins, L. N., Frossard, A. A., Quinn, P. K., and Bates, T. S.: Carbohydrate-like composition of submicron atmospheric particles and their production from ocean bubble bursting, *Proceedings of the National Academy of Sciences*, 107, 6652–6657, <https://doi.org/10.1073/pnas.0908905107>, 2010.
- Sakaguchi, K. and Zeng, X.: Effects of soil wetness, plant litter, and under-canopy atmospheric stability on ground evaporation in the Community Land Model (CLM3.5), *Journal of Geophysical Research: Atmospheres*, 114, <https://doi.org/https://doi.org/10.1029/2008JD010834>,
- 1125 [2009](https://doi.org/https://doi.org/10.1029/2008JD010834).
- Saliba, G., Chen, C.-L., Lewis, S., Russell, L. M., Rivellini, L.-H., Lee, A. K. Y., Quinn, P. K., Bates, T. S., Haëntjens, N., Boss, E. S., Karp-Boss, L., Baetge, N., Carlson, C. A., and Behrenfeld, M. J.: Factors driving the seasonal and hourly variability of sea-spray aerosol number in the North Atlantic, *Proceedings of the National Academy of Sciences*, 116, 20 309–20 314, <https://doi.org/10.1073/pnas.1907574116>, 2019.
- 1130 Salisbury, D. J., Anguelova, M. D., and Brooks, I. M.: On the variability of whitecap fraction using satellite-based observations, *Journal of Geophysical Research: Oceans*, 118, 6201–6222, <https://doi.org/https://doi.org/10.1002/2013JC008797>, 2013.
- Salisbury, D. J., Anguelova, M. D., and Brooks, I. M.: Global distribution and seasonal dependence of satellite-based whitecap fraction, *Geophysical Research Letters*, 41, 1616–1623, <https://doi.org/https://doi.org/10.1002/2014GL059246>, 2014.



- 1135 Salter, M. E., Nilsson, E. D., Butcher, A., and Bilde, M.: On the seawater temperature dependence of the sea spray aerosol generated by a continuous plunging jet, *Journal of Geophysical Research: Atmospheres*, 119, 9052–9072, 2014.
- Salter, M. E., Zieger, P., Acosta Navarro, J. C., Grythe, H., Kirkevåg, A., Rosati, B., Riipinen, I., and Nilsson, E. D.: An empirically derived inorganic sea spray source function incorporating sea surface temperature, *Atmospheric Chemistry and Physics*, 15, 11 047–11 066, <https://doi.org/10.5194/acp-15-11047-2015>, 2015.
- 1140 Sato, Y., Miura, H., Yashiro, H., Goto, D., Takemura, T., Tomita, H., and Nakajima, T.: Unrealistically pristine air in the Arctic produced by current global scale models, *Scientific Reports*, 6, <https://doi.org/10.1038/srep26561>, 2016.
- Schmale, J., Arnold, S., Law, K. S., Thorp, T., Anenberg, S., Simpson, W., Mao, J., and Pratt, K.: Local Arctic air pollution: A neglected but serious problem, *Earth's Future*, 6, 1385–1412, 2018.
- Schroeder, W. H., Anlauf, K., Barrie, L., Lu, J., Steffen, A., Schneeberger, D., and Berg, T.: Arctic springtime depletion of mercury, *Nature*, 394, 331–332, 1998.
- 1145 Schwarz, J. P., Weinzierl, B., Samset, B. H., Dollner, M., Heimerl, K., Markovic, M. Z., Perring, A. E., and Ziemba, L.: Aircraft measurements of black carbon vertical profiles show upper tropospheric variability and stability, *Geophysical Research Letters*, 44, 1132–1140, <https://doi.org/https://doi.org/10.1002/2016GL071241>, 2017.
- Seinfeld, J. and Pandis, S.: *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change*, 1998.
- Seinfeld, J. H.: *Atmospheric chemistry and physics of air pollution*, California, California Institute of Technology Pasadena, 1986.
- 1150 Sellers, P., Randall, D., Collatz, G., Berry, J., Field, C., Dazlich, D., Zhang, C., Collelo, G., and Bounoua, L.: A revised Land Surface parameterization (SiB2) for atmospheric GCMs. Part I: Model Formulation, *Journal of Climate*, 9, 676–705, [https://doi.org/10.1175/1520-0442\(1996\)009<0676:ARLSPF>2.0.CO;2](https://doi.org/10.1175/1520-0442(1996)009<0676:ARLSPF>2.0.CO;2), 1996.
- Sharma, S., Barrie, L. A., Magnusson, E., Brattström, G., Leaitch, W., Steffen, A., and Landsberger, S.: A factor and trends analysis of multidecadal lower tropospheric observations of Arctic aerosol composition, black carbon, ozone, and mercury at Alert, Canada, *Journal of Geophysical Research: Atmospheres*, 124, 14 133–14 161, 2019.
- 1155 Shaw, P. M., Russell, L. M., Jefferson, A., and Quinn, P. K.: Arctic organic aerosol measurements show particles from mixed combustion in spring haze and from frost flowers in winter, *Geophysical Research Letters*, 37, <https://doi.org/https://doi.org/10.1029/2010GL042831>, 2010.
- Shrivastava, M., Fast, J., Easter, R., Gustafson Jr., W. I., Zaveri, R. A., Jimenez, J. L., Saide, P., and Hodzic, A.: Modeling organic aerosols in a megacity: comparison of simple and complex representations of the volatility basis set approach, *Atmospheric Chemistry and Physics*, 11, 6639–6662, <https://doi.org/10.5194/acp-11-6639-2011>, 2011.
- 1160 Simpson, W., Carlson, D., Hönninger, G., Douglas, T., Sturm, M., Perovich, D., and Platt, U.: First-year sea-ice contact predicts bromine monoxide (BrO) levels at Barrow, Alaska better than potential frost flower contact, *Atmospheric Chemistry and Physics*, 7, 621–627, 2007.
- 1165 Slinn, S. and Slinn, W.: Predictions for particle deposition on natural waters, *Atmospheric Environment (1967)*, 14, 1013–1016, 1980.
- Slinn, W.: Predictions for particle deposition to vegetative canopies, *Atmospheric Environment (1967)*, 16, 1785–1794, [https://doi.org/https://doi.org/10.1016/0004-6981\(82\)90271-2](https://doi.org/https://doi.org/10.1016/0004-6981(82)90271-2), precipitation chemistry, 1982.
- Sofiev, M., Soares, J., Prank, M., de Leeuw, G., and Kukkonen, J.: A regional-to-global model of emission and transport of sea salt particles in the atmosphere, *Journal of Geophysical Research: Atmospheres*, 116, <https://doi.org/https://doi.org/10.1029/2010JD014713>, 2011.



- 1170 Spada, M., Jorba, O., Pérez García-Pando, C., Janjic, Z., and Baldasano, J. M.: Modeling and evaluation of the global sea-salt aerosol distribution: sensitivity to size-resolved and sea-surface temperature dependent emission schemes, *Atmospheric Chemistry and Physics*, 13, 11 735–11 755, <https://doi.org/10.5194/acp-13-11735-2013>, 2013.
- Stohl, A., Klimont, Z., Eckhardt, S., Kupiainen, K., Shevchenko, V. P., Kopeikin, V. M., and Novigatsky, A. N.: Black carbon in the Arctic: the underestimated role of gas flaring and residential combustion emissions, *Atmospheric Chemistry and Physics*, 13, 8833–8855, <https://doi.org/10.5194/acp-13-8833-2013>, 2013.
- 1175 Stone, R. S., Sharma, S., Herber, A., Eleftheriadis, K., Nelson, D. W., Helmig, D., and Bottenheim, J. W.: A characterization of Arctic aerosols on the basis of aerosol optical depth and black carbon measurementsA characterization of Arctic aerosols, *Elementa: Science of the Anthropocene*, 2, 2014.
- Stroeve, J., Serreze, M., Holland, M., Kay, J., Malanik, J., and Barrett, A.: The Arctic’s rapidly shrinking sea ice cover: a research synthesis, *Climatic Change*, 110, 1005–1027, <https://doi.org/10.1007/s10584-011-0101-1>, 2012.
- 1180 Wang, S., McNamara, S. M., Moore, C. W., Obrist, D., Steffen, A., Shepson, P. B., Staebler, R. M., Raso, A. R., and Pratt, K. A.: Direct detection of atmospheric atomic bromine leading to mercury and ozone depletion, *Proceedings of the National Academy of Sciences*, 116, 14 479–14 484, 2019a.
- Wang, X., Jacob, D. J., Eastham, S. D., Sulprizio, M. P., Zhu, L., Chen, Q., Alexander, B., Sherwen, T., Evans, M. J., Lee, B. H., et al.: The role of chlorine in global tropospheric chemistry, *Atmospheric Chemistry and Physics*, 19, 3981–4003, 2019b.
- 1185 Whaley, C. H., Makar, P. A., Shephard, M. W., Zhang, L., Zhang, J., Zheng, Q., Akingunola, A., Wentworth, G. R., Murphy, J. G., Kharol, S. K., et al.: Contributions of natural and anthropogenic sources to ambient ammonia in the Athabasca Oil Sands and north-western Canada, *Atmospheric Chemistry and Physics*, 18, 2011–2034, 2018.
- Whaley, C. H., Mahmood, R., von Salzen, K., Winter, B., Eckhardt, S., Arnold, S., Beagley, S., Becagli, S., Chien, R.-Y., Christensen, J., Damani, S. M., Eleftheriadis, K., Evangeliou, N., Faluvegi, G. S., Flanner, M., Fu, J. S., Gauss, M., Giardi, F., Gong, W., Hjorth, J. L., Huang, L., Im, U., Kanaya, Y., Krishnan, S., Klimont, Z., Kühn, T., Langner, J., Law, K. S., Marelle, L., Massling, A., Olivíe, D., Onishi, T., Oshima, N., Peng, Y., Plummer, D. A., Popovicheva, O., Pozzoli, L., Raut, J.-C., Sand, M., Saunders, L. N., Schmale, J., Sharma, S., Skov, H., Taketani, F., Thomas, M. A., Traversi, R., Tsigaridis, K., Tsyro, S., Turnock, S., Vitale, V., Walker, K. A., Wang, M., Watson-Parris, D., and Weiss-Gibbons, T.: Model evaluation of short-lived climate forcers for the Arctic Monitoring and Assessment Programme: a multi-species, multi-model study, *Atmospheric Chemistry and Physics Discussions*, 2022, 1–88, <https://doi.org/10.5194/acp-2021-975>, 2022.
- 1195 Wild, O., Zhu, X., and Prather, M.: Fast-J: Accurate Simulation of In- and Below-Cloud Photolysis in Tropospheric Chemical Models, *Journal of Atmospheric Chemistry*, 37, 245–282, 2000.
- Winiger, P., Barrett, T., Sheesley, R., Huang, L., Sharma, S., Barrie, L. A., Yttri, K. E., Evangeliou, N., Eckhardt, S., Stohl, A., et al.: Source apportionment of circum-Arctic atmospheric black carbon from isotopes and modeling, *Science advances*, 5, eaau8052, 2019.
- 1200 Witek, M. L., Flatau, P. J., Quinn, P. K., and Westphal, D. L.: Global sea-salt modeling: Results and validation against multicampaign shipboard measurements, *Journal of Geophysical Research:Atmospheres*, 112, 2007.
- Xu, J.-W., Martin, R. V., Morrow, A., Sharma, S., Huang, L., Leaitch, W. R., Burkart, J., Schulz, H., Zanatta, M., Willis, M. D., Henze, D. K., Lee, C. J., Herber, A. B., and Abbatt, J. P. D.: Source attribution of Arctic black carbon constrained by aircraft and surface measurements, *Atmospheric Chemistry and Physics*, 17, 11 971–11 989, <https://doi.org/10.5194/acp-17-11971-2017>, 2017.
- 1205



- Xu, L., Russell, L. M., Somerville, R. C. J., and Quinn, P. K.: Frost flower aerosol effects on Arctic wintertime longwave cloud radiative forcing, *Journal of Geophysical Research: Atmospheres*, 118, 13,282–13,291, <https://doi.org/https://doi.org/10.1002/2013JD020554>, 2013.
- Xu, L., Russell, L. M., and Burrows, S. M.: Potential sea salt aerosol sources from frost flowers in the pan-Arctic region, *Journal of Geophysical Research: Atmospheres*, 121, 10,840–10,856, <https://doi.org/https://doi.org/10.1002/2015JD024713>, 2016.
- 1210 Yang, R. and Friedl, M. A.: Modeling the effects of three-dimensional vegetation structure on surface radiation and energy balance in boreal forests, *Journal of Geophysical Research: Atmospheres*, 108, <https://doi.org/https://doi.org/10.1029/2002JD003109>, 2003.
- Yang Z.-L., R. E. Dickinson, A. R. and Vinnikov., K. Y.: Validation of the snow sub-model of the Biosphere-Atmosphere Transfer Scheme with Russian snow cover and meteorological observational data, *J. Clim.*, 10, 353—373, [https://doi.org/10.1175/1520-0442\(1997\)010<0353:VOTSSO>2.0.CO;2](https://doi.org/10.1175/1520-0442(1997)010<0353:VOTSSO>2.0.CO;2), 1997.
- 1215 yue Niu, G. and liang Yang, Z.: Effects of frozen soil on snowmelt runoff and soil water storage at a continental scale, *J. Hydrometeorol.*, pp. 937–952, 2006.
- Zaveri, R. A., Easter, R. C., Fast, J. D., and Peters, L. K.: Model for Simulating Aerosol Interactions and Chemistry (MOSAIC), *Journal of Geophysical Research: Atmospheres*, 113, <https://doi.org/https://doi.org/10.1029/2007JD008782>, 2008.
- 1220 Zhang, L., Gong, S., Padro, J., and Barrie, L.: A size-segregated particle dry deposition scheme for an atmospheric aerosol module, *Atmospheric Environment*, 35, 549–560, [https://doi.org/https://doi.org/10.1016/S1352-2310\(00\)00326-5](https://doi.org/https://doi.org/10.1016/S1352-2310(00)00326-5), 2001.
- Zwaafink, C. G., Grythe, H., Skov, H., and Stohl, A.: Substantial contribution of northern high-latitude sources to mineral dust in the Arctic, *Journal of Geophysical Research: Atmospheres*, 121, 13–678, 2016.