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► To cite this version:

Peter Bergamaschi, Arjo Segers, Dominik Brunner, Jean-Matthieu Haussaire, Stephan Henne, et al.. High-resolution inverse modelling of European CH₄ emissions using the novel FLEXPART-COSMO TM5 4DVAR inverse modelling system. Atmospheric Chemistry and Physics, 2022, 22, pp.13243-13268. 10.5194/acp-22-13243-2022 . insu-03824352

HAL Id: insu-03824352

<https://insu.hal.science/insu-03824352>

Submitted on 21 Oct 2022

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High-resolution inverse modelling of European CH₄ emissions using the novel FLEXPART-COSMO TM5 4DVAR inverse modelling system

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Received: 18 February 2022 – Discussion started: 16 March 2022

Revised: 27 August 2022 – Accepted: 4 September 2022 – Published: 17 October 2022

Abstract. We present a novel high-resolution inverse modelling system (“FLEXVAR”) based on FLEXPART-COSMO back trajectories driven by COSMO meteorological fields at 7 km × 7 km resolution over the European

COSMO-7 domain and the four-dimensional variational (4DVAR) data assimilation technique. FLEXVAR is coupled offline with the global inverse modelling system TM5-4DVAR to provide background mole fractions (“baselines”) consistent with the global observations assimilated in TM5-4DVAR. We have applied the FLEXVAR system for the inverse modelling of European CH₄ emissions in 2018 using 24 stations with in situ measurements, complemented with data from five stations with discrete air sampling (and additional stations outside the European COSMO-7 domain used for the global TM5-4DVAR inversions). The sensitivity of the FLEXVAR inversions to different approaches to calculate the baselines, different parameterizations of the model representation error, different settings of the prior error covariance parameters, different prior inventories, and different observation data sets are investigated in detail. Furthermore, the FLEXVAR inversions are compared to inversions with the FLEXPART extended Kalman filter (“FLEXKF”) system and with TM5-4DVAR inversions at $1^\circ \times 1^\circ$ resolution over Europe. The three inverse modelling systems show overall good consistency of the major spatial patterns of the derived inversion increments and in general only relatively small differences in the derived annual total emissions of larger country regions. At the same time, the FLEXVAR inversions at $7\text{ km} \times 7\text{ km}$ resolution allow the observations to be better reproduced than the TM5-4DVAR simulations at $1^\circ \times 1^\circ$. The three inverse models derive higher annual total CH₄ emissions in 2018 for Germany, France, and BENELUX compared to the sum of anthropogenic emissions reported to UNFCCC and natural emissions estimated from the Global Carbon Project CH₄ inventory, but the uncertainty ranges of top-down and bottom-up total emission estimates overlap for all three country regions. In contrast, the top-down estimates for the sum of emissions from the UK and Ireland agree relatively well with the total of anthropogenic and natural bottom-up inventories.

1 Introduction

Atmospheric methane (CH₄) is the second most important anthropogenic greenhouse gas (GHG) after carbon dioxide (CO₂), with an estimated contribution of $\sim 16.3\%$ (0.520 W m^{-2}) to the direct anthropogenic radiative forcing of all long-lived GHGs in 2020 (NOAA Annual Greenhouse Gas Index (AGGI), evaluated relative to 1750; Butler and Montzka, 2022). Including also additional indirect effects (e.g. production of tropospheric ozone), however, the total radiative forcing of CH₄ is considerably higher, with current estimates of the emission-based effective radiative forcing (ERF) of 1.21 (0.90 to 1.51) W m^{-2} (Szopa et al., 2021). The current global average CH₄ mole fraction is 162% higher than preindustrial levels in 1750 (WMO, 2021) and continues to increase, with recent growth rates (2014–2020: $10.1 \pm 3.2\text{ ppb yr}^{-1}$) being again close to the high growth rates observed during the 1980s (1984–1989: $11.9 \pm 0.9\text{ ppb yr}^{-1}$), while lower growth rates were observed during the 1990s and almost zero growth rates during 2000–2006 (Dlugokencky, 2022).

Reducing CH₄ emissions plays an essential role in mitigating climate change, especially in the near term (Shindell et al., 2012, 2017; United Nations Environment Programme and Climate and Clean Air Coalition, 2021), due to CH₄’s relatively short atmospheric lifetime of around 10 years combined with its high radiative efficiency (resulting in a global warming potential (GWP) around 80 times higher compared to CO₂ on a 20-year timescale; Forster et al., 2021). The global emissions pathways to limit global warming to 1.5°C , compiled by IPCC (2018), include significant reductions of CH₄ emissions after 2020 (for scenarios with no or lim-

ited overshoot of temperature above the 1.5°C target). The recognition of the importance of CH₄ emission reductions to mitigate climate change has also led to the recent “Global Methane Pledge” (European Commission, 2021), with the collective goal to reduce methane emissions by 2030 by at least 30% compared to 2020.

The development of emission reduction pathways as well as the control of international climate agreements requires the accurate quantification of current (and past) GHG emissions. For CH₄, however, the quantification of emissions and sinks is particularly challenging, mainly owing to the large spatial and temporal variability of emissions from many source sectors and consequently large uncertainties in assumed mean emission factors (e.g. for natural emissions from wetlands and anthropogenic emissions from fugitive sources such as fossil fuels (coal, oil, gas) (e.g. Brandt et al., 2014) or emissions from the waste sector). Therefore, bottom-up inventories of CH₄, which are compiled by scaling up emissions using activity data and emission factors, have significant uncertainties. Complementary to bottom-up inventories, inverse modelling provides top-down emission estimates using atmospheric measurements and atmospheric transport models, by optimizing emissions from emission inventories (used as prior estimates) to get an optimal agreement between simulated and observed CH₄ mole fractions, taking into account the uncertainties of prior emission estimates, measurements, and model simulations (e.g. Bergamaschi et al., 2018b; Houweling et al., 2017). The Global Carbon Project CH₄ (GCP-CH₄) provides synthesis analyses of the global CH₄ cycle based on comprehensive sets of different bottom-up and top-down estimates from the international science community working on this topic (Jackson

et al., 2020; Saunio et al., 2020; Stavert et al., 2021). The global inverse models used in these analyses have generally relatively coarse horizontal resolution, in the range of 2.5–6.0° (longitude) $\times$ 1.9–4.0° (latitude) (Saunio et al., 2020). Therefore, such models are mainly suitable to provide information on the global and larger regional scales.

In order to analyse regional emissions in more detail (and more accurately), specific regional inversions have been performed, employing regional atmospheric transport models at higher horizontal resolution (typically in the range of $\sim$ 20–100 km) and making use of the increasing number of regional in situ GHG measurements, which have become available in recent years in particular in Europe and North America (e.g. Bergamaschi et al., 2018a; Ganesan et al., 2015; Lunt et al., 2021; Manning et al., 2011; Miller et al., 2013). The regional models generally require global boundary conditions, which are usually provided from global inverse models. Alternatively, global models with zooming option for the specific region of interest have been employed (Bergamaschi et al., 2018a). A specific purpose of such regional inversions is to verify national bottom-up emission inventories reported to the United Nations Framework Convention on Climate Change (UNFCCC), finally aiming at an emission monitoring and verification system to support the international climate agreements using in situ and satellite observations (Bergamaschi et al., 2018b; Deng et al., 2022; National Research Council, 2010; Pinty et al., 2017, 2019). Within the European project VERIFY (<https://verify.lsce.ipsl.fr/>, last access: 4 October 2022) a pre-operational GHG verification system has been developed, employing various state-of-the-art global and regional atmospheric transport models (Petrescu et al., 2021a, b).

In order to further improve the atmospheric modelling, it is essential to further increase the spatial resolution, aiming at further improving the simulation of regional monitoring stations. A pioneering high-resolution study has been reported by Henne et al. (2016), using the FLEXPART-COSMO back trajectories driven by meteorological fields from the Swiss national weather service (MeteoSwiss) at horizontal resolution of approximately 7 km $\times$ 7 km, analysing the CH₄ emissions from Switzerland using continuous measurements from six atmospheric monitoring stations. The authors generated the FLEXPART-COSMO sensitivity fields (based on the sampling of released particles) at even higher resolution (than the resolution of the COSMO meteorological fields) of 0.02° $\times$ 0.015° ($\approx$ 1.7 km) over the Alpine domain (but coarser horizontal resolution of 0.16° $\times$ 0.12° ($\approx$ 13 km) outside the Alpine domain). Since they solved the inverse problem analytically (denoted by the authors as the “Bayesian method”), however, they applied a reduced grid by merging model grid cells in areas with smaller average source sensitivities in order to reduce the size of the inversion problem to about 1000 unknowns. As an alternative to the Bayesian method, Henne et al. (2016) also applied the extended Kalman filter method described by Brunner et

al. (2012), which – in contrast to the Bayesian method – assimilates the observations sequentially but for computational reasons also requires the application of a reduced grid.

Aiming at high-resolution inversions of larger regions (such as the European domain), we have therefore developed a novel inversion framework (denoted “FLEXVAR”) based on the four-dimensional variational (4DVAR) data assimilation technique (Meirink et al., 2008; Talagrand and Courtier, 1987), which allows for the optimization of a much larger number of parameters and therefore also avoids the need to apply reduced grids. As in the Henne et al. (2016) study, the new system uses FLEXPART-COSMO back trajectories driven by COSMO meteorological fields at 7 km $\times$ 7 km resolution but is optimizing emissions from individual grid cells over the whole European COSMO-7 domain with 393 (longitude) $\times$ 338 (latitude) grid cells. Furthermore, the new system uses background mole fractions (baselines) from global TM5-4DVAR inversions (using two different approaches); i.e. it is coupling the FLEXPART-COSMO inversions with the global/European TM5-4DVAR inversions offline.

The objective of this paper is to present the new FLEXVAR system and its application to the inversion of European CH₄ emissions for 2018 using a comprehensive data set from 24 stations with in situ measurements, complemented with data from 5 stations with discrete air sampling (and additional stations outside the European COSMO-7 domain used for the global TM5-4DVAR inversions). We analyse in detail the sensitivity of the FLEXVAR inversions to internal parameterizations and model settings, as well as the sensitivity to the main model input data, i.e. prior inventories and observational data. Furthermore, we compare the FLEXVAR inversions with the extended Kalman filter (FLEXKF) method and with TM5-4DVAR inversions (at 1° $\times$ 1° resolution over Europe). Finally, we present an overall analysis of derived European CH₄ emissions and comparison with emissions reported to UNFCCC for some major countries (or group of countries) which are best constrained by the available observations (Germany, France, BENELUX, and UK and Ireland).

2 Modelling

2.1 FLEXPART-COSMO back trajectories

FLEXPART is a Lagrangian atmospheric transport model that simulates the advective, turbulent, and convective transport by tracking the positions of a large number of infinitesimally small air parcels, so-called particles, either forward or backward in time (Pisso et al., 2019; Stohl et al., 2005). FLEXPART is an offline model that requires meteorological fields such as 3D wind fields from a numerical weather prediction (NWP) model as input. FLEXPART-COSMO is a version of the model that is driven by the output of the NWP model COSMO, which was jointly developed by a consortium of European weather services under the lead of the German meteorological service DWD (Baldauf et al., 2011).

Different from all other FLEXPART versions, FLEXPART-COSMO operates on the native vertical grid of the driving model COSMO, which avoids potential loss of information and inaccuracies associated with the interpolation onto a different grid. More details on the model are provided in Henne et al. (2016) and Pisso et al. (2019).

In the backward mode, particles are released at the locations of individual observations and followed backwards in time over typically a few days. By sampling the near-surface residence times of the particles along their paths, a so-called source–receptor sensitivity matrix or “footprint” is computed, which describes the relationship between the change in mole fraction at the observation site and the fluxes discretized in space and time (Seibert and Frank, 2004). A time series of simulated mole fractions can be obtained by integrating the time series of source–receptor matrices with a discretized flux estimate. The simulations were driven by hourly output from the operational COSMO-7 analyses of the Swiss weather service MeteoSwiss at a horizontal spatial resolution of about 7 km × 7 km, largely covering western and central Europe (Fig. 1). In the simulations used here, 50 000 virtual particles were released at all observation locations every 3 h (evenly distributed over each 3 h time interval) and traced backwards in time for 10 d (or until individual particles left the COSMO-7 domain). Despite the high spatial resolution, the orography of the COSMO-7 model is smoothed for complex terrain, leading to differences between the model surface altitude and the altitude of the observation site. When this difference is greater than 200 m, the release height of the particles has been chosen as the average between the measurement height above sea level and the model surface altitude. This approach has been used in previous studies (Henne et al., 2016) and was found to be the most representative release height (Brunner et al., 2013).

2.2 Coupled FLEXPART-COSMO TM5 4DVAR inverse modelling system FLEXVAR

2.2.1 Inversion framework

The new coupled FLEXPART-COSMO TM5 4DVAR inverse modelling system, denoted FLEXVAR, allows for the optimization of emissions at grid scale using the four-dimensional variational (4DVAR) data assimilation technique (Meirink et al., 2008; Talagrand and Courtier, 1987), the FLEXPART-COSMO back trajectories described in Sect. 2.1, and background mole fractions (“baselines”) from TM5-4DVAR (described in Sect. 2.2.2 and 2.4). The system follows the classical Bayesian approach minimizing the cost function $J(\mathbf{x})$:

$$J(\mathbf{x}) = \frac{1}{2}(\mathbf{x} - \mathbf{x}_b)^T \mathbf{B}^{-1}(\mathbf{x} - \mathbf{x}_b) + \frac{1}{2}(\mathbf{H}(\mathbf{x}) - \mathbf{y})^T \mathbf{R}^{-1}(\mathbf{H}(\mathbf{x}) - \mathbf{y}), \quad (1)$$

where $\mathbf{x}$ is the state vector, $\mathbf{x}_b$ the prior estimate of the state vector (in data assimilation usually called the “background”), $\mathbf{y}$ the set of observations (measurements) to be assimilated, $\mathbf{H}(\mathbf{x})$ the observation operator (or model operator), representing the model simulation of the observations, and $\mathbf{B}$ and $\mathbf{R}$ the error covariance matrices of the prior estimate and the observations, respectively. For the regular inversions, a semi-lognormal probability density function (pdf) is applied for the emissions $\mathbf{e}$ to be optimized (Bergamaschi et al., 2010), optimizing the emission deviation factors $x_{i,j,t}$,

$$e_{i,j,t}(x_{i,j,t}) = \begin{cases} e_{b,i,j,t} \exp(x_{i,j,t}) & \text{for } x_{i,j,t} < 0 \\ e_{b,i,j,t} (1 + x_{i,j,t}) & \text{for } x_{i,j,t} \geq 0 \end{cases}, \quad (2)$$

for each element $e_{i,j,t}$, representing the emissions of an individual grid cell with longitude index i , latitude index j , and at emission time step t (and $e_{b,i,j,t}$ the prior estimate of the emission $e_{i,j,t}$). In contrast, a linear expansion of the emission deviation factors is used (resulting in a Gaussian pdf of the emissions) for additional inversions for the evaluation of posterior uncertainties (as will be described in more detail below):

$$e_{i,j,t}(x_{i,j,t}) = e_{b,i,j,t} (1 + x_{i,j,t}). \quad (3)$$

For the prior estimate, the emission deviation factors $x_{i,j,t}$ are set to zero, i.e. $e_{i,j,t}(x_{i,j,t}) = e_{b,i,j,t}$ (both for the semi-lognormal and Gaussian pdf). In this study we present inversions optimizing the monthly total emissions of all FLEXPART-COSMO grid cells (at 7 km × 7 km resolution) for 1 year (2018); hence the dimension of the state vector is

$$n_{\text{state}} = n_{\text{longitude}} \times n_{\text{latitude}} \times n_{\text{time}} = 393 \times 338 \times 12 = 1\,594\,008. \quad (4)$$

The background covariance matrix $\mathbf{B}$ is parameterized as the product

$$\mathbf{B} = \mathbf{S} \mathbf{C} \mathbf{S}, \quad (5)$$

where $\mathbf{S}$ is a diagonal matrix with the diagonal elements containing the uncertainty of emissions (standard deviations), and $\mathbf{C}$ is the correlation matrix, which is parameterized as a Kronecker product of a horizontal correlation matrix $\mathbf{C}_{\text{hor}}$ and a temporal correlation matrix $\mathbf{C}_t$ (as in TM5-4DVAR, Meirink et al., 2008):

$$\mathbf{C} = \mathbf{C}_{\text{hor}} \otimes \mathbf{C}_t. \quad (6)$$

The spatial covariance between two grid cells is parameterized using a Gaussian function:

$$c_{\text{hor}}(i_1, i_2, j_1, j_2) = \exp\left(-\frac{1}{2}\left(\frac{d(i_1, i_2, j_1, j_2)}{L_{\text{corr}}}\right)^2\right), \quad (7)$$

where $d(i_1, i_2, j_1, j_2)$ is the distance between two grid cells (with longitude indices i_1 and i_2 and latitude indices j_1 and

j_2) and L_{corr} a predefined correlation length constant. The temporal correlation uses an exponential decay function:

$$c_t(t_1, t_2) = \exp\left(-\frac{d_t(t_1, t_2)}{t_{\text{corr}}}\right), \quad (8)$$

where $d_t(t_1, t_2)$ is the temporal distance between two emission time steps and t_{corr} a predefined temporal correlation scale constant.

The observation error covariance matrix $\mathbf{R}$ considers only diagonal elements (i.e. ignores any error correlation between different observations) and takes into account the uncertainties of the measurements and the model representation error:

$$\mathbf{R} = \mathbf{R}_{\text{obs}} + \mathbf{R}_{\text{mod}}. \quad (9)$$

We use two different approaches to parameterize the model representation error, which are described in more detail in Sect. 2.2.3.

The observation operator $\mathbf{H}(\mathbf{x})$ simulates the measurements as a function of the state vector (i.e. as a function of emission deviation factors). For a given measurement m in time interval t_m the simulation is computed as

$$H_m(\mathbf{x}, t_m) = \sum_{t \in T_m} \sum_{i,j} \sum_{k \in K} e_{i,j,t}(x_{i,j,t}) \cdot \frac{w_k}{\delta h_k} \cdot \frac{M_{\text{air}}}{M_{\text{CH}_4}} \cdot 10^9 \cdot \phi_{i,j,k}^m(t, t + \delta t) [\text{ppb}] \quad (10)$$

using the 3-dimensional FLEXPART-COSMO footprints $\phi_{i,j,k}^m(t, t + \delta t)$ (units: $[1/(\text{kg air m}^{-3} \text{s}^{-1})]$), described in Sect. 2.1. $e_{i,j,t}(x_{i,j,t})$ denotes the CH₄ emissions (computed according Eq. 2 or Eq. 3) in units of kilograms of methane per square metre per second ($\text{kg CH}_4 \text{ m}^{-2} \text{s}^{-1}$). M_{air} and M_{CH_4} are the molecular masses of air and CH₄, respectively. δh_k is the layer thickness of vertical layer k (here we use the two lowermost layers, each with thickness of 50 m), and w_k is the weighting of layer k (here 0.5 for the applied two layers). T_m represents the time interval of the applied footprints (i.e. 10 d prior to the measurements) and δt the averaging time (3 h) for the single footprints (computed for the time interval between t and $t + \delta t$). $H_m(\mathbf{x}, t_m)$ represents the simulated enhancement of the CH₄ mole fraction above the baseline (which is evaluated using two different approaches as described in Sect. 2.2.2).

The minimization of the cost function Eq. (1) requires the evaluation of the gradient of the cost function with respect to the state vector:

$$\nabla_{\mathbf{x}} J(\mathbf{x}) = \mathbf{B}^{-1}(\mathbf{x} - \mathbf{x}_b) + \mathbf{H}^T \mathbf{R}^{-1}(\mathbf{H}(\mathbf{x}) - \mathbf{y}), \quad (11)$$

where $\mathbf{H}^T$ is the adjoint model operator, describing the sensitivity of the simulated observations with respect to changes of the state vector. $\mathbf{H}^T$ can be directly computed using the FLEXPART-COSMO footprints $\phi_{i,j,k}^m(t, t + \delta t)$.

In order to achieve better convergence of the minimization algorithm, pre-conditioning is applied, transforming the state

vector $\mathbf{x}$ to the control vector $\mathbf{w}$ (similarly as, for example, in TM5-4DVAR, Meirink et al., 2008),

$$\mathbf{w} = \mathbf{B}^{-1/2}(\mathbf{x} - \mathbf{x}_b), \quad (12)$$

and in reverse direction,

$$\mathbf{x} = \mathbf{x}_b + \mathbf{B}^{1/2} \mathbf{w}. \quad (13)$$

The square root of the submatrix $\mathbf{C}_t$ is calculated by eigenvalue decomposition. However, this is not possible for the submatrix $\mathbf{C}_{\text{hor}}$ (in contrast to TM5-4DVAR) due to the large size of this matrix of

$$n_{\text{C}_{\text{hor}}} = (n_{\text{longitude}} \times n_{\text{latitude}})^2 = 132\,834^2. \quad (14)$$

Therefore, an Arnoldi factorization (Arnoldi, 1951) is used, computing only the largest eigenvalues/eigenvectors. Consequently, the square root matrix $\mathbf{B}^{1/2}$ and its inverse in Eqs. (13) and (12) are replaced by corresponding factorizations. As default setting, we use a fraction of 1 % of the eigenpairs, which is reducing the size of the control vector to 1 % of the size of the state vector. Test inversions with higher fractions of eigenpairs (up to 3 %) showed that a fraction of 1 % is generally sufficient for the range of correlation lengths ($L_{\text{corr}} = 50 \text{ km}$ – 200 km) used in this study.

For the regular inversions, we use the limited memory quasi-Newton algorithm mlqn3 developed by Gilbert and Lemaréchal (1989), which also allows for the optimization of non-linear problems (as the semi-lognormal pdf (Eq. 2) introduces non-linearity to our optimization problem). In order to evaluate the posterior uncertainties, additional inversions are performed using a conjugate gradient algorithm (Fisher and Courtier, 1995; Lanczos, 1950; Meirink et al., 2008) for minimization and the linear expansion of the emission deviation factors (Eq. 3). The posterior covariance is then computed from the leading eigenvalues of the Hessian of the cost function:

$$\mathbf{B}_{\text{apos}} = \mathbf{B} + \sum_{k=1}^K \left(\frac{1}{\theta_k} - 1 \right) (\mathbf{B}^{1/2} \mathbf{v}_k) (\mathbf{B}^{1/2} \mathbf{v}_k)^T, \quad (15)$$

with $\mathbf{v}_k$ and θ_k the eigenvectors and eigenvalues of the Hessian of the cost function. The inversions based on the conjugate gradient algorithm yield in general rather similar posterior emissions as the regular inversions using the mlqn3 algorithm. Derived aggregated annual total CH₄ emissions agree on average (average over all sensitivity inversions) within -1.0% to 2.1% for the “country regions” discussed in this paper (Germany, France, BENELUX, UK + Ireland; Fig. 2) and within -6.4% to 6.5% for individual inversions.

2.2.2 Baselines

Since FLEXPART-COSMO is a limited domain model, providing the source–receptor relationship due to emissions

in the European COSMO-7 domain, it requires the provision of background mole fractions (baselines), representing the CH₄ mole fractions of the air masses entering the COSMO-7 domain. Here, we simulate the baselines using the global TM5-4DVAR inverse modelling system (described in Sect. 2.4), using two different approaches to couple the regional FLEXPART-COSMO inversion with TM5-4DVAR. The first approach applies the method described by Rödenbeck et al. (2009) (denoted in the following as “Rödenbeck baselines”), which is a flexible nesting scheme allowing for the offline coupling of regional models with global inversions. For this purpose, we use a manipulated version of TM5 which simulates only the “cis” part of the CH₄ fields (representing the CH₄ enhancement due to emissions in the area to the FLEXPART-COSMO model transported directly to the European measurement stations, without leaving the COSMO-7 domain), denoted $\Delta c_{\text{mod } 1, \text{ cis}}$. The baseline mole fractions, c_{baseline} , are then computed as difference of the TM5-4DVAR posterior simulation ($c_{\text{mod } 1}$) and the “cis” part (Rödenbeck et al., 2009), both sampled at the location of the corresponding station:

$$c_{\text{baseline}} = c_{\text{mod } 1} - \Delta c_{\text{mod } 1, \text{ cis}}. \quad (16)$$

The second approach to couple FLEXPART-COSMO with TM5-4DVAR uses the particle positions of the FLEXPART-COSMO back trajectories at the time of termination, i.e. either at the applied maximum termination time (set to 10 d in our study) or when they leave the COSMO-7 domain (which can be well before the 10 d). For each individual particle position, the CH₄ mole fraction is then extracted from the 3-dimensional TM5-4DVAR CH₄ fields. Since each 3 h average FLEXPART-COSMO footprint is based on the release of 50 000 particles, the corresponding (3 h average) baseline mole fraction is computed as the average of the CH₄ mole fraction at the termination points of all individual 50 000 particles. In the following, this approach is denoted “particle position baselines”.

2.2.3 Model representation error

We applied two different approaches to estimate the model representation error. The first approach, denoted “OBS”, is similar to the method described by Henne et al. (2016), evaluating the residuals (difference between observations and model simulations) as a function of CH₄ enhancement. However, some details of our method are different from Henne et al. (2016). Here, we use the following function to parameterize the model representation error, $\Delta y_{\text{MOD } k, i}$, as a function of the absolute observed CH₄ enhancement (i.e. observed CH₄ mole fraction minus CH₄ background), $|y_{k, i}|$:

$$\Delta y_{\text{MOD } k, i} = f_{0, k} + \rho_k \left(|y_{k, i}| + \frac{\exp(-a_k |y_{k, i}|) - 1}{a_k} \right), \quad (17)$$

where k is the station index, i the index of the individual observational data point of the time series of station k , and

$f_{0, k}$, ρ_k , and a_k the three fit parameters evaluated for each station. These fit parameters are determined by calculating the fit curve of the absolute residuals, $|y_{k, i} - (\mathbf{H}(\mathbf{x}))_{k, i}|$ as a function of $|y_{k, i}|$. For large $y_{k, i}$, the curve becomes linear with slope ρ_k ; the minimum value of the fit function is $f_{0, k}$, and a_k defines how fast the function becomes linear. The evaluation of the fit function is generally performed iteratively. In a first step, the fit parameters are computed using the prior model simulations, while in subsequent iterations the posterior simulations of the previous iteration are used. FLEXPART includes an outer loop system, which allows for an arbitrary number of iterations. However, tests have shown that changes after the second iteration are usually very small. Therefore, for the inversions presented in this paper, we used generally only two iterations. Figure S1 in the Supplement (left column) illustrates the fit functions for some selected stations.

As alternative to the model representation error OBS described above, a second approach, denoted “METEO”, has been developed, parameterizing the model representation error as a function of wind speed:

$$\Delta y_{\text{MOD } k, i} = f_{0, k} + \rho_k \exp(-a_k w_{k, i}), \quad (18)$$

where $w_{k, i}$ is the wind speed (in m s^{-1}) extracted from the COSMO model corresponding to the data point $y_{k, i}$, and $f_{0, k}$, ρ_k , and a_k are the three fit parameters for station k . The rationale is that if wind speed at the measurement location is low, the observed mole fraction might be more influenced by local sources of emissions and therefore less representative of the modelled mole fractions. Again, the fit parameters are determined to get an optimal fit through the absolute residuals, $|y_{k, i} - (\mathbf{H}(\mathbf{x}))_{k, i}|$ and they are evaluated in two iterations. For large wind speeds, $\Delta y_{\text{MOD } k, i}$ converges towards $f_{0, k}$, while $f_{0, k} + \rho_k$ represents the model representation error at wind speed zero. The right column of Fig. S1 shows the fit functions for the METEO model representation error for some selected stations.

2.3 FLEXPART extended Kalman filter inverse modelling system

FLEXPART is an inverse modelling system based on an extended Kalman filter as described in detail in Brunner et al. (2012, 2017). Observations are assimilated sequentially (here day by day) to provide the best linear unbiased estimate of the emissions and their uncertainties based on measurements up to the present time of the assimilation. Rather than estimating monthly or annual mean emissions, the system adjusts the emissions continuously as the assimilation proceeds and in this way creates a smoothly varying emission field that later can be averaged to monthly or annual means. The filter includes a forecast step, which predicts the evolution of the emissions from one assimilation time step to the next. The simplest assumption is persistence (i.e. no change with time), but to incorporate seasonally varying prior emissions,

a non-zero forecast update was implemented that follows the linear change in prior emissions from 1 month to the next. Since the forecast step is associated with an uncertainty, the posterior uncertainty can become larger than the prior uncertainty in regions that are poorly constrained by observations. In order to avoid an unrealistic growth of the uncertainties in these regions, the posterior uncertainties are reset to the prior uncertainty whenever they become larger.

The state vector consists of two components: (i) emissions on a reduced grid with a total of 3608 elements for inversions with observation data set O1 and 6497 elements for inversions with observation data set O2 (Sect. 3.1 and Table 1) and (ii) coefficients of an AR(1) autoregressive process describing temporal correlations in the residuals at each individual site. As described in Brunner et al. (2012), the reduced grid has high spatial resolution near the measurement sites and lower resolution further away, reflecting the reduced contribution of emissions at larger distance to observed CH₄ variations. It was constructed based on the combined total annual footprint of all measurement sites. The state vector may contain the emissions directly or the logarithm of the emissions. The latter option was chosen here to enforce a positive solution, i.e. positive methane emissions in each grid cell.

Another option in FLE_xKF is to optimize baseline mole fractions at each observation site in addition to the gridded emission field. However, in the configuration used here, the baselines (Rödenbeck baselines) described in Sect. 2.2.2 were used directly without optimization. The results of FLE_xKF should be readily comparable to those of FLEXVAR, since the same FLEXPART-COSMO back trajectories, baselines, and observations were used. Spatial correlations in the prior emission uncertainties were represented in the prior error covariance matrix with a correlation length scale of 200 km (exponential decay as in Eq. 8). The matrix was scaled such that the prior uncertainty of the total domain emissions was 20 % (1σ).

2.4 TM5-4DVAR inverse modelling system

TM5-4DVAR is a global inverse modelling system based on the 4DVAR data assimilation technique and has been described in detail by Meirink et al. (2008), while subsequent updates have been reported in Bergamaschi et al. (2018a, 2010). TM5-4DVAR uses the Eulerian atmospheric chemistry transport model TM5 (Krol et al., 2005), a two-way nested zoom model, which allows the system to zoom in over specific regions of interest. Here, we apply the $1^\circ \times 1^\circ$ zooming over the European domain -18° to 42° (longitude) $\times 32^\circ$ to 64° (latitude) and an intermediate 3° (longitude) $\times 2^\circ$ (latitude) zooming over the extended European domain -30° to 48° (longitude) $\times 26^\circ$ to 70° (latitude), while the remaining global domain is simulated at a horizontal resolution of 6° (longitude) $\times 4^\circ$ (latitude). TM5 is an offline transport model, driven by external meteorological data, preprocessed to provide consistent meteorological fields for the different TM5

resolutions (Krol et al., 2005). In this study, 3 h interpolated meteorological fields from the European Centre for Medium-Range Weather Forecasts (ECMWF) ERA-Interim reanalysis (Dee et al., 2011) have been applied, using 25 vertical layers (defined as a subset of the 60 layers of the ERA-Interim reanalysis). About 5 layers represent the boundary layer, 10 layers the free troposphere, and 10 layers the stratosphere (Krol et al., 2005).

As for the regular FLEXVAR inversions, a semi-lognormal pdf has been used (Eq. 2). Minimization of the cost function Eq. (1) is performed using the mlqn3 algorithm (Gilbert and Lemaréchal, 1989) and the adjoint of the tangent linear TM5 model (Krol et al., 2008; Meirink et al., 2008) for evaluation of the gradient of the cost function Eq. (11). Four groups of CH₄ emissions are optimized independently: (1) wetlands, (2) rice, (3) biomass burning, and (4) all remaining sources (Bergamaschi et al., 2018a). Uncertainties of 100 % per grid cell and month were applied for each source group with a spatial correlation length scale of 200 km (Eq. 7). The temporal correlation timescales (Eq. 8) are set to 12 months for the “remaining” CH₄ sources (which are assumed to have no or only small seasonal variations) and to zero for the other source groups (wetlands, rice, and biomass burning), which have pronounced seasonal cycles. The model representation error is parameterized as a function of local emissions (i.e. emissions of the grid cell in which the corresponding monitoring station is located) and 3-dimensional gradients of simulated mole fractions (Bergamaschi et al., 2010). The photochemical sinks of CH₄ in the troposphere (OH) and stratosphere (OH, Cl, and O(¹D)) are simulated as described in Bergamaschi et al. (2010). For the coupling of the FLEXPART-COSMO inversions over the European COSMO domain with the global TM5-4DVAR inversions, the baselines at the monitoring stations (within the COSMO-7 domain) have been computed using the two different approaches described in Sect. 2.2.2.

3 Model input data and inversions

3.1 Atmospheric observations

The atmospheric observations used in this study (within the COSMO-7 domain) include ground-based CH₄ data for 2018 from 24 stations with in situ measurements, complemented with data from 5 stations with discrete air sampling, as compiled in Table 1. Most of the in situ measurements (15 stations) are from the atmosphere network of the Integrated Carbon Observation System (ICOS) (Heiskanen et al., 2022), a pan-European infrastructure providing harmonized atmospheric measurements which are rigorously standardized in terms of instrumentation, calibration, air sampling, and quality control, including centralized data processing and data evaluation at the ICOS Atmospheric Thematic Centre (<https://icos-atc.lsce.ipsl.fr/>, last access: 5 October 2022) (Hazan et al., 2016; ICOS RI, 2020; Yver-Kwok

et al., 2021). Here, we use the ICOS Atmosphere Release of final, quality-controlled data 2021-1 (ICOS RI, 2021). For station Lütjehad (LUT), the ICOS data only start on 13 August 2018. Data before that date were provided by the University of Groningen, with data processing very similar to the ICOS data. Data from the ICOS station Ispra (IPR) have been further processed using the robust extraction of baseline signal (REBS) spike detection algorithm (El Yazidi et al., 2018; Ruckstuhl et al., 2012) in order to filter out data affected by nearby farming activities. In addition to the ICOS measurements, further in situ measurements have been used from the UK Deriving Emissions linked to Climate Change (DECC) network (Bilsdale (BSD), Tacolneston (TAC), Ridge Hill (RGL), Heathfield (HFD)), from the Advanced Global Atmospheric Gases Experiment (AGAGE) (Mace Head (MHD)), from the University of East Anglia (Weybourne (WAO)), from the Netherlands Organisation for Applied Scientific Research (TNO) (Cabauw (CBW)), from Empa (Lägeren Hochwacht (LHW)), and from the University of Bern (Beromünster (BRM)). The in situ measurements are complemented by discrete air samples (which are usually collected weekly) from the NOAA Earth System Research Laboratory (ESRL) global cooperative air sampling network (Dlugokencky et al., 2021), with five stations within the COSMO-7 domain. Additional atmospheric data used for the TM5-4DVAR inversions are compiled in Table S1 and include six further ICOS stations with in situ measurements and 31 NOAA discrete air sampling sites located outside the COSMO-7 domain. The selected stations outside the European $1^{\circ} \times 1^{\circ}$ or $3^{\circ} \times 2^{\circ}$ TM5 zoomed-in regions are mostly global background stations in remote areas which can be reasonably well reproduced with the coarse global TM5 resolution of $6^{\circ} \times 4^{\circ}$.

The atmospheric CH₄ data are reported on the WMO X2004A calibration scale (Dlugokencky et al., 2005; NOAA, 2021), except the AGAGE MHD data which are reported on the Tohoku University (TU) CH₄ standard scale (Aoki et al., 1992; Prinn et al., 2000). Comparison of parallel measurements by NOAA and AGAGE at five global sites over more than 25 years showed that the two calibration scales are in close agreement, with an average ratio of 1.0002 ± 0.0007 . Therefore, no scale correction has been applied.

For the in situ measurements (which are available quasi-continuously in time), we assimilate only early afternoon data for stations in the boundary layer and night-time data for mountain stations (Bergamaschi et al., 2015), selecting the 3 h time interval of the FLEXPART back trajectories (which are provided for [00:00–03:00, 03:00–06:00, ...] UTC), which is closest to the time interval [12:00–15:00] LT for the stations in the boundary layer and [00:00–03:00] LT for the mountain stations (indicated in Table 1 by column “M”), respectively. This procedure ensures consistent averaging of the FLEXPART back trajectories and the assimilated observations over the same 3 h time intervals. Discrete air samples were taken as available, i.e. without any

temporal selection. The measurement uncertainty is set to 3 ppb for all observations (for observational part $\mathbf{R}_{\text{obs}}$ of the observation error covariance matrix Eq. 9).

In this study, we investigate two observation data sets (Table 1). The first data set, denoted O1, is considered the observational base data set and uses only the ICOS and NOAA data, while the second data set, O2, also includes all additional in situ measurements. The largest difference between the two data sets is the much better observational coverage of the British Isles in O2 with six in situ measurement stations located in that area, compared to only one station with discrete air sampling (MHD/NOAA) in O1.

3.2 Emission inventories

Three different emission inventories are used alternatively as prior estimates of the major anthropogenic CH₄ emissions (Table 2). The first inventory is the Emissions Database for Global Atmospheric Research (EDGAR) v6.0 (EDGAR v6.0, 2021), which provides monthly sector-specific global grid maps of emissions at a horizontal resolution of $0.1^{\circ} \times 0.1^{\circ}$ for 2000–2018. The second inventory, TNO-VERIFYv3.0, is the third version of the TNO greenhouse gas and co-emitted species (GHGco) emission database, developed by TNO within the VERIFY project. TNO-VERIFYv3.0 provides annual European CH₄ emissions at a horizontal resolution of $\sim 6\text{ km} \times 6\text{ km}$ for the years 2005–2018 but includes monthly emission profiles. The third emission inventory has been provided by GCP-CH₄ (Saunio et al., 2020) globally at a horizontal resolution of $1^{\circ} \times 1^{\circ}$ for 2000–2017. In the absence of emission data for 2018 in the GCP-CH₄ inventory, we use here the 2017 data of this inventory. The resulting error of this 1-year inconsistency, however, is considered to be much smaller compared to the overall uncertainties of the emission inventories.

Natural CH₄ emissions were used from the GCP-CH₄ data set (Saunio et al., 2020), providing estimates of the climatological mean emissions of the major natural source categories. Furthermore, CH₄ emissions from biomass burning were taken from the Global Fire Emissions Database (GFED) version 4.1 (Van Der Werf et al., 2017). However, these were included only when using the EDGAR v6.0 or TNO-VERIFYv3.0 inventories, while the GCP-CH₄ (anthropogenic) data set already includes emissions from biomass burning.

Using the above emission inventories, we have assembled the emission data sets E1, E2, and E3 as compiled in Table 2 and used as prior estimates for the different inversions described in Sect. 3.4. All emission data sets have been mapped on the COSMO-7 grid, using the Python package “emiproc” (Jähn et al., 2020), which has been integrated into the FLEXPART inverse modelling system.

Table 1. European monitoring stations used in this study. “Alt” is the surface altitude (m above sea level), “SH” is the sampling height (m) above ground, “ST” specifies the sampling type (“T”: in situ measurements; “D”: discrete air sample measurements). “FLEX” is the release height (m) used for calculation of the FLEXPART-COSMO back trajectories (“a.g.l.”: release height above model surface; “a.s.l.”: release height above sea level). The column “M” indicates the stations which have been classified as mountain stations. The last two columns indicate the use of the corresponding station data in the observation data sets O1 and O2.

ID	Station name	Data provider	Lat.	Long.	Alt.	SH	ST	FLEX	M	O1	O2
HTM	Hyltemossa	ICOS	56.10	13.42	115	150	I	150 a.g.l.		•	•
BSD	Bilsdale	DECC	54.36	−1.15	380	248	I	248 a.g.l.			•
LUT	Lutjewad	RUG/ICOS ¹	53.40	6.35	1	60	I	60 a.g.l.		•	•
MHD	Mace Head	NOAA	53.33	−9.90	5	21	D	10 a.g.l.		•	•
		AGAGE	53.33	−9.90	5	8	I	10 a.g.l.			•
GAT	Gartow	ICOS	53.07	11.44	70	341	I	341 a.g.l.		•	•
WAO	Weybourne Atmospheric Observatory	UEA	52.95	1.12	15	10	I	10 a.g.l.			•
TAC	Tacolneston	DECC	52.52	1.14	56	100	I	100 a.g.l.			•
LIN	Lindenberg	ICOS	52.17	14.12	73	98	I	98 a.g.l.		•	•
RGL	Ridge Hill	DECC	52.00	−2.54	204	90	I	90 a.g.l.			•
CBW	Cabauw	TNO	51.97	4.93	0	200	I	200 a.g.l.			•
TOH	Torfhaus	ICOS	51.81	10.53	801	147	I	147 a.g.l.		•	•
HFD	Heathfield	DECC	50.98	0.23	150	100	I	100 a.g.l.			•
OXK	Ochsenkopf	NOAA	50.03	11.82	1022	163	D	1035 a.s.l.		•	•
KRE	Křešín u Pacova	ICOS	49.57	15.08	534	250	I	250 a.g.l.		•	•
KIT	Karlsruhe	ICOS	49.09	8.42	110	200	I	200 a.g.l.		•	•
SAC	Saclay	ICOS	48.72	2.14	160	100	I	100 a.g.l.		•	•
OPE	Observatoire Perenne ³	ICOS	48.56	5.50	390	120	I	120 a.g.l.		•	•
TRN	Trainou	ICOS	47.96	2.11	131	180	I	180 a.g.l.		•	•
HPB	Hohenpeissenberg	ICOS	47.80	11.02	934	131	I	938 a.s.l.		•	•
		NOAA	47.80	11.01	985	5	D	840 a.s.l.		•	•
LHW	Lägern Hochwacht	EMPA	47.48	8.40	840	32	I	699 a.s.l.			•
BRM	Beromünster	UBE	47.19	8.18	797	212	I	212 a.g.l.			•
HUN	Hegyhátsál	NOAA	46.95	16.65	248	96	D	96 a.g.l.		•	•
JFJ	Jungfraujoch	ICOS	46.55	7.98	3580	5	I	3137 a.s.l.	•	•	•
IPR	Ispira	ICOS/JRC ²	45.81	8.64	210	100	I	100 a.g.l.		•	•
PUY	Puy de Dôme	ICOS	45.77	2.97	1465	10	I	1168 a.s.l.	•	•	•
CMN	Monte Cimone	ICOS	44.19	10.70	2165	8	I	1753 a.s.l.	•	•	•
CIB	CIBA ⁴	NOAA	41.81	−4.93	845	5	D	5 a.g.l.		•	•

¹ Data since 13 August 2018 from ICOS data release and before that date from the University of Groningen. ² Data filtered with REBS spike detection algorithm (see Sect. 3.1).

³ Observatoire Perenne de l’Environnement. ⁴ Centro de Investigación de la Baja Atmósfera.

3.3 Post-processing of gridded emission data

In order to extract from gridded emission data (on COSMO-7 grid) total emissions of countries (or group of countries), country masks have been generated using the “Natural Earth dataset” (<https://www.naturalearthdata.com/>, last access: 30 July 2021), attributing each 7 km × 7 km COSMO-7 grid cell to a certain country (or sea). Offshore emissions over the sea are not included in the country totals.

Since the COSMO-7 domain does not cover the upper northern part of the UK, a correction factor of 1.057 is applied to estimate the total emissions of the country region “UK + Ireland”; i.e. the UK + Ireland emissions extracted from the corresponding grid cells within the COSMO-7 domain are multiplied by this factor (for further details, see Sect. S1 in the Supplement). Furthermore, small correction factors are applied when extracting country total emissions

from the gridded emissions of data set E3 (at horizontal resolution of 1° × 1°), since sampling of coastal 1° × 1° grid cells with the corresponding 7 km × 7 km COSMO-7 grid cells leads to a loss of emissions attributed to the countries, if the emissions of the coastal 1° × 1° grid cell originate mainly from land (for further details, see Sect. S1).

3.4 Sensitivity inversions

Table 3 compiles the different FLEXVAR inversions presented in this paper. INV-E1-O1 represents the base inversion, using the emission data set E1 as prior estimates, the observation data set O1, the METEO model representation error (Sect. 2.2.3), the Rödenbeck baselines (Sect. 2.2.2), and our default settings for the prior error covariance. A first set of sensitivity inversions investigates the impact of using alternatively the particle position baselines and the alternative

Table 2. Emission inventories used in this study. The second column (“Total”) lists the total CH₄ emissions over the COSMO-7 domain in units of teragrams of methane per year (Tg CH₄ yr^{−1}) for the individual categories and the totals of each inventory. The last three columns indicate the use of the corresponding inventory data in the emission data sets E1, E2, and E3. Data for EDGARv6.0, TNO-VERIFYv3.0, and GFEDv4.1 are for 2018, and GCP-CH₄ (anthropogenic) data are for 2017, while GCP-CH₄ (natural) data represent climatological mean values.

Inventory/category	Total	E1	E2	E3
EDGARv6.0				
Total	17.94			
Coal	0.74	•		
Oil	0.24	•		
Gas	2.20	•		
Enteric fermentation	6.84	•		
Manure management	2.18	•		
Rice agriculture	0.08	•		
Solid waste (landfills and incineration)	3.22	•		
Wastewater treatment	1.41	•		
Energy for buildings	0.71	•		
Further minor anthropogenic sources	0.33	•		
TNO-VERIFYv3.0				
Total	15.67			
Fugitive emissions	1.86		•	
Waste	4.29		•	
Agriculture: livestock	8.37		•	
Agriculture: other	0.15		•	
Other stationary combustion	0.57		•	
Further minor anthropogenic sources	0.44		•	
GCP-CH ₄ (anthropogenic)				
Total	21.00			
Coal	0.98			•
Oil gas industry	2.78			•
Livestock	9.39			•
Agriculture: rice	0.10			•
Waste	7.12			•
Biofuels and biomass burning	0.63			•
GCP-CH ₄ (natural)				
Total	2.15			
Wetlands	1.76	•	•	•
Geological	0.48	•	•	•
Termites	0.10	•	•	•
Ocean	0.55	•	•	•
Soil sink	−0.75	•	•	•
GFEDv4.1				
Biomass burning	0.02	•	•	

parameterization OBS of the model representation error (and the combination of both). In a further inversion series, we analyse the sensitivity of the inversions to the main settings of the prior error covariance matrix, i.e. for the spatial correlation length constant, L_{corr} , the temporal correlation scale

constant, t_{corr} , and the assumed uncertainty of emissions per grid cell and month. Furthermore, we examine the sensitivity of the inversions to the use of the alternative emission inventories E2 and E3 as prior estimates instead of E1 and the use of the extended observational data set O2 instead of O1.

In addition to the FLEXVAR inversions compiled in Table 3, inversions with the FLEExKF system (described in Sect. 2.3) and with TM5-4DVAR (described in Sect. 2.4) have been performed for comparison with FLEXVAR (and will be discussed in Sect. 4.3). These inversions have been made for both observational data sets, O1 and O2, using the emission inventory E3 as prior estimates. Furthermore, additional FLEExKF inversions have been performed using alternatively E1 as prior estimates.

4 Results and discussion

4.1 Sensitivity of FLEXVAR inversions to internal parameterizations and model settings

4.1.1 Sensitivity of FLEXVAR inversions to baselines

Figure 1 shows maps of European CH₄ emissions derived for the base inversion INV-E1-O1 and the sensitivity inversion INV-E1-O1-S1, in which the particle position baselines were used instead of the Rödenbeck baselines. Both inversions display in general similar spatial patterns of the inversion increments (i.e. difference between posterior and prior emissions); however in most regions INV-E1-O1-S1 shows somewhat lower CH₄ emissions than INV-E1-O1, visible in the slightly larger areas with negative inversion increments and slightly smaller areas with positive inversion increments. Consequently, also the derived country total emissions (shown in Fig. 2 and compiled in Table S4) are lower in INV-E1-O1-S1, e.g. −6.6% lower over Germany and −12.8% lower over France compared to INV-E1-O1.

Figure 3 illustrates the two different baselines at some example stations during the 3-month period from 1 April until 1 July 2018. In general, both baselines are rather similar, including their synoptic variability. However, there are certain periods during which the particle position baselines are somewhat higher than the Rödenbeck baselines, for example, at KIT, SAC, and OPE during the period between day 140 and day 162. Consequently, the observational forcing (i.e. the enhancement of the observations above baseline) is lower during such periods for the particle position baselines, resulting in lower derived emissions. One major difference between both approaches is that in the case of the Rödenbeck baselines, the background mole fractions are transported to the stations by TM5, while in the case of the particle position baselines, they are transported by FLEXPART. In order to further investigate which baselines are more realistic, we have compared model simulations and observations for “background conditions”, defined as events when the contribution of European emissions (evaluated by Eq. 10) is lower

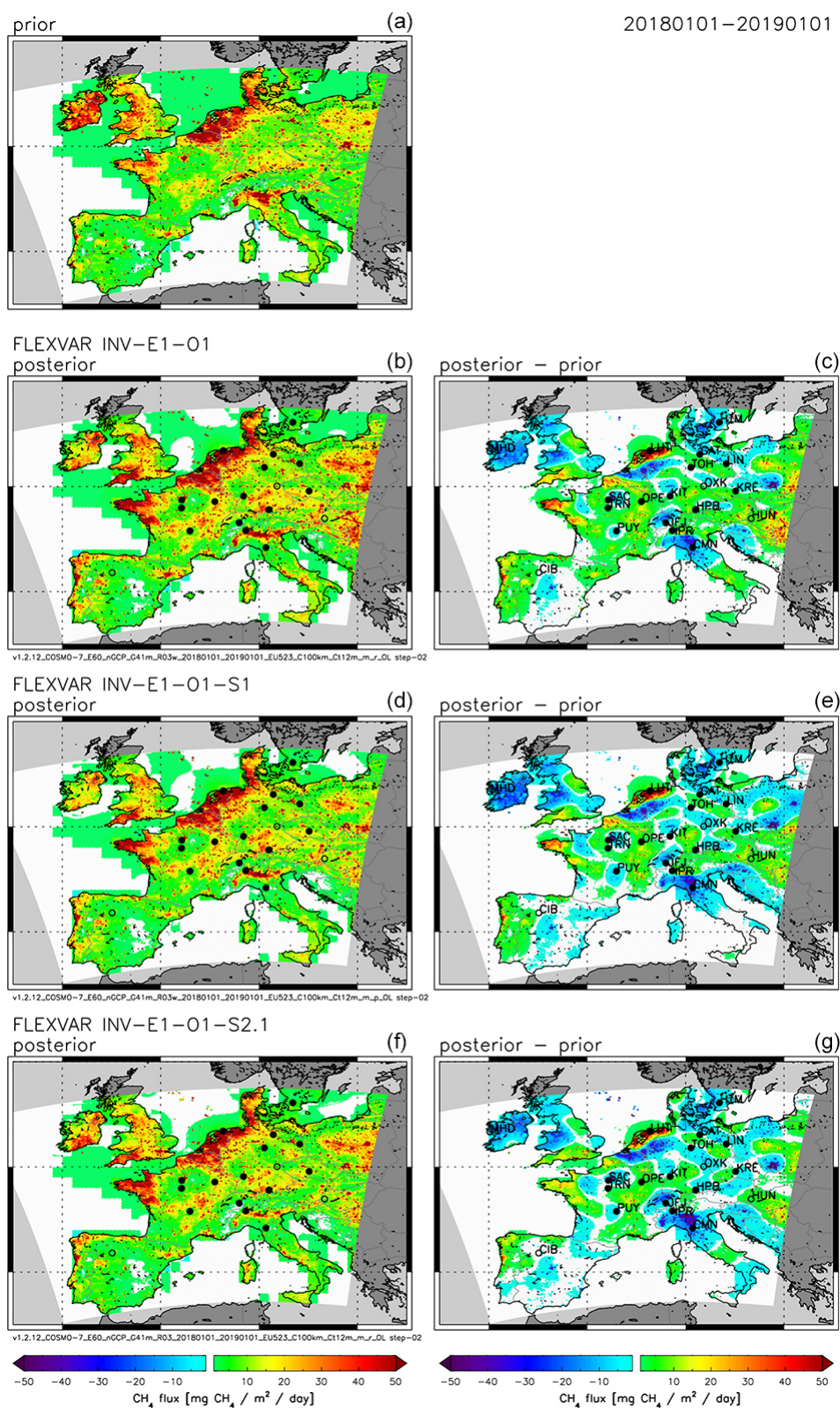


Figure 1. Sensitivity of FLEXVAR inversions to different approaches to calculate the baselines and to parameterize the model representation error. **(a)** Prior emissions (emission data set E1). **(b, c)** Posterior emissions (left) and difference between posterior and prior emissions (right) for base inversion INV-E1-O1 (using the Rödénbeck baselines and the METEO model representation error). **(d, e)** Inversion INV-E1-O1-S1 using the particle position baselines. **(f, g)** Inversion INV-E1-O1-S2.1 using the OBS model representation error. All figures show annual average CH₄ emissions (or change in emissions, respectively) for 2018.

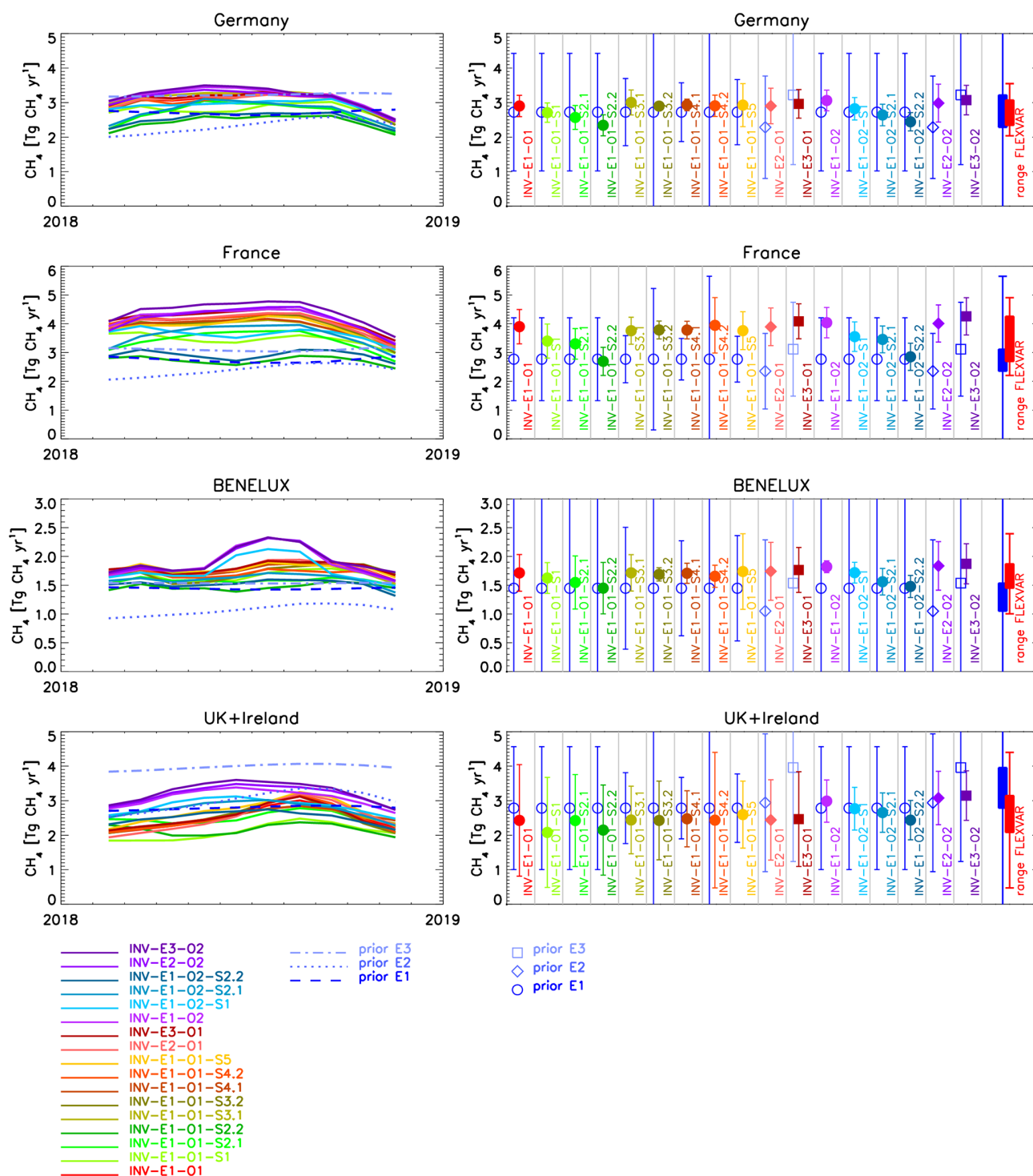


Figure 2. Total CH₄ emissions for Germany, France, BENELUX, and UK + Ireland derived for different sensitivity inversions (Table 3). Left: 3-month running mean total CH₄ emissions of the corresponding country regions. Right: Annual total CH₄ emissions. Open symbols show prior emissions, solid symbols show posterior emissions, and the error bars are the 2 σ uncertainties of prior and posterior emissions, respectively. The solid blue and red rectangles on the right side of the figures show the prior and posterior range from all individual inversions and the error bars on these rectangles the minimum and maximum values of the 2 σ uncertainties of the individual inversions.

Table 3. FLEXVAR sensitivity inversions. The column “Prior” lists the emission data set used (Table 2) and column “Obs” the observation data set (Table 1). “MRE” is the applied model representation error (Sect. 2.2.3), and “Baseline” lists the applied approach to calculate the baselines (Sect. 2.2.2). “ L_{corr} ” is the applied spatial correlation length constant, “ t_{corr} ” the temporal correlation scale constant, and “Unc” the assumed 1σ uncertainty of total emissions per grid cell and month (Sect. 2.2.1).

Inversion	Prior	Obs	MRE	Baseline	L_{corr}	t_{corr}	Unc
INV-E1-O1	E1	O1	METEO	Rödenbeck	100 km	12 m	100 %
Sensitivity to baselines							
INV-E1-O1-S1	E1	O1	METEO	particle position	100 km	12 m	100 %
Sensitivity to model representation error							
INV-E1-O1-S2.1	E1	O1	OBS	Rödenbeck	100 km	12 m	100 %
INV-E1-O1-S2.2	E1	O1	OBS	particle position	100 km	12 m	100 %
Sensitivity to spatial correlation length							
INV-E1-O1-S3.1	E1	O1	METEO	Rödenbeck	50 km	12 m	100 %
INV-E1-O1-S3.2	E1	O1	METEO	Rödenbeck	200 km	12 m	100 %
Sensitivity to prior uncertainty							
INV-E1-O1-S4.1	E1	O1	METEO	Rödenbeck	100 km	12 m	50 %
INV-E1-O1-S4.2	E1	O1	METEO	Rödenbeck	100 km	12 m	200 %
Sensitivity to temporal correlation length							
INV-E1-O1-S5	E1	O1	METEO	Rödenbeck	100 km	1 m	100 %
Sensitivity to prior inventories							
INV-E2-O1	E2	O1	METEO	Rödenbeck	100 km	12 m	100 %
INV-E3-O1	E3	O1	METEO	Rödenbeck	100 km	12 m	100 %
Sensitivity to observations							
INV-E1-O2	E1	O2	METEO	Rödenbeck	100 km	12 m	100 %
INV-E1-O2-S1	E1	O2	METEO	particle position	100 km	12 m	100 %
INV-E1-O2-S2.1	E1	O2	OBS	Rödenbeck	100 km	12 m	100 %
INV-E1-O2-S2.2	E1	O2	OBS	particle position	100 km	12 m	100 %
INV-E2-O2	E2	O2	METEO	Rödenbeck	100 km	12 m	100 %
INV-E3-O2	E3	O2	METEO	Rödenbeck	100 km	12 m	100 %

than a certain threshold (here set to 5 ppb). Figure S2 shows the comparison for eight stations, for which a sufficient number (> 20) of events with background conditions has been found. For the Rödenbeck baselines, six of these eight stations show posterior biases close to zero (< 2 ppb), while PUY shows a small negative bias (-5.3 ppb) and CMN a small positive bias (3.7 ppb). In contrast, the particle position baselines result in a smaller negative bias at PUY (-3.3 ppb) but larger positive biases at WAO (2.5 ppb), JFJ (5.5 ppb), and CMN (8.6 ppb). This analysis suggests that the performance of the Rödenbeck baselines is slightly better compared to the particle position baselines under background conditions. However, we note that differences of the baselines shown in Fig. 3 are mainly during periods of elevated CH₄ enhancements, for which it is more difficult to evaluate (based on the observations) which baselines are more realistic.

4.1.2 Sensitivity of FLEXVAR inversions to parameterization of model representation error

Figure 1 illustrates the sensitivity of the derived emissions to the applied parameterization of the model representation error. Inversion INV-E1-O1-S2.1, for which the OBS model representation error has been used, results in overall lower CH₄ emissions compared to the base inversion INV-E1-O1 with the METEO parameterization, again reflected in the larger extension of the areas with negative inversion increments and smaller extension (and magnitude) of the areas with positive inversion increments. Accordingly, the annual total emissions derived in INV-E1-O1-S2.1 are lower compared to INV-E1-O1 for all countries or group of countries (denoted in the following as country regions) shown in Fig. 2.

The OBS model representation error increases with increasing observed CH₄ enhancement (i.e. observed CH₄

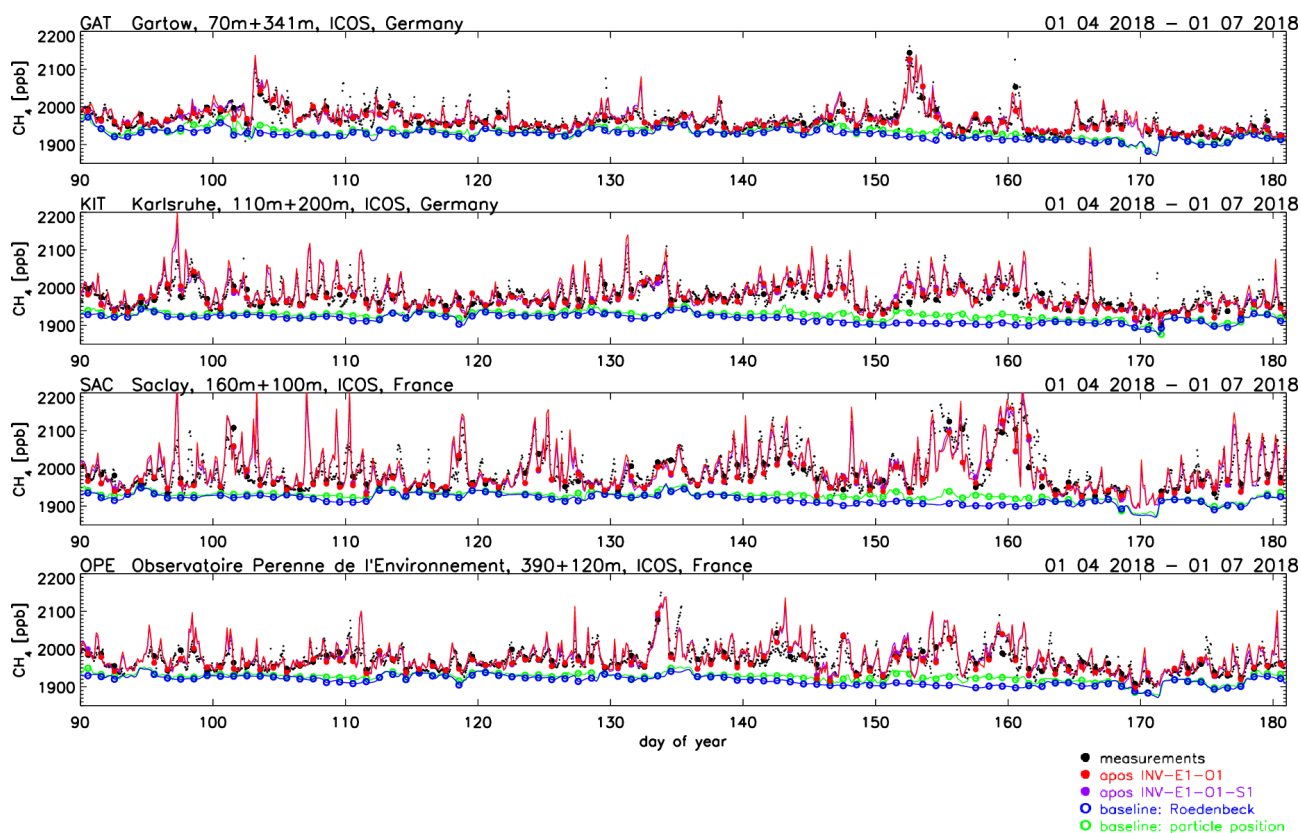


Figure 3. Time series of simulated and observed CH₄ mole fractions at stations GAT, KIT, SAC, and OPE for the 3-month period from 1 April until 1 July 2018. Blue curve shows the Rödenbeck baselines, green the particle position baselines, red the posterior CH₄ mole fractions for inversion INV-E1-O1 (using the Rödenbeck baselines), and violet the posterior CH₄ mole fractions for inversion INV-E1-O1-S1 (using the particle position baselines) (owing to the similarity of both posterior simulations, however, the results of INV-E1-O1-S1 are largely overlaid by those of INV-E1-O1). Small black dots: hourly averaged observations. Solid black circles: assimilated observations. Coloured symbols show the corresponding assimilated values (solid circles: assimilated posterior mole fractions; open circles: baseline values used for the assimilation).

mole fraction minus CH₄ background) (Sect. 2.2.3 and Fig. S1) and shows a large dynamic range at most stations, resulting in a generally relative low weighting in the inversion of events with larger CH₄ enhancements. In contrast, the dynamic range of the METEO model representation error is smaller at most stations, leading to a generally more equal weighting of all data points. Using the METEO model representation error, the observations can be better reproduced, achieving a higher average correlation coefficient ($r = 0.85$) and lower average root mean square difference ($\text{rms} = 30.0$ ppb) compared to the OBS model representation error ($r = 0.80$; $\text{rms} = 35.4$ ppb), as shown in Fig. S3. Apart from the better statistical performance, the METEO model representation is probably better at estimating the capability of the model to reproduce the observations (which largely depends on the specific meteorological situation), since wind speed might be a better indicator of the representativeness of a certain data point than the observed CH₄ enhancement, as the latter not only depends on the meteorological situation, but also on the regional CH₄ emissions.

Given the relatively large impact of the parameterization of the model representation error and the baselines, we have also performed an inversion combining the OBS model representation error and the particle position baselines (inversion INV-E1-O1-S2.2), which yields further reduced country total emissions compared to INV-E1-O1-S2.1 and INV-E1-O1-S1 (Fig. 2).

4.1.3 Sensitivity of FLEXVAR inversions to model covariance settings

In the following, the sensitivity of the FLEXVAR inversions to the main parameters of the prior covariance is investigated, i.e. horizontal correlation length constant, temporal correlation scale constant, and assumed uncertainties of emissions per grid cell and emission time step. Figure S4 shows inversions for horizontal correlation length constants L_{corr} (Eq. 7) of 50 km (INV-E1-O1-S3.1), 100 km (default value; INV-E1-O1), and 200 km (INV-E1-O1-S3.2). As expected, the spatial dimension of the inversion increments is increasing with

increasing L_{corr} . Despite these clearly visible differences in the spatial patterns of the inversion increments, the impact on the annual total emissions of the country regions shown in Fig. 2 is relatively small, since apparently the differences in the smaller-scale spatial patterns are largely averaged out over larger areas. Associated with the increase of the horizontal correlation length constant is a significant increase of the prior uncertainties of the annual total emissions per country, since increasing horizontal correlation length constant implies larger error correlations between neighbouring grid cells and hence increasing aggregated uncertainties (as uncertainties per grid cell and month were kept constant (at 100 %) in this sensitivity inversion series). Analogously, the decrease in the temporal correlation scale constant, t_{corr} (Eq. 8), results in a decrease of the aggregated annual prior uncertainty, as illustrated by inversion INV-E1-O1-S5, in which t_{corr} has been set to 1 month (instead of the default value of 12 months applied in all other inversions). Again, however, the effect on the derived annual emissions of the country regions remains very small (Fig. 2).

Figure S5 shows the dependence of the inversions on the assumed uncertainties of prior emissions per grid cell and month for values of 50 % (INV-E1-O1-S4.1), 100 % (default value; INV-E1-O1), and 200 % (INV-E1-O1-S4.2). The increase of the assumed prior uncertainty leads to a significant increase of the derived regional inversion increments. This effect is most pronounced at larger distances from the monitoring stations, where observational constraints are relatively weak. Especially the large inversion increments visible in INV-E1-O1-S4.2 at the eastern domain boundary are probably an artefact, since the inversion may generate such patterns in regions far from the observations to compensate for systematic errors, for example, in model transport and with little penalty in the cost function in the case of prior uncertainties that are assumed very high.

Despite the dependence of the smaller-scale regional inversion increments on the assumed prior uncertainties, the impact on the derived annual total emissions remains again very small for the country regions shown in Fig. 2, since their emissions are relatively well constrained by the available observations and since differences in the smaller-scale inversion increments are averaged out over larger areas.

4.2 Sensitivity of FLEXVAR inversions to model input data

4.2.1 Sensitivity of FLEXVAR inversions to prior emission inventories

Figure 4 shows maps of the European CH₄ emissions for INV-E1-O1, INV-E2-O1, and INV-E3-O1, which use the three different emission data sets E1, E2, and E3 (Sect. 3.2; Table 2) as prior emissions. While the major patterns of the spatial prior emission distribution look relatively similar for the three inventories (e.g. the high emissions over the

BENELUX countries and the Po valley), there are significant differences in the prior country region total emissions (Fig. 2; Table S4). E2 has lower emissions over Germany (16.1 %), France (15.1 %), and BENELUX (27.4 %) compared to E1 (and 11.3 % lower over the whole COSMO-7 domain; Table 2), while E3 has higher total emissions over the COSMO-7 domain (15.1 % higher than E1) and very high emissions in particular for UK + Ireland (42.1 % higher than E1). Despite these considerable differences in the prior emissions, the annual total posterior emissions of the country regions shown in Fig. 2 are very similar for the three inversions. This indicates that the inversions are largely driven by the observations. For UK + Ireland this is somewhat surprising, since only one measurement station (MHD/NOAA) is located in this country region in the applied observation data set O1, but apparently the continental stations provide some constraints for the emissions from UK + Ireland. We will see in the next section, however, that including additional stations has a significant impact on the CH₄ emissions derived for UK + Ireland.

4.2.2 Sensitivity of FLEXVAR inversions to assimilated observations

While the base observation data set O1 uses only the ICOS in situ stations, complemented by the NOAA discrete air sampling sites, nine further in situ stations from other networks/institutions are added in observation data set O2 (Table 1). Six of the additional stations are located on the British Isles, two in Switzerland, and one in the Netherlands. Figure 5 displays the inversions INV-E1-O1 and INV-E1-O2 using the two different observation data sets. As expected, the largest differences are visible in the regions around the additional stations. For UK + Ireland, the annual total emissions are 23.0 % higher in INV-E1-O2 ($2.99 \text{ CH}_4 \text{ yr}^{-1}$) compared to INV-E1-O1 ($2.43 \text{ CH}_4 \text{ yr}^{-1}$) (Fig. 2; Table S4). The significant additional observational constraints for UK + Ireland are also reflected in the significantly lower posterior uncertainty for INV-E1-O2 (2σ uncertainty: $0.6 \text{ Tg CH}_4 \text{ yr}^{-1}$) compared to INV-E1-O1 (2σ uncertainty: $1.6 \text{ Tg CH}_4 \text{ yr}^{-1}$; Fig. 2; Table S4). For the BENELUX country region, only a moderate change in the annual total emissions is calculated (INV-E1-O1: $1.71 \text{ Tg CH}_4 \text{ yr}^{-1}$; INV-E1-O2: $1.82 \text{ Tg CH}_4 \text{ yr}^{-1}$; Fig. 2; Table S4), but the spatial distribution of posterior emissions is somewhat different, with higher emissions around the additional station CBW in INV-E1-O2 (Fig. 5). For Switzerland, a larger (relative) difference of posterior emissions is calculated, with annual total emission increasing from $0.15 \text{ Tg CH}_4 \text{ yr}^{-1}$ (INV-E1-O1) to $0.22 \text{ Tg CH}_4 \text{ yr}^{-1}$ (INV-E1-O2).

Using the extended observation data set O2, we have performed additional inversions, using alternatively the prior emission data sets E2 or E3 instead of E1. As for observation data set O1 (discussed in Sect. 4.2.1), the sensitivity of

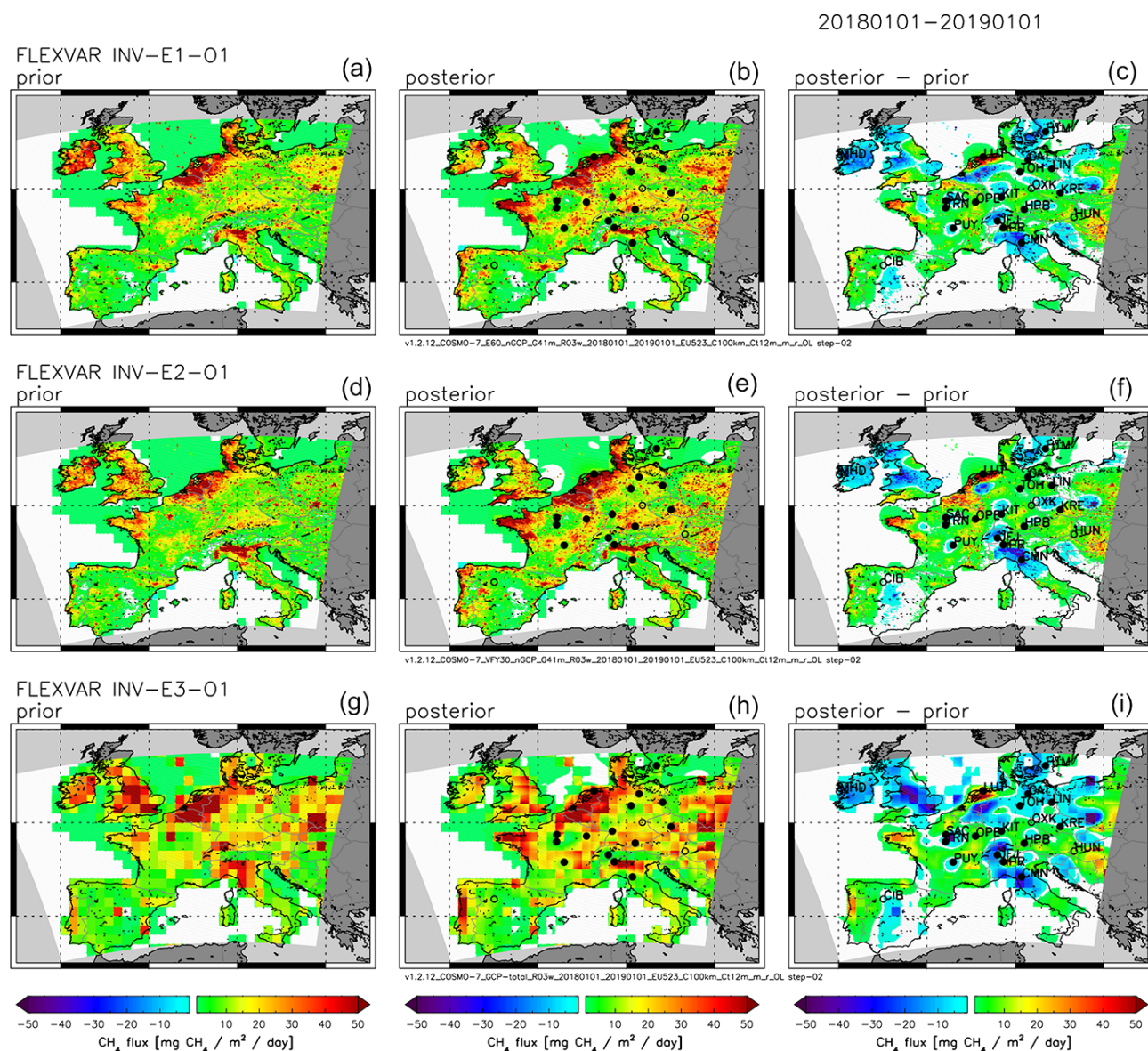


Figure 4. Sensitivity of FLEXVAR inversions to applied prior emission inventories. (a, b, c) Inversion INV-E1-O1 using emission data set E1 as prior estimate. (d, e, f) Inversion INV-E2-O1 using emission data set E2 as prior estimate. (g, h, i) Inversion INV-E3-O1 using emission data set E3 as prior estimate. (a, d, g) Prior emissions. (b, e, h) Posterior emissions. (c, f, i) Difference between posterior and prior emissions. All figures show annual average CH₄ emissions (or change in emissions, respectively) for 2018.

derived annual total emissions to the applied prior emission data set is relatively small (Fig. 2).

Furthermore, additional inversions (of observation data set O2) have been performed using alternatively the particle position baselines (INV-E1-O2-S1) or the alternative parameterization OBS of the model representation error (INV-E1-O2-S2.1). In a similar way, as shown with observation data set O1 (discussed in Sect. 4.1.1. and 4.1.2.), the use of these alternative parameterizations results in generally lower posterior emissions, with lowest posterior emission calculated in

inversion INV-E1-O2-S2.2 (combining the OBS model representation error and the particle position baselines).

4.3 Model comparison and analysis of European CH₄ emissions

In the following we compare the FLEXVAR inversions with inversions using the extended Kalman filter (FLEXKF) system (Sect. 2.3) and TM5-4DVAR (Sect. 2.4). Figure 6 shows the results of these three models using the emission data set E3 as prior and the observation data set O2. Overall, all three inverse models show relatively good consistency of the ma-

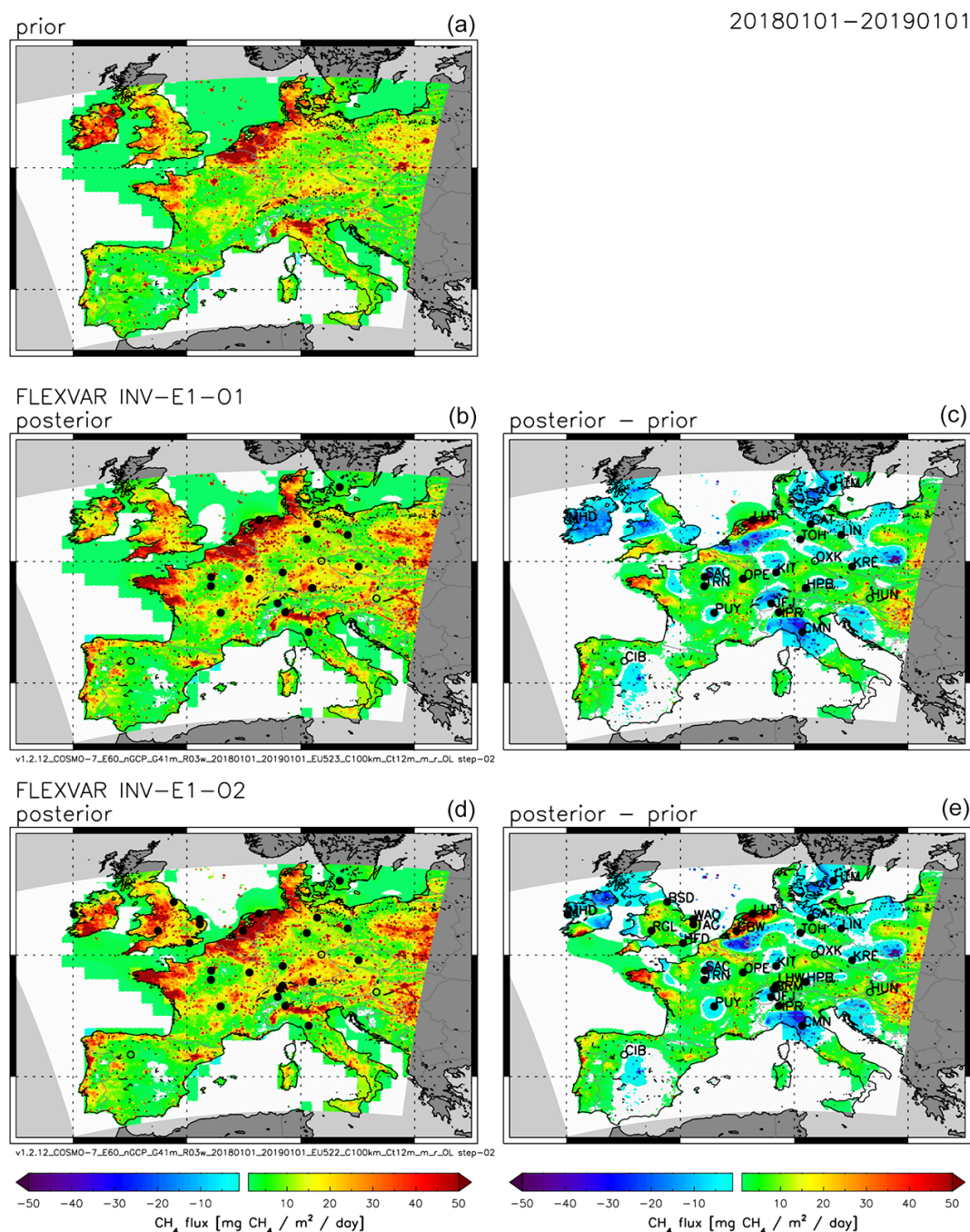


Figure 5. Sensitivity of FLEXVAR inversions to assimilated observations. (a) Prior emission. (b, c) Posterior emissions (left) and difference between posterior and prior emissions (right) for inversion INV-E1-O1 (using observation data set O1). (d, e) Inversion INV-E1-O2 (using observation data set O2). Solid black circles show locations of stations with in situ data and open circles locations of stations with discrete air sampling. All figures show annual average CH₄ emissions (or change in emissions, respectively) for 2018.

for spatial patterns of the derived inversion increments, for example, the increase of emissions over the BENELUX region and north-western France, the decrease of emissions around Paris, and the decrease of offshore emissions over the North Sea compared to the prior emissions. Since FLEXKF uses the same atmospheric transport as FLEXVAR, it is to

be expected that the inversions of these two models should give similar results. Nevertheless, there are also some significant differences visible between the two models, especially for the southern part of France, for which FLEXKF yields overall lower emissions than FLEXVAR. This difference is also clearly visible in the derived country to-

tal emissions (Fig. 7; Table S4), with 10.3 % lower annual total CH₄ emission for France calculated by FLE_xKF (FLE_xKF E3-O2: 3.82 CH₄ yr⁻¹) compared to FLEXVAR (INV-E3-O2: 4.26 CH₄ yr⁻¹). In contrast, FLE_xKF derives somewhat higher CH₄ emissions for BENELUX (6.3 %) and UK + Ireland (6.8 %) than FLEXVAR, while emissions derived for Germany are very similar (within 1.4 %). One major difference between FLE_xKF and FLEXVAR is the different parameterization of the model representation error, leading to a different weighting of the individual observational data points, which can cause differences in the calculated regional inversion increments as shown for FLEXVAR in Sect. 4.1.2. Another difference is the magnitude of the prior uncertainties, though this was shown for FLEXVAR to have a rather small impact on total emissions for the country regions presented in Fig. 2. Furthermore, it is likely that the different inversion techniques have some impact on the calculated solutions. For example, FLE_xKF yields generally smoother seasonal variations of derived emissions, while FLEXVAR shows larger month-to-month variability. The latter are, however, largely filtered out by the use of 3-month running mean values for the seasonal variation of the total emissions of country regions shown in Fig. 7 (left column).

The spatial emission patterns derived by TM5-4DVAR are in general similar to those calculated by FLEXVAR and FLE_xKF (Fig. 6) but show also some differences, for example, around the stations PUY and HPB, where TM5-4DVAR calculates higher emissions than FLEXVAR and FLE_xKF, probably related to the particular challenge to simulate mountain sites and sites in complex topography. Further differences between the models are the different derived seasonal variations of emissions, with larger variations calculated by TM5-4DVAR for Germany, France, and UK + Ireland compared to FLEXVAR and FLE_xKF (while the FLEXVAR inversions using the observation data set O2 show larger variations for BENELUX than the other models). In addition to the different model representation error in TM5-4DVAR, very likely the fundamentally different nature of the models (Eulerian vs. Lagrangian) and the related different simulation of transport play an essential role.

Nevertheless, the differences in the annual total emissions for the country regions are only moderate. For Germany, somewhat higher emissions are calculated by TM5-4DVAR compared to FLEXVAR and FLE_xKF, while the posterior emissions for France, BENELUX, and UK + Ireland derived by TM5-4DVAR are in the range of emissions calculated by FLEXVAR and FLE_xKF.

Figure 7 also includes inversions of the three models using the base observation data set O1. As discussed for FLEXVAR in Sect. 4.2.2., FLE_xKF and TM5-4DVAR also show higher emissions for UK + Ireland, when using O2 instead of O1 due to the six additional stations in data set O2 in that area. Furthermore, FLE_xKF inversions have also been performed using E1 instead of E3 as prior emissions (Fig. 7). As for

FLEXVAR (Sect. 4.2.1), the impact on derived emissions is relatively small.

In order to evaluate the quality of the derived emissions, it is useful to analyse how well the observations are reproduced by the models. Figure S6 compares the statistics (correlation coefficient and rms difference) for the three models (using prior emission data set E3 and observation data set O2). At most stations relatively high correlation coefficients and low rms differences are obtained by all three models. However, stations with larger regional emissions (e.g. LUT, CBW, BRM, IPR) or complex topography (e.g. OXK, IPR) show generally poorer statistical performance. Figure S6 also shows that the best statistical performance is achieved by FLEXVAR with a mean correlation coefficient of $r = 0.86$ (FLE_xKF: $r = 0.84$, TM5-4DVAR: $r = 0.81$) and a mean rms difference of 28.21 ppb (FLE_xKF: 30.53 ppb, TM5-4DVAR: 31.82 ppb). This finding demonstrates that the high spatial resolution of FLEXVAR and FLE_xKF at 7 km × 7 km allows the observations to be somewhat better reproduced than the TM5-4DVAR simulations at 1° × 1°, although – beside the different spatial resolution – other factors (such as fundamental differences in the modelling of transport) are also likely to play a role. The slightly better statistical performance of FLEXVAR compared to FLE_xKF could be due to the higher degree of freedom to optimize the emissions in FLEXVAR but may also be partly related to other factors, such as different weighting of observations due to different parameterizations of the model representation error and differences in the model covariance settings. For example, the FLEXVAR inversion INV-E3-O2 used for the model comparison applies a smaller spatial correlation length ($L_{\text{corr}} = 100$ km) compared to FLE_xKF ($L_{\text{corr}} = 200$ km). Comparison of FLEXVAR inversions INV-E1-O1 and INV-E1-O1-S3.2 shows that increasing the correlation length from 100 to 200 km is indeed slightly deteriorating the statistical performance (mean correlation coefficient and mean rms difference), but nevertheless FLEXVAR (INV-E1-O1-S3.2) still performs slightly better compared to FLE_xKF (inversion FLE_xKF E1-O1). On the other hand, FLE_xKF applies a higher prior uncertainty than FLEXVAR (Table S4) in the model comparison discussed in the paper. For FLEXVAR, increasing the prior uncertainty from 100 % to 200 % (INV-E1-O1-S4.2 vs. INV-E1-O1), slightly improves the statistical performance, i.e. partly compensating for the effect of a larger correlation length (results not shown).

Figure S7 shows the time series of observed and simulated CH₄ mole fractions for all stations (inversion INV-E1-O2), illustrating that in general the synoptic variability is well reproduced at most sites. Furthermore, FLEXVAR also simulates the average diurnal cycle at most sites realistically.

In the following, we compare the annual total CH₄ emissions derived by the inverse models with the anthropogenic CH₄ emissions reported by the countries to UNFCCC (UNFCCC, 2021). For a consistent comparison, it is necessary also to take into account estimates of the natural CH₄ emis-

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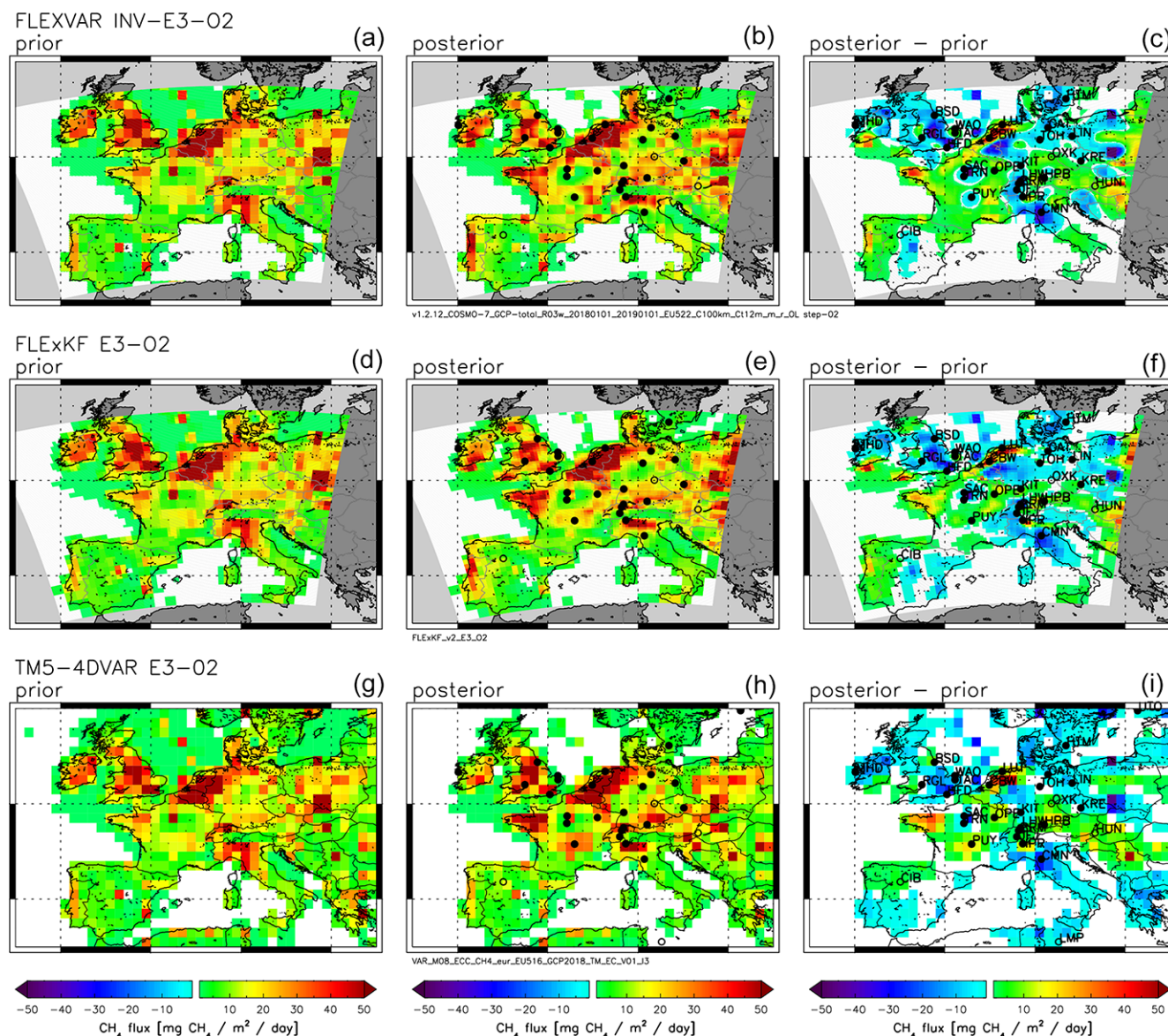


Figure 6. Annual average CH₄ emissions derived for year 2018 using FLEXVAR (a, b, c), FLEXKF (d, e, f), and TM5-4DVAR (g, h, i). (a, d, g) Prior emissions. (b, e, h) Posterior emissions. (c, f, i) Difference between posterior and prior emissions. All three inversions shown here use the same inventory data set E3 as prior and the observation data set O2.

sions, for which we use the bottom-up inventories of natural sources from the GCP-CH₄ data set (Saunio et al., 2020) (Table 2). Furthermore, the comparison of top-down and bottom-up emission estimates requires the inclusion of estimates of their uncertainties. For the uncertainty estimate of the inverse models, we use the range of results from the individual inversions (shown by the solid red rectangles in Fig. 7) and the minimum–maximum values of the 2 σ uncertainty ranges based on the uncertainties computed for the individual inversions (shown by the error bars). The total uncertainty ranges are evaluated separately (1) for the whole

set of FLEXVAR sensitivity inversions (as shown in Fig. 2) and (2) for the whole set of all inversions, i.e. also including all FLEXKF and TM5-4DVAR inversions. The uncertainties of the UNFCCC emissions are based on the uncertainties reported by the countries for the major CH₄ source categories, while estimates of the uncertainty of total CH₄ emissions are not provided by the countries. As in Bergamaschi et al. (2015), we estimate the total uncertainties from the reported uncertainties per category, assuming – among other things – uncorrelated uncertainties for the different major source categories (for further details, see Sect. S2).

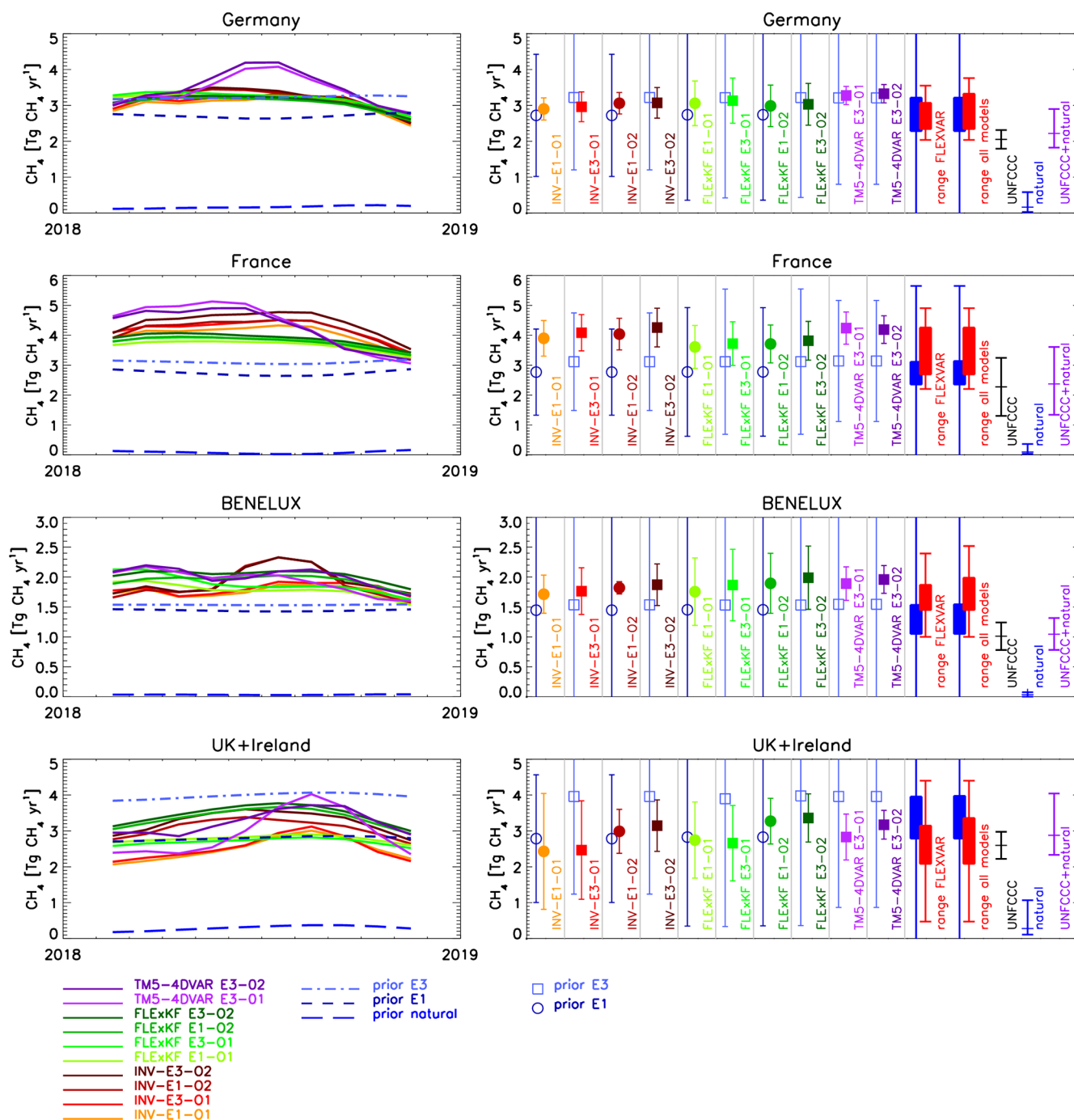


Figure 7. Total CH₄ emissions for Germany, France, BENELUX, and UK + Ireland derived by the three different inverse modelling systems FLEXVAR, FLExKF, and TM5-4DVAR. For FLEXVAR only a subset of inversions is displayed here, while the whole range from all FLEXVAR sensitivity inversions is shown by the first pair (from left to right) of solid rectangles, which is identical to the pair of rectangles shown in Fig. 2. The second pair of rectangles shows the range of prior (blue) and posterior (red) CH₄ emissions from all three models (and the error bars the minimum and maximum values of the 2σ uncertainties of all individual inversions). The black symbols show the anthropogenic CH₄ emissions reported to UNFCCC (and their estimated 2σ uncertainties), blue symbols the natural emissions estimated from the GCP CH₄ inventory, and the violet symbols the sum of anthropogenic and natural bottom-up inventories.

The uncertainties of natural CH₄ emissions from wetlands were estimated from the ensemble of wetland models used for the GCP-CH₄ wetland emissions, taking the minimum–maximum range of the 11 individual wetland models (Poulter et al., 2017). For other natural CH₄ emissions, we assume an uncertainty of 100 %.

Figure 7 shows that the CH₄ emissions estimated by the inverse models are higher than the sum of anthropogenic (UNFCCC) and natural bottom-up inventories for Germany, France, and BENELUX, but the uncertainty ranges of top-down and bottom-up estimates overlap for all three country regions. The smallest overlap, however, is found for BENELUX. In contrast, the top-down estimates for UK + Ireland agree relatively well with the total of anthropogenic and natural bottom-up inventories. A tendency to higher top-down emissions compared to the total (anthropogenic and natural) bottom-up inventories for Germany, France, and BENELUX has also been found in the analysis reported by Bergamaschi et al. (2018a) for the period 2006–2012, but also in that study uncertainty ranges of bottom-up and top-down estimates were overlapping. Similar tendencies to higher top-down emissions are apparent in the VERIFY analyses for the period 2005–2017 (VERIFY, 2021) using a larger ensemble of regional inversions, while global inversions (with coarser resolution) showed in general lower emissions, closer to the UNFCCC estimates for these country regions. Based on the observation that several models showed clear seasonal cycles of the derived emissions with maximum during summer, Bergamaschi et al. (2018a) suggested that higher natural emissions could explain the difference between top-down and bottom-up estimates. The FLEXVAR and FLEXKF inversions analysed in this study, however, show in general only relatively small seasonal variations for Germany, France, and UK + Ireland compared to TM5-4DVAR. The use of seasonal cycles to disentangle anthropogenic and natural sources is further hampered by the fact that the seasonal cycles of major anthropogenic sources are still not well characterized. Also, the anthropogenic emission inventories used in this study show rather different seasonal cycles. Most of the anthropogenic GCP-CH₄ emission categories (which are largely based on EDGARv4.3.2, except biomass burning, Saunio et al., 2020) have no seasonality, except emissions from rice agriculture and biomass burning, which however play only a minor role in Europe. EDGARv6.0 (used for E1) has small seasonal variations of most energy-related source categories but assumes constant emissions for the agricultural sources (except rice) and for waste emission. In contrast, most sectors of the TNO-VERIFYv3.0 inventory (used for E2) show seasonal variations, including significant seasonal variations of all agricultural sources, resulting in significant seasonal variations of the total anthropogenic emissions with maximum emissions in September (Fig. 2).

5 Conclusions

We have presented the novel inverse modelling system FLEXVAR based on the 4DVAR assimilation technique and FLEXPART-COSMO back trajectories driven by COSMO meteorological fields at 7 km × 7 km resolution over the European COSMO-7 domain. A major advantage of the 4DVAR technique is that it allows a much larger number of variables to be constrained (in our study about 1.6 million) compared to analytical inversion techniques. The offline coupling with TM5-4DVAR ensures that the background mole fractions (baselines) used in FLEXVAR are consistent with the global observations assimilated in TM5-4DVAR. We have applied the FLEXVAR system for inversions of European CH₄ emissions in 2018 using 24 stations with in situ measurements, complemented with data from 5 stations with discrete air sampling (and additional stations outside the European COSMO-7 domain used for the global TM5-4DVAR inversions).

We have investigated the sensitivity of the FLEXVAR inversions to internal parameterizations, model settings, and main model input data. Using the particle position baselines yields in general lower derived emissions compared to inversions which apply the Rödenbeck baselines, resulting in differences in the annual total emissions of 5 %–14 % for the analysed country regions (Germany, France, BENELUX, and UK + Ireland). Furthermore we found a significant impact of the applied parameterization of the model representation error. Inversions using the OBS model representation error derive, over large parts of the domain, somewhat lower emissions compared to the METEO model representation error, with differences in the annual total emissions of 0 %–15 % for the analysed country regions. Varying the main parameters of the prior covariance (i.e. horizontal correlation length constant, temporal correlation scale constant, and assumed uncertainties of emissions per grid cell and month) has clearly visible effects on the smaller-scale regional inversion increments, but the impact on the derived annual total emissions remains very small for the analysed country regions, since the differences in the smaller-scale spatial patterns are largely averaged out over larger areas. Furthermore, the dependence of derived emissions on the applied prior emission inventory has been found to be relatively small for the country regions, which are well constrained by the observations. Changing these observational constraints by including additional sites, however, has a significant impact on the inversions, especially in the vicinity of these sites. Using the extended observation data set O2 (which includes six additional in situ stations located on the British Isles) yields 23 %–28 % higher emissions for UK + Ireland compared to inversions using only the base observation data set O1. At the same time, the calculated uncertainty of the posterior emissions for UK + Ireland is significantly reduced by these additional observational constraints.

The comparison of the FLEXVAR inversions with inversions using the extended Kalman filter (“FLEXKF”) system (which both use the same atmospheric transport model) shows overall good consistency of major spatial patterns of the derived inversion increments but some difference (7 %–10 %) for the derived total CH₄ emission of France, probably mostly related to the use of different parameterizations of the model representation error. TM5-4DVAR also shows in general similar inversion increments and derives posterior emissions for France, BENELUX, and UK + Ireland in the range of emissions calculated by FLEXVAR and FLEXKF. For Germany, however, TM5-4DVAR estimates 5 %–11 % higher emissions than the other two models.

The FLEXVAR and FLEXKF inversions at high spatial resolution of 7 km × 7 km allow for a better reproduction of the observations compared to the TM5-4DVAR simulations at 1° × 1°, reflected in the achieved higher correlation coefficient and lower rms difference between simulations and observations. Furthermore, the statistical performance of FLEXVAR is slightly better than that of FLEXKF, which could be due to the higher degree of freedom to optimize the emissions in FLEXVAR but could be partly related also to other differences in the inversions, for example, the different parameterizations of the model representation error and differences in the model covariance settings.

The inverse models derive higher annual total CH₄ emissions in 2018 for Germany, France, and BENELUX compared to the sum of emissions reported to UNFCCC and natural emissions (estimated from the GCP-CH₄ inventory), but the uncertainty ranges of top-down and bottom-up estimates overlap for all three country regions. In contrast, the top-down estimates for UK + Ireland agree relatively well with the total of anthropogenic and natural bottom-up inventories.

Our study demonstrates that the new FLEXVAR system can be applied for verification of reported emissions, as planned, for example, by Empa for its quasi-operational system to estimate Switzerland’s annual CH₄ emissions as a contribution to the Swiss National Inventory Reporting. FLEXVAR inversions with the configuration presented in this paper could be performed for the years 2002 to 2021, the period for which meteorological fields from the COSMO-7 model at 7 km × 7 km resolution are available. For analysis periods after 2021, the use of different high-resolution meteorological input fields could be considered, such as the operational analysis data from the ECMWF IFS model at high resolution (0.1° × 0.1°) or the operational MeteoSwiss COSMO-1 analysis at horizontal resolution of 1 km × 1 km. COSMO-1, however, is limited to the larger Alpine area but can be nested into FLEXPART-IFS. A FLEXPART-COSMO modelling system using COSMO-1 has already been developed by Empa, including a modification of the turbulence parameterization (Katharopoulos et al., 2022), which is required owing to the very high resolution of 1 km × 1 km.

While the relatively good agreement among the three models used in this study gives some confidence in the robust-

ness of the inverse modelling results, further specific studies should be performed to assess the quality of the top-down estimates independently. Such assessments should include the comparison with further inverse models, comparison with independent regional emission estimates (e.g. based on aircraft or satellite measurements), and a more detailed validation of the applied atmospheric transport models (especially regarding the simulation of boundary layer height dynamics and vertical transport).

Code and data availability. The code of the FLEXVAR inverse modelling system is available upon request. The atmospheric observations from ICOS are available at <https://www.icos-cp.eu/data-products/atmosphere-release> (ICOS RI, 2021). NOAA data are available at <https://doi.org/10.15138/VNCZ-M766> (Dlugokencky et al., 2021), AGAGE data at <https://doi.org/10.15485/1781803> (Prinn et al., 2021), and UK DECC data at <https://archive.ceda.ac.uk/> (O’Doherty et al., 2022).

Supplement. The supplement related to this article is available online at: <https://doi.org/10.5194/acp-22-13243-2022-supplement>.

Author contributions. PB led the FLEXVAR development. AS designed the FLEXVAR concept and performed the technical implementation of the FLEXVAR code. PB performed the FLEXVAR and TM5-4DVAR inversions. DB performed the FLEXKF inversions and contributed to the integration of the FLEXPART-COSMO back trajectories into FLEXVAR. JMH generated the FLEXPART-COSMO back trajectories and contributed to the development of the FLEXVAR emission preprocessing. SH contributed to the development of the FLEXPART-COSMO modelling system. PB prepared the paper and figures with contributions from DB, AS, and JMH. SH, MR, TA, TB, HC, SC, MD, GF, AF, DK, XL, MLi, MLo, GM, JMW, SOD, BS, MS, PT, GV, and CYK provided atmospheric observation data.

Competing interests. The contact author has declared that none of the authors has any competing interests.

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Acknowledgements. The VERIFY project has received funding from the European Union’s Horizon 2020 Research and Innovation programme under grant agreement no. 776810. We are grateful to ECMWF for providing computing resources under the special projects “Improve European and global CH₄ and N₂O flux inversions (2018–2020)” and “Extend and improve CH₄ flux inversions at global and European scale (2021)”. Furthermore, we thank Bradley Matthews for the compilation of uncertainties of emissions

reported to UNFCCC by EU member states. ICOS Switzerland was funded by the Swiss National Science Foundation, in-house contributions, and the State Secretariat for Education, Research and Innovation.

Financial support. This research has been supported by Horizon 2020 (grant no. VERIFY (776810)).

Review statement. This paper was edited by Ilse Aben and reviewed by two anonymous referees.

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