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Bridging the spatial gaps of the Ammonia Monitoring Network using satellite ammonia measurements

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Abstract. Ammonia (NH_3) is a key precursor to fine particulate matter ($PM_{2.5}$) and a primary form of reactive nitrogen. The limited observations of NH_3 hinders further understanding of its impacts on air quality, climate, and biodiversity. Currently, NH_3 ground monitoring networks are limited in number across the globe, and even in the most established networks, large spatial gaps exist between sites and only a few sites have records that span longer than a decade. Satellite NH_3 observations can be used to discern trends and fill spatial gaps in networks, but many factors influence the syntheses of the vastly different spatiotemporal scales between surface network and satellite measurements. To this end, we intercompared surface NH_3 data from the Ammonia Monitoring Network (AMoN) and satellite NH_3 total columns from the Infrared Atmospheric Sounding Interferometer (IASI) in the contiguous United States (CONUS) and then performed trend analyses using both datasets. We explored the sensitivity of correlations between the two datasets to factors such as satellite data availability and distribution over the surface measurement period as well as agreement within selected spatial and temporal windows. Given the short lifetime of atmospheric ammonia and consequently sharp gradients, smaller spatial windows show better agreement than larger ones except in areas of relatively uniform, low concentrations where large windows and more satellite measurements improve the signal-to-noise ratio. A critical factor in the comparison is having satellite measurements across most of the measurement period of the monitoring site. When IASI data are available for at least 80% days of AMoN's 2-week sampling period within a 25 km spatial window of a given site, IASI NH_3 column concentrations and the AMoN NH_3 surface concentrations have a correlation of 0.74, demonstrating the feasibility of using satellite NH_3 columns to bridge the spatial gaps existing in the surface network NH_3 concentrations. Both IASI and AMoN show increasing NH_3 concentrations across CONUS (median: 6.8%·yr^{-1} vs. 6.7%·yr^{-1}) in the last decade (2008 - 2018), stressing the rising importance of NH_3 in terms of nitrogen deposition. NH3

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trends for AMoN sites correlates with IASI NH_3 trend IASI and AMoN NH_3 trend (r = 0.66) and show a similar spatial pattern, with the highest increases in the Midwest and eastern U.S., and NH_3 trend for AMoN sites correlates with IASI NH_3 trend (r = 0.66). In spring and summer, increases of NH_3 were larger than 10%·yr^{-1} in the eastern U.S. and Midwest (cropland dominated) and western U.S. (pastureland dominated), respectively. In terms of trend in NH_3 hotpots (defined as regions where the IASI NH_3 column is larger than the 95th percentile of 11-year CONUS map, 6.7 × 10^{15} molec/cm^2), these largest emissions sources are also experiencing increasing concentrations over time with the median of NH_3 trend is 4.7% · yr^{-1}. IASI data show large NH_3 increases in urban areas (8.1%·yr^{-1}), including 8 of the top 10 most populous regions in the CONUS, where AMoN sites are sparse. The increasing NH_3 could have detrimental effects on nearby eco-sensitive regions through nitrogen deposition and on aerosol chemistry in the densely populated urban areas, hence needs immediate attention.

1 Introduction

Gas phase ammonia (NH_3) is the most abundant alkaline gas in the atmosphere, mainly emitted from agricultural activities such as nitrogen fertilizer applications and livestock waste volatilization (Bouwman et al., 1997; Paulot et al., 2014). As a major precursor to fine particulate matter ($PM_{2.5}$), NH_3 critically affects aerosol heterogeneous chemistry, air quality, visibility, human health, and climate (Hauglustaine et al., 2014; Hill et al., 2019; Lawal et al., 2018; Malm et al., 2004). Ammonia neutralizes sulfuric acid (H_2SO_4) and nitric acid (HNO_3) in the atmosphere to form ammoniated aerosols, ammonium sulfate ($(NH_4)_2SO_4$) and ammonium nitrate (NH_4NO_3), which in total can contribute to more than 50 % of total $PM_{2.5}$ mass (Feng et al., 2020). NH4NO3 is critical during wintertime haze periods because the cold and humid condition favor its formation (Shah et al., 2018; Zhai et al., 2021). Besides, NH_3 plays an important role in the nitrogen cycle. Wet deposition of NH_4^+ dominates the wet inorganic nitrogen deposition at nearly 70% of monitoring sites in the United States (Li et al., 2016). Total NH_x (≡NH_3(g) + NH_4^+ (aq)) deposition is expected to become even more dominant in the future because NO_x emissions decrease under pollution control while NH_3 emissions are predicted to continue to increase with the rising global food demands (Erisman et al., 2008; Goldberg et al., 2021; Pinder et al., 2008). Excessive NH_3 deposition in the non-agricultural ecosystems can reduce biodiversity, result in soil acidification, and increase eutrophication, especially in the sensitive ecosystems (Ellis et al., 2013; Phoenix et al., 2006).

Although NH_3's importance has been well recognized, routine NH_3 observations are lacking even in countries with comprehensive monitoring networks, partly due to the difficulty of measuring gas phase NH_3 (von Bobrutzki et al., 2010; Fehsenfeld et al., 2002). The Ammonia Monitoring Network (AMoN) (Puchalski et al., 2015) is the only routine set of NH_3 measurements in the United States, with 110 active AMoN sites in the contiguous United States (CONUS) in 2021, providing high-quality surface observations of NH_3. AMoN data have been used widely for model evaluation and long-term trend analysis (Butler et al., 2016; Nair et al., 2019; Yao and Zhang, 2016, 2019). AMoN only provides bi-weekly NH_3 observations,

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in contrast to monitoring networks for two other important gas phase precursors of $PM_{2.5}$, SO_2 and NO_2, which provide hourly or daily scale observations. $PM_{2.5}$, SO_2, and NO_2 are directly regulated as criteria pollutants, however contributions from NH_3 emissions sources must be considered in State Implementation Plan (SIP) demonstrations for areas out of attainment for $PM_{2.5}$, which can be a challenge for areas lacking NH_3 measurements (EPA 2023).

Population weighted $PM_{2.5}$ concentrations are widely used to estimate the health effects of $PM_{2.5}$, however, the sparse number of NH_3 sites with only biweekly or monthly resolution makes it difficult to derive population weighted $PM_{2.5}$ precursor datasets. Gas phase NH_3 is critical to determine the partitioning of the total NH_x (Hennigan et al., 2015), and the lack of gas phase NH_3 observations hampers the evaluation of chemistry models. The ISORROPIA-II thermodynamic model has been extensively adopted to compute the equilibrium composition for the inorganic aerosol systems (Fountoukis and Nenes, 2007) and requires both gas and aerosol phase data as input to provide accurate and robust results (Hennigan et al., 2015). However, the limited number of NH_3 ground monitoring sites currently prevents synthesizing the AMoN NH_3 data with other ground monitoring networks, e.g., IMPROVE, as input for ISORROPIAII (Pan et al., 2020). GEOS-Chem implemented with ISORROPIA-II was found to significantly underestimate gas phase NH_3 and overestimate NH_4^+ in winter (Holt et al., 2015; Nair et al., 2019; Walker et al., 2012), with the normalized NH_4^+ mean biases as high as 86% in January at sites for the Interagency Monitoring of Protected Visual Environments (IMPROVE) (Holt et al., 2015). The lifetime of NH_3 ranges from hours to days, hence large spatiotemporal variability exists (Golston et al., 2020; Miller et al., 2015; Wang et al.; 2021), and large spatial gaps exist in the current AMoN. Currently there are no AMoN sites in some states, e.g., North Dakota and South Dakota, and only 12 sites are within the characteristic length scale (12 km) of NH_3 hotspots regions (Wang et al., 2021). Ten national parks in the U.S. are within 100 km of an NH_3 hotspot, and more observations are needed to quantify the impacts of these hotspots on dry NH_3 deposition in these regions (Pan et al., 2021). A lack of long-term AMoN data also hinders the possibility of investigating NH_3 trends in the CONUS. Increasing NH_3 concentrations are observed using AMoN data, yet all of the previous trend analyses are limited to fewer than 20 AMoN sites that may not be representative of NH_3 trends in the CONUS (Butler et al., 2016; Yao and Zhang, 2016, 2019).

Satellite NH_3 observations are on a global and daily basis, providing long-term trends and ubiquitous coverage. Instruments that measures NH_3 include the Infrared Atmospheric Sounding Interferometer (IASI) on the MetOp satellites, Cross-track Infrared Sounder (CrIS) on NOAA and NASA Suomi National Polar-orbiting Partnership (S-NPP), Tropospheric Emission Spectrometer (TES) on NASA Aura satellite, Atmospheric Infrared Sounder (AIRS) on NASA EOS Aqua satellite, and Thermal and Near Infrared Sensor for Carbon Observations – Fourier Transform Spectrometer (TANSO-FTS) on the Greenhouse Gases Observing SATellite (GOSAT) (Clarisse et al., 2009; Shephard et al., 2011; Shephard & Cady-Pereira, 2015; Someya et al., 2020; Warner et al., 2016). Satellite NH_3 data have been widely used to constrain NH_3 emissions, estimate NH_3 deposition, and analyze NH_3 trends (Cao et al., 2020; Chen et al., 2020; Kharol et al., 2018; Van Damme et al., 2021). Van Damme et al. (2021) utilized 11-year IASI NH_3 observations a and found a worldwide NH_3 increase (12.8 ± 1.3 %) from

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2008 to 2018 with especially large increases in east Asia (75.7 ± 6.3 %) and North America (26.8 ± 4.5 %). Warner et al. (2017) used 14-year AIRS NH_3 measurements and found statistically significant NH_3 increase (2.61%·yr^{-1}) in the U.S. from 2002 to 2016.

The global daily coverage and long-term data record make it possible for satellite observations to fill the spatial and temporal gaps of the current ground monitoring networks. Although limited in numbers, the validations of satellite NH_3 observations with in-situ measurements provide confidence in integrating the two datasets (Guo et al., 2021; Sun et al., 2015). Sun et al. (2015) performed the first daily and pixel scale satellite NH_3 validations using TES NH_3 columns and airborne NH_3 observations in the San Joaquin Valley of California, USA, showing that the differences between the total NH_3 column and the in-situ total column were within 6 %. However, the validation included only 9 TES pixels, and TES is no longer in operation now. Guo et al. (2021) showed that IASI NH_3 columns and NH_3 columns derived from airborne and ground-based NH_3 observations were indistinguishable from one another on daily and pixel bases in Colorado, USA, in summer. All of these validation works were performed in certain seasons and were limited to source regions with high NH_3 concentrations (Guo et al., 2021; Sun et al., 2015; Warner et al., 2016). Ground-based FTIR NH_3 observations provided a better temporal coverage for evaluating IASI and CrIS NH_3 retrievals, however, low concentration sites were excluded from the evaluation and only ~ 10 sites were included across the globe (Dammers et al., 2016; Dammers et al., 2017). Furthermore, FTIR-based measurements also have not been directly validated against in-situ measurements of NH_3 vertical profile themselves.

To capitalize on the benefits of both surface and satellite observations and synthesize these datasets, a detailed understanding of the comparison between IASI NH_3 column concentrations and AMoN NH_3 surface concentrations is necessary. Here we focus on IASI NH_3 measurements because it offers the longest data record (2008 - present) among the satellite NH_3-measuring instruments. The comparison between AMoN and IASI is complex because AMoN is a ground-based, point measurement integrated over fourteen days, whereas IASI is a space-borne volumetric measurement averaged over the pixel footprint at the instantaneous overpass time. There are several factors that need to be taken into consideration:

(1) The extent to which the IASI NH_3 column represents the surface AMoN NH_3 concentration: Knowledge of NH_3 vertical profiles in the atmosphere are limited due to the lack of observational data, and model simulated NH_3 vertical profiles are often biased compared with the airborne measurements (Schiferl et al., 2016). Ammonia is mostly concentrated in the planetary boundary layer (PBL) because of its short lifetime (~hours to days) and surface emission sources (Dentener & Crutzen, 1994; Guo et al., 2021; Sun et al., 2015; Seinfeld & Pandis, 2016). Sun et al. (2015) showed that NH_3 was almost well mixed in the lower PBL, and the TES NH_3 columns were strongly correlated ($R^2 = 0.82$) with the median NH_3 mixing ratios measured at the surface, demonstrating that satellite NH_3 columns could represent the ground NH_3 concentrations. Van Damme et al. (2015) converted IASI NH_3 columns to surface NH_3 concentrations using fixed NH_3 profiles generated by GEOS-Chem, then performed monthly comparisons with ground monitoring networks. IASI derived surface NH_3 observations are in fair

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agreement with ground observations in Europe, China, and Africa, but are limited to a small number of sites in each region for a short time range, e.g., 27 sites in Europe in 2011 (Van Damme et al., 2015). Furthermore, the latest IASI NH_3 products have switched to a new algorithm and no longer use a fixed NH_3 profile (Whitburn et al., 2016; Van Damme et al., 2017).

(2) Optimal spatial window for comparing and integrating satellite pixels and AMoN sites: Previous comparisons of satellite NH_3 retrievals with observations from ground monitoring networks simply averaged the data from the monitoring site within a coarse model grid (~ 100 km) with the averaged modeling/satellite NH_3 concentration of the whole grid (Kharol et al., 2018; Nair et al., 2019; Van Damme et al., 2015). If NH_3 concentrations are uniformly distributed within the spatial window, increasing the spatial window will increase the number of IASI pixels and decrease the signal-to-noise ratio. However, the spatial heterogeneity of NH_3 is quite large near hotspots due to its short lifetime (Golston et al., 2020; Miller et al., 2015; Wang et al., 2021; Warner et al., 2016). The relationship between spatial window size and satellite/surface measurements agreement needs to be examined in more details.

(3) Temporal distribution of satellite measurements across the two-week AMoN sampling period: Previous comparisons of model or satellite products against surface observations did not consider the distribution of IASI measurements during the two-week sampling period (Kharol et al., 2018; Nair et al., 2019; Van Damme et al., 2015). AMoN measures continuously, whereas a series of cloudy days would preclude any valid satellite measurements. Therefore, any AMoN/satellite comparison is intrinsically biased towards clear sky days on the satellite side but includes all conditions for the AMoN site.

(4) Number of available IASI pixels in the comparison: Guo et al. (2021) has shown that, even at low column amounts, IASI NH_3 has no known biases. AMoN is an extremely sensitive measurement of NH_3, far more precise than any satellite NH_3 product (NADP, 2023; Van Damme et al., 2017). Therefore, increasing the number of satellite measurements within a certain spatiotemporal window is expected to improve the signal-to-noise ratio in the satellite measurements and may lead to improved agreements with AMoN under clean conditions.

(5) Regional and seasonal variabilities: Different regional and seasonal patterns are expected to influence the comparison. The performances of thermal infrared sounders are highly affected by the thermal contrast between the surface air temperature and skin temperature (Clarisse et al., 2010). In winter, low thermal contrast results in low sensitivity, which explains the low number of IASI pixels in winter compared to summer (Clarisse et al., 2010; Guo et al., 2021). Kharol et al. (2018) showed that CrIS surface NH_3 concentrations had an overall mean CrIS–AMoN difference of ~+15%, however, they only averaged CrIS data over the warm season in 2013.

In this study, to demonstrate the capabilities of using IASI NH_3 observations to augment the ground monitoring network, we performed a comprehensive comparison between IASI and AMoN on weekly/seasonal scales. We directly compare the

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correlation between IASI NH_3 columns with AMoN surface NH_3. We avoided direct comparisons when converting column NH_3 into surface concentrations because of possible biases introduced by assuming vertical profiles, boundary layer heights at local sites, and gas phase - aerosol partitioning. The impacts of the different factors on the comparison are examined in the context of points raised above. After identifying the most optimal method for comparison, we examined NH_3 trends over AMoN sites and the larger applicability of using satellite retrievals to discern NH_3 trends over regions and seasons lacking AMoN data.

2 Data and methods

2.1 Satellite NH_3 observations

IASI is an infrared sounder launched on board of the MetOp-A, MetOp-B, and MetOp-C platforms in sun-synchronous orbits since October 2006, September 2012, and November 2018, respectively. IASI has a swath of 2200 km and provides global coverage twice per day at around 09:30 and 21:30 mean local solar time. At nadir, the IASI footprint has a 12-km diameter. The first IASI NH_3 product was developed by Clarisse et al. (2009) by converting the brightness temperature differences into total NH_3 columns. Later on, a flexible and robust retrieval algorithm based on an artificial neural network for IASI (ANNI) (Whitburn et al., 2016) was developed. The latest version consists of a reanalyzed dataset provided with the European Centre for Medium-Range Weather Forecasts Re-Analysis v5 (ERA5) as its meteorological input (Van Damme et al., 2017; Van Damme et al., 2021). Because the meteorological input for reanalysis data is coherent in time, it is the more appropriate dataset to be used to study trends. For the present analyses, we used IASI version 3.1 reanalysis (v3.1r) retrieval product data from the MetOp/A (2008-2018) and MetOp/B (2013-2018) satellites (limited to cloud fraction ≤ 25 %). Only the morning orbits were analyzed because of higher sensitivity than the evening overpasses (Clarisse et al., 2010).

2.2 Ground-based observations

AMoN is the only network providing a consistent, long-term record of NH_3 gas concentrations across the United States. AMoN was established by the National Atmospheric Deposition Program (NADP) in October 2007 and expanded to 19 sites in 2010 and 105 sites in 2018. AMoN deploys Radiello® passive samplers that rely upon diffusion theory, where gas phase NH_3 is adsorbed onto a cylindrical interior filter and extracted as NH_4^+ to be analyzed by Flow Injection Analysis (FIA). AMoN provides biweekly surface NH_3 concentrations, and the network detection limit is 0.083 mg NH_4^+ L^{-1} (~0.078 µg NH_3 m^{-3}) for the 2-week samples in 2020 (NADP, 2023). The Radiello passive samplers were found to be biased low by 37% against denuders used as reference method (Puchalski et al., 2011). In this study, we are comparing the relative variations instead of absolute concentrations of IASI and AMoN, therefore the low bias of AMoN measurements is not as relevant to the outcome.

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We incorporated data from all AMoN sites with one notable exception. Using satellite imagery, we identified that the AMoN site in Logan, Utah (UT01), is located only ~ 100 m away from a livestock farm. Ammonia concentrations downwind of a beef/dairy feedlot at this distance are far above background levels and unrepresentative of those at the local-regional scales (1-10 km) (Golston et al., 2020; Miller et al., 2015; Sun et al., 2018). Concentrations at UT01 are expected to be strongly dependent upon the extent to which local winds blow directly from that farm to the AMoN site throughout the two-week integration period. Not surprisingly, the UT01 site has the highest annual mean concentration (16.2 μg/m^3) in the entire AMoN network (three times higher than the next one). Furthermore, this AMoN site may be particularly susceptible to trends in animal operations or management practices at the farm. While it is possible the measurements of UT01 are representative of the local region, it is beyond the scope of this work to make such an assessment of its representativeness.

2.3 Trend analyses

2.3.1 Oversampled NH_3 maps

From 2008 to 2018, a 0.02° × 0.02° (~2 km) annual mean NH_3 map in the CONUS was created each year based on a physical oversampling algorithm that represents the satellite spatial response functions as generalized 2-D super Gaussian functions (Sun et al., 2018). This algorithm weighs IASI measurements by their uncertainties which include varying sensitivities to thermal contrast as described in Sun et al. (2018) and Wang et al. (2021). For each year, seasonally averaged oversampling maps were also generated for spring (March, April, and May, MAM), summer (June, July, and August, JJA), fall (September, October, and November, SON), and winter (December, January, and February, DJF). For each season, we were able to achieve sufficiently overlapped IASI pixels through calculating the sum of the unnormalized spatial response function (SRF) of the oversampling results (Sun et al., 2018; Wang et al., 2021).

2.3.2 Mann-Kendall test and Theil-Sen's slope estimator

We use the Mann-Kendall (MK) test and Theil-Sen's slope estimator for NH_3 trend analyses. The non-parametric Mann-Kendall test and Theil-Sen's slope estimator are widely used in detecting trends of variables in meteorology and hydrology fields (Ahn and Merwade, 2014; Kendall, 1975; Yue and Wang, 2004). The Kendall rank correlation coefficient, commonly referred to as Kendall's τ coefficient, is a statistic used to measure the rank correlation. An MK test is a non-parametric hypothesis test for statistical dependence based on the Kendall's τ coefficient. The Theil–Sen's slope estimator is commonly used to fit a line to data points by calculating the median of the slopes of all lines through pairs of points.

Different from simple linear regression, the Mann-Kendall test and Theil-Sen's slope estimator do not require the data to follow normal distribution and therefore are more robust to any outliers (Yue and Wang, 2004). This method is computationally

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efficient and is insensitive to outliers. For skewed and heteroskedastic data, the Theil-Sen estimator can be significantly more accurate than linear least squares regression. For normally distributed data, the Theil-Sen estimator competes well against the least squares in terms of statistical power (Yue and Wang, 2004).

3 IASI & AMoN comparison

3.1 Sensitivity to spatial windows

For the initial analysis, we first show the simplest way of comparing the satellite measurements with ground observations. In other words, we center on each AMoN site, average all IASI observations within a given radius of the AMoN site for the sampling time frame (2 weeks) for comparison, and refer to that radius as a spatial window. If the distribution of NH_3 pixels is spatially uniform, increasing the spatial window may improve the correlation between the two because of a larger number of IASI pixels. Larger spatial windows include more IASI pixels than smaller spatial windows but at the expense of potentially not being representative of the AMoN site. In addition, a larger region is likely to encompass NH_3 spatial gradients. In contrast, small spatial windows may only include a limited number of IASI pixels, encompassing more inherent noise in the satellite measurements, especially if close to the detection limit. Each integrated 2-week AMoN measurement for each site was correlated with any relevant satellite data within the spatial window (total of 104 AMoN sites with 16,093 measurements). Correlations between IASI and AMoN for different spatial windows (15 km, 25 km, 50 km, and 100 km) are summarized in Table 1. The minimum spatial window radius of 15 km is based upon an approximate scale for NH_3 hotspots (Wang et al. 2021).

As the spatial window becomes larger, mean temporal coverage (defined as the percentage of days with available IASI data of the 2-week AMoN sampling period) and number of IASI pixels both have significant increases, but the Pearson's r coefficient only increases slightly from 0.35 at a 15 km spatial window to 0.44 at a 100 km spatial window. Indeed, doubling the spatial window from 50 km to 100 km yields an almost tripled mean number of IASI pixels, yet maintains the almost the same correlation with r = 0.45 and r = 0. 44, respectively. This indicates that including IASI pixels at longer distances from the AMoN site may not be representative of the AMoN site, especially near sources or regions with complex topography. The slightly increased r value over spatial window range may result from a tradeoff between averaging spatial gradients versus integrating a larger number of IASI pixels to improve the signal-to-noise ratio of the satellite measurements. To balance these competing effects, we select 25 km as the nominal spatial window for the further comparisons.

Table 1. AMoN & IASI comparison results for different spatial windows

Spatial window	15 km	25 km	50 km	100 km

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Pearson's r	0.35	0.41	0.45	0.44
Mean temporal coverage per pair (%)	31	44	57	71
Mean # IASI pixels per pair	7	17	69	278
# AMoN & IASI pairs	14734	15543	15933	16022

3.2 Sensitivities to temporal coverage and number of IASI pixels

NH_3 is a short-lived species with a complicated diurnal profile (Nair and Yu, 2020) and the potential for large day-to-day concentration changes because of the variability in emissions, wind speed, temperature, PBL height, and aerosol partitioning (Golston et al., 2020; Miller et al., 2015). Thus, the temporal distribution of satellite measurements within the AMoN measurement period may impact the comparison. Fig. 1 illustrates four examples where the number of IASI pixels, and their relative distribution throughout the 2-week AMoN integration period, may impact the comparison (25 km spatial window). An ideal comparison case would have a uniform number of IASI measurements on each day during the approximate 14-day AMoN measurement period, similar to the case shown in Fig. 1a. In this case, there is no specific day having more weight than the other when calculating the biweekly mean. More common, however, are cases where some days have no satellite measurements due to clouds or low thermal contrast. For example, Fig. 1b has one missing day (N=23 satellite measurements) but with an otherwise even distribution throughout the remainder of the period, while Fig. 1c (N=24) has nearly the same number of satellite measurements as Fig. 1b but clustered on only 8 of the 15 days. Finally, there are also many cases where selected day(s) have few or no IASI measurements at all (Fig. 1d). When neither temporal coverage nor the number of IASI pixels are high, one can still calculate the matched IASI NH_3 column for this AMoN sample, but the result is unlikely to be more representative than a more temporally distributed comparison.

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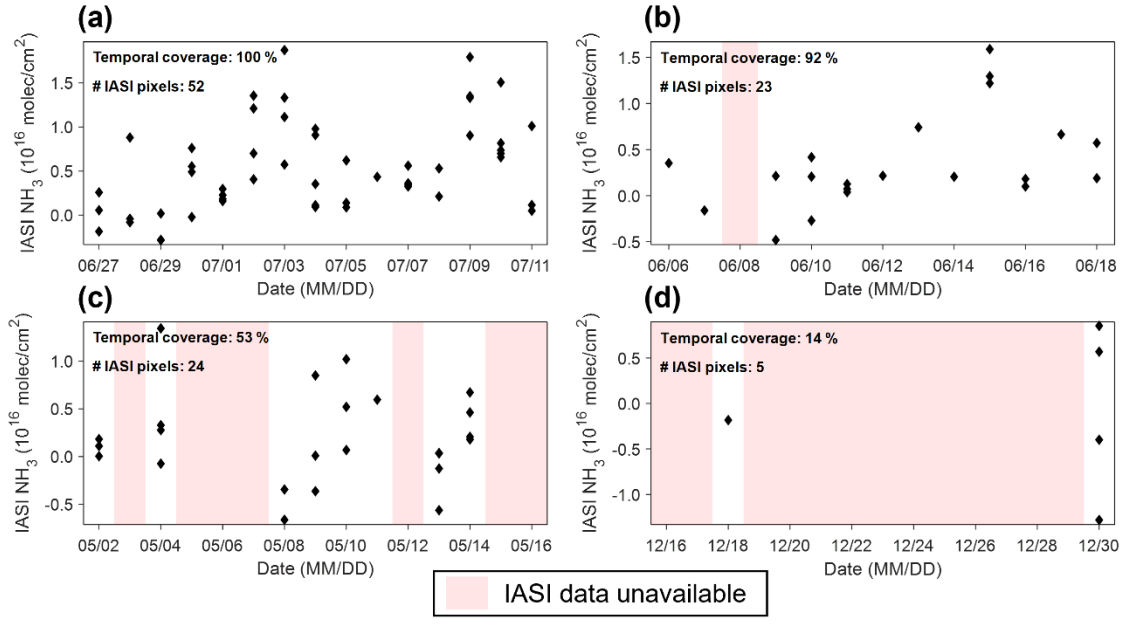

Figure 1. Examples of IASI data temporal coverage over the biweekly AMoN sampling period: **(a)** several IASI measurements every day during the 2-week sampling period; **(b)** a few IASI measurements for most time of the 2-week sampling period; **(c)** many IASI measurements but only in several days during the 2-week sampling period; **(d)** sparse IASI measurements for only several days during the 2-week sampling period.

To this end, we explore the correlation with IASI data's temporal coverage of the 2-week sampling period and total number of IASI pixels within the 2-week AMoN sampling period using the 25 km spatial window. For example, the temporal coverages for Fig. 1 are 100%, 92%, 53%, and 14%, respectively, and the number of IASI pixels are 52, 23, 24, and 5, respectively. The impact of different temporal averaging and number of IASI pixels requirements are summarized in Table 2 and Table 3, respectively. Increasing temporal coverage and number of IASI pixels both yield higher r values than any of the simple spatial windows alone. Table 2 shows that the correlation improves to $r = 0.74$ when the temporal coverage is $\geq 80\%$, suggesting a significant impact of temporal coverage of the IASI data. The IASI and AMoN correlations also increase over a simple spatial window with increasing numbers of IASI pixels, yet the impact is not as strong ($r = 0.63$ for $N \geq 40$) as the sensitivity to temporal coverage.

Table 2. The impact of IASI data's temporal coverage for the 2-week AMoN sampling period (25 km spatial window)

IASI temporal coverage per pair (%)	[0, 20)	[20, 50)	[50, 80)	[80, ∞)

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r	0.17	0.29	0.47	0.74
Mean # IASI pixels per pair	3	13	26	38
# AMoN & IASI pairs	1766	7641	5137	999

Table 3. The impact of # IASI pixels (25 km spatial window)

# IASI pixels per pair	**[0, 10)**	**[10, 20)**	**[20, 40)**	**[40, ∞)**
r	0.16	0.37	0.50	0.63
Mean temporal coverage per pair (%)	22	42	61	80
# AMoN & IASI pairs	4533	5025	5309	676

Because the temporal coverage and number of IASI pixels are not independent variables, additional analyses are conducted to study the sensitivity of these two effects using Monte-Carol method. First, the available dataset is filtered to cases when at least one of the fourteen days have multiple IASI measurements per AMoN measurement, at least 7 days of the 14-day sampling period had at least one IASI measurement, and the total number of IASI pixels is at least 20. The number of days with available IASI measurement is denoted by T. Two opposite approaches are explored for 104 qualified AMoN sites:

(1) Maximized temporal coverage (TC_max): only one IASI pixel is randomly selected to represent that day, and the total number of IASI pixels equals T (T≤14). In this case, the temporal coverage is maximized.

(2) Minimized temporal coverage (TC_min): only days with the largest number of IASI pixels are selected until the total number of IASI pixels equals T (T≤14). In this case, the temporal coverage is minimized, and the total number of selected IASI pixels is same with TC_max.

For each AMoN site, we performed the two different sampling strategies for 100 times, then calculated the median r value to represent each site using the maximum and minimum coverage approaches. Fig. 2a shows the histogram and normalized fit of change in r (Δr = TC_max-TC_min) for each site between the two scenarios with the number of bins determined by Sturge's

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rule. The increased correlation of $\Delta r = 0.45 \pm 0.28$ shows the large impact of temporal coverage. The total number of IASI pixels used for the two strategies were identical.

To further investigate the impact of including more IASI pixels after maximizing temporal coverage, we also test the process described in (1) and then randomly added (20-T) more IASI pixels from the remaining IASI pixels and referred to it as TC_max_add. Fig. 2b shows that the changes Δr between TC_max and TC_max_add are small (-0.00 ± 0.05). For the TC_max strategy, the initial number of IASI pixels was between 7 and 14, which means using TC_max_add strategy result in a 43 ~ 186 % increase in the number of IASI pixels compared to TC_max alone. Adding more IASI pixels does not have a significant impact on the r values, indicating that maximized temporal coverage alone is the most important factor when comparing IASI to AMoN stations.

After applying a temporal coverage requirement (temporal coverage $\geq$ 80 %) to filter the overall dataset, we revisit the sensitivity of the agreement between spatial windows. The smaller spatial window now yields better agreement than the larger spatial windows (Table 4). Compared with Table 1 which has no filter for temporal coverage, the r values in Table 4 increase for all spatial windows. The correlations are clearly better for smaller spatial windows ($r = 0.74$ for 25 km versus $r = 0.48$ for 100 km). In this way, the use of a larger spatial window is indeed a tradeoff between the increasing temporal coverage versus incorporating a larger spatial gradient. The results further demonstrate that the IASI pixels far from the AMoN sites may not be representative to the AMoN site.

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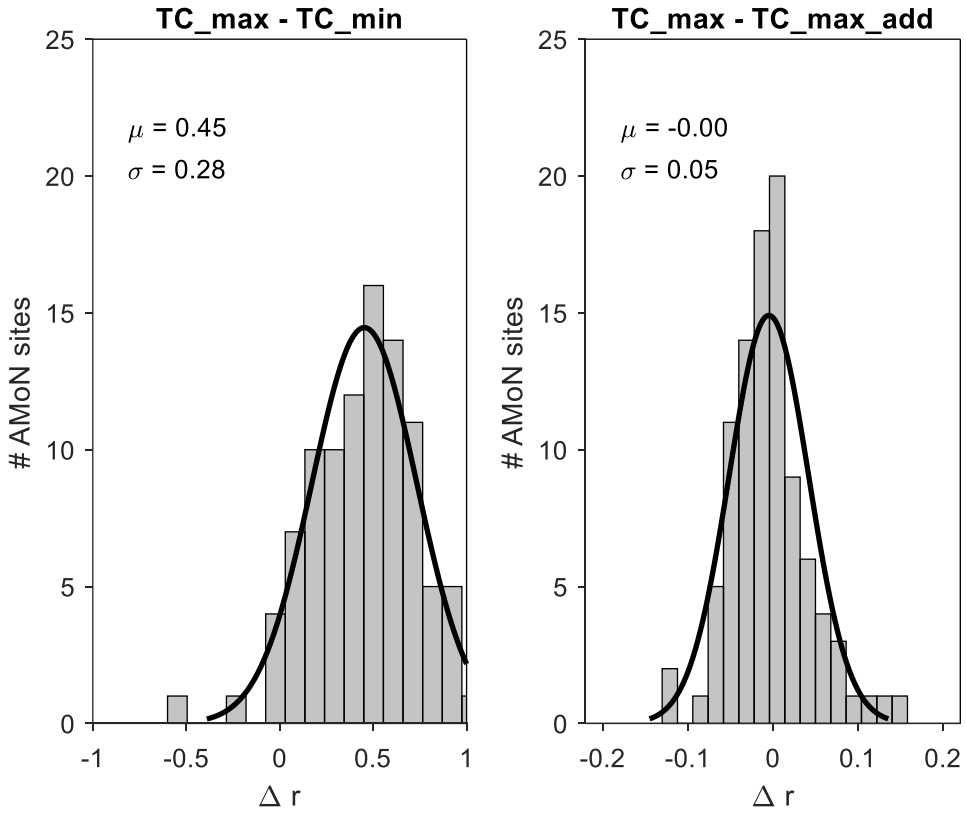

Figure 2. The change in r values for individual AMoN sites using different sampling strategies: **(a)** maximized temporal coverage (TC_max); minimized temporal coverage (TC_min) and **(b)** maximized temporal coverage & randomly adding more pixels (TC_max_add).

Table 4. AMoN & IASI comparison results for different spatial windows (temporal coverage ≥ 80 %)

Spatial window	15 km	25 km	50 km	100 km
Pearson's r	0.76	0.74	0.58	0.48
Mean # IASI pixels per pair	19	38	119	392
# AMoN & IASI pairs	105	999	3138	6899

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3.3 Sensitivity to seasons and temporal averaging

AMoN has similar numbers of measurements in spring (March, April, May), summer (June, July, August), autumn (September, October, November), and winter (December, January, February), while the mean number of IASI pixels (# IASI pixels) per pair in winter is only around half of other seasons (Fig. 3). In winter, low thermal contrasts result in low sensitivity of thermal infrared sounder, which explains the low number of IASI pixels in winter (Clarisse et al., 2010; Guo et al., 2021). The lower sensitivity of the infrared thermal sounder measurements in winter results in higher uncertainties, and thus comparisons between IASI and AMoN are especially important. When temporal coverage is at least 80%, IASI wintertime data still have good agreement with AMoN ($r = 0.61$) although the comparison are limited to only a few AMoN & IASI pairs ($N = 33$). IASI in general only provides a small number of pixels in winter, however, it indeed has the capability of reflecting surface NH_3 variations even in winter.

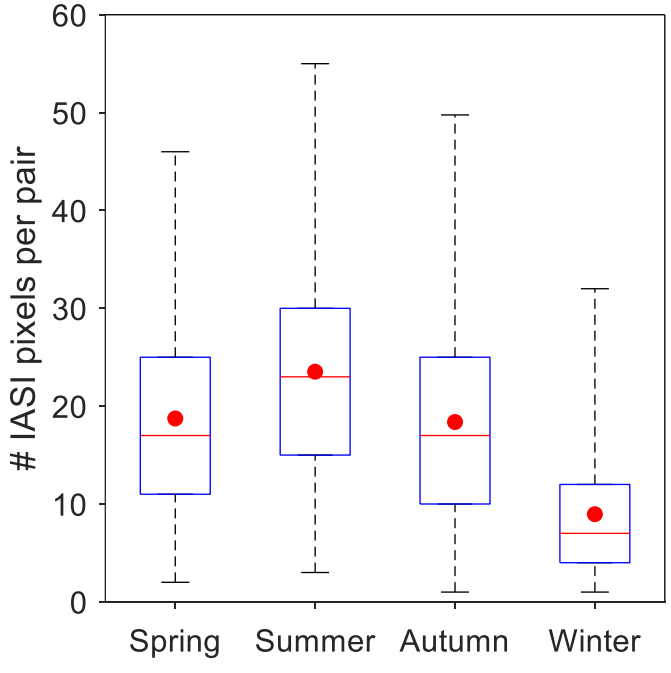

Figure 3. Boxplot of number of IASI pixels per pair for spring, summer, autumn, and winter. The boxes denote the 25th and 75th percentiles, the whiskers denote the 1st and 99th percentiles, and the red dot denotes the mean.

The results in 3.1 and 3.2 have already shown the importance of spatial window and temporal coverage. The temporal averaging, such as the tessellation oversampling and physical oversampling, is a common method to achieve a high spatial resolution map by sacrificing the temporal resolution (Sun et al., 2018; Van Damme et al., 2018; Wang et al., 2021). Here we neglect the interannual variability and calculate the multi-year averaged seasonal IASI NH_3 concentrations using the 25 km

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spatial window. By averaging the multi-year IASI data, the impacts of temporal coverage are alleviated because both temporal coverages and numbers of IASI pixels increase. Among the 101 AMoN sites with at least one full year data and available IASI v3.1r NH_3 data, 49 sites show strong agreement with IASI with $r > 0.8$, 29 sites have moderate agreement of $0.5 < r \leq 0.8$, while 23 sites do not have statistically significant agreements (Fig. 4). If taking all data into consideration, the overall r value for the CONUS is 0.69. The AMoN sites with higher NH_3 concentrations tend to show better agreements between AMoN and IASI. The median AMoN NH_3 annual mean concentrations for all sites is 0.86 $\mu g/m^3$. Most sites with no statistically significant agreements have a low NH_3 concentration (median: 0.48 $\mu g/m^3$). Currently, most AMoN sites are located in low or moderate NH_3 concentration regions with a lack of sites in the NH_3 hotspots (Wang et al., 2021) and urban areas, complicating the comparison between AMoN and IASI.

The above agreement demonstrates that IASI NH_3 column reflects the variation of the surface NH_3 concentration at seasonal resolution. For regions without any available ground measurements, IASI NH_3 observations can be used to help better understand the NH_3 variations. However, large differences exist among the relationships between IASI and AMoN NH_3 concentrations over different AMoN sites (an example of linear regression plot in Fig. 5b). Even for AMoN sites with excellent correlation ($r > 0.8$), the slopes vary a lot, ranging from $0.08 – 1.4 \times 10^{16}$ molec/cm^2 per $\mu g/m^3$. For instance, two AMoN sites in California, Joshua Tree National Park (CA 67) and Sequoia & Kings Canyon National Park (CA 83), both exhibit great seasonality agreements with IASI ($r = 0.97$ and $r = 0.99$, respectively) but the slope for CA 83 is 44 % higher than CA 67. The difference between the slopes suggests that although IASI is able to capture the general seasonality, the relationship between NH_3 column and surface NH_3 is distinctly different due to complicated topography, meteorology, and other factors at different AMoN sites.

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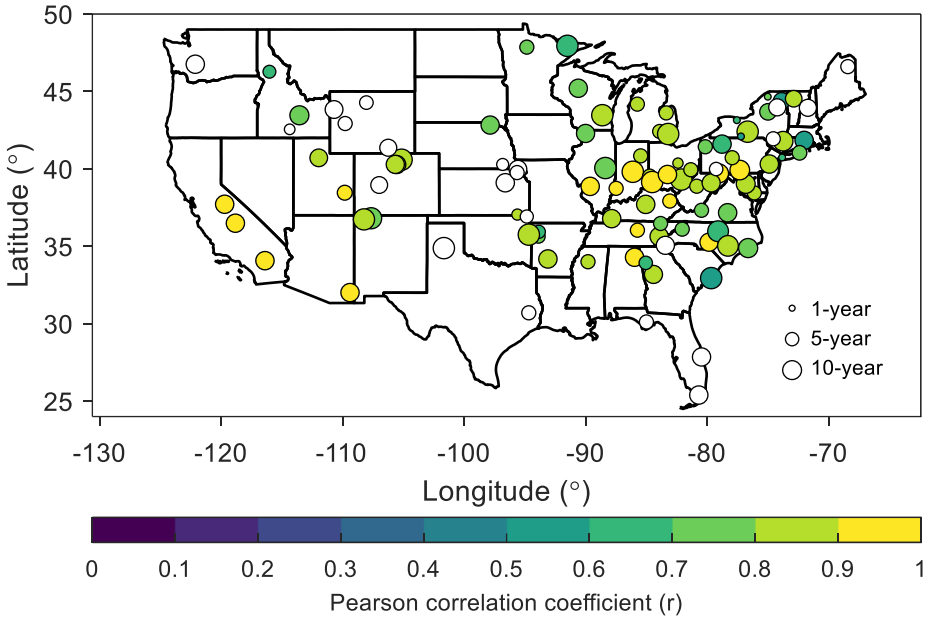

Figure 4. Multi-year averaged comparison results between AMoN sites and the IASI observations within 25 km of the AMoN sites at monthly resolution. Circles without filled color denote the AMoN sites with no statistically significant correlation with IASI. The circle sizes denote the length of AMoN data record.

4 Trend analysis

4.1 Trend in the CONUS

The methodology and comparison results in section 3 demonstrate that IASI NH_3 can be used to estimate regional NH_3 trends over the last decade. In this regard, satellite NH_3 observations will be used to augment the AMoN observed NH_3 trends in the CONUS over the last decade. We include AMoN trend analysis only for sites with full year coverage during 2008 - 2018 (N=13). Strong evidence of increasing NH_3 concentrations in the U.S. comes from both ground-based observations and satellite measurements (Van Damme et al., 2021; Warner et al., 2017; Yao and Zhang, 2016; Yao and Zhang, 2019; Yu et al., 2018). Fig. 5a shows monthly IASI and AMoN timeseries in from Indianapolis, Indiana, USA (IN 99). The strong correlation (r = 0.96) between the two measurements is shown in Fig. 5b. Although the NH_3 seasonality remain consistent from 2008 to 2018 - namely spring maxima and secondary maxima in fall with lowest values in winter - both AMoN and IASI also show increasing trends of NH_3 concentrations over the entire timeseries. AMoN shows a trend of 6.5%·yr^{-1} while IASI shows a trend of 7.0%·yr^{-1}.

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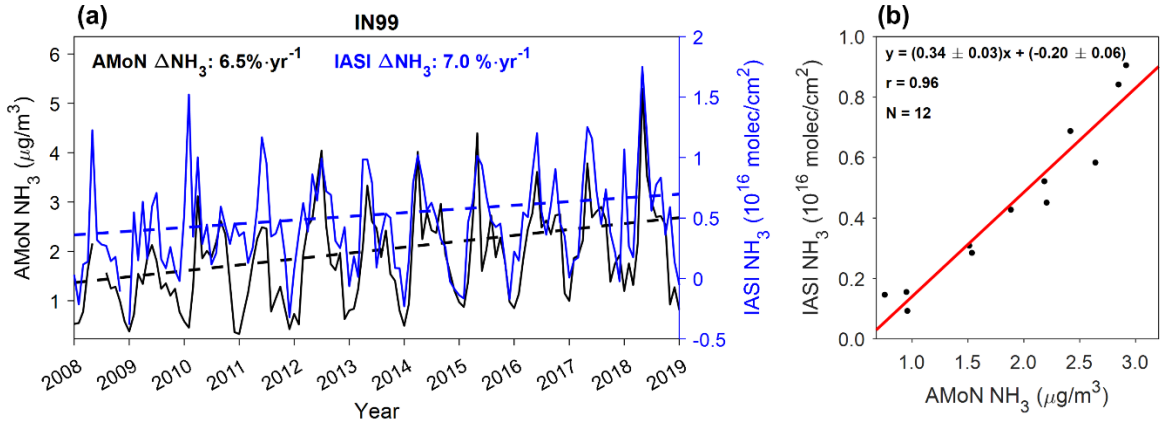

Figure 5. (a) 2008 – 2018 monthly averaged NH_3 trends for AMoN site in Indianapolis, Indiana, U.S. (IN 99) and IASI NH_3 observations within 25 km of IN 99; **(b)** seasonality correlation between AMoN and IASI NH_3 for IN 99.

A long-term trend analysis was performed using AMoN and IASI data to examine the agreement between the datasets and explore any regional differences. IASI NH_3 columns smaller than the 5th percentile (0.5×10^{15} molec/cm^2) of the 11-year NH_3 average in the CONUS region were excluded to avoid spurious trend results caused by the higher noise in these measurements. To perform the interannual trend analysis, we require each region or site to have at least one valid measurement in each season to alleviate the possible bias due to seasonal variations. Fig. 6 shows the annual percentage change for both IASI and AMoN. Most regions in the CONUS have increasing NH_3 concentrations based on the 11-year IASI observations (median: 6.8% · yr^{-1}), including eastern U.S., Midwest, and parts of the western U.S. 10 out of 13 AMoN sites have statistically significant NH_3 increases. AMoN data in general suggest similar increases (median: 6.7% · yr^{-1}). When plotting the trends of AMoN sites against the median of IASI trends within a 25 km spatial window (Fig. 7), a moderate correlation (r = 0.66) was found between IASI and AMoN NH_3 trends. IASI in general suggested a higher NH_3 increase compared to AMoN (slope: 1.26 ± 0.51) with the ratio larger than one for most sites. The absolute NH_3 change also is in correspondence with the previous study, with significant NH_3 increases across the CONUS regions, especially in the Midwest (Van Damme et al., 2021).

The spatial consistency across the datasets differs significantly. Both AMoN and IASI suggest ~ 5% · yr^{-1} NH_3 increases in the Great Lake Region, while IASI suggests a higher NH_3 increase in the eastern US compared with AMoN. The IASI trend analysis results suggest a significant NH_3 increase in the northern Great Plains, e.g., North Dakota, South Dakota, and Montana, yet there are no AMoN sites in this region. Furthermore, the trends are consistent with the NH_3 emissions increases caused by increased N fertilizer usage in the northern Great Plains (Cao et al., 2020b). McHale et al. (2021) showed that wet-precipitation NH_4^+ concentrations based on NADP observations suggested the highest increases in the Great Plains, the Rocky Mountain Region, and the Great Lake Region from 2000 to 2017, which is geographically consistent with the NH_3 trends observed by

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both AMoN and IASI. If considering the CONUS as a whole and calculating the annual mean NH_3 for the whole CONUS during 2008 – 2018 to derive the overall trend in CONUS, the IASI NH_3 change for 2008 – 2018 is (3.9 ± 2.2) % · yr^{-1} and (1.3 ± 0.8) × 10^{14} molec/cm^2·yr^{-1}, similar with the trend in the previous study (3.4 ± 0.6) % · yr^{-1} and (1.1 ± 0.4) × 10^{14} molec/cm^2·yr^{-1}) (Van Damme et al., 2021).

In terms of trend in NH_3 hotpots, which are here defined as regions where the IASI NH_3 column is larger than the 95th percentile of 11-year CONUS map (6.7 × 10^{15} molec/cm^2), the median of NH_3 trend is 4.7% · yr^{-1}, indicating that the regions of the largest emissions sources are also realizing increasing concentrations over time. Although the percent changes in the regions with the highest concentrations are smaller compared with the trend in CONUS median (8.0% · yr^{-1}), in terms of the absolute changes, the median trend of NH_3 columns over these NH_3 hotspots are higher compared with the CONUS median (3.7 × 10^{14} molec/cm^2·yr^{-1} vs. 2.8 × 10^{14} molec/cm^2·yr^{-1}). The top 10 NH_3 hotspots in CONUS regarding column-areal weighting all exhibit increasing NH_3 concentrations from 2008 to 2018 (Table 5). Within these hotspots, the central Great Plains experience the largest NH_3 increase (median: 5.0% · yr^{-1}, 4.0 × 10^{14} molec/cm^2·yr^{-1}) while the San Joaquin Valley (median: 2.0% · yr^{-1}, 1.6 × 10^{14} molec/cm^2·yr^{-1}) and Imperial County, California (median: 2.1% · yr^{-1}, 1.9 × 10^{14} molec/cm^2·yr^{-1}) have a smaller change change.

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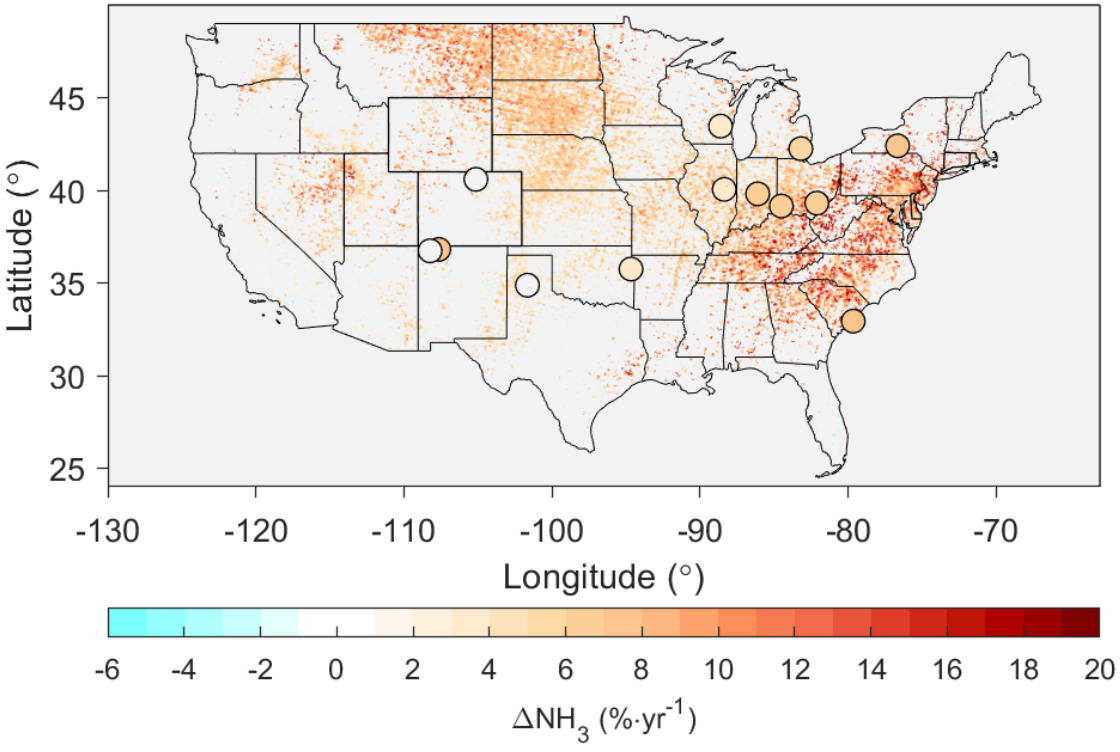

Figure 6. Trend analysis for IASI NH_3 (2008 - 2018) and AMoN NH_3 measurements in the contiguous U.S. The gray color indicates no statistically significant change. The circle size denotes the length of AMoN data record.

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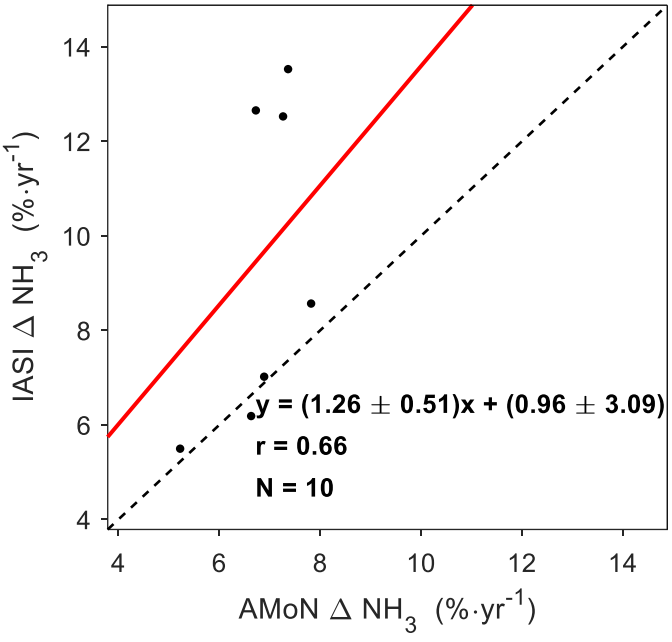

Figure 7. Comparison between AMoN and IASI NH_3 trends (25 km spatial window) for AMoN sites with available nearby IASI trend data

Table 5. 2008 – 2018 IASI observed NH_3 trend in the top 10 NH_3 hotspots (column-areal weighting) in CONUS

Hotspots	% · yr^{-1}	10^{14} molec/cm^2 · yr^{-1}
Central Great Plains	5.0	4.0
The San Joaquin Valley	2.0	1.6
North Oklahoma	3.9	2.9
Texas panhandle	3.6	2.8
Central Iowa	4.4	3.3
The Snake River Valley	3.8	3.3
Southeast Iowa	5.2	3.9

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Beadle County, South Dakota	8.3	6.0
Weld County, Colorado	3.6	2.9
Imperial County, California	2.1	1.9

To provide a detailed insight of the increasing NH_3 over the CONUS, we further perform trend analyses for different seasons (Fig. 8). In spring, significant NH_3 increases are found in the Midwest and Eastern US. In summer, NH_3 increases shift to the western US and part of the eastern US. AMoN and IASI seasonality clustering results show that the Midwest and eastern United States, dominated by fertilizer NH_3 emissions, have a broad, spring maximum of NH_3, while the western United States, dominated by volatilization of livestock waste NH_3 emissions, in contrast, show a narrower midsummer peak (Wang et al., 2021). The spatial patterns of spring and summer NH_3 trends are in agreement with the seasonality clustering results, indicating that increasing NH_3 emissions caused by agricultural activities may contribute to NH_3 concentration increase. The increasing wildfire activities in the western U.S. may also contribute to NH_3 increases (Lindaas et al., 2021a, b). In fall and winter, most regions in the U.S. do not have statistically significant IASI NH_3 trends, and a decreasing NH_3 trend is observed by IASI in the Southwest US in fall. In contrast, AMoN data suggest a notable NH_3 increase in Northeast and the Corn Belt region in winter. Again, IASI data are susceptible to low thermal contrasts in winter, which to some extent explains the disagreement between IASI and AMoN in winter as discussed in Section 3.3.

Wintertime NH_3 plays an important role in haze episodes through the formation of aerosol phase NH_4NO_3 (Shah et al., 2018; Zhai et al., 2021), and increasing NH_3 concentrations in winter may affect aerosol acidity and aerosol chemistry (Lawal et al., 2018; Zheng et al., 2020). In the past decades, NO_x and SO_2 emissions reductions have resulted in less NH_x partitioning into particle phase NH_4^+ (Shah et al., 2018), however, the partitioning alone is not able to fully explain the significant NH_3 concentration increases (Yao and Zhang, 2019; Yu et al., 2018). The change of meteorological conditions, such as increasing air temperatures may also contribute to the increasing NH_3 trends (Warner et al., 2017; Yao and Zhang, 2019). No matter the reason for increasing NH_3 concentrations across the CONUS regions, the fact that both NH_3 surface concentrations and NH_3 column concentrations are increasing during the past decade will have significant impacts on air quality and nitrogen deposition. EPA is reviewing the 2020 $PM_{2.5}$ National Ambient Air Quality Standard (NAAQS) currently set at 12.0 $\mu g \cdot m^{-3}$ and if the NAAQS is lowered, NH_3 controls will become increasingly important for meeting the standard. Additionally, Pan et al. (2021) demonstrates that NH_3 transported from Colorado significantly increased the dry NH_3 deposition the Rocky Mountain National Park. Increasing gas phase NH_3 may result in longer spatiotemporal scales for dry nitrogen deposition, leading to adverse impacts on remote regions and sensitive ecosystems (Phoenix, et al., 2006). Reduction of NH_3 emissions is critical to protect human health and the biodiversity in sensitive ecosystems (Ellis et al., 2013, Hill et al., 2019).

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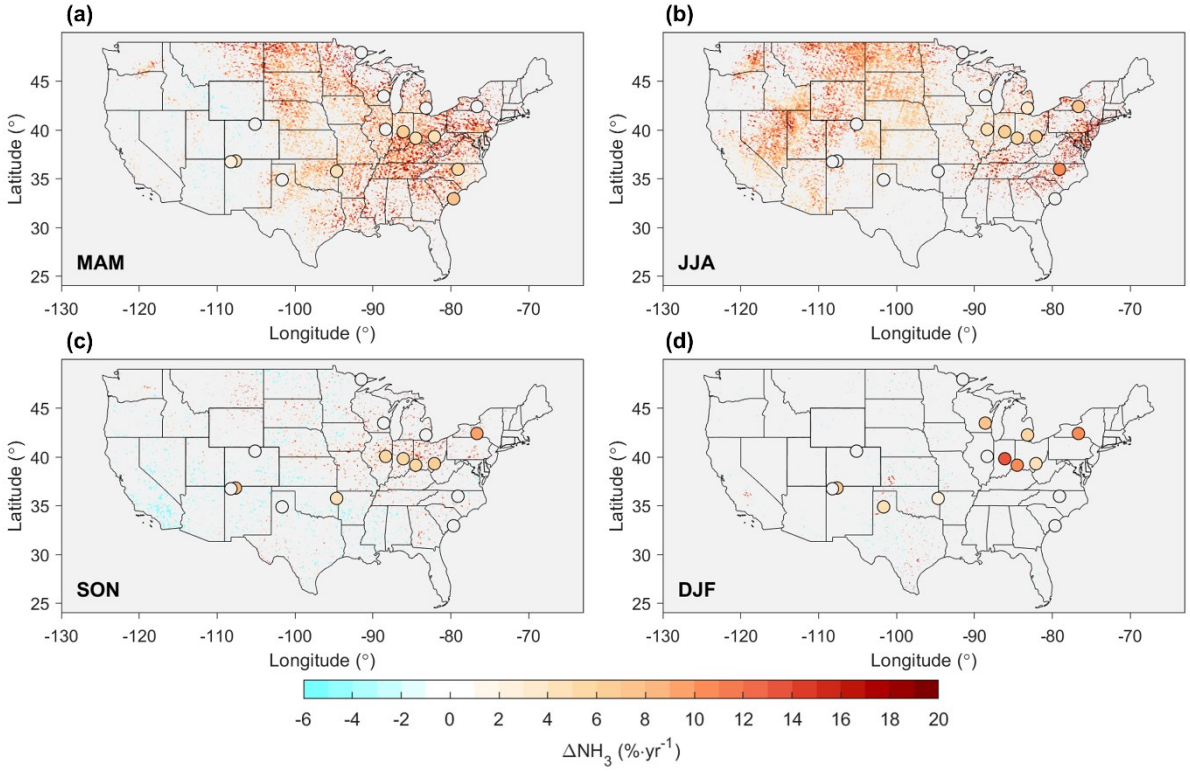

Figure 8. 2008 – 2018 NH_3 trend for different seasons based on IASI NH_3 measurements in the contiguous U.S. **(a)** spring (March, April, May); **(b)** summer (June, July, August); **(c)** autumn (September, October, November); **(d)** winter (December, January, February)

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4.2 Trend in the urbanized areas

The short lifetime of NH_3 leads to strong spatial variabilities of NH_3 concentrations, and most AMoN sites are not located in highly populated urban regions (Wang et al., 2021). Fig. 9 shows population coverage of AMoN in the CONUS region. Population data were retrieved from the Gridded Population of the World, Version 4 (GPWv4) (Center for International Earth
490 Science Information Network – Columbia University, 2018). More than half of the CONUS population is at least 100 km away from an AMoN site. As mentioned in the previous discussion of spatial windows, AMoN may best represent the NH_3 variations for regions within ~ 10 km radius, and less than 2% of CONUS population are within 10 km of an AMoN site. More urban AMoN sites are needed to represent the urban areas and better quantify NH_3 emissions from mobile sources, trends in the urban areas. Satellite observations are the only dataset that can currently be used to investigate source contributions and trends
495 in population centers (Cao et al., 2022).

We retrieved urban area data from the 2010 US Census, which includes two different types of urban areas: Urbanized Areas (UAs) of 50,000 or more people and Urban Clusters (UCs) of at least 2,500 and less than 50,000 people (U.S. Census Bureau,

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2012). The urban areas have a similar NH_3 trend compared with CONUS (8.1% · yr^{-1} vs. 8.0% · yr^{-1}), suggesting a simultaneous NH_3 increase in both urban and rural areas. The top ten most populous urbanized areas almost all exhibit significant NH_3 increases with the exception of Miami, Florida, which has a negative trend and Dallas, Texas, without any significant trend (Table 6). These ten areas in total accommodate more than seventy million population, making up more than one fifth of the total population in the CONUS. The NH_3 increase in these densely populated areas and its impact on the aerosol chemistry need to be further addressed.

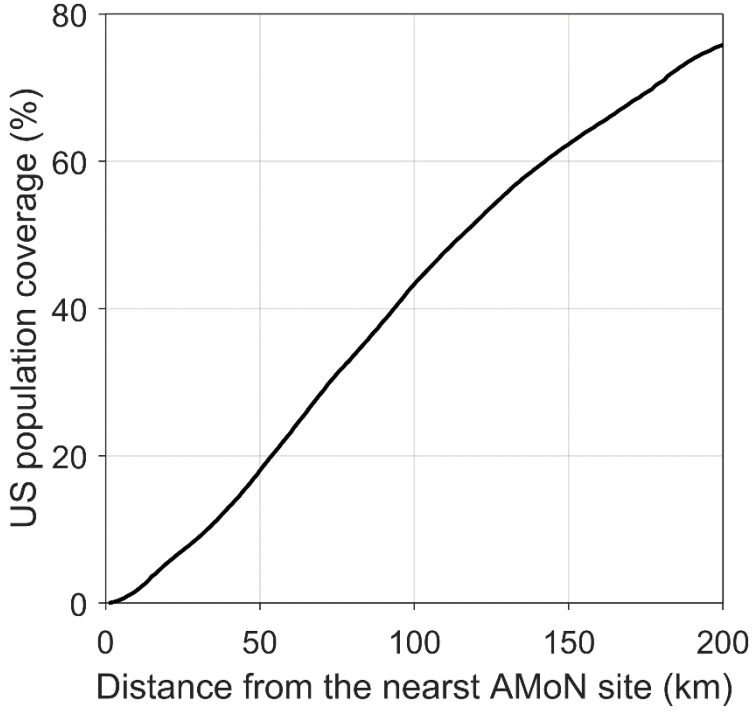

Figure 9. The population coverage of AMoN sites.

Table 6. 2008 – 2018 IASI NH_3 trend in the top 10 most populous urbanized areas

Urbanized Area	Population (million)	% · yr^{-1}	10^{14} molec/cm^2 · yr^{-1}
New York--Newark, NY--NJ--CT	18.0	10.8	2.0
Los Angeles--Long Beach--Anaheim, CA	12.0	4.3	2.1

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Chicago, IL--IN	8.6	5.2	2.5
Miami, FL	5.5	-25.2	-1.5
Philadelphia, PA--NJ--DE--MD	5.4	10.9	2.6
Dallas--Fort Worth--Arlington, TX	5.1	/	/
Houston, TX	4.9	7.9	2.0
Washington, DC--VA--MD	4.6	9.0	2.2
Atlanta, GA	4.5	9.4	2.2
Boston, MA--NH--RI	4.2	10.5	1.4

5 Implications

Under favorable conditions, IASI NH_3 columns correlate with AMoN NH_3 surface concentrations even at the 2-week scale and for low concentration regions ($r = 0.74$ when temporal coverage ≥ 80 %). IASI measurements' temporal coverage of AMoN's 2-week sampling period dominates the agreement presumably because of the larger day-to-day variability of NH_3. The agreement demonstrates the strong potential for using IASI NH_3 columns to bridge the spatial gaps of the AMoN network. The global coverage of satellite measurements enables IASI NH_3 product to serve as an alternative dataset in countries and regions that do not have any NH_3 monitoring networks, particularly in developing countries. For example, India is the second most populated country in the world with a sixth of the world's population, and recent study has shown that the unique role of NH_3 in forming massive chloride aerosols (up to 40 $\mu g/m^3$) in India (Gunthe et al., 2021). However, there are currently no long-term NH_3 ground monitoring networks in India, impeding the efforts to estimate and control NH_3 emissions (Beale et al., 2022). IASI's low sensitivity to wintertime NH_3 shows the value of the more sensitive AMoN sites. Extra attention is needed when using IASI data in such circumstances.

The increasing NH_3 in the CONUS (median: 6.8%·yr^{-1}, 2.8 × 10^{14} molec/cm^2·yr^{-1}), including the hotspots region (median: 4.7%·yr^{-1}, 3.7 × 10^{14} molec/cm^2·yr^{-1}), highlights the more important role of NH_3 in $PM_{2.5}$ formation and nitrogen deposition in the future. AMoN suggests a similar NH_3 increase (6.7%·yr^{-1}) as well as a similar spatial pattern with IASI. Both IASI and

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AMoN show largest NH_3 increase in the Midwest and eastern U.S., with a moderate agreement for AMoN sites (r = 0.66). More co-located measurements of $PM_{2.5}$ mass and NH_3 concentrations would help assess the impact increasing trends of NH_3 will have on human health. The integrated satellite and ground-based measurements are already playing a role in our understanding of under-represented NH_3 emissions sources in the inventories. NH_3 already dominates the reactive nitrogen deposition in the majority areas in the U.S., with the continuing efforts on NO_x emission reductions, NH_3 is expected to become the key species for nitrogen deposition (Li et al., 2016) and poses adverse impacts on the nearby ecosystem regions, e.g., the National Parks (Benedict et al., 2013; Pan et al., 2021). The changing partitioning of NH_x between NH_3 and NH_4^+ is likely to impact the lifetime of NH_x due to differences between the removal velocity of gas phase NH_3 via dry deposition and particle phase NH_4^+ wet deposition. The trends vary in different seasons, with NH_3 increases mainly in spring in the Midwest and eastern U.S. (cropland dominated) while in summer in the western U.S. (feedlot dominated), suggesting the impacts from agricultural activities and the necessity of developing regionally-specific emission control strategies.

Because of the scarcity of the ground monitoring sites in the urban areas, satellite NH_3 measurements are extremely valuable to characterize NH_3 magnitude, seasonality, and trend in densely populated areas. Satellite observations suggests NH_3 increases across the U.S. urban areas (median: 8.1%). New York—Newark, NY--NJ—CT alone has more than eighteen million population, experiencing an 10.8 % · yr^{-1} NH_3 increase. Measurements from satellites will help inform where ground based NH_3 samplers could be located to better understand local air quality in overburdened communities that lack resources for continuous monitors. In addition, NH_3 sources in the urban areas and the related atmospheric chemistry are both poorly understood (Gu et al., 2022; Sun et al., 2017) and could be constrained by satellite NH_3 observations (Cao et al., 2022). However, satellite observations alone are not able to answer all questions under the complex urban atmospheric conditions. For instance, gas phase NH_3 and HNO_3 can nucleate directly to form NH_4NO_3 particles in cold atmospheric conditions and is likely to result in rapid growth of new atmospheric particles in winter in urban areas (Wang et al., 2020). To provide accurate and fine spatial scale NH_3 observations in the urban areas, more routine ground monitoring sites are needed both in urban areas and high NH_3 emission source regions.

6 Data availability

The AMoN data were downloaded from the National Atmospheric Deposition Program/National Trends Network (NADP/NTN): https://nadp.slh.wisc.edu/networks/ammonia-monitoring-network/. The authors acknowledge the AERIS data infrastructure (https://www.aeris-data.fr) for providing access to the IASI Level 2 NH3 data used in this study. Population data were retrieved from Center for International Earth Science Information Network, Columbia University: https://sedac.ciesin.columbia.edu/data/collection/gpw-v4/. The urban areas data are downloaded from the U.S. Census Bureau: https://www.census.gov/geographies/mapping-files.html.

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Author contributions

MAZ and RW designed the research; RW led the analysis; KS, DP, and XG contributed to data analysis; LC, MV, LP, and CC helped with the usage of IASI data; MP helped with the usage of AMoN data; and RW wrote the paper with contributions from all co-authors.

Competing interests

Competing interests. The contact author has declared that none of the authors has any competing interests.

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