

First ozone reanalysis on Mars using SPICAM data

James A. Holmes, Stephen R Lewis, Manish R. Patel, Franck Lefèvre

To cite this version:

James A. Holmes, Stephen R Lewis, Manish R. Patel, Franck Lefèvre. First ozone reanalysis on Mars using SPICAM data. Scientific Workshop: "From Mars Express to ExoMars", ESA, Feb 2018, Madrid, Spain. insu-04392364

HAL Id: insu-04392364 <https://insu.hal.science/insu-04392364v1>

Submitted on 13 Jan 2024

HAL is a multi-disciplinary open access archive for the deposit and dissemination of scientific research documents, whether they are published or not. The documents may come from teaching and research institutions in France or abroad, or from public or private research centers.

L'archive ouverte pluridisciplinaire **HAL**, est destinée au dépôt et à la diffusion de documents scientifiques de niveau recherche, publiés ou non, émanant des établissements d'enseignement et de recherche français ou étrangers, des laboratoires publics ou privés.

First ozone reanalysis on Mars using SPICAM data

J. A. Holmes¹, S. R. Lewis¹, M. R. Patel^{1,2} and F. Lefèvre³
School of Physical Sciences, Open University, UK, ² Space Science and Technology De¹ *partment, STFC, RAL, UK, ³ LATMOS, CNRS/Université Pierre et Marie Curie, France*

Introduction: To further our understanding of important photochemical processes in the Martian atmosphere, a synthesis can be used to investigate the temporal and spatial agreement between model and observations and determine any possible causes of identified differences. In this study [1], we have assimilated, for the first time, total ozone into a Mars Global Circulation model (GCM) to study the ozone cycle.

Model Setup & Assimilation: We use the UK version of the Laboratoire de Météorologie Dynamique (LMD) Mars GCM. This 4D-model uses the physical parameterisations [2] and LMD photochemical module [3] shared with a recent version of the LMD Mars GCM coupled to a UK-only spectral dynamical core and semi-Lagrangian advection scheme [4].It has been developed in a collaboration between the LMD, the Open University, the University of Oxford and the Instituto de Astrofisica de Andalucia. The model was run at T31 resolution in the horizontal, corresponding to a resolution of 5° latitude by 5° longitude, with 32 vertical levels in the range 0–105 km. The Mars GCM includes the latest sub-models to provide the most realistic modelling of the planetary boundary layer and water and dust cycles [5,6,7]. The photochemical module provides multiple photolytic and chemical reactions with up-to-date reaction rates between 16 advected species including carbon dioxide, water vapour and ozone. It also includes heterogeneous processes removing odd hydrogen radicals, a process which has been shown to improve the agreement between models and observations [8].

The assimilation is performed using a form of the Analysis Correction scheme [9] converted to martian con- ditions, and has been shown in the past to be a computation- ally inexpensive and robust method [10]. Using this methodology, observations of short-lived (and long-lived) species can be supplemented by knowledge of the transport and atmospheric chemistry from a Mars GCM.

SPICAM retrievals: The retrievals assimilated in this study (displayed in Figure 1a) are from the SPICAM instrument on Mars Express presented in [11]. Maximum total ozone values are seen around northern hemisphere spring equinox at northern polar latitudes, as a result of polar night conditions in the recent past and a lack of odd hydrogen species. The SPICAM instrument is able to observe large ozone abundance before water vapour and associated odd hydrogen species are transported into the northern polar regions resulting in destruction of ozone.

Figure 1: SPICAM nadir retrievals for $L_s = 341^\circ$ MY 26 to $L_s = 69^\circ$ MY 28 as a function of latitude and solar longitude displaying a) the total ozone abundance and b) number of retrievals. The number of SPICAM retrievals are binned every 5° L_S and 5° latitude.

For the first assimilation of ozone for Mars, the dynamically determined wind field from the Mars GCM is used to transport ozone and all other chemical species, but future investigations could potentially also assimilate thermal profiles alongside ozone retrievals. This highlights another benefit of data assimilation, in that data sparse regions are potentially influenced by SPICAM retrievals of total ozone being advected into the local unobserved region.

Results of the ozone reanalysis: To investigate the effect of total ozone assimilation on the seasonal cycle, we ran a 'control' simulation with no assimilation and a second simulation in which total ozone retrievals from SPICAM were assimilated (hereafter the SPICAM ozone rea-

nalysis). Differences in the modelled and observed ozone cycle (Figure 2b) are efficiently identified by the assimilation process, with the largest total ozone percentage differences identified between 45° S-10°S from $L_S = 135-180^{\circ}$ and at northern polar (60°N–90°N) latitudes from $L_S = 150-195$ °, at the onset of northern polar winter. An asymmetry in the magnitude of total ozone percentage differences in northern fall $(L_S = 150-195^o)$ and early northern spring $(L_S = 0-30^o)$ at northern polar latitudes is also identified. Through investigation of the water cycle, underestimated amounts of total ozone in the control simulation were found to be due to excessive northward transport of water vapour west of the Tharsis region and over Arabia Terra throughout northern polar winter resulting in increased increments in total ozone necessary in the SPICAM ozone reanalysis. Neither modelling biases in water vapour or heterogeneous pro- cesses on water ice clouds can explain outstanding differences in total ozone between the SPICAM ozone reanalysis and the control simulation at northern polar latitudes from L_S = 150–195°. In this region in particular, investigation of the vertical profiles of water vapour/ice and ozone are needed to try and explain the difference in total ozone between the control simulation and SPICAM ozone reanalysis, and determine whether it is a further unidentified process which is not adequately represented.

Figure 2: (a) Assimilated zonally averaged ozone column at 12 p.m. everywhere and (b) the percentage difference between the SPICAM ozone reanalysis and control simulation for MY 27 and start of MY 28. Shading indicates regions where there are no nearby SPICAM retrievals at the corresponding time and space.

The assimilation of SPICAM retrievals can also provide potential information on the polar maximum, with chemical changes as a result of the assimilation process shown to affect the southern polar high altitude ozone layer. For a more realistic study of the polar dynamics, assimilation of total ozone, and ideally water vapour too, would provide a more accurate ozone spatial distribution.

In the future, observations from NOMAD [12] on the ExoMars TGO spacecraft, will provide a more comprehensive dataset to assimilate, with the SPICAM observational dataset extremely useful also for cross-validation purposes. Assimilation of vertical profiles of ozone and water vapour/ice alongside column retrievals, such as will be possible from the NOMAD instrument on the ExoMars TGO, will also be of great use to further study the origin of differences in total ozone between the SPICAM ozone reanalysis and global circulation models.

References:

- [1] Holmes, J. A. et al., *Icarus 302*, 308-318, 2018.
- [2] Forget, F. et al., *JGR 104,* 24155-24175, 1999.
- [3] Lefèvre, F. et al., *JGR 109,* E07004, 2004.
- [4] Newman, C. E. et al., *JGR 107,* 5123, 2002.
- [5] Colaïtis, A. et al., *JGR (Planets) 118,* 1468- 1487, 2013.
- [6] Navarro, T. et al., *JGR (Planets) 119,* 1479- 1495, 2014.
- [7] Madeleine, J.-B. et al., *JGR (Planets) 116,* E11010, 2011.
- [8] Lefèvre, F. et al., *Nature 454,* 971-975, 2008.
- [9] Lorenc, A. C. et al., *QJRMS 117,* 59-89, 1991.
- [10] Lewis, S. R. et al., *Icarus 192*, 327-347, 2007.
- [11] Perrier, S. et al., *JGR 111,* E09S06, 2006.
- [12] Patel, M. R. et al., *Appl. Opt. 56 (10)*, 2771-2782, 2017.

Acknowledgements: All authors acknowledge support as part of the project UPWARDS-633127, funded by the European Union Horizon 2020 Programme. JAH, SRL and MRP also acknowledge the support of the UK Space Agency/STFC. We are grateful for an ongoing collaboration with François Forget and coworkers at LMD.