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1 Tetrapod tracks from the Middle/Late Triassic of the Ourika and Yagour basins

2 (Morocco): new ichnotaxonomical data based on the Biron collection

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Abstract: Palaeontological prospecting organized in the Ourika and Yagour basins 14 (Marrakech High Atlas, Morocco) in the 1980's, have yielded abundant and exquisitely 15 16 preserved tetrapod tracks, some of them preserving thinly detailed skin traces (Collection Biron, Cadi Ayyad, Marrakech, Morocco). Based on this material, diverse morphologies of 17 18 tracks were described and six new ichnotaxa were erected. However, these trackways have 19 never been reinvestigated in detail from an ichnotaxonomic point of view. Here, the Biron 20 collection is reevaluated using modern photogrammetric imaging techniques. The fine-21 grained sandstone slabs bear trackways ascribed to Apatopus lineatus, Isochirotherium isp. and Rhynchosauroides isp. whose trackmakers are possible phytosaurs, crocodylian-stem 22 archosaurs or non-crown-group archosauriforms and lepidosauromorphs/archosauromorphs, 23 respectively. "Chirotherium atlensis", "Enigmatopus atlensis", "Hyloidichnus triasicus", 24 "Prochirotherium triasicum", "Pseudochirotherium oukaimedensis" and "Rectilinetopus 25 ourikensis" are considered as invalid. The palaeoichnological assemblage includes a mixture 26 27 of Early to the Middle Triassic (Isochirotherium) and Late Triassic (Apatopus lineatus) tracks. This result corroborates the biostratigraphic data recently obtain from the base of the 28 Oukaïmeden Sandstone Formation, suggesting a Middle-Late Triassic age. The living 29 environment was a semi-arid alluvial plain with ponds and ephemeral rivers. 30

31 Keywords: tetrapod tracks; Mesozoic; Ladinian; Carnian; Marrakech High Atlas;

32 photogrammetry

33

34 **1. Introduction**

In Morocco, significant ichnofossils including abundant and diverse vertebrate and 35 invertebrate tracks were reported from Triassic deposits of the Argana Basin (Klein et al., 36 2010, 2011; Tourani et al., 2010; Lagnaoui et al., 2012, 2016; Zouheir et al., 2020), the Sidi 37 Saïd Maachou Basin (Hminna et al., 2013, 2020, 2021) and the Yagour Plateau (Biron and 38 39 Dutuit, 1981; Biron, 1982). They constitute a key for the reconstruction of continental early Mesozoic palaeoecosystems (Biron and Dutuit, 1981; Klein et al., 2010, 2011; Tourani et al., 40 41 2010; Lagnaoui et al., 2012, 2016; Zouheir et al., 2020; Hminna et al., 2013, 2020, 2021). Whereas tracks from the Triassic deposits of the Argana Basin were abundantly studied 42 43 (Biron and Dutuit, 1981; Klein et al., 2010, 2011; Tourani et al., 2010; Lagnaoui et al., 2012, 2016; Zouheir et al., 2020), those from the Ourika and Yagour basins clearly appear to be 44 underestimated. In 1981, Pierre-Eymard Biron discovered several tracksites (Biron and 45 Dutuit, 1981; Biron, 1982). This material was described by Biron and Dutuit (1981) who 46 dated it as Late Triassic. In this study, they described two slabs bearing exquisitely preserved 47 tracks that they attributed to Apatopus lineatus and a new species of Chirotherium: "C. 48 atlensis". Then, based on isolated pes or manus tracks, they briefly described five additional 49 new ichnotaxa: "Enigmatopus atlensis", "Hyloidichnus triasicus", "Prochirotherium 50 triasicum", "Pseudochirotherium oukaimedensis" and "Rectilinetopus ourikensis". Except for 51 a tentative discussion by Klein and Lucas (2013, 2021) of some of these ichnospecies, the 52 tracks from the Ourika and Yagour basins have never been reinvestigated in detail. Here, we 53 revised Biron's material housed at the Cadi Ayyad University (Marrakech). We describe it 54 using photogrammetric imaging techniques, update the ichnotaxonomy of the tracks and 55 56 discuss their age.

57

58 2. Geographical and geological setting

The study area is located in the Marrakech High Atlas (Fig. 1A), 60 to 80 km 60 southeast of Marrakech. At the south of the Carboniferous Ourika-Zat anticline (Fig. 1B), the 61 Permian and Triassic outcrops overlie the deformed Variscan basement, represented by 62 Precambrian volcanic rocks, Cambrian series, granitoids and metamorphic rocks, or 63 Carboniferous deposits. The Permian and Triassic rift area is related to the NW-SE-oriented 64 extension, which generated N-S trending normal faults (Van Houten, 1977; El Arabi et al., 65 2003; Qarbous et al., 2003; Baudon et al., 2009) and reactivated pre-existing NE-SW to ENE-66 WSW Variscan structures (Fig. 1B). These dip-slip and strike-slip faults are reactivated 67 (Baudon et al., 2009; El Arabi, 1988; El Arabi et al., 2003; Laville and Petit, 1984; Laville 68 and Piqué, 1991) and controlled the Triassic sedimentation (Benaouiss et al., 2023). The 69 Permian and Triassic sedimentary successions are represented by elongated and narrow 70 tectonic belts with an ENE-WSW orientation (Fig. 1B), up to 17 km wide and up to 35 km 71 72 long. These Permian and Triassic deposits are subdivided into six formations (fms), noted F1 73 to F6 (Biron, 1982; Fig. 1C), corresponding to the Permian Anrar Conglomerate (F1), dated 74 by radioactive and zircon method (Van Houten, 1977 and Perez et al., 2019) and the Cham-El-Houa Siltstone (F2), located below the Permian-Triassic unconformity, is attributed to the 75 Permian age by correlation with the Argana t2 Unit (e.g. Benaouiss et al., 2023). The F3 76 (Timalizène Conglomerate), F4 (Anouffig Siltstone), F5 (Oukaïmeden Sandstone), and F6 77 (Tafilalt Siltstone) fms of Biron (1982) correspond respectively to the Basal Conglomerate, 78 Ramuntcho Siltstone, the Oukaïmeden Sandstone and Upper Siltstone fms of Mattis (1977). 79 The F3 Formation (Fm) is attributed to Early Triassic (e.g. Benaouiss et al., 2023) by 80 correlation with the t3 Unit in the Argana Basin, dated by Tourani et al. (2010) with tetrapod 81 tracks. Based on pollen analyses, the lower part of F4 Fm is middle Anisian age (El Arabi et 82 al., 2006) and the F5 Fm is dated middle Carnian at its lower (Biron and Courtinat, 1982) and 83 upper parts (Cousminer and Manspeizer, 1976) (Fig. 1C). The base of the Oukaïmeden 84 Sandstone Fm was recently dated by pollen grains and conchostracans as Ladinian-early 85 Carnian (Benaouiss et al., 2018). This part of F5 Fm represents the equivalent of the lower 86 87 part of the t5 Unit in the Argana Basin which is dated as Julian-Tuvalian (mid-Carnian) by Tourani et al. (2023). The F6 Fm is attributed to the Norian, by correlation to the t8 Unit in 88 89 the Argana Basin which is dated by conchostracans according to Weems and Lucas (2015). The Triassic series is overlain by Early Jurassic basalts, representing a good stratigraphic 90 91 marker across Morocco (Bertrand et al., 1982; Blackburn et al., 2013; De Pachtère, 1983; Marzoli et al., 2004). 92

Two main Triassic outcrops are of interest to this study (Fig. 1B); (1) those of the 93 Ourika Basin, which is limited at the north by the Sidi Ali-Ou-Fars fault and at the south by 94 the Oukaïmeden fault (Fig. 1B), (2) those of the Yagour Basin, which is limited by the Sidi 95 Ali-Ou-Fars and Tafrawt faults at the north and the Meltsene fault at the south (Fig. 1B). 96 Biron and Dutuit (1981) reported three tracksites on the Ourika and Yagour basins, at the East 97 of Aghbalou, at the North-East of Anammer and at the North-West of the Yagour Plateau 98 (Star symbol in Fig. 1B). The historical tracksites are located at the base of the Oukaïmeden 99 100 Sandstone Fm (Fig. 1B and C). This formation is up to 600 m thick and outcrops mainly as cliffs, accessible along streams and tracks, and is mainly characterized by channel sandstone 101 bodies alternating with lenticular fluvial bar sandstones and overbank mudstone deposits 102 103 (Benaouiss et al., 1996; Fabuel-Perez et al., 2009). The Oukaïmeden Sandstone Fm is subdivided into three units (Benaouiss et al., 1996; Fabuel-Perez et al., 2009), called lower-, 104 middle- and upper Okaïmeden units. The lower unit, which yields the fossil tracks, is 30 to 105 110 m thick. It is richer, than the middle and upper units, in red clay-stones and siltstones, and 106 107 was deposited in an ephemeral fluvial system under dry climate. The middle unit is 25 to 80 m thick, and corresponds to a distal braided system associated with more humid conditions 108 (Fabuel-Perez et al., 2009). The upper unit, which is 30 to 80 m thick, corresponds to the 109 comeback of ephemeral conditions with the installation of aeolian dunes corresponding to arid 110 conditions (Benaouiss et al., 1996; Fabuel-Perez et al., 2009). 111

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113 **3. Material and methods**

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The palaeontological collection of the department of Geology (Cadi Ayyad 115 University) includes two large fine-grained sandstone slabs bearing tetrapod tracks, described 116 and figured in Biron and Dutuit (1981) (Figs. 2 and 3). These slabs called "Dalle 1" and 117 118 "Dalle 2" in Biron and Dutuit (1981), bear currently the collection numbers UCA.82 and UCA.83, respectively. The slab UCA.82 (Fig. 2), was collected at the base of the Yagour 119 120 Plateau whereas the geographical origin of the slab UCA.83 (Fig. 3) was not indicated in Biron and Dutuit (1981). The slab UCA.82 (63 cm long and 58 cm wide) bears eight large and 121 exquisitely preserved tracks briefly described by Biron and Dutuit (1981) for which they 122 erected the new ichnospecies "Chirotherium atlensis". The slab UCA.82 also bears numerous 123 124 (around one hundred) small lacertoid tracks (Rhynchosauroides) undescribed by Biron and

Dutuit (1981) but briefly mentioned in Biron (1982). The slab UCA.83 (44 cm long and 41 125 cm wide) bears eight large tracks briefly described by Biron and Dutuit (1981) who ascribed 126 them to Apatopus lineatus (Fig. 3). This slab also bears four isolated and partial small 127 lacertoid tracks ascribed to Rhynchosauroides (Fig. 4). On both slabs, tetrapod ichnofossils 128 are preserved as convex hyporeliefs. The tracks that Biron and Dutuit (1981) described as 129 "Enigmatopus atlensis", "Hyloidichnus triasicus", "Prochirotherium triasicum", and 130 "Rectilinetopus ourikensis" (Fig. 5) were preserved on large isolated blocs left in place on the 131 132 field (unfortunately not precisely located in Biron and Dutuit, 1981). The descriptive terminology and biometric parameters used in this study follow those 133 of Haubold (1971) and Leonardi (1987). We used the following standard abbreviations: ML, 134 manus track length; MW, manus track width; ML/MW, length to width ratio of manus track; 135 PL., pes track length; PW, pes track width; and PL/PW, length to width ratio of pes track. 136 3D photogrammetric reconstructions were produced with the software Agisoft 137 Metashape Professional 2.0. Pictures of the surfaces were taken with a Nikon D5200 camera 138 139 coupled with an AF-S Nikkor 18–105 128 mm f/3.5–5.6G ED lens. The same software was used to produce photogrammetric digital elevation models and contours. 140 141

142 **4. Systematic palaeoichnology**

143

144 Isochirotherium isp.

145 Fig. 2A–F

146 *Material*. Eight tracks preserved as convex hyporeliefs on slab UCA.82, Cadi Ayyad

147 University (Marrakech).

148 *Description.* The slab UCA.82 bears eight semi-plantigrade and quadrupedal tracks, including

149 five pes tracks and three manus imprints (including three pes/manus sets; Fig. 2A–C). Two

150 pes/manus sets are part of a trackway (Fig. 2A–C). Pes and manus tracks from the same set

151 form a small divergent angle of 12 to 15°. Pes and manus track paces are 34.5 cm and 36.0 cm

152 long, respectively. The pes and manus tracks are strongly oriented outward.

Pes imprints are pentadactyl, longer than wide, 16–17 cm long and 10.7–11.8 cm wide
(Fig. 2D–F). The PL/PW ratio varies from 1.4 to 1.5. Digit imprints I-V are elongated, robust,

straight and apically pointed. They form a group oriented outward, whereas imprint of digit V
is curved rearward. The position of the impression of digit V is more proximal than those of
digits I-IV. Traces of digit V show a well-imprinted, oval and longer than wide proximal pad
that is in line with impression of digit IV (Fig. 2D–F). The impressions of the digit II and III
(that are similar in length) are the longest. The impressions of digits I and IV are shorter than
those of II and III. The divarication angle of digits I-IV and digits I-V varies from 20 to 35°

161 and 52 to 54° , respectively.

Manus imprints are pentadactyl, quite symmetrical, smaller than pes tracks (PL/ML= 162 2.6), wider than long (ML=6.5 cm, MW=7.3 cm, ML/MW ratio = 0.9; Fig. 2D–F). The digit 163 imprints are moderately elongated, straight and apically pointed. The impression of the digit 164 III is the longest. The impression of digit V is shallower than others. The impression of the 165 digit I is the shortest. The divarication angle of digits I-V is up to 120°. Proximal pads of 166 digits I–V form a posteriorly concave margin. Manus tracks sometimes show a poorly marked 167 sole. On pes and manus tracks, phalangeal nodes, skin detailed-structures and large triangular 168 claws are very well-preserved (Fig. 2D-F). The slab UCA.82 bears thin, elongated and 169 170 straight to curved ridges that could be possible tail marks (Fig. 2C).

Remarks. Biron and Dutuit (1981) identified these tracks as a new species, they named 171 "Chirotherium atlensis". The gross morphologies of Chirotherium and Isochirotheriumare 172 very similar. However, Isochirotherium differs from Chirotherium by the size of pedal digits 173 II and III that are quite similar in length (Klein et al., 2011, 2013). In *Chirotherium* the digit 174 III is the longest. For this reason, we attribute the tracks from the slab UCA.82 to 175 Isochirotherium and consider the ichnospecies "Chirotherium atlensis" as an invalid name. 176 Chirotherian tracks (Brachychirotherium, Parachirotherium, Protochirotherium, and 177 Synaptichnium), were abundantly described from the Triassic of Morocco: in the Sidi Saïd 178 179 Maachou (Hminna et al., 2013, 2021) and Argana basins (Klein et al., 2010, 2011; Tourani et al., 2010; Lagnaoui et al., 2012, 2016; Zouheir et al., 2020). They all differ from the tracks of 180 the slab UCA.82: Brachychirotherium (Fig. 6A) differs by broad pes that has short and broad 181 digits and the oval basal pad of digit V that is in a postero-lateral position relatively to the 182 anterior digit group I–IV; Parachirotherium (Fig. 6C) differs by the pentadactyl, functionally 183 tridactyl pes traks with a wide divarication angle I-V, reduced imprints of digits II and IV, and 184 a posteriorly shifted digit I. Lagnaoui et al. (2012, 2016) mentioned the occurrence of 185 *Parachirotherium* in the Late Triassic of the Argana Basin. Nevertheless, according to Klein 186 and Lucas (2021) Parachirotherium should be regarded as a junior synonym of Chirotherium; 187

Protochirotherium (emended in Klein and Lucas, 2021 after Fichter and Kunz, 2004) differs 188 from the tracks of the slab UCA.82 by larger manus tracks compared to the size of pes track 189 and digit V of pes with strongly elongated pad proximally (Fig. 6D); Synaptichnium differs by 190 a more slender overall shape, a less massive pedal digit V, and a pedal digit IV that is 191 subequal with or even longer than digit III (Fig. 6E). In Morocco, Tourani et al. (2010) and 192 Klein et al., (2011) reported Isochirotherium cf. gierlinskiiand Isochirotherium courelifrom 193 the Triassic of the Argana Basin. The tracks described as Isochirotherium cf. gierlinskii in 194 195 Tourani et al. (2010) were tentatively renamed Protochirotherium by Klein et al. (2013). The tracks of the slab UCA.82 has no distinctive character supporting the creation of a new 196 species nor can it be confidently assigned to any of the known species of Isochirotherium. 197 198 According to Klein and Lucas (2021), the valid Isochirotherium ichnospecies are I. coltoni, I. coureli, I. delicatum, I. herculis, I. lomasi, I. marshalli and I. soergeli. I. coureli differs from 199 200 the tracks of the slab UCA.82 in showing higher PL/PW ratio (up to 1.9). I. soergeli differs in showing pes tracks with as long as wide to slightly longer than wide group I-IV. I. delicatum 201 202 shows more elongated pes tracks with long traces of digits II-III as well as very short traces of digits I and IV. I. marshalli differs from the tracks of the slab UCA.82 in showing a bigger 203 204 ovoid metatarsal-phalangeal pad on the digit V of pes tracks. *I. herculis* shows strongly bigger pes (PL can be more than 30 cm) and a group I-IV which is wider than long. I. coltoni and I. 205 soergeli are clearly smaller than the tracks of the slab UCA.82. Similar to the tracks we 206 describe here, the pes track of *I. lomasi* shows a longer than wide group I-IV. However, *I.* 207 lomasi displays a higher PL/PW ratio (more than 1.6) and a shorter trace of digit IV on pes 208 209 track.

210

- 211 Apatopus lineatus Bock, 1952
- 212 Fig. 3A–G

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214 *Material*. Eight tracks preserved as convex hyporeliefs on slab UCA.83, Cadi Ayyad

215 University (Marrakech).

- 216 *Description.* The slab UCA.83 bears eight plantigrade to semi-plantigrade and quadrupedal
- tracks, with four pes tracks and three manus imprints (including three pes/manus sets; Fig.
- 218 3A–C). Two pes/manus sets are part of a trackway (Fig. 3C). Pes and manus tracks from the

same set form divergent angle of 22 to 40°. Pes and manus track paces are 22.5 cm and 19.0 219

220 cm long, respectively. The pes tracks are oriented outward whereas manus tracks are oriented forward to inward. 221

Pes imprints are pentadactyl, longer than wide, 10.0–10.5 cm long and 4.7–6.0 cm wide (Fig. 222

3D–G). The PL/PW ratio varies from 1.7 to 2.1. Digit imprints are elongated, straight and 223

224 apically pointed. The impressions of digit increase in length from I to IV. The impression of

digit V is shallower than others, nearly absent. Digits I-IV shows a sub-parallel arrangement. 225

The divarication angle of digits I-V varies from 40 to 41°. Pes tracks show a large and 226

triangular plantar-like impression (Fig. 3D–G). 227

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Manus imprints are pentadactyl, smaller than pes tracks, longer than wide to as long as wide (ML=5.1-6.0 cm, MW=4.4-5.2 cm; Fig. 3D-G). The ML/MW ratio varies from 0.8 to 1.2. 230 Digit imprints are short, straight and apically pointed. The impression of the digit III is the 231

longest. The divarication angle of digits I-V varies from 68 to 70°. Manus tracks sometimes 232 show a poorly impressed sole. Marks of phalangeal nodes and claws are very well-preserved 233

234 on both pes and manus tracks (Fig. 3D–G).

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Remarks. Based on the following characters, these tracks match with the diagnostic features 237 of *Apatopus lineatus*: quadrupedal trackways; pes long and narrow, plantigrade, pentadactyl 238 239 with slender digits increasing in length from I to IV; digits with well-developed articular swellings and sharp claw impressions; manus pentadactyl, semi-plantigrade, short, and quite 240 241 symmetrical around digit III, which is the longest. Biron and Dutuit (1981) correctly ascribed these tracks to Apatopus lineatus. However they committed several mistakes in their 242 243 interpretations of the tracks: (1) their interpretative sketch of the trackway (Biron and Dutuit, 1981; fig. 9B p. 412) is not concordant with what can be seen on the slab - although Biron and 244 Dutuit (1981) noted a problem of laterally of right /left sets, it appears that they erroneously 245 associated two manus/pes sets to a same trackway - this misinterpretation explained why the 246 order of digits I-V is inverted on their stretch (Biron and Dutuit, 1981; fig. 9B p. 412; digit I 247 towards the outside of the trackway whereas digit V was located towards the inside); (2) 248 although they mentioned five digits on manus tracks, they drawed only up to four digits on 249 their interpretative sketch (Biron and Dutuit, 1981; fig. 9 p. 412) of the trackway. The gross 250 251 morphologies of Apatopus and Synaptichnium show many similarities. However,

- 252 Synaptichnium displays a compact metatarsal-phalangeal axis, which is absent amongst
- 253 *Apatopus* imprints (Klein and Lucas, 2013). *Apatopus lineatus* was described from the
- Triassic deposits of the Argana Basin (Lagnaoui et al., 2012, 2016).
- 255

256 Rhynchosauroides isp.

257 Fig. 4A–D

Material. Numerous tracks (around a hundred) preserved as convex hyporeliefs on slabs
UCA.82 and UCA.83, Cadi Ayyad University (Marrakech).

260 Description. The tracks are small, pentadactyl, digitigrade to plantigrade and consist of numerous isolated pes and manus imprints (Fig. 4A–D) forming a trampled surface on slab 261 262 UCA.82 (Fig. 2C). Tracks are so abundant that it was not possible to distinguish a trackway nor to clearly identify pes/manus sets. Morphology of tracks varies a lot. Only some tracks 263 264 display well marked impression of all digits. In most of tracks only two or three traces of digits are present (II-IV mainly). Tracks showing five traces of digits are longer than wide, up 265 to 2.3 cm long and up to 1.6 cm wide (Fig. 4A–D). The L/W ratio varies from 1.4 to 1.9. Digit 266 imprints are slender, elongated, curved (sometimes straight) and show sharp claw traces. The 267 impressions of digit increase in length from I to IV. When preserved, the impression of digit 268 V is clearly more curved than others (Fig. 4A–D). Digits I-IV are curved inward showing a 269 sub-parallel arrangement. Digit V is oriented outward compared to the group I-IV. The 270 position of the impression of digit V is more proximal than those of digits I-IV. The 271 divarication angle of digits I-V varies from 63 to 88°. 272

273

274*Remarks.* These tracks can be confidently ascribed to *Rhynchosauroides* on the basis of the275following characters: small, pentadactyl, digitigrade to plantigrade tracks; length impressions276of digits increasing from I to IV; impressions of digits I-IV usually curved inward; postero-277laterally positioned short digit V impression that is oriented antero-laterally. *Gwyneddichnium*278differs from *Rhynchosauroides* in showing a digit proportion III > II ≥ IV > I (Klein and279Lucas, 2021).

280

281 **5. Discussion**

5.1. Ichnotaxonomic remarks on the other tetrapod tracks described by Biron and Dutuit
(1981)

Based on a single, small and pentadactyl lacertoid footprint, Biron and Dutuit (1981) 285 286 described a new ichnotaxon, "Enigmatopus atlensis" (Fig. 5A). Showing some resemblance to 287 Rhynchosauroides and Procolophonichnium-like latertoid forms, Klein and Lucas (2021) 288 considered this track as a *nomen dubium*. We agree with this proposition. Based on a track with bifurcated digit impressions, Biron and Dutuit (1981) erected the ichnospecies 289 "Hyloidichnus triasicus" (Fig. 5B). Hyloidichnus being a Permian ichnogenus (Voigt and 290 Lucas, 2018) with a different digit arrangement and given the fact that the description and the 291 illustrations of Biron and Dutuit (1981) are of very poor quality without any valid arguments 292 supporting their hypothesis, we do not consider "*H. triasicus*" as a valid ichnospecies. Biron 293 and Dutuit (1981) described an isolated track they ascribed to a new species, "Rectilinetopus 294 ourikensis" (Fig. 5C), in spite of similarities with the tracks they identified as Apatopus 295 lineatus. Klein and Lucas (2013) synonymized "Rectilinetopus ourikensis" with Apatopus 296 *lineatus*. We agree with this proposition since this track shares strong similarities with pes 297 298 tracks of the slab UCA.83 (pes imprints pentadactyl, longer than wide; with digit imprints elongated, strait and apically pointed; impressions of digit increasing in length from I to IV; 299 digits I-IV with a sub-parallel arrangement; large and triangular plantar-like impression). 300 Biron and Dutuit (1981) erected the ichnospecies "Prochirotherium triasicum" (Fig. 5D), for 301 an isolated, small, pentadactyl, longer than wide, quite symmetrical track with a wide sole, 302 303 and with straight and apically pointed digit impressions (impression of the digit III being the longest). They justified this assignment by similarities with tracks described as 304 *Prochirotherium* from Permian deposits of Italy. Due to the poorness of the material 305 described by Biron and Dutuit (1981) and the absence of diagnostic features, we consider "P. 306 307 triasicum" as a nomen dubium. Based on a single isolated semi-plantigrade to plantigrade tetradactyl imprint, Biron and Dutuit (1981) erected the new species "Pseudochirotherium 308 309 oukaimedensis" (Fig. 5E). Due to the partial preservation of tracks and the absence of diagnostic features, we consider this ichnotaxon as a nomen dubium. Klein and Lucas (2021) 310 proposed that it consists of a chirotheriid track of indeterminate affinity. 311

312

313 5.2. Age of the tracksite

All tracksites reported by Biron and Dutuit (1981) and Biron (1982) are 314 stratigraphically located at the base of the Oukaïmeden Sandstone Fm (F5 Fm sensu Biron, 315 1982). Based on pollen analyses (Cousminer and Manspeizer, 1976; Biron and Courtinat, 316 1982), these deposits were considered as Late Triassic. However, based on conchostracans 317 and new palynological data, Benaouiss et al. (2018) proposed a Ladinian-early Carnian age 318 for the base of the Oukaïmeden Sandstone Fm. In Morocco, Rhynchosauroides was reported 319 from Middle to Late Triassic in the Argana Basin (Klein et al., 2011; Lagnaoui et al., 2012, 320 321 2016; Zouheir et al., 2020). Possible Rhynchosauroides were also described in the Early Triassic from the same basin (Klein et al., 2010). Although Rhynchosauroides shows a broad 322 stratigraphic range, extending from the middle Permian (e.g. Marchetti et al., 2022) to the 323 Late Jurassic (Avanzini et al., 2010), the stratigraphic distribution of Apatopus lineatus and 324 *Isochirotherium* are more restricted. *Apatopus lineatus* is characteristic of the Late Triassic 325 being known from the Carnian to the Rhaetian (Klein and Lucas, 2013). In Morocco, this 326 ichnospecies was reported from the Carnian deposits of the Argana Basin (Lagnaoui et al., 327 2012, 2016). Isochirotherium is characteristic of the Lower to the Middle Triassic with a large 328 geographic distribution (France, Italy, Germany, Great Britain, Spain, and the USA; Klein and 329 330 Haubold, 2007; Klein and Lucas, 2021 and references therein). In Morocco, Isochirotherium *coureli* was reported by Klein et al. (2011) from the Anisian-Ladinian (Middle Triassic) 331 deposits of the Argana Basin. In conclusion, the palaeoichnological assemblage described 332 here, as well as the palynological and conchostracean content (Benaouiss et al., 2018) from 333 the base of the Oukaïmeden Sandstone Fm suggest a Middle-Late Triassic age. The 334 ichnoassemblage described here displays the first documented co-occurrence of Apatopus and 335 336 Isochirotherium.

337

338 5.3. Trackmakers and their paleoenvironment

Trackmakers of *Apatopus* are considered as phytosaurs (Klein and Lucas, 2013). Morphology of *Apatopus* shares similarities with crocodylian imprints, suggesting a semiaquatic mode of life (Klein and Lucas, 2013). This contrast with the compact metatarsalphalangeal axis typically observed amongst *Synaptichnium* whose trackmakers are supposed to be a terrestrial archosauriforms (Klein and Lucas, 2013). *Isochirotherium* is usually referred to possible crocodylian-stem archosaurs or non-crown-group archosauriforms (Klein

and Lucas, 2021). Trackmakers of *Rhynchosauroides* is considered as lepidosauromorph
and/or archosauromorph trackmakers (Avanzini and Renesto, 2002).

As mentionned by Biron and Dutuit (1981), vertebrate tracks from the Triassic 347 deposits of the Ourika and Yagour basins are commonly associated with Scovenia. This 348 ichnogenus is interpreted as a feeding or a combined feeding and locomotion trace, produced 349 by arthropods or annelids (Bromley and Asgaard 1979; Frey et al., 1984). The co-occurrence 350 of these protostomian burrows and traces with tetrapod tracks matches with the Scoyenia 351 352 ichnofacies. This latter is observed in non-marine formations commonly interpreted as subaqueous freshwater sediments (Buatois and Mángano, 1995; Buatois et al., 1997). 353 According to Fabuel-Perez et al. (2009), the lower part of the Oukaïmeden Sandstone Fm was 354 355 deposited in an ephemeral fluvial system. As a result, the living environment was semi-arid, with ponds and ephemeral sediment supply. 356

357

358 **6.** Conclusions

-The two slabs from the Cadi Ayyad University collection bear tetrapods tracks ascribed to

360 *Apatopus lineatus, Isochirotherium* isp. and *Rhynchosauroides* isp. whose trackmakers are

361 possible phytosaurs, crocodylian-stem archosaurs or non-crown-group archosauriforms and

362 lepidosauromorphs/archosauromorphs, respectively.

-All ichnotaxa created by Biron and Dutuit (1981) ("*Chirotherium atlensis*", "*Enigmatopus*

364 atlensis", "Hyloidichnus triasicus", "Prochirotherium triasicum", "Pseudochirotherium

365 *oukaimedensis*" and "*Rectilinetopus ourikensis*") are considered as invalid.

The lithological features of the base of the Oukaïmeden Sandstone Fm as well as the
presence of tetrapod walking tracks indicate that the depositional environment was a semiarid alluvial plain with ponds, periodically flooded by ephemeral fluvial system.

- The palaeoichnological assemblage includes a mixture of Early to the Middle Triassic

370 (Isochirotherium) and Late Triassic (Apatopus lineatus) tracks. This result corroborates the

biostratigraphic data recently obtained from the base of the Oukaïmeden Sandstone Fm,

372 suggesting a Middle-Late Triassic age.

373

374 Declaration of competing interest

375	The authors declare that they have no known competing financial interests or personal
376	relationships that could have appeared to influence the work reported in this paper.
377	
378	Data availability
379	Data will be made available on request.
380	
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383	Geology at Cadi Ayyad University Marrakech for the access to the collection room housing
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385	constructive and thoughtful reviews on the manuscript.
386	
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542 Figure captions

543

- Fig. 1. Geological and stratigraphical setting of Marrakech High Atlas around the Ourika and
 Yagour basins. A, Permian and Triassic outcrops throughout the High Atlas:
- 546 CHA = Central High Atlas, EHA = Eastern High Atlas, MA = Middle Atlas, MHA =
- 547 Marrakech High Atlas and WHA = Western High Atlas (modified after Beauchamp, 1988;
- Laville and Piqué, 1991). Rectangle indicates the study area. **B**, geological map modified after
- 549 Biron (1982), showing tetrapod trackway sites marked by stars with number: (1) East of
- 550 Aghbalou, (2) North-East of Anammer, and (3) NW of Yagour Plateau. C, Simplified
- 551 lithostratigraphic column of the MHA area (F1 to F6 Fm). The studied interval is localised at
- the base of the F5 Fm. (1) early Permian (radioactive and zircon method ,Van Houten, 1977);
- (2) middle Anisian (from pollens, El Arabi et al., 2006); (3) middle Carnian (from pollens,
- Biron and Courtinat, 1982); (4) Ladinian-early Carnian (from pollens and conchostracans;
- 555 Benaouiss et al., 2018); (5) middle Carnian (from pollens; Cousminer and Manspeizer, 1976).

556

Fig. 2. *Isochirotherium*. A–C, Slab UCA.82 bearing several consecutive pes/manus sets (P.,
pes track; M., manus track), photograph (A), photogrammetric 3D mesh (B) and interpretative
sketch (C); the red dotted line indicates the axis of the partial trackway. D–F, Pes/manus set;
photograph (D), digital elevation model in false colours (E) and interpretative sketch (F).
Convex hyporeliefs. G–H, details of the skin trace; photograph (G), and interpretative sketch
(H).

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Fig. 3. *Apatopus lineatus*. A–C, Slab UCA.83. bearing several consecutive pes/manus sets,
two of them forming a trackway (blue dots) (LP. left pes track, LM. left manus track, RP.
right pes track, RM. right manus track), photograph (A), photogrammetric 3D mesh (B) and
interpretative sketch (C). D–G, Two pes/manus sets; photograph (D, F) and interpretative
sketches (E, G). Convex hyporeliefs.

- 570 Fig. 4. *Rhynchosauroides* isp., two tracks from the slab UCA.82. A–D, Isolated tracks,
- 571 photographs (A, C) and interpretative sketch (B, D). Convex hyporeliefs.

572

- 573 Fig. 5. Invalid or problematic ichnotaxa reported by Biron and Dutuit (1981) in the Triassic of
- the Ourika-Yagour basins. A, "Enigmatopus atlensis" (invalid ichnotaxon in this study). B,
- 575 "Hyloidichnus triasicus" (nomen dubium in this study). C, "Rectilinetopus ourikensis"
- 576 (Apatopus lineatus in this study). **D**, "Prochirotherium triasicum" (nomen dubium in this
- 577 study). **E**, "*Pseudochirotherium oukaimedensis*" (nomen dubium in this study).

578

- 579 Fig. 6. Comparison between some Triassic Chirotherian tracks. A, *Brachychirotherium*, Late
- 580 Triassic, Germany. **B**, *Isochirotherium lomasi*, Middle Triassic, Great Britain. **C**,
- 581 *Chirotherium ("Parachirotherium")*, Middle Triassic, Germany. **D**, *Protochirotherium*, Early
- 582 Triassic, Germany. E, *Synaptichnium*, Middle Triassic, Great Britain. Modified after Klein

583 and Lucas (2021).













-Triassic tetrapod tracks from the Ourika-Yagour basins (Morocco) are studied.

-The historical collection of Pierre-Eymard Biron is ichnotaxonomically revised.

-These exquisitely preserved trackways are studied using 3D photogrammetric technique.

-Tracks include Apatopus lineatus, Isochirotherium isp. and Rhynchosauroides isp.

-They suggests a Middle-Late Triassic age for the Oukaïmeden Sandstone Formation.

Journal Proproof



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To Editor-in-Chief of Journal of African Earth Sciences

Dear Editor-in-Chief,

Declaration of interests

 \boxtimes The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

□The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

Yours sincerely,

Dr. Jean-David Moreau for the authors