

# Tetrapod tracks from the Middle/Late Triassic of the Ourika and Yagour basins (Morocco): new ichnotaxonomical data based on the Biron collection

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- 1 Tetrapod tracks from the Middle/Late Triassic of the Ourika and Yagour basins
- 2 (Morocco): new ichnotaxonomical data based on the Biron collection
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- 13
- 14 **Abstract:** Palaeontological prospecting organized in the Ourika and Yagour basins
- 15 (Marrakech High Atlas, Morocco) in the 1980's, have yielded abundant and exquisitely
- preserved tetrapod tracks, some of them preserving thinly detailed skin traces (Collection
- Biron, Cadi Ayyad, Marrakech, Morocco). Based on this material, diverse morphologies of
- tracks were described and six new ichnotaxa were erected. However, these trackways have
- 19 never been reinvestigated in detail from an ichnotaxonomic point of view. Here, the Biron
- 20 collection is reevaluated using modern photogrammetric imaging techniques. The fine-
- 21 grained sandstone slabs bear trackways ascribed to *Apatopus lineatus*, *Isochirotherium* isp.
- and *Rhynchosauroides* isp. whose trackmakers are possible phytosaurs, crocodylian-stem
- archosaurs or non-crown-group archosauriforms and lepidosauromorphs/archosauromorphs,
- respectively. "Chirotherium atlensis", "Enigmatopus atlensis", "Hyloidichnus triasicus",
- 25 "Prochirotherium triasicum", "Pseudochirotherium oukaimedensis" and "Rectilinetopus
- ourikensis" are considered as invalid. The palaeoichnological assemblage includes a mixture
- of Early to the Middle Triassic (*Isochirotherium*) and Late Triassic (*Apatopus lineatus*) tracks.
- 28 This result corroborates the biostratigraphic data recently obtain from the base of the
- 29 Oukaïmeden Sandstone Formation, suggesting a Middle-Late Triassic age. The living
- 30 environment was a semi-arid alluvial plain with ponds and ephemeral rivers.

31	Keywords:	tetrapod tracks	; Mesozoic;	Ladinian;	Carnian:	Marrakech	High Atlas:
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32 photogrammetry

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### 1. Introduction

35	In Morocco, significant ichnofossils including abundant and diverse vertebrate and
36	invertebrate tracks were reported from Triassic deposits of the Argana Basin (Klein et al.,
37	2010, 2011; Tourani et al., 2010; Lagnaoui et al., 2012, 2016; Zouheir et al., 2020), the Sidi
38	Saïd Maachou Basin (Hminna et al., 2013, 2020, 2021) and the Yagour Plateau (Biron and
39	Dutuit, 1981; Biron, 1982). They constitute a key for the reconstruction of continental early
40	Mesozoic palaeoecosystems (Biron and Dutuit, 1981; Klein et al., 2010, 2011; Tourani et al.,
41	2010; Lagnaoui et al., 2012, 2016; Zouheir et al., 2020; Hminna et al., 2013, 2020, 2021).
42	Whereas tracks from the Triassic deposits of the Argana Basin were abundantly studied
43	(Biron and Dutuit, 1981; Klein et al., 2010, 2011; Tourani et al., 2010; Lagnaoui et al., 2012,
44	2016; Zouheir et al., 2020), those from the Ourika and Yagour basins clearly appear to be
45	underestimated. In 1981, Pierre-Eymard Biron discovered several tracksites (Biron and
46	Dutuit, 1981; Biron, 1982). This material was described by Biron and Dutuit (1981) who
47	dated it as Late Triassic. In this study, they described two slabs bearing exquisitely preserved
48	tracks that they attributed to Apatopus lineatus and a new species of Chirotherium: "C.
49	atlensis". Then, based on isolated pes or manus tracks, they briefly described five additional
50	new ichnotaxa: "Enigmatopus atlensis", "Hyloidichnus triasicus", "Prochirotherium
51	triasicum", "Pseudochirotherium oukaimedensis" and "Rectilinetopus ourikensis". Except for
52	a tentative discussion by Klein and Lucas (2013, 2021) of some of these ichnospecies, the
53	tracks from the Ourika and Yagour basins have never been reinvestigated in detail. Here, we
54	revised Biron's material housed at the Cadi Ayyad University (Marrakech). We describe it
55	using photogrammetric imaging techniques, update the ichnotaxonomy of the tracks and
56	discuss their age.

57

58

### 2. Geographical and geological setting

59

60	The study area is located in the Martakech High Atias (Fig. 1A), 60 to 80 km
61	southeast of Marrakech. At the south of the Carboniferous Ourika-Zat anticline (Fig. 1B), the
62	Permian and Triassic outcrops overlie the deformed Variscan basement, represented by
63	Precambrian volcanic rocks, Cambrian series, granitoids and metamorphic rocks, or
64	Carboniferous deposits. The Permian and Triassic rift area is related to the NW-SE-oriented
65	extension, which generated N-S trending normal faults (Van Houten, 1977; El Arabi et al.,
66	2003; Qarbous et al., 2003; Baudon et al., 2009) and reactivated pre-existing NE-SW to ENE
67	WSW Variscan structures (Fig. 1B). These dip-slip and strike-slip faults are reactivated
68	(Baudon et al., 2009; El Arabi, 1988; El Arabi et al., 2003; Laville and Petit, 1984; Laville
69	and Piqué, 1991) and controlled the Triassic sedimentation (Benaouiss et al., 2023). The
70	Permian and Triassic sedimentary successions are represented by elongated and narrow
71	tectonic belts with an ENE-WSW orientation (Fig. 1B), up to 17 km wide and up to 35 km
72	long. These Permian and Triassic deposits are subdivided into six formations (fms), noted F1
73	to F6 (Biron, 1982; Fig. 1C), corresponding to the Permian Anrar Conglomerate (F1), dated
74	by radioactive and zircon method (Van Houten, 1977 and Perez et al., 2019) and the Cham-
75	El-Houa Siltstone (F2), located below the Permian-Triassic unconformity, is attributed to the
76	Permian age by correlation with the Argana t2 Unit (e.g. Benaouiss et al., 2023). The F3
77	(Timalizène Conglomerate), F4 (Anouffig Siltstone), F5 (Oukaïmeden Sandstone), and F6
78	(Tafilalt Siltstone) fms of Biron (1982) correspond respectively to the Basal Conglomerate,
79	Ramuntcho Siltstone, the Oukaïmeden Sandstone and Upper Siltstone fms of Mattis (1977).
80	The F3 Formation (Fm) is attributed to Early Triassic (e.g. Benaouiss et al., 2023) by
81	correlation with the t3 Unit in the Argana Basin, dated by Tourani et al. (2010) with tetrapod
82	tracks. Based on pollen analyses, the lower part of F4 Fm is middle Anisian age (El Arabi et
83	al., 2006) and the F5 Fm is dated middle Carnian at its lower (Biron and Courtinat, 1982) and
84	upper parts (Cousminer and Manspeizer, 1976) (Fig. 1C). The base of the Oukaïmeden
85	Sandstone Fm was recently dated by pollen grains and conchostracans as Ladinian-early
86	Carnian (Benaouiss et al., 2018). This part of F5 Fm represents the equivalent of the lower
87	part of the t5 Unit in the Argana Basin which is dated as Julian-Tuvalian (mid-Carnian) by
88	Tourani et al. (2023). The F6 Fm is attributed to the Norian, by correlation to the t8 Unit in
89	the Argana Basin which is dated by conchostracans according to Weems and Lucas (2015).
90	The Triassic series is overlain by Early Jurassic basalts, representing a good stratigraphic
91	marker across Morocco (Bertrand et al., 1982; Blackburn et al., 2013; De Pachtère, 1983;
92	Marzoli et al., 2004).

Two main Triassic outcrops are of interest to this study (Fig. 1B); (1) those of the
Ourika Basin, which is limited at the north by the Sidi Ali-Ou-Fars fault and at the south by
the Oukaïmeden fault (Fig. 1B), (2) those of the Yagour Basin, which is limited by the Sidi
Ali-Ou-Fars and Tafrawt faults at the north and the Meltsene fault at the south (Fig. 1B).
Biron and Dutuit (1981) reported three tracksites on the Ourika and Yagour basins, at the East
of Aghbalou, at the North-East of Anammer and at the North-West of the Yagour Plateau
(Star symbol in Fig. 1B). The historical tracksites are located at the base of the Oukaïmeden
Sandstone Fm (Fig. 1B and C). This formation is up to 600 m thick and outcrops mainly as
cliffs, accessible along streams and tracks, and is mainly characterized by channel sandstone
bodies alternating with lenticular fluvial bar sandstones and overbank mudstone deposits
(Benaouiss et al., 1996; Fabuel-Perez et al., 2009). The Oukaïmeden Sandstone Fm is
subdivided into three units (Benaouiss et al., 1996; Fabuel-Perez et al., 2009), called lower-,
middle- and upper Okaïmeden units. The lower unit, which yields the fossil tracks, is 30 to
110 m thick. It is richer, than the middle and upper units, in red clay-stones and siltstones, and
was deposited in an ephemeral fluvial system under dry climate. The middle unit is 25 to 80
m thick, and corresponds to a distal braided system associated with more humid conditions
(Fabuel-Perez et al., 2009). The upper unit, which is 30 to 80 m thick, corresponds to the
comeback of ephemeral conditions with the installation of aeolian dunes corresponding to aric
conditions (Benaouiss et al., 1996; Fabuel-Perez et al., 2009).

#### 3. Material and methods

The palaeontological collection of the department of Geology (Cadi Ayyad University) includes two large fine-grained sandstone slabs bearing tetrapod tracks, described and figured in Biron and Dutuit (1981) (Figs. 2 and 3). These slabs called "Dalle 1" and "Dalle 2" in Biron and Dutuit (1981), bear currently the collection numbers UCA.82 and UCA.83, respectively. The slab UCA.82 (Fig. 2), was collected at the base of the Yagour Plateau whereas the geographical origin of the slab UCA.83 (Fig. 3) was not indicated in Biron and Dutuit (1981). The slab UCA.82 (63 cm long and 58 cm wide) bears eight large and exquisitely preserved tracks briefly described by Biron and Dutuit (1981) for which they erected the new ichnospecies "Chirotherium atlensis". The slab UCA.82 also bears numerous (around one hundred) small lacertoid tracks (*Rhynchosauroides*) undescribed by Biron and

125	Dutuit (1981) but briefly mentioned in Biron (1982). The slab UCA.83 (44 cm long and 41
126	cm wide) bears eight large tracks briefly described by Biron and Dutuit (1981) who ascribed
127	them to Apatopus lineatus (Fig. 3). This slab also bears four isolated and partial small
128	lacertoid tracks ascribed to Rhynchosauroides (Fig. 4). On both slabs, tetrapod ichnofossils
129	are preserved as convex hyporeliefs. The tracks that Biron and Dutuit (1981) described as
130	"Enigmatopus atlensis", "Hyloidichnus triasicus", "Prochirotherium triasicum", and
131	"Rectilinetopus ourikensis" (Fig. 5) were preserved on large isolated blocs left in place on the
132	field (unfortunately not precisely located in Biron and Dutuit, 1981).
133	The descriptive terminology and biometric parameters used in this study follow those
134	of Haubold (1971) and Leonardi (1987). We used the following standard abbreviations: ML,
135	manus track length; MW, manus track width; ML/MW, length to width ratio of manus track;
136	PL., pes track length; PW, pes track width; and PL/PW, length to width ratio of pes track.
137	3D photogrammetric reconstructions were produced with the software Agisoft
138	Metashape Professional 2.0. Pictures of the surfaces were taken with a Nikon D5200 camera
139	coupled with an AF-S Nikkor 18–105 128 mm f/3.5–5.6G ED lens. The same software was
140	used to produce photogrammetric digital elevation models and contours.
141	
142	4. Systematic palaeoichnology
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143	
144	Isochirotherium isp.
145	Fig. 2A–F
146	Material. Eight tracks preserved as convex hyporeliefs on slab UCA.82, Cadi Ayyad
147	University (Marrakech).
148	Description. The slab UCA.82 bears eight semi-plantigrade and quadrupedal tracks, including
149	five pes tracks and three manus imprints (including three pes/manus sets; Fig. 2A–C). Two
150	pes/manus sets are part of a trackway (Fig. 2A–C). Pes and manus tracks from the same set
151	form a small divergent angle of 12 to 15°. Pes and manus track paces are 34.5 cm and 36.0 cm
152	long, respectively. The pes and manus tracks are strongly oriented outward.
153	Pes imprints are pentadactyl, longer than wide, 16–17 cm long and 10.7–11.8 cm wide
154	(Fig. 2D-F). The PL/PW ratio varies from 1.4 to 1.5. Digit imprints I-V are elongated, robust,

straight and apically pointed. They form a group oriented outward, whereas imprint of digit V 155 is curved rearward. The position of the impression of digit V is more proximal than those of 156 digits I-IV. Traces of digit V show a well-imprinted, oval and longer than wide proximal pad 157 that is in line with impression of digit IV (Fig. 2D-F). The impressions of the digit II and III 158 (that are similar in length) are the longest. The impressions of digits I and IV are shorter than 159 those of II and III. The divarication angle of digits I-IV and digits I-V varies from 20 to 35° 160 and 52 to 54°, respectively. 161 Manus imprints are pentadactyl, quite symmetrical, smaller than pes tracks (PL/ML= 162 2.6), wider than long (ML=6.5 cm, MW=7.3 cm, ML/MW ratio = 0.9; Fig. 2D-F). The digit 163 imprints are moderately elongated, straight and apically pointed. The impression of the digit 164 III is the longest. The impression of digit V is shallower than others. The impression of the 165 digit I is the shortest. The divarication angle of digits I-V is up to 120°. Proximal pads of 166 digits I–V form a posteriorly concave margin. Manus tracks sometimes show a poorly marked 167 sole. On pes and manus tracks, phalangeal nodes, skin detailed-structures and large triangular 168 claws are very well-preserved (Fig. 2D-F). The slab UCA.82 bears thin, elongated and 169 170 straight to curved ridges that could be possible tail marks (Fig. 2C). Remarks. Biron and Dutuit (1981) identified these tracks as a new species, they named 171 "Chirotherium atlensis". The gross morphologies of Chirotherium and Isochirotheriumare 172 very similar. However, *Isochirotherium* differs from *Chirotherium* by the size of pedal digits 173 II and III that are quite similar in length (Klein et al., 2011, 2013). In Chirotherium the digit 174 III is the longest. For this reason, we attribute the tracks from the slab UCA.82 to 175 *Isochirotherium* and consider the ichnospecies "Chirotherium atlensis" as an invalid name. 176 Chirotherian tracks (Brachychirotherium, Parachirotherium, Protochirotherium, and 177 Synaptichnium), were abundantly described from the Triassic of Morocco: in the Sidi Saïd 178 179 Maachou (Hminna et al., 2013, 2021) and Argana basins (Klein et al., 2010, 2011; Tourani et al., 2010; Lagnaoui et al., 2012, 2016; Zouheir et al., 2020). They all differ from the tracks of 180 the slab UCA.82: Brachychirotherium (Fig. 6A) differs by broad pes that has short and broad 181 digits and the oval basal pad of digit V that is in a postero-lateral position relatively to the 182 anterior digit group I–IV; *Parachirotherium* (Fig. 6C) differs by the pentadactyl, functionally 183 tridactyl pes traks with a wide divarication angle I-V, reduced imprints of digits II and IV, and 184 a posteriorly shifted digit I. Lagnaoui et al. (2012, 2016) mentioned the occurrence of 185 Parachirotherium in the Late Triassic of the Argana Basin. Nevertheless, according to Klein 186 and Lucas (2021) Parachirotherium should be regarded as a junior synonym of Chirotherium; 187

188	Protochirotherium (emended in Klein and Lucas, 2021 after Fichter and Kunz, 2004) differs
189	from the tracks of the slab UCA.82 by larger manus tracks compared to the size of pes track
190	and digit V of pes with strongly elongated pad proximally (Fig. 6D); Synaptichnium differs by
191	a more slender overall shape, a less massive pedal digit V, and a pedal digit IV that is
192	subequal with or even longer than digit III (Fig. 6E). In Morocco, Tourani et al. (2010) and
193	Klein et al., (2011) reported <i>Isochirotherium</i> cf. gierlinskii and <i>Isochirotherium coureli</i> from
194	the Triassic of the Argana Basin. The tracks described as <i>Isochirotherium</i> cf. <i>gierlinskii</i> in
195	Tourani et al. (2010) were tentatively renamed <i>Protochirotherium</i> by Klein et al. (2013). The
196	tracks of the slab UCA.82 has no distinctive character supporting the creation of a new
197	species nor can it be confidently assigned to any of the known species of Isochirotherium.
198	According to Klein and Lucas (2021), the valid <i>Isochirotherium</i> ichnospecies are <i>I. coltoni</i> , <i>I</i> .
199	coureli, I. delicatum, I. herculis, I. lomasi, I. marshalli and I. soergeli. I. coureli differs from
200	the tracks of the slab UCA.82 in showing higher PL/PW ratio (up to 1.9). I. soergeli differs in
201	showing pes tracks with as long as wide to slightly longer than wide group I-IV. I. delicatum
202	shows more elongated pes tracks with long traces of digits II-III as well as very short traces of
203	digits I and IV. I. marshalli differs from the tracks of the slab UCA.82 in showing a bigger
204	ovoid metatarsal-phalangeal pad on the digit V of pes tracks. I. herculis shows strongly bigger
205	pes (PL can be more than 30 cm) and a group I-IV which is wider than long. I. coltoni and I.
206	soergeli are clearly smaller than the tracks of the slab UCA.82. Similar to the tracks we
207	describe here, the pes track of <i>I. lomasi</i> shows a longer than wide group I-IV. However, <i>I.</i>
208	lomasi displays a higher PL/PW ratio (more than 1.6) and a shorter trace of digit IV on pes
209	track.
210	
211	Apatopus lineatus Bock, 1952
212	Fig. 3A–G
213	
214	Material. Eight tracks preserved as convex hyporeliefs on slab UCA.83, Cadi Ayyad
215	University (Marrakech).
216	Description. The slab UCA.83 bears eight plantigrade to semi-plantigrade and quadrupedal
217	tracks, with four pes tracks and three manus imprints (including three pes/manus sets; Fig.
218	3A-C). Two pes/manus sets are part of a trackway (Fig. 3C). Pes and manus tracks from the

219	same set form divergent angle of 22 to 40°. Pes and manus track paces are 22.5 cm and 19.0
220	cm long, respectively. The pes tracks are oriented outward whereas manus tracks are oriented
221	forward to inward.
222	Pes imprints are pentadactyl, longer than wide, 10.0–10.5 cm long and 4.7–6.0 cm wide (Fig.
223	3D-G). The PL/PW ratio varies from 1.7 to 2.1. Digit imprints are elongated, straight and
224	apically pointed. The impressions of digit increase in length from I to IV. The impression of
225	digit V is shallower than others, nearly absent. Digits I-IV shows a sub-parallel arrangement.
226	The divarication angle of digits I-V varies from 40 to 41°. Pes tracks show a large and
227	triangular plantar-like impression (Fig. 3D–G).
228	
229	Manus imprints are pentadactyl, smaller than pes tracks, longer than wide to as long as wide
230	(ML=5.1-6.0 cm, MW=4.4-5.2 cm; Fig. 3D-G). The ML/MW ratio varies from 0.8 to 1.2.
231	Digit imprints are short, straight and apically pointed. The impression of the digit III is the
232	longest. The divarication angle of digits I-V varies from 68 to 70°. Manus tracks sometimes
233	show a poorly impressed sole. Marks of phalangeal nodes and claws are very well-preserved
234	on both pes and manus tracks (Fig. 3D-G).
235	
236	
237	Remarks. Based on the following characters, these tracks match with the diagnostic features
238	of Apatopus lineatus: quadrupedal trackways; pes long and narrow, plantigrade, pentadactyl
239	with slender digits increasing in length from I to IV; digits with well-developed articular
240	swellings and sharp claw impressions; manus pentadactyl, semi-plantigrade, short, and quite
241	symmetrical around digit III, which is the longest. Biron and Dutuit (1981) correctly ascribed
242	these tracks to Apatopus lineatus. However they committed several mistakes in their
243	interpretations of the tracks: (1) their interpretative sketch of the trackway (Biron and Dutuit,
244	1981; fig. 9B p. 412) is not concordant with what can be seen on the slab - although Biron and
245	Dutuit (1981) noted a problem of laterally of right /left sets, it appears that they erroneously
246	associated two manus/pes sets to a same trackway - this misinterpretation explained why the
247	order of digits I-V is inverted on their stretch (Biron and Dutuit, 1981; fig. 9B p. 412; digit I
248	towards the outside of the trackway whereas digit V was located towards the inside); (2)
249	although they mentioned five digits on manus tracks, they drawed only up to four digits on
250	their interpretative sketch (Biron and Dutuit, 1981; fig. 9 p. 412) of the trackway. The gross
251	morphologies of Apatopus and Synaptichnium show many similarities. However,

252	Synaptichnium displays a compact metatarsal-phalangeal axis, which is absent amongst
253	Apatopus imprints (Klein and Lucas, 2013). Apatopus lineatus was described from the
254	Triassic deposits of the Argana Basin (Lagnaoui et al., 2012, 2016).
255	
256	Rhynchosauroides isp.
257	Fig. 4A–D
258	Material. Numerous tracks (around a hundred) preserved as convex hyporeliefs on slabs
259	UCA.82 and UCA.83, Cadi Ayyad University (Marrakech).
260	Description. The tracks are small, pentadactyl, digitigrade to plantigrade and consist of
261	numerous isolated pes and manus imprints (Fig. 4A-D) forming a trampled surface on slab
262	UCA.82 (Fig. 2C). Tracks are so abundant that it was not possible to distinguish a trackway
263	nor to clearly identify pes/manus sets. Morphology of tracks varies a lot. Only some tracks
264	display well marked impression of all digits. In most of tracks only two or three traces of
265	digits are present (II-IV mainly). Tracks showing five traces of digits are longer than wide, up
266	to 2.3 cm long and up to 1.6 cm wide (Fig. 4A-D). The L/W ratio varies from 1.4 to 1.9. Digit
267	imprints are slender, elongated, curved (sometimes straight) and show sharp claw traces. The
268	impressions of digit increase in length from I to IV. When preserved, the impression of digit
269	V is clearly more curved than others (Fig. 4A-D). Digits I-IV are curved inward showing a
270	sub-parallel arrangement. Digit V is oriented outward compared to the group I-IV. The
271	position of the impression of digit V is more proximal than those of digits I-IV. The
272	divarication angle of digits I-V varies from 63 to 88°.
273	
274	Remarks. These tracks can be confidently ascribed to Rhynchosauroides on the basis of the
275	following characters: small, pentadactyl, digitigrade to plantigrade tracks; length impressions
276	of digits increasing from I to IV; impressions of digits I-IV usually curved inward; postero-
277	laterally positioned short digit V impression that is oriented antero-laterally. Gwyneddichnium
278	differs from Rhynchosauroides in showing a digit proportion III $>$ II $\ge$ IV $>$ I (Klein and
279	Lucas, 2021).
280	
281	5. Discussion

283 284	5.1. Ichnotaxonomic remarks on the other tetrapod tracks described by Biron and Dutuit (1981)
285	Based on a single, small and pentadactyl lacertoid footprint, Biron and Dutuit (1981)
286	described a new ichnotaxon, "Enigmatopus atlensis" (Fig. 5A). Showing some resemblance to
287	Rhynchosauroides and Procolophonichnium-like latertoid forms, Klein and Lucas (2021)
288	considered this track as a nomen dubium. We agree with this proposition. Based on a track
289	with bifurcated digit impressions, Biron and Dutuit (1981) erected the ichnospecies
290	"Hyloidichnus triasicus" (Fig. 5B). Hyloidichnus being a Permian ichnogenus (Voigt and
291	Lucas, 2018) with a different digit arrangement and given the fact that the description and the
292	illustrations of Biron and Dutuit (1981) are of very poor quality without any valid arguments
293	supporting their hypothesis, we do not consider "H. triasicus" as a valid ichnospecies. Biron
294	and Dutuit (1981) described an isolated track they ascribed to a new species, "Rectilinetopus
295	ourikensis" (Fig. 5C), in spite of similarities with the tracks they identified as Apatopus
296	lineatus. Klein and Lucas (2013) synonymized "Rectilinetopus ourikensis" with Apatopus
297	lineatus. We agree with this proposition since this track shares strong similarities with pes
298	tracks of the slab UCA.83 (pes imprints pentadactyl, longer than wide; with digit imprints
299	elongated, strait and apically pointed; impressions of digit increasing in length from I to IV;
300	digits I-IV with a sub-parallel arrangement; large and triangular plantar-like impression).
301	Biron and Dutuit (1981) erected the ichnospecies "Prochirotherium triasicum" (Fig. 5D), for
302	an isolated, small, pentadactyl, longer than wide, quite symmetrical track with a wide sole,
303	and with straight and apically pointed digit impressions (impression of the digit III being the
304	longest). They justified this assignment by similarities with tracks described as
305	Prochirotherium from Permian deposits of Italy. Due to the poorness of the material
306	described by Biron and Dutuit (1981) and the absence of diagnostic features, we consider "P.
307	triasicum" as a nomen dubium. Based on a single isolated semi-plantigrade to plantigrade
308	tetradactyl imprint, Biron and Dutuit (1981) erected the new species "Pseudochirotherium
309	oukaimedensis" (Fig. 5E). Due to the partial preservation of tracks and the absence of
310	diagnostic features, we consider this ichnotaxon as a nomen dubium. Klein and Lucas (2021)
311	proposed that it consists of a chirotheriid track of indeterminate affinity.
312	

5.2. Age of the tracksite

All tracksites reported by Biron and Dutuit (1981) and Biron (1982) are
stratigraphically located at the base of the Oukaïmeden Sandstone Fm (F5 Fm sensu Biron,
1982). Based on pollen analyses (Cousminer and Manspeizer, 1976; Biron and Courtinat,
1982), these deposits were considered as Late Triassic. However, based on conchostracans
and new palynological data, Benaouiss et al. (2018) proposed a Ladinian-early Carnian age
for the base of the Oukaïmeden Sandstone Fm. In Morocco, <i>Rhynchosauroides</i> was reported
from Middle to Late Triassic in the Argana Basin (Klein et al., 2011; Lagnaoui et al., 2012,
2016; Zouheir et al., 2020). Possible Rhynchosauroides were also described in the Early
Triassic from the same basin (Klein et al., 2010). Although <i>Rhynchosauroides</i> shows a broad
stratigraphic range, extending from the middle Permian (e.g. Marchetti et al., 2022) to the
Late Jurassic (Avanzini et al., 2010), the stratigraphic distribution of Apatopus lineatus and
Isochirotherium are more restricted. Apatopus lineatus is characteristic of the Late Triassic
being known from the Carnian to the Rhaetian (Klein and Lucas, 2013). In Morocco, this
ichnospecies was reported from the Carnian deposits of the Argana Basin (Lagnaoui et al.,
2012, 2016). Isochirotherium is characteristic of the Lower to the Middle Triassic with a large
geographic distribution (France, Italy, Germany, Great Britain, Spain, and the USA; Klein and
Haubold, 2007; Klein and Lucas, 2021 and references therein). In Morocco, Isochirotherium
coureli was reported by Klein et al. (2011) from the Anisian-Ladinian (Middle Triassic)
deposits of the Argana Basin. In conclusion, the palaeoichnological assemblage described
here, as well as the palynological and conchostracean content (Benaouiss et al., 2018) from
the base of the Oukaïmeden Sandstone Fm suggest a Middle-Late Triassic age. The
ichnoassemblage described here displays the first documented co-occurrence of Apatopus and
Isochirotherium.

#### 5.3. Trackmakers and their paleoenvironment

Trackmakers of *Apatopus* are considered as phytosaurs (Klein and Lucas, 2013). Morphology of *Apatopus* shares similarities with crocodylian imprints, suggesting a semi-aquatic mode of life (Klein and Lucas, 2013). This contrast with the compact metatarsal-phalangeal axis typically observed amongst *Synaptichnium* whose trackmakers are supposed to be a terrestrial archosauriforms (Klein and Lucas, 2013). *Isochirotherium* is usually referred to possible crocodylian-stem archosaurs or non-crown-group archosauriforms (Klein

as is considered as lanidessymmers and
es is considered as lepidosauromorph
Renesto, 2002).
vertebrate tracks from the Triassic
only associated with Scoyenia. This
I feeding and locomotion trace, produced
79; Frey et al., 1984). The co-occurrence
ood tracks matches with the Scoyenia
rmations commonly interpreted as
gano, 1995; Buatois et al., 1997).
art of the Oukaïmeden Sandstone Fm was
t, the living environment was semi-arid,
ection bear tetrapods tracks ascribed to
osauroides isp. whose trackmakers are
non-crown-group archosauriforms and
'Chirotherium atlensis", "Enigmatopus
'Chirotherium atlensis", "Enigmatopus
'Chirotherium atlensis'', "Enigmatopus m triasicum'', "Pseudochirotherium
"Chirotherium atlensis", "Enigmatopus m triasicum", "Pseudochirotherium e considered as invalid.
"Chirotherium atlensis", "Enigmatopus m triasicum", "Pseudochirotherium e considered as invalid.
"Chirotherium atlensis", "Enigmatopus m triasicum", "Pseudochirotherium e considered as invalid.  deden Sandstone Fm as well as the depositional environment was a semi-
"Chirotherium atlensis", "Enigmatopus m triasicum", "Pseudochirotherium e considered as invalid.  deden Sandstone Fm as well as the depositional environment was a semi-
"Chirotherium atlensis", "Enigmatopus m triasicum", "Pseudochirotherium e considered as invalid.  deden Sandstone Fm as well as the depositional environment was a semi- oy ephemeral fluvial system.  ture of Early to the Middle Triassic

373

374

### **Declaration of competing interest**

375	The authors declare that they have no known competing financial interests or personal
376	relationships that could have appeared to influence the work reported in this paper.
377	
378	Data availability
379	Data will be made available on request.
380	
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383	Geology at Cadi Ayyad University Marrakech for the access to the collection room housing
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386	
387	References
388	Avanzini, M., Renesto, S., 2002, A review of Rhynchosauroides tirolicus Abel, 1926
389	ichnospecies (Middle Triassic: Anisian-Ladinian) and some inferences on
390	Rhynchosauroides trackmaker. Riv. Ital. di Paleontol. e Stratigr. 108, 51-66.
391	
392	Avanzini, M., Piñuela, L., Garcia-Ramos, J.C., 2010. First report of a Late Jurassic lizard-like
393	footprint (Asturias, Spain). J. Iber. Geol. 36, 175–180.
394	
395	Baudon, C., Fabuel-Perez, I., Redfern, J., 2009. Structural style and evolution of a Late
396	Triassic rift basin in the Central High Atlas, Morocco: controls on sediment
397	deposition. Geol. J. 44, 677–691.
398	Beauchamp, J., 1988. Triassic sedimentation and rifting in the High Atlas (Morocco). Dev.
399	Geotecton. 22, 477–497.
400	Benaouiss, N., Courel, L., Beauchamp, J., 1996. Rift-controlled fluvial/tidal transitional series
401	in the Oukaïmeden Sandstones, High Atlas of Marrakesh (Morocco). Sed. Geol. 107,
402	21–36.

403	Benaouiss, N, Juncal, M., Bourquin, S., Tourani, A., Broutin, J., Diez, JB., 2018. New
104	biostratigraphical records in the Permian and Triassic basins, High Atlas of
405	Marrakech, Morocco. 2nd International Congress on Permian and Triassic, 19–20.
106	Benaouiss, N., Tourani, A., Bourquin, S., Jalil, NE., 2023. Late Permian to early Middle
407	Triassic palaeoenvironmental reconstructions of the High Atlas, Morocco:
408	Geodynamic and climate implications in the intertropical western Peri-Tethyan
109	domain. Palaeogeogr. Palaeoclimatol. Palaeoecol. 630, 111809.
410	Bertrand, H., Dostal, J., Dupuy, C., 1982. Geochemistry of early Mesozoic tholeiites from
411	Morocco. Earth Planet. Sci. Lett. 58, 225–239.
112	Biron, P. E., 1982. Le Permo-Trias de la région de l'Ourika (Haut-Atlas de Marrakech,
413	Maroc): lithostratigraphie, sédimentologie, tectonique et minéralisations. Université de
114	Grenoble, France. Unpublished PhD diss.
415	Biron, P.E., Dutuit, J., 1981. Figurations sédimentaires et traces d'activité au sol dans le Trias
416	de la formation d'Argana et de l'Ourika (Maroc): Bull. Mus. Hist. Nat. Paris 4, 399-
417	427.
118	Biron, P.E., Courtinat, B., 1982. Contribution palynologique à la connaissance du Trias du
119	Haut Atlas de Marrakech. Geobios 15, 231–235.
120	Blackburn, T.J., Olsen, P.E., Bowring, S.A., McLean, N.M., Kent, D.V., Puffer, J., McHone,
121	G., Rasbury, E.T., Et-Touhami, M., 2013. Zircon U-Pb Geochronology Links the End-
122	Triassic Extinction with the Central Atlantic Magmatic Province. Science 340, 941-
123	945.
124	Bock, W., 1952. Triassic reptilian tracks and trends of locomotive evolution. J. Paleontol. 26,
125	395–433.
126	Bromley, R., Asgaard, U., 1979. Triassic freshwater ichnocoenoses from Carlsberg Fjord, east
127	Greenland. Palaeogeogr. Palaeoclimatol. Palaeoecol. 28, 39–80.
128	Buatois, L. A., Mángano, M.G., 1995. The paleoenvironmental and paleoecological
129	significance of the lacustrine Mermia ichnofacies: an archetypical subaqueous
430	nonmarine trace fossil assemblage. Ichnos 4, 151–161.

431	Buatois, L.A., Mangano, M.G., Maples, C.G., 1997. The paradox of nonmarine ichnofaunas
132	in tidal rhythmites; integrating sedimentologic and ichnologic data from the Late
433	Cretaceous of eastern Kansas, USA. Palaios 12, 467–481.
134	Cousminer, H.L., Manspeizer, W., 1976. Triassic pollen date Moroccan High Atlas and the
435	incipient rifting of Pangea as middle Carnian. Science 191, 943–945.
136	De Pachtère, P., 1983. Le volcanisme permien et fini-triasique dans le Haut-Atlas de
137	Marrakech (Maroc): Approche pétrologique et géochimique. Université de Grenoble,
138	France. Unpublished PhD diss.
139	El Arabi, E.H., 1988. Le Permo-Trias du flanc sud du Haut Atlas de Marrakech.
140	Sédimentologie, cartographie et paléogéographie. Université de Grenoble, France.
141	Unpublished PhD diss.
142	El Arabi, E.H., Ferrandini, J., Essamoud, R., 2003. Triassic stratigraphy and structural
143	evolution of a rift basin: the Eç Çour basin, High Atlas of Marrakech, Morocco. J. Afr
144	Earth Sci. 36, 29–39.
145	El Arabi, E.H., Diez, J.B., Broutin, J., Essamoud, R., 2006. Première caractérisation
146	palynologique du Trias moyen dans le Haut Atlas ; implications pour l'initiation du
147	rifting téthysien au Maroc. C. R. Geoscience 338, 641-649.
148	Fabuel-Perez, I., Redfern, J., Hodgetts, D., 2009. Sedimentology of an intra-montane rift-
149	controlled fluvial dominated succession: the Upper Triassic Oukaimeden Sandstone
450	Formation, Central High Atlas, Morocco. Sediment. Geol. 218, 103-140.
451	Fichter, J., Kunz, R., 2004. New genus and species of chirotheroid tracks in the Detfurth-
452	Formation (Middle Bunter, Lower Triassic) of Central Germany. Ichnos 11, 183-193.
453	Frey, R.W., Pemberton, S.G., Fagerstrom, J.A., 1984. Morphological, ethological, and
154	environmental significance of the ichnogenera Scoyenia and Ancorichnus. J. Paleonto
455	58, 511–528.
456	
457	Haubold, H. 1971. Ichnia Amphibiorum et Reptiliorum fossilium. Handb. Paläoherpetol. 18,
458	1–124.

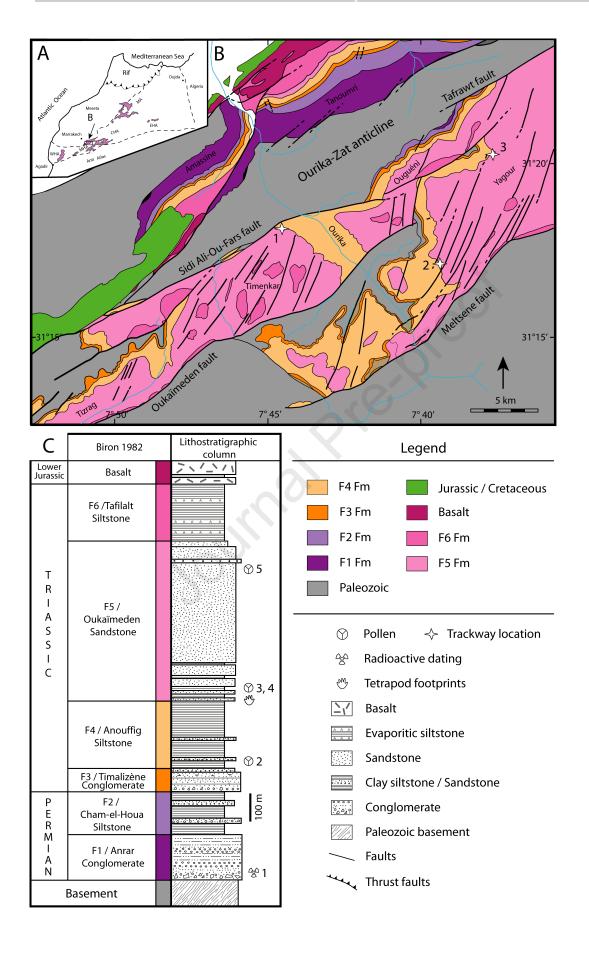
459	Hminna, A., Voigt, S., Klein, H., Saber, H., Schneider, J.W., Hmich, D., 2013. First
460	occurrence of tetrapod footprints from the continental Triassic of the Sidi Said
461	Maachou area (Western Meseta, Morocco). J. Afr. Earth Sci. 80, 1–7.
462	Hminna, A., Lagnaoui, A., Zouheir, T., Saber, H., Schneider, J.W., 2020. Late Triassic
463	ichnoassemblage from a playa-lake system of the Coastal Meseta, Morocco:
464	Palaeoenvironmental and palaeoecological implications. J. Afr. Earth Sci. 172,
465	103995.
466	Hminna, A., Klein, H., Zouheir, T., Lagnaoui, A., Saber, H., Lallensack, J. N., Oukassou, M.,
467	2021. The Late Triassic archosaur ichnogenus Brachychirotherium: first complete step
468	cycles from Morocco, North Africa, with implications for trackmaker identification
469	and ichnotaxonomy. Hist. Biol. 33, 723–736.
470	Klein, H., Haubold, H., 2007. Archosaur footprints – Potential for biochronology of Triassic
471	continental sequences. New Mexico Mus. Nat. Hist. Sci. Bull. 41, 120-130.
472	Klein, H., Lucas, S.G., 2013. The Late Triassic tetrapod ichnotaxon <i>Apatopus lineatus</i> (Bock
473	1952) and its distribution. New Mexico Mus. Nat. Hist. Sci. Bull. 61, 313–324.
474	Klein, H., Lucas, S.G., 2021. The Triassic tetrapod footprint record. New Mexico Mus. Nat.
475	Hist. Sci. Bull. 83, 1–194.
476	Klein, H., Voigt, S., Hminna, A., Saber, H., Schneider, J., Hmich, D., 2010. Early Triassic
477	archosaur-dominated footprint assemblage from the Argana Basin (western High
478	Atlas, Morocco). Ichnos 17, 215–227.
479	Klein, H., Voigt, S., Saber, H., Schneider, J.W., Hminna, A., Fischer, J., Lagnaoui, A.,
480	Brosig, A., 2011. First occurrence of a Middle Triassic tetrapod ichnofauna from the
481	Argana Basin (western High Atlas, Morocco). Palaeogeogr. Palaeoclimatol.
482	Palaeoecol. 307, 218–231.
483	Klein, H., Niedźwiedzki, G., Voigt, S., Lagnaoui, A., Hminna, A., Saber, H., Schneider, J.W.,
484	2013. The tetrapod ichnogenus <i>Protochirotherium</i> Fichter and Kunz 2004, a
485	characteristic Early Triassic morphotype of central Pangea. Ichnos 20, 24-30.

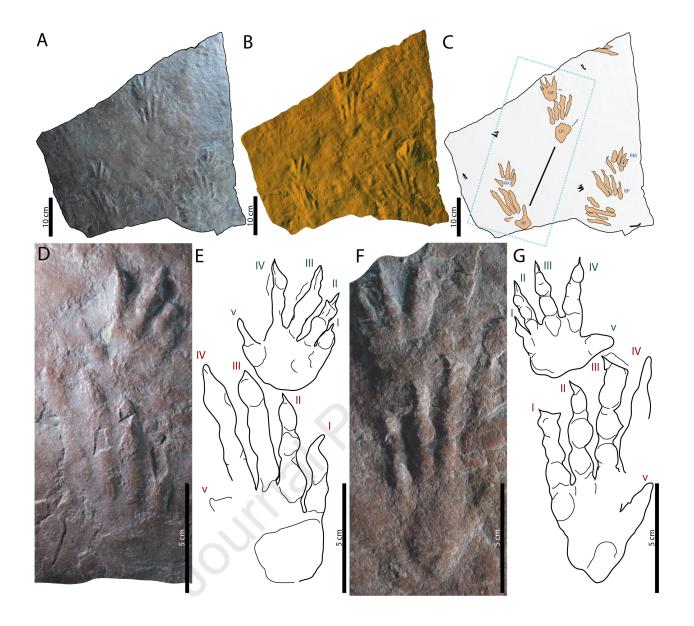
486	Lagnaoui, A., Klein, H., Voigt, S., Hminna, A., Saber, H., Schneider, J. W., Werneburg, R.,
487	2012. Late Triassic tetrapod-dominated ichnoassemblages from the Argana Basin
488	(Western High Atlas, Morocco). Ichnos 19, 238–253.
489	Lagnaoui, A., Klein, H., Saber, H., Fekkak, A., Belahmira, A., Schneider, J.W., 2016. New
490	discoveries of archosaur and other tetrapod footprints from the Timezgadiouine
491	Formation (Irohalene Member, Upper Triassic) of the Argana Basin, western High
492	Atlas, Morocco–Ichnotaxonomic implications. Palaeogeogr. Palaeoclimatol.
493	Palaeoecol. 453, 1–9.
494	Laville, E., Petit, JP., 1984. Role of synsedimentary strike-slip faults in the formation of
495	Moroccan Triassic basins. Geology 12, 424–427.
496	Laville, E., Piqué, A., 1991. La Distension crustale atlantique et atlasique au Maroc au début
497	du Mésozoïque; le rejeu des structures hercyniennes. Bull. Soc. Géol. Fr. 162, 1161-
498	1171.
499	Leonardi, G., 1987. Glossary and manual of tetrapod footprint palaeoichnology. Ministerio
500	Minas Energie, Departemento Nacional da Producao Mineral, Brasilia.
501	Marchetti, L., Logghe, A., Mujal, E., Barrier, P., Montenat, C., Nel, A., J-M, Pouillon, R.
502	Garrouste, J.S., Steyer, 2022. Vertebrate tracks from the Permian of Gonfaron
503	(Provence, Southern France) and their implications for the late Capitanian terrestrial
504	extinction event. Palaeogeography, Palaeoclimatology, Palaeoecology, 599, 111043.
505	Marzoli, A., Bertrand, H., Knight, K.B., Cirilli, S., Buratti, N., Vérati, C., Nomade, S.,
506	Renne, P.R., Youbi, N., Martini, R., 2004. Synchrony of the Central Atlantic
507	magmatic province and the Triassic-Jurassic boundary climatic and biotic crisis.
508	Geology 32, 973–976.
509	Mattis, A.F., 1977. Nonmarine Triassic sedimentation, central High Atlas Mountains,
510	Morocco. J. Sediment. Petrol. 47, 107–119.
511	Perez, N.D., Teixell, A., Gómez-Gras, D., Stockli, D.F., 2019. Reconstructing Extensional
512	Basin Architecture and Provenance in the Marrakech High Atlas of Morocco:
513	Implications for Rift Basins and Inversion Tectonics. Tectonics 38, 1584–1608.

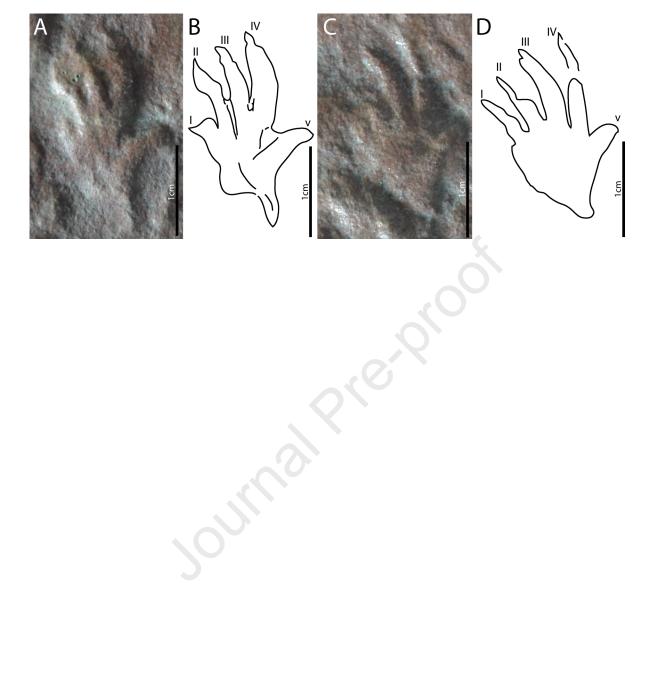
514	exemple d'évolution d'un segment oblique au rift de l'Atlantique central au Trias.
516	Can. J. Earth Sci. 40, 949–964.
517	Tourani, A., Benaouiss, N., Gand, G., Bourquin, S., Jalil, NE., Broutin, J., Battail, B.,
518	Germain, D., Khaldoune, F., Sebban, S., Steyer, JC., Vacant, R., 2010. Evidence of
519	an Early Triassic age (Olenekian) in Argana Basin (High Atlas, Morocco) based on
520	new chirotherioid traces. C. R. Palevol 9, 201–208.
521	Tourani, A., Benaouiss, N., De la Horra, R., Barrenechea, J. F., Lopez-Gomez, J., Gallego, O.
522	F., Bourquin, S., 2023. Characterization of the Carnian Pluvial Episode in the Argana
523	Basin (Western High Atlas, Morocco): An approach based on sedimentology, clay
524	mineralogy and paleosols. Palaeogeogr. Palaeoclimatol. Palaeoecol. 627, 111720.
525	Van Houten, F.B., 1977. Triassic-Liassic deposits of Morocco and East North America: A
526	comparison. Bull. Am. Assoc. Pet. Geol. 61, 79–99.
527	Voigt, S., Lucas, S.G., 2018. Outline of a Permian tetrapod footprint ichnostratigraphy. Geol.
528	Soc. Spec. Publ. 450, 387–404.
529	Weems, R. E., Lucas, S.G., 2015. A revision of the Norian conchostracan zonation in North
530	America and its implications for Late Triassic North American tectonic history. New
531	Mexico Mus. Nat. Hist. Sci. Bull. 67, 303–318.
532	Zouheir, T., Hminna, A., Klein, H., Lagnaoui, A., Saber, H., Schneider, J.W., 2020. Unusual
533	archosaur trackway and associated tetrapod ichnofauna from Irohalene member
534	(Timezgadiouine Formation, late Triassic, Carnian) of the Argana Basin, Western
535	High Atlas, Morocco. Hist. Biol. 32, 589-601.
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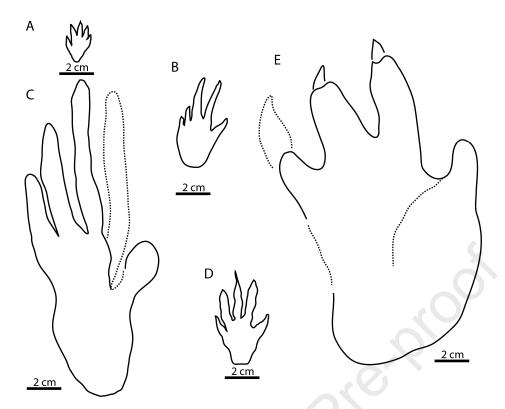
542	Figure captions
543	
544 545	<b>Fig. 1</b> . Geological and stratigraphical setting of Marrakech High Atlas around the Ourika and Yagour basins. <b>A</b> , Permian and Triassic outcrops throughout the High Atlas:
343	ragour basins. A, remitan and rhassic outcrops unoughout the riigh Atlas.
546	CHA = Central High Atlas, EHA = Eastern High Atlas, MA = Middle Atlas, MHA =
547	Marrakech High Atlas and WHA = Western High Atlas (modified after Beauchamp, 1988;
548	Laville and Piqué, 1991). Rectangle indicates the study area. ${\bf B}$ , geological map modified after
549	Biron (1982), showing tetrapod trackway sites marked by stars with number: (1) East of
550	Aghbalou, (2) North-East of Anammer, and (3) NW of Yagour Plateau. C, Simplified
551	lithostratigraphic column of the MHA area (F1 to F6 Fm). The studied interval is localised at
552	the base of the F5 Fm. (1) early Permian (radioactive and zircon method ,Van Houten, 1977);
553	(2) middle Anisian (from pollens, El Arabi et al., 2006); (3) middle Carnian (from pollens,
554	Biron and Courtinat, 1982); (4) Ladinian-early Carnian (from pollens and conchostracans;
555	Benaouiss et al., 2018); (5) middle Carnian (from pollens; Cousminer and Manspeizer, 1976).
556	
557	Fig. 2. Isochirotherium. A–C, Slab UCA.82 bearing several consecutive pes/manus sets (P.,
558	pes track; M., manus track), photograph (A), photogrammetric 3D mesh (B) and interpretative
559	sketch (C); the red dotted line indicates the axis of the partial trackway. <b>D</b> – <b>F</b> , Pes/manus set;
560	photograph (D), digital elevation model in false colours (E) and interpretative sketch (F).
561	Convex hyporeliefs. <b>G–H</b> , details of the skin trace; photograph (G), and interpretative sketch
562	(H).
563	
564	Fig. 3. Apatopus lineatus. A–C, Slab UCA.83. bearing several consecutive pes/manus sets,
565	two of them forming a trackway (blue dots) (LP. left pes track, LM. left manus track, RP.
566	right pes track, RM. right manus track), photograph (A), photogrammetric 3D mesh (B) and
567	interpretative sketch (C). <b>D–G</b> , Two pes/manus sets; photograph (D, F) and interpretative
568	sketches (E, G). Convex hyporeliefs.
569	

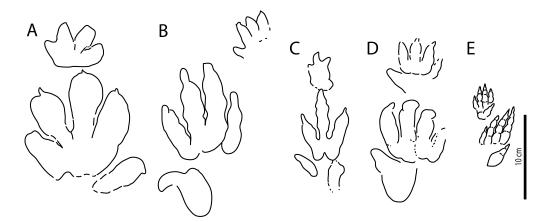
570	Fig. 4. Rhynchosauroides isp., two tracks from the slab UCA.82. A–D, Isolated tracks,
571	photographs (A, C) and interpretative sketch (B, D). Convex hyporeliefs.
572	
573	Fig. 5. Invalid or problematic ichnotaxa reported by Biron and Dutuit (1981) in the Triassic of
574	the Ourika-Yagour basins. A, "Enigmatopus atlensis" (invalid ichnotaxon in this study). B,
575	"Hyloidichnus triasicus" (nomen dubium in this study). C, "Rectilinetopus ourikensis"
576	(Apatopus lineatus in this study). D, "Prochirotherium triasicum" (nomen dubium in this
577	study). <b>E</b> , "Pseudochirotherium oukaimedensis" (nomen dubium in this study).
578	
579	Fig. 6. Comparison between some Triassic Chirotherian tracks. A, Brachychirotherium, Late
580	Triassic, Germany. B, Isochirotherium Iomasi, Middle Triassic, Great Britain. C,
581	Chirotherium ("Parachirotherium"), Middle Triassic, Germany. <b>D</b> , Protochirotherium, Early
582	Triassic, Germany. E, Synaptichnium, Middle Triassic, Great Britain. Modified after Klein
583	and Lucas (2021).
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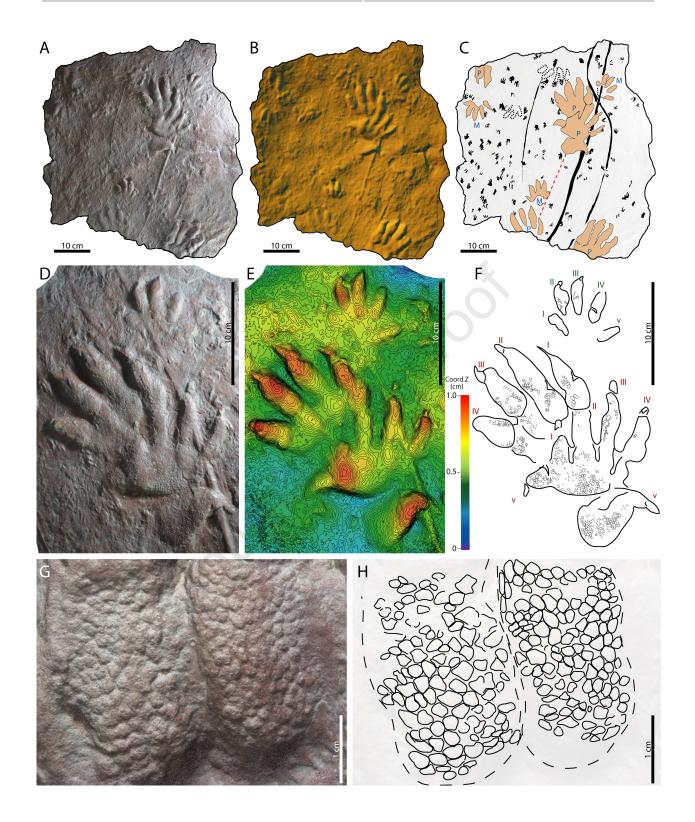












- -Triassic tetrapod tracks from the Ourika-Yagour basins (Morocco) are studied.
- -The historical collection of Pierre-Eymard Biron is ichnotaxonomically revised.
- -These exquisitely preserved trackways are studied using 3D photogrammetric technique.
- -Tracks include Apatopus lineatus, Isochirotherium isp. and Rhynchosauroides isp.
- -They suggests a Middle-Late Triassic age for the Oukaïmeden Sandstone Formation.



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To Editor-in-Chief of Journal of African Earth Sciences

Dear Editor-in-Chief,

### **Declaration of interests**

Yours sincerely,

Declaration of interests
oxtimes The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.
□The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

Dr. Jean-David Moreau for the authors